The authors thank both reviewers for their favorable reception of the revised manuscript. There was only one open issue in the recevt version:

Original Comment: Related to the above point, I have to wonder, is the covariance profile useful beyond a correction when applying the mean averaging kernel? My (perhaps wrong) interpretation is that when the covariance profile is 0, the mean of the retrieved profile is a smoothed version of the true mean atmospheric state. I suppose what I am asking is that if the covariance profile is not 0, is it wrong to interpret the retrieved mean as a smoothed version of the true atmospheric mean? If so, I would like to see a discussion of this included in the manuscript.

Original Reply: We are not aware of any other interpretation beyond the one offered. The interpretation that zero covariance means that the mean of the retrieved profiles is a smoothed version of the true mean atmospheric state is not true, at least not in a general sense. Assume a case with infinite noise, i.e., no measurement information. The retrieval will then be identical to the a priori. Assume further that a constant (e.g. climatological) a priori has been chosen. The result will not vary at all. Thus also the covariance will be zero although the mean result is fully determined by the a priori, in shape and values.

New Comment: The author's are certainly correct with their example, but I think the confusion comes from my (mis)use of the word "smoothing". If you take Eq. 6 from the manuscript and set the covariance terms to 0, you obtain

$$\langle \hat{\vec{x}} \rangle = \langle \vec{x}_a \rangle + \langle \mathbf{A} \rangle (\langle \vec{x} \rangle - (\langle \vec{x}_a \rangle)).$$

This is the same as the standard "smoothing" equation that you would get for a single profile, with all of the quantities replaced by their means. So the point that I was trying to make is that if the covariance terms are zero there is an analogous interpretation for the mean profile compared to a single profile, which is no longer the case if the covariance terms are non-zero. I think that a statement to this effect in the manuscript would be helpful for the reader to better understand the implications of the covariance terms being zero or non-zero.

New Reply: Ok, now we've got it. The reviewer is obviously right. When writing the paper this issue was so obvious to us that we simply forgot to mention is explicitly but we agree that a statement on this should be made.

Action: We have included the following paragraph after Eq. (9):
Often these correlations are close to zero, e.g., in the case of almost linear radiative transfer. In this case, Equation (9) reduces structurally to Eq. (4), and can be re-interpreted in the sense of Eq. (5), applied to mean averaging kernels and profiles, as

$$\langle \vec{x}_{degraded} \rangle = (\mathbf{I} - \langle \mathbf{A} \rangle) \langle \vec{x}_{\mathrm{a}} \rangle + \langle \mathbf{A} \rangle \langle \vec{x}_{\mathrm{high-resolved}} \rangle. \tag{1}$$

For all other cases, i.e., when the covariance terms $\mathrm{cov}(\mathbf{A}, \vec{x}_\mathrm{a})$ and $\mathrm{cov}(\mathbf{A}, \vec{x})$ are non-zero, the respective additive corrections are necessary.

This implied another change in the following paragraph: "This" had to be replaced with "Equation (8)".

The application of mean averaging kernels to mean trace gas distributions

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Abstract. To avoid unnecessary data traffic it is sometimes desirable to apply mean averaging kernels to mean profiles of atmospheric state variables. Unfortunately, application of averaging kernels and averaging are not commutative in cases when averaging kernels and state variables are correlated. That is to say, the application of individual averaging kernels to individual profiles and subsequent averaging will, in general, lead to different results than averaging of the original profiles prior to the application of the mean averaging kernels unless profiles and averaging kernels are fully independent. The resulting error, however, can be corrected by subtraction of the covariance between the averaging kernel and the vertical profile. Thus it is recommended to calculate the covariance profile along with the mean profile and the mean averaging kernel.

1 Introduction

More often than not satellite data retrievals are constrained because the unconstrained profile retrieval on a given altitude[1] grid would lead to an ill-posed inverse problem. The constrained retrieval is more robust but the price to pay typically is, among other effects, a certain loss in vertical resolution. The effect of the constraint is characterized by the averaging kernel matrix (Rodgers, 2000).

Many applications of remotely sensed data involve comparison with independent model or independent measurement data. If these comparison data are better resolved than the remotely sensed data, the averaging kernel of the latter has to be applied to the former to make the comparison meaningful (Connor et al., 1994). Otherwise differences caused by the different altitude resolution would mask scientifically significant differences. Unfortunately, for a vertical profile of n values of an atmospheric state variable, the related averaging kernel matrix is of the size $n \times n$, that is to say, the data traffic is dominated by the averaging kernel data while the data product of interest, namely the profile, could be communicated with much less effort. Often the data users are not interested in the individual measurements but prefer to work, e.g., with monthly zonal mean profiles (e.g. Hegglin and Tegtmeier, 2011). In this case, it would be convenient if the data user could simply apply monthly zonal mean averaging kernels to their better resolved monthly zonal mean data to make them comparable to the coarser resolved zonal monthly mean measurements. Unfortunately averaging and application of the averaging kernel are not commutative. As soon as the data

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[1]We use the term altitude in a generic sense, not limited to geometric altitude but for any vertical coordinate. Alternative vertical coordinates can be, e.g., pressure or potential temperature.

and the averaging kernels covary, the application of the mean averaging kernel to mean profiles gives a different result than the application of individual averaging kernels prior to averaging. We solve this problem by providing statistically inferred covariance terms which can be used to correct the related error. In the next section we describe the theoretical framework used. As a case study, covariances applicable to trace gas profiles retrieved from MIPAS (Michelson Interferometer for Passive

5 Atmospheric Sounding, Fischer et al. 2008) measurements are inferred in Section 3. The varying importance of the covariance effect is illustrated in Section 4. Section 5 is an interlude where we investigate into pitfalls regarding the applicability of averaging kernels to comparison data, before a critical discussion of the applicability of our suggested appproach concludes the paper (Section 6).

2   The formal concept

10 We borrow the formal concept of retrieval theory from Rodgers (2000). The intended application of our study is, at worst, moderately nonlinear retrievals. That is to say, linear theory is assumed to be adequate for the characterization of the retrieval in terms of error estimation, assessment of vertical resolution, and so forth. Thus, we ignore all complication which may arise from non-linearity and thus do not discuss the retrievals in an iterative setting. Within the framework of moderately nonlinear problems our results are still applicable to the results of iterative retrievals.

15   The vertical resolution of a profile of an atmospheric state variable, e.g., temperature or the volume mixing ratio of a trace gas, with n gridpoints, is usually characterized by the averaging kernel matrix $\mathbf{A}$ of size $n \times n$. Its elements are the partial derivatives $\frac{\partial \hat{x}_i}{\partial x_j}$ of the estimated state variables $\hat{x}_i$ with respect to the true state variable x_j. While the indices i and j typically run over altitude levels of one vertical profile, the concept as such has a much wider range of applicability, e.g., horizontal averaging kernels (von Clarmann et al., 2009a), or characterization of cross-dependence of multiple species. In this study, we

20 restrict ourselves to averaging kernels of vertical profiles of single species. For a constrained retrieval of the type

$$\hat{\boldsymbol{x}} = \boldsymbol{x}_{\mathrm{a}} + \left(\mathbf{K}^T \mathbf{S}_{\mathrm{y}}^{-1} \mathbf{K} + \mathbf{R}\right)^{-1} \mathbf{K}^T \mathbf{S}_{\mathrm{y}}^{-1} \left(\boldsymbol{y} - \boldsymbol{F}(\boldsymbol{x}_{\mathrm{a}})\right), \tag{1}$$

or any equivalent formulation of it, where the $\hat{\boldsymbol{x}}$ vector represents the estimated profile, $\boldsymbol{x}_{\mathrm{a}}$ is an a priori profile, $\mathbf{K}$ is the Jacobian matrix $\frac{\partial y_i}{\partial x_j}$, ^{T} indicates a transposed matrix, $\mathbf{S}_{\mathrm{y}}$ is the measurement error covariance matrix, $\mathbf{R}$ is a regularization matrix, $\boldsymbol{F}$ is the radiative transfer function, and $\boldsymbol{y}$ is the vector of measurements (von Clarmann et al. 2003a, building largely

25 upon Rodgers 2000), the averaging kernel matrix is

$$\mathbf{A} = \left(\mathbf{K}^T \mathbf{S}_{\mathrm{y}}^{-1} \mathbf{K} + \mathbf{R}\right)^{-1} \mathbf{K}^T \mathbf{S}_{\mathrm{y}}^{-1} \mathbf{K}. \tag{2}$$

The state dependence of the averaging kernel is largely due to the state dependence of the Jacobian $\mathbf{K}$. With the averaging kernel matrix introduced above, and using the linearization

$$\boldsymbol{y} - \mathbf{F}(\boldsymbol{x}_{\mathrm{a}}) \approx \mathbf{K}(\boldsymbol{x} - \boldsymbol{x}_{\mathrm{a}}), \tag{3}$$

30 Eq. 1 can be rewritten as

$$\hat{\boldsymbol{x}} = (\mathbf{I} - \mathbf{A})\boldsymbol{x}_{\mathrm{a}} + \mathbf{A}\boldsymbol{x} \tag{4}$$

The most common application of the averaging kernel matrix is the degradation of high-resolved vertical profiles to make them comparable to poorer resolved profiles by application of the averaging kernel matrix of the poorer resolved profile to the high-resolved profile (Connor et al., 1994).

$$\boldsymbol{x}_{degraded} = (\mathbf{I} - \mathbf{A})\boldsymbol{x}_{\mathrm{a}} + \mathbf{A}\boldsymbol{x}_{\mathrm{high-resolved}} \tag{5}$$

5   Here $\mathbf{A}$ and $\boldsymbol{x}_{\mathrm{a}}$ refer to the poorer resolved profile. It goes without saying that the high-resolved profile has to be resampled on the grid on which the application of the averaging kernel is performed, and, if applicable, transformed to the same units (volume mixing ratio, number density, etc.).

For applications where the a priori profiles are all zero, as being the case for most trace gas profiles retrieved from MIPAS (von Clarmann et al., 2009b), which often is appropriate if a smoothing regularization (Steck and von Clarmann 2001, building
10   on Tikhonov 1963) is used instead of an inverse a priori covariance matrix as suggested by Rodgers (1976, 2000), this reduces to

$$\boldsymbol{x}_{\mathrm{degraded}} = \mathbf{A}\boldsymbol{x}_{\mathrm{high-resolved}}. \tag{6}$$

The same is true if for all retrievals the same altitude-constant prior is used in combination with an averaging kernel with unity row sums as associated with purely smoothing constraints.
15   Using

$$\mathrm{cov}(\mathbf{a}, \boldsymbol{x}) = \langle \mathbf{A}\boldsymbol{x} \rangle - \langle \mathbf{A} \rangle \langle \boldsymbol{x} \rangle, \tag{7}$$

calculation of, e.g., zonal averages over L profiles renders[2]

$$
\begin{aligned}
\langle \hat{\boldsymbol{x}} \rangle &= \langle (\mathbf{I} - \mathbf{A})\boldsymbol{x}_{\mathrm{a}} + \mathbf{A}\boldsymbol{x} \rangle \\
&= \langle \boldsymbol{x}_{\mathrm{a}} \rangle - \langle \mathbf{A} \rangle \langle \boldsymbol{x}_a \rangle - \\
&\quad \mathrm{cov}(\mathbf{A}, \boldsymbol{x}_{\mathrm{a}}) + \langle \mathbf{A} \rangle \langle \boldsymbol{x} \rangle + \mathrm{cov}(\mathbf{A}, \boldsymbol{x}),
\end{aligned} \tag{8}
$$

20   where

$$\mathrm{cov}(\mathbf{A}, \boldsymbol{x}) = \frac{1}{L} \sum_{l=1}^{L} (\mathbf{A}_l - \langle \mathbf{A} \rangle)(\boldsymbol{x}_l - \langle \boldsymbol{x} \rangle), \tag{9}$$

and in an analogue way for $\mathrm{cov}(\mathbf{A}, \boldsymbol{x}_{\mathrm{a}})$.

Often these correlations are close to zero, e.g., in the case of almost linear radiative transfer. In this case, Equation (8) reduces
25   structurally to Eq. (4), and can be re-interpreted in the sense of Eq. (5), applied to mean averaging kernels and profiles, as

$$\langle \boldsymbol{x}_{degraded} \rangle = (\mathbf{I} - \langle \mathbf{A} \rangle)\langle \boldsymbol{x}_{\mathrm{a}} \rangle + \langle \mathbf{A} \rangle \langle \boldsymbol{x}_{\mathrm{high-resolved}} \rangle. \tag{10}$$

---

[2]Here a caveat is in order. The average of profiles which are 'optimal' in the sense of maximum a posteriori information and where the a priori information is the same for all averaged profiles is not the optimal average. This is, because the weight of the a priori information will be too large in the average. A more thorough discussion of this issue, however, is beyond the scope of this paper.

For all other cases, i.e., when the covariance terms $\mathrm{cov}(\mathbf{A}, \boldsymbol{x}_\mathrm{a})$ and $\mathrm{cov}(\mathbf{A}, \boldsymbol{x})$ are non-zero, the respective additive corrections are necessary.

For a retrieval with $\boldsymbol{x}_\mathrm{a} = \mathbf{0}$ (or $\boldsymbol{x}_\mathrm{a}$ constant with altitude and a purely smoothing constraint) ~~this~~Equation (8) simplifies to

$$\langle \hat{\boldsymbol{x}} \rangle \quad = \quad \langle \mathbf{A}\boldsymbol{x} \rangle \tag{11}$$

5
$$= \quad \langle \mathbf{A} \rangle \langle \boldsymbol{x} \rangle + \mathrm{cov}(\mathbf{A}, \boldsymbol{x}).$$

$\mathrm{cov}(\mathbf{A}, \boldsymbol{x})$ can be approximated by $\mathrm{cov}(\mathbf{A}, \hat{\boldsymbol{x}})$ which can easily be evaluated statistically from the available results and distributed to the data user along with the mean averaging kernel $\langle \mathbf{A} \rangle$ and the mean profile $\langle \boldsymbol{x} \rangle$ and used to correct profiles of averaged comparison data. All this is valid only with some qualification. Related problems will be discussed in Section 6.

For a retrieval with constant climatological $\boldsymbol{x}_\mathrm{a}$ for the entire sample of profiles we get

10
$$\langle \hat{\boldsymbol{x}} \rangle \quad = \quad \langle (\mathbf{I} - \mathbf{A})\boldsymbol{x}_\mathrm{a} + \mathbf{A}\boldsymbol{x} \rangle \tag{12}$$
$$= \quad \boldsymbol{x}_\mathrm{a} - \langle \mathbf{A} \rangle \boldsymbol{x}_a +$$
$$+ \langle \mathbf{A} \rangle \langle \boldsymbol{x} \rangle + \mathrm{cov}(\mathbf{A}, \boldsymbol{x}).$$

For a retrieval where an individual prior $\boldsymbol{x}_a$ is used for each profile retrieval, i.e., a prior which represents the best available information on the current state not in a climatological sense, but, e.g., from independent measurements specific to each

15 measurement of the ensemble, it may also be adequate to assume

$$\mathrm{cov}(\mathbf{A}, \boldsymbol{x}) \approx \mathrm{cov}(\mathbf{A}, \boldsymbol{x}_\mathrm{a}), \tag{13}$$

i.e., that the prior information is a good representation of the true atmospheric state and variability. In this case the correction by the covariance terms becomes approximately obsolete, because

$$\langle (\mathbf{I} - \mathbf{A})\boldsymbol{x}_\mathrm{a} + \mathbf{A}\boldsymbol{x} \rangle \quad = \tag{14}$$

20
$$\langle \boldsymbol{x}_\mathrm{a} \rangle - \langle \mathbf{A} \rangle \langle \boldsymbol{x}_a \rangle - \mathrm{cov}(\mathbf{A}, \boldsymbol{x}_a)$$
$$+ \langle \mathbf{A} \rangle \langle \boldsymbol{x} \rangle + \mathrm{cov}(\mathbf{A}, \boldsymbol{x}) \quad \approx$$
$$\langle \boldsymbol{x}_\mathrm{a} \rangle - \langle \mathbf{A} \rangle \langle \boldsymbol{x}_a \rangle - \mathrm{cov}(\mathbf{A}, \boldsymbol{x}_a) +$$
$$\langle \mathbf{A} \rangle \langle \boldsymbol{x} \rangle + \mathrm{cov}(\mathbf{A}, \boldsymbol{x}_\mathrm{a}) \quad =$$
$$\langle \boldsymbol{x}_\mathrm{a} \rangle - \langle \mathbf{A} \rangle \langle \boldsymbol{x}_a \rangle + \langle \mathbf{A} \rangle \langle \boldsymbol{x} \rangle$$

25 For retrievals performed in the log-space, all this becomes slightly more complicated (e.g., Stiller et al., 2012). Equation (5) then reads

$$\boldsymbol{x}_\mathrm{degraded} = \exp\left((\mathbf{I} - \mathbf{A}) \ln \boldsymbol{x}_\mathrm{a} + \mathbf{A} \ln \boldsymbol{x}_\mathrm{high\text{-}resolved} \right), \tag{15}$$

where $\mathbf{A}$ is $\frac{\ln \hat{\boldsymbol{x}}_i}{\ln \boldsymbol{x}_j}$. For log retrievals there is no obvious way to correct for the averaging artefacts as long as the averaging is performed linearly in the vmr-space. Since averaging of logarithmic rerievals in the logarithmic domain has its own problems

30 (Funke and von Clarmann, 2012), we do not pursue this option any further.

4

The issues discussed in this Section have to be considered if mean averaging kernels are to be applied to mean profiles in the spirit of Eq. 5, in order to make mean profiles of different sources comparable.

3   Covariances

The covariances between the averaging kernel matrices and the state vectors are calculated as

5
$$\mathrm{cov}(\mathbf{A}, \boldsymbol{x}) = \tag{16}$$
$$\frac{1}{L}\left(\sum_{l=1}^{L}(\mathbf{A}_l - \langle \mathbf{A}\rangle)(\boldsymbol{x}_l - \langle \boldsymbol{x}\rangle)\right) =$$
$$\frac{1}{L}\left(\sum_{l=1}^{L}\mathbf{A}_l\boldsymbol{x}_l - \frac{1}{L}\sum_{l=1}^{L}\mathbf{A}_l\sum_{l=1}^{L}\boldsymbol{x}_l\right),$$

where L denotes the sample size; we divide by L instead of $L-1$ because the latter would entail an inconsistency with Eq. (8) and Eqs. (11–14). The formulation in the lowermost line of Eq. (16) is computationally more efficient. For our case
10   study, averaging kernel matrices and state vectors retrieved from limb emission spectra measured by the MIPAS are used. The general processing scheme is described by von Clarmann et al. (2003b, 2009b). We study covariances for MIPAS O_3 and HCN profiles (Laeng et al. 2018 and Glatthor et al. 2015, respectively).

To illustrate the relevance of the correction terms, we also present the normalized covariance term $\tilde{r}$ for each profile element:

15   $$\tilde{r}_n = \mathrm{cov}(\mathbf{A}, \boldsymbol{x})_{\mathrm{n}}/(\langle \mathbf{A}\rangle\langle \boldsymbol{x}\rangle)_{\mathrm{n}} \tag{17}$$

where index n runs over the profile elements. The ˜ symbol shall avoid confusion with the product moment correlation coefficient established by Pearson (1895) for which r is often used as a symbol and which is widely used for normalization of covariances but which causes some headache when applied to correlations of matrices with vectors. For simplicity, we still call the normalized covariance 'correlation', however without claiming equivalence with its scalar counterpart.

4   Results

Case studies have been performed using ozone and HCN vertical profiles retrieved from MIPAS measurements of February 10, 2009. The test data set consists of 1385 geolocations. This day was characterized by a significantly disturbed Arctic vortex. Figure 1 shows the covariances between the profiles and the averaging kernel matrices of ozone globally (black solid line) and for various latitude bands of different size (dashed and dotted lines). In general the values are largest at the extreme ends of the
25   profiles, where the effect of the constraint on the retrieved profile is typically largest.

These results suggest that for MIPAS ozone in the middle and upper stratosphere the effect studied here can be safely ignored. Problems are limited to the upper troposphere and lower stratosphere and the mesosphere. The relevance of this effect can better be judged on the basis of the correlation profiles (Fig. 2). From 20 to about 60 km the effect is negligibly small for
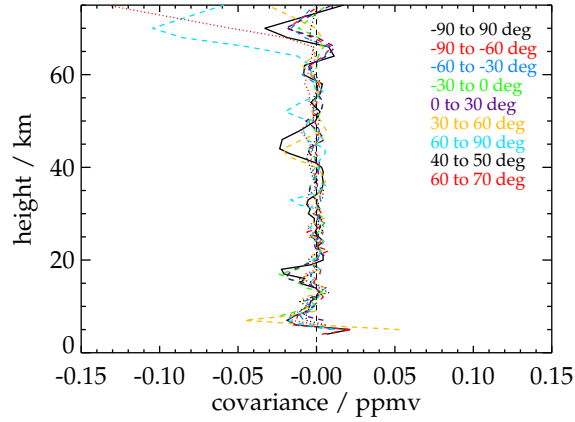
Figure 1. : Covariance of the averaging kernel and ozone mixing ratio for various latitude bands. The black solid line refers to global data. The dashed lines refer to 30-degrees latitude bands, and the dotted lines to 10-degrees latitude bands.
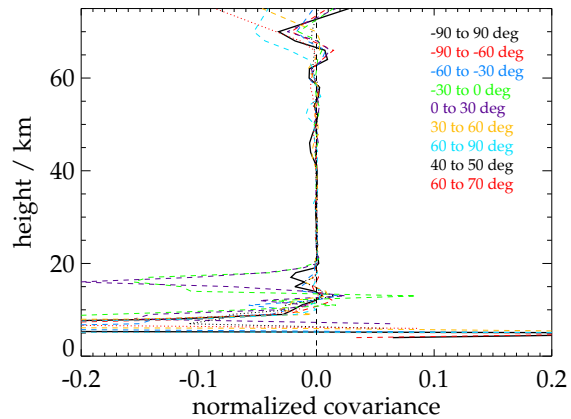

Figure 2. : Correlation of the averaging kernel and ozone mixing ratio for various latitude bands.
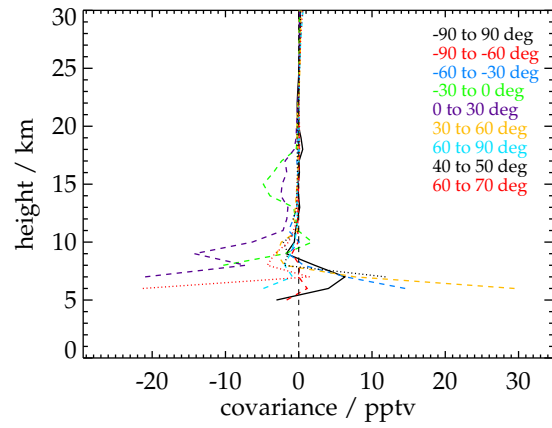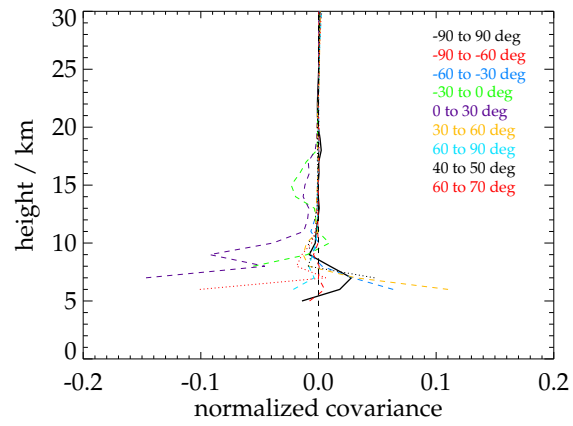
Figure 3. : As Fig 1, but for HCN.

Figure 4. : As Fig 2, but for HCN.

all latitude bands investigated in this case study. Only at the uppermost and lowermost altitudes the effect becomes relevant. The large effects at lower altitudes are simply caused by normalization of the original covariances by low ozone mixing ratios.

To study HCN is particularly interesting in the tropical upper troposphere and lower stratosphere. This is because HCN has tropospheric sources, and its pathway into the stratosphere is a particular research issue. The covariance effects can exceed 10% (violet and yellow dashed lines) and thus need to be considered when mean profiles are used for quatitative analysis and mean averaging kernels are applied.

These case studies are not meant to be representative for other gases or other instruments. Instead, they are shown to give a flavour about the order of magnitude this kind of effect can reach. Unless $\mathrm{cov}(\mathbf{A}, \boldsymbol{x})$ can be shown to be small, we recommend to use this covariance term for an additive correction when mean averaging kernels are applied to averaged comparison data.

5 An Important Side Remark

The issue of the limited applicability of averaging kernels to independent comparison data deserves awareness. When averaging kernels of a measurement are applied to better resolved comparison data, it is almost always tacitly assumed that the atmospheric state represented by the measurement is the same as that of the comparison data and that thus the averaging kernel of the measurement can be safely applied to the comparison data. However, since averaging kernels are in general state-dependent, a caveat is in order.

Application of the formalism of Connor et al. (1994) (our Eq. 5) has its own specific problems which fully apply to our proposed scheme. The application of the averaging kernel matrix of a poorly resolved profile $\boldsymbol{x}_{\mathrm{coarse}}$ to a better resolved profile $\boldsymbol{x}_{\mathrm{fine}}$ is only adequate if both data sets describe approximately the same atmospheric state, i.e., if the one profile is in the linear domain of the other. That is to say, if the same Jacobians apply to both profiles. Otherwise it would be necessary to construct an averaging kernel using the Jacobian $\mathbf{K}$ evaluated for the atmospheric state represented by the profile $\boldsymbol{x}_{\mathrm{fine}}$ but with the measurement covariance matrix $\mathbf{S}_{\mathrm{y}}$ and the regularization matrix $\mathbf{R}$ corresponding the the retrieval producing $\boldsymbol{x}_{\mathrm{coarse}}$. Within linear theory, the state-dependence of the Jacobian $\mathbf{K}$ and with this the state dependence of the averaging kernel matrix $\mathbf{A}$ are often ignored. To do so is justifiable as long as the profiles to be intercompared are sufficiently similar. In this case the comparison will show reasonable agreement.

If, in turn, the profiles are very different, two components contribute to the disagreement seen after application of the Connor-method. First the genuine difference of the profiles, and second any artefact caused by the inadequate averaging kernels. Thus, in the logic of a testing scheme, good apparent agreement hints at good genuine agreement even more, because it is extremely unlikely that genuine differences which would survive the application of the Connor-method with the correct averaging kernel are 'convolved away'[3] with the averaging kernel evaluated for the wrong atmosphere.

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[3]We put this term in quotes to highlight that this method is, mathematically speaking, not a convolution. It is not even a numerical approximation to a convolution.

6 Discussion and Conclusion

We have identified the problem that it is not generally allowable to apply mean averaging kernels to mean atmospheric profiles in situations where the averaging kernels and the profiles covary. The relevance of this effect, however, depends on the instrument, the species, the latitude band and the altitude under investigation. To solve this problem, we have proposed a statistical correction scheme which involves the covariance between the averaging kernel and the profile. With this correction in place, the scheme suggested by Connor et al. (1994) to make better resolved vertical profiles of atmospheric state variables comparable to coarser resolved ones can be applied also to averaged profiles.

For data producers who distribute, besides their original retrievals, also zonal mean data or similar data products, we recommend the following: Along with the generation of zonal mean data and averaging kernels, the correlation profiles should be calculated. Compared to averaging kernels and covariance matrices they need negligible storage and cause negligible data traffic. In cases when zonal mean data have already been generated but when mean covariance matrices and covariance profiles are not available, the huge I/O load associated with reading all individual averaging kernels may be prohibitive. In these cases one might consider to estimate the mean averaging kernel and the covariance profile on the basis of a limited random sample out of the measurements which went into the zonal mean.

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