Resolving the size of ice-nucleating particles with a balloon deployable aerosol sampler: the SHARK

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Abstract. Ice-nucleating particles (INPs) affect cloud development, lifetime and radiative properties, hence it is 11 important to know the abundance of INPs throughout the atmosphere. A critical factor in determining the lifetime 12 13 and transport of INPs is their size, however very little size-resolved atmospheric INP concentration information 14 exists. This is especially so in the free troposphere. Here we present the development and application of a radio-15 controlled payload capable of collecting size-resolved aerosol from a tethered balloon for the primary purpose of 16 offline INP analysis. This payload, known as the SHARK (Selective Height Aerosol Research Kit), consists of 17 two complementary cascade impactors for aerosol size-segregation from 0.25 to 10 µm, with an after-filter and 18 top stage to collect particles below and above this range at flow rates up to 100 L min⁻¹. The SHARK also contains 19 an optical particle counter to quantify aerosol size distribution between 0.38 and 10 µm, and a radiosonde for the measurement of temperature, pressure, GPS altitude, and relative humidity. This is all housed within a 20 21 weatherproof box, can be run from batteries for up to 11 h and has a total weight of 9 kg. The radio control and 22 live data link with the radiosonde allow the user to start and stop sampling depending on meteorological conditions 23 and height, which can, for example, allow the user to avoid sampling in very humid or cloudy air, even when the 24 SHARK is out of sight. While the collected aerosol could, in principle, be studied with an array of analytical 25 techniques, this study demonstrates that the collected aerosol can be analysed with an off-line droplet freezing 26 instrument to determine size-resolved INP concentrations, activated fractions and active site densities, producing 27 similar results to those obtained using a standard PM₁₀ aerosol sampler when summed over the appropriate size range. Test data is presented from four contrasting locations having very different size resolved INP spectra: 28 29 Hyytiälä (Southern Finland), Leeds (Northern England), Longyearbyen (Svalbard), and Cardington (Southern England). 30

31 1 Introduction

Atmospheric ice-nucleating particles (INPs) are not well understood, with knowledge of their concentration, sources, temporal variability, transport and size in its infancy (Kanji et al., 2017; Murray et al., 2012). This is of importance because clouds between 0 °C and around -35 °C can exist in a supercooled liquid, mixed-phase (ice and water) or glaciated (ice only) state depending in part on the presence or absence of INPs (Kanitz et al., 2011; Vergara-Temprado et al., 2018). In the absence of INP, cloud droplets can supercool to below ~-35°C (Herbert et al., 2015), but INP can trigger freezing at much higher temperatures (Kanji et al., 2017). These particles usually

38 have concentrations that are orders of magnitude smaller than cloud condensation nuclei (CCN), and have a 39 disproportionate impact on clouds because the nucleated ice crystals grow rapidly and precipitate out (Lohmann, 2017; Murray, 2017). In a shallow cloud, heterogeneous ice nucleation can result in dramatic reductions in cloud 40 41 albedo by removal of supercooled liquid water (Storelvmo, 2017; Vergara-Temprado et al., 2018), whereas in deep convective clouds it can influence a web of microphysical processes in a complex way (Deng et al., 2018; 42 43 Kanji et al., 2017; Rosenfeld et al., 2011). Hence, a greater understanding of INP lifetime, transport and 44 distribution in the vertical profile is needed in order to better understand and model cloud processes and their 45 response to a changing climate.

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The size of an aerosol particle significantly affects its lifetime and therefore transport in the atmosphere, with 47 particles of a few hundred nanometres having a lifetime of weeks in the free troposphere, whereas particles of 10 48 49 µm have a lifetime of only hours (Jaenicke, 2007). While composition is recognised to be an important controller 50 of ice nucleation ability (Kanji et al., 2017), it has also been generally thought that the larger an aerosol particle, 51 the more likely it is to serve as an INP (Pruppacher and Klett, 1997). However, the lifetime of coarse mode aerosol 52 particles decreases rapidly with increasing size. Consistent with larger particles being better ice nucleators, 53 parameterisations of INPs in the atmosphere have been proposed wherein the INP concentration is related to the 54 concentration of aerosol particles larger than 0.5 µm (DeMott et al., 2010, 2015; Tobo et al., 2013). However, most atmospheric measurements of INPs report the sum of INPs below some threshold size set by an inlet or size 55 56 cut, specified by the aerosol sampler used. For instance, DeMott et al. (2017) provides a comparison between a 57 selection of instruments for the collection and subsequent INP analysis of aerosol, where the aerosol samplers 58 have either a defined size cut-off or have collection efficiencies that decrease in magnitude above a defined size. Nevertheless, there are examples of field studies in which INPs have been size-resolved (Berezinski et al., 1988; 59 60 Creamean et al., 2018b; Huffman et al., 2013; Mason et al., 2016; Reicher et al., 2018; Santachiara et al., 2010; 61 Si et al., 2018; Welti et al., 2009). These studies generally show that while the fine mode aerosol particles are 62 more abundant, coarse mode aerosol particles often contribute more to the INP population. In addition, the 63 activated fraction (n_n) of coarse mode aerosol is usually greater than fine mode aerosol. However, in some field 64 studies (Mason et al., 2016; Si et al., 2018), fine aerosol sometimes contributes more to the INP population than 65 the coarse mode. Therefore, there is a need to determine INP sizes when quantifying atmospheric INP 66 concentrations, as size is important for transport and lifetime and is therefore required to accurately model global 67 INP populations.

69 Measurements of INPs in and above the boundary layer are crucial to understanding the contribution of local 70 sources to the ice-nucleating activity in clouds, compared to transported aerosol. Aircraft measurements (e.g. Price 71 et al., 2018; Rogers et al., 2001) and mountaintop observatories (e.g. Conen et al., 2015) have been used to quantify 72 INP populations above the boundary layer. For example, it has been shown that there are differences in the INP 73 concentrations measured when in and out of the boundary layer at the High Altitude Research Station Jungfraujoch 74 (Switzerland) (Conen et al., 2015; Lacher et al., 2018). While these measurements are undoubtedly useful, 75 mountaintop measurements are only possible in locations with sufficiently tall yet accessible mountains, and 76 aircraft sampling is expensive and not necessarily possible in remote regions. It is therefore essential that 77 instrumentation is available that can be used to sample aerosol at selected altitudes (including ground level) in

order to determine INP concentrations throughout the vertical profile. Unmanned aerial vehicles (UAVs) are becoming more widely used in atmospheric science; these allow the collection of aerosol at altitude at significantly lower cost than with manned aircraft, but are <u>limited by relatively short battery lives</u>, <u>usually under 1 h</u>, and <u>potential</u>, propeller interference (Jacob et al., 2018; Villa et al., 2016).

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83 Tethered kite and balloon systems have historically been used to make atmospheric measurements and collect 84 aerosol samples with much longer sampling times (many hours are readily achievable) at altitudes up to 2 km and 85 5 km for tethered balloons and kites respectively (Armstrong et al., 1981; Balsley et al., 1998). An advantage of 86 a balloon or kite system is that an instrument can be held at a chosen altitude for many hours without the balloon interfering with measurements, as the instrument can be suspended on a line many meters below the balloon. They 87 can also stay inflated and in use for periods of many weeks, making them ideal for longer campaigns in remote 88 environments. A new instrument called the Honing On VERtical Cloud and Aerosol properTies (HOVERCAT) 89 90 (Creamean et al., 2018a) provides the capability to sample aerosol for subsequent INP analysis on a tethered 91 balloon or UAV, allowing both variable altitudes and static collection of non-size resolved aerosol smaller than 92 $10 \,\mu\text{m}$ at $1.2 \,L \,\text{min}^{-1}$. In the past, aerosols have been size-segregated using cascade impactors on a tethered balloon system (Hara et al., 2013; Reagan et al., 1984), but balloon-borne cascade impactor systems have not yet been 93 94 adapted for the purpose of size-resolved INP analysis. The downsides of balloon-based platforms include the need for wind speeds below around 64.4 km h⁻¹ to avoid damage to the balloon, and the possibility of 'icing' of the 95 96 balloon and lines when deployed in a cold and humid environment, which could add to the weight of the payload and cause the system to sink, or fall slowly. Nevertheless, balloon and kite-borne measurements remain a valuable 97 98 way to obtain continuous, high resolution measurements over a period of many hours in a single location at a 99 range of altitudes.

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101 In this paper, the design, testing and operation of a payload named the Selective Height Aerosol Research Kit 102 (SHARK) is presented. It consists of two separate cascade impactor systems, operating at 9 and 100 L min⁻¹, for 103 the size-sorting of ambient aerosol particles from 0.25 to 10 μ m, with an after-filter and top stage to collect particles below and above this range for offline INP (or other) analysis. The SHARK also features an optical 104 particle counter (OPC) and a radiosonde, which provides real-time measurements of relative humidity (RH), 105 temperature, Global Positioning System (GPS) altitude and pressure. Weighing 9 kg, the payload is suitable for 106 107 use with a 21 m³ or larger tethered balloon such as in Figure 1 a where the SHARK is shown in-flight. The use of 108 a tethered balloon and a high-capacity battery allow aerosol to be collected for up to 11 h at a user-selected altitude.

109 2 The design and development of the SHARK

110 2.1 Instrument description

The SHARK, shown in Figure 1, comprises two cascade impactors and corresponding pumps, alongside an OPC (OPC-N2, Alphasense, UK) and radiosonde (S1H2-R, Windsond, Sweden), all mounted within a weatherproof enclosure with a tail fin to orient it into the wind. A photograph of the internal components of the SHARK are shown in Figure 1b. The two cascade impactors were employed to collect particles across different size bins:

115 Impactor 1 from 0.25-2.5 μm and Impactor 2 from 1-10 μm. Impactor 1 is a cascade impactor (U.S. Patent No.

Deleted: limited by relatively short battery lives of 10s of minutes and potential

118 6,786,105, Sioutas, SKC, UK), which requires a flow rate of 9 L min⁻¹ and operates with a portable pump (Leland Legacy, SKC, UK). Impactor 2 is also a cascade impactor (MSP Model 128, TSI, USA), which requires a flow 119 120 rate of 100 L min⁻¹ at a pressure drop of 0.6 kPa (Marple et al., 1991; Misra et al., 2002), and for which a radial 121 flow impeller (Radial Blower U51, Micronel, UK) was used in reverse as a lightweight pump (~120 g). These 122 pumps maintain the volumetric flow rate through the impactors as temperature and atmospheric pressure change 123 with altitude. The pump for Impactor 1 was calibrated to apply this adjustment to at least 2.3 km (Leland Legacy 124 Sample Pump: Operating Instructions, SKC), although the presence of the after-filter may reduce the battery life 125 at this altitude. The pump for Impactor 2 is supplied by a larger battery and should be able to maintain flow to at 126 least the same altitude as the Impactor 1 pump, and over a longer period of time. The SHARK records the volume 127 of air sampled through Impactor 1 during the flight, and so if the pump battery was depleted, or the pressure drop 128 became too great before Impactor 2 had finished sampling, the Impactor 1 pump would shut down and store the 129 recorded value for later analysis. Further testing of the SHARK would be required to define a maximum altitude 130 limit that each SHARK component could operate at. In order to provide RH, temperature, GPS altitude and 131 pressure data in real-time, the sensors and transmitter from a radiosonde were integrated into the system. The OPC 132 measured aerosol size distributions, which were saved in the on-board memory. Servo-controlled caps covered 133 the sample inlets and outlets to reduce contamination during ascent and descent, as well as to protect the 134 components from cloud water. The operation of the SHARK components was controlled remotely via a radio link 135 using an Arduino microcontroller board (16 km range); once the SHARK was at the desired altitude according to 136 the constantly transmitting radiosonde, the inlet caps opened 10 s prior to the pumps and OPC starting in order to initiate aerosol sampling and monitoring. The payload components, including the servo inlet covers and Arduino 137 control boards, were powered by a 5000 mAh battery (4S 14.8 V LiPo, Overlander, UK). The components were 138 139 assembled into the SHARK payload with the static (i.e. no wind) weight budget of 10 kg for a 21 m3 balloon 140 (Skyhook Helikite, Allsopp Helikites Ltd., UK) in mind, hence the SHARK weighs 9 kg when fully instrumented. 141

The cascade impactors allow for the collection of size-segregated aerosol (further details are provided in Section 2.2) onto thin films (0.25 mm thickness) for subsequent off-line analysis, which can be used alongside information about the aerosol size distributions obtained via the OPC and atmospheric conditions from the radiosonde. Our initial focus concerns the analysis of the ice-nucleating properties of the collected aerosol, but an array of analytical techniques could be applied to characterise the size-selected aerosol, including mass spectrometry, DNA analysis, scanning electron microscopy (SEM) and transmission electron microscopy (TEM) (Ault and Axson, 2017; Garcia et al., 2012; Huffman et al., 2013; Laskin et al., 2018).

149 2.2 Size-segregated collection of aerosol

Two separate cascade impactors were installed, each operating over different size ranges. This enabled sizeresolved aerosol sampling onto substrates across both the fine and coarse modes at high flow rates, while keeping power consumption low enough to be run from batteries. Single impactor systems designed to operate across the accumulation and coarse modes simultaneously require a relatively large pressure drop that would typically require a prohibitively large (and heavy) pump and battery for this application.

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156 Impactor 1 sorts aerosol into five size categories: <0.25 µm (this size bin is defined by the impactor after-filter and is hereafter referred to as 1a), 0.25-0.5 µm (from stage 1b), 0.5-1.0 µm (from stage 1c), 1.0-2.5 µm (from 157 stage 1*d*), and >2.5 μ m (from stage 1*e*). The size categories *b* to *e* correspond to the impactor stages where the 50 158 159 % collection cut-off diameter (d50) is the lower bound of each bin. The size bins and collection efficiencies for each impactor were digitised from data provided by the manufacturers, (Misra et al., 2002; Product Information 160 161 Sheet - MSP) and are shown in Figure 2. Several collection substrates were tested by Misra et al. (2002), and the 162 dataset from the Teflon substrates was chosen to represent Impactor 1 here as that substrate most closely resembled 163 those used in this study. For Impactor 1, the particles were collected on 25 mm diameter filters of pore size 0.05 164 µm (Nuclepore Track-Etched Membrane polycarbonate filters, Whatman, UK). Filters were used as impactor substrates rather than films since they have very low background contamination and are easier to obtain. Size 165 category 1a corresponds to an after-filter situated after Impactor 1, which comprised a 47 mm diameter 166 polycarbonate filter with a pore size of 5 µm (Nuclepore Track-Etched Membrane) to maintain the flow rate. The 167 168 collection efficiency of the after-filter was estimated to be 50-100 % at 0.25 µm and below (Soo et al., 2016). 169 Impactor 2 collected aerosol particles into three size categories: 1.0-2.5 μ m (2*d*), 2.5-10 μ m (2*e*), and >10 μ m 170 (2f), also illustrated in Figure 2. 75 mm diameter filters of pore size 0.05 µm (Nuclepore Track-Etched Membrane polycarbonate filters) were used in Impactor 2. An after-filter could not be used with this impactor since its 171 172 inclusion increased the required pressure drop to beyond what the pump could supply at 100 L min⁻¹.

A further benefit of using these two impactors in tandem is that, in the size ranges where they overlap of 1.0-2.5 μ m (stage *d*) and 2.5-10 μ m (stage *e*), the impaction efficiencies are very similar, allowing a direct comparison between the two impactors in this size range. The stages are labelled *a* through *f* for the smallest to largest impactor stage sizes (including the after-filter), such that 1*d* and 2*d* refer to stage *d* (1.0-2.5 μ m) on Impactors 1 and 2, respectively (see Figure 2). Background runs were produced by placing the substrates in the SHARK as if setting up to sample, before removing and analysing them as normal to determine the contamination introduced through the installation and recovery of the substrates.

182 Particle bounce, the bouncing of particles off the impaction substrate and the collection of these particles on the 183 lower stages, has previously been identified as a factor that can cause biases when aerosol is collected by cascade 184 impactors (Cheng and Yeh, 1979; Dzubay et al., 1976). The collection efficiency curves shown in Figure 2 for 185 Impactor 1 already account for some degree of particle bounce, having been determined experimentally by Misra 186 et al., (2002) using monodispersed polymer particles on a variety of substrates. However, the efficiency curves 187 for Impactor 2 are based on theoretical predictions (Rader and Marple, 1985) and so do not account for any bounce 188 effects. Since two of the stages of Impactors 1 and 2 overlap (stages d and e), it is possible to comment on the 189 possible effects, or lack thereof, of particle bounce, based on the results obtained using each of the comparable 190 stages. This is briefly addressed in section 3.4 where we show good agreement between these two impactors.

191 2.3 Size distribution measurements

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192 The OPC produced binned particle size distributions from 0.38-17 µm every 1.<u>38 s at a typical flow rate of 1.2 L</u>

<u>min⁻¹</u>. The OPC was remotely operated through the use of its serial link via an Arduino microcontroller board.
 Particle size, surface area and mass concentration data were produced from the raw OPC data, and these then used

to calculate the fraction of the aerosol that act as an INP (activated fraction, $n_n(T)$), and to weight the INP data to particle surface area, generating the ice-active site density per surface area ($n_s(T)$) of aerosol. The particle density used was 1.65 g cm⁻³, as assumed by the OPC software, and they were assumed to be spherical. No correction was made for the hygroscopic growth of aerosol particles as this required assumptions about the chemical nature of the particles, and hygroscopic growth effects were minimised by avoiding sampling when the RH was above about 80 % (see next section).

201 2.4 Radiosonde data

Utilising the radio control built into the payload, real-time data informed decisions of when to turn the pumps on and off to sample. Continuous monitoring of the radiosonde data allows the user to avoid sampling under conditions where RH approached 100 %, at which point aerosol particles become excessively swollen with water or activated to cloud droplets. Hence, the influence of hygroscopic growth or cloud droplets on the collected aerosol could be minimised. The temperature and pressure measurements allowed the volume of air sampled by the impactors and OPC to be corrected to standard conditions (1 atm at 0 °C).

208 2.5 Housing and instrument orientation

The weatherproof housing consisted of an acrylonitrile butadiene styrene (ABS) polymer box with dimensions of 209 210 560 mm x 380 mm x 180 mm (IP67, Fibox). Holes to mount the impactors and OPCs were drilled so that Impactor 211 2 sat vertically upright and Impactor 1 was oriented 180° to Impactor 2 so that it faced downwards, ensuring that both impactors were always oriented 90° to the wind. The OPC was at 90° to both impactors and facing towards 212 213 the front of the box, into the wind (see Figure 3a-c) See section 2.6 for the rationale of the positioning of the OPC 214 and impactor inlets. The tail fin, which is mounted to the lid of the box, was designed to keep the SHARK 215 orientated into the wind, and was fabricated from rigid polyvinyl chloride (PVC) sheet. Impactor 1 had its own 216 mounting screws by which it was attached to the box, whilst for Impactor 2 a custom mount was built. Securing 217 ropes were threaded through reinforced holes in the box and connected via a carabiner for quick and easy attachment to the balloon instrument line, as seen in Figure 1a. Modular foam was used to keep all components 218 219 in place during flight.

220 2.6 Inlet sampling efficiencies via particle loss modelling

221 Calculation of the particle losses associated with the instrument inlets due to excessive wind speeds in various 222 configurations were used to inform the design of the SHARK and to minimise sampling biases in higher wind 223 conditions. The calculations were done using an open source particle loss calculator program in Igor Pro, the 224 details and assumptions for which are presented in Von Der Weiden et al., (2009). The particle loss characteristics 225 of the impactor and OPC inlets at their required flow rates were calculated for a wind speed of 0 and 24 km h⁻¹, 226 the latter used as a maximum representative wind speed for operation. The wind speeds required for optimum 227 performance are <8 km h⁻¹ for the impactors and OPC, but the system may experience higher wind speeds. Hence, 228 we use this modelling to guide our choice of positioning of the instrument relative to wind direction in order to 229 minimise sampling biases at the inlets. The modelling also allows us to better understand which impactor stages 230 (and OPC size bins) will be most affected by such biases. We make no attempt to correct the measurements for sampling biases, since this correction itself would carry substantial uncertainty, but used the calculations to informus of the best configuration for the various inlets.

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234 The inlet sampling efficiencies in the orientations chosen for the final design of the SHARK are shown in Figure 3. It is important to note that, due to their dissimilar inlet dimensions and operational flow rates, Impactors 1 and 235 236 2 are affected differently by the wind. The particle losses for the largest stages of each impactor are the most 237 affected. Stages a to d on both impactors are only minimally affected by losses. The losses are more significant 238 in stage e on both impactors, but the losses on 1e are greater than on 2e with a 50% cut off at around 5.5 μ m and a negligible sampling efficiency above about 8 µm on 1e. These calculations also demonstrate that the losses are 239 wind-speed dependent, but that in situations where there is significant wind, the results from Impactor 2 will be 240 241 less influenced by losses than Impactor 1 at sizes above 2.5 µm

The OPC suffers up to 1.6 times oversampling for 10 μ m particles when sampling into 24 km h⁻¹ wind, but when oriented at 90° to the wind the collection efficiency of >6 μ m particles approaches 0 % (see Figure 3c). Therefore, the OPC has been positioned in the SHARK to be oriented into the wind to ensure data is collected for the whole

size range, with the caveat of a sub-isokinetic oversampling of larger particles.

247 3 Results and Discussion

The SHARK has been deployed at ground level and on a tethered balloon during development and testing at four locations for the collection and monitoring of aerosol: Cardington (UK), Hyytiälä (Finland), Leeds (UK), and Longyearbyen (Svalbard). Details of the sampling locations, periods, and instrumentation can be found in Table S1 of the Supplementary Information (SI). In this section, we present the results for this set of four SHARK deployments to illustrate the capabilities of the SHARK for quantifying ice-nucleating particle spectra as well as demonstrating that the technique is consistent with more established methods.

254 3.1 Meteorological and aerosol size distribution data from a SHARK flight

255 An example of the radiosonde and OPC data that was collected during a SHARK flight is shown in Figure 4. The 256 data was from a sampling event in the High Arctic in the summer of 2018, during which the meteorological data 257 from the radiosonde and aerosol particle data from the OPC were collected alongside impactor films for INP 258 analysis (the INP results will be published elsewhere). Throughout the 4.5 h flight the altitude, humidity and 259 temperature were closely monitored to inform decisions on sampling. The sampling start and end times are 260 indicated as solid lines in Figure 4. The SHARK reached 450 m above Mean Sea Level (MSL) and in the last hour of flight lowered to 350 m due to ice formation on the balloon, instrument and tether. The RH during the flight 261 262 was monitored to ensure the SHARK did not sample in humidity approaching saturation; the impactor and OPC 263 manufacturers' specified thresholds for the components is 95 % RH, but we aim to only sample with the RH below 264 this value (~80 %) in order to reduce the influence of hygroscopic growth on aerosol size. After sampling was 265 stopped, the SHARK was brought down to ground level, resulting in the humidity rising. The ability to stop the sampling during the flight meant the impactors were covered and the pumps turned off during the descent and so 266 267 did not sample the more humid environment. The ambient temperature was monitored alongside the dewpoint temperature to follow the surface inversions. The temperature inversion was used to determine where to stabilise the SHARK and begin sampling, as sampling was desired above the surface inversion for this run.

- 271 The total particle counts per 1.38 s interval from the OPC are shown in Figure 4d. Processing of the OPC data 272 yielded the results shown in Figure 5 for the particle number $(dN/dlogD_p)$, particle surface area $(dS/dlogD_p)$ and 273 particle mass $(dM/dlog D_p)$ size distribution data for the sampling period, where D_p is the particle diameter. We 274 present this data to demonstrate that the OPC produces reasonable data when used facing into wind while 275 suspended from a balloon at altitude. Unfortunately, there is no direct comparison with other aerosol size 276 distribution measurements at the sampling location. While the particle number concentration decreases roughly 277 linearly with size, the surface and mass concentration curves have a mode at around 4 μm in Figure 5b and 5c. 278 This is consistent with previous studies conducted within the boundary layer in the Arctic (Freud et al., 2017; 279 Hegg et al., 1996; Seinfeld and Pandis, 2016).
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281 3.2 Deriving size-resolved INP concentrations from the SHARK samples

The ability to measure INP concentrations and properties using samples collected via the SHARK was tested by 282 performing immersion mode droplet freezing assays on the sampled aerosols. Following a flight, impactor films 283 were removed from both cascade impactors of the SHARK, then each immersed in 5 mL of water and mixed on 284 285 a vortex mixer for 5 min to wash the collected particles into suspension (O'Sullivan et al., 2018). This suspension was then analysed via a droplet freezing assay using the microlitre Nucleation by Immersed Particle Instrument 286 (µL-NIPI) (Whale et al., 2015), in which 40-50 droplets of 1 µL volume were pipetted onto a hydrophobic glass 287 slide atop a cold plate. A Perspex shield was placed over the cold stage and N2 gas introduced to purge the chamber 288 of moisture as the cold plate was cooled to -40 °C at 1 °C min⁻¹. The temperatures at which droplets froze were 289 290 recorded using video analysis until the entire population had frozen. This allowed the fraction of droplets frozen as a function of temperature, $f_{ice}(T)$, to be calculated (O'Sullivan et al., 2018; Whale et al., 2015) using the equation 291 $f_{\text{ice}}(T) = N_f / N_t$, where N_f is the number of frozen droplets at temperature T, and N_t is the total number of droplets. 292 293 The INP concentration per volume of sampled air as a function of temperature, $[INP]_T$, was then calculated for each film using $f_{ice}(T)$, according to Equation 1 adapted from (Vali, 1971) to include weighting to the volume of 294 295 air sampled:

$$[INP]_T = -\frac{\ln\left(1 - f_{ice}(T)\right)}{v_{droplet}} \cdot \frac{v_{wash}}{v_{air}},$$
(1)

where $V_{droplet}$ is the droplet volume (i.e. 1 µL), V_{wash} is the amount of water into which the filter is immersed to produce the suspension for analysis (i.e. 5 mL), and V_{air} is the volume of air sampled.

300 3.3 Testing the SHARK INP concentrations against a standard aerosol sampler

In order to test whether the SHARK impactors were sampling in a representative manner, the SHARK was run
 concurrently with a filter-based particle sampler (BGI PQ100, Mesa Labs) and which is used as an EPA Federal
 Reference Method for PM₁₀ (designation no. RFPS-1298-124). This sampler was equipped with a PM10 head and
 an optional cyclone impactor which provided a size cut at 2.5 µm. Aerosol was collected onto 0.4 µm pore size

305 Nuclepore Track-Etched Membrane polycarbonate filters at a flow rate of 16.7 L min⁻¹ (i.e. 1 m³ h⁻¹). This type 306 of filter collects particles across the full range of available aerosol sizes, even at sizes smaller than the pore 307 diameter, with high collection efficiencies (Lindsley, 2016; Soo et al., 2016). These polycarbonate filters have 308 also been successfully employed in other ice nucleation field measurements (DeMott et al., 2016; Harrison et al., 2018; Huffman et al., 2013; McCluskey et al., 2016; Reicher et al., 2019; Tarn et al., 2018). These substrates are 309 310 known to have a low ice-nucleating ability and allow the collected particles to be released into suspension for 311 subsequent INP analysis (O'Sullivan et al., 2018). The filters were analysed using the µL-NIPI in the same manner as for the impactor films collected using the SHARK. The PQ100 filter sampler was deployed alongside the 312 313 SHARK in Cardington (UK) and in Hyytiälä (Finland).

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315 In order to compare the SHARK-derived, size-resolved INP data with the results of the PM10 or PM2.5 PQ100 316 filter sampler, the INP concentrations determined across the appropriate SHARK size categories were summed. 317 In Figure 6a, data is presented from Cardington, where the sum of 2d and 2e from SHARK is compared with the filter sampler fitted with a PM₁₀ head (Impactor 1 was not available during this test). The SHARK was suspended 318 319 from a tethered balloon roughly 20 m from the ground, whereas the filter sampler was on the ground (inlet ~150 cm above the surface), where both samplers were within the well-mixed boundary layer. The agreement is very 320 good apart from two highest temperature points from the filter sampler, but note that the Poisson uncertainties on 321 these points are substantial and also that the two samplers were separated vertically by 20 m. 322

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324 We then show data from Hyytiälä in Figure 6b where we compare the INP spectrum from the filter sampler, with 325 a PM_{2.5} cut-off installed, with the sum of stages 1b, 1c and 1d (the after-filter, stage 1a was not used on Impactor 326 1 in this case). Here, both samplers were positioned within a few metres above the ground. Again, the agreement 327 between the SHARK and the filter sampler was very good. For both Cardington and Hyytiälä, the smallest 328 particles (<0.25 µm) were not sampled using the SHARK, but the agreement between the filter sampler and the 329 SHARK implies that, in these cases, the smallest particles made a minor contribution to the overall INP 330 population, which is what we would generally anticipate from the literature (Berezinski et al., 1988; Huffman et 331 al., 2013; Mason et al., 2016; Santachiara et al., 2010; Si et al., 2018; Welti et al., 2009). The consistency between 332 the SHARK and the filter sampler indicates that there are no major losses of aerosol in the SHARK sampler, at 333 least relative to the PQ100 filter sampler.

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335 3.4 Consistency of INP concentrations between SHARK impactors

An example of data from the size-resolved collection and analysis of INPs is shown in Figure 7, from a sampling 336 run performed in Leeds (UK). The $f_{ice}(T)$ curves for each impactor stage are illustrated in Figure 7a. As discussed 337 in section 2.2, there are two stages, d and e, which have similar size cuts on both stages. Using stage e as an 338 339 example, it can be seen that while the fraction frozen curves for the two samplers are shifted by about 3 $^\circ$ C (Figure 340 7a), normalising to the volume of air sampled to yield $[INP]_T$ in Figure 7b shows that the INP spectra derived 341 from stages 1e and 2e are consistent with one another. Stage 2e covers a lower range of INP concentrations than 342 stage 1e by about 1 order of magnitude, because the flow rate through this impactor was more than a factor of 343 11.1 (100 L min⁻¹/9 L min⁻¹) higher and the probability of collecting rarer INP was increased by this factor. The

344 agreement between the two impactors indicates that aerosol was collected with no significant losses/enhancements 345 due to factors like particle bounce or wind observed. Based on the inlet particle loss calculations in Figure 3, 346 higher losses may have been expected in impactor stage 1*e*, but these are not apparent here.

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348 3.5 Size-resolved ice-nucleating particle (srINP) spectra at four locations

349 The derived size-resolved INP (srINP) concentrations for all four test sites are shown in Figure 8 and Figure 9. 350 The $f_{ice}(T)$ curves for these test sites can be found in Figures S1 to S4 of the S1. Figure 8 shows the INP 351 concentration spectra in the classic form, wherein INP concentrations are plotted against temperature for each size 352 bin, whereas Figure 9 shows the same data in novel srINP plots to allow more intuitive comparison of the INP 353 concentration contribution from each stage with respect to temperature. In Figure 9, where there were 354 measurements from two impactors for the same stage (e.g. d and e), the INP concentrations were merged by taking 355 an average at temperature intervals of 0.5 °C (also for Figure 6). The colour gradient in Figure 9 represents the temperature dependant concentration for each size bin and the overall steepness of the $d[INP]_T/dT$ curve. The 356 357 steepness of the INP spectra can be useful in discriminating between different INP species. On inspection of Figure 8 and Figure 9, it can be seen that the spectra in the four locations have very different characteristics. Not 358 only does the general shape of the spectra vary, but the size-dependence is also very different in the four locations. 359 360 Due to the sample size, these variations could be attributed to the different aerosol population in each location, 361 the time of year and meteorology, which could affect the INP concentrations and spectra (Kanji et al., 2017; Šantl-362 Temkiv et al., 2019; Tobo et al., 2019; Wex et al., 2019). We now discuss the size-resolved INP concentration 363 spectra from these tests, bearing in mind that these four tests were one-offs and should not be regarded as characteristic of those sampling sites, but rather illustrative of the importance of making size-resolved 364 365 measurements.

The first site testing of a prototype of the SHARK in which all of the components were installed was conducted 367 in Cardington (UK) on the 15th of May 2018, but only Impactor 2 was used (see Figure 6a and Figure 8a). The 368 369 Cardington site is an airfield, with large areas of grassy land near a main road, and the sampling was conducted 370 during spring. In order to demonstrate the utility of the SHARK to make balloon-borne INP measurements whilst 371 providing a comparison with a commercial ground-based sampler, the SHARK was sampling whilst suspended 372 from a tethered balloon, flying roughly 20 m above the ground. The INP spectra (Figure 8a and 9a) in this location 373 are steep, increasing two orders of magnitude within 2.5 °C, and are centred around -18 to -20 °C; the [INP]_T for 374 2f and 2e increases by an order of magnitude in just ~1 °C. The INPs in this location were dominated by particles 375 greater than 2.5 µm, whereas particles between 1-2.5 µm made a smaller contribution and show a shallower 376 $d[INP]_{T}/dT$, seen in Figure 9a as a larger spread of data. We speculate that the coarse mode INPs at this site were 377 of biological origin, such as fungal material, pollen or bacteria with a steep INP spectrum (Kanji et al., 2017). The 378 steepness of the curve and the temperature are consistent with ice nucleation by pollen (O'Sullivan et al., 2015; 379 Pummer et al., 2012; Tarn et al., 2018). Although the size of whole pollen grains are often larger than 10 µm, 380 pollen is known to release nanoscale materials that nucleate ice, which might be internally mixed with aerosol in 381 this size bin.

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383 In Hyytiälä (Finland), a field site in the boreal forest, the INP spectra contrast quite strongly with those in Cardington (see Figure 6b and Figure 8b). Sampling took place on the 11th of March 2018, when the Hyytiälä site 384 was snow-covered and sampling was performed at the surface (inlet ~150 cm above surface). In this case only 385 386 Impactor 1 was used without the after-filter installed. The complex nature of the size-dependence of INP is clear here. Intriguingly, in this location, the INP concentration was greatest for the smallest stage used (1b; 0.25-0.5 387 388 μ m), and accounted for the majority of the INPs between -17 and -22 °C. The fewest INP came from the next 389 smallest stage 1c (0.5-1 µm), while at temperatures below -23 °C, stage 1e contained the majority of the INPs. 390 These results indicate that the INP spectra are complex, and that concentrations of INPs do not always increase with increasing size as might be expected. Huffman et al. (2013) reported INP concentration measurements in a 391 392 forest ecosystem, where the particles between 1.8 and 5.6 µm enhanced during rain. Hence, as in the present study, 393 Huffman et al. (2013) showed that INP activity does not always increase with size. The highest INP concentrations 394 in Hyytiälä were measured for aerosol sizes of 0.25-0.5 µm, and we note that these accumulation mode INPs would have lifetimes of many days to weeks in the atmosphere and could therefore be transported to locations and 395 396 altitudes where they may influence clouds. Clearly, this would be an interesting location for more measurements with the full SHARK payload to gain further information on the long term INP concentration variations and the 397 aerosol sizes responsible for them. 398 399

400 The testing in Leeds (UK) used both impactors at ground level with the SHARK suspended from a frame to allow 401 orientation into wind. The Leeds sampling was conducted within the University of Leeds campus on a patch of 402 grass on the 7th of June 2018 in close proximity to the School of Earth and Environment. In this test the full suite 403 was deployed, including the impactors, after-filter and OPC. The particle number, surface and mass size 404 distribution data for this test can be found in Figures S5 to S7 of the SI. It can be seen in Figure 8c that generally, 405 the larger bins contained more active INP. The only exception to this occurred with the after-filter ($\leq 0.25 \ \mu m$), 406 which had slightly higher INP concentrations below about -25 °C than the next two size bins (0.25-1.0 μ m). As 407 with the measurements in Hyytiälä, clearly more measurements illuminating the contribution of the smaller 408 particles in similar environments would be beneficial since the atmospheric lifetime of these fine particles is 409 relatively long. We note that a substantial proportion of INPs quantified just outside of Leeds in a previous study 410 were heat-sensitive and therefore most likely of biological origin (O'Sullivan et al. 2018). In the future, conducting heat tests, as well as using Mass Spectrometry, SEM and DNA analysis with the size-resolved INP samples may 411 412 help to identify the INP types in the various size fractions and highlight any differences between size ranges. 413

414 The final test was in Longyearbyen (Svalbard) from the 7th deck of the icebreaker Oden, 25 m above the surface, 415 when moored ~ 200 m from the shore, overnight from the 23^{rd} to the 24^{th} of September 2018. The full SHARK 416 payload was used in this case, with the OPC, both impactors and the after-filter on Impactor 1. The particle 417 number, surface and mass size distribution data for this test can be found in Figures S8 to S10 of the SI. The INP 418 spectra in this location, shown in Figure 8d was quite distinct from the other three locations in that all size fractions 419 contributed similarly to the INP population and there is a very shallow slope of $dln[INP]_T/dT$ (Figure 9d). We 420 detected INPs at temperatures of up to -10 °C with concentrations of around 0.01 INP L⁻¹. These high-temperature 421 INP concentrations are consistent with the summertime measurements reported at other Arctic locations, including 422 Ny-Ålesund (Svalbard) (Wex et al., 2019). The INP in this region potentially originate from a range of sources.

Tobo et al. (2019) recently reported that dust and biological material from glacial valleys in Svalbard may be an important source of INPs in the region. We also note that we sampled while the Oden was moored in the port of

- 425 Longyearbyen where local pollution sources may have been significant (Zhao et al., 2019).
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427 **3.6** Ice-active surface site density, $n_s(T)$, and the activated fraction, $n_n(T)$

The addition of size distribution information to the INP concentration spectra allowed the calculation of the number of active sites per unit surface area, $n_s(T)$, and the activated fraction, $n_n(T)$, of the size-resolved samples. These quantities are determined by weighting the srINP concentrations to the total surface area and the aerosol number in each size bin, respectively, as shown in Equations 2 and 3.

432 $n_{\rm s}(T) = -\frac{\ln\left(1 - f_{\rm ice}(T)\right)}{A_{\rm s}},$ (2)

433 where A_s is the total surface area of the particles per droplet in a μ L-NIPI droplet freezing assay. This was 434 calculated for each impactor size range, using data from the relevant size bins of the OPC data.

435
$$n_{\rm n}(T) = -\frac{\ln\left(1 - f_{\rm ice}(T)\right)}{N}$$
, (3)

where N is the total number of particles sampled by the impactor in each size bin, calculated using the number
concentration in each size category as measured by the OPC, and the volume of air sampled by the impactor. The
size bins from the OPC which have been included in the calculations were matched to those in the impactors. The
bin boundaries for the OPC calculations were within tens of nanometres of the impactor bin boundaries.

Calculating the $n_s(T)$ and $n_n(T)$ values from the INP data was only possible for some of the size ranges due to the sampling ranges of the instrumentation employed. The smallest particle diameter measured by the OPC is 0.38 µm, i.e. above the lower limit of impactor stage 1*b*, while the largest impactor stage, 2f (>10 µm) has no defined upper bound. Therefore, the three bins (i.e. impactor stages) that were used to produce $n_s(T)$ and $n_n(T)$ were *c* (0.5-1.0 µm), *d* (1.0-2.5 µm) and *e* (2.5-10 µm). The $n_s(T)$ and $n_n(T)$ data were calculated for the field tests in Leeds and Longyearbyen; data from Cardington and Hyytiälä is not provided as the OPC was not in use at these sites.

The plots of activated fraction shown in Figure 10 are addressed first. For the Leeds sample, there is a difference in the $n_n(T)$ values between bins *c* to *e* (Figure 10a), where the smallest bin is 1-3 orders of magnitude lower than the largest bin, with the middle bin in the centre of the two. In Longyearbyen (Figure 10b), the $n_n(T)$ for bin *e* is about a factor of 10 larger than bin *c*, but bins *c* and *d* produce very similar values of $n_n(T)$. Overall, these $n_n(T)$ plots show that the coarse mode aerosol generally have a higher fraction of aerosol that serve as INPs than the fine mode, but there is variability in the dependence on size between the two samples. In contrast to the $n_n(T)$ values, the size resolved $n_s(T)$ data for both Leeds and Longyearbyen show that the data from the three size

455 categories are all within a factor of 2-10 (close to our uncertainty estimates). Given the activity of aerosol across 456 these bins scales with surface area, this data might indicate the same INP species is active across each bin at these

457 sites.

Deleted: where *N* is the total number of particles sampled during the sampling period in each size category measured by the OPC...

461 4 Conclusions

This paper describes a lightweight and portable payload, the SHARK, that is capable of collecting size-resolved 462 463 aerosol particles alongside measurements of ambient temperature, relative humidity, pressure, GPS coordinates, aerosol number distribution and aerosol size distribution. The 9 kg payload was designed for use on a tethered 464 balloon for measurements at user-selected altitudes for up to 11 h via radio controlled instrumentation, but can be 465 466 used wherever it can be suspended. During a SHARK flight, the atmospheric conditions the SHARK experiences can be monitored in real-time via a radiosonde and sampling is controlled remotely, allowing the SHARK to be 467 468 held at a user-defined height and to only sample under specific conditions (for instance above the surface boundary 469 layer).

470

471 The SHARK samples aerosol onto filter/film substrates using two cascade impactors to allow aerosol size-472 segregation from 0.25 to 10 $\mu m,$ with an after-filter and top stage to collect particles below and above this range. 473 One impactor samples at 9 L min⁻¹, while the other samples at 100 L min⁻¹. The filters were collected here for the 474 offline analysis of INP concentrations and properties, but they could equally be used for other analyses such as 475 mass spectrometry, DNA analysis, SEM, TEM and ion chromatography. A comparison of ambient INP 476 concentrations measured using the SHARK to those measured using PM10 and PM2.5 aerosol samplers at ground 477 level demonstrated excellent agreement between the instruments. Field testing was conducted in four locations to demonstrate the capabilities of the SHARK. 478

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The size resolved INP concentration spectra reveal complex behaviour. For example in Hyytiälä the 0.25-0.5 μ m aerosol size fraction had the most active INP, whereas in Leeds the INP concentration generally decreased with decreasing particle size. Ambient aerosol size distribution measured using the on-board OPC allowed the calculation of the activated fraction ($n_n(T)$) and ice-active surface site density ($n_s(T)$) data for the sampled INPs in the tests at Leeds and Longyearbyen. It was shown that $n_s(T)$ was consistent between 0.5 and 10 μ m in these two locations at the times of sampling. It will be interesting to make similar measurements in other locations in the future.

487

488 Generally, it is expected that larger aerosol are more likely to nucleate ice (Pruppacher, H.R. and Klett, 1997) and 489 our results are consistent with other size resolved INP measurements which indicate that the size distribution of INP varies spatially and temporarily e.g. (Mason et al., 2016; Si et al., 2018). Quantifying the size of INP, possibly 490 491 in conjunction with other analytical techniques, is a useful means of identifying different INP types and their sources (Huffman et al., 2013). In addition, knowledge of their size will allow the improved representation of INP 492 493 in global aerosol models where size is key determinant of lifetime and transport (Atkinson et al., 2013; Perlwitz et al., 2015; Vergara-Temprado et al., 2017). Clearly, more systematic and widespread measurements of INP size 494 495 is needed in the future in a range of target locations.

496

The high sample flow rate, choice of low contamination aerosol collection substrates and long sampling durations mean that the payload is well suited for INP measurements, including those in low aerosol environments and locations with relatively low INP concentrations (down to below ~0.01 INP L⁻¹ and at temperatures down to about -25 °C and below). The SHARK is an accessible tool for quantifying size-resolved atmospheric INP 501 concentrations through the vertical profile, both within and above the atmospheric boundary layer. This will allow 502 improved determination of INP sizes, properties, and sources, towards ultimately improving model 503 representations of atmospheric INP distributions.

504 Data availability

505 The data sets for this paper will be made publicly available in the University of Leeds Data Repository upon 506 publication.

507 Author contribution

GCEP led the development of the SHARK, performed the bulk of the experiments and led the writing of the paper. The initial instrument concept was conceived by GCEP, SNFS and BJM with advice from IMB. The building and testing of the SHARK and its electrical components was done by SNFS with the assistance of GCEP. The collection and analysis of field samples was performed by GCEP, MPA, UP, ADH, MDT and IMB. All authors contributed to the writing of this paper. BJM oversaw this project as part of his MarineIce ERC fellowship.

513 Competing interest

514 The authors declare that they have no conflict of interest.

515 Acknowledgements

516 The personnel of Hyytiälä forestry station, the HyIce project team, the Cardington meteorological research unit, 517 and those aboard the icebreaker Oden during 2018 are sincerely thanked for support during field testing. One of 518 the field tests was completed as part of the Arctic Ocean (AO) 2018 expedition. The Swedish Polar Research 519 Secretariat (SPRS) provided access to the icebreaker Oden and logistical support in collaboration with the U.S. 520 National Science Foundation. We are grateful to the Chief Scientists Caroline Leck and Patricia Matrai for 521 planning and coordination of AO2018, as well as to the SPRS logistical staff and icebreaker Oden's Captain 522 Mattias Peterson and his crew for expert field support. The authors thank the European Research Council for 523 funding (H2020 ERC; 648661 MarineIce) and the Natural Environment Research Council (NERC, 524 NE/M010473/1, NE/R009686/1). We are grateful to the EU's H2020 ACTRIS-2 for a mobility grant to access 525 the Hyytiälä forestry station as part of the HyIce project (SMR7 RP3 HyICE18, 654109). Anthony Windross and Stephen Burgess are thanked for help with the fabrication of the SHARK housing. 526 527

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528 References

- Armstrong, J. A., Russell, P. A., Sparks, L. E. and Drehmel, D. C.: Tethered Balloon Sampling Systems for
 Monitoring Air Pollution, J. Air Pollut. Control Assoc., 31(7), 735–743, doi:10.1080/00022470.1981.10465268,
- 531 1981.
- 532 Atkinson, J. D., Murray, B. J., Woodhouse, M. T., Whale, T. F., Baustian, K. J., Carslaw, K. S., Dobbie, S.,
- 533 O'Sullivan, D., Malkin, T. L., O'sullivan, D. and Malkin, T. L.: The importance of feldspar for ice nucleation by
- 534 mineral dust in mixed-phase clouds., Nature, 498(7454), 355–8, doi:10.1038/nature12278, 2013.
- Ault, A. P. and Axson, J. L.: Atmospheric Aerosol Chemistry: Spectroscopic and Microscopic Advances, Anal.
 Chem., 89(1), 430–452, doi:10.1021/acs.analchem.6b04670, 2017.
- Balsley, B. B., Jensen, M. L. and Frehlich, R. G.: The use of state-of-the-art kites for profiling the lower
 atmosphere, Boundary-Layer Meteorol., 87(1), 1–25, doi:10.1023/A:1000812511429, 1998.
- Berezinski, N. A., Stepanov, G. V. and Khorguani, V. G.: Ice-forming activity of atmospheric aerosol particles of
 different sizes, Atmos. Aerosols Nucleation, 309, 709–712, doi:https://doi.org/10.1007/3-540-50108-8_1167,
 1988.
- Cheng, Y.-S. and Yeh, H.-C.: Particle bounce in cascade impactors., Environ. Sci. Technol., 13(11), 1392–1396,
 doi:10.1021/es60159a017, 1979.
- 544 Conen, F., Rodríguez, S., Hülin, C., Henne, S., Herrmann, E., Bukowiecki, N., Alewell, C., Rodri'guez, S.,
- Rodri'guez, R., Hu"glin, C. and Hu"glin, H.: Atmospheric ice nuclei at the high-altitude observatory Jungfraujoch,
 Chem. Phys. Meteorol., 67(1), doi:10.3402/tellusb.v67.25014, 2015.
- Creamean, J. M., Primm, K., Tolbert, M. A., Hall, E. G., Wendell, J., Jordan, A., Sheridan, P. J., Smith, J. and
 Schnell, R. C.: HOVERCAT: A novel aerial system for evaluation of aerosol-cloud interactions, Atmos. Meas.
 Tech., 11(7), 3969–3985, doi:10.5194/amt-11-3969-2018, 2018a.
- 550 Creamean, J. M., Kirpes, R. M., Pratt, K. A., Spada, N. J., Maahn, M., De Boer, G., Schnell, R. C. and China, S.:
- Marine and terrestrial influences on ice nucleating particles during continuous springtime measurements in an
 Arctic oilfield location, Atmos. Chem. Phys., 18(24), 18023–18042, doi:10.5194/acp-18-18023-2018, 2018b.
- 553 DeMott, P. J., Prenni, A. J., Liu, X., Kreidenweis, S. M., Petters, M. D., Twohy, C. H., Richardson, M. S.,
- Eidhammer, T. and Rogers, D. C.: Predicting global atmospheric ice nuclei distributions and their impacts on
 climate, Proc. Natl. Acad. Sci., 107(25), 11217–11222, doi:10.1073/pnas.0910818107, 2010.
- 556 DeMott, P. J., Prenni, A. J., McMeeking, G. R., Sullivan, R. C., Petters, M. D., Tobo, Y., Niemand, M., Möhler,
- O., Snider, J. R., Wang, Z. and Kreidenweis, S. M.: Integrating laboratory and field data to quantify the immersion
 freezing ice nucleation activity of mineral dust particles, Atmos. Chem. Phys., 15(1), 393–409, doi:10.5194/acp 15-393-2015, 2015.
- 560 DeMott, P. J., Hill, T. C. J., McCluskey, C. S., Prather, K. A., Collins, D. B., Sullivan, R. C., Ruppel, M. J.,
- 561 Mason, R. H., Irish, V. E., Lee, T., Hwang, C. Y., Rhee, T. S., Snider, J. R., McMeeking, G. R., Dhaniyala, S.,
- 562 Lewis, E. R., Wentzell, J. J. B., Abbatt, J., Lee, C., Sultana, C. M., Ault, A. P., Axson, J. L., Diaz Martinez, M.,
- 563 Venero, I., Santos-Figueroa, G., Stokes, M. D., Deane, G. B., Mayol-Bracero, O. L., Grassian, V. H., Bertram, T.

- H., Bertram, A. K., Moffett, B. F. and Franc, G. D.: Sea spray aerosol as a unique source of ice nucleating particles,
 Proc. Natl. Acad. Sci., 113(21), 5797–5803, doi:10.1073/pnas.1514034112, 2016.
- 566 DeMott, P. J., Hill, T. C. J., Petters, M. D., Bertram, A. K., Tobo, Y., Mason, R. H., Suski, K. J., Mccluskey, C.
- 567 S., Levin, E. J. T., Schill, G. P., Boose, Y., Rauker, A. M., Miller, A. J., Zaragoza, J., Rocci, K., Rothfuss, N. E.,
- 568 Taylor, H. P., Hader, J. D., Chou, C., Huffman, J. A., Pöschl, U., Prenni, A. J. and Kreidenweis, S. M.:
- 569 Comparative measurements of ambient atmospheric concentrations of ice nucleating particles using multiple
- immersion freezing methods and a continuous flow diffusion chamber, Atmos. Chem. Phys., 17(18), 11227–
 11245, doi:10.5194/acp-17-11227-2017, 2017.
- Deng, X., Xue, H. and Meng, Z.: The effect of ice nuclei on a deep convective cloud in South China, Atmos. Res.,
 206, 1–12, doi:10.1016/j.atmosres.2018.02.013, 2018.
- 574 Dzubay, T. G., Hines, L. E. and Stevens, R. K.: Particle bounce errors in cascade impactors, Atmos. Environ.,
 575 10(3), 229–234, doi:10.1016/0004-6981(76)90095-0, 1976.
- Freud, E., Krejci, R., Tunved, P., Leaitch, R., Nguyen, Q. T., Massling, A., Skov, H. and Barrie, L.: Pan-Arctic
 aerosol number size distributions: Seasonality and transport patterns, Atmos. Chem. Phys., 17(13), 8101–8128,
 doi:10.5194/acp-17-8101-2017, 2017.
- Garcia, E., Hill, T. C. J., Prenni, A. J., DeMott, P. J., Franc, G. D. and Kreidenweis, S. M.: Biogenic ice nuclei in
 boundary layer air over two U.S. High Plains agricultural regions, J. Geophys. Res. Atmos., 117(D18), n/a-n/a,
 doi:10.1029/2012JD018343, 2012.
- Hara, K., Osada, K. and Yamanouchi, T.: Tethered balloon-borne aerosol measurements: Seasonal and vertical
 variations of aerosol constituents over Syowa Station, Antarctica, Atmos. Chem. Phys., 13(17), 9119–9139,
 doi:10.5194/acp-13-9119-2013, 2013.
- Harrison, A. D., Whale, T. F., Rutledge, R., Lamb, S., Tarn, M. D., Porter, G. C. E., Adams, M. P., McQuaid, J.
 B., Morris, G. J. and Murray, B. J.: An instrument for quantifying heterogeneous ice nucleation in multiwell plates
 using infrared emissions to detect freezing, Atmos. Meas. Tech., 11(10), 5629–5641, doi:10.5194/amt-11-56292018, 2018.
- Hegg, D. A., Hobbs, P. V, Gassó, S., Nance, J. D. and Rangno, A. L.: Aerosol measurements in the Arctic relevant
 to direct and indirect radiative forcing, J. Geophys. Res. Atmos., 101(D18), 23349–23363,
 doi:10.1029/96jd02246, 1996.
- Herbert, R. J., Murray, B. J., Dobbie, S. J. and Koop, T.: Sensitivity of liquid clouds to homogenous freezing
 parameterizations, Geophys. Res. Lett., 42(5), 1599–1605, doi:10.1002/2014GL062729, 2015.
- 594 Huffman, J. A., Prenni, A. J., Demott, P. J., Pöhlker, C., Mason, R. H., Robinson, N. H., Fröhlich-Nowoisky, J.,
- 595 Tobo, Y., Després, V. R., Garcia, E., Gochis, D. J., Harris, E., Müller-Germann, I., Ruzene, C., Schmer, B., Sinha,
- 596 B., Day, D. A., Andreae, M. O., Jimenez, J. L., Gallagher, M., Kreidenweis, S. M., Bertram, A. K., Pöschl, U., M
- ⁵⁹⁷ Uller-Germann, I., Ruzene, C., Schmer, B., Sinha, B., Day, D. A., Andreae, M. O., Jimenez, J. L., Gallagher,
- 598 M., Kreidenweis, S. M., Bertram, A. K. and Pöschl, U.: High concentrations of biological aerosol particles and
 - ice nuclei during and after rain, Atmos. Chem. Phys., 13(13), 6151-6164, doi:10.5194/acp-13-6151-2013, 2013.

599

- Jacob, J., Chilson, P., Houston, A., Smith, S., Jacob, J. D., Chilson, P. B., Houston, A. L. and Smith, S. W.:
 Considerations for Atmospheric Measurements with Small Unmanned Aircraft Systems, Atmosphere (Basel).,
 9(7), 252, doi:10.3390/atmos9070252, 2018.
- 603 Jaenicke, R.: Aerosol Physics and Chemistry. In: Landolt-Börnstein Numerical Data and Functional Relationships
- 604 in Science and Technology New Series Group V: Geophysics and Space Research Volume 4 Meteorology
- 605 Subvolume b, Physical and Chemical Properties of the Air, edited by G. Fischer., 2007.
- 606 Kanitz, T., Seifert, P., Ansmann, A., Engelmann, R., Althausen, D., Casiccia, C. and Rohwer, E. G.: Contrasting

the impact of aerosols at northern and southern midlatitudes on heterogeneous ice formation, Geophys. Res. Lett.,

608 38(17), n/a-n/a, doi:10.1029/2011GL048532, 2011.

607

- Kanji, Z. A., Ladino, L. A., Wex, H., Boose, Y., Burkert-Kohn, M., Cziczo, D. J. and Krämer, M.: Overview of
 Ice Nucleating Particles, Meteorol. Monogr., 58, 1.1-1.33, doi:10.1175/AMSMONOGRAPHS-D-16-0006.1,
 2017.
- Lacher, L., Steinbacher, M., Bukowiecki, N., Herrmann, E., Zipori, A. and Kanji, Z. A.: Impact of air mass
 conditions and aerosol properties on ice nucleating particle concentrations at the High Altitude Research Station
 Jungfraujoch, Atmosphere (Basel)., 9(9), 363, doi:10.3390/atmos9090363, 2018.
- Laskin, J., Laskin, A. and Nizkorodov, S. A.: Mass Spectrometry Analysis in Atmospheric Chemistry, Anal.
 Chem., 90(1), 166–189, doi:10.1021/acs.analchem.7b04249, 2018.
- 617 Lindsley, W. G.: Filter Pore Size and Aerosol Sample Collection, in NIOSH Manual of Analytical Methods, pp.
- 618 1–14. [online] Available from: https://www.cdc.gov/niosh/docs/2014-151/pdfs/chapters/chapter-fp.pdf
 619 (Accessed 30 July 2018), 2016.
- Lohmann, U.: Anthropogenic Aerosol Influences on Mixed-Phase Clouds, Curr. Clim. Chang. Reports, 3(1), 32–
 44, doi:10.1007/s40641-017-0059-9, 2017.
- Marple, V. A., Rubow, K. L. and Behm, S. M.: A microorifice uniform deposit impactor (moudi): Description,
 calibration, and use, Aerosol Sci. Technol., 14(4), 434–436, doi:10.1080/02786829108959504, 1991.
- 624 Mason, R. H., Si, M., Chou, C., Irish, V. E., Dickie, R., Elizondo, P., Wong, R., Brintnell, M., Elsasser, M., Lassar,
- 625 W. M., Pierce, K. M., Leaitch, W. R., MacDonald, A. M., Platt, A., Toom-Sauntry, D., Sarda-Estève, R., Schiller,
- 626 C. L., Suski, K. J., Hill, T. C. J., Abbatt, J. P. D., Huffman, J. A., DeMott, P. J. and Bertram, A. K.: Size-resolved
- measurements of ice-nucleating particles at six locations in North America and one in Europe, Atmos. Chem.
 Phys., 16(3), 1637–1651, doi:10.5194/acp-16-1637-2016, 2016.
- 629 McCluskey, C. S., Hill, T. C. J., Malfatti, F., Sultana, C. M., Lee, C., Santander, M. V, Beall, C. M., Moore, K.
- 630 A., Cornwell, G. C., Collins, D. B., Prather, K. A., Jayarathne, T., Stone, E. A., Azam, F., Kreidenweis, S. M. and
- 631 DeMott, P. J.: A dynamic link between ice nucleating particles released in nascent sea spray aerosol and oceanic
- biological activity during two mesocosm experiments, J. Atmos. Sci., JAS-D-16-0087.1, doi:10.1175/JAS-D-16-
- 633 0087.1, 2016.
- 634 Misra, C., Singh, M., Shen, S., Sioutas, C. and Hall, P. M.: Development and evaluation of a personal cascade
- 635 impactor sampler (PCIS), Aerosol Sci., 33(7), 1027–1047, doi:10.1016/S0021-8502(02)00055-1, 2002.

- Murray, B. J.: Cracking the problem of ice nucleation, Science (80-.)., 355(6323), 346–347,
 doi:10.1126/science.aam5320, 2017.
- Murray, B. J., O'Sullivan, D., Atkinson, J. D. and Webb, M. E.: Ice nucleation by particles immersed in
 supercooled cloud droplets, Chem. Soc. Rev., 41(19), 6519, doi:10.1039/c2cs35200a, 2012.
- 640 O'Sullivan, D., Murray, B. J., Ross, J. F., Whale, T. F., Price, H. C., Atkinson, J. D., Umo, N. S. and Webb, M.
- E.: The relevance of nanoscale biological fragments for ice nucleation in clouds, Sci. Rep., 5(1), 8082,
 doi:10.1038/srep08082, 2015.
- 643 O'Sullivan, D., Adams, M. P., Tarn, M. D., Harrison, A. D., Vergara-Temprado, J., Porter, G. C. E., Holden, M.
- 644 A., Sanchez-Marroquin, A., Carotenuto, F., Whale, T. F., McQuaid, J. B., Walshaw, R., Hedges, D. H. P., Burke,
- I. T., Cui, Z. and Murray, B. J.: Contributions of biogenic material to the atmospheric ice-nucleating particle
 population in North Western Europe, Sci. Rep., 8(1), 13821, doi:10.1038/s41598-018-31981-7, 2018.
- Perlwitz, J. P., Pérez García-Pando, C. and Miller, R. L.: Predicting the mineral composition of dust aerosols-Part
 Representing key processes, Atmos. Chem. Phys, 15, 11593–11627, doi:10.5194/acp-15-11593-2015, 2015.
- 649 Price, H. C., Baustian, K. J., McQuaid, J. B., Blyth, A., Bower, K. N., Choularton, T., Cotton, R. J., Cui, Z., Field,
- 650 P. R., Gallagher, M., Hawker, R., Merrington, A., Miltenberger, A., Neely, R. R., Parker, S. T., Rosenberg, P. D.,
- 651 Taylor, J. W., Trembath, J., Vergara-Temprado, J., Whale, T. F., Wilson, T. W., Young, G. and Murray, B. J.:
- 652 Atmospheric Ice-Nucleating Particles in the Dusty Tropical Atlantic, J. Geophys. Res. Atmos., 123(4), 2175-
- 653 2193, doi:10.1002/2017JD027560, 2018.
- Product Information Sheet MSP: Models 128,129,130 and 131—High Flow Impactors (HFI), [online] Available
- from: http://www.mspcorp.com/resources/msp-pi-130-revb-us-high-flow-impactors-hfi-128-129-130-131.pdf
 (Accessed 3 October 2019), n.d.
- Pruppacher, H.R. and Klett, J. D.: Microphysics of Clouds and Precipitation, 2nd Editio., Kluwer Academic
 Publishers, Dordrecht., 1997.
- Pummer, B. G., Bauer, H., Bernardi, J., Bleicher, S. and Grothe, H.: Suspendable macromolecules are responsible
 for ice nucleation activity of birch and conifer pollen, Atmos. Chem. Phys., 12(5), 2541–2550, doi:10.5194/acp12-2541-2012, 2012.
- Rader, D. J. and Marple, V. A.: Effect of Ultra-Stokesian Drag and Particle Interception on Impaction
 Characteristics, Aerosol Sci. Technol., 4(2), 141–156, doi:10.1080/02786828508959044, 1985.
- Reagan, J. A., Apte, M. V., Bruhns, T. V. and Youngbluth, O.: Lidar and Balloon-Borne Cascade Impactor
 Measurements of Aerosols: A Case Study, Aerosol Sci. Technol., 3(3), 259–275,
 doi:10.1080/02786828408959014, 1984.
- Reicher, N., Segev, L. and Rudich, Y.: The WeIzmann Supercooled Droplets Observation on a Microarray
 (WISDOM) and application for ambient dust, Atmos. Meas. Tech., 11(1), 233–248, doi:10.5194/amt-11-2332018, 2018.
- 670 Reicher, N., Budke, C., Eickhoff, L., Raveh-Rubin, S., Kaplan-Ashiri, I., Koop, T. and Rudich, Y.: Size-dependent
- 671 ice nucleation by airborne particles during dust events in the Eastern Mediterranean, Atmos. Chem. Phys.

- 672 Discuss., 1–26, doi:10.5194/acp-2019-511, 2019.
- Rogers, D. C., Demott, P. J. and Kreidenweis, S. M.: Airborne measurements of tropospheric ice-nucleating
 aerosol particles in the Arctic spring, J. Geophys. Res., 106(D14), 15053–15063, doi:10.1029/2000JD900790,
 2001.
- 676 Rosenfeld, D., Yu, X., Liu, G., Xu, X., Zhu, Y., Yue, Z., Dai, J., Dong, Z., Dong, Y. and Peng, Y.: Glaciation
- temperatures of convective clouds ingesting desert dust, air pollution and smoke from forest fires, Geophys. Res.
 Lett., 38(21), n/a-n/a, doi:10.1029/2011GL049423, 2011.
- Santachiara, G., Di Matteo, L., Prodi, F. and Belosi, F.: Atmospheric particles acting as Ice Forming Nuclei in
 different size ranges, Atmos. Res., 96(2–3), 266–272, doi:10.1016/j.atmosres.2009.08.004, 2010.
- Šantl-Temkiv, T., Lange, R., Beddows, D., Rauter, U., Pilgaard, S., Dall'osto, M., Gunde-Cimerman, N.,
 Massling, A. and Wex, H.: Biogenic Sources of Ice Nucleating Particles at the High Arctic Site Villum Research
 Station, Environ. Sci. Technol., 53(18), 10580–10590, doi:10.1021/acs.est.9b00991, 2019.
- 5 Station, Environ. Sci. Technol., 55(16), 10500–10590, doi:10.1021/dcs.est.9000991, 2019.
- Seinfeld, J. H. and Pandis, S. N.: Atmospheric chemistry and physics : from air pollution to climate change, Third
 edition. [online] Available from: https://www.worldcat.org/title/atmospheric-chemistry-and-physics-from-air pollution-to-climate-change/oclc/929985301 (Accessed 5 July 2019), 2016.
- 687 Si, M., Irish, V. E., Mason, R. H., Vergara-Temprado, J., Hanna, S. J., Ladino, L. A., Yakobi-Hancock, J. D.,
- Schiller, C. L., Wentzell, J. J. B., Abbatt, J. P. D., Carslaw, K. S., Murray, B. J. and Bertram, A. K.: Ice-nucleating
 ability of aerosol particles and possible sources at three coastal marine sites, Atmos. Chem. Phys, 18, 15669–
 15685, doi:10.5194/acp-18-15669-2018, 2018.
- 691 SKC: Leland Legacy Sample Pump: Operating Instructions, , (Form 40075 Rev 1910) [online] Available from:
 692 https://www.skcinc.com/catalog/pdf/instructions/40075.pdf (Accessed 25 February 2020), n.d.
- Soo, J. C., Monaghan, K., Lee, T., Kashon, M. and Harper, M.: Air sampling filtration media: Collection efficiency
 for respirable size-selective sampling, Aerosol Sci. Technol., 50(1), 76–87, doi:10.1080/02786826.2015.1128525,
- 695
 2016.
- Storelvmo, T.: Aerosol Effects on Climate via Mixed-Phase and Ice Clouds, Annu. Rev. Earth Planet. Sci., 45(1),
 199–222, doi:10.1146/annurev-earth-060115-012240, 2017.
- 698 Tarn, M. D., Sikora, S. N. F., Porter, G. C. E., O'Sullivan, D., Adams, M., Whale, T. F., Harrison, A. D., Vergara-
- Temprado, J., Wilson, T. W., Shim, J. uk and Murray, B. J.: The study of atmospheric ice-nucleating particles via
 microfluidically generated droplets, Microfluid. Nanofluidics, 22(5), doi:10.1007/s10404-018-2069-x, 2018.
- Tobo, Y., Prenni, A. J., Demott, P. J., Huffman, J. A., McCluskey, C. S., Tian, G., Pöhlker, C., Pöschl, U. and
 Kreidenweis, S. M.: Biological aerosol particles as a key determinant of ice nuclei populations in a forest
 ecosystem, J. Geophys. Res. Atmos., 118(17), 10100–10110, doi:10.1002/jgrd.50801, 2013.
- Tobo, Y., Adachi, K., Demott, P. J., Hill, T. C. J., Hamilton, D. S., Mahowald, N. M., Nagatsuka, N., Ohata, S.,
- Uetake, J., Kondo, Y. and Koike, M.: Glacially sourced dust as a potentially significant source of ice nucleating
 particles, Nat. Geosci., 12, doi:10.1038/s41561-019-0314-x, 2019.
- 707 Vali, G.: Quantitative Evaluation of Experimental Results an the Heterogeneous Freezing Nucleation of

- Supercooled Liquids, J. Atmos. Sci., 28(3), 402–409, doi:10.1175/1520-0469(1971)028<0402:qeoera>2.0.co;2,
 1971.
- 710 Vergara-Temprado, J., Murray, B. J., Wilson, T. W., O'Sullivan, D., Browse, J., Pringle, K. J., Ardon-Dryer, K.,
- 711 Bertram, A. K., Burrows, S. M., Ceburnis, D., Demott, P. J., Mason, R. H., O'Dowd, C. D., Rinaldi, M. and
- 712 Carslaw, K. S.: Contribution of feldspar and marine organic aerosols to global ice nucleating particle
- 713 concentrations, Atmos. Chem. Phys., 17(5), 3637–3658, doi:10.5194/acp-17-3637-2017, 2017.
- 714 Vergara-Temprado, J., Miltenberger, A. K., Furtado, K., Grosvenor, D. P., Shipway, B. J., Hill, A. A., Wilkinson,
- J. M., Field, P. R., Murray, B. J. and Carslaw, K. S.: Strong control of Southern Ocean cloud reflectivity by ice nucleating particles, Proc. Natl. Acad. Sci., 115(11), 201721627, doi:10.1073/pnas.1721627115, 2018.
- Villa, T., Salimi, F., Morton, K., Morawska, L. and Gonzalez, F.: Development and Validation of a UAV Based
 System for Air Pollution Measurements, Sensors, 16(12), 2202, doi:10.3390/s16122202, 2016.
- Von Der Weiden, S.-L., Drewnick, F. and Borrmann, S.: Particle Loss Calculator a new software tool for the
 assessment of the performance of aerosol inlet systems, Atmos. Meas. Tech, 2, 479–494 [online] Available from:
 www.atmos-meas-tech.net/2/479/2009/ (Accessed 12 January 2018), 2009.
- Welti, A., Lüönd, F., Stetzer, O. and Lohmann, U.: Influence of particle size on the ice nucleating ability of
 mineral dusts, Atmos. Chem. Phys., 9(18), 6705–6715, doi:10.5194/acp-9-6705-2009, 2009.
- 724 Wex, H., Huang, L., Zhang, W., Hung, H., Traversi, R., Becagli, S., Sheesley, R. J., Moffett, C. E., Barrett, T. E.,
- 725 Bossi, R., Skov, H., Hünerbein, A., Lubitz, J., Löffler, M., Linke, O., Hartmann, M., Herenz, P. and Stratmann,
- 726 F.: Annual variability of ice-nucleating particle concentrations at different Arctic locations, Atmos. Chem. Phys.,
- 727 19(7), 5293–5311, doi:10.5194/acp-19-5293-2019, 2019.
- 728 Whale, T. F., Murray, B. J., O'Sullivan, D., Wilson, T. W., Umo, N. S., Baustian, K. J., Atkinson, J. D., Workneh,
- D. A. and Morris, G. J.: A technique for quantifying heterogeneous ice nucleation in microlitre supercooled water
 droplets, Atmos. Meas. Tech., 8(6), 2437–2447, doi:10.5194/amt-8-2437-2015, 2015.
- 731 Zhao, B., Wang, Y., Gu, Y., Liou, K.-N., Jiang, J. H., Fan, J., Liu, X., Huang, L. and Yung, Y. L.: Ice nucleation
- 732 by aerosols from anthropogenic pollution, Nat. Geosci., 12(8), 602–607, doi:10.1038/s41561-019-0389-4, 2019.
- 733

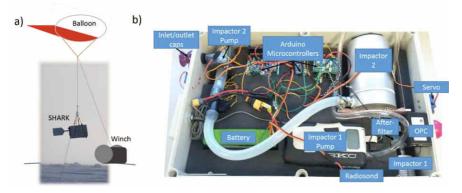


Figure 1. The SHARK. (a) The SHARK payload on a tethered balloon connected to ground by a winch. The photograph was taken during deployment in the High Arctic. (b) The components inside the SHARK payload labelled on a photograph. The payload featured a large impactor inlet at the top of the platform for Impactor 2, with the OPC inlet facing the front, and a small impactor inlet at the bottom for Impactor 1. The radiosond was at the bottom of the box, and the outlet valve for the pump system is shown at the back of the SHARK, where the 100 L min⁻¹ pump for Impactor 2 vents.

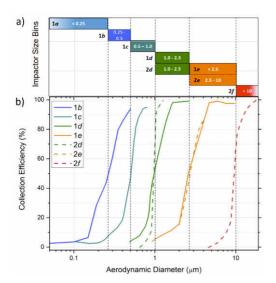


Figure 2. Collection efficiencies of each size bin of the two cascade impactors in the SHARK. (a) The size bins for each stage of Impactor 1 and 2 at flow rates of 9 and 100 L min⁻¹, respectively. (b) Impactor efficiency curves for each stage. Impactor 1 has four stages (1*b-e*) and one after-filter (1*a*), while Impactor 2 has three stages (2*d-f*). Stages 1*d* and 2*d* as well as 1*e* and 2*e* should be approximately equivalent in terms of the aerosol size ranges collected.

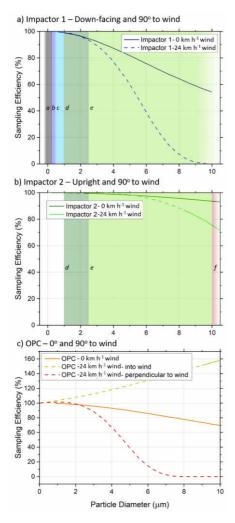


Figure 3. SHARK sampling efficiencies (a) The sampling efficiencies of Impactor 1, with and without wind, when sampling at 90° to the wind direction. (b) The sampling efficiencies of Impactor 2, with and without wind, when sampling at 90° to the wind direction. (c) The sampling efficiency of the OPC, with and without wind, when sampling at 0° and 90° to the wind direction (the OPC was deployed at 0° to the wind, based on this calculation). Solid lines denote model predictions within the formulas' validity range, and dotted lines represent approximations (Von Der Weiden et al., 2009).

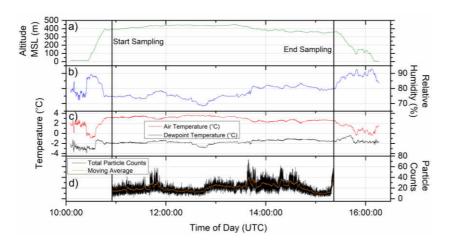


Figure 4. Windsond and optical particle counter (OPC) data for a flight during a campaign to the High Arctic. (a) The altitude of the SHARK payload throughout the 4.5 hour flight. The sampling start and end times are indicated as solid lines. The SHARK reached 450 m above Mean Sea Level (MSL) and in the last hour of flight was lowered to 350 m due to ice formation on the balloon, instrument and tether. (b) The humidity during the flight was monitored to ensure the SHARK was not sampling during unfavourable conditions. The SHARK was not sampling during unfavourable conditions. The SHARK was monitored alongside the dewpoint temperature. (d) Total particle counts throughout the sampling period, as monitored by the OPC.

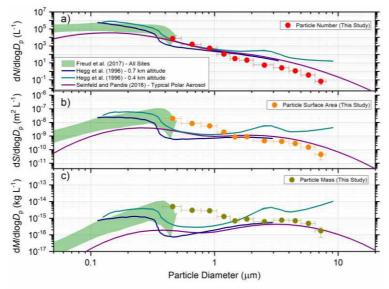


Figure 5. Size distribution data produced from OPC measurements. (a) Particle number, (b) particle surface area and (c) particle mass size distribution data above the surface temperature inversion during a test run of the SHARK suite whilst deployed on a tethered balloon in the High Arctic. Comparisons to previous studies at Arctic sites are shown (Freud et al., 2017; Hegg et al., 1996; Seinfeld and Pandis, 2016). The August aerosol number size distributions for all listed sites in Freud et al., including Zeppelin, Nord, Alert, Barrow and Tiksi are shown. The data from Hegg et al., at altitudes of 0.7 km and 0.4 km in **Prudhoe Bay**, Alaska, are presented. The size distributions from Seinfeld and Pandis are calculated given the parameters for multimode distributions given in Table 8.3.

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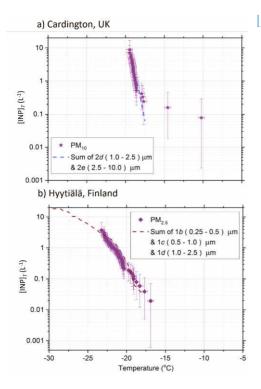


Figure 6. The sum of INP concentrations, [INP]*r*, for labelled stages measured at: (a) Cardington (UK) and (b) Hyytiälä (Finland) alongside data from a standard sampler. Cardington data was taken from Impactor 2 whilst on a tethered balloon at 20 m above ground level, and is shown against a PM₁₀ sampler at ground level, Hyytiälä data was collected using Impactor 1 at ground level, alongside a PM_{2.5} sampler. The dotted lines indicate the sum of the INP concentrations for the SHARK impactor stages, calculated by weighting $f_{ice}(T)$ to the volume of sampled air, and summing the concentrations in each temperature bin.

I

Commented [GP[1]: Changed legend to include size ranges

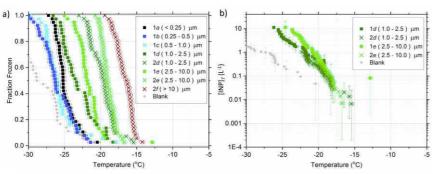


Figure 7. Ice-nucleating particle (INP) analysis of samples collected in Leeds (UK) using the SHARK. (a) The fraction of droplets frozen as a function of temperature, $f_{ice}(T)$, for each stage of Impactors 1 and 2. The handling blank is shown in grey. (b) The INP concentrations for stage 'e' of both impactors (2.5-10 µm), highlighting their excellent agreement.

Commented [GP[2]: Figure changed to include stage d in b)

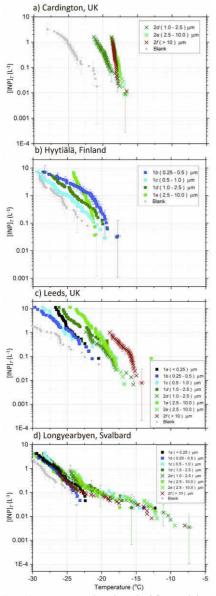


Figure 8. INP concentrations determined from each impactor stage of the SHARK at the four testing sites: (a) Cardington (UK), (b) Hyytiälä (Finland), (c) Leeds (UK) and (d) Longyearbyen (Svalbard). Handling blank data, which determine the baseline of the results, are shown in grey. Samples of the error bars are shown.

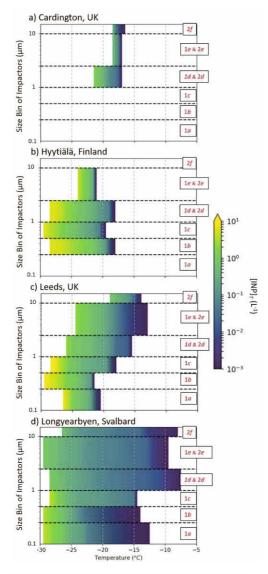
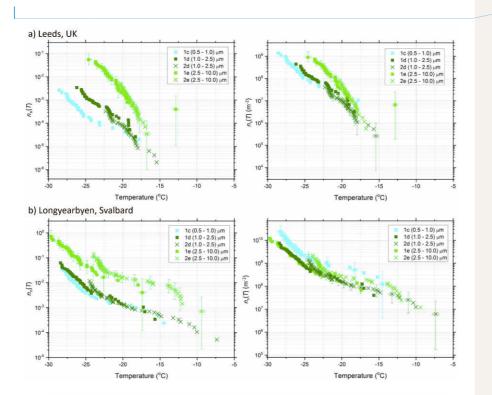


Figure 9. Size-resolved ice-nucleating particle concentrations $(sr[INP]_T)$ for the four test sites: (a) Cardington (UK), (b) Hyytiälä (Finland), (c) Leeds (UK) and (d) Longyearbyen (Svalbard). The colour bars indicate the INP concentration. The dotted lines on the y-axis indicate the size cuts of the impactors. The data from Figure 8 is presented here in an alternative format, which has the advantage of more clearly and concisely displaying the features of the INP spectrum in each size bin than the plots in Fig 8.



Commented [GP[3]: Figure changed to show $n_s(T)$ and $n_n(T)$ on y axis.

Figure 10. Plots showing (left) the activated fraction of aerosol ($n_n(T)$) and (right) the number of active sites per surface area ($n_s(T)$) for samples tested from two measurement sites: (a) Leeds (UK) and (b) Longyearbyen (Svalbard). The colours of the data points indicate the size bins of each impactor, and the different symbols represent the two impactors. Samples of the error bars are shown.

Supplementary information for:

Resolving the size of ice-nucleating particles with a balloon deployable aerosol sampler: the SHARK

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Site	Date (dd/mm/yy)	Sampling period	Impactors	Volume sampled by Impactor 1 at 9 L min ⁻¹ (L)	Volume sampled by Impactor 2 at 100 L min ⁻¹ (L)	Impactor 1 after-filter?	Windsond?	OPC?
Cardington (UK)	15/05/2018	14:15-16:15 (2 h)	Impactor 1 and 2 installed, only Impactor 2 sampling	-	12,000	No	No	Installed but not sampling
Hyytiälä (Finland)	11/03/2018	10:45-16:00 (5 h 15 min)	Impactor 1 only	2,835	-	No	No	No
Leeds (UK)	07/06/2018	12:21-15:21 (3 h)	Both impactors	1,620	18,000	Yes	Yes	Yes
Longyearbyen (Svalbard)	23/09/18-24/09/18	20:00-04:30 (8 h 30 min)	Both impactors	4,590	51,000	Yes	Yes	Yes
High Arctic	20/08/2018	10:40-15:30 (4 h 50 min)	-	-	-	-	Yes	Yes

Table S1. Details of sampling dates, times, locations and SHARK component information. S1 Sampling information

S2 Fraction frozen curves for collected samples

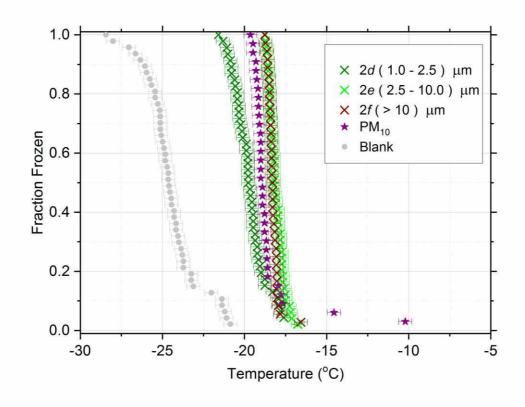


Figure S1. Fraction frozen curves for samples collected in Cardington (UK).

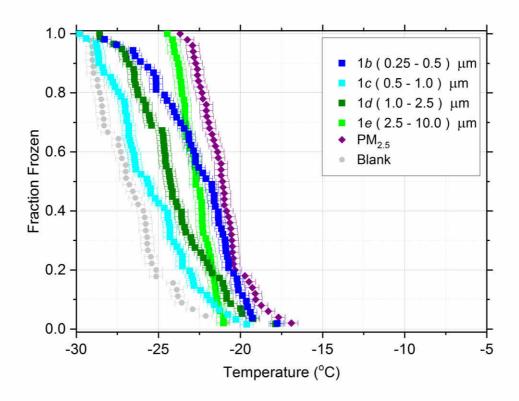


Figure S2. Fraction frozen curves for samples collected in Hyytiälä (Finland).

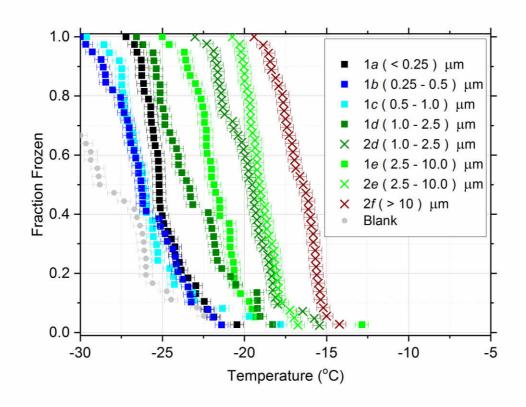


Figure S3. Fraction frozen curves for samples collected in Leeds (UK).

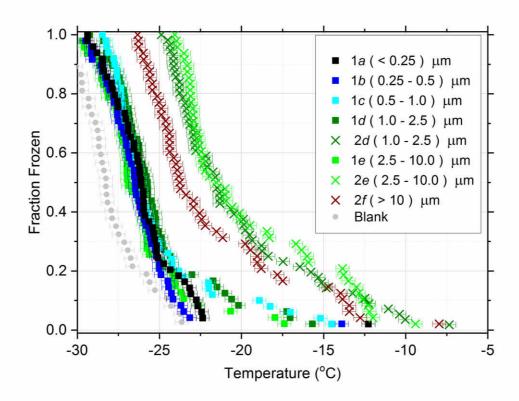


Figure S4. Fraction frozen curves for samples collected in Longyearbyen (Svalbard).

S3 Aerosol data from sampling in Leeds (UK)

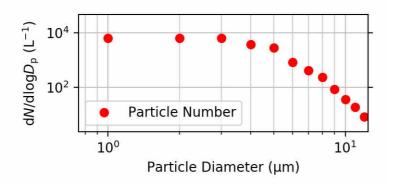


Figure S5. Particle number size distribution data for samples collected in Leeds (UK).

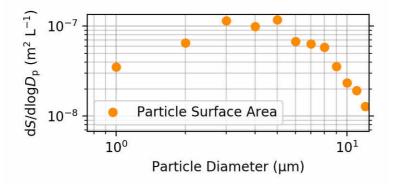


Figure S6. Particle surface area size distribution data for samples collected in Leeds (UK).

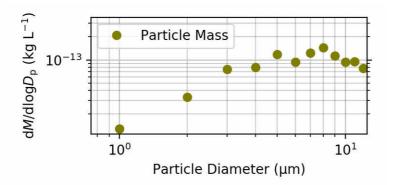


Figure S7. Particle mass size distribution data for samples collected in Leeds (UK).

S4 Aerosol data from sampling in Longyearbyen (Svalbard)

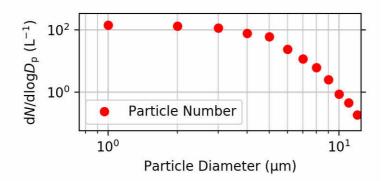


Figure S8. Particle number size distribution data for samples collected in Longyearbyen (Svalbard).

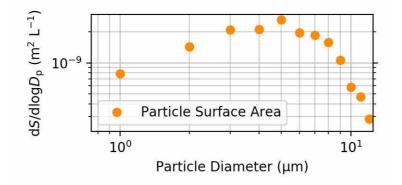


Figure S9. Particle surface area size distribution data for samples collected in Longyearbyen (Svalbard).

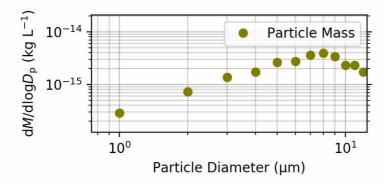


Figure S10. Particle mass size distribution data for samples collected in Longyearbyen (Svalbard).

Response to Referee #1

We would like to thank the referee for their useful comments and have responded below. The referee comments are highlighted in red and our responses are in black.

This work reports on a new technique for measuring size-resolved INP number concentrations using an aerosol sampling system, SHARK. The SHARK system uses two different types of impactors for collecting aerosol samples for INP measurements. Then, they compare the sized-resolved INP number concentrations obtained using these two impactors and demonstrate that the results were almost the same. This result is very important and valuable, because one of the impactors use a high flow rate up to 100 L min-1, indicating the possibility that this impactor may be useful for collecting sufficient amounts of aerosol samples for INP measurements efficiently during a limited sampling time of tethered balloon measurements. However, it is still unclear whether the SHARK system is indeed a reliable technique for measuring size-resolved INP number concentrations at high latitudes, because the results presented here are based on only several measurements near the ground level at Hyytiala, Leeds, Longyearbyen, and Cardington. I would strongly suggest that the authors include the results of balloonborne measurements at higher altitudes (~100 m above ground level or more) in the revised paper.

General comments:

1) I couldn't find any detailed descriptions regarding a campaign in the High Arctic. When and where were the samples shown in Figures 4 and 5 collected? Why didn't you report any results of INP number concentrations using the SHARK during a campaign to the High Arctic?

We have added to Table S1 in the supplementary material to present the sample information for samples collected in Figures 4 and 5 including the sampling location and run time for the OPC.

In this paper we have demonstrated the capability of the SHARK payload in four very different locations, at ground level and suspended from a balloon at 20 m, comparing samples from the SHARK with those from ground-based samplers mounted nearby. Data from the Arctic field campaign are shown purely to demonstrate the effective working of all the SHARK sensors. It is not the purpose of this technical paper to present a scientific analysis of INPs at any of the locations from which measurements were obtained, only to demonstrate that the system produces reliable measurements.

During the campaign to the High Arctic we made what we believe to be the first ever airborne measurements of INPs at the North Pole, as part of the MOCCHA campaign. These results are exciting, and we are preparing a separate publication with a detailed analysis, as described at the start of Section 3.1. It is also unclear what these results would add to the current manuscript. We have already shown that: we can sample when the SHARK is in flight, the INP concentrations are consistent with another filter based sampler, we can communicate with the airborne SHARK when it is above 100 m (Figures 4 and 5), and have described the instrument and the analysis in detail. Since there are no other measurements against which to compare the high-altitude balloon-borne samples from the Arctic, they cannot be used to demonstrate the reliability of the system.

2) Because the objective of this paper must be to report a new technique that can measure size-resolved INP number concentrations using a balloon deployable aerosol sampler, the authors need to evaluate the performance of the impactors used for the SHARK system at high altitudes and provide evidence that it can indeed be useful for INP measurements even at higher altitudes (at least, more than 100 m above ground level).

A height of 20 m was chosen for sampling to demonstrate the utility of the SHARK to make balloon-borne INP measurements, whilst providing comparison with another sampler. We chose to work at 20 m rather than 100 m or more in order to have a comparison with a commercial ground based sampler, as this showed that INP measurements made during flight were comparable to those made on the surface.

We have edited the text in Section 3.5 to emphasise this:

"...The Cardington site is an airfield, with large areas of grassy land near a main road, and the sampling was conducted during spring. In order to demonstrate the utility of the SHARK to make balloon-borne INP measurements whilst providing a comparison with a commercial ground-based sampler, the SHARK was sampling whilst suspended from a tethered balloon, flying roughly 20 m above the ground. The INP spectra..."

Additionally, the usefulness of the system should not change with altitude as long as the pumps maintain the flow rate through the impactors. The pumps in the SHARK were chosen because they allow the volumetric flow rate to be maintained while temperature and pressure change with altitude.

We have added the following statement to the text in Section 2.1:

"...used in reverse as a lightweight pump (~120 g). These pumps maintain the volumetric flow rate through the impactors as temperature and atmospheric pressure change with altitude. The pump for Impactor 1 was calibrated to apply this adjustment to at least 2.3 km (Leland Legacy Sample Pump: Operating Instructions, SKC), although the presence of the after-filter may reduce the battery life at this altitude. The pump for Impactor 2 is supplied by a larger battery and should be able to maintain flow to at least the same altitude as the Impactor 1 pump, and over a longer period of time. The SHARK records the volume of air sampled through Impactor 1 during the flight, and so if the pump battery was depleted, or the pressure drop became too great before Impactor 2 had finished sampling, the Impactor 1 pump would shut down and store the recorded value for later analysis. Further testing of the SHARK would be required to define a maximum altitude limit that each SHARK component could operate at. In order to provide RH, temperature..."

Specific/technical comments:

3) What do you mean by the description "short battery lives of 10s of minutes (Line 79)"?

We have re-worded the sentence to read, "...limited by relatively short battery lives, usually under 1 h, and potential..."

4) Line 264: "While the particle number concentration increases roughly linearly with size": It seems that the concentrations decrease with increasing the size.

This has been corrected.

5) Although the authors describe that "the spectra in the four locations have very different characteristics (Lines 345-346)", the spectra may also show some seasonal variations. For examples, recent field studies (e.g., Santl-Temkiv et al., Environ. Sci. Technol. 2019; Tobo et

al., Nat. Geosci. 2019; Wex et al., Atmos. Chem. Phys. 2019) show the seasonal variation of INP number concentrations in the Arctic.

We have included the following statement highlighting that there are multiple reasons for the different characteristics of the INP spectra.

"...but the size-dependence is also very different in the four locations. Due to the sample size, these variations could be attributed to the different aerosol population in each location, the time of year and meteorology, which could affect the INP concentrations and spectra (Kanji et al., 2017; Šantl-Temkiv et al., 2019; Tobo et al., 2019; Wex et al., 2019). We now discuss..."

6) Lines 358-361: Why did you speculate that the course mode INPs at this site were possibly pollen? In general, pollen grains have much higher sizes (>10 um). Also, why did you rule out the possibility of other possible sources, such as fungal spores, fertile soils, etc.?

We have adjusted this sentence to read:

"We speculate that the coarse mode INPs at this site were of biological origin, such as fungal material, pollen or bacteria with a steep INP spectrum (Kanji et al., 2017). The steepness of the curve and the temperature are consistent with ice nucleation by pollen (O'Sullivan et al., 2015; Pummer et al., 2012; Tarn et al., 2018). Although the size of whole pollen grains are often larger than 10 μ m, pollen is known to release nanoscale materials that nucleate ice, which might be internally mixed with aerosol in this size bin"

7) I couldn't find the results of ice-active site density per mass (n_m) (Line 185), while those per surface area are shown in Figure 10.

We have removed n_m from line 185.

8) The authors should provide the more detailed information on the field samplings (locations, time, periods, etc.) at at Hyytiala, Leeds, Longyearbyen, and Cardington in Section 2 and/or table, and not Section 3.

The table describing the field sampling is in the supplementary material section. We do not believe that it fits in Section 2, which is focused on the design and development of the SHARK, but have edited the text in Section 3 to point the reader towards this information:

"...Longyearbyen (Svalbard). Details of the sampling locations, periods, and instrumentation can be found in Table S1 of the Supplementary Information (SI). In this section..."

9) Figure 2 caption: "100 µL min-1" => "100 L min-1" (?)

This has been corrected.

10) Figure 5: Where were the data from Hegg et al. (1996) obtained?

We have added the location that Hegg et al. acquired their data as "Prudhoe Bay, Alaska".

11) The y axis of Figures 6, 7b, and 8: [INPs] => INP concentration (?)

Our convention is to use square brackets to represent concentration, as is done in chemistry. We have done this for previous publications (Harrison et al., 2019, Wilson et al., 2015, Vergara-Temprado et al., 2017), this is also now stated in the figure caption.

12) I would like to suggest that the authors show the comparison of INP number concentrations from stages 1d and 2d in Figure 7.

We have added this to the figure.

13) Figure 8: I would like to suggest including the information on the size ranges of "Sum of 2d & 2e" and "Sum of 1b, 1c, and 1d" in the figure legend.

We have added this to the figure.

14) As also mentioned by the authors, it seems that Figures 8 and 9 are essentially the same. I would like to suggest merging Figure 9 into Figure 8, or simply removing Figure 9 from this manuscript.

While Figure 8 is included as the standard format of reporting INP concentrations, we would like to introduce the srINP plot in figure 9 as an alternative, more intuitive means of presenting data of this sort, and so we have not changed this. We find that when the INP spectra of the different size bins fall on top of each other in Fig 8, Fig 9 makes the distinction between them a lot clearer. We have added the following text to the figure caption of figure 9:

"The data from Figure 8 is presented here in an alternative format, which has the advantage of more clearly and concisely displaying the features of the INP spectrum in each size bin than the plots in Fig 8."

15) The y axis of Figure 10: "Activated fraction" => "n_n", and "Active site density" => "n_s" (?)

This has been changed in the figure.

Response to Referee #2

We would like to thank the referee for their useful comments and have responded below. The referee comments are highlighted in red and our responses are in black.

In this paper, the authors describe the SHARK platform to measure the aerosol size and to sample the aerosols for INP analysis. The platform also deploys meteorological sensors. Size-resolved aerosol and INP measurements within the boundary layer are missing, and I think such a platform in the future can be very useful. The paper is well written. I have a few minor comments that I suggest the authors address before the paper can be published.

1) The importance of aerosol composition towards INP efficiency should be mentioned. Although size is important for transport/dispersion and residence time within the atmosphere; it should be noted that INP efficiency in addition to the size also depends upon the other factors (e.g., composition: e.g., organics vs. dust, particle type: e.g., spherical vs. non-spherical, etc.). Currently, it reads like size is the most important factor that determines the INP efficiency.

We fully agree that size is not the only determinant of ice nucleation and do not suggest it is. We have adjusted the text in the introduction to read:

"While composition is recognised to be an important controller of ice nucleation ability (Kanji et al., 2017), it has also been generally thought that the larger an aerosol particle, the more likely it is to serve as an INP (Pruppacher and Klett, 1997). However, the lifetime of coarse mode aerosol particles decreases rapidly with increasing size."

2) It should be acknowledged that the SHARK technique does not provide spatial and temporal measurements of INP.

It is implicitly acknowledged that instrument does not produce high temporal resolution, since it is a filter-based technique. However, samples from the SHARK can be used to produce a time-series of INP concentrations with a resolution on the order of hours in specific locations. Similarly, when the SHARK is deployed on a tethered balloon, there are no means of making measurements beyond the length of the tether, but the altitude of the SHARK can be controlled within that range.

3) It is not clear regarding the use of equation 3 to calculate Ns. Fice and N are determined using different techniques. It is not clear how the measurements from both techniques can be combined. N quantity (line 414) is the total number of particles, which depends upon the volume of air sampled, duration time, and some particle concentration (#/cc). Is it possible that the number of particles that enter the impactor (section 3.2) might be different than OPC (line 415) because of losses within the impactor?

The following statement has been added to the paper:

"where *N* is the total number of particles sampled by the impactor in each size bin, calculated using the number concentration in each size category as measured by the OPC, and the volume of air sampled by the impactor (see Table S1). The size bins from the OPC which have been included in the calculations were matched to those in the impactors. The bin boundaries for the OPC calculations were within tens of nanometres of the impactor bin boundaries."

In addition, we made considerable effort to quantify losses in the OPC and impactors in order to understand them, as demonstrated in Figure 3. We also plot the collection efficiencies for both impactors in Figure 2 and discuss effects such as bounce in the text.