

Application of time-of-flight aerosol mass spectrometry for the real-time measurement of particle phase organic peroxides: An online redox derivatization – aerosol mass spectrometer (ORD-AMS)

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Abstract. Aerosol mass spectrometers (AMS) are frequently applied in atmospheric aerosol research in connection with climate, environmental or health-related projects. This is also true for the measurement of the organic fraction of particulate matter, still the least understood group of components contributing to atmospheric aerosols. While quantification of the organic/inorganic aerosol fractions is feasible, more detailed information about individual organic compounds or compound classes can usually not be provided by AMS measurements. In this study, we present a new method to detect organic peroxides in the particle phase in real-time using an AMS. Peroxides (ROOR') are of high interest to the atmospheric aerosol community due to their potentially high mass contribution to SOA, their important role in new particle formation and their function as 'reactive oxygen species' in aerosol-health-related topics. To do so, supersaturated gaseous triphenylphosphine (TPP) was continuously mixed with the aerosol flow of interest in a condensation/reaction volume in front of the AMS inlet. The formed triphenylphosphine oxide (TPPO) from the peroxide/TPP reaction was then detected by an aerosol mass spectrometer (AMS), enabling the quantitative determination of peroxide with a time resolution of two minutes. The method was tested with freshly formed and aged biogenic VOC/ozone SOA as well as in a short proof-of-principle study with ambient aerosol.

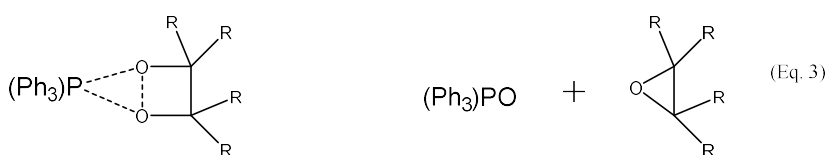
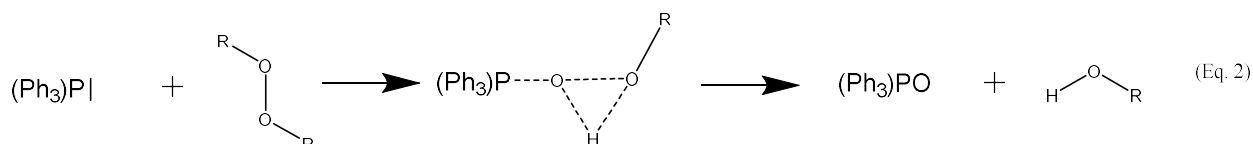
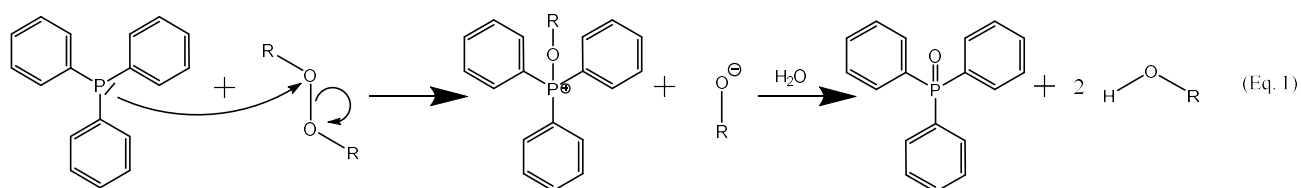
1 Introduction

Atmospheric aerosols attract attention in atmospheric research owing to their effects on climate and human health. The climate relevance of aerosol particles arises from their influence on the transmissivity of the atmosphere for solar radiation and on cloud properties (McNeill, 2017). Secondary organic aerosols (SOAs) are a major contributor to tropospheric particulate matter (Hallquist et al., 2009; Poschl, 2005). One highly relevant class of compounds which is assumed to play an essential role in SOA formation, including participation in new particle formation, are organic peroxides (Bonn et al., 2004; Mochida et al., 2006; Reinnig et al., 2008; Reinnig et al., 2009; Tobias et al., 2000; Tobias and Ziemann, 2002; Zhou et al., 2018; Bianchi et al., 2016; Bianchi et al., 2017; Bianchi et al., 2019; Ziemann, 2002, 2003; Kahnt et al., 2018; Zhang et al., 2017; Lee et al., 2019). Besides their significant contribution to SOA mass (Docherty et al., 2005; Ziemann and Atkinson, 2012), particle phase peroxides are especially interesting due to their often very low saturation vapor pressure, particularly when they contain several hydroperoxy groups induced by autoxidation (Tröstl et al., 2016; Krapf et al., 2016). Also their chemical properties such as their oxidizing character (Claflin et al., 2018) or their function as an indicator for the oxidizing capacity of the troposphere (Thompson, 1992; Lelieveld et al., 2008) create a special interest in the analysis of organic peroxides. Finally, peroxide functionalities are likely to play an important role in health-related effects of SOA, since organic peroxides are believed to contribute to the formation of reactive organic species (ROS) by transporting oxidants on/in particles into the respiratory system (Verma et al., 2015; Wragg et al., 2016; Fuller et al., 2014; Steimer et al., 2017; Tong et al., 2018).

The fact that organic peroxides are oxidizing agents is often used in their analysis. One of the oldest and most frequently applied techniques to measure and quantify peroxides is iodometry (Mertes et al., 2012; Docherty et al., 2005; Mutzel et al.,

2013; Zhao et al., 2018). Alternatively, triphenylphosphine (TPP) can be used as a reacting agent. TPP rapidly reacts with
 40 hydrogen peroxide, organic (hydro)peroxides and peroxy acids to triphenylphosphine oxide (TPPO) (Docherty et al., 2004;
 Ruiz et al., 2001), which can be detected spectrometrically (Porter et al., 1979). There are various suggestions about the
 mechanism of the reaction. One of the mechanisms involves a bimolecular nucleophilic displacement of TPP on the O-O bond
 resulting in a bipolar transition state (Eq. 1.) (Jarvie et al., 1971; Chellamani and Suresh, 1988). Two other suggested
 mechanisms proceed via a cyclic peroxide intermediates (Eq.2 + 3) (Davies and Feld, 1958)(Jayanthi and Koshore, 1997)
 45 The main application of the TPP/TPPO-method is the determination of hydroperoxides from the autoxidation of oils and fats
 (Gotoh et al., 2011; Nakamura and Maeda, 1991; Wang et al., 2016). Compared to iodometric peroxide determination, the
 TPP/TPPO method is insensitive to moisture and the reaction with oxygen is much slower.

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Triphenylphosphine (TPP) MW 262 g mol⁻¹
 R = hydrogen, alkyl, acyl
 R ≠ tert-alkyl

Triphenylphosphine oxide (TPPO) MW 278 g mol⁻¹

Another technique that has been used for the analysis of organic peroxides is mass spectrometry (Krawczyk and Baj, 2014),
 for example for the measurement of SOA related organic peroxides by direct infusion of liquid samples into an atmospheric
 55 pressure chemical ionization–tandem mass spectrometer (Zhou et al., 2018). However, all the above-mentioned techniques are
 offline methods relying on filter sampling techniques, which have disadvantages such as the evaporation of semi-volatile
 species or chemical alterations of labile compounds during sampling. Obviously also the temporal resolution of offline methods
 is limited. Therefore, online methods for organic peroxide measurements would be beneficial.

One of the main online techniques that is used in the field of aerosol research is the aerosol mass spectrometer (AMS), which
 60 is commercially available from Aerodyne Research Inc. (Billerica, MA, USA). Aerosol particles are introduced into a vacuum
 chamber through an aerodynamic lens that focuses the sampled aerosol particles to a beam (Jayne et al., 2000). The particle
 beam is directed onto an evaporation unit, which is heated to 600 °C, ideally volatilizing the entire particle mass. The
 evaporated compounds are subsequently ionized by electron ionization (EI) at 70 eV and analyzed by MS (Pratt and Prather,
 2012a, 2012b; Jayne et al., 2000). The application of EI allows the analysis of a wide range of chemical species and provides
 65 robust quantitative information on various non-refractory ambient aerosol components since the ionization efficiency of EI is
 much less analyte and matrix dependent than, for example, chemical ionization techniques (Allan et al., 2003b; Allan et al.,
 2003a; Canagaratna et al., 2007). However, EI induces a high degree of fragmentation of organic compounds due to excess
 internal energy imparted during the ionization processes. Usually, these fragmentation steps impede the identification and

quantification of individual particle phase organics, thus leading to a reduced level of chemical information (Zhang et al., 2011). Therefore, recent mass spectrometric instrumental developments for organic aerosol analysis focus on softer ionization techniques such as atmospheric pressure chemical ionization (APCI-MS) (Vogel et al., 2013b; Vogel et al., 2013a; Brüggemann et al., 2014; Warscheid and Hoffmann, 2001a, 2001b), flowing atmospheric-pressure afterglow mass spectrometry (AeroFAPA-MS) (Brueggemann et al., 2015; Brueggemann et al., 2016; Brueggemann et al., 2017) or extractive electrospray ionization (EESI-MS) (Gallimore et al., 2017). However, while these mass spectrometric techniques are applicable for a wide spectrum of organic compounds and functionalities, for example, carboxylic acids (Hallquist et al., 2009; Yatavelli et al., 2015), unambiguous identification of organic peroxides using mass spectrometric techniques is still challenging, since in contrast to other functionalities the molecular ions usually fragment resulting from the cleavage of the peroxide bond and/or further fragmentation (Reinnig et al., 2009; Guascito et al., 2015).

Here, we have combined the concept of selective peroxide determination by a suitable redox-reaction (the TPP – TPPO system) to use this reaction as an online derivatization method to monitor the educts and products with an aerosol mass spectrometer (online redox derivatization-AMS, ORD-AMS). The development and application of the method presented in this paper is essentially a proof-of-principle study, however, the results clearly show the potential of online derivatization as an extension of regular AMS measurements, in this case for the understanding of particle phase organic peroxides. Similar experimental setups with different derivatization agents are also conceivable in the future to further extend the usability of AMS systems.

2 Experimental

2.1 Online redox derivatization

Figure 1 shows a schematic diagram of the experimental setup for the AMS online measurements. Gas phase organics were firstly removed by two charcoal filled denuders (inner tube: diameter = 2 cm, length 40 cm, outer tube: diameter 13.5 cm, length 40 cm) and the particle size distribution was measured using a scanning mobility particle sizer (SMPS, Grimm Aerosol Technik, Ainring, Germany).

When the instrument was used in the ORD-AMS mode (blue path in Fig. 1) the particles to be analysed were introduced into a 500 mL glass tube (35 cm length, max. diameter 4.5 cm), the condensation/reaction volume. Here, the aerosol was mixed with the gaseous reactant (TPP, 99%, Acros Organics BVBA, Geel, Belgium), which was supplied from a Sinclair-Lamer setup (see Figure 1, blue triangle). The TPP source consisted of a 100 mL two-necked round bottom glass filled with TPP and was heated to a temperature of 90 °C. Nitrogen was used as a carrier gas to deliver the TPP to the condensation/reaction volume, where TPP condensed on the available particle surface. After passing the condensation/reaction volume the aerosol was led through a third denuder filled with active charcoal to remove gas phase TPP and finally introduced into the AMS and a condensation particle counter (CPC, PortaCountPlus, TSI In., Shoreview, MN, USA). A total flow rate of 1.74 L/min was obtained using an additional pump. The residence time of the aerosol in the reaction volume and the denuder to the AMS was about 35 s.

When the instrument was used in the AMS mode (red path in Figure 1), the aerosol was led directly to the AMS and the CPC. Switching between the two modes was achieved by manually changing the PTFE connections to avoid losses or artifacts from additional valves. The data acquisition of the AMS and SMPS was interrupted during this time and continued (1-2 min) after stabilisation of the particle number concentration measured by the CPC. As mentioned above, a nitrogen stream enriched with TPP was added to the aerosol stream in the ORD-AMS mode resulting in a dilution of the aerosol. The correction for this dilution was done by comparing the particle number concentrations, measured in the ORD-AMS and AMS mode using the CPC. The ratio between the particle number concentrations of the two modes was used as a dilution factor.

2.2 ToF-AMS

The aerosol particles were introduced into the Time-of-Flight (ToF)-AMS (Aerodyne Research Inc., Billerica, MA, USA) at a sampling flow rate of 1.23 mL s^{-1} . An aerodynamic lens system focuses the particles in a size range of $\sim 35\text{--}1000 \text{ nm}$ into the vacuum system. While most of the remaining gas phase is removed, the particle beam passes into the particle sizing chamber. The ToF-AMS was either operated in the Mass Spectrum (MS) or the Particle Time-of-Flight (PToF) mode. The first mode provides information about the overall composition of the total aerosol mass without any information about the particle size, the second mode results in the size-segregated analysis of the particle composition. For the latter one, a chopper allows the transmission of only a small part of the aerosol into the PToF chamber. Due to increasing inertia with larger particle size, the particle package is separated by their time of flight and the size distribution based on the mean vacuum aerodynamic diameter of the aerosol can be determined. In this PToF-mode, the EI-ToF-MS functions as a detector for the particle time of flight. To determine the concentration in the MS-mode, the beam passes the chopper alternately blocked or unblocked for background corrections. After passing the particle sizing chamber, the aerosol particles are vaporized on a heated porous tungsten surface at $600 \text{ }^\circ\text{C}$ followed by immediate ionization in electron impact (EI) mode at 70 eV . The generated ions are extracted by an orthogonal extractor into the time ToF-MS, which was operated in the V-mode to give higher sensitivity. The MS spectra acquisition was performed in the positive ion mode. The AMS data were processed by the Igor (Version Pro 6.37) based programs "SQUIRREL 1.57 G" (unit mass resolution) and "PIKA 1.16H" (high resolution) (DeCarlo et al., 2006; Canagaratna et al., 2007).

2.3 Physical aerosol characterization and generation

A Scanning Mobility Particle Sizer was used to measure the size distributions of the particle number concentration of the incoming aerosol, which was used to obtain particle surface, volume and mass concentration. The conversion of the volume concentration into the mass concentration for the ozonolysis experiments is described in detail in the additional information. Besides the measurement of the particle size distribution, the instrument was used to generate monodisperse particles for instrument calibration. The EI-spectra of the derivatization agent TPP and its reaction product with peroxides TPPO ($> 98 \%$, Merck KGaA, Darmstadt, Germany) were recorded using a nebulizer as particle generator (Fig. 1). Pure TPP and TPPO particles were produced from solutions in methanol (Optima® LC/MS grade, Fisher Scientific, Loughborough, UK) and *n*-heptane (HPLC grade, Carl Roth GmbH + Co. KG, Karlsruhe, Germany), respectively. For these experiments, the particles were size-selected using the DMA in a setup as described in SI.

2.4 Reaction chamber as SOA generator

For the proof-of-principle studies with SOA, a laboratory-scale reaction chamber in continuous-flow tank reactor mode was used. Synthetic air was supplied with a total flow of 7.6 L min^{-1} into the 100 L chamber, resulting in a residence time of 13.5 min . The total flow was composed of a gas stream through a diffusion test gas source (Thorenz et al., 2012) (0.600 L min^{-1}), where the VOC of interest was added to the synthetic air supply. A second gas stream (2.6 L min^{-1}) was humidified using a gas washing bottle and third gas flow was led through an ozone generator (Dasibi Environmental Corp. Model 1008 RS O3 analyzer, Glendale, CA, USA, 2.7 L min^{-1}). To add seed particles an aqueous solution of ammonium sulfate (ca 0.5 g/L , $> 99.5\%$, Merck KGaA, Darmstadt, Germany) was nebulized using technical nitrogen and introduced into the chamber (1.7 L min^{-1}). Temperature and relative humidity were $25 \text{ }^\circ\text{C}$ and 61% for the α -pinene ozonolysis and $22 \text{ }^\circ\text{C}$ and 59% for the β -pinene ozonolysis. They were monitored by a thermo/hygrometer (Amarell, ama-digit ad 910 h, Kreuzwertheim, Germany). The chamber was set under slight overpressure to prevent the entry of laboratory air into the setup. The reaction chamber was wrapped with aluminum foil to avoid photochemical reactions. For the ozonolysis of α -pinene (99% , Fluka, Seelze, Germany),

150 the chamber was first equilibrated with the terpene before the ozone generator was switched on. For the β -pinene (99%, Sigma Aldrich Chemie GmbH, Steinheim, Germany) experiments, the order of the addition of the individual compounds was reversed, i.e. ozone was added before the terpene. The ozone concentration was estimated to be about 1 ppm, the α -pinene concentration was set to 290 ppbv and the β -pinene concentration to 380 ppbv. To remove gaseous organic compounds the SOA formed in the reaction chamber was led through two diffusion denuders filled with activated charcoal.

155 3 Results and discussion

3.1 EI mass spectra

Figure 2 shows the obtained mass spectra of TPPO and TPP. The EI spectra of the two reference compounds measured by the AMS are very similar to the spectra measured by regular MS setup (Williams et al., 1968; Stein and Slawson, 1963). The base peak for TPPO is at m/z 277 resulting from a hydrogen abstraction ($[M-1]^+$) from the molecular ion (MW (TPPO) = 278 g mol⁻¹). The second-largest peak is m/z 77 formed from the abstraction of a phenyl cation. The base peak for TPP is located at m/z 183 originating from the abstraction of two hydrogens and a phenyl group, probably a result of an aromatic stabilization of the resulting dibenzophosphole cation (Williams et al., 1968). However, also the molecular ion of TPP at m/z 262 is visible. More important is the small but still observable $[M-1]^+$ ion of TPPO in the TPP spectra. Since the peroxide determination is based on the measurement of TPPO, this background signal has to be taken into account (see below). Finally, aerosol species typically found in ambient air and measured by the AMS have mass spectra with signals mainly below m/z 100 due to the fragmentation after EI ionization. As seen also in Figure S4 and S5 in SI, the signals at m/z 277 and 262 are specific for TPPO and TPP, respectively. These signals can be used for the quantification of both compounds.

Experiments with peroxides (R-O-O-R) as reference compounds were also performed, using commercially available organic peroxides. e.g. lauryl peroxide and benzoyl peroxide. While these experiments clearly showed the formation of TPPO from TPP, the use of these compounds for quantification was not possible, since the peroxide standards studied undergo a substantial change in volatility after the reaction with TPP. Therefore, lauryl peroxide and benzoyl peroxide were used to demonstrate the general proof of feasibility of the method, however, not for calibration.

3.2 Quantification

The mass concentration c_m of a single compound measured by an AMS can be expressed as (Canagaratna et al., 2007):

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$$c_m = \frac{10^{12} \cdot MW(\text{compound})}{IE(\text{compound}) \cdot Q \cdot N_A} \cdot \sum_{\text{all},i} I_{\text{compound},i} \quad (\text{Eq. 6})$$

where MW and IE is the molar weight and the ionization efficiency of an individual compound. The ionization efficiency (IE) is a dimensionless quantity that describes the ionization and detection efficiency and is defined as the ratio of the ions detected to the parent molecules vaporized. Q is the flowrate into the AMS, N_A the Avogadro's number and 10^{12} a unit conversion factor. $I_{\text{compound},i}$ is the ion rate of an ion i formed by ionization and fragmentation of a compound. Since the fragmentation pattern can be assumed to be robust, it is possible to use the ion yield of a single ion $I_{m/z,i}$ to calculate the concentration. The molar concentration c_n of peroxides is according to the reaction equation the same as the concentration of TPPO and can be calculated from the signal at m/z 277 as:

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$$c_n(\text{TPPO}) = \frac{10^{12}}{IE\left(\text{TPPO}, \frac{m}{z} i\right) \cdot Q \cdot N_A} \cdot I_{\frac{m}{z} i} \quad (\text{Eq. 7})$$

$$c_n(TPPO) = \frac{10^{12}}{IE\left(TPPO, \frac{m}{z} 277\right) \cdot Q \cdot N_A} \cdot I_{\frac{m}{z} 277} \quad (\text{Eq. 8})$$

$$c_n(TPP) = \frac{10^{12}}{IE\left(TPPO, \frac{m}{z} 262\right) \cdot Q \cdot N_A} \cdot I_{\frac{m}{z} 262} \quad (\text{Eq. 9})$$

190 The experimental procedure to estimate IE is discussed in more detail in SI.

3.3 Background correction

As mentioned above also TPP particles, as well as pure ammonium sulfate aerosol particles passing the TPP-condensation/reaction chamber, show a small but non-negligible TPPO signal (m/z 277). This is either a consequence of TPPO impurities in the TPP itself or due to traces of peroxides in the solvents used to generate the aerosol particles when the nebulizer was used. Consequently, a background correction procedure has to be introduced to account for these effects. To do so, pure AS aerosol particles were investigated with different initial ammonium sulfate concentrations with the ORD-AMS setup as described above. These experiments showed a constant background contribution of 1,222 (+/-0.1) % TPPO from the conversion of TPP (m/z 262 \Rightarrow C_n (detected TPP)). Therefore, the concentration of the analyte of interest (C_n (peroxide)) was calculated as follows:

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$$c_n(\text{backgr. TPPO}) = 0,01215 \cdot c_n(\text{detected TPP}) \quad (\text{Eq.10})$$

$$c_n(\text{produced TPPO}) = c_n(\text{detected TPPO}) - c_n(\text{backgr. TPPO}) \quad (\text{Eq.11})$$

$$c_n(\text{peroxide}) = c_n(\text{produced TPPO}) \quad (\text{E1.12})$$

3.4 SOA experiments

205 The ozonolysis of α - or β -pinene was performed in the presence of ammonium sulfate (AS) seed particles and the results of a typical chamber run with α -pinene is shown in Figure 3. First, AS seed particles were injected into the chamber. After a certain period of time, when the particle concentration became constant, the biogenic VOC was introduced into the chamber (e.g. after 34 min in the run shown in Fig. 3). Following an initial period to monitor background peroxide concentrations in the AS seed aerosol / VOC system, ozone was introduced (e.g. after 135 min in the α -pinene run shown in Fig. 3 (green area)). Periods, when TPP was not added (AMS mode in Fig. 1), are marked in grey and were used to determine the dilution factor and the concentration of SOA by the AMS. As can be seen in the figure, the concentrations of both TPPO and TPP rose quickly during the first 20-25 min after the ozone addition (Fig. 3a), followed by a slow decrease towards the end of the experimental run. In fact, peroxides were only detected when SOA formation began, as indicated by the time series of peroxide concentration (Fig. 3b, black) and SOA mass concentration (Fig. 3b, blue line). Figure 3c shows the temporal behavior of the relative contribution of peroxides to SOA, whereby an average molar mass of 250 g mol⁻¹ for SOA was assumed. The course of the peroxide quantity per SOA quantity (Fig. 3c) shows a continuous decrease during the entire experimental run. Such concentration profiles of the organic peroxides are indicative of consecutive reactions, in which the peroxides are formed as first-generation products from the initial ozone/alkene reaction in the gas phase, followed by gas-to-particle conversion and further reactions of the peroxides to non-peroxidic products in the particle phase. Indeed, several studies suggest that organic peroxides are instantly formed in the VOC-ozone reaction (Bianchi et al., 2019; Mertes et al., 2012) and the rapid increase of the peroxide signal as shown in Figure 3b confirms these studies. However, since the O-O bond is relatively weak, organic peroxides tend to decompose or react with other aerosol components. Possible reactions to non-peroxidic products could be reactions with alcohols or aldehydes forming aldehydes or carboxylic acids, or hydrolysis followed by the loss of H₂O₂ (Zhao et al., 2018; Clafin et al., 2018; Epstein et al., 2014). In any case, the temporal evolution of the peroxide concentration (i.e. the

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225 formation/decline profile) indicates consecutive reactions and demonstrates the intermediate character of the peroxides within the VOC oxidation system.

As can be seen in Figure 3c, the contribution of peroxides to SOA was almost 70% after about 20 minutes. During the last 80 min of the experiment, the SOA concentration was about $69 \mu\text{g m}^{-3}$ ($0,28 \mu\text{mol m}^{-3}$), assuming again an average molar mass
230 of the SOA components of 250 g mol^{-1} . During this time the measured peroxide concentration was about $0.0470 \mu\text{mol m}^{-3}$.
Zieman and coworkers (Docherty et al. 2005) proposed an average molar mass of 300 g mol^{-1} for SOA-peroxides. Based on
this value the average peroxide mass concentration was $14,1 \mu\text{g m}^{-3}$, assuming the same density of peroxides as for other SOA
components (1.2 g cm^{-3}). Thus the average peroxide/SOA yield within this period is estimated to be about 20 % ($\pm 4\%$).
Actually, this value is very close to the value of 22 % (Epstein et al. 2014) or the value of 21 % (Li et al. 2016). Mertes and
235 coworkers (Mertes et al. 2012) observed 34, 17 and 12 % for 15-35 min, 4 hours and 6 hours filter samples, respectively.
Taking into account the influence of sampling time, duration of sampling, initial ozone and terpene concentrations as well as
temperature and relative humidity, these values agree quite well.

As can be seen in Figure 3a, the TPP concentration also increased after the start of ozonolysis. This is understandable,
240 considering that TPP is only detected by the AMS when it condenses on particles. The more particles are introduced into the
ORD setup, the higher the TPP concentration measured by the AMS. Figure 4 shows the particle size distribution measured
by the SMPS system before (black) and within the first 10 minutes after (red) ozone was introduced into the reaction chamber.
Despite the presence of AS seed particles, new particle formation from nucleation of organic compounds was observed,
resulting in a bimodal size distribution at the beginning of the experiment. In principle, new particle formation is not surprising
245 due to the high particle formation potential of biogenic VOCs as well as the relatively high concentrations of the reactants.
The same behavior was also visible in the AMS data. One of the main advantages of the AMS is the possibility to determine
the size distributions of individual compounds. If the m/z is characteristic for an aerosol component, such as m/z 30 and 46 for
nitrate, for example, the size distribution of these components can be obtained. Figure 5 shows the m/z values as a function of
particle size for TPP (m/z 262 (blue)), TPPO (m/z 277 (black)), SOA (m/z 43 (green)) and sulfate (red) for the first 10 minutes
250 after ozone was introduced, all normalized to the respective maximum value. Although the determined size distributions by
SMPS and AMS are not identical, likely due to scanning time effects, differences in aerodynamic and mobility diameters as
well as differences in the averaging procedure, the bimodal size distribution of pure organic particles and pre-existing AS
containing particles are also visible in the AMS data. While the occurrence of organic nucleation simultaneously with
condensational growth resulting in two externally mixed particle populations is not surprising, a more quantitative description
255 of the size distribution of selected compound groups is challenging. As can be seen in Figure 5, the size profiles of SOA and
TPPO are very similar with a clear preference to the smaller particle sizes. Besides the obvious explanation of the existence of
a pure organic particle population formed by homogeneous nucleation, an additional explanation is a size-dependent
condensation of the organics also on the inorganic seed particles. A unique characteristic of condensational growth is that it is
proportional to the Fuchs-corrected surface area of particles and not the particle volume (“surface-limited” vs “volume-limited”
260 growth). Because the surface area to volume ratio of particles increases as their diameter decreases, condensation tends to have
a larger effect on smaller particles. Therefore, after the condensational growth of AS seed particles the final distribution of the
condensed organics can be expected to favor the smaller particles (Donahue et al., 2014; Riipinen et al., 2011). Therefore, both
nucleation of organics and condensational growth will contribute to the size distribution shown in Fig. 5. Nevertheless, the
results clearly show that the formation of peroxidic compounds is directly connected to SOA formation. More difficult to
265 explain is the quite similar size distribution of TPP and ammonium sulfate seed aerosol since also TPP moves to the particle
phase via condensation. One possible explanation could be attractive forces between the nonpolar TPP molecules and AS,

similar to a 'salting in' effect (Hyde et al., 2017). However, possibly also other effects, such as the high supersaturation of TPP in the condensation/reaction volume, might be considered to explain the observations.

270 Figure 6 a-d show the size distributions of m/z 277 (black), 262 (blue), 43 (green) and sulfate (red) for the periods before (top) and during the last 80 min of the ozonolysis (bottom) with (right) and without (left) added TPP. Again, each of the individual ratios is normalized to its respective maximum. M/z 262 corresponds to the detected TPP, m/z 277 to TPPO and m/z 43 to SOA species (see figure S6 in SI). Figure 6a shows the size distribution of pure AS particles. As can be seen, by comparison of 6a and b, the diameter of sulfate remained almost constant due to the condensation of TPP, if anything then the diameter
275 slightly decreased. Generally, the vacuum-aerodynamic diameter depend on the particle density (DeCarlo et al., 2004). The larger the density, the larger is d_{va} for particles with the same physical diameter. The condensation of TPP (density 1.1 g cm^{-3}) will lower the particle density (density AS 1.77 g cm^{-3}), which compensates the increase of the physical diameter. Nevertheless, the condensation of SOA material finally led to an increase in the sulfate diameter (Fig. 6c). In Figure 6d the size distributions of the four different components are shown. First to notice is again the clear overlap of m/z 277 and m/z 43,
280 demonstrating the simultaneous presence of TPPO and SOA species in the particle phase. The slight shift in the size distribution to larger sizes for SOA (m/z 43) and TPPO is likely a result of slightly faster growth of the pure organic particles formed by nucleation as discussed above compared to the ammonium sulfate containing particles.

A second biogenic VOC, β -pinene, was investigated with the online redox derivatization AMS system, this time with a reverse
285 introduction of the VOC and the oxidant to examine ozone as a direct oxidant for TPP. Again, the ozonolysis of β -pinene was performed in the presence of AS seed aerosol. Figure 7 shows the temporal evolution of the peroxide concentration together with the total aerosol mass concentration determined with the SMPS. As in the case of α -pinene, the peroxide concentration increased rapidly with the formation of SOA, once again demonstrating the direct connection of SOA and peroxide formation via the detection of TPPO. The comparison of the profiles of SOA (blue) and peroxide (black) demonstrates the very rapid
290 peroxide formation followed by the decreasing contribution of peroxides to the particle phase in the later stages of the experiment as a consequence of consecutive reactions of the peroxides to non-peroxidic products. The strong decrease of total SOA after 240 min is induced by the end of β -pinene addition to the reaction system. The observed shift between the peroxide and SOA concentration supports the concept of organic peroxides as reaction intermediates, which contribute especially to the composition of fresh SOA and convert to non-peroxidic products in aged SOA particles. This is also clearly visible in the
295 bottom part of Figure 7, which shows the relative contribution of peroxides to the total particle volume in the course of the experiment. However, most important for the overall method evaluation are the very low and rather constant TPPO concentrations in the period between the ozone and β -pinene addition (between 86 min and 180 min) when just AS particles were introduced into the ORD-AMS system, a result that clearly excludes ozone as a direct oxidant for TPP.

3.5 Discussion

300 As discussed above, the validation of the quantitative performance of the presented measurement method was not performed with peroxide reference compounds. Instead, the ratio between the concentration of peroxides and total SOA at the end of the ozonolysis of α -pinene was compared with the results of Epstein et al. (2014), Li et al. (2016) and Mertes et al. (2012). In these studies the peroxide concentration formed from the ozonolysis of α -pinene using filter or impactor sampling, followed by the extraction and iodometric analysis of the peroxide content, was measured. Several experimental parameters, such as the terpene
305 and ozone concentration, their ratio to each other, the temperature and the relative humidity, were slightly different. However, the comparison show a very good agreement with values close to 20 %. Another aspect of all SOA experiments which include chemistry, as here the peroxide-TPP reaction, that have to be considered is the question of the phase state of SOA. According to Koop et al. (2011) SOA is expected to be a liquid under the condition used for our experiments ($22\text{-}25 \text{ }^\circ\text{C}$, r.h. 60%). Under

these conditions also the AS seed particles can be expected to behave liquid-like (Matthew et al., 2008). As a consequence the
310 diffusion time of TPP within the particles should be approximately one second (Koop et al., 2011), which is much shorter than
the residence time of about 35 seconds of the aerosol in the ORD-setup (Fig. 1). In conclusion, based on the agreement of the
peroxide/SOA ratio at a later stage of the experiments measured in this work and the results of former studies, the estimated
diffusion times and the excess of the reactant TPP to SOA, a quantitative reaction of TPP with the particle phase peroxides to
TPPO can be assumed. However, in future investigations, the method should be revalidated using as aerosol particles composed
315 of pure single component peroxides.

3.6 Ambient air measurements

Finally, the ORD-AMS method was tested with ambient air directly outside the chemistry building located at the campus of
the university in Mainz (49.9915928 N, 8.2312864 E). The measurements were performed between May 4th (starting time 9:30
320 am) and May 9th 2018 (5:30 pm). Meteorological data were measured and provided by the Institute of Physics, also located at
the campus of the university. During the measurement period, the air masses came from north, northeast or east and air mass
composition can be expected to be especially influenced by the Frankfurt Rhine-Main Metropolitan Region with approximately
5.8 million inhabitants. Figure S9 in the SI shows 72 h backward trajectories of the air masses (every 24 hours) from May 4th
to May 10th 2018. The blank value for the seeded SOA experiments was used for the background correction and values below
325 the quantification limit were set to zero. Figure 8a shows the concentration ratios of TPPO to total TPP during the measurement
period and Figure 8b the course of the particle phase peroxide concentration. Actually, the smallest values of detected TPPO
to total TPP shown in Fig. 8a lie very close to the estimated blank value level (red line). This indicates that blank values of
ambient air measurements and the blank value from the seeded SOA experiments are almost identical. Values of the peroxide
concentration above the limit of quantification (LOQ) (blue line) were only obtained during weekdays (May 5th Saturday; May
330 6th Sunday) and at daytime. The enhanced concentration during weekdays at daylight conditions indicates local sources of
peroxides in combination with photochemistry, an observation that is in agreement with the result from previous peroxide
measurements (Hua et al., 2008; Guo et al., 2014). However, due to highly variable emission sources directly outside the
chemistry building the selected measurement site was certainly not ideal for general studies about particle phase peroxide
chemistry. Nevertheless, this short proof-of-principle study demonstrates that peroxides in atmospheric aerosols can be
335 detected in real time with the new ORD-AMS system. .

4 Conclusions

The real-time detection of peroxides by adding gaseous TPP to aerosol particles was successfully realized using an aerosol
mass spectrometer in combination with a specifically designed online derivatization unit. The ozonolysis of α - and β -pinene in
the presence of ammonium sulfate seed aerosol was used to demonstrate the potential of the analytical procedure. The results
340 clearly show the conversion of TPP to TPPO and hence the presence of organic peroxides in biogenic SOA. Especially in the
early phase of SOA formation the contribution of peroxides is exceptionally high (up to ca. 80% v/v). In addition, the time-
resolved measurements show a rapid further reaction of the peroxides to non-peroxidic compounds indicative of consecutive
reactions of the peroxides within the particle phase. The developed instrumental setup can also be used to receive particle size-
resolved information about particle phase peroxides. Again, the important role of organic peroxides in the early stages of
345 particle formation, which was probably induced by very low volatile multifunctional organic peroxides (HOMs), became clear.
Finally, the new online redox derivatisation method was successfully tested with ambient air outside the laboratory. The results
presented here represent the first attempt to better characterize the role of organic peroxides in atmospheric aerosols. Future
laboratory and field studies are certainly necessary to better understand the chemistry of peroxides in organic aerosols.

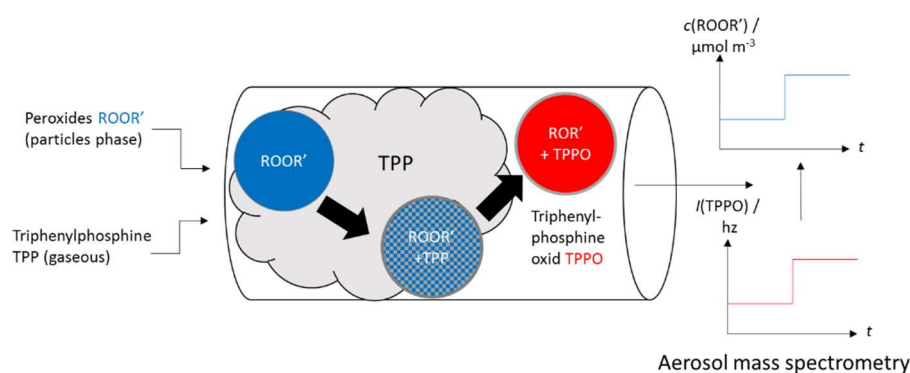
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Data availability. The data obtained by ORD-AMS measurements from the SOA experiments are available on request from Marcel Weloc.

Author contributions. MW and TH designed the study. MW performed the experiments to develop ORD-AMS and performed the laboratory work, calibration as well as the data analysis. MW and TH examined the results, provided input on the data analysis and interpretation and prepared the manuscript.

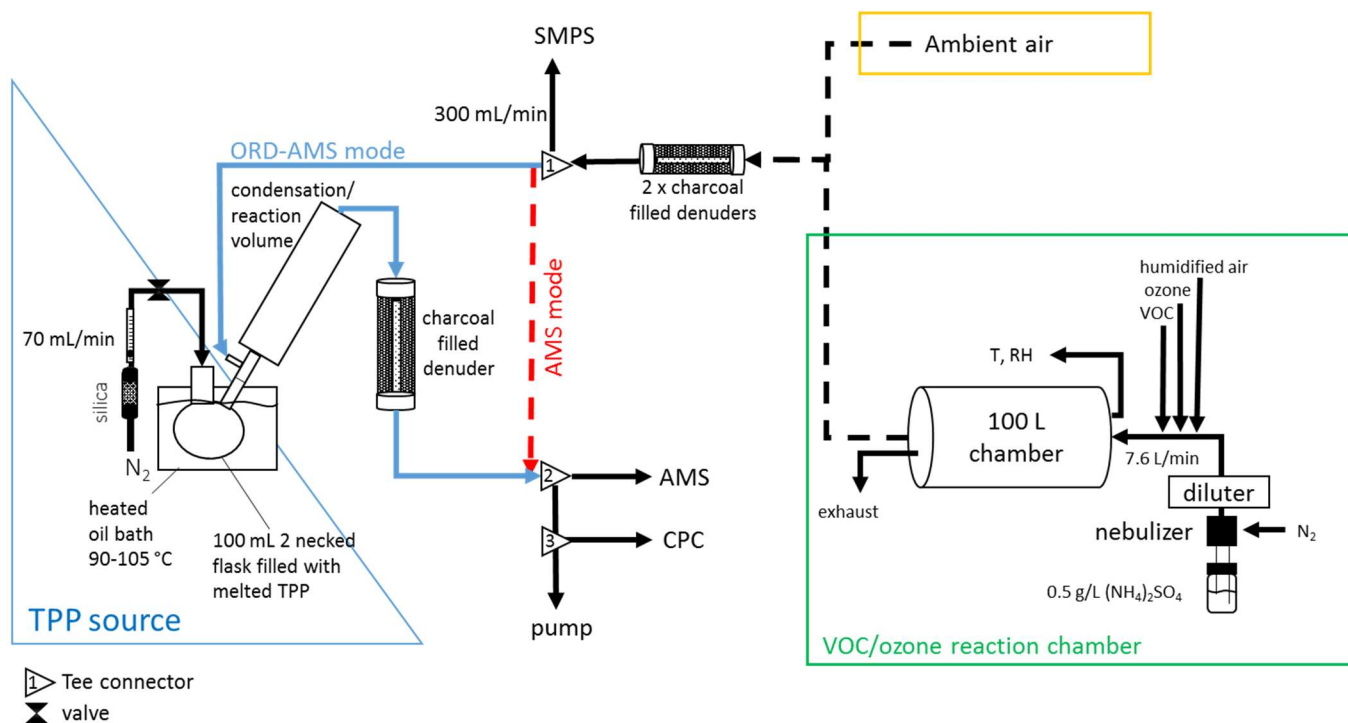
Competing interests. The authors declare that they have no conflict of interest.

360 *Financial support.* This work was supported by the DFG, grant number HO 1748/16-1, ‘Development and application of a real-time method for the determination of selected compound classes (especially organic peroxides) in atmospheric aerosol particles using an aerosol mass spectrometer (AMS)’.



Graphical abstract.

365



370 **Figure 1:** Experimental set up: The aerosol is either generated with a 100 L reaction chamber via VOC/O₃ reaction, usually also introducing AS seed particles or supplied from outside of the chemistry building. The aerosol is then led to the aerosol mass spectrometer, either directly (AMS mode, red arrow) or through the online derivatization unit (ORD-AMS mode) in which the test aerosol is mixed with a nitrogen stream containing supersaturated gaseous TPP. Remaining gaseous TPP is removed by charcoal filled denuder after the condensation/reaction volume. Most of the time the aerosol is characterized by the HR-ToF-AMS, an SMPS system and a CPC.

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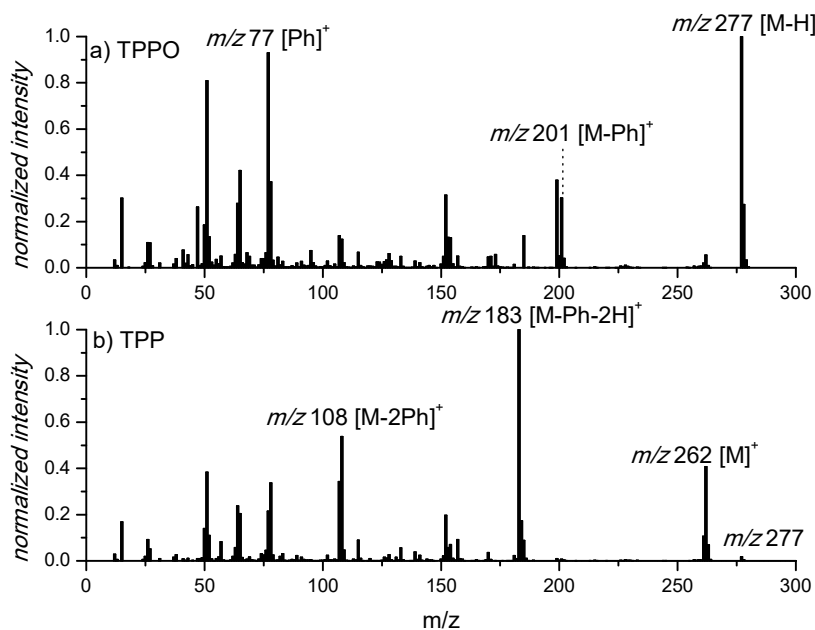
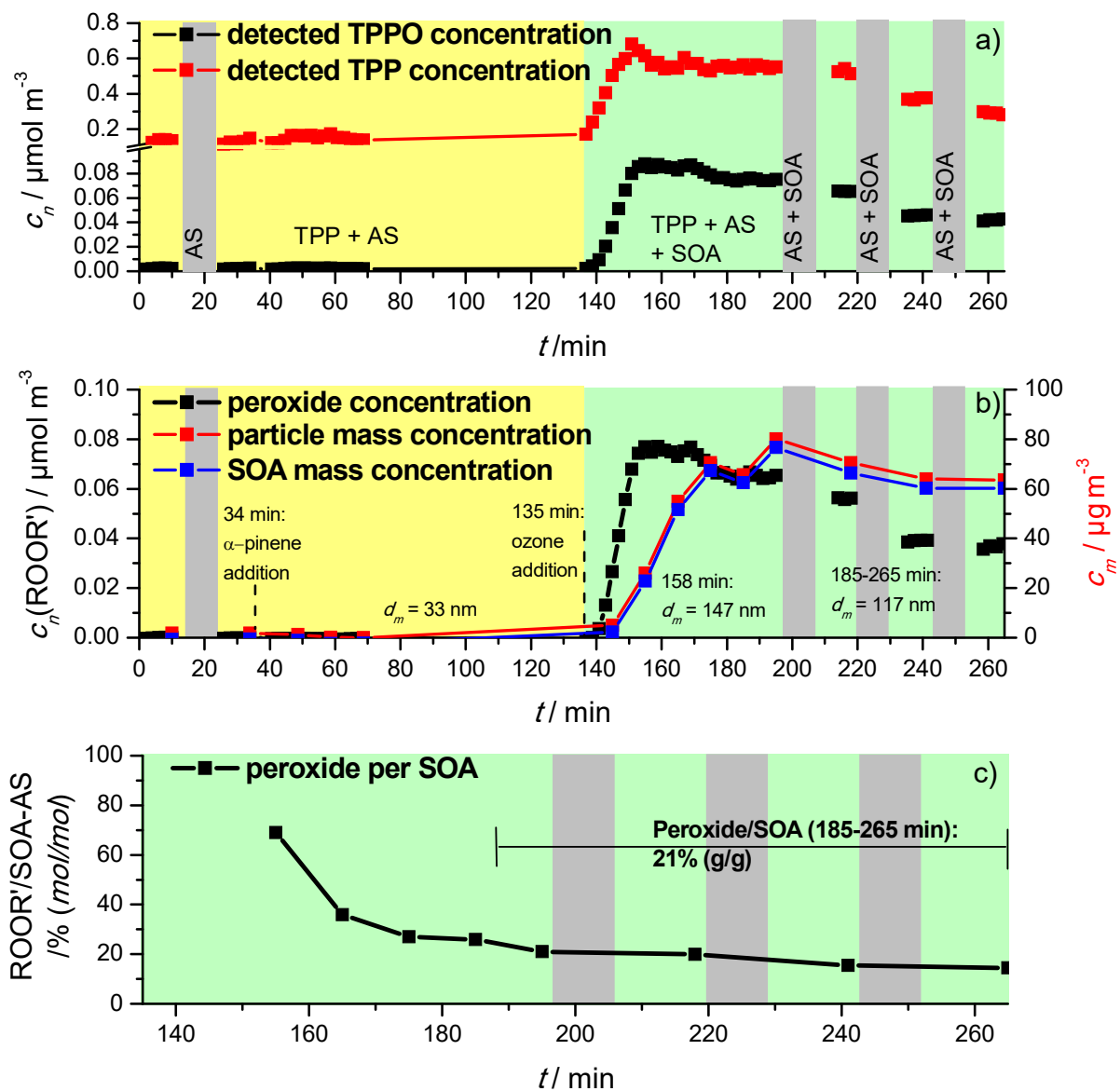


Figure 2: Mass spectra of a) triphenylphosphine oxide (TPPO) and b) triphenylphosphine (TPP).



380 **Figure 3:** Time series of a) detected TPPO (black) and detected TPP (red), of b) peroxide concentration (black) after background correction and particle mass concentration of the SOA-AS aerosol (red) and of SOA (blue), and of c) relative contribution of peroxide to the total aerosol volume (black, time resolution 10 min). Periods without the addition of TPP (AMS mode) are marked with a grey background, the period before the addition of ozone is marked in yellow and during the ozonolysis in green (with TPP, ORD-AMS mode). α -pinene was continuously added after 34 min and ozone after 135 min.

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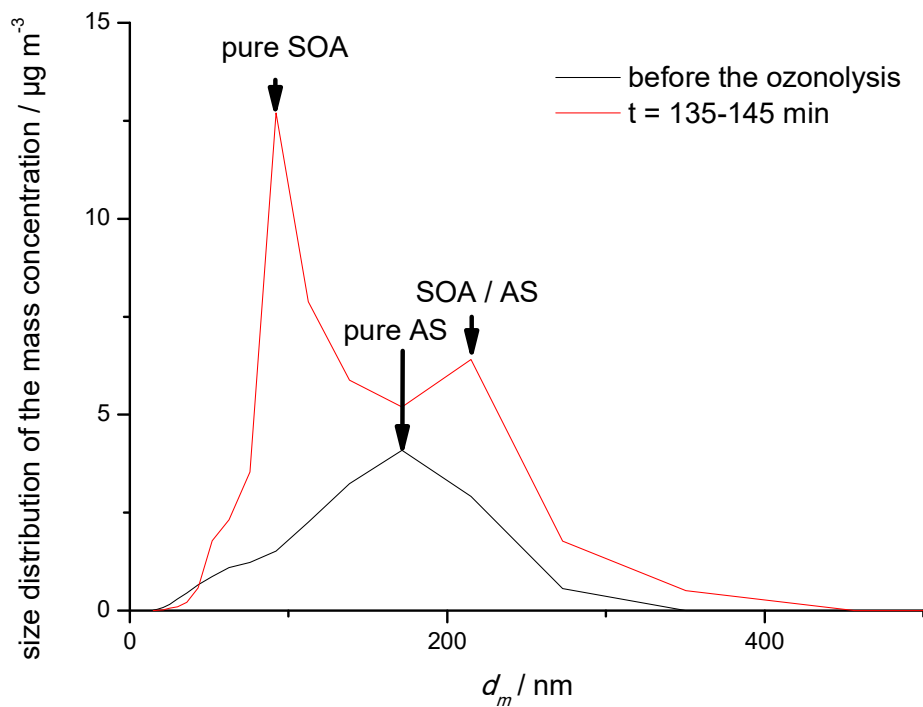


Figure 4: Particle mass size distribution

before and ten minutes after the addition of ozone.

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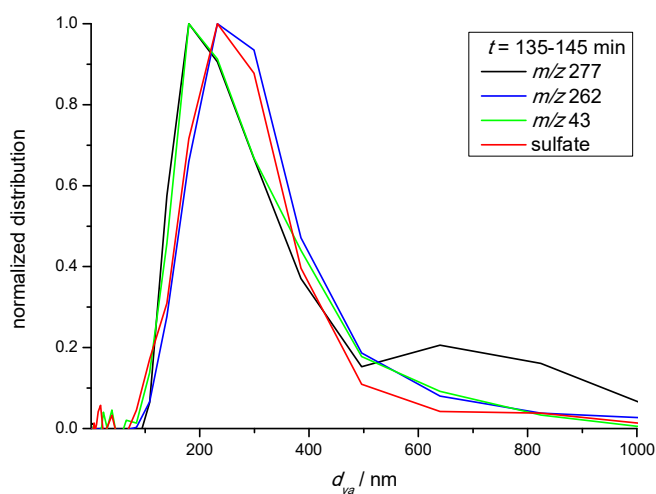
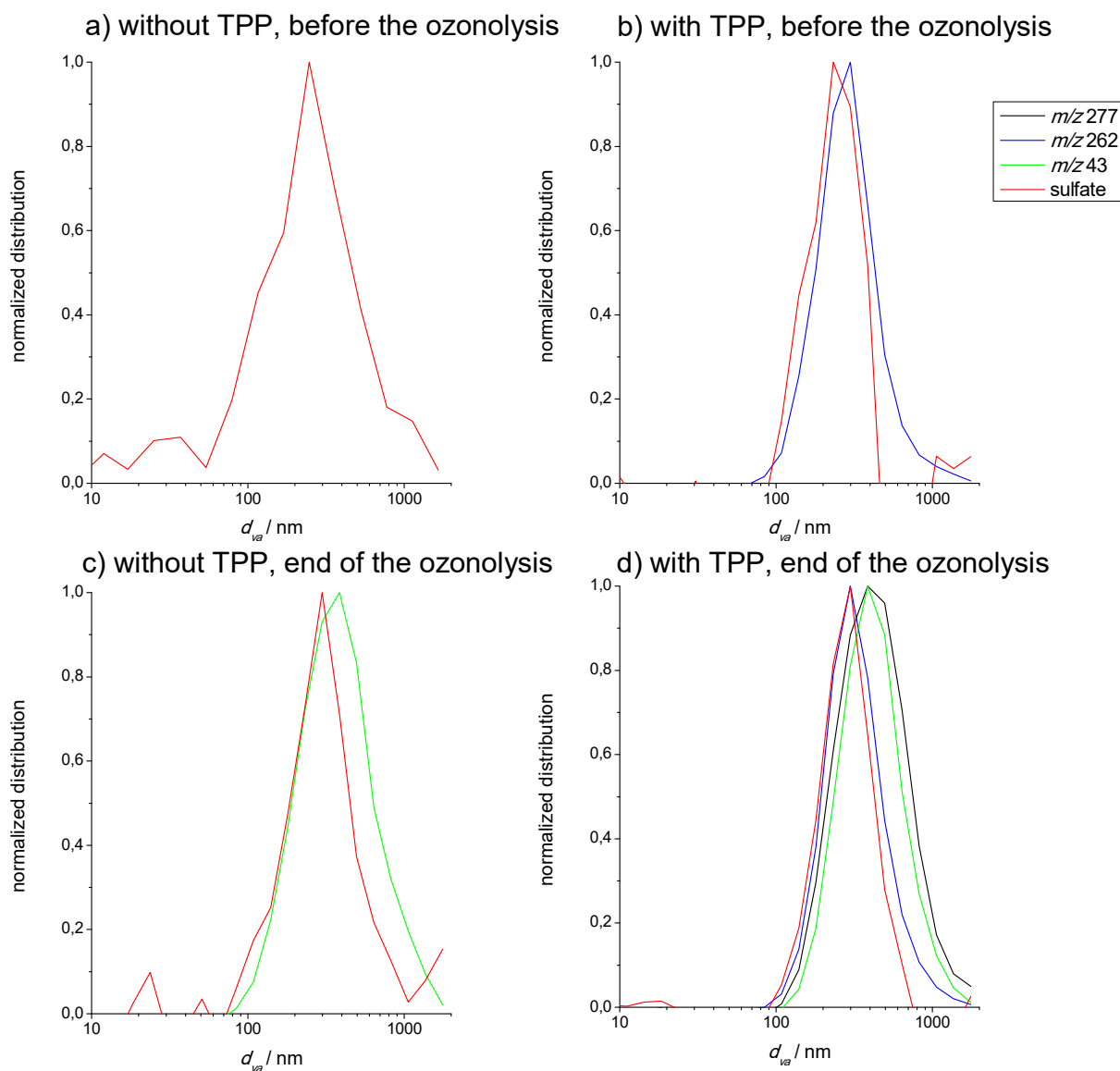
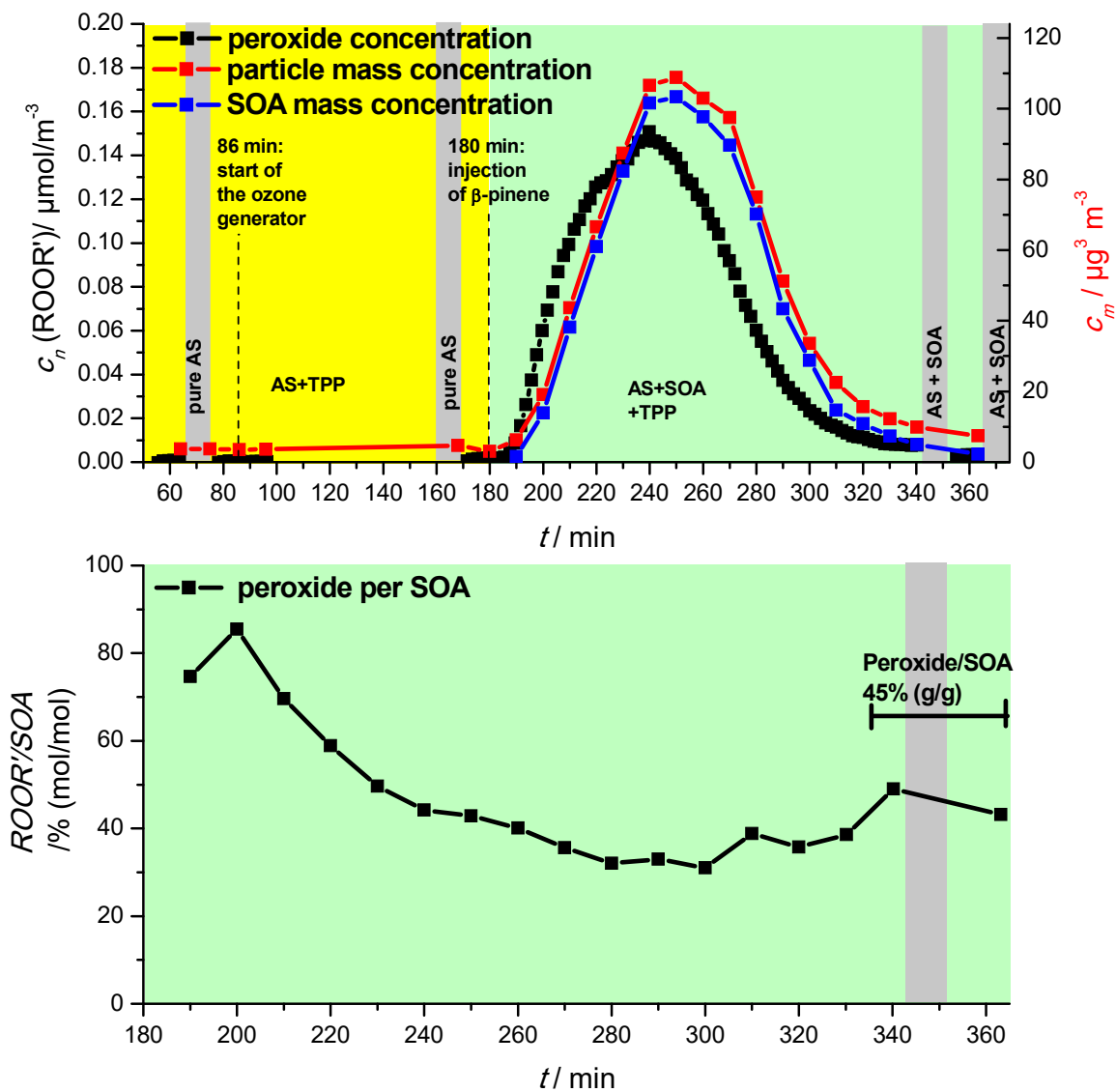


Figure 5: Size distribution of individual particle phase compounds (TPPO, TPP, SOA, sulfate) within the first 10 minutes after the addition of ozone.



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Figure 6: Size distribution of *m/z* 277 (black = TPPO), *m/z* 262 (blue = residual TPP), *m/z* 43 (green = SOA), sulfate (red) for following periods: Before the start of the ozonolysis a) without and b) with adding TPP, at the end (185-265 min) of the ozonolysis c) without and d) with adding TPP. All distributions were normalized to their respective maximum.



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Figure 7: Ozonolysis of β -pinene in the presence of AS aerosol. Time series of a) peroxide (black), particle mass concentration (red), SOA mass concentration (blue) and b) relative contribution of peroxide to the amount of SOA. Ozone was added at 86 min, β -pinene at 180 min.

405 Averages molar masses of 250 g mol^{-1} and 300 g mol^{-1} for SOA and peroxides were assumed, respectively.

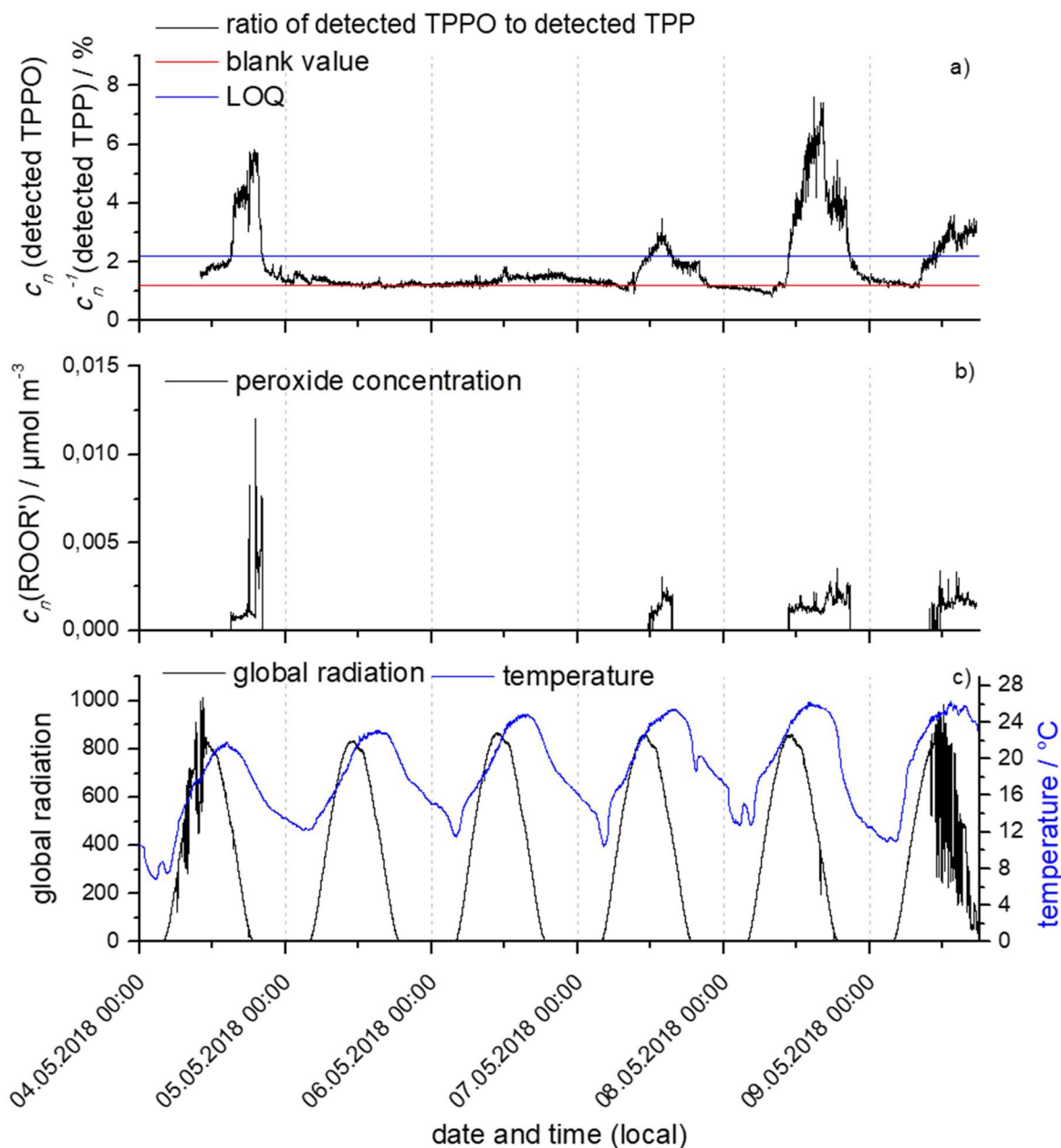


Figure 8: a) Time series of the relative amount of detected TPPO to detected TPP, LOQ and blank value, b) diurnal variation of the estimated peroxide concentration and c) time series of the global radiation and temperature.

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References

- Allan, J. D., Alfarra, Bower, K. N., Williams, P. I., Gallagher, M. W., Jimenez, J. L., McDonald, A. G., Nemitz, E., Canagaratna, Jayne, J. T., Coe, H., and Worsnop: Quantitative sampling using an Aerodyne aerosol mass spectrometer - 2. Measurements of fine particulate chemical composition in two U.K. cities, *Journal of Geophysical Research-Atmospheres*, 108, doi:10.1029/2002JD002359, 2003a.
- Allan, J. D., Jimenez, J. L., Williams, P. I., Alfarra, Bower, K. N., Jayne, J. T., Coe, H., and Worsnop: Quantitative sampling using an Aerodyne aerosol mass spectrometer - 1. Techniques of data interpretation and error analysis, *Journal of Geophysical Research-Atmospheres*, 108, doi:10.1029/2002JD002358, 2003b.
- Bianchi, F., Garmash, O., He, X., Yan, C., Iyer, S., Rosendahl, I., Xu, Z., Rissanen, M. P., Riva, M., Taipale, R., Sarnela, N., Petaja, T., Worsnop, D. R., Kulmala, M., Ehn, M., and Junninen, H.: The role of highly oxygenated molecules (HOMs)

420

- in determining the composition of ambient ions in the boreal forest, *Atmos. Chem. Phys.*, 17, 13819–13831, doi:10.5194/acp-17-13819-2017, 2017.
- 425 Bianchi, F., Kurten, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crouse, J. D., Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R., Thornton, J. A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly Oxygenated Organic Molecules (HOM) from Gas-Phase Autoxidation Involving Peroxy Radicals: A Key Contributor to Atmospheric Aerosol, *Chemical Reviews*, 119, 3472–3509, doi:10.1021/acs.chemrev.8b00395, 2019.
- 430 Bianchi, F., Trostl, J., Junninen, H., Frege, C., Henne, S., Hoyle, C. R., Molteni, U., Herrmann, E., Adamov, A., Bukowiecki, N., Chen, X., Duplissy, J., Gysel, M., Hutterli, M., Kangasluoma, J., Kontkanen, J., Kuerten, A., Manninen, H. E., Muench, S., Perakyla, O., Petaja, T., Rondo, L., Williamson, C., Weingartner, E., Curtius, J., Worsnop, D. R., Kulmala, M., Dommen, J., and Baltensperger, U.: New particle formation in the free troposphere: A question of chemistry and timing, *Science*, 352, 1109–1112, doi:10.1126/science.aad5456, 2016.
- Bonn, B., Kuhlmann, R. von, and Lawrence, M. G.: High contribution of biogenic hydroperoxides to secondary organic aerosol formation, *Geophys. Res. Lett.*, 31, doi:10.1029/2003GL019172, 2004.
- 435 Brueggemann, M., Karu, E., and Hoffmann, T.: Critical assessment of ionization patterns and applications of ambient desorption/ionization mass spectrometry using FAPA-MS, *Journal of Mass Spectrometry*, 51, 141–149, doi:10.1002/jms.3733, 2016.
- Brueggemann, M., Karu, E., Stelzer, T., and Hoffmann, T.: Real-Time Analysis of Ambient Organic Aerosols Using Aerosol Flowing Atmospheric-Pressure Afterglow Mass Spectrometry (AeroFAPA-MS), *Environ. Sci. Technol.*, 49, 5571–5578, doi:10.1021/es506186c, 2015.
- 440 Brueggemann, M., Poulain, L., Held, A., Stelzer, T., Zuth, C., Richters, S., Mutzel, A., van Pinxteren, D., Iinuma, Y., Katkevica, S., Rabe, R., Herrmann, H., and Hoffmann, T.: Real-time detection of highly oxidized organosulfates and BSOA marker compounds during the F-BEACH 2014 field study, *Atmos. Chem. Phys.*, 17, 1453–1469, doi:10.5194/acp-17-1453-2017, 2017.
- 445 Brüggemann, M., Vogel, L. A., and Hoffmann, T.: Analysis of organic aerosols using a micro-orifice volatilization impactor (MOVI) coupled to an atmospheric-pressure chemical ionization mass spectrometer (APCI-MS), *European Journal of Mass Spectrometry*, 20, 31–41, doi:10.1255/ejms.1260, 2014.
- Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Chemical and microphysical characterization of ambient aerosols with the aerodyne aerosol mass spectrometer, *Mass Spectrometry Reviews*, 26, 185–222, doi:10.1002/mas.20115, 2007.
- Chellamani, A. and Suresh, R.: Kinetics and mechanism of oxidation of triphenylphosphine by hydrogen peroxide, *React. Kinet. Catal. Lett.*, 37, 501–505, 1988.
- 455 Claflin, M. S., Krechmer, J. E., Hu, W., Jimenez, J. L., and Ziemann, P. J.: Functional Group Composition of Secondary Organic Aerosol Formed from Ozonolysis of alpha-Pinene Under High VOC and Autoxidation Conditions, *Acs Earth and Space Chemistry*, 2, 1196–1210, doi:10.1021/acsearthspacechem.8b00117, 2018.
- Davies, A. G. and Feld, R.: 934. Organic peroxides. Part IX. A stereochemical study of the formation of alkyl hydroperoxides from hydrogen peroxide, *J. Chem. Soc., (Journal of the Chemical Society)*, 4637–4643, 1958.
- 460 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.: Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, *Analytical chemistry*, 78, 8281–8289, doi:10.1021/ac061249n, 2006.

- DeCarlo, P. F., Slowik, J. G., Worsnop, D. R., Davidovits, P., and Jimenez, J. L.: Particle Morphology and Density Characterization by Combined Mobility and Aerodynamic Diameter Measurements. Part 1: Theory, *Aerosol Science and Technology*, 38, 1185–1205, doi:10.1080/027868290903907, 2004.
- 465 Docherty, K. S., Kumboonlert, K., Lee, I. J., and Ziemann, P. J.: Gas chromatography of trimethylsilyl derivatives of α -methoxyalkyl hydroperoxides formed in alkene–O₃ reactions, *Journal of Chromatography A*, 1029, 205–215, doi:10.1016/j.chroma.2003.12.014, 2004.
- Docherty, K. S., Wu, W., Lim, Y. B., and Ziemann, P. J.: Contributions of Organic Peroxides to Secondary Aerosol Formed from Reactions of Monoterpenes with O₃, *Environ. Sci. Technol.*, 39, 4049–4059, doi:10.1021/es050228s, 2005.
- 470 Donahue, N. M., Robinson, A. L., Trump, E. R., Riipinen, I., and Kroll, J. H.: Volatility and Aging of Atmospheric Organic Aerosol, *ATMOSPHERIC AND AEROSOL CHEMISTRY*, 339, 97–143, doi:10.1007/128_2012_355, 2014.
- Epstein, S. A., Blair, S. L., and Nizkorodov, S. A.: Direct Photolysis of α -Pinene Ozonolysis Secondary Organic Aerosol: Effect on Particle Mass and Peroxide Content, *Environ. Sci. Technol.*, 48, 11251–11258, doi:10.1021/es502350u, 2014.
- 475 Fuller, S. J., Wragg, F. P. H., Nutter, J., and Kalberer, M.: Comparison of on-line and off-line methods to quantify reactive oxygen species (ROS) in atmospheric aerosols, *Atmos. Environ.*, 92, 97–103, doi:10.1016/j.atmosenv.2014.04.006, 2014.
- Gallimore, P. J., Giorio, C., Mahon, B. M., and Kalberer, M.: Online molecular characterisation of organic aerosols in an atmospheric chamber using extractive electrospray ionisation mass spectrometry, *Atmos. Chem. Phys.*, 17, 14485–14500, doi:10.5194/acp-17-14485-2017, 2017.
- 480 Gotoh, N., Miyake, S., Takei, H., Sasaki, K., Okuda, S., Ishinaga, M., and Wada, S.: Simple Method for Measuring the Peroxide Value in a Colored Lipid, *Food Anal. Methods*, 4, 525–530, doi:10.1007/s12161-011-9193-5, 2011.
- Guo, J., Tilgner, A., Yeung, C., Wang, Z., Louie, P. K. K., Luk, C. W. Y., Xu, Z., Yuan, C., Gao, Y., Poon, S., Herrmann, H., Lee, S., Lam, K. S., and Wang, T.: Atmospheric peroxides in a polluted subtropical environment: Seasonal variation, sources and sinks, and importance of heterogeneous processes, *Environmental science & technology*, 48, 1443–1450, doi:10.1021/es403229x, 2014.
- 485 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., and Goldstein, A. H.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155–5236, doi:10.5194/acp-9-5155-2009, 2009.
- Hua, W., Chen, Z. M., Jie, C. Y., Kondo, Y., Hofzumahaus, A., Takegawa, N., Lu, K. D., Miyazaki, Y., Kita, K., Wang, H. L., Zhang, Y. H., and Hu, M.: Atmospheric hydrogen peroxide and organic hydroperoxides during PRIDE-PRD'06, China: their concentration, formation mechanism and contribution to secondary aerosols, *Atmos. Chem. Phys. Discuss.*, 8, 10481–10530, doi:10.5194/acpd-8-10481-2008, 2008.
- 490 Hyde, A. M., Zultanski, S. L., Waldman, J. H., Zhong, Y.-L., Shevlin, M., and Peng, F.: General Principles and Strategies for Salting-Out Informed by the Hofmeister Series, *Org. Process Res. Dev.*, 21, 1355–1370, doi:10.1021/acs.oprd.7b00197, 2017.
- 495 Jarvie, A. W.P., Moore, C. G., and Skelton, D.: Reactions of some organic peroxides with triphenylphosphine and sodium dialkyl phosphites, *Journal of Polymer Science: Part A-1*, 9, 3105–3114, 1971.
- Jayanthi, S. and Koshore, K.: Reaction of polyperoxides with triphenylphosphine, *J. Polym. Sci. Pol. Phys (Journal of Polymer Science A: Polymer Chemistry)*, 35, 1167–1172, 1997.
- 500 Jayne, J. T., Leard, D. C., Zhang, X. F., Davidovits, P., Smith, K. A., Kolb, C. E., and Worsnop, D. R.: Development of an aerosol mass spectrometer for size and composition analysis of submicron particles, *Aerosol Sci. Technol.*, 33, 49–70, doi:10.1080/027868200410840, 2000.
- Kahnt, A., Vermeulen, R., Iinuma, Y., Shalamzari, M. S., Maenhaut, W., and Claeys, M.: High-molecular-weight esters in α -pinene ozonolysis secondary organic aerosol: Structural characterization and mechanistic proposal for their

- 505 formation from highly oxygenated molecules, *Atmos. Chem. Phys.*, 18, 8453–8467, doi:10.5194/acp-18-8453-2018, 2018.
- Koop, T., Bookhold, J., Shiraiwa, M., and Pöschl, U.: Glass transition and phase state of organic compounds: Dependency on molecular properties and implications for secondary organic aerosols in the atmosphere, *Physical chemistry chemical physics PCCP*, 13, 19238–19255, doi:10.1039/c1cp22617g, 2011.
- 510 Krapf, M., El Haddad, I., Bruns, E. A., Molteni, U., Daellenbach, K. R., Prévôt, A. S.H., Baltensperger, U., and Dommen, J.: Labile Peroxides in Secondary Organic Aerosol, *Chem*, 1, 603–616, doi:10.1016/j.chempr.2016.09.007, 2016.
- Krawczyk, T. and Baj, S.: REVIEW: ADVANCES IN THE DETERMINATION OF PEROXIDES BY OPTICAL AND SPECTROSCOPIC METHODS, *Analytical Letters*, 47, 2129–2147, doi:10.1080/00032719.2014.900781, 2014.
- Lee, S.-H., Gordon, H., Yu, H., Lehtipalo, K., Haley, R., Li, Y., and Zhang, R.: New Particle Formation in the Atmosphere: From Molecular Clusters to Global Climate, *Journal of Geophysical Research-Atmospheres*, 124, 7098–7146, doi:10.1029/2018JD029356, 2019.
- 515 Lelieveld, J., Butler, T. M., Crowley, J. N., Dillon, T. J., Fischer, H., Ganzeveld, L., Harder, H., Lawrence, M. G., Martinez, M., Taraborrelli, D., and Williams, J.: Atmospheric oxidation capacity sustained by a tropical forest, *Nature*, 452, 737–740, doi:10.1038/nature06870, 2008.
- 520 Matthew, B. M., Middlebrook, A. M., and Onasch, T. B.: Collection Efficiencies in an Aerodyne Aerosol Mass Spectrometer as a Function of Particle Phase for Laboratory Generated Aerosols, *Aerosol Science and Technology*, 42, 884–898, doi:10.1080/02786820802356797, 2008.
- McNeill, V. F.: Atmospheric Aerosols: Clouds, Chemistry, and Climate, *Annual Review of Chemical and Biomolecular Engineering*, Vol 8, 8, 427–444, doi:10.1146/annurev-chembioeng-060816-101538, 2017.
- 525 Mertes, P., Pfaffenberger, L., Dommen, J., Kalberer, M., and Baltensperger, U.: Development of a sensitive long path absorption photometer to quantify peroxides in aerosol particles (Peroxide-LOPAP), *Atmospheric Measurement Techniques*, 5, 2339–2348, doi:10.5194/amt-5-2339-2012, 2012.
- Mochida, M., Katrib, Y., Jayne, J. T., Worsnop, D. R., and Martin, S. T.: The relative importance of competing pathways for the formation of high-molecular-weight peroxides in the ozonolysis of organic aerosol particles, *Atmos. Chem. Phys.*, 6, 4851–4866, doi:10.5194/acp-6-4851-2006, 2006.
- 530 Mutzel, A., Rodigast, M., Iinuma, Y., Böge, O., and Herrmann, H.: An improved method for the quantification of SOA bound peroxides, *Atmospheric Environment*, 67, 365–369, doi:10.1016/j.atmosenv.2012.11.012, 2013.
- Nakamura, T. and Maeda, H.: A simple assay for lipid hydroperoxides based on triphenylphosphine oxidation and high-performance liquid chromatography, *Lipids*, 26, 765–768, doi:10.1007/BF02535628, 1991.
- 535 Porter, N. A., Logan, J., and Kontoyiannidou, V.: Preparation and purification of arachidonic acid hydroperoxides of biological importance, *J.Org.Chem*, 3177-3181, 1979.
- Poschl, U.: Atmospheric aerosols: Composition, transformation, climate and health effects, *Angewandte Chemie-International Edition*, 44, 7520, 2005.
- Pratt, K. A. and Prather, K. A.: Mass spectrometry of atmospheric aerosols: Recent developments and applications. Part I: Off-line mass spectrometry techniques, *Mass Spectrometry Reviews*, 31, 1–16, doi:10.1002/mas.20322, 2012.
- 540 Reinnig, M.-C., Müller, L., Warnke, J., and Hoffmann, T.: Characterization of selected organic compound classes in secondary organic aerosol from biogenic VOCs by HPLC/MSn, *Analytical and bioanalytical chemistry*, 391, 171–182, doi:10.1007/s00216-008-1964-5, 2008.
- Reinnig, M.-C., Warnke, J., and Hoffmann, T.: Identification of organic hydroperoxides and hydroperoxy acids in secondary organic aerosol formed during the ozonolysis of different monoterpenes and sesquiterpenes by on-line analysis using atmospheric pressure chemical ionization ion trap mass spectrometry, *Rapid communications in mass spectrometry RCM*, 23, 1735–1741, doi:10.1002/rcm.4065, 2009.
- 545

- Riipinen, I., Pierce, J. R., Yli-Juuti, T., Nieminen, T., Hakkinen, S., Ehn, M., Junninen, H., Lehtipalo, K., Petaja, T., Slowik, J., Chang, R., Shantz, N. C., Abbatt, J., Leaitch, W. R., Kerminen, V.-M., Worsnop, D. R., Pandis, S. N., Donahue, N. M., and Kulmala, M.: Organic condensation: A vital link connecting aerosol formation to cloud condensation nuclei (CCN) concentrations, *Atmos. Chem. Phys.*, 11, 3865–3878, doi:10.5194/acp-11-3865-2011, 2011.
- 550 Ruiz, A., Canada, M. J.A., and Lendl, B.: A rapid method for peroxide value determination in edible oils based on flow analysis with Fourier transform infrared spectroscopic detection, *Analyst*, 126, 242–246, doi:10.1039/b008688f, 2001.
- Steimer, S. S., Kourtchev, I., and Kalberer, M.: Mass Spectrometry Characterization of Peroxycarboxylic Acids as Proxies for Reactive Oxygen Species and Highly Oxygenated Molecules in Atmospheric Aerosols, *Anal. Chem.*, 89, 2873–2879, doi:10.1021/acs.analchem.6b04127, 2017.
- 555 Stein, R. and Slawson, V.: Spectrophotometric hydroperoxide determination by the use of triphenylphosphine, *Anal. Chem.*, 35, 1008–&, doi:10.1021/ac60201a027, 1963.
- Thompson, A.M.: The Oxidizing Capacity of the Earth's Atmosphere - Probable Past and Future Changes, *Science*, 256, 560 1157–1165, 1992.
- Thorenz, U. R., Kundel, M., Müller, L., and Hoffmann, T.: Generation of standard gas mixtures of halogenated, aliphatic, and aromatic compounds and prediction of the individual output rates based on molecular formula and boiling point, *Analytical and bioanalytical chemistry*, 404, 2177–2183, doi:10.1007/s00216-012-6202-5, 2012.
- Tobias, H. J., Docherty, K. S., Beving, D. E., and Ziemann, P. J.: Effect of relative humidity on the chemical composition of secondary organic aerosol formed from reactions of 1-tetradecene and O₃, *Environ. Sci. Technol.*, 34, 2116–2125, doi:10.1021/es991057s, 2000.
- 565 Tobias, H. J. and Ziemann, P. J.: Kinetics of the gas-phase reactions of alcohols, aldehydes, carboxylic acids, and water with the C₁₃ stabilized Criegee intermediate formed from ozonolysis of 1-tetradecene, *J. Phys. Chem. A*, 105, 6129–6135, 2002.
- 570 Tong, H., Lakey, P. S. J., Arangio, A. M., Socorro, J., Shen, F., Lucas, K., Brune, W. H., Pöschl, U., and Shiraiwa, M.: Reactive Oxygen Species Formed by Secondary Organic Aerosols in Water and Surrogate Lung Fluid, *Environmental science & technology*, 52, 11642–11651, doi:10.1021/acs.est.8b03695, 2018.
- Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C., Bianchi, F., Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J., Adamov, A., Almeida, J., Bernhammer, A.-K., 575 Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, A., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lawler, M., Leiminger, M., Mathot, S., Möhler, O., Nieminen, T., Onnela, A., Petäjä, T., Piel, F. M., Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Sipilä, M., Smith, J. N., Steiner, G., Tomè, A., Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen, 580 J., Kirkby, J., Kulmala, M., Riipinen, I., Worsnop, D. R., Donahue, N. M., and Baltensperger, U.: The role of low-volatility organic compounds in initial particle growth in the atmosphere, *Nature*, 533, 527–531, doi:10.1038/nature18271, 2016.
- Verma, V., Fang, T., Xu, L., Peltier, R. E., Russell, A. G., Ng, N. L., and Weber, R. J.: Organic aerosols associated with the generation of reactive oxygen species (ROS) by water-soluble PM_{2.5}, *Environmental science & technology*, 49, 4646–4656, doi:10.1021/es505577w, 2015.
- 585 Vogel, A. L., Aijala, M., Brüeggemann, M., Ehn, M., Junninen, H., Petaja, T., Worsnop, D. R., Kulmala, M., Williams, J., and Hoffmann, T.: Online atmospheric pressure chemical ionization ion trap mass spectrometry (APCI-IT-MSn) for measuring organic acids concentrated bulk aerosol - a laboratory and field study, *Atmospheric Measurement Techniques*, 6, 431–443, doi:10.5194/amt-6-431-2013, 2013a.

- 590 Vogel, A. L., Aijala, M., Corrigan, A. L., Junninen, H., Ehn, M., Petaja, T., Worsnop, D. R., Kulmala, M., Russell, L. M., Williams, J., and Hoffmann, T.: In situ submicron organic aerosol characterization at a boreal forest research station during HUMPPA-COPEC 2010 using soft and hard ionization mass spectrometry, *Atmos. Chem. Phys.*, 13, 10933–10950, doi:10.5194/acp-13-10933-2013, 2013b.
- Wang, N., Ma, T., Yu, X., Xu, L., and Zhang, R.: Determination of Peroxide Values of Edible Oils by Ultraviolet
595 Spectrometric Method, *Food Anal. Methods*, 9, 1412–1417, doi:10.1007/s12161-015-0322-4, 2016.
- Warscheid, B. and Hoffmann, T.: On-line measurements of α -pinene ozonolysis products using an atmospheric pressure chemical ionisation ion-trap mass spectrometer, *Atmos. Environ.*, 35, 2927–2940, doi:10.1016/S1352-2310(00)00513-6, 2001a.
- Warscheid, B. and Hoffmann, T.: Structural elucidation of monoterpene oxidation products by ion trap fragmentation using
600 on-line atmospheric pressure chemical ionisation mass spectrometry in the negative ion mode, *Rapid communications in mass spectrometry RCM*, 15, 2259–2272, doi:10.1002/rcm.504, 2001b.
- Williams, D. H., Ward, R. S., and Cooks, R. G.: Mass spectrometry. XXIV. A study of the reactions induced in triphenylphosphine, triphenylphosphine oxide, and related substances upon electron impact, *J. Am. Chem. Soc.*, 90, 966–972, doi:10.1021/ja01006a021, 1968.
- 605 Wragg, F. P. H., Fuller, S. J., Freshwater, R., Green, D. C., Kelly, F. J., and Kalberer, M.: An automated online instrument to quantify aerosol-bound reactive oxygen species (ROS) for ambient measurement and health-relevant aerosol studies, *Atmos. Meas. Tech.*, 9, 4891–4900, doi:10.5194/amt-9-4891-2016, 2016.
- Yatawelli, R. L. N., Mohr, C., Stark, H., Day, D. A., Thompson, S. L., Lopez-Hilfiker, F. D., Campuzano-Jost, P., Palm, B., B., Vogel, A. L., Hoffmann, T., Heikkinen, L., Aijala, M., Ng, N. L., Kimmel, J. R., Canagaratna, M.I.R., Ehn, M.,
610 Junninen, H., Cubison, M. J., Petaja, T., Kulmala, M., Jayne, J. T., Worsnop, D. R., and Jimenez, J. L.: Estimating the contribution of organic acids to northern hemispheric continental organic aerosol, *Geophys. Res. Lett.*, 42, 6084–6090, doi:10.1002/2015GL064650, 2015.
- Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., and Sun, Y.: Understanding atmospheric organic aerosols via factor analysis of aerosol mass spectrometry: a review, *Anal. Bioanal. Chem.*, 401,
615 3045–3067, doi:10.1007/s00216-011-5355-y, 2011.
- Zhang, X., Lambe, A. T., Upshur, M. A., Brooks, W. A., Be, A. G., Thomson, R. J., Geiger, F. M., Surratt, J. D., Zhang, Z., Gold, A., Graf, S., Cubison, M. J., Groessl, M., Jayne, J. T., Worsnop, D. R., and Canagaratna, M. R.: Highly Oxygenated Multifunctional Compounds in α -Pinene Secondary Organic Aerosol, *Environ. Sci. Technol.*, 51, 5932–5940, doi:10.1021/acs.est.6b06588, 2017.
- 620 Zhao, R., Kenseth, C. M., Huang, Y., Dalleska, N. F., and Seinfeld, J. H.: Iodometry-Assisted Liquid Chromatography Electro-spray Ionization Mass Spectrometry for Analysis of Organic Peroxides: An Application to Atmospheric Secondary Organic Aerosol, *Environmental science & technology*, 52, 2108–2117, doi:10.1021/acs.est.7b04863, 2018.
- Zhou, S., Rivera-Rios, J. C., Keutsch, F. N., and Abbatt, J. P. D.: Identification of organic hydroperoxides and peroxy acids using atmospheric pressure chemical ionization-tandem mass spectrometry (APCI-MS/MS): Application to secondary
625 organic aerosol, *Atmos. Meas. Tech.*, 11, 3081–3089, doi:10.5194/amt-11-3081-2018, 2018.
- Ziemann, P. J.: Evidence for low-volatility diacyl peroxides as a nucleating agent and major component of aerosol formed from reactions of O₃ with cyclohexene and homologous compounds, *J. Phys. Chem. A*, 106, 4390–4402, 2002.
- Ziemann, P. J.: Formation of alkoxyhydroperoxy aldehydes and cyclic peroxyhemiacetals from reactions of cyclic alkenes with O₃ in the presence of alcohols, *Journal of Physical Chemistry A*, 107, 2048–2060, 2003.
- 630 Ziemann, P. J. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic aerosol formation, *Chem. Soc. Rev.*, 41, 6582–6605, doi:10.1039/c2cs35122f, 2012.