Authors' response to comments and suggestions from referees and the editor *Black, italic: Referee's comments* Blue: Author's reply Red: Changes in the original discussion paper

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1. in the manuscript, you state in line 289 that "the NO2 sources in Feicheng are mainly related to traffic and concentrated along the S104". I frankly cannot see that in the figure and suggest to remove this statement.

Removed.

#### 10

2. in the caption of Figure 10, you write "NO2 emission". Please replace by "NOx emission" as emissions are usually in the form of NO, not NO2.

Replaced NO2 emission with NOx emission.

#### 15

3. when introducing the car-DOAS measurements, please add that the retrieval window in the UV differs from the one used for the airborne observations.

We added this statement in line 338.

## 20

It is worth noting that the retrieval window in the mobile DOAS observations differs from the one used for the airbrine observations.

4. please add acknowledgements and data sources for TROPOMI and LANDSAT 8 data.

## 25

We added acknowledgements and data sources for TROPOMI and LANDSAT 8 data.

Acknowledgments. We would like to thank Thomas Danckaert, Caroline Fayt and Michel Van Roozendael for help on QDOAS software. We are thankful to the following agencies for providing the satellite data: The Sentinel 5 Precursor

30 TROPOMI Level 2 NO<sub>2</sub> product is developed by KNMI with funding from the Netherlands Space Office (NSO) and processed with funding from the European Space Agency (ESA). TROPOMI data can be downloaded from https://s5phub.copernicus.eu. Landsat 8 OLI data have been produced, archived, and distributed by the U.S. Geological Survey (USGS). The original Landsat surface reflectance algorithm was developed by Dr. Eric Vermote, NASA Goddard Space Flight Center (GSFC). Landsat 8 OLI data are available at https://earthexplorer.usgs.gov/.

5. in the introduction, it would be good to also make reference to some other imaging DOAS instruments that have been flown for NO2 measurements:

We added these references of other imaging DOAS instruments in section 1.

40

For the purpose of retrieval of urban NO<sub>2</sub> horizontal distribution, Popp et al. (2012), General et al. (2014), Schönhardt et al. (2015), Lawrence et al. (2015), Nowlan et al. (2016), and Lamsal et al. (2017) performed their measurements separately in Zürich, Switzerland, Indianapolis and Barrow, USA, Ibbenbüren, Germany, Leicester, England, Houston, USA, and Maryland, USA.

Lamsal, L. N., S. J. Janz, N. A. Krotkov, K. E. Pickering, R. J. D. Spurr, M. G. Kowalewski, C. P. Loughner, J. H. Crawford, W. H. Swartz, and J. R. Herman (2017), High-resolution NO<sub>2</sub> observations from the Airborne Compact Atmospheric Mapper: Retrieval and validation, J. Geophys. Res. Atmos., 122, 1953–1970, doi:10.1002/2016JD025483.

50 Nowlan, C. R., Liu, X., Leitch, J. W., Chance, K., González Abad, G., Liu, C., Zoogman, P., Cole, J., Delker, T., Good, W., Murcray, F., Ruppert, L., Soo, D., Follette-Cook, M. B., Janz, S. J., Kowalewski, M. G., Loughner, C. P., Pickering, K. E., Herman, J. R., Beaver, M. R., Long, R. W., Szykman, J. J., Judd, L. M., Kelley, P., Luke, W. T., Ren, X., and Al-Saadi, J. A.: Nitrogen dioxide observations from the Geostationary Trace gas and Aerosol Sensor Optimization (GeoTASO) airborne instrument: Retrieval algorithm and measurements during DISCOVER-AQ Texas 2013, Atmos. Meas. Tech., 9, 2647–2668,

55 https://doi.org/10.5194/amt-9-2647-2016, 2016.

Schönhardt, A., Altube, P., Gerilowski, K., Krautwurst, S., Hartmann, J., Meier, A. C., Richter, A., and Burrows, J. P.: A wide field-of-view imaging DOAS instrument for two-dimensional trace gas mapping from aircraft, Atmos. Meas. Tech., 8, 5113-5131, doi:10.5194/amt-8-5113-2015, 2015.

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General, S., Pöhler, D., Sihler, H., Bobrowski, N., Frieß, U., Zielcke, J., Horbanski, M., Shepson, P. B., Stirm, B. H., Simpson, W. R., Weber, K., Fischer, C. and Platt, U.: The Heidelberg Airborne Imaging DOAS Instrument (HAIDI) – a novel imaging DOAS device for 2-D and 3-D imaging of trace gases and aerosols, Atmos. Meas. Tech., 7(10), 3459–3485, doi:10.5194/amt-7-3459-2014, 2014.

<sup>45</sup> 

6. Line 133: Can the authors quantify in the text the radiance threshold applied?

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We used a radiance threshold of 12.8  $\mu$ W·cm<sup>-2</sup>·sr<sup>-1</sup>·nm<sup>-1</sup> at 450 nm.

Furthermore, a radiance threshold of 12.8  $\mu$ W·cm<sup>-2</sup>·sr<sup>-1</sup>·nm<sup>-1</sup> at 450 nm is set to neglect some over-illuminated ground pixels inside the flight area, which are usually caused by the presence of clouds or water mirror reflection.

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7. Equation 1 is valid for a total VCD, but not tropospheric VCD as is referred to below this equation. The stratospheric slant column must also be subtracted from the dSCD. In lines Line 167-170, it is stated that the stratospheric column cancels out during flight from the reference but in the comment back to the reviewers it was quantified as a difference up to 8e14 due to the changes in the slant path through the atmosphere, which is not stated in the text. This is an impact and should either be added to this product or by quantified in the error estimate. It is a change of 23% over the time period of the flight...

We added this stratospheric correction in Sect 4.2.

Changes in the stratospheric NO<sub>2</sub> could also propagate to the measured tropospheric columns of UVHIS. Under the assumption of a constant stratosphere in time and space during the flight, the changes in the SZA impact the column difference between the measurement and the reference. To correct the change in the stratospheric NO<sub>2</sub> SCD, we apply a geometric approximation of the stratospheric AMF with a stratospheric VCD of  $3.5 \times 10^{15}$  molec cm<sup>-2</sup> from TROPOMI product. The maximum change in the stratospheric SCD with respect to the reference, was  $8 \times 10^{14}$  molec cm<sup>-2</sup>.

90 8. Line 195: State the range of AOD measurements from MODIS for this flight day.

The MODIS AODs in the flight area range from 0.142 to 0.365.

9. Line 228: Inconsistencies are expected because one is radiance and the other is surface reflectance. They aren't the same thing thought they should be similar in spatial pattern. I don't see differences that would be caused by time differences with how it is displayed here. The one detail that does stick out though is that roadway isn't visible in the radiance measurements even though UVHIS should be higher resolution than Landsat, so can this be explained just by the color bar range choice? Perhaps but the unweighted averaging within the grid as discussed in Section 4.4? It would be nice to have a sentence or two about this.

Actually, the roadway is visible in Fig. 9 (a), but the width of the road looks thinner compared to Fig. 9 (b). We replaced this figure with a higher resolution one. We modified the last sentence in this paragraph.

Fig. 9 also shows several slight inconsistencies between the UVHIS measured radiance and the Landsat 8 surface reflectance

105 product. For example, the east-west main road looks thinner in Fig. 7 (a) compared to Figs. 7 (b) and (c). This could be explained by the relatively higher spatial resolution performance of the UVHIS and the resampling of Landsat 8 pixels.

10. Line 278: The authors refer to this 'weaker' plume, but it has a higher column density than the last described plume.

110 We deleted the word 'weaker'.

11. Is the noise in Figure 11 primarily attributable to surface reflectivity?

We thought the noise in Fig. 11 is primarily attributable to random noise, because the spatial patterns of VCD and surface 115 reflectance (or radiance) are inconsistent.

12. Line 297: This is not correct. Areas with NO2 sources can vary dramatically with time. This is a sample size of 2. Emissions from sources aren't always constant and their footprints can change dramatically due to meteorology as well. In the case of region B, the variability is almost just as large as it is in region a (maybe even more so). The only consistent region is region C.

We agreed this correction and modified the last few sentences in this paragraph.

In these three regions, only region C is temporally consistent with relatively low NO<sub>2</sub> columns, whilst a large temporal variety of NO<sub>2</sub> VCDs exists in region A and region B because of inconstant emission sources and changing meteorology.

13. Line 315: The authors refer to AMF of 2 and AMF of 1.8 for the reference in the same sentence for the reference. This is confusing. I think we need the AMF for TROPOMI in this case to convert 1e15 VCD to its SCD then divide by the AMF from UVHIS.

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We thought the TROPOMI AMF of the reference may be inappropriate considering the different viewing angles. The AMF of 1.8 is near the mean value of the UVHIS AMFs in the reference area, and the AMF of 2.0 is near the mean value of all UVHIS AMFs.

- 14. Line 365: This comment is in relation to the low bias by UVHIS and high bias by mobile DOAS. It seems that the mobile 135 DOAS would probably be more sensitive to near surface pollution than the aircraft. Is the assumption of the a priori profile a reason for the aircraft being lower than the mobile DOAS and is there an optimal profile between the two that would result in the two being closer to 1:1. I am not suggesting to redo all work with a new a priori but perhaps doing the sensitivity test like was done on the airborne work on the mobile DOAS and see if that is a plausible cause for the biases. Then reflect on 140 that in the text.

## We did a sensitivity test, the detail is in next answer.

15. Last sentence in Section 6: We haven't been given reason to believe that the spatial footprint of the mobile DOAS is 145 smaller than the UVHIS pixels. Therefore, this sentence is not a valid reason for these results. The above comment seems more plausible.

## We deleted this sentence and modified to

- 150 A sensitivity test of the AMF on the NO<sub>2</sub> profile was performed for all co-located measurements, using a box profile of 500 m. Compared to the box profile of 2 km, the UVHIS AMFs decreased by an average of 17 %, whilst the mobile DOAS AMFS decreased by an average of 2.7 %. This results suggest that a more realistic profile with the NO<sub>2</sub> layer closer to the ground could improve the slope closer to unity.
- 155 16. Line 20 + final paragraph in Section 6: It is nice to see this improvement in the correlation between Fig 14 a and b. Is the correlation different because of the time difference between measurements or because of the location near the steel factory? It seems more likely to focus on is the difference in time between measurements in the text rather than focusing on the location near the steel plant. This should be detailed in the caption of Figure 14 as well.

We added the time difference information in the caption of Fig. 14. 160

17. Title has a misspelled word: Imaing = imaging

#### Fixed

<sup>18.</sup> Line 44: Change 'lower' to 'coarser'

## Fixed

170 19. Delete Lines 120-122. These details are on the next page and fit better there.

## Deleted

20. Line 215: Delete Other RTM parameters used in the AMF calculations are also provided in Fig. 7. Because there are no other parameters shown that I can see.

## Deleted

21. Line 236: 'estimated' PBL height.

## 180

Fixed

22. Should section 4.4 go before 4.3 or even before 4.1 because geolocation and other steps would have to be performed before AMF analysis to link it to surface reflectivity and other attitude parameters?

#### 185

We added a geo-referencing part in Sect 4.1 and changed the title of Sect 4.4 to 'Resampling and mapping'.

23. Line 264: refer to Figure 10.

190 We added 'As shown in Fig. 10' in this sentence.

24. Line 310: the mobile DOAS system hasn't been introduced yet so it seems inappropriate to bring up here.

## Deleted

## 195

25. Table 3 and Table 5: In the text, please explain what the intensity offset is.

This is the offset parameter in the DOAS fitting for possible instrumental and/or atmospheric stray light or residual dark current signal. We changed the 'Intensity offset' to 'Offset' to reduce misunderstandings.

26. Figure 4, 10, 12, and 13: Please make the numbers indicating the point sources larger. They're hard to read.

# We updated these figures.

*27.* The authors included water vapour in the DOAS analysis, unfortunately this did not improve the fit. Nevertheless the water cross section should be added in Table 5.

In the manuscript, Table 5 lists the parameters of mobile DOAS system. We did not add a water vapour cross section because the absorption of water vapour is weak in the 356-376 nm fit window of the mobile DOAS.

Relevant changes:

## 235

- 1. Modified the title.
- 2. Modified the affiliation.
- 3. Added a stratospheric correction in the DOAS analysis.
- 4. Added a sensitivity test of the AMF on the NO2 profile for the co-located measurements.
- 240 5. Updated Figs. 4, 9, 10, 12, 13.
  - 6. Fixed grammatical problems in the text.

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# **The first**<u>First</u> high-resolution tropospheric NO<sub>2</sub> observations from the Ultraviolet Visible Hyperspectral <u>Imaing Imaging</u> Spectrometer (UVHIS)

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Abstract. We present a novel airborne imaging differential optical absorption spectroscopy (DOAS) instrument: Ultraviolet Visible Hyperspectral Imaging Spectrometer (UVHIS), which is developed for trace gas monitoring and pollution mapping. Within a broad spectral range from of 200 to 500 nm - operated and operating in three channels, the spectral resolution of UVHIS is better than 0.5 nm. The optical design of each channel comprises a fore-optics with a field of view (FOV) of 40°, 280 an Offner imaging spectrometer, and a charge-coupled device (CCD) array detector of 1032 × 1072 pixels. A first demonstration flight using UVHIS was undertaken conducted on 23 June 2018, above an aera of approximate 600 km<sup>2</sup> area in Feicheng, China, with a spatial resolution of about  $25 \times 22$  m<sup>2</sup>. Measurements of nadir backscattered solar radiation of channel 3 are used to retrieve tropospheric vertical column densities (VCDs) of NO<sub>2</sub> with a mean total error of  $3.0 \times 10^{15}$ molec cm<sup>-2</sup>. The UVHIS instrument clearly detected several emission plumes transporting from south to north, with a peak 285 value of  $3 \times 10^{16}$  molec cm<sup>-2</sup> in the dominant one. The UVHIS NO<sub>2</sub> vertical columns are well correlated consistent with the ground-based mobile DOAS observations, with a correlation coefficient of 0.65 for all co-located measurements, a correlation coefficient of 0.86 for the co-located measurements that only circled the steel factory, and a slight underestimation for the polluted observations. This study demonstrates the capability of UVHIS for NO<sub>2</sub> local emission and 290 transmission monitoring.

#### 1 Introduction

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Nitrogen oxides (NO<sub>x</sub>), the sum of nitrogen monoxide (NO) and nitrogen dioxide (NO<sub>2</sub>), <u>plays-play</u> a key role in the chemistry of the atmosphere, such as the ozone destruction in the stratosphere (Solomon, 1999), and the secondary aerosol formation in the troposphere (Seinfeld and Pandis, 2016). In the troposphere, despite lightning, soil emissions and other natural processes, the main sources of NO<sub>x</sub> are anthropogenic activities like fossil fuel combustion by power plants, factories, and road transportation, especially in the urban and polluted regions. As an indicator of anthropogenic pollution which leads to negative effects both on the environment and human health, the amounts and spatial distributions of NO<sub>x</sub> attract largehave attracted significant attention. For example, China becomes has become one of the largest NO<sub>x</sub> emitters in

the world due to its fast industrialization industrialisation, meanwhile China is has also experiencing experienced a series of

300 severe air pollution problems in recent years (Crippa et al., 2018; An et al., 2019). Therefore, measuring the NO<sub>x</sub> distribution by application of applying different techniques, would benefit the pollutant emission detection and the air quality trend forecast (Liu et al., 2017; Zhang et al., 2019).

Compared to NO, Nitrogen dioxide (NO<sub>2</sub>) is more stable in the atmosphere. Based on the characteristic absorption structures of NO<sub>2</sub> in the ultraviolet-visible spectral range, the differential optical absorption spectroscopy (DOAS) technique has been

- 305 applied to retrieve the light path integrated densities from different platforms (Platt and Stutz, 2008). Combining Combined the imaging spectroscopy technique, imaging DOAS instruments were developed in recent years to determine the trace gases temporal variation as well as and the two dimensional distribution of trace gases. The Global global horizontal distribution of tropospheric NO<sub>2</sub> and other trace gases has been mapped and studied by several space-borne sensors, including SCIAMACHY (Scanning Imaging Absorption Spectrometer for Atmospheric CHartographY; Bovensmann et al., 1999),
- GOME (Global Ozone Monitoring Experiment; Burrows et al., 1999), GOME-2 (Munro et al., 2016), OMI (Ozone Monitoring Instrument; Levelt et al., 2006) and TROPOMI (TROpospheric Ozone Monitoring Instrument; Veefkind et al., 2012). The Environmental trace gases Monitoring Instrument (EIMEMI; Zhao et al., 2018; Cheng et al., 2019; Zhang et al., 2020), as the first designed space-borne trace gas sensor in China, was launched on 9 May 2018, on-board the Chinese GaoFen-5 (GF5) satellite. In terms of The spatial resolution, the majority of these of most space-borne sensors are is lower
- 315 <u>coarser</u> than 10 × 10 km<sup>2</sup>, except <u>for that of</u> TROPOMI with a relative higher resolution of which is 3.5 × 5.5 km<sup>2</sup>. In order toTo achieve <u>a</u> spatial resolution higher than 100 × 100 m<sup>2</sup> to investigate for investigating the spatial distribution in urban areas and individual source emissions, several researchers have applied <u>the</u>-imaging DOAS instruments on airborne platforms. The airborne imaging DOAS measurement was <u>firstly-first</u> performed by Heue et al. (2008) over <u>the</u> South African Highveld plateau. For the purpose of retrieval of To retrieve urban NO<sub>2</sub> horizontal distribution, Popp et al. (2012),
- <u>General et al. (2014), Schönhardt et al. (2015), and Lawrence et al. (2015), Nowlan et al. (2016), and Lamsal et al. (2017)</u> performed their measurements separatelyrespectively took measurements in Zürich, Switzerland, <u>; Indianapolis and Barrow,</u> <u>USA; Ibbenbüren, Germany; and Leicester, England; Houston, USA; and Maryland and Washington DC, USA</u>. In 2013, an airborne measurement focusing on source emissions was <u>performed taken</u> in China, over Tianjin, Tangshan and <u>the Bohai</u> <u>gulf Bay</u> (Liu et al., 2015). An inter-comparison study of four airborne imaging DOAS instruments over Berlin, Germany;
   suggests a good agreement between different sensors, and the effectiveness of imaging DOAS to reveal<u>in revealing</u> the fine-

scale horizontal variability in tropospheric NO<sub>2</sub> in an-urban context (Tack et al., 2019). Here we present a novel airborne imaging DOAS instrument: Ultraviolet Visible Hyperspectral Imaging Spectrometer (UVHIS), <u>which was</u> designed and developed by Anhui Institute of Optics and Fine Mechanics, Chinese Academy of Sciences (AIOFM, CAS). As a hyperspectral imaging sensor with <u>a</u> high spectral and spatial resolution, UVHIS is designed

330 to be operated for operation on an aircraft platform for the purpose of atmospheric trace gas measurements and pollution monitoring over large area areas in a relative short time frame. By using the DOAS technique and geo-referencing, the two

<u>two-</u>dimensional spatial distribution of tropospheric  $NO_2$  of its first demonstration flight over Feicheng, China is also presented in this paper.

<u>This paper is organised as follows</u> The organization of this paper is listed as follows: Sect.Section 2 presents a technical description of <u>the</u> UVHIS system, as well as and its preflight calibration results. Section 3 introduces the detailed information of its first research flight above over Feicheng, China. Section 4 describes the developed algorithm for <u>the</u> retrieval and geographical mapping of tropospheric NO<sub>2</sub> vertical column densities from hyperspectral data. Section 5 presents the retrieved NO<sub>2</sub> column densities, and Section 6 compares <u>the</u> airborne measurements with <u>the</u> correlative ground-based data sets from a mobile DOAS system.

## 340 2 Instrument Details

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#### 1.1 The UVHIS instrument

UVHIS is a hyperspectral instrument measuring nadir backscattered solar radiation in the ultraviolet and visible wavelength region from 200 to 500 nm. The instrument is operated in three channels with the wavelength ranges of 200-276 nm (channel 1), 276-380 nm (channel 2) and 380-500 nm (channel 3) for minimal stray light effects and highest spectral performance. The main characteristics of UVHIS are summarized-summarised in Table 1.

- Figure 1 shows the optical bench of channel 3 and that those of the other two are similar. The optical design of each channel comprises a telecentric fore-optics, an Offner imaging spectrometer, and a two-two-dimensional charge coupled device (CCD)CCD array detector. The Offner imaging spectrometer consists of a concave mirror and a convex grating. The backscattered Backscattered light below the aircraft is collected by a wide-field telescope with a FOV of 40° in the across-
- track dimension. After passing through a bandpass filter and a 12.5 mm long entrance slit in the focal plane, <u>the light is</u> reflected and diffracted by a concave mirror and a convex grating. The dispersed light is imaged onto a frame transfer CCD detector, which consists of 1032 × 1072 individual pixels. For the <u>purpose of alignment and slight adjustment of the</u> spectrometer, only the central 1000 rows of pixels are well illuminated in the across-track dimension. In the wavelength dimension, the image covers central 1024 columns of pixels on the CCD detector-<u>while, whilst</u> the left and right edges are used to monitor dark current. The spectral sampling and spectral resolution of all three channels can be found in Table 1.
- To reduce dark current and improve the signal-to-noise ratio (SNR) of the instrument, the CCD detector is thermally stabilized stabilised at -20 °C with a temperature stability of  $\pm 0.05$  °C (Zhang et al., 2017). However, the optical bench is not thermally controlled, because the instrument is mounted inside the aircraft platform which is temperature consistent at has a constant temperature of 20 °C. The UVHIS is mounted on a Leica PAV-80 gyro-stabilized stabilised platform that provides
- 360 angular motion compensation. A high-grade Applanix navigation system on-board is used to receive the position (i.e. latitude, longitude, and elevation) and orientation (i.e. pitch, roll, and heading) information, which is are required for accurate geo-referencing. The UVHIS instrument telescope collects the solar radiation backscattered from the surface and

atmosphere through a fused silica window <u>on-at</u> the bottom of the aircraft. In the case of  $NO_2$  measurement, all observations in this study only use the channel 3.

#### 365 2.2 Preflight calibration

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Spectral and radiometric calibration in the laboratory were done were performed in the laboratory prior to the flights to reduce errors in spectral analysis.

For radiometric calibration, we used an integrating sphere with a tungsten halogen lamp for channel 2 and channel 3. While for For channel 1, a diffuser plate with a Newport xenon lamp is-was used for a-sufficient ultraviolet output. With the help of a well-well-calibrated spectral radiometer to monitor the radiance of calibration system, the digital numbers (DNs) from the CCD detectors of the three channels can be converted to radiance correctly. The uncertainty Uncertainty of absolute radiance calibration of the UVHIS is 4.89 % for channel 1, 4.67 % for channel 2 and 4.42 % for channel 3.

The preflightPreflight wavelength calibration was also performed in the laboratory, using a mercury–argon lamp and a tunabletuneable laser as light sources. We modelled the slit function of the UVHIS using a symmetric Gaussian function.

375 Spectral registration and slit function calibration are-were achieved by least square fitting of the characteristic lines in the collected spectra. Table 2 lists the retrieved full-width at half maximums maximum values (FWHMs) for channel 3. Figure 2 shows the measured slit functions at 450.504 nm for 9-nine viewing angles (i.e. -20°, -15°, -10°, -5°, 0°, 5°, 10°, 15°, 20°), and the respective retrieved slit function shapes using a symmetric Gaussian function. These Gaussian fit results suggest that a symmetric Gaussian function is a reasonable assumption for the slit shape in all viewing directions.

## 380 3 Research flight

The first demonstration flight <u>above\_over</u>Feicheng <u>eityCity</u>, Shiheng <u>town\_Town\_</u>and neighbouring rural areas was <u>performed\_conducted</u> on 23 June 2018, aiming at producing tropospheric NO<sub>2</sub> field maps of <u>a</u> large area in a relatively short time frame. Feicheng is a county-level city in Shandong province, <u>approximatelyabout</u> 410 km away from Beijing. Figure 3 shows the TROPOMI NO<sub>2</sub> tropospheric observation on 23 June<sub>7</sub> 2018, with the background Google map and the location of Feicheng. The flight area is located on the south bank of the Yellow River, at the western foot of Mount Tai. <u>The\_</u>UVHIS was operated from the Y-5 aircraft at <u>an</u> altitude of 3 km above sea level,—<u>which is</u> higher than the height of planetary boundary layer (PBL), with an average aircraft ground speed of 50 m/s. An overview of the observation area and <u>the flight</u> lines are provided in Fig. 4. The aircraft took off at local noon from the airfield in Pingyin <u>countyCounty</u>, <u>approximatelyabout</u> 19 km northwest of the centre of the field. An area of approximately 600 km<sup>2</sup> was covered in 3 hourh, under clean sunny and cloudless conditions with low-speed southerly winds.

Research The research flight consists of included 13 parallel lines in the east-west direction, starting from the lower left corner in Fig. 2. The distance between adjacent lines is was 1.5 km, while whilst the swath width of each individual line is aboutwas approximately 2.2 km. Gapless coverage between adjacent lines can be guaranteed in this pattern because of the

adequate overlap. To validate the NO<sub>2</sub> column densities retrieved from the UVHIS by comparison to ground measurements,

395 mobile DOAS measurements were <u>performed\_taken</u> inside the research area on the same day. As shown in Fig.\_4, the measurements of the mobile DOAS system circle<u>d</u> around the steel factory and the power plant, which are the presumed major emission sources inside the observation area.

In the condition of spatial binning by 10 pixels across track, the across track spatial resolution of the ground pixel is about 22 m. At typical aircraft ground speed of 50 m/s and integration time of 0.5 s, the along track spatial resolution of the ground pixel is about 25 m.

#### 4 Data processing chain

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The NO<sub>2</sub> tropospheric vertical column density (VCD) retrieval algorithm of <u>the</u>UVHIS consists of four major steps. First, some necessary <u>preprocessingpre-processing</u> procedures are required before any spectral analysis of <u>the</u>UVHIS data. Next, <u>the</u>UVHIS spectral data after <u>preprocessing-pre-processing</u> are <u>analyzed\_analysed\_in</u> a suitable wavelength region by application\_applying of the well-established DOAS technique. After thatThen, the air mass factors (AMFs) are calculated for every observation based on <u>the</u>SCIATRAN radiative transfer model to convert the slant column densities (SCDs) to tropospheric vertical column densities. In the final step, <u>the geo-referenced combining with sensor position and orientation information</u>, NO<sub>2</sub> VCDs are <u>geo-referenced</u>resampled and overlaid onto Google satellite map layers.

## 4.1 Pre-processing

- 410 The preprocessing processing procedure before spectral analysis includes data selection, geo-referencing, dark current correction, spatial binning<sub>7</sub> and in-flight calibration. First, the spectral data acquired during <u>U turns of aircraft U-turns</u> are removed in the processing because of the large and changing orientation angles. Also Furthermore, a threshold of radiance values radiance threshold of 12.8 μW cm<sup>-2</sup> sr<sup>-1</sup> nm<sup>-1</sup> at 450 nm is set to neglect some over-illuminated ground pixels inside the flight area, which are usually caused by the presence of clouds or water mirror reflection. During the entire flight, the sun
- 415 glinted on water occurred several times in the southern part of the flight area, especially above the river near the reference area. However, cloud was not present due to the clean clear-sky weather condition.
  Accurate geo referencing is essential for emission source locating and data comparison, and can be achieved with the sensor position and orientation information recorded by the navigation system and IMU on-board.

The dark<u>Dark</u> current correction is performed based on the measurement at the start of the entire flight by blocking the foreoptics, which is necessary to improve <u>the</u> instrument performance and reduce <u>the</u> analysis error in DOAS fit.

In order to increase the SNR of the instrument and the sensitivity to NO<sub>2</sub>, the raw pixels of <u>the</u> imaging DOAS are usually aggregated in <u>the</u> across-track direction and along-track directions. According to photon statistics when only shot noise is considered, the SNR should rise with <u>the</u> square root of <u>the</u> number of binned spectra. However, this improved SNR of <u>the</u>

instrument results a reduction of in reduced spatial resolution. In the data analysis of the Feicheng flight, we use the binning

- 425 of 10 pixels in <u>the across-track direction</u>, resulting in a ground pixel size of about approximately  $25 \times 22 \text{ m}^2$ .
- Since-Given that the wavelength-to-pixel registration and the slit function shape of the UVHIS could change compared to laboratory calibration results, in-flight wavelength calibration is essential for the next DOAS analysis. This in-flight wavelength calibration is achieved by fitting the measured spectra to a high-resolution solar reference (Chance and Kurucz, 2010) with slit function convolution and wavelength shift. The nominal wavelength-to-pixel registration determined in
- 430 laboratory calibration, is used as initial values in the iteratively fitting procedure for convergence to the optimal solution. The effective shifts and FWHMs of different across-track positions are plotted in Fig. 5. Results The results at three wavelengths are shownpresented as follows: blue for 430 nm (the start of the analysis wavelength region), green for 450 nm (the middle of the analysis wavelength region) and red for 470 nm (the end of the analysis wavelength region).

## 4.2 DOAS analysis

- After-preprocessing\_pre-processing, the observed UVHIS spectra are analysed by the application of using the QDOAS (Danckaert et al., 2020) software in order to retrieve the NO<sub>2</sub> slant column densities. The basic idea of the DOAS approach is to separate broadband signals like surface reflectance and Rayleigh scattering, and narrow-band signals like trace gas molecular absorption. The fitting fit window is within 430 and \_470 nm-wavelength region, considered to contain strongly structured NO<sub>2</sub> absorption features, and with low interference of other trace gases such as O<sub>3</sub>, O<sub>4</sub>, and water vapor. The Absorption-absorption cross-sections of NO<sub>2</sub> and other trace gases and a synthetic Ring spectrum are simultaneously fitted to the logarithm of the ratio of the observed spectrum and ato the reference spectrum. These cross sections are made by convolving the high-resolution cross sections with the in-flight wavelength calibration results for all across-track positions. Further details of the DOAS analysis setting can be found in Table 3.
- For each analysed spectrum, the direct result of the DOAS fit is <u>the</u> differential slant column density (dSCD)<sub>5</sub> which is the difference of NO<sub>2</sub> integrated concentration <u>difference</u> along the effective light path between the studied spectrum and the selected reference spectrum (SCD<sub>ref</sub>). Reference spectra were acquired over a clean rural area upwind of the urban and factory areas, in the lower left corner of Fig. 4. In the quite homogeneous background area, several spectra were averaged to increase the SNR of the reference spectrum. For the purpose of avoidingTo avoid across-track biases, a reference spectrum is required for each across-track position because of its intrinsic spectral response. Under the assumption that the spatial and
- 450 temporal variability of the stratospheric  $NO_2$  field is negligible during the test flight, this approach eliminates the stratospheric  $NO_2$  contribution, making the results only sensitive to the tropospheric portion. According to\_TROPOMI tropospheric  $NO_2$  product of <u>the</u> reference area on the same day, the residual  $NO_2$  amount in the background spectra is estimated to be 3 × 10<sup>15</sup> molec cm<sup>-2</sup>. <u>Changes in the stratospheric  $NO_2$  could also propagate to the measured tropospheric</u> <u>columns of UVHIS. Under the assumption of a constant stratosphere in time and space during the flight, the changes in the</u>
- 455 <u>SZA impact the column difference between the measurement and the reference. To correct the change in the stratospheric NO<sub>2</sub> SCD, we apply a geometric approximation of the stratospheric AMF with a stratospheric VCD of  $3.5 \times 10^{15}$  molec cm<sup>-2</sup></u>

from TROPOMI product. The maximum change in the stratospheric SCD with respect to the reference, was  $8 \times 10^{14}$  molec cm<sup>-2</sup>.

A sample NO<sub>2</sub> DOAS fit result and <u>the</u> corresponding residual of UVHIS spectra <u>is-are</u> illustrated in Fig. 6, with a differential slant column density (dSCD)dSCD of  $4.95 \pm 0.34 \times 10^{16}$  molec cm<sup>-2</sup> and a RMS on the residuals of  $4.27 \times 10^{-3}$ .

#### 4.3 Air mass factor calculations

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SCD is the integrated concentration along the effective light path of observation, which is strongly dependent on the viewing geometry and the properties that influence radiative transfer of light through the atmosphere. VCD is the integrated concentration along a single vertical transect from the Earth's surface to the top of the atmosphere, which is independent of the changes in the light path length of the SCD.

$$VCD_{i}^{t} = \frac{dSCD_{i} + dSCD_{i}^{s} + SCD_{ref}^{t}}{AMF_{i}^{t}} = \frac{dSCD_{i} + dSCD_{i}^{s} + VCD_{ref}^{t} \times AMF_{ref}^{t}}{AMF_{i}^{t}}.$$
(1)

As shown in Eq. (1), the dSCD<sub>i</sub> dSCD<sub>i</sub> from the DOAS fit can be converted to tropospheric-VCD<sub>i</sub>\_VCD<sub>i</sub><sup>t</sup> by dividing the AMF<sub>i</sub>\_AMF<sub>i</sub><sup>t</sup> -which accounts for the enhancements in the light path (Solomon et al., 1987). The dSCD<sub>i</sub><sup>s</sup> is the stratospheric SCD difference between the measurement and the reference, the SCD<sub>ref</sub>, the VCD<sub>ref</sub> and the AMF<sub>ref</sub> are the tropospheric 470 SCD, VCD and AMF of the reference. In this study, tropospheric NO<sub>2</sub> AMFs have been computed using the SCIATRAN (Rozanov et al., 2014) radiative transfer model (RTM). The SCIATRAN model numerically calculates AMFs based on a priori information on the parameters that change the effective light path, such as sun and viewing geometry, trace gas and aerosol vertical profiles- and surface reflectance.

#### 4.3.1 Parameters in RTM

- (1) During the flight, the viewing geometry is retrieved from the orientation information of the aircraft. Solar The solar position defined by the solar zenith angle (SZA) and the solar azimuth angle (SAA), as well as the relative azimuth angle (RAA) can be calculated, based on the time information, and the latitude and longitude position of each observation. (2) Since the flight is performed under a clear-sky condition, the effect of cloud presence can be neglected ignored in AMF computation. (3) Surface The surface reflectance used in AMF calculation is the product of the Landsat 8 Operational Land
- Imager (OLI) space-borne instrument (Barsi et al., 2014). Coastal The coastal aerosol band (433 to 450 nm) is selected because its bandwidth is relatively narrow, and this band is basically inside the DOAS fitting window (Vermote et al., 2016).
  (4) Since no accurate trace gas vertical profile is available during the flight, a well-mixed vertical distribution (box profile) of NO<sub>2</sub> in the PBL is assumed. However, accurate PBL height is also unavailable, a so the typical height of 2 km is a reasonable guess in the case of a sunny summer day in the mid-latitude area in China. (5) Aerosol optical DepthThe aerosol optical depth (AOD) information used in AMF calculation is the MODIS AOD product MYD04 at 470 nm on the same day

with resampling for every ground UVHIS pixel (Remer et al., 2005), because neither ground-based aerosol measurement is not performed, nor anyand no AERONET station data near the flight area are is available. The MODIS AOD measurements inside the flight area ranges from 0.14 to 0.36. Like the NO<sub>2</sub> profile, the aerosol extinction box profile is constructed from the PBL height and the AOD. Single A single scattering albedo (SSA) is assumed to be 0.93, and an asymmetry factor is assumed to be 0.68 for aerosol extinction profile, based on previous studies of typical urban/industrial aerosols (Li et al.,

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2018).

The Landsat 8 surface reflectance is retrieved through atmospheric correction, using the Second Simulation of the Satellite Signal in the Solar Spectrum Vectorial (6SV) model (Vermote et al., 1997). Since there is no overpass on the same day <u>existed</u> inside <u>the</u> UVHIS research flight area, we <u>choose-selected the</u> surface reflectance product on 3 May 2018, considering the sunny weather condition and no cloud presence. The spatial resolution of Landsat is <u>about-approximately</u> 30 m, which is slightly larger than <u>that of the</u> UVHIS. A resampling of <u>the</u> Landsat 8 surface reflectance product based on nearest neighbour interpolation is performed for every UVHIS ground pixel.

The radiative transfer equation in SCIATRAN is solved in a pseudo-spherical multiple scattering atmosphere, using the scalar discrete ordinate technique. Simulations are-were performed for the sensor altitude of 3 km above sea level, and the wavelength of the middle of the NO<sub>2</sub> fitting windows, i.e. 450 nm. A NO<sub>2</sub> AMF look-up table (LUT) was computed, with the different RTM parameter settings provided in Table 4. For each retrieved dSCD, an AMF was linear-linearly interpolated from the LUT based on the sun geometry, the viewing geometry, and the surface reflectance.

#### 4.3.2 RTM dependence study

1. AMF dependence on the surface reflectance

- 505 As shown in Fig. 7, a time series of computed AMFs are plotted for the research flight on 23 June 2018, as well as the corresponding surface reflectance, solar zenith angles, and relative azimuth angles. Other RTM parameters used in the AMF calculations are also provided in Fig. 7. Note that only data of nadir observations are plotted for a clearer clear display, and the time gaps between adjacent flight lines can be observed. Despite the greater great degree of varieties in viewing and sun geometries, it is obvious that the AMFs strongly depend on the surface reflectance. Previous studies reported in by Lawrence et al. (2015), Meier et al. (2017) and Tack et al. (2017) suggest a similar conclusion. A sensitivity test was carried out to investigate the integer of provided method.
- investigate the impact of surface reflectance on the AMF calculations, based on <u>the</u>\_SCIATRAN model, with varying values of surface reflectance, and <u>the</u> fixed values of other parameters. <u>Results-The results</u> of this test are shown in Fig. 8 (a), and indicate that the relation between <u>the</u> surface reflectance and <u>the</u> AMF is non-linear. Especially when <u>the</u> surface reflectance is <u>blow-below</u> 0.1, <u>the</u> AMF increase with <u>the</u> surface reflectance rapidly.
- 515 Generally speaking, the AMF should be higher in the case of a bright surface reflectance, because more sunlight is reflected from the ground back to <u>the</u> atmosphere and then recorded by the airborne sensor. Compared to rural areas, urban and

industrial areas usually exhibit an enhancement in value of enhanced surface reflectance and subsequently an<u>a</u> subsequent increment of in AMF. As shown in Fig. 9, the dependency of the AMF on the surface reflectance is very strong. Also Moreover, a strong variability of the surface reflectance and the AMF can be observed in these areas. From Fig. 9, we can also observe several slight inconsistencies between UVHIS measured radiance and Landsat 8 surface reflectance product,

- 520 also observe several slight inconsistencies between UVHIS measured radiance and Landsat 8 surface reflectance product, due to time offset and spatial resolution difference.Fig. 9 also shows several slight inconsistencies between the UVHIS measured radiance and the Landsat 8 surface reflectance product. For example, the east-west main road looks thinner in Fig. 7 (a) compared to Figs. 7 (b) and (c). This could be explained by the relatively higher spatial resolution performance of the UVHIS and the resampling of Landsat 8 pixels.
- 525 2. AMF dependence on profiles

Based on airborne UVHIS retrieval product, the horizontal distribution of NO<sub>2</sub> can be detected, but the vertical distribution of NO<sub>2</sub> in the atmosphere is <u>not availableunavailable</u>. The assumptions we made for <u>the profile</u> shape of <u>the trace</u> gas and aerosol extinction do not consider the effective variability during research flight; which can be expected in an urban area. Focusing on the impact of different profile shapes on the AMF computation, sensitivity tests of two different NO<sub>2</sub> profiles

- 530 which are closer to ground surface are-were performed: well-mixed NO<sub>2</sub> box profiles of 0.5 and 1 km heights. Compared to the box profile of 2 km which is close tonear the estimated height of PBL, the AMFs decreased by an average of 13 % in the case of a \_box profile of 1.0 km, while whilst the AMFs decreased by an average of 22 % in the case of a box profile of 0.5 km.
- Depending on the relative position of <u>the</u> aerosol and trace gas layer, the optical thickness and <u>the</u> scattering properties, aerosols can enhance or reduce the AMF in different ways(Meier et al., 2017). If an aerosol layer is located above the majority of the trace gas, <u>the</u> aerosols with high SSA have a shielding effect as less scatter light passes through the trace gas layer, leading to a shorter light path. On the other hand, if aerosols and the trace gas are present in the same layer, <u>the</u> aerosols can lead to multiple scattering effects which extend the light path and result <u>in</u> a larger AMF. According to the simulations of a well-mixed aerosol box profile of 2 km and a pure Rayleigh atmosphere, AMFs are slightly higher (<del>about by</del> **540** approximately 1 %) compared to than those of the pure Rayleigh scenario.

3. AMF dependence on sun and viewing geometries

As can be seen in Fig. 7, Figure 7 shows that the effect of sun and viewing geometries on AMFs is very small. Based on a previous study from by Tack et al. (2017), the changing SZA have the greatest effect on the AMFs, in comparisoncompared to other sun and viewing geometries. In this study, we also did-performed an AMF dependence analysis on SZAs and VZAs.

545 The SZA <u>varies\_varied</u> from 12.8° to 37.4° during the 3 <u>hour\_h</u> research flight, <u>while whilst</u> the VZA <u>ranges ranged</u> from 0° to 30° in most cases. As shown in Figs. 8 (b) and (c), the changes <u>of in</u> AMF <u>are-were</u> less than 10% and 7% respectively, when other parameters <u>are-were</u> set <u>as to the mean values</u>. Generally, a larger SZA or a larger VZA could result <u>in</u> a longer light path through the atmosphere and thus a larger AMF.

#### 4. AMF dependence on the analysis wavelength

550 The dependence of AMF on the analysis wavelength is shown in Fig.  $9_{-}$ . The AMF increases with the analysis wavelength. This could be explained by the Rayleigh scattering characteristics. That is, photons at shorter wavelengths are more likely to be scattered than photons at longer wavelengths, leading to the reduced sensitivity to AMF at shorter wavelengths. In the DOAS analysis wavelength window of 430-470 nm, the increase in AMF is about approximately 2 %.

## 4.4 Geo-referencingResampling and mapping

555 Accurate geo referencing is essential for emission source locating and data comparison, and can be achieved with sensor position and orientation information recorded by navigation system and IMU on board. After geo referencing, the The georeferenced NO<sub>2</sub> VCDs are gridded to combine overlapped adjacent measurements, with a spatial resolution of  $0.0003^{\circ} \times$  $0.0002^{\circ}$ . Corresponding to  $27 \times 22$  m<sup>2</sup>, the grid size used is slightly larger than the effective spatial resolution of the UVHIS for the purpose of reducing to reduce the number of empty grid cells. All VCDs are assigned to a grid cell based on its centre 560 coordinates, and several VCDs in one grid cell are unweighted averaged. As shown in Fig. 10, the The final NO<sub>2</sub> VCD distribution map is plotted over the satellite Maps-maps layers in QGIS 3.8 software (QGIS development team, 2020).

#### **5** Results

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The tropospheric NO<sub>2</sub> VCD two-dimensional distribution map is shown in Fig. 10 for the research flight on 23 June 2018. With a-the high performance of UVHIS in spectral and spatial resolution, Figure 10Fig. 10 shows fine-scale NO<sub>2</sub> spatial variability to resolve individual emission sources. In general, the NO<sub>2</sub> distribution is dominated by several exhaust plumes with enhanced NO<sub>2</sub> concentration in the northwest part, which that share a transportation pattern from south to north that is consistent with the wind direction. These sources include a power plant, a steel factory, two cement factories, and several carbon factories. The largest plume with peak values of up to  $3 \times 10^{16}$  molec cm<sup>-2</sup>, originates originated from an emitter inside a steel factory (number 3 in Fig. 10). This dominant plume reaches its peak value outside at a small valley about 570 approximately 1 km north of the factory, and is was transporting at least 9 km and seems to be continuing outside the flight region. This enhanced level of NO<sub>2</sub> may be caused by the terrain factor which contributes to the accumulation of pollution gases.

Numbers 4 to 6 represent other emitters inside the steel factory. While, whilst the exhaust plumes originated from numbers 4 and 5 merged with the dominant plume, the plume from number 6 transports transported to north individually with a peak

575 value of  $1.4 \times 10^{16}$  molec cm<sup>-2</sup>. A weaker-plume with peak values of  $1.5 \times 10^{16}$  molec cm<sup>-2</sup> is-was also detected by UVHIS, which seems seemed to originate from the power plant. Indicated by number 2 in Fig. 10, this power plant is less than 2 km south of the steel factory. Number 1 in Fig. 10 indicates several carbon factories, which are located on the left side of the flight area. Several plumes with peak values of  $\frac{1.51.6}{1.51.6} \times 10^{16}$  molec cm<sup>-2</sup>, gradually merged together during transportation downwind. Numbers 7 and Number-8 in Fig. 10 represent two different cement factories. Peak-The peak values of these two plumes are  $1.5 \times 10^{16}$  molec cm<sup>-2</sup> and  $1.4 \times 10^{16}$  molec cm<sup>-2</sup> respectively.

Compared to the industrial areas mentioned above, the pollution levels of the rural areas are much lower due to the lack of contributing sources, ranging from 2 to  $6 \times 10^{15}$  molec cm<sup>-2</sup>. The urban area of Feicheng <u>eity-City</u> is located on the right side of the flight area. Figure 11 is an enlarged map of <u>the UVHIS NO<sub>2</sub></u> observations over Feicheng <u>eityCity</u>, with a colour scale <u>that</u> only extends to  $7 \times 10^{15}$  molec cm<sup>-2</sup>. Two-The two black lines in Fig. 11 represent the truck roads in this city. The S104

585 is a provincial highway that crosses Feicheng from north to south, while the whilst S330 crosses Feicheng from east to west. Although lots of noise can be observed in Fig. 11, the NO<sub>2</sub> sources in Feicheng are mainly related to traffic and concentrated along the S104.

Due to temporal discontinuity of <u>the</u> flight lines and <u>the</u> dynamic characteristics of <u>the</u> tropospheric NO<sub>2</sub> field, artefacts can be observed between adjacent flight lines. Figure 12 shows three flight lines that pass through the steel factory, at local times <u>590</u> of 13:26 (a), 13:32 (c), and 14:57 (b). Panel<u>s</u> (a) and (b) represent <u>the</u> flight lines that cover the same area with a 1.5 hour <u>h</u> time gap, <u>and panels</u> (a) and (c) represent adjacent flight lines with a 6 <u>minutes-min</u> time gap. These flight lines can be divided into three regions: region A covers no NO<sub>2</sub> source but is affected by <u>the</u> carbon factories <u>about-approximately</u> 3 km away; region B covers the steel factory as <u>the</u> dominant NO<sub>2</sub> source; region C covers no NO<sub>2</sub> source and is not affected by other sources. <u>Compared to region B</u>, there is a large temporal variety of NO<sub>2</sub>. VCDs in region A between three flight lines.

595 Region C is temporally consistent with relatively low NO<sub>2</sub> columns. From these observations it may be concluded that largest temporal variability could occur where there is no local NO<sub>2</sub> source but is down wind of other sources, especially when wind direction is changing. In these three regions, only region C is temporally consistent with relatively low NO<sub>2</sub> columns, whilst a large temporal variety of NO<sub>2</sub> VCDs exists in region A and region B because of inconstant emission sources and changing meteorology.

#### 600 6 NO<sub>2</sub> VCD assessment

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### 6.1 Uncertainty analysis

The total uncertainty on the retrieved tropospheric NO<sub>2</sub> VCDs is composed of three parts: (1) uncertainties in the retrieved dSCDs;-, (2) uncertainties in reference column SCD<sub>ref</sub>; and (3) uncertainties in computed AMFs. Assuming that these uncertainties originating from independent steps are sufficiently uncorrelated, the total uncertainty of <u>the</u> tropospheric NO<sub>2</sub> VCD can be quantified as follows:

$$\sigma_{\text{VCD}_i} = \sqrt{\left(\frac{\sigma_{\text{dSCD}_i}}{\text{AMF}_i}\right)^2 + \left(\frac{\sigma_{\text{SCD}_{ref}}}{\text{AMF}_i}\right)^2 + \left(\frac{\text{SCD}_i}{\text{AMF}_i^2} \times \sigma_{\text{AMF}_i}\right)^2}.$$
(2)

The first uncertainty source,  $\sigma_{dSCDi}$ , originates from <u>the</u> DOAS fit residuals and is a direct output in <u>the</u> QDOAS software. This dSCD uncertainty is dominated by <u>the</u> shot noise from radiance, <u>the</u> electronic noise from the instrument, <u>the</u> systematic uncertainties from the cross sections and <u>the</u> errors from wavelength calibration. In this study, spatial binning of 10 pixels is

610 performed to reduce this-these\_DOAS fit residuals, with a mean slant error of  $4.8 \times 10^{15}$  molec cm<sup>-2</sup>. Compared to mobile DOAS system, this DOAS fit error is approximately two times larger. Unlikely situation of mobile DOAS system, it becomes a major contributor to the total uncertainty in the case of an airborne sensor, especially in a clean area.

The second uncertainty source,  $\sigma_{SCDref}$ , is caused by the NO<sub>2</sub> residual amount in the reference spectra. Since we use <u>the</u> TROPOMI tropospheric NO<sub>2</sub> product of the clean reference area as <u>the</u> background amount, the uncertainty of NO<sub>2</sub> vertical column is estimated to be  $1 \times 10^{15}$  molec cm<sup>-2</sup> directly from TROPOMI product. Assuming a<u>A</u> tropospheric AMF of 2.0 and a tropospheric AMF over the reference spectra of 1.8, <u>this resultsresult in</u> an uncertainty  $9 \times 10^{14}$  molec cm<sup>-2</sup> to the tropospheric vertical column.

The third uncertainty source,  $\sigma_{AMFi}$ , derives from the uncertainties in the parameter assumptions of radiative transfer model inputs. According to previous studies (Boersma et al., 2004; Pope et al., 2015), it-oAMFL is treated as systematic and depends 620 on the surface albedo, the NO<sub>2</sub> profile, the aerosol parameters, and the cloud fraction. (1) Since the research flight took place under cloudless conditions, cloud fraction is neglected in this case. The cloud fraction is neglected in this case because the research flight was under cloudless conditions. (2) Results-The results of the dependence tests in Sect. Section 4.3.2 suggest that the surface albedo has the most significant effect on the AMF. According to Vermote et al. (2016), the uncertainty of the LANDSAT 8 surface reflectance product of band 1 is 0.011. (3) Based on According to the sensitivity study performed in Sect. Section 4.3.2, the uncertainty related to the a priori NO<sub>2</sub> profile shape is lower than 22 %. (4) According to the 625 performed simulations of a pure Rayleigh atmosphere, the uncertainty related to the aerosol state is estimated to be less than 1 %. (5) Because of the high accuracy of the viewing and sun geometry, geometries and its-their low impact on the AMF computation revealed in the previous section, the uncertainty related to the viewing and sun geometry geometries is expected to be <u>ignorable</u>negligible. Therefore, combining all the uncertainty sources in the quadrature, a mean relative uncertainty of 630 24 % on the  $\sigma_{AMFi}$  is obtained.

Based on above discussion, the total uncertainties on the retrieved tropospheric NO<sub>2</sub> VCDs of all <u>the</u> observations of the research flight are calculated. They, typically range between ranging from  $1.5 \text{ and } 1.5 \times 10^{15}$  to  $5.9 \times 10^{15}$  molec cm<sup>-2</sup>, with a mean value of  $3.0 \times 10^{15}$  molec cm<sup>-2</sup>.

#### 6.2 Comparison to mobile DOAS measurements

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635 In order to compare <u>the UVHIS NO<sub>2</sub> VCDs</u> to <u>the ground-based measurements</u>, mobile DOAS observations were performed on 23 June 2018 as well. This mobile DOAS system is composed of a spectrum acquisition unit and a GPS module. The spectrum collection unit consists of a spectrometer, a telescope, an optical <u>fiberfibre</u>, and a workbench. The FOV of this telescope is 0.3°, and its focal length is 69 mm. The spectrometer used is a Maya 2000 Pro spectrometer, with a wavelength range of 290-420 nm and a spectral resolution of 0.55 nm. Zenith sky The zenith-sky observations of the mobile DOAS are

640 <u>were</u> adopted for minimal blocking of buildings and trees in this research. The important properties of the mobile DOAS system and its NO<sub>2</sub> retrieval approach are shown in Table 5. <u>It is worth noting that the retrieval window in the mobile DOAS observations differs from the one used for the airborne observations.</u>

For better comparison with <u>the UVHIS NO<sub>2</sub></u> observations, assumptions and parameters in <u>the tropospheric NO<sub>2</sub></u> retrieval method for the mobile DOAS were <u>similarly</u> set to <u>the same asthose of</u> the UVHIS. For example, <u>the</u> residual amount of  $NO_2$ 

645 in <u>the</u> reference spectra was set to  $3 \times 10^{15}$  molec cm<sup>-2</sup> with an error of  $1 \times 10^{15}$  molec cm<sup>-2</sup>; <u>the</u> mobile DOAS observations only focused on <u>the</u> tropospheric portion of <u>the</u> NO<sub>2</sub> columns, <u>assumed assuming</u> that the difference <u>of in</u> the stratospheric NO<sub>2</sub> columns between <u>the</u> observed <u>spectra</u> and reference spectra is negligible; <u>the</u> vertical profiles of NO<sub>2</sub> and aerosol extinction, albedo, and aerosol properties in the AMF calculation were <u>similarly</u> set to <u>the same asthose of the</u> UVHIS.

Like the uncertainty analysis of <u>the</u> UVHIS NO<sub>2</sub> columns, the total uncertainty on the retrieved mobile tropospheric VCD is composed of three parts: (1) the mean uncertainty on <u>the</u> dSCD of <u>the</u> mobile DOAS is  $1.4 \times 10^{15}$  molec cm<sup>-2</sup>; (2) the uncertainty of reference vertical column is estimated to be  $1 \times 10^{15}$  molec cm<sup>-2</sup>. In the case that the tropospheric AMFs of <u>the</u> measured and reference spectra are very close, this part results <u>in</u> an uncertainty <u>of</u>  $1 \times 10^{15}$  molec cm<sup>-2</sup> to the total uncertainty; (3) the mean relative uncertainty on the AMF calculation is 22 % by <u>the</u> square root of the quadratic sum of <u>the</u> individual uncertainties like UVHIS. Combining these uncertainties together, the mean total uncertainties on the retrieved

655 tropospheric NO<sub>2</sub> VCD is  $2.1 \times 10^{15}$  molec cm<sup>-2</sup>.

Basically, the route of the mobile DOAS was designed to encircle the power plant and <u>the</u> steel factory which are supposed to be predominant sources. For the comparison, <u>the</u> mobile DOAS observations are first gridded to the same sampling of <u>the</u> UVHIS pixels,...<u>then-Thern</u> the VCD of <u>the</u> UVHIS NO<sub>2</sub> results is extracted for each co-located mobile measurement. An overview of the mobile DOAS measurements over <u>the</u> UVHIS NO<sub>2</sub> layer is shown in Fig. 13. <u>The</u> NO<sub>2</sub> distributions of <u>the</u> 660 mobile DOAS system and <u>the</u> UVHIS exhibit similar spatial characteristics, <u>which-i.e.</u> low values are <u>located</u> in the south of

the steel factory and power plant, and high values are inside the plumes.

Figure 14 (a) shows scatter plots with <u>the VCDs</u> retrieved by <u>the UVHIS</u> on the x-axis and <u>the mobile DOAS VCDs</u> on the y-axis, for all co-located measurements. The corresponding results of <u>the linear</u> regression analysis are also provided in Fig.14 (a), with a correlation coefficient of 0.69, a slope of 1.30, and an intercept of -9.01 × 10<sup>14</sup>. The absolute time offset between <u>the mobile DOAS</u> and airborne observations can be up to 1 hourh, which means indicating that both instruments cannot sample the NO<sub>2</sub> column at certain geolocations simultaneously. As shown in Fig. 14 (b), when only comparing UVHIS VCDs to mobile measurements that circled the steel factory, the correlation coefficient <u>improves-improved</u> to 0.86. In this case, all mobile measurements occurred inside the swath of one flight line of aircraft, and the time offset between <u>two</u> instruments shortened to 15 minutesmin. In general, an underestimation of <u>the UVHIS VCDs</u> of increased value can be
670 observed in Figs 14 (a) and (b). Considering the variability in local emissions and meteorology, it is reasonable that the

differences between these two instruments exist. Besides, the averaging effect of the area inside an UVHIS pixel can also lead to the underestimation of UVHIS compared to mobile DOAS system.<u>A</u> sensitivity test of the AMF on the NO<sub>2</sub> profile was performed for all co-located measurements, using a box profile of 500 m. Compared to the box profile of 2 km, the UVHIS AMFs decreased by an average of 17 %, whilst the mobile DOAS AMFS decreased by an average of 2.7 %. This results suggest that a more realistic profile with the NO<sub>2</sub> layer closer to the ground could improve the slope closer to unity.

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## 7 Conclusions

In this paper, we present the newly developed <u>Ultraviolet Visible Hyperspectral Imaging Spectrometer (UVHIS)UVHIS</u> <u>instrument</u>, with a broad spectral region <u>ranging</u> from 200 to 500 nm, and a high spectral resolution better than 0.5 nm. The instrument is operated in three channels at wavelength 200 to 276 nm (channel 1), 276 to 380 nm (channel 2), and 380 to 500

- 680 nm (channel 3) for minimal stray light effects and <u>the highest spectral performance</u>. The optical design of each channel consists of a fore-optics with a FOV of 40°, an Offner imaging spectrometer, and a CCD array detector of  $1032 \times 1072$  pixels.
  - <u>Also weWe also</u> present the first tropospheric NO<sub>2</sub> retrieval results from the UVHIS airborne observation in June 2018. The research flight above over Feicheng, China, covered an area of <u>approximatelyabout 20-30 × 30-20</u> km<sup>2</sup> within 3 <u>hourh</u>, with a high spatial resolution about approximately 25 × 22 m<sup>2</sup>. We first retrieved the differential NO<sub>2</sub> slant column densities from
- 685 high spatial resolution about approximately  $25 \times 22$  m<sup>2</sup>. We first retrieved the differential NO<sub>2</sub> slant column densities from nadir observed spectra by application of applying the DOAS technique, to a mean reference spectra over a clean area. Then we converted those NO<sub>2</sub> slant columns to tropospheric vertical columns using the air mass factors derived from the SCIATRAN model with the Landsat 8 surface reflectance product. Total-The total uncertainties of the tropospheric NO<sub>2</sub> vertical columns are in range of range from  $1.5 \times 10^{15}$ -to  $5.9 \times 10^{15}$  molec cm<sup>-2</sup>, with a mean value of  $3.0 \times 10^{15}$  molec cm<sup>-2</sup>.
- 690 The two-dimensional distribution map of <u>the</u> tropospheric NO<sub>2</sub> VCD demonstrates that <u>the</u> UVHIS is adequate for trace gas pollution monitoring over a large area in a relative<u>ly</u> short time frame. With the high spatial resolution of <u>the</u> UVHIS, different local emission sources can be distinguished, fine-scale horizontal variability can be revealed, and trace gas emission and transmission can be understood. For the flight on 23 June 2018, <u>the</u> NO<sub>2</sub> distribution <u>is-was</u> dominated by several exhaust plumes which exhibit<u>ed the</u> same south to north direction of transmission, with a peak value of  $3 \times 10^{16}$  molec cm<sup>-2</sup>
- 695 in the dominant plume. <u>Comparisons-The comparisons of the UVHIS NO<sub>2</sub> vertical columns to-with the mobile DOAS observations show a good overall agreement-overall</u>, with a correlation coefficient of 0.65 for all <u>the co-located measurements</u>, and a correlation coefficient of 0.86 for <u>the co-located measurements</u> that only circled the steel factory. However, an underestimation of the high NO<sub>2</sub> columns of <u>the UVHIS</u> is observed relative to the mobile DOAS measurements.
- 700 High resolution The high-resolution information about the NO<sub>2</sub> horizontal distribution, generated from UVHIS airborne data, is unique and valuable compared to those from ground-based instruments and space-borne sensors. In future study, the

UVHIS could be applied in <u>the validation</u> of satellite trace gas instruments, and in <u>the connection</u> between local point observations, air quality models, and global monitoring from space.

705 Data availability. The datasets in the present work are available from the corresponding author upon reasonable request.

*Author Contributions*. ConceptualizationConceptualisation, F.S.; methodology, Y.J. and H.Z.; software, Z.C.; validation, X.Q. and D.Y.; formal analysis, L.X.; resources, K.Z.; writing—original draft preparation, L.X.; writing—review and editing, F.S.

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715 (ESA). TROPOMI data can be downloaded from https://s5phub.copernicus.eu. Landsat 8 OLI data have been produced, archived, and distributed by the U.S. Geological Survey (USGS). The original Landsat surface reflectance algorithm was developed by Dr. Eric Vermote, NASA Goddard Space Flight Center (GSFC). Landsat 8 OLI data are available at https://earthexplorer.usgs.gov.

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Characteristic	Channel 1	Channel 2	Channel 3
Wavelength range	200 <u>–</u> -276 nm	276 <u>–</u> -380 nm	380 <u>-</u> -500 nm
Spectral sampling	0.074 nm	0.10 nm	0.12 nm
Spectral resolution	0.34 nm	0.46 nm	0.49 nm
FOV	40°	40°	40°
Focal length	18 mm	18 mm	18 mm
Across-track angular resolution	0.5 mrad	0.5 mrad	0.5 mrad
f-number	3.4	3.6	3.6
Detector size	$1032 \times 1072$	$1032 \times 1072$	$1032 \times 1072$

Table 1: UVHIS instrument characteristics of three channels.

 Table 2. Preflight wavelength calibration results (FWHMs) of UVHIS channel 3 for 9 viewing angles. Light sources used in the

 calibration are a mercury-argon lamp and a tunabletuneable laser. Slit function shapes are retrieved by least square fitting of characteristic

 spectral lines, using a symmetric Gaussian function.

FOV	379.887 nm	404.656 nm	450.504 nm	500.566 nm
-20°	0.35 nm	0.35 nm	0.39 nm	0.50 nm
-15°	0.33 nm	0.31 nm	0.33 nm	0.43 nm
-10°	0.31 nm	0.29 nm	0.29 nm	0.41 nm
-5°	0.31 nm	0.30 nm	0.29 nm	0.34 nm
0°	0.31 nm	0.32 nm	0.30 nm	0.30 nm
5°	0.34 nm	0.36 nm	0.34 nm	0.30 nm
10°	0.38 nm	0.39 nm	0.38 nm	0.32 nm
15°	0.40 nm	0.44 nm	0.42 nm	0.35 nm
20°	0.45 nm	0.46 nm	0.47 nm	0.38 nm

Parameter	Settings	
Wavelength calibration	Solar atlas, (Chance and Kurucz, 2010)	
Fitting interval	430 <u>-</u> -470 nm	
Cross sections		
NO <sub>2</sub>	298 K, Vandaele et al. (1998)	
O3	223 K, Serdyuchenko et al. (2014)	
O4	293 K, Thalman and Volkamer (2013)	
$H_2O$	293 K, Rothman et al. (2013)	
Ring effect	Chance and Spurr (1997)	
Polynomial term	Order 5	
Intensity offsetOffset	Order 1	

Table 3. Main analysis parameters and absorption cross sections for NO<sub>2</sub> DOAS retrieval.

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Table 4. Overview of the input parameters in the SCIATRAN RTM, characterizing the AMF LUT.

<b>RTM Parameter</b>	Grid settings	
Wavelength	450 nm	
Sensor altitude	3 km	
Surface reflectance	0.01 <u>-</u> 0.4 (steps of 0.01)	
Solar zenith angle	10 <u>-</u> -40° (steps of 10°)	
Viewing zenith angle	0 <u>-</u> -40° (steps of 10°)	
Relative azimuth angle	0 <u>-</u> -180° (steps of 30°)	
Aerosol optical depth	0 <u>-</u> -1 (steps of 0.1)	
Aerosol extinction profile	Box of 2.0 km	
NO <sub>2</sub> profile	Box of 2.0 km	

Parameter	Settings	
Elevation angle	zenith	
Fitting interval	356 <u>–</u> -376 nm	
Wavelength calibration	Mercury lamp	
Cross sections		
$NO_2$	298 K, Vandaele et al. (1998)	
$O_3$	223 K, Serdyuchenko et al. (2014)	
$O_4$	293 K, Thalman and Volkamer (2013)	
Ring effect	Chance and Spurr (1997)	
Polynomial term	Order 5	
Intensity offsetOffset	Order 1	

Table 5. Properties of the mobile DOAS system and its NO<sub>2</sub> fit.





Figure 1. Optical layout of the UVHIS channel 3. Optical design of channel 1 and channel 2 is similar.



Figure 2. Measured slit functions (dots) at 450.504 nm and retrieved slit function shapes (lines) using a symmetric Gaussian function for 9 viewing angles.



Figure 3. TROPOMI observation of tropospheric NO<sub>2</sub> over China on 23 June, 2018. The location of UVHIS flight (Feicheng city) is also plotted in the map.



Figure 4. Overview of the Feicheng demonstration flight on 23 June, 2018. Flight lines are shown in blue. Two orange circles represent the routes of mobile DOAS system. White dots numbered from 1 to 8 represent the major emission sources. Number 1: several carbon factories; number 2: a power plant; numbers 3-6: individual emitters inside the steel factories, while numbers 4 and 5 are inside the circle of one mobile DOAS route; numbers 7-8: two cement factories. White dashed box represents the reference area.



Figure 5. In-flight spectral calibration: (a) the spectral resolution (FWHM); (b) the spectral shift on different across-track position. Results at three wavelengths are plotted: blue for 430 nm, green for 450 nm and red for 470 nm.



Figure 6. Sample DOAS fit result for NO<sub>2</sub>: (a) observed (black dashed line) and fitted (blue line) optical depths from measured spectra; (b) the remaining residuals of DOAS fit.



Figure 7. Time series of NO<sub>2</sub> AMF compared with (a) surface reflectance; (b) SZA and RAA for the research flight on 23 June 2018, computed with SCIATRAN model based on the RTM parameters from the UVHIS instrument. Only data of the nadir observations in each flight line are plotted.



Figure 8. AMF dependence analysis results (a): on the surface reflectance; (b): on the SZAs; (c): on the VZAs; (d): on the wavelength. 



Figure 9. (a) UVHIS <u>Measured measured</u> radiance; (b) Landsat 8 Surface reflectance; (c) computed AMFs, for one flight line of the Feicheng data set. A strong dependency of the AMF on the surface reflectance can be observed.



6010 Figure 10. Tropospheric NO<sub>2</sub> VCD map retrieved from UVHIS over Feicheng on 23 June 2018. The major contributing  $\frac{NO_2 NO_x}{NO_x}$  emission sources are indicated by numbers 1 to 8.



Figure 11. Enlargement of UVHIS NO<sub>2</sub> VCD map over Feicheng city with a <u>colorcolour</u> scale only extends to  $7 \times 10^{15}$  molec cm<sup>-2</sup>. Two black lines in the map represent two truck roads that cross Feicheng city: S104, and S330.



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Figure 12. Three flight lines that pass through the steel factory, at local time 13:26 (a), 13:32 (c), and 14:57 (b). Panel (a) and (b) represent flight lines that cover the same area with a 1.5 hour h time gap, panel (a) and (c) represent adjacent flight lines with a 6 minutes min time



Figure 13. Overview of VCDs retrieved from ground-based mobile DOAS system (circle marks), and VCDs retrieved by UVHIS (background layer), measured on 23 June 2018.



Figure 14. Scatter plot and linear regression analysis of the co-located NO<sub>2</sub> VCDs, retrieved from UVHIS and mobile DOAS system, (a) for all co-located measurements, with a time offset of 1 h, (b) for co-located measurements that only circled the steel factory, with a time offset of 15 min.