



## Validation of TROPOMI Surface UV Radiation Product

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**Abstract.** The TROPospheric Monitoring Instrument (TROPOMI) onboard the Sentinel-5 Precursor (S5P) satellite was launched on 13 October 2017 to provide the atmospheric composition for atmosphere and climate research. The S5P is a sun-synchronous polar-orbiting satellite providing global daily coverage. The TROPOMI swath is 2600 km wide, and the ground resolution for most data products is 7.2x3.5 km<sup>2</sup> (5.6x3.5 km<sup>2</sup> since 6 August 2019) at nadir. The Finnish Meteorological Institute (FMI) is responsible for the development and processing of the TROPOMI Surface Ultraviolet (UV) Radiation Product which includes 36 UV parameters in total. Ground-based data from 25 sites located in arctic, subarctic, temperate, equatorial and antarctic areas were used for validation of TROPOMI overpass irradiance at 305, 310, 324 and 380 nm, overpass erythemally weighted



dose rate / UV index and erythemally weighted daily dose for the period from 1 January 2018 to 31 August 2019. The validation results showed that for most sites 60–80% of TROPOMI data was within  $\pm 20\%$  from ground-based data for snow free surface conditions. The median relative differences to ground-based measurements of TROPOMI snow free surface daily doses were within  $\pm 10\%$  and  $\pm 5\%$  at two thirds and at half of the sites, respectively. At several sites more than 90% of clear sky TROPOMI data were within  $\pm 20\%$  from ground-based measurements. Generally median relative differences between TROPOMI data and ground-based measurements were a little biased towards negative values, but at high latitudes where non-homogeneous topography and albedo/snow conditions occurred, the negative bias was exceptionally high, from -30% to -65%. Positive biases of 10–15% were also found for mountainous sites due to challenging topography. The TROPOMI Surface UV Radiation Product includes quality flags to detect increased uncertainties in the data due to heterogeneous surface albedo and rough terrain which can be used to filter the data retrieved under challenging conditions.

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## 1 Introduction

The Tropospheric Monitoring Instrument (TROPOMI) is a nadir-viewing imaging spectrometer measuring in the ultraviolet, visible, near-infrared, and the shortwave infrared wavelengths onboard the Sentinel-5 Precursor (S5P) polar-orbiting satellite. The S5P was launched on 13 October 2017 as part of the EU Copernicus programme to monitor atmospheric composition with nominal life time of seven years. The mission is a cooperative undertaking between the European Space Agency (ESA) and the Netherlands. The S5P satellite is on a sun-synchronized afternoon orbit with an ascending node equatorial crossing at 13:30, which provides global daily observations of the sunlit part of the Earth for air quality and climate applications. The S5P is the first Copernicus mission dedicated to atmospheric observations and it will be complemented by Sentinel 4 with geostationary orbit and Sentinel 5 on sun-synchronous morning orbit with planned launches in the coming years. The TROPOMI Level 2 data products include information of aerosols, carbon monoxide, clouds, formaldehyde, methane, nitrogen oxide, sulphur dioxide, ozone and surface ultraviolet (UV) radiation. Other products are generated within the Copernicus ground system, while the surface UV radiation is generated through the Finnish Sentinel collaborative ground segment.

Solar UV radiation at short wavelengths (280–400 nm) is harmful for the whole ecosystem including humans, animals, plants, aquatic environments and materials (e.g., EEAP, 2019, and references therein). For humans the well known harmful effects of UV radiation are sunburns and other skin problems, increased risk of skin cancer and cataract, premature aging of the skin and weakening of the immune system (EEAP, 2019). On the other hand UV radiation initiates vitamin D production in the skin (Webb, 2006) and has many more positive effects (Juzeniene and Moan, 2012). The ozone layer in the stratosphere protects the Earth from the most dangerous UV wavelengths by absorbing the shortest part of the spectrum. In the late 1970s the ozone layer was found to decrease at an alarming speed above Antarctica (Farman et al., 1985; WMO, 1990). Later, also in the Arctic the total ozone was found to decrease in the spring and ozone trends turned to negative at mid-latitudes (WMO,



1999). The international Montreal Protocol was signed in 1987 to restrict ozone depleting substances, which has led to the start  
40 of the recovery of the ozone layer (WMO, 2018). Despite the slow ozone recovery at middle latitudes and in the polar regions,  
ozone loss occurrences can still be experienced related to circulation patterns, which in turn affect solar UV levels at ground  
(Petkov et al., 2014; Fountoulakis et al., 2020).

Ground-based UV monitoring started to increase in the late 1980s to respond to the concerns about increased surface UV  
levels due to the depleting ozone layer (Solomon et al., 1986). However, the ground-based UV monitoring network is sparse  
45 from a global point of view and many places are not covered. The advantage of retrievals from space is that satellites provide  
global coverage of biologically effective UV parameters. The disadvantage is that for polar-orbiting satellites there is only one  
overpass per day for most sites. However, daily doses can be estimated using combination of radiative transfer calculations and  
measurements from satellite instrument during the overpass (e.g., Kalliskota et al., 2000; Tanskanen et al., 2007).

The Finnish Meteorological Institute is responsible for the development, processing and archiving of the TROPOMI Surface  
50 UV Radiation Product, which continues UV records started by NASA Total Ozone Mapping Spectrometer (TOMS) instrument  
in 1978 (Eck et al., 1995; Krotkov et al., 2001) and followed by the Dutch-Finnish Ozone Monitoring Instrument (OMI)  
onboard NASA's Aura satellite launched in 2004 (Levelt et al., 2006; Tanskanen et al., 2006). Compared to the preceding  
instruments, TROPOMI has an increased spatial resolution with a swath of 2600 km including 450 across-track pixels. The  
ground resolution of the UV product was 7.2x3.5 km<sup>2</sup> at nadir until 6 August 2019, and is 5.6x3.5 km<sup>2</sup> since then. The  
55 TROPOMI Surface UV Radiation Product responds to the increasing need of information regarding the tropospheric chemistry  
and biologically active wavelengths of the solar spectrum reaching the surface. In this paper, overpass irradiances at 305, 310,  
324 and 380 nm, overpass erythemally weighted dose rates / UV index and daily doses are validated against well maintained  
and calibrated ground-based spectroradiometer, broadband and multichannel radiometer measurements from 25 sites.

## 2 Data

### 60 2.1 TROPOMI surface UV radiation product

The TROPOMI surface UV algorithm is explained in detail in Lindfors et al. (2018) and Kujanpää et al. (2020). It is based on  
the heritage of the surface UV algorithms for the TOMS (Eck et al., 1995; Krotkov et al., 2001; Herman et al., 2009), the OMI  
(Levelt et al., 2006; Tanskanen et al., 2006; Arola et al., 2009) and the offline UV product (OUV) of the EUMETSAT Satellite  
Application Facility on Atmospheric Composition Monitoring (AC SAF) (Kujanpää and Kalakoski, 2015). Satellite surface  
65 UV products are based on radiative transfer modelling using as main inputs: solar zenith angle (SZA), total ozone column,  
cloud optical depth, aerosol optical properties, surface pressure and surface albedo. For the TROPOMI product, the VLIDORT  
radiative transfer model (Spurr, 2006) is used for the radiative transfer calculations.

The TROPOMI UV algorithm is based on two pre-computed lookup tables (LUT) in order to save computing time compared  
to runtime radiative transfer calculations. The first LUT is used to retrieve the cloud optical depth from the measured 354 nm  
70 reflectance using SZA, viewing zenith angle, relative azimuth angle, surface pressure and surface albedo as other inputs. The  
measured 354 nm reflectance together with the angles and surface pressure are obtained from the TROPOMI L2 aerosol index



(AI) product (Stein Zweers, 2018) while the surface albedo is obtained from the surface albedo climatology used in the AC SAF OUV product (Kujanpää and Kalakoski, 2015). More details can be found in Sect. 3.3 of Lindfors et al. (2018). The second LUT stores the irradiances and dose rates as a function of total ozone column, surface pressure, surface albedo, cloud  
75 optical depth and SZA. The irradiances and dose rates are obtained by Lagrange polynomial interpolation using the total ozone column from the offline version of the TROPOMI L2 total ozone column product (Garane et al., 2019). Surface albedo and pressure are the same as in the first step. The SZA is either the overpass time value or calculated for the solar noon time. A post-correction for the effect of absorbing aerosols based on an aerosol climatology by Kinne et al. (2013) is applied to the irradiances (Arola et al., 2009). The correction for erythemal and vitamin D synthesis weightings is the same as for the 310 nm  
80 irradiance. A correction for the variation in Sun-Earth distance is also applied in the post-processing step.

The TROPOMI L2 UV product contains 36 UV parameters in total (Table 1), including irradiances at four different wavelengths and dose rates for erythemal (Commission Internationale de l'Eclairage, 1998) and vitamin D synthesis (Bouillon et al., 2006) action spectra. All parameters are calculated for overpass time, solar noon time, and for theoretical clear sky conditions with no clouds or aerosols. Daily doses and accumulated irradiances are also calculated by integrating over the sunlit part of  
85 the day. As the cloud optical depth is retrieved at the overpass time, the uncertainties in the final cumulative product (daily dose and accumulated irradiances) increases especially for changing cloudiness. For rapidly changing cloudiness the effect is seen also in noon parameters. In addition to UV parameters, quality flags related to the UV product and processing are generated (Kujanpää, 2020). The processing quality flags are a standard set included in all TROPOMI L2 products while the product quality flags are specific to the surface UV product. A continuous overall quality value number ( $UVQAV \in [0,1]$ , over 0.5  
90 representing the most reliable data) computed from the product quality flags indicates increasing product quality and can be used together with the quality flags to filter out problematic data.

The level 2 data are stored in netCDF-4/HDF5 format. One file is *ca.* 250 MB (190 MB before 6 Aug 2019) in size. UV product version 1.02.02 was used for the current study. The input total ozone and aerosol index files were collected from the reprocessed and offline data in order to construct as homogeneous a time-series as possible. However, the total ozone product  
95 version varies from 1.01.02 (starting from 7 Nov 2017) to 1.01.05 (15 Apr 2018) to 1.07.07 (30 Apr 2018) while the aerosol index product version goes from 1.00.01 (7 Nov 2017) to 1.02.02 (15 Apr 2018) to 1.03.01 (30 Apr 2018) to 1.03.02 (27 Jun 2019).

To facilitate the validation work, ground station overpass text files containing the UV parameters and supporting input and quality flag data were extracted from the large L2 files. The extractor (version 1.02.00) computes the great-circle distance  
100 between the ground station and TROPOMI pixel centre coordinates using the haversine formula and the Earth radius at the ground station coordinates. When the great-circle distance is smaller than a pre-defined limit, here set to 10 km, the data for the TROPOMI ground pixel are stored. No interpolation between the ground pixels is performed.



**Table 1.** TROPOMI surface UV parameters

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Overpass and solar noon time irradiance at 305, 310, 324 and 380 nm [ $\text{W}/\text{m}^2/\text{nm}$ ]
Overpass and solar noon time dose rate for erythemal and vitamin D synthesis action spectra [ $\text{W}/\text{m}^2$ ]
Daily accumulated irradiances at 305, 310, 324 and 380 nm [ $\text{J}/\text{m}^2$ ]
Daily dose for erythemal and Vitamin D synthesis action spectra [ $\text{J}/\text{m}^2$ ]
Overpass and solar noon time UV index (dimensionless)
+ all parameters for clear sky conditions (no clouds nor aerosols)
+ quality flags (UV product and processing flags, and UV quality value (UVQAV))

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## 2.2 Ground-based reference data

The TROPOMI surface UV radiation product is compared with ground-based UV measurements. The location and characteristics of the sites are shown in Fig. 1 and Table 2 in which they are listed from North to South. The sites were chosen to represent different latitudes, longitudes and topography. The sites are located in arctic, subarctic, temperate, equatorial and antarctic areas including inland, coastal and mountainous stations. At few stations, snow occurs during some period of the year. Ground-based UV measurements are performed using spectroradiometers, broadband and multiband radiometers. The instrumentation and its characteristics are shown in Tables 3 and 4. Many of the spectroradiometers have participated in on-site quality assurance of spectral solar UV measurements performed by the traveling reference spectroradiometer QASUME since 2002 (Gröbner et al., 2005). The reports of the site visits can be found at <https://www.pmodwrc.ch/en/world-radiation-center-2/wcc-uv/qasume-site-audits/>.

The Norwegian UV Monitoring Program includes UV measurements at 9 sites throughout Norway. It is a cooperation between the Norwegian Radiation and Nuclear Safety Authority (DSA), Norwegian Institute for Air Research (NILU) and the University of Oslo. Four sites were chosen for this study based on their latitude and topography. Ny-Ålesund is the northernmost site and located in Svalbard. Measurements from the GUV-instrument reveals snow cover typically from mid of September to early July (albedo  $>0.2$ ). The seasonal maximum albedo is 0.8, but during the later years the albedo is now 0.5-0.6. Andøya is located at the tip of a long island, locally influenced by snow in winter and spring. The sea around the site is usually open. Finse is a mountain village at an altitude of 1200 m, close to the Hardanger-Jøkulen glacier. Measurements from the GUV instrument reveal snow cover typically lasting from 20 September to mid of July (albedo  $>0.2$ ), but the timing of the melting season may be shifted by  $\pm 1$  month (2015 and 2018), interspersed with periods with wet snow (end of April 2019). The maximum albedo exceeds 0.90. Blindern is located at the suburban area of the city of Oslo. At all sites, the cloudiness is characterized by rapidly moving clouds. The network is equipped with GUV multifilter radiometers which measure UV irradiance at five channels as one minute averages. The data is used to retrieve the UV index and UV dose rates using several action spectra (Bernhard et al., 2005; Johnsen et al., 2002, 2008) and is freely available at <https://github.com/uvnrpa/>. The quality assurance of the network includes transfer of the absolute calibration using a regularly calibrated traveling reference. The data is corrected for drift and



for angular dependency. Intercomparisons of UVI against the QASUME reference (2003, 2005, 2009, 2010, 2014, 2019) show an interquartile range within  $\pm 5\%$  for all GUV instruments and campaigns performed within the period 2003-2019.

The FMI performs spectral UV measurements with Brewer spectroradiometers in the South of Finland in Helsinki and in the  
130 North in Sodankylä. The spectral time series of Sodankylä is one of the longest in the Arctic (Lakkala et al., 2003). The site in Helsinki is located in the vicinity of the city centre, but characterized by urban green area. The measurements are performed at the roof of the FMI main building and the horizon is free except in the North side. The weather is characterized by convective cloudiness in summer afternoons and humid winters. UV measurements in Sodankylä are part of the research infrastructure of the Arctic Space Centre. The research centre is located 5 km from the village by the river Kitinen and surrounded by swamps  
135 and boreal forest. Snow occurs from October to April/May. Temperatures can reach  $-40^{\circ}\text{C}$  and  $+30^{\circ}\text{C}$  in winter and summer, respectively. The Sun is below the horizon for a couple of weeks during winter, and stays above the horizon during a couple of weeks around mid-summer. The FMI Brewer spectroradiometers are calibrated every second or third month using 1 kW lamps in the laboratory (Lakkala et al., 2016). The primary calibration lamps are calibrated yearly at the National Standard Laboratory MIKES-Aalto (Heikkilä et al., 2016; Kübarsepp et al., 2000). The quality assurance of the measurements includes  
140 corrections for temperature dependence and cosine error (Lakkala et al., 2008; Mäkelä et al., 2016; Lakkala et al., 2018) and data are submitted to the European UV data base (Heikkilä et al., 2016). Data is regularly compared to the QASUME reference and differences of less than 6% have been found for wavelengths  $> 305$  nm (Lakkala et al., 2008).

The Royal Meteorological Institute of Belgium operates two Brewer spectrophotometers on the roof of its building at Uccle, a residential suburb of Brussels about 100 km from the shore of the North Sea. The climate is influenced by the Gulf stream  
145 with mild winters and warm summers. Cloudiness is most of the time variable. The measurements of the Brewer no. 178 were used in this study. It is a double monochromator Mk III which was installed in September 2001. The raw UV counts are converted to counts per second and corrected for instrument dead time, dark count and temperature. Brewer measurements are calibrated with 50 W tungsten halogen lamps on a monthly basis and with 1 kW lamps during less frequent but regular intercomparisons. The instruments were also compared with the traveling QUASUME (Quality Assurance of Spectral UV  
150 Measurements in Europe) unit in 2004 (Gröbner et al., 2006).

The Laboratoire d'Optique Atmosphérique (LOA) performs spectral UV measurements with Bentham spectroradiometers at three French sites, in metropolitan and overseas regions (Brogniez et al., 2016). The first site, Villeneuve d'Ascq (VDA), is a semi-urban site located in a flat region of the North of France close to Lille. It is characterized by an oceanic midlatitude climate (warm summers, mild humid winters). The second site, Observatory of Haute-Provence (OHP), is a rural mountainous  
155 site located in the French Southern Alps. It is characterized by a mountainous Mediterranean climate (warm summers, harsh winters). The third site, Saint-Denis (OPA) is a coastal urban site located on the Moufia campus in the small mountainous island of La Réunion in the Indian Ocean. This environment leads to frequent occurrence of orographic clouds forming in early afternoon especially in summer. OPA is characterized by a tropical climate (hot-humid summers, mild-warm winters). At the tropical site UV radiation level in summer is much higher around noon than at the two metropolitan sites due to a higher  
160 sun elevation and a lower total ozone column. Note that, at VDA and OHP sites, absorbing aerosols are present, and need to be accounted for in satellite UV algorithms (Arola et al., 2009). Due to its oceanic and mountainous surroundings, OPA



is a challenging site for satellite validation, since there might be a large spatial variability of cloud cover and surface type in a satellite pixel. The three instruments are affiliated with NDACC (Network for the Detection of Atmospheric Composition Change), thus to meet the requirements of this network they are calibrated every 2-4 months with 1kW lamps traceable to  
165 National Institute of Standards and Technology (NIST) and the measurements are corrected from wavelength misalignment and cosine response. Following Bernhard and Seckmeyer (1999), the expanded uncertainties ( $k=2$ ) are 5.3% at VDA and OHP and 5% at OPA. At OHP and OPA global irradiance measurements are available every 15 min. At VDA, scans are performed every 30 min. Spectroradiometer's data have been already used for OMI validation (Buchard et al., 2008; Brogniez et al., 2016).

Central European mountainous sites are Davos in Switzerland and Aosta in Italy. Both sites are located in the Alps: Aosta  
170 (570 m a.s.l.) being located in a large valley floor with a wide field of view, surrounded by mountains (as high as 3500 m a.s.l.), and Davos a mountainous site stretching from around 1500 m a.s.l. to just above 3000 m. a.s.l. in altitude. UV measurements in Aosta are maintained by ARPA and performed with a Bentham DTMc300 spectroradiometer, which is calibrated on a monthly basis using a set of three 200 W lamps, recently complemented with a setup including two 1 kW lamps. The spectroradiometer is additionally compared to the world calibration reference QASUME every second year. Average differences are generally  
175 within  $\pm 2\%$ , with a diurnal variability below 4%. The whole dataset has been subjected to QA/QC and has been recently re-evaluated and homogenized. The expanded uncertainty for wavelengths above 305 nm and SZAs below 70 degrees is 4%. For larger SZAs and shorter wavelengths the uncertainty is larger, reaching 11% at 300 nm when SZA is 75 degrees. The spectroradiometer is preprogrammed to take measurements every 15 minutes. Weather in Aosta is characterized by warm summer, when convective clouds usually develop along the mountain slopes, and dry winter. Snowfalls occur at the station  
180 during some winter days in December-March, while the mountains around the station are covered by snow for most part of the year (October-June).

The measurements of the Swiss site Davos are part of the Physikalisch-Meteorologisches Observatorium Davos, World Radiation Centre (PMOD-WRC). They include spectroradiometer measurements performed with the World reference spectro-  
185 radiometer QASUMEII and with the double Brewer #163 using an optimized diffuser (Gröbner, 2003). The spectral solar UV irradiance measurements are traceable to the SI using a set of transfer standards (Gröbner and Sperfeld, 2005). The expanded uncertainty of the spectral solar UV irradiance measurements ( $k=2$ ) is 1.7% for overcast situations (diffuse sky), and 2.0% for clear sky situations (Hülse et al., 2016). In addition, 5 broadband UV radiometers (SL1492, SL3860, SL1492, YES010938, KZ 560) measure solar UV irradiance continuously at the site. The average of these measurements is used in this study. The estimated expanded uncertainty ( $k=2$ ) throughout the year for clear sky measurements of these radiometers is 3.6%, while for  
190 all sky conditions the expanded uncertainty is increased to 6.6% due to the increased uncertainty for broken cloud conditions and the corresponding uncertainty of the angular response cosine correction applied to the radiometers. In Davos, mountains limit the field of view so the diffuse radiation is reduced approximately by 5%. There is snow cover from November to March.

Rome and Thessaloniki are both urban sites at the Mediterranean coast. The climate at both sites is characterized by mild humid winters and warm dry summers. Both sites are occasionally under the influence of Saharan dust (Amiridis et al., 2005;  
195 Gobbi et al., 2019), which is seen as increased aerosol concentration. The aerosol load can also be increased due to pollution (Fountoulakis et al., 2019). In the summer, most of the days are sunny. In Thessaloniki measurements are performed by



the Aristotle University of Thessaloniki with a Brewer MKIII spectroradiometer. The quality assurance of the measurements include: 1 kW lamp calibrations traceable to PTB, temperature and cosine correction (Bais et al., 1998; Garane et al., 2006; Fountoulakis et al., 2017). Detailed information on data quality control and analysis can be found in Fountoulakis et al. (2016).

200 The measurements of Rome are maintained by Sapienza Università di Roma and are performed with a Brewer spectroradiometer. UV irradiance and total ozone content have been measured since 1992 at Rome by the Brewer Mk IV spectrophotometer No. 067. The overall performance of Brewer 067 has been controlled every 2 years since 1992 by the intercomparison with the traveling standard reference Brewer 017 from International Ozone Services Inc. (IOS), (Siani et al., 2018). The last calibration was performed in July 2019 and the UV calibration was completed using IOS 1 kW lamp. UV data are processed  
205 using cosine and temperature correction. The instrument was also compared with the traveling spectroradiometer QUASUME unit during the UV intercomparison campaign in Arosa (Switzerland) in 2012. UV measurements are taken every 30 minutes.

Measurements at the Israeli sites, Bet-Dagan, Jerusalem and Eilat are maintained by the Israel Meteorological service (IMS). Bet-Dagan station is located in open shrublands near Tel-Aviv metropolis on the coast of the Mediterranean Sea. It is characterized by hot and humid summers and mild winters. The city of Jerusalem is located on the Judean Mountains with hot and  
210 dry summers and cold winters. Most of the rain occurs between October and May. Eilat is located on the north coast of the Red Sea surrounded by the mountains of Eilat. The climate there is typical for deserts with hot and arid conditions, the maximum temperature in summer are often over 40°C with constant clear skies conditions between June and September. Winter is also relatively hot with maximum temperatures around 20°C and with an annual average precipitation of 25 mm. UV index is monitored every minute by calibrated Yankee Environmental Systems (YES) UVB-1 radiometers, and the data is saved as 10  
215 minute averages.

The Izaña Atmospheric Observatory is a high-mountain station located on the island of Tenerife (Canary Islands, Spain; 2373 m a.s.l.). The observatory is thus located in the region below the descending branch of the Hadley cell, typically above a stable inversion layer and on an island far away from any significant industrial activities. This ensures clean-air and clear-sky conditions all year. This predominant meteorological conditions of trade wind inversion give rise to the presence of a dense  
220 stratocumulus layer of clouds lying below the observatory (García et al., 2016). The surroundings of the observatory is characterized by low bushes and rocks (García et al., 2019). The UV measurements reported are performed with a Brewer no. 183 from the European Brewer Calibration Centre (RBCC-E) maintained by the Spanish State Meteorological Agency (AEMET). The RBCC-E triad is calibrated annually from of 1 KW (NIST traceable) lamps used the observatory facilities (Guirado et al., 2012). The UV response of each instrument is checked regularly used a 200W portable lamp system (Sierra Ramos, 2012). In  
225 addition, during the RBCC-E campaign, the travelling reference Brewer no. 185 is compared every year with the QASUME unit from PMOD-WRC (Egli, 2019; Gröbner et al., 2006). The comparison has been shown to be within 2% with a daily variation of less than 5%. Then, in the Izaña Observatory, the UV measurements of Brewer no. 183 and no. 157 are intercompared with those obtained by the Brewer no. 185 to check its calibration. The difference between Brewer no. 183 used in this comparison and Brewer no. 185 is around 1%.

230 The University of La Réunion monitors UV radiation with Kipp&Zonen UVS-E-T radiometers at four sites: Mahé - Seychelles, Antananarivo - Madagascar, Anse Quitor - Rodriguez, and Saint-Denis - Reunion Island. The stations are part of the



UV-Indien network. The objective of this network is to monitor and study UV radiation over on the southwestern basin of the Indian Ocean. This region has very few measurements of solar UV irradiance and shows extreme UV Index (UVI) throughout the year. In the context of climate change, this region of the world (southern hemisphere tropics) could be affected by a decrease in ozone and an increase in UVR levels through the 21st century (Lamy et al., 2019). UV-Indien measurement sites correspond to various environments (seaside, altitude, urban) and are homogeneously distributed throughout the Western Indian Ocean. These radiometers are calibrated every 2 years, either at the WRC Davos Switzerland, or directly from the measurements of the Bentham DM300 spectroradiometer installed on the site of the University of la Réunion Island and managed jointly with the University of Lille (see the section on the sites of the University of Lille for a description of the Moufia site). The more recent instruments (MAH, ANT and ROD) used the manufacturer's calibration. Raw data are corrected according to the calibration. The calibration coefficient depends on the SZA and the ozone total column. For the ozone total column, the OMI total ozone column OMTO3 product is used.

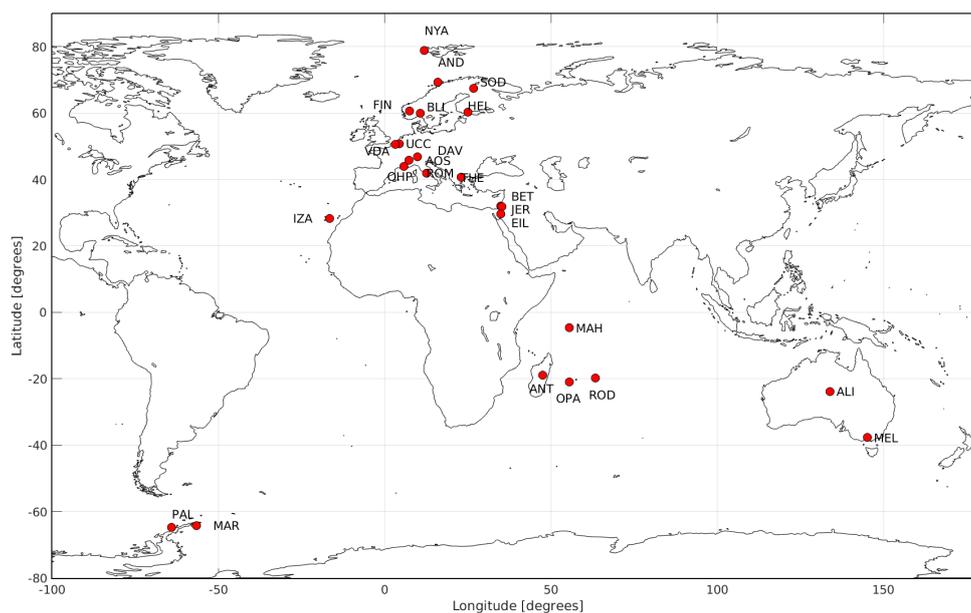
The Australian sites, Alice Springs and Melbourne are maintained by the Australian Radiation Protection and Nuclear Safety Agency (ARPANSA). Melbourne is a city of 5 million inhabitants located in the southeastern part of Australia on the shores of Port Phillip Bay. Like all Australian cities, Melbourne is sprawling and has a low population density by world standards. The climate is oceanic with hot summers and mild winters. The weather can change rapidly, especially during summers, due to the location of the city between hot inland and cold southern ocean. Heavy storms and rain associated to cold fronts are typical during summers, while winters are more stable but cloudy. Measurements in Melbourne are performed using a Bentham DTMc300 spectroradiometer. This instrument is calibrated for irradiance twice a year using a 1 kW QTH lamp whose output is traceable to NIST and the wavelength calibration is based on the UV spectral lines of a mercury lamp. Alice Springs was selected to represent inland Australian site. The site is located in the Northern Territory of Australia and it is surrounded by deserts. Summers are extremely hot and dry while winters are short and mild. The average temperatures during summer are over 30°C, and the minimum temperatures can drop below 10°C during winter. There are typically more than 200 cloud-free sunny days per year in Alice Springs. The UV Index is monitored using a radiometer manufactured by sglux GmbH (Berlin, Germany). The sensor is a hybrid SiC photodiode model UV-Cosine\_UVI or ERYCA. A logger records data every minute and the average over ten minutes is calculated during post-processing. The radiometer is exchanged every second year for an equivalent sensor that has been calibrated at the ARPANSA laboratory in Melbourne against the Bentham spectroradiometer. All data for Alice Springs reported in this paper was collected with a single UV sensor.

Marambio Base is located on the highest part of the Seymour/Marambio Island, surrounded by the Weddell sea on the north-east side of the Antarctic Peninsula. As a cooperation between the Argentinian National Meteorological Service and the FMI, GUV-radiometer, model GUV-2511, measurements started in 2017. Near real time data is shown in [http://fmiarc.fmi.fi/sub\\_sites/GUVant/](http://fmiarc.fmi.fi/sub_sites/GUVant/) for the last five days. The temperatures at the site are around 10°C in summer and can drop down to -30°C in winter. The soil is frozen and covered with snow most of the year and the Weddell Sea in the East is frozen during the winter, but the coast at Marambio is free from ice the whole year. In the summer heavy cloudiness and fog are common. The station is part of the Global Atmospheric Watch (GAW) program of the World Meteorological Organization (WMO). Two radiometers rotate so that one is measuring at the site while the other is calibrated by Biospherical Instruments, Inc, U.S. and



also participates in solar comparisons in Sodankylä, Finland. The expanded ( $k=2$ ) uncertainty of the GUV measurements in Marambio is 9% at SZAs smaller than  $80^\circ$ .

Palmer Station is located on Anvers Island on the west coast of the Antarctic continent. It is a research station of the United States, operated by the U.S. National Science Foundation. UV measurements are performed with a SUV-100 spectroradiometer and are part of the NOAA Antarctic UV monitoring network. The effective albedo at Palmer Station is about 0.8 in winter and 0.4 in summer (Bernhard et al., 2005). The sea adjacent to the station is frozen during winter and open during summer. Temperatures can fall below  $-20^\circ\text{C}$  in winter and can reach up to  $10^\circ\text{C}$  in summer. Heavy winds are frequent during winter time. The quality assurance of the spectroradiometer was described by Bernhard et al. (2005) and includes comparisons with results of radiative transfer models and measurements of a GUV-511 multi-channel filter radiometer that is deployed next to the SUV-100 instrument.



**Figure 1.** Location of ground-based reference sites. See Table 2 for explanation of site acronyms.



**Table 2.** Validation sites ordered according to latitude from North to South.

Site	site's acronym	Affiliation	Lat., °N	Long., °E	Elev., m	Characteristics
Ny-Ålesund	NYA	NILU/DSA	78.924	11.930	10	Arctic coast
Andøya	AND	NILU/DSA	69.279	16.009	380	Arctic coast
Sodankylä	SOD	FMI	67.367	26.630	179	rural, subarctic
Finse	FIN	DSA/NILU	60.593	7.524	1200	mountainous
Helsinki	HEL	FMI	60.203	24.961	43	urban, subarctic coast
Blindern	BLI	NILU/DSA	59.938	10.717	90	urban, subarctic coast
Uccle	UCC	RMIB	50.797	4.358	100	suburban
Villeneuve d'Ascq	VDA	Univ. Lille	50.611	3.140	70	suburban
Davos	DAV, DBR, DBB	PMOD-WRC	46.813	9.844	1610	mountainous
Aosta	AOS	ARPA	45.742	7.357	570	valley, mountainous
Haute-Provence	OHP	Univ. Lille	43.935	5.712	688	rural, mountainous
Rome	ROM	Univ. Rome	41.901	12.516	70	urban
Thessaloniki	THE	Aristotle Univ.	40.634	22.956	60	urban, Mediterranean coast
Bet Dagan	BET	IMS	32.008	34.815	31	rural, shrublands
Jerusalem	JER	IMS	31.770	35.197	770	urban
Eilat	EIL	IMS	29.553	34.952	22	urban
Izana	IZA	AEMET	28.308	-16.499	2372	Top of mountain
Mahé	MAH	Univ. Réunion	-4.679	55.531	15	coast
Antananarivo	ANT	Univ. Réunion	-18.916	47.565	1370	urban, medium mountain
Anse Quito	ROD	Univ. Réunion	-19.758	63.368	32	coast
Saint-Denis	OPA, STD	Univ. Lille/Univ. Réunion	-20.902	55.485	82	coast, mountainous
Alice Springs	ALI	ARPANSA	-23.796	133.889	550	desert
Melbourne	MEL	ARPANSA	-37.728	145.100	60	coast
Marambio	MAR	FMI/SMN	-64.241	-56.627	198	Antarctic coast
Palmer	PAL	NSF/NOAA	-64.774	-64.051	21	Antarctic coast



**Table 3.** Spectroradiometers used in the study. Eryth. act. denotes which erythemal action spectrum is used for retrieving erythemally weighted dose rates and daily doses. 1987 denotes the McKinlay and Diffey (1987) and 1998 the Commission Internationale de l’Eclairage (1998) action spectrum. See Table 2 for explanation of site acronyms.

Site	Instrument	data period	Eryth. act.	Traceability
SOD	Brewer Mk II #037	1.1.2018-31.8.2019	1987	MIKES-Aalto
HEL	Brewer MK III #107	1.1.2018-31.8.2019	1987	MIKES-Aalto
UCC	Brewer MKIII #178	1.1.2018-31.8.2019	1987	NIST via Kipp&Zonen
VDA	Bentham DTMc300	1.1.–15.5.2018, 14.9.2018–31.8.2019	1998	NIST
DAV	Bentham QASUMEI	15.6.-26.10.2018, 21.3-27.3 and 17.5-23.8.2019	1998	PTB
DBR	Brewer MK III #163	1.1.-26.7.2018, 29.9.2018 - 31.8.2019	1998	PTB
AOS	Bentham DTMc300	1.1.2018-31.8.2019	1998	PTB
OHP	Bentham DTMc300	1.1.2018-31.12.2018	1998	NIST
ROM	Brewer #067	5.7.2019-26.8.2019	—	NIST via IOS
THE	Brewer MKIII #086	1.1.2018-31.8.2019	1998	PTB
IZA	Brewer #183	1.1.2018-31.8.2019	1998	PTB via PMOD-WRC
OPA	Bentham DTMc300	1.1.2018-31.8.2019	1998	NIST
MEL	Bentham DTMc300	1.1.2018-31.8.2019	1998	NIST
PAL	SUV-100	1.1.2018-31.8.2019	1987	NIST

**Table 4.** Broadband and multiband radiometers used in the study. Data period is from 1 Jan 2018 to 31 Aug 2019. Eryth. act. denotes which erythemal action spectrum is used for retrieving erythemally weighted dose rates and daily doses. 1987 denotes the McKinlay and Diffey (1987) and 1998 the Commission Internationale de l’Eclairage (1998) action spectrum. See Table 2 for explanation of site acronyms.

Site	Instrument	data frequency	bandwidth	Eryth. act.	Traceability
NYA	GUV-541	1 min ave	5 channels, FWHM 10nm	1998	PTB via PMOD-WRC
AND	GUV-541	1 min ave	5 channels, FWHM 10nm	1998	PTB via PMOD-WRC
FIN	GUV-541	1 min ave	5 channels, FWHM 10nm	1998	PTB via PMOD-WRC
BLI	GUV-511	1 min ave	5 channels, FWHM 10nm	1998	PTB via PMOD-WRC
DBB	average KZ560, YES010938, SL501A	10 min ave	broadband	1998	PTB
BET	YES UVB-1	10 min ave	broadband 280-320 nm	1987	Kipp&Zonen
JER	YES UVB-1	10 min ave	broadband 280-320 nm	1987	Kipp&Zonen
EIL	YES UVB-1	10 min ave	broadband 280-320 nm	1987	Kipp&Zonen
MAH	Kipp&Zonen UVS-E-T	5 min ave	ISO 17166/CIE S007/E-1999	1998	Kipp&Zonen
ANT	Kipp&Zonen UVS-E-T	5 min ave	ISO 17166/CIE S007/E-1999	1998	Kipp&Zonen
ROD	Kipp&Zonen UVS-E-T	2 min ave	ISO 17166/CIE S007/E-1999	1998	Kipp&Zonen
STD	Kipp&Zonen UVS-E-T	5 min ave	ISO 17166/CIE S007/E-1999	1998	PTB via PMOD-WRC
ALI	sglux ERYCA	10 min ave	ISO 17166	1998	NIST via ARPANSA
MAR	GUV-2511	1 min ave	5 channels, FWHM 10nm	1987	NIST via BSI



### 3 Validation methods

TROPOMI overpass irradiance at 305, 310, 324 and 380 nm, overpass erythemally weighted dose rate, overpass UV index and erythemally weighted daily dose were compared to ground-based measurements. The ground-based data were used as such, as provided by operators, and no conversion between UV index and dose rate was done. The TROPOMI UV parameters are calculated using the erythemal action spectrum from Commission Internationale de l'Eclairage (1998). Most of ground-based measurements used the same action spectrum, while a couple of sites had still in use the action spectrum from McKinlay and Diffey (1987). The effect of using different action spectrum was modeled (results not shown), and they were in agreement with those of Webb et al. (2011). The uncertainty related to the choice of action spectrum was assumed to be less than 1% for low and middle latitudes sites and less than 2% for high latitude sites.

Spectroradiometers measure complete UV spectra and their data were used for the validation of irradiances. Each spectrum was first deconvoluted and then convoluted using a triangular slit of 1 nm at full-width-at-half-maximum using the Shicrivm package (Slaper et al., 1995) freely available at <https://www.rivm.nl/en/uv-ozone-layer-and-climate/shicrivm> (last visited 19 Mar 2020) as the TROPOMI irradiance is calculated using that standard slit. Data from Palmer were processed with the algorithm described by (Bernhard et al., 2004).

The validation of overpass erythemally weighted dose rate was performed against both spectroradiometer and broadband radiometer measurements. The measurement frequency of these instruments is different: a spectroradiometer may take from 3 to 6 minutes to scan the whole UV range, while a broadband radiometer can measure every second, even if the final product is saved as 1, 2, 5 or 10 minute average. This means that e.g. during changing cloudiness, the cloudiness conditions in the beginning of the spectrum (at short wavelengths) measured by a spectroradiometer may differ from those at the end of the spectrum (at longer wavelengths). The time stamp of spectroradiometer dose rate can differ between sites. For some sites the time stamp is set at the beginning of the spectrum and for some it is set at the most effective wavelength regarding erythemally weighted UV irradiance, at around 308–311 nm. Most of the spectroradiometers measure only 2–4 scans per hour. In order to get enough points between satellite overpasses and ground-based measurements, the allowed time difference between the satellite overpass and the spectroradiometer measurement was set to less than 5 minutes.

Recording frequencies of broadband and multichannel radiometers are listed in Table 4: averages were made over 1, 2, 5 or 10 minutes. The allowed time difference between the satellite overpass and the ground-based measurement was set to be less than half of the recording frequency. Eg. if ground-based data were recorded every minute, then the allowed time difference was set to less than 30 seconds. If the ground-based data was ten minute average, then the maximum time difference was set to be less than 5 minutes.

For the validation of overpass UV index both broadband and multichannel radiometers were used and the time difference between satellite overpass and ground-based data was limited to half of the recording frequency, as for the dose rate validation. All type of instruments (spectroradiometer, broadband and multichannel radiometer) were used for the validation of the erythemally weighted daily dose.



310 For all TROPOMI overpass pixels a ground-based measurement was chosen if found within the allowed time difference. No  
quality filtering was performed for the TROPOMI data. The distance between the TROPOMI pixel and the ground station was  
filtered to be less than 5 km, the SZA less than  $80^\circ$ , and following Tanskanen et al. (2007) the altitude difference between the  
altitude of the site and the TROPOMI pixel less than 500 m. For mountainous sites of Davos and Aosta the maximum distance  
between the TROPOMI pixel and the site was limited to be less than 3 km. The SZA of  $80^\circ$  was chosen to avoid very low UV  
315 irradiances, as for very low irradiances the ratio between satellite and ground-based data becomes unstable (Tanskanen et al.,  
2007). Also at SZAs smaller than  $80^\circ$  the effect of stray light in single monochromator Brewsters (Bais et al., 1996) is avoided.

The relative difference  $\rho$  between satellite data and ground-based data was calculated for each pair of satellite data (sat) and  
ground-based data (gr) using the following equation:

$$\rho = 100\% * [(sat - gr)/gr] \quad (1)$$

320 The median and 25th and 75th percentiles of the  $\rho$  values were calculated for each site. The  $W_{10}$  and  $W_{20}$  from Tanskanen  
et al. (2007) were calculated also in this study. The  $W_X$  is determined as percentage of satellite data which is within  $X\%$  from  
ground-based data:

$$W_X = P(-X < \rho < X) \quad (2)$$

Similarly to Tanskanen et al. (2007) data sets were divided into subsets according to albedo. Snow cover was considered,  
325 when the albedo used by the TROPOMI UV processor (Kujanpää et al., 2020) was higher than 0.1, and the data set was divided  
into snow cover (SC) and snow free (SF) ground conditions. The albedo used by the TROPOMI UV processor is derived from  
albedo climatology (Kujanpää and Kalakoski, 2015), thus it may differ from the true albedo conditions of the site (Tanskanen  
et al., 2007; Bernhard et al., 2015). In addition, a subset of clear sky data was selected. Clear skies were assumed when the  
cloud optical depth retrieved by the TROPOMI UV processor was lower than 0.5. Clear sky data sets included both snow cover  
330 and snow free conditions. Here again, one needs to keep in mind, that it is the cloud optical depth as derived from the LUT of  
the TROPOMI UV processor, not the cloudiness observation from the site.

The spatial resolution of TROPOMI data is very high compared to older generation satellite instruments. This leads to huge  
amount of data and at most sites several satellite pixels fulfilling the selection criteria were colocated with the same ground-  
based measurement. Thus, the sensitivity of the results was studied by comparing three different data selection methods for  
335 Villeneuve d'Ascq measurements: 1) Each TROPOMI pixel was treated as individual measurement, 2) the pixel nearest of  
the site was chosen, 3) the average of the TROPOMI pixels fulfilling the chosen limitations (time difference, SZA, altitude,  
distance) was used. Results are shown in the Fig. S13 and Table S6 of the Supplement material.

## 4 Results

Results for validation of overpass spectral irradiances, dose rates, UV index and daily doses are discussed separately in the  
340 following sections. Scatter plots, histograms and tables including the statistics were prepared for all studied UV parameters,  
and they are shown in the Supplement material of this paper. Here they are shown only for dose rate / UV index.

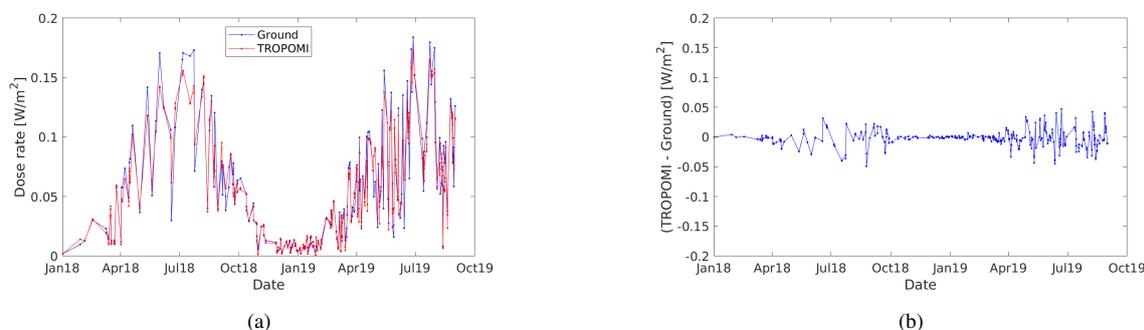


#### 4.1 Spectral irradiances

TROPOMI overpass irradiances were compared with the following spectroradiometers listed in Table 3: SOD, HEL, VDA, DAV, DBR, AOS, OHP, ROM, THE, IZA, OPA, MEL and PAL. The statistics are shown in the Tables S1–S4 of the Supplement material. Scatter plots and histograms are showed in Figs. S1–S8. For irradiances at 305, 310, 324 and 380 nm the median  $\rho$  was within  $\pm 10\%$  at 11, 9, 10 and 6 sites from the 13 sites (7 sites for 380 nm), respectively, for snow free ground conditions. For the four wavelengths, at all sites except one, more than 50% of satellite data were within  $\pm 20\%$  from ground measurements. During snow conditions the percentage of satellite data being within  $\pm 20\%$  from ground measurements decreased in Davos from more than 60% to around 20%. This is seen as a shift of  $\rho$  towards negative values when comparing snow cover data set to snow free data set. For the other four sites which had data sets during snow cover, Sodankylä, Helsinki, Aosta and Palmer, no significant difference was observed between snow cover and snow free surfaces. At Palmer, a systematic underestimation of irradiances occurred at all wavelengths. The median  $\rho$  at 305 nm was -46% and -56% for snow free and snow covered surface, respectively. Satellite data had a positive bias at Davos, Aosta and Izaña, while at other sites the bias was more randomly distributed. The spread of the scatter plot was larger at 380 nm than at 305 nm, which is related to the influence of clouds: radiation distribution at short UV wavelengths is less affected by clouds than at longer wavelengths. For all stations, over 50% of clear sky satellite data were within  $\pm 20\%$  from ground measurements.

#### 4.2 Erythemally weighted dose rate and UV index

An example of TROPOMI overpass and ground-based erythemally weighted dose rate time series is shown in Fig. 2 together with the absolute difference for Uccle. TROPOMI data follow well daily variations in UV dose rates, and the absolute difference was less than  $0.05 \text{ W/m}^2$ . The validation results are shown by comparisons against spectroradiometers, broadband and multichannel radiometers in the following subsections.



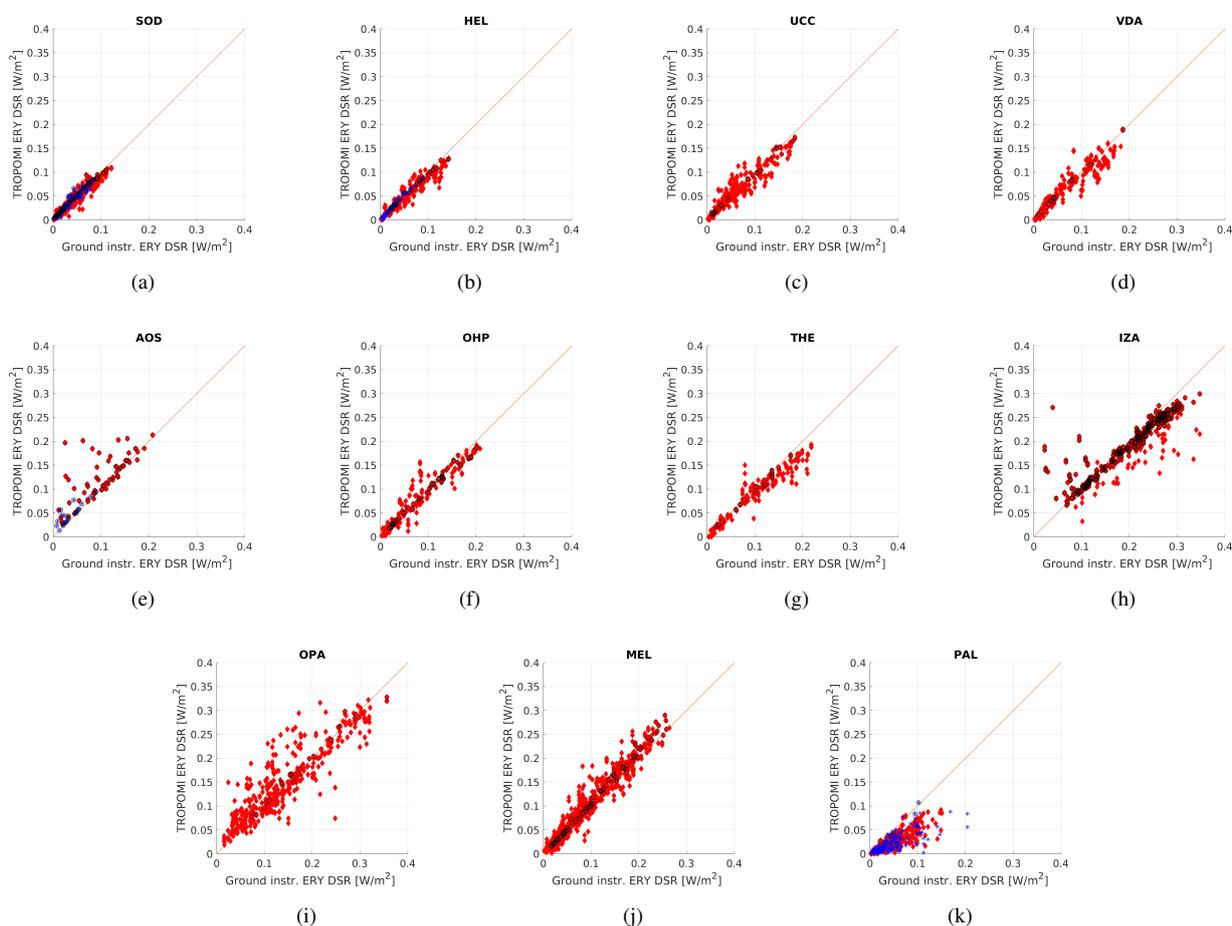
**Figure 2.** TROPOMI overpass and Brewer spectroradiometer a) erythemal dose rates, b) absolute difference of erythemal dose rates at Uccle, Belgium, during Jan 2018 – Aug 2019.



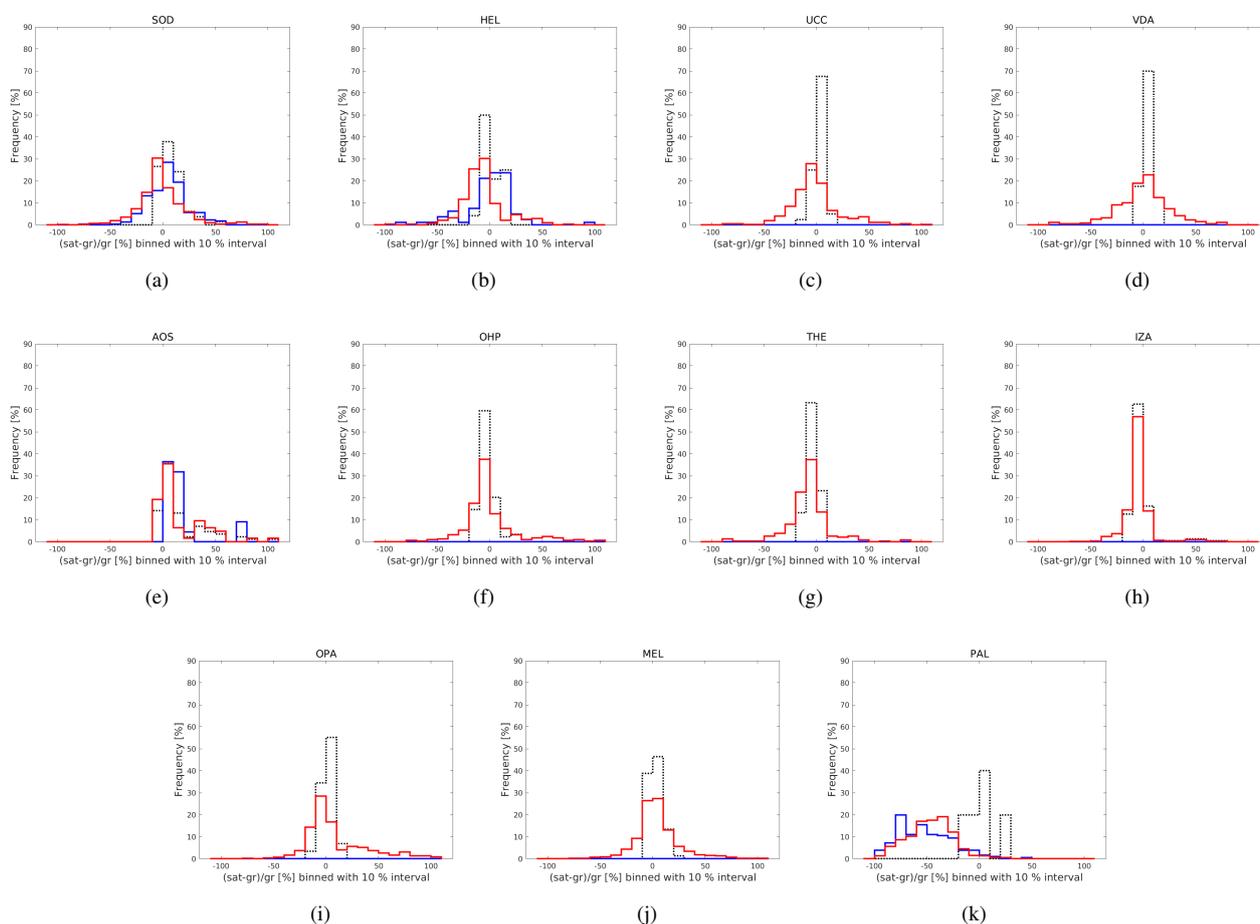
#### 4.2.1 Validation against spectroradiometers

The comparison of TROPOMI overpass erythemally weighted dose rates against spectroradiometer measurements showed similar patterns as the comparison of single irradiances at 305 and 310 nm. The scatter plot and histograms are showed in  
365 Figs 3 and 4, respectively, and the statistics in Table 5. At ten and seven sites the median  $\rho$  was within  $\pm 8\%$  and  $\pm 5\%$ , respectively, for snow free conditions. As for irradiances, TROPOMI UV dose rates show a systematic negative bias at Palmer, with median  $\rho$  of  $-45\%$  during snow cover conditions. The histograms of  $\rho$  are similar for snow cover and snow free conditions. Also at the other sites which have data sets for snow cover conditions, there are no noticeable differences between snow cover and snow free conditions.

370 At all sites, except Palmer, over 60% of TROPOMI data are within  $\pm 20\%$  from ground-based measurements. In Aosta and Izaña there is a large positive bias for some pixels. The feature is more pronounced when small dose rates are measured, which suggests the reason to be how TROPOMI interpret cloudiness: underestimation of cloudiness leads to positive bias in TROPOMI data.



**Figure 3.** Erythemally weighted dose rates from spectroradiometer measurements and retrieved from satellite overpass at a) Sodankylä, b) Helsinki, c) Uccle, d) Villeneuve d' Ascq , e) Aosta, f) Haute-Provence, g) Thessaloniki, h) Izaña, i) Saint-Denis, j) Melbourne and k) Palmer. Red diamond denotes snow free surface, blue star snow cover and black circle clear sky.



**Figure 4.** Histograms of relative difference between spectroradiometer measurements and satellite overpass erythemally weighted dose rates at a) Sodankylä, b) Helsinki, c) Uccle, d) Villeneuve d'Ascq, e) Aosta, f) Haute-Provence, g) Thessaloniki, h) Izaña, i) Saint-Denis, j) Melbourne and k) Palmer. Black dotted line denotes clear sky, red snow free surface and blue snow cover on the ground.



**Table 5.** TROPOMI overpass erythemally weighted dose rates compared to spectroradiometer measurements. The percentage relative differences  $100\% \cdot (\text{sat-gr})/\text{gr}$ , their medians (median) and 25th (p25) and 75th (p75) percentiles were calculated. N is the number of measurement days included in the study.  $W_{10}$  and  $W_{20}$  are percentage of satellite data which are within 10% and 20% from ground measurements. CS=Clear sky, SCAS=Snow cover at ground, all sky, SFAS=Snow free ground, all sky.

Station	Conditions	N	median [%]	p25 [%]	p75 [%]	$W_{10}$	$W_{20}$
SOD	CS	132	5.92	-0.45	12.48	64.39	88.64
SOD	SCAS	211	4.09	-6.49	14.50	44.08	76.78
SOD	SFAS	451	-3.19	-12.22	8.82	47.23	71.62
HEL	CS	48	-0.86	-4.24	8.80	70.83	100.00
HEL	SCAS	80	2.43	-6.48	11.99	45.00	76.25
HEL	SFAS	275	-8.11	-15.22	0.29	40.00	67.64
UCC	CS	40	3.01	-1.09	4.84	92.50	100.00
UCC	SFAS	399	-4.42	-13.89	6.64	46.87	69.42
VDA	CS	40	6.65	3.60	7.83	87.50	100.00
VDA	SFAS	337	1.63	-13.17	14.14	41.84	64.39
AOS	CS	84	10.19	1.96	43.24	50.00	63.10
AOS	SCAS	22	15.00	6.68	73.36	36.36	68.18
AOS	SFAS	62	7.22	1.43	41.87	54.84	61.29
OHP	CS	89	-2.44	-7.19	0.25	79.78	96.63
OHP	SFAS	280	-4.11	-11.07	3.40	50.36	73.93
THE	CS	30	-2.57	-8.61	-0.50	86.67	100.00
THE	SFAS	235	-7.68	-14.10	-0.62	51.06	76.17
IZA	CS	386	-4.50	-7.34	-0.11	79.02	91.97
IZA	SFAS	454	-5.38	-9.02	-1.15	70.93	86.12
OPA	CS	29	1.29	-1.18	6.32	89.66	100.00
OPA	SFAS	424	0.05	-8.14	28.19	45.28	63.68
MEL	CS	142	2.34	-2.02	7.56	85.21	98.59
MEL	SFAS	1074	2.24	-6.36	12.34	53.91	76.16
PAL	CS	5	3.13	-7.07	9.87	60.00	80.00
PAL	SCAS	180	-56.64	-72.29	-34.24	5.56	10.56
PAL	SFAS	393	-44.88	-60.74	-31.96	2.80	7.89

#### 4.2.2 Validation against broadband and multiband radiometers

375 The scatter plots and histograms of TROPOMI overpass UV index and dose rates comparison against broadband and multi-channel radiometers are shown in Figs 5 and 6, respectively, and the statistics in Table 6. The number of collocated pixels is

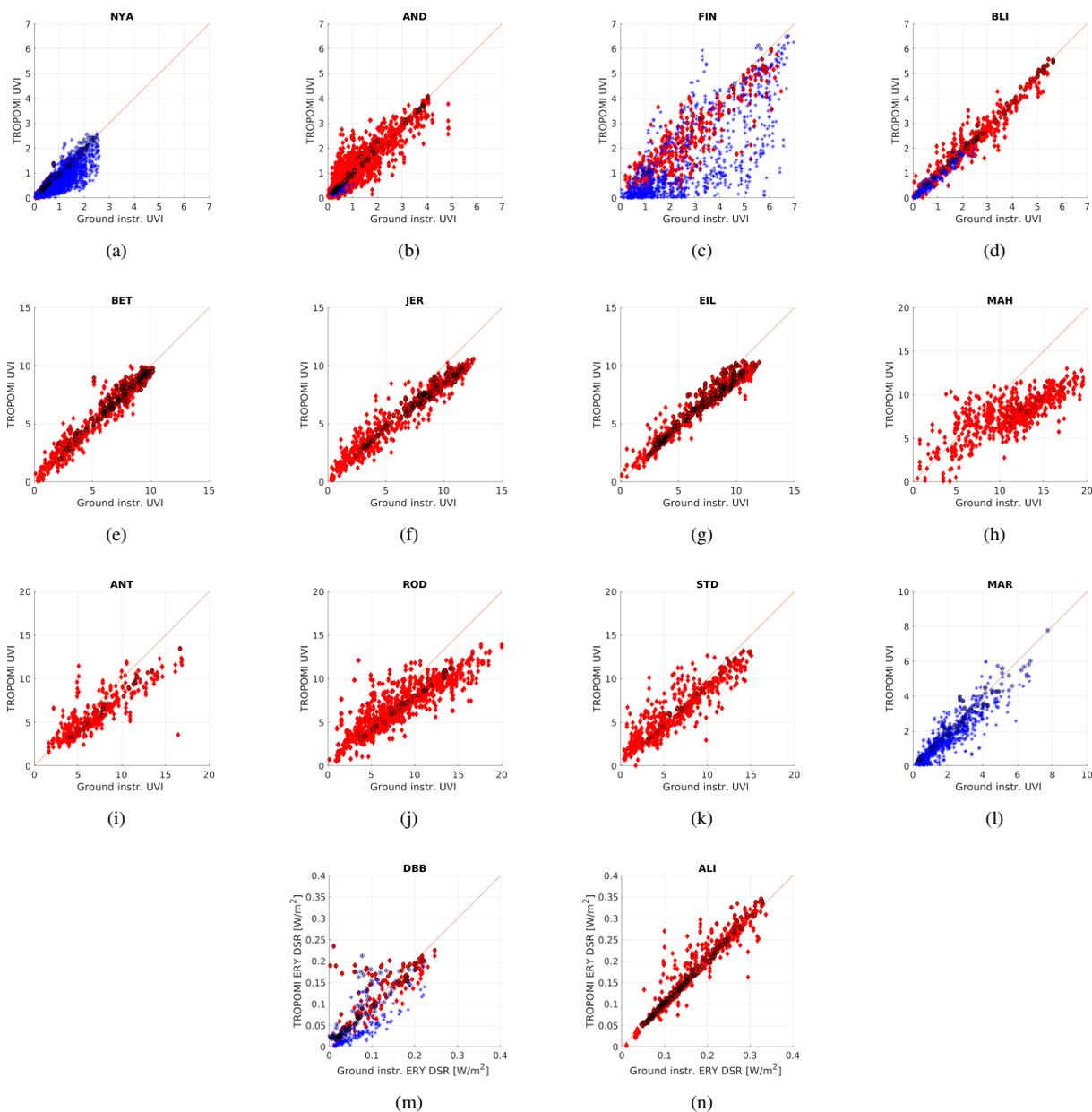


much higher for broadband instruments than for spectroradiometers, as they measure continuously. At several sites (Jerusalem, Mahé, Antananarivo, Anse Quito, Saint-Denis and Alice Springs), the UV index can be higher than 11, categorized as "extreme" UV (WMO, 1997). These extreme values are underestimated by TROPOMI, except in Alice Springs. The feature is pronounced in the Indian Ocean sites Mahé, Antananarivo and Anse Quito. The strongest underestimation is seen in Mahé, where the median  $\rho$  was -34% with the 25th and 75th percentiles of -40% and -20%, respectively.

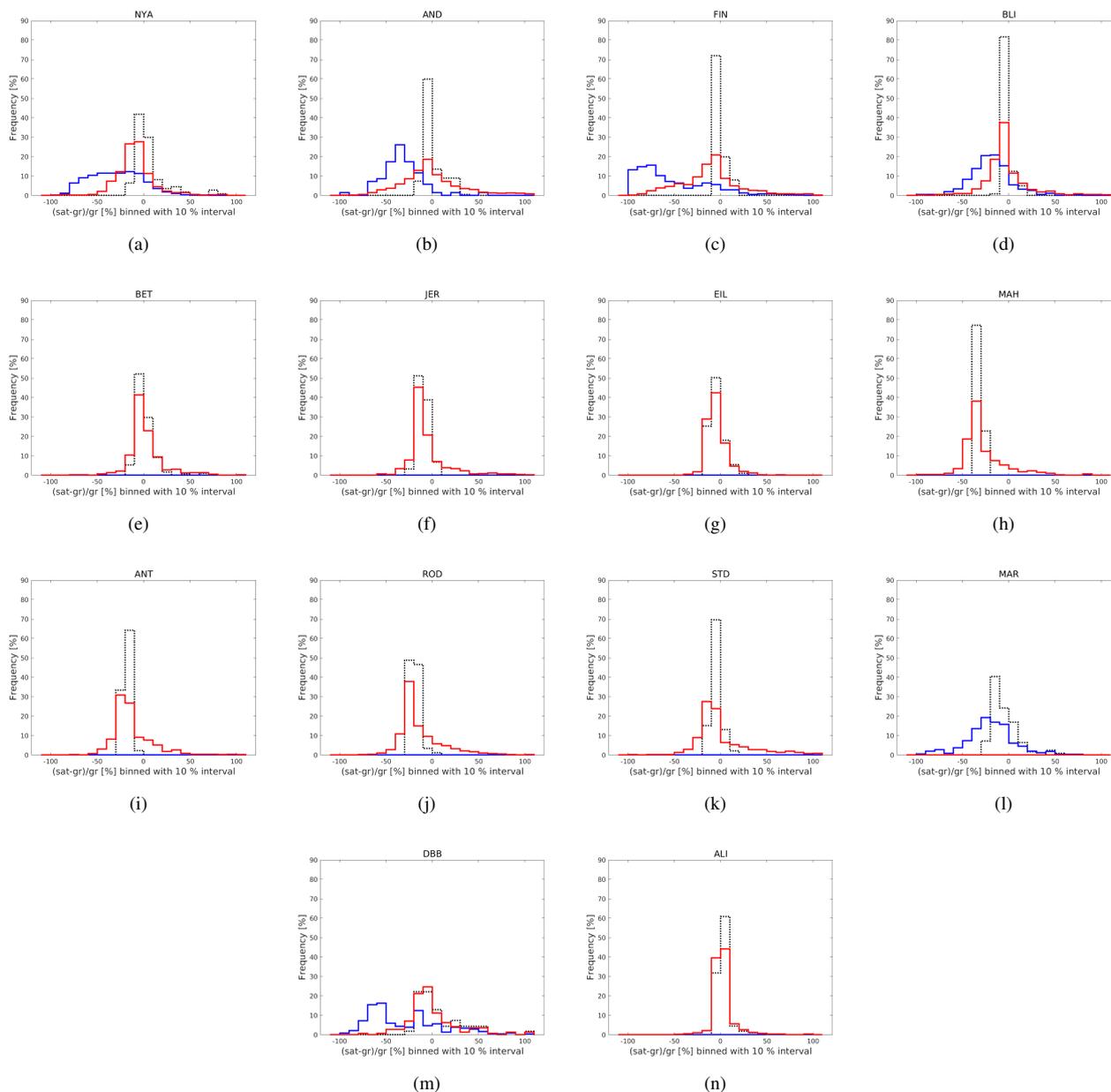
For the other sites, the median  $\rho$  for snow free conditions was between -1 and -10%. At the high latitude site of Ny-Ålesund where snow covers the surface almost half a year, and at the mountainous site of Finse, similar underestimation to Palmer is seen. However, at Ny-Ålesund and Finse, differences occur between snow cover and snow free data sets. The medians  $\rho$  for snow free conditions are -10% at both sites, and for snow cover conditions -30% and -65% at Ny-Ålesund and Finse, respectively. The difference of snow cover and snow free conditions is distinctly seen in the histogram (Fig. 6a and 6c). Also at Davos and Andøya, underestimation occurred during snow cover, and  $\rho$  differed between snow cover and snow free data sets. The median  $\rho$  was approximately -5% and -35% for snow free and snow covered conditions, respectively, at both sites. At Blindern, the same feature was seen, but with a smaller difference between the two conditions: -5% and -20% for snow free and snow cover conditions, respectively. At Marambio, the Antarctic station which has snow cover all year round, the underestimation was similar to Blindern with median  $\rho$  of -20%.

The median  $\rho$  for clear sky conditions was within  $\pm 10\%$  for all sites except the Indian Ocean sites, Mahé, Antananarivo and Anse Quito, and the Israeli site of Jerusalem. At 8 sites the median was within  $\pm 5\%$  for clear skies.

A study on the effect of taking into account quality flags was done for the site of Davos. Data for which the quality value number UVQAV was less than 0.5 were excluded (see Section 2.1 for explanation of UVQAV). This removed e.g., unreliable values when the cloud optical depth was 0. Indeed, as mentioned in Section 2.2, Davos is a mountainous site with heterogeneous albedo during the winter. As shown in Supplement material Figs. S14, setting a limit of 0.5 for the UVQAV, results in removing satellite observations with at least two of the following warnings: "rough\_terrain", "alb\_hetero" or "clearsky\_assumed".



**Figure 5.** UV index from radiometer measurements and retrieved from satellite overpass at a) Ny-Ålesund, b) Andøya, c) Finse, d) Blindern, e) Bet Dagan, f) Jerusalem, g) Eilat, h) Mahé, i) Antananarivo, j) Anse Quito, k) Saint-Denis, l) Marambio and erythemally weighted dose rates at m) Davos and n) Alice Springs. Red diamond denotes snow free surface, blue star snow cover and black circle clear sky.



**Figure 6.** Histograms of relative difference between radiometer measurements and satellite overpass UV index at a) Ny-Ålesund, b) Andøya, c) Finse, d) Blindern, e) Bet Dagan, f) Jerusalem, g) Eilat, h) Mahé, i) Antananarivo, j) Anse Quitor, k) Saint-Denis, l) Marambio and overpass erythemally weighted dose rates at m) Davos and n) Alice Springs. Black dotted line denotes clear sky, red snow free surface and blue snow cover on the ground.



**Table 6.** TROPOMI overpass UV index/erythemally weighted dose rates compared to broadband and multichannel radiometer measurements. The percentage relative differences  $100\% \cdot (\text{sat-gr})/\text{gr}$ , their medians (median) and 25th (p25) and 75th (p75) percentiles were calculated. N is the number of measurement days included in the study.  $W_{10}$  and  $W_{20}$  are percentage of satellite data which are within 10% and 20% from ground measurements. CS=Clear sky, SCAS=Snow cover at ground, all sky, SFAS=Snow free ground, all sky.

Station	Conditions	N	median [%]	p25 [%]	p75 [%]	$W_{10}$	$W_{20}$
NYA	CS	110	0.64	-4.18	8.55	71.82	86.36
NYA	SCAS	3333	-30.28	-52.24	-9.25	18.33	34.26
NYA	SFAS	832	-9.96	-18.93	-1.49	38.94	70.07
AND	CS	177	-2.71	-5.53	1.48	73.45	89.83
AND	SCAS	69	-33.66	-43.95	-22.00	7.25	18.84
AND	SFAS	2378	-3.23	-18.94	22.13	29.48	49.54
FIN	CS	25	-2.17	-4.52	0.19	92.00	100.00
FIN	SCAS	1099	-64.96	-82.27	-21.73	8.55	17.93
FIN	SFAS	446	-9.86	-28.60	3.79	29.60	50.67
BLI	CS	120	-3.76	-5.84	-1.11	94.17	100.00
BLI	SCAS	345	-20.00	-32.41	-8.06	20.87	44.93
BLI	SFAS	620	-5.26	-14.33	2.77	49.03	74.03
DBB	CS	163	1.90	-9.26	39.66	34.97	61.35
DBB	SCAS	233	-37.33	-60.86	2.90	10.30	24.03
DBB	SFAS	142	-4.88	-12.92	16.17	35.92	63.38
BET	CS	305	-0.85	-4.06	4.13	81.97	96.39
BET	SFAS	886	-1.21	-6.28	6.76	64.33	84.09
JER	CS	250	-10.82	-14.94	-6.86	45.60	96.80
JER	SFAS	899	-12.10	-16.83	-2.98	27.81	77.09
EIL	CS	406	-5.31	-10.10	-0.25	68.23	99.01
EIL	SFAS	872	-5.99	-11.48	0.63	59.06	92.66
MAH	CS	22	-32.84	-36.91	-31.24	0.00	0.00
MAH	SFAS	719	-33.81	-39.83	-19.79	8.48	17.80
ANT	CS	42	-18.16	-21.70	-16.43	2.38	66.67
ANT	SFAS	431	-17.81	-25.34	-5.12	16.71	48.49
ROD	CS	88	-19.69	-21.81	-17.78	4.55	51.14
ROD	SFAS	1527	-20.82	-26.54	-3.69	15.46	35.30
STD	CS	46	-6.69	-9.01	-4.31	82.61	100.00
STD	SFAS	610	-8.11	-14.96	14.35	30.33	63.11
ALI	CS	286	1.79	-0.72	5.59	92.66	97.20
ALI	SFAS	898	0.75	-2.62	5.33	83.74	91.54
MAR	CS	124	-9.68	-14.84	1.85	41.13	87.90
MAR	SCAS	1169	-20.00	-33.53	-4.76	21.90	43.37



### 4.3 Erythemally weighted daily dose

400 TROPOMI erythemally weighted daily doses were compared against daily doses derived from spectroradiometer measurements (SOD, HEL, UCC, VDA, AOS, OHP, IZA, OPA, MEL and PAL), multichannel radiometers (NYA, AND, FIN, BLI and MAR) and a broadband radiometer (ALI). The scatter plots, histograms and statistics are showed in the Figs. S9-S10, S11-S12 and Table S5 of the Supplement material, respectively.

As the satellite daily dose is calculated using the assumption that the cloudiness retrieved by the satellite during the overpass would be the same during the whole day, which is an oversimplified assumption for many sites, larger deviation in the results were expected than for overpass dose rates. At all sites except Palmer, Aosta, Finse and Ny-Ålesund, the increased deviation was seen in both positive and negative biases. However, the median  $\rho$  was within  $\pm 10\%$  and  $\pm 5\%$  at 11 and 8 sites, respectively. The total number of sites providing daily doses for this study was 16. At all sites, except Palmer, over 50% of satellite data were within  $\pm 20\%$  from ground-based measurements during snow free surface conditions. Marambio was always snow-covered.

410 At Palmer, the pattern was similar to results of the dose rate validation indicating large underestimation with median  $\rho$  of -49% and -62% for snow free and snow covered conditions. Also at Ny-Ålesund there was significant underestimation, but daily doses differed from dose rates by having also several overestimation cases. At Ny-Ålesund the median  $\rho$  was -12% and -33% for snow free and snow cover conditions, while during clear skies there was an overestimation with a median  $\rho$  of 6%. This could occur at situations, when the sky was clear during the overpass, but later changed towards more cloudy conditions, or would have been more cloudy before the overpass. Also at Andøya, Blindern and Finse large differences between the median of snow free ( $\rho = -2\%$ ,  $\rho = -5\%$  and  $\rho = -12\%$ , respectively) and snow cover ( $\rho = -36\%$ ,  $\rho = -22\%$  and  $\rho = -66\%$ , respectively) condition occurred. At Sodankylä and Helsinki the difference between the median of snow free and snow cover conditions was less than 10% and 4%, respectively.

420 At Aosta and Izaña there were large clear sky overestimations. At Aosta, the reason is the non-homogeneous topography around this mountainous site. At Aosta almost all satellite data is flagged with the "rough\_terrain" and "clearsky\_assumed" flags, and the data agree with measurements only during real clear sky conditions. At Izaña there were similar overestimation cases but also several underestimations. The underestimations can be due to situations, when the station is above the clouds, but the satellite interprets surrounding clouds, which in reality are below the station, as cloudiness of the site.

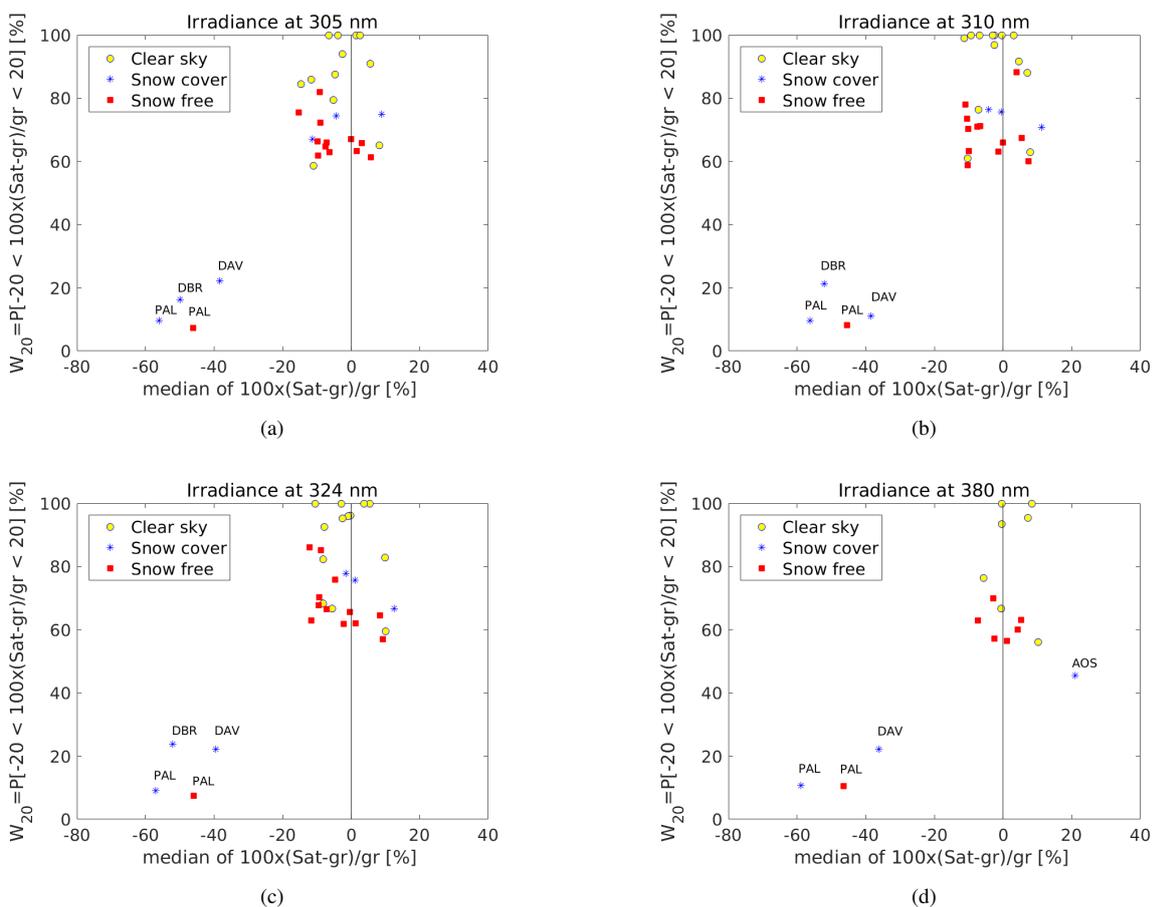
## 5 Discussion

425 Tanskanen et al. (2007) summarized the key validation statistics of OMI daily dose by plotting  $W_{20}$  in function of median  $\rho$  (Fig. 6 in Tanskanen et al. (2007)). The same was done for this validation. Figure 7 shows results for overpass irradiance validation, and Fig. 8 for overpass dose rates and daily doses. One notable difference between the OMI results from Tanskanen et al. (2007) and the results of the TROPOMI validation study is that positive bias due to tropospheric extinction is missing from TROPOMI results. That is due to the correction for absorbing aerosols which was not implemented in the OMI data in the study of Tanskanen et al. (2007). The current OMI UV algorithm is updated with the absorbing aerosol correction method described in Arola et al. (2009) and the same method is used in the TROPOMI UV algorithm. Thessaloniki is a site for which aerosols are

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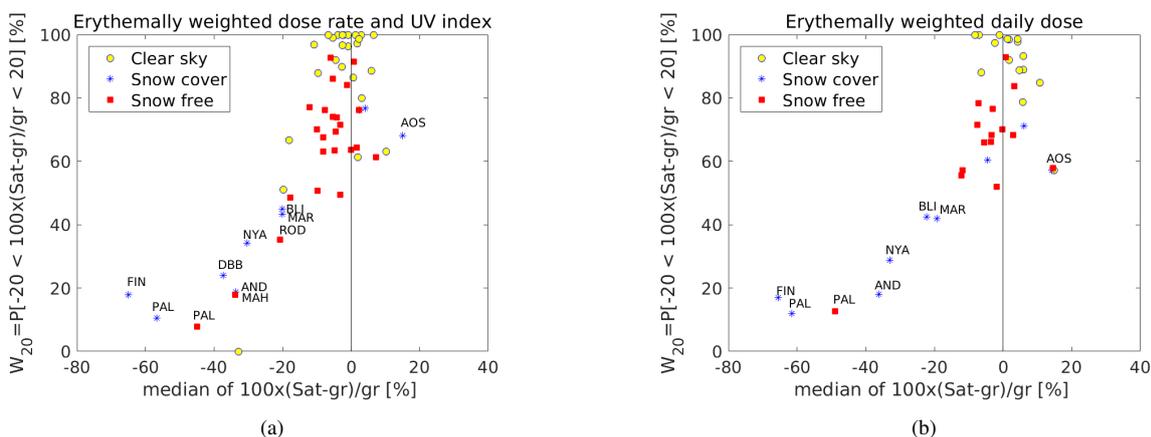


an important factor affecting UV radiation (e.g., Fountoulakis et al., 2016). Tanskanen et al. (2007) found a median difference of 10% (OMI/ground -1) between OMI and ground based erythemally weighted daily dose, while the TROPOMI validation showed an underestimation of -8% with 15% more satellite retrievals within  $\pm 20\%$  from ground-based measurements.



**Figure 7.**  $W_{20}$  in function of  $\rho$  for validation sites measuring overpass irradiance at a) 305, b) 310, c) 324 and d) 380 nm.

435 Kalliskota et al. (2000) found an underestimation of TOMS UV daily dose at Ushuaia, Argentina, and Palmer, and an overestimation in San Diego, U.S. Also TROPOMI underestimated the daily dose at Palmer (median  $\rho$  of -49%) for snow free surface, while agreed quite well (median  $\rho$  of -8%) during clear sky conditions. The monthly average underestimation of erythemally weighted daily dose was -35% for TOMS and the median underestimation (OMI/ground-1) was -49% for OMI. The results of OMI refer to those calculated using the cloud correction method based on the Lambertian equivalent reflectivity  
 440 (Tanskanen et al., 2007). OMI overestimated erythemally weighted daily doses by 9% (median of (OMI/ground -1)) for snow covered conditions at Sodankylä. For TROPOMI, the corresponding overestimation was 6%. For snow free conditions at



**Figure 8.**  $W_{20}$  in function of  $\rho$  for a) overpass erythemally weighted dose rate and UV index b) erythemally weighted daily dose.

Sodankylä, the median differences between satellite retrievals and ground-based measurements were 2% and -3% for OMI and TROPOMI, respectively.

Bernhard et al. (2015) studied in detail comparison of the OMI UV data against ground-based measurements at high latitudes and focused on the albedo effect. The sites of Ny-Ålesund, Sodankylä, Finse and Blindern were also included in their study. For Sodankylä, the results agreed with those of Tanskanen et al. (2007) (median  $\rho = 11\%$  for snow free and 6% for snow covered conditions). At Ny-Ålesund, Finse and Blindern the median differences of erythemally weighted daily dose (OMI/ground - 1) were -42%, -53% and -6%, respectively, for snow covered conditions, and 6%, 1% and 7%, respectively, for snow free conditions. The corresponding values from the TROPOMI validation were -33%, -66% and -22%, respectively, for snow covered conditions, and -12%, -12% and -6% for snow free conditions. The results from Bernhard et al. (2015) were calculated from two months of data (one month in spring when the ground was snow covered and one month in summer when it was snow free), while the TROPOMI validation results were calculated from the data set covering the whole year.

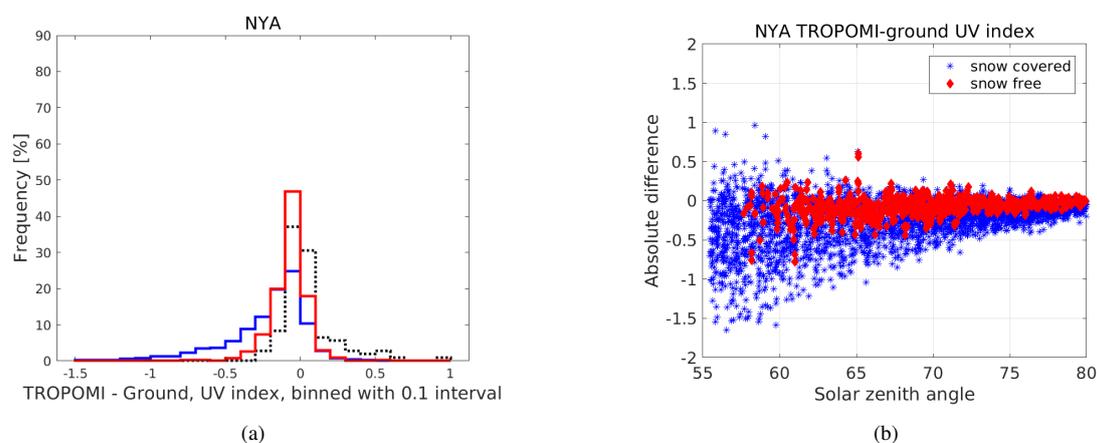
As discussed in Tanskanen et al. (2007), the OMI UV underestimation in Palmer is mostly due to unreasonably small values of surface albedo used by the OMI UV processor which results in an overestimate of the cloud optical thickness by misinterpreting reflections from the surface. This is also the dominant reason why TROPOMI underestimates UV radiation at Palmer and other sites with challenging snow albedo conditions (e.g., Ny-Ålesund, Finse, Davos). The albedo used by TROPOMI was compared with those calculated from SUV-100 spectra at Palmer. The SUV-100 albedo ranged from 0.2 to 0.7 while the TROPOMI albedo ranged from 0.05 to 0.4. The largest differences between the two albedo data sets was around 0.3. Surface albedo in Antarctica has a strong impact on UV radiation by increasing reflections from the surface and by multiple scattering between surface, clouds and atmosphere, which reduce the attenuation by clouds. According to Nichol et al. (2003) "a cloud with an optical depth of 10 reduces the UV irradiance, relative to clear-sky conditions, by 40% when the surface albedo is 0.05 as compared with reductions of 20% and 10% for surface albedos of 0.80 and 0.96, respectively."



At sites with homogeneous albedo conditions, even at locations with high surface albedo like Sodankylä, the differences between satellite retrievals and ground-based data are much smaller than for non-homogeneous conditions. In this study even  
465 the two Antarctic sites, Palmer and Marambio, differed from each other regarding the underestimation during snow cover conditions. At Marambio, the median  $\rho$  of erythemally weighted dose rate and daily doses was -20%, while at Palmer it was -62% for snow covered conditions. The smaller bias at Marambio can be attributed to the fact that the albedo values used in the TROPOMI UV algorithm (ranging from 0.2 to 0.7) are more realistic than at Palmer.

At high latitudes also long-term changes in the effective albedo resulting from climate change have to be considered in  
470 the future. Already now, at some Arctic sites the length of snow cover period has shortened by several weeks compared to a couple of decades ago (Bernhard, 2011; Luomaranta et al., 2019; Takala et al., 2011). That might be one reason for the positive median  $\rho$  of Sodankylä and Helsinki dose rates compared to snow free negative median  $\rho$ : If the climatological albedo used by TROPOMI is too high it can lead to underestimation of cloudiness. The results of Bernhard et al. (2015) showed that for the OMI instrument, overestimations and underestimations of up to 55% and -59%, respectively, were due to errors in the albedo  
475 climatology used in the OMI UV algorithm.

As discussed also by Tanskanen et al. (2007) validation results become unstable when UV irradiance is low, which is the case when SZAs or cloud optical depths are very large. Then, even small absolute differences are seen as large relative differences. This is frequent at high latitudes where large SZAs are present even at noon (overpass time) in winters. An example is shown in Fig. 9 for Ny-Ålesund, a site where UV index underestimations of more than -50% were found. Most of differences in the  
480 UV index are less than 0.5 when the whole data set is studied (including snow covered and snow free conditions) (Fig. 9a). The maximum differences are 1.5 for the lowest SZAs during snow cover conditions (9b).



**Figure 9.** The absolute difference between the TROPOMI overpass UV index and ground-based measurement at Ny-Ålesund as a) histogram binned with 0.1 interval and b) a function of SZA.

For many sites with uniform topography, stable albedo conditions, or predictable changes in albedo over the course of a year, TROPOMI UV data products agreed with ground-based measurements to within  $\pm 5\%$ . Aosta is a good example of



485 problems faced when retrieving satellite UV for mountainous sites. The non-homogeneous topography leads to uncertainties in  
the retrieval of cloud optical depth, which often result in a cloud optical depth of zero when in fact clouds were present, which  
in turn leads to overestimations of UV radiation by TROPOMI. Quality flags related to topography and cloud optical thickness  
are included in the TROPOMI surface UV product and could be used to identify such cases of heterogeneous topography.

490 As the TROPOMI pixel is between  $7 \times 3.5 \text{ km}^2$  (nadir) and  $9 \times 14 \text{ km}^2$  (edge of swath), also the surroundings of a site are  
included in the pixel. Coastal high-latitude sites are particularly challenging (e.g., NYA, AND, PAL and MAR) because both  
open ocean and land covered by snow can be within the area of a pixel. The albedo is determined by the climatology of the  
central point of the pixel which can be open ocean even if the radiation field of the site is characterized by the surrounding  
snow cover. As for heterogeneous topography, the quality flagging also includes a flag for heterogeneous surface albedo.

495 The same applies for the impact of cloudiness when clouds are non-uniformly located around the site due to topography  
or changes in surface (e.g. sea/ground). For example the site itself is free from clouds but there can be small cumulus clouds  
at the edge of the pixel which increase the reflection towards the satellite. In that case the TROPOMI would most probably  
underestimate the UV irradiance as the small fraction of clouds is considered as cloudiness of the whole pixel. Under scattered  
clouds, the UV radiation at the surface can be larger than during clear skies (Calbó et al., 2005; Jégou et al., 2011). This  
phenomenon occurs when the direct radiation from the Sun is not obstructed and additional radiation is scattered by clouds to  
the radiometer at the surface. The TROPOMI algorithm does not consider these situations, resulting UV levels that are too low.  
500 This phenomenon is likely one reasons for the underestimation found in the TROPOMI UV dose rates during high UV levels  
at tropical sites.

The calculation of daily UV doses is based on the cloud optical depth at the time of the satellite overpass and the assumption  
that cloud cover is constant throughout this day. While this simplification leads to uncertainties, day-to-day variations in daily  
doses measured by the radiometers at the ground are generally well reproduced by TROPOMI. One could expect systematic  
505 biases at sites in which a diurnal cycle occurs, e.g. differences in cloudiness between morning and afternoon like orographic  
clouds forming in the afternoon. This is the case for Sait-Denis, however, the effect on the bias of TROPOMI data is not very  
large. Rapidly changing cloudiness can be seen as large deviation between UV irradiances measured at successive pixels.

510 The smaller pixel size of TROPOMI increases the amount of data compared to previous satellite instruments. Only TROPOMI  
pixels with center points less than 5 km away from a ground station were included in this study. For certain sites, especially  
at high latitudes, this increased the number of data to include more than 5 pixels for each overpass. In this study the results  
were calculated for each pixel separately. The results did not differ significantly when changing the method to use either 1) the  
average of the pixels or 2) the nearest pixel of the site.

## 6 Conclusions

515 The TROPOMI Surface UV Radiation Product was validated against ground-based measurements at 25 sites for the period  
from 1 Jan 2018 to 31 Aug 2019. TROPOMI overpass irradiances at 305, 310, 324 and 380 nm were compared against spec-  
roradiometer measurements from 13 sites (7 sites for 380 nm). No major differences between results of different wavelengths



were found, except that cloudiness affected the irradiance at 380 nm by enlarging the spread of the deviation from ground-based measurements. The median relative difference between TROPOMI data and ground-based measurements was within  $\pm 10\%$  at 11, 9, 10 and 6 sites for irradiance at 305, 310, 324 and 380 nm, respectively, during snow free surface conditions. More than  
520 half of the satellite data were within  $\pm 20\%$  from ground-based measurements.

TROPOMI overpass erythemally weighted dose rates were compared against dose rates measured by spectroradiometers at 11 sites and by broadband or multiband radiometers at two sites. The TROPOMI overpass UV index, which is directly proportional to the erythemally weighted dose rate, was compared against broadband and multichannel radiometer measurements at 12 sites. These validation results showed that the median relative difference between TROPOMI and ground-based dose rates  
525 was within  $\pm 10\%$  and  $\pm 5\%$  at 18 and 10 sites, respectively, for snow free surface conditions.

TROPOMI erythemally weighted daily doses were compared against spectroradiometer, broadband and multichannel radiometer measurements at 16 sites. The median relative difference was within  $\pm 10\%$  and  $\pm 5\%$  at 11 and 8 sites, respectively, for snow free surface conditions. For both dose rates and daily doses, 60 – 80% of TROPOMI data were within  $\pm 20\%$  from ground-based data at most of the sites for snow free surface conditions.

530 For all UV parameters an increased over- or underestimation was found at challenging conditions for satellite retrievals. Those are related to non-homogeneous topography (large altitude differences within the satellite pixel), non-homogeneous surface albedo (e.g. open water + snow cover on land within the satellite pixel), surface albedo differing from the albedo climatology used by the TROPOMI (at high latitudes where year-to-year changes are large) and high UV levels in tropics. The TROPOMI UV parameters include quality flagging to help identifying many of the above mentioned cases.

535 Retrieving the correct cloud optical depth is still challenging over snow albedo when reflections from snow and clouds are misinterpreted or confused with each other. The underestimation in satellite retrievals is related to the albedo climatology used by TROPOMI, which should be updated. Also, the current albedo climatology does not change from year to year while the actual albedo (e.g., timing of snow melt) can change a lot. The challenge is that the albedo climatology should be retrieved in at least as small pixel size as used by the TROPOMI, preferably from the TROPOMI data itself.

540 The TROPOMI Surface UV Radiation Product continues the former TOMS and OMI UV time series with an upgraded spatial resolution. The nominal life time is 7 years after which the S5P mission hopefully could be extended as has been the case for e.g. OMI with already over 15 years of observations. Together these UV time series based on satellite retrievals form a unique 30 years long global data set which can be used for multiple UV impact studies all over the World. For this purpose special efforts to develop homogenized long term satellite time series are needed.

545 *Data availability.* The station overpass files and documentation are available from [https://nsdc.fmi.fi/data/data\\_s5puv.php](https://nsdc.fmi.fi/data/data_s5puv.php). Ground-based data are available from the authors.



- Author contributions.* K. Lakkala analyzed the data and led the manuscript preparation. She was responsible for the QC/QA and processing of Sodankylä data, and supervised the QC/QA and monitoring of Marambio data.
- J. Kujanpää implemented the TROPOMI UV data processor and processing system, processed the L2 data and extracted the station overpass files, participated in data analysis and contributed to the writing of the manuscript.
- 550 C. Brogniez supervised the monitoring, processing and QC/QA of French spectroradiometer's data and contributed to the writing of the manuscript.
- N. Henriot contributed to the sensitivity study regarding the pixel selection and the UVQAV flag.
- A. Arola contributed to the development of the TROPOMI UV algorithm, data analysis and the writing of the manuscript.
- 555 M. Aun processed Marambio UV data and contributed to the writing of the manuscript.
- F. Auriol: Responsible of calibration of the three French spectroradiometers.
- A. F. Bais supervised the monitoring, calibration and quality control of THE data.
- G. Bernhard: Responsible for QC/QA and processing of Palmer SUV-100 data, and contributed to the writing of the manuscript.
- V. De Bock is co-responsible (together with Hugo De Backer) for the QC/QA and processing of Uccle Brewer spectroradiometer data.
- 560 M. Catalfamo contributed in calibration of VDA and OHP spectroradiometers.
- C. Deroo: Responsible of automatic processing of the three French spectroradiometer's data.
- H. Diémoz: Responsible of QC/QA and processing of Aosta data and contributing to the writing of the manuscript.
- L. Egli: Responsible for QC/QA and processing of the Davos Brewer spectroradiometer data.
- J.-B. Forestier: Technical manager of MAH,ANT,ROD and STD.
- 565 I. Fountoulakis: Contribution in QC/QA and processing of Aosta data and contributing to the writing of the manuscript.
- R. D. Garcia responsible for QC/QA and processing of data of the Izaña BSRN.
- J. Gröbner: Responsible for QC/QA of the Davos measurements and contributed to the writing of the manuscript.
- S. Hassinen: Overseeing the work in the group and contributed to the TROPOMI data handling.
- A. Heikkilä: Responsible for the maintenance of Helsinki Brewer and QC/QA and processing of Helsinki Brewer data.
- 570 S. Henderson: Responsible for QC/QA and processing of data of ALI and MEL and contributed to the writing of the manuscript.
- G. Hülsen: Responsible for QC/QA and processing of the Davos Qasume spectroradiometer and broadband data and contributed to the writing of the manuscript.
- B. Johnsen: Responsible for the QC/QA of the Norwegian UV-network sites and contributed to the writing of the manuscript.
- N. Kalakoski contributed to the development of the TROPOMI UV algorithm.
- 575 A. Karanikolas contributed in calibration and analysis of THE data.
- T. Karppinen contributed to the QC/QA of the Sodankylä Brewer data.
- K. Lamy: Responsible for QC/QA and processing of data of MAH, ANT, ROD and STD and contributed to the writing of the manuscript.
- S. Leon: Co-Responsible for QC/QA and processing of IZO Brewer spectroradiometer data.
- A. V. Lindfors contributed to the development of the TROPOMI UV algorithm and writing of the manuscript.
- 580 J.-M. Metzger contributed in calibration and maintenance of OPA spectroradiometer.
- F. Minvielle contributed in calibration of VDA and OHP spectroradiometers.
- H. B. Muskatel: is responsible for the QC/QA and data processing of the measurements from BET, JER and EIL.
- T. Portafaix: P.I. of UV-Indien network and responsible for QC/QA of MAH,ANT,ROD and STDR.
- A. Redondas: Team leader of Regional Brewer Calibration Center-Europe (RBCC-E). Co-Responsible for QC/QA and processing of IZO



585 Brewer spectroradiometer data.

R. Sanchez: Responsible in SMN for Marambio UV measurements.

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T. Svendby: Responsible for GUV measurements at Blindern, Andøya and Ny-Ålesund.

J. Tamminen participated in the development of TROPOMI UV data product, discussion of the validation results and manuscript preparation.

590

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