In-situ cloud ground based measurements in Finnish sub-Arctic: Intercomparison of three cloud spectrometers setups.

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Abstract. Continuous, semi-long term, ground based, in-situ cloud measurements were conducted during the Pallas Cloud Experiment (PaCE) in 2013. The measurements were carried out in Finnish sub-Arctic region at Sammaltunturi station (67°58′N, 24°07′E, and 560 m a.s.l.), the part of Pallas Atmosphere - Ecosystem Supersite and Global Atmosphere Watch (GAW) programme. The main motivation of the campaign was to conduct in-situ cloud measurements with three different cloud spectrometer probes and perform an evaluation of their ground based setups. Therefore, we mutually compared the performance of the Cloud and Aerosol Spectrometer (CAS), the Cloud Droplet Probe (CDP) and the Forward Scattering Spectrometer Probe (FSSP-100), (DMT, Boulder, CO, USA). We investigated how different meteorological parameters affect each instrument instruments' ground based setup operation and quantified possible biases and discrepancies of different microphysical cloud properties. Based on obtained results we suggested limitations for further use of the instruments setups in campaigns where focus is on investigating aerosol cloud interactions. Measurements in this study were made by Finnish Meteorological Institute owned instruments and results concern their operation in sub-Arctic conditions with frequently occurring super-cooled clouds.

Measured parameter from each instrument was the size distribution and additionally we derived the number concentration, the effective diameter, the median volume diameter and the liquid water content. A complete intercomparison between the CAS probe and the FSSP-100 ground setups and additionally between the FSSP-100 and the CDP probe ground setups was made and presented. Unfortunately, there was not sufficient amount of common data to compare all three probes together due to operational problems of the CDP ground setup in sub-zero conditions. The CAS probe that was fixed to one direction lost significant number of cloud droplets when the wind direction was out of wind iso axial conditions in comparison with the FSSP-100 and the CDP which were both placed on a rotating platform. We revealed that CAS and FSSP-100 had good agreement in deriving sizing parameters (effective diameter and median volume diameter from 5 to 35 µm) even though CAS was losing a significant amount of cloud droplets. The most sensitive derived parameter was liquid water content which was strongly connected to the wind direction and temperature.

1 Introduction

Clouds and their interaction with aerosol particles provide some of the greatest uncertainties in predictions of climate change (Boucher et al., 2013). Therefore, in situ measurements of clouds play a key factor for further investigation of the aerosol cloud interaction area. Many of the climatically important cloud properties (e.g. albedo, precipitation rate and lifetime) depend, among other factors, on the number concentration of aerosol particles and on their chemical composition (Komppula et al., 2005; Lihavainen et al., 2008). Measuring the cloud size distribution is an important parameter for identification and description of clouds; their microphysical properties (Pruppacher and Klett, 1977; Rosenfeld and Ulbrich, 2003), and their lifetime (Albrecht 1989; Small et al., 2009).

One major category of instruments that it is commonly used for in situ cloud measurements are known as cloud spectrometers (e.g. Knollenberg, 1976; Dye and Baumgardner, 1984; Wendish, 1996; Baumgardner et al., 2001; Lance et al., 2010; Baumgardner et al., 2014). Cloud spectrometers are single particle counters that use the forward scattering, usually with the angles between 4 and 12° of a laser beam to detect and classify in different size bins individual particles. The main theory used for the particle sizing based on the scattering of light from single particles is the Lorenz - Mie theory (Mie, 1908). Several experiments were conducted with those instruments; they mainly cover: laboratory (e.g. Wagner et al., 2006; Smith et al., 2015; Nichmann et al., 2017), ground based (e.g. Mertes et al., 2001; Bukarel et al., 2002; Henning et al., 2002; Eugster et al., 2006; Lihavainen et al., 2008; Guyot et al., 2015; Loyd et al., 2015; Lowenthal et al. 2019) and airborne measurements (e.g. Knollenberg et al.,1981; Heymsfeld et al., 2004; Bromwich et al., 2012; Johnson et al., 2012, Jones et al., 2012, Briswick et al., 2014; Korolev et al., 2014; Petäjä et al., 2016; Webke et al., 2016; Vogt et al., 2017; Faber et al., 2018)

In addition to above mentioned experiments, many studies were done to quantify biases, uncertainties and limitations of cloud spectrometers while they were used in measurement campaigns. Uncertainties were usually a result of different meteorological conditions. Baumgardner (1983) presented a comparison of five water droplet instruments, included the axially scattering spectrometer (ASSP) and the forward scattering spectrometer probe (FSSP). He concluded that scattering probes had an accuracy of 17% in number concentration and size measurements. A full description and evaluation of the FSSP was made by Baumgardner et al., (1985) investigating coincidence and dead-time losses and by Baumgardner et al., (1990) investigating time response and laser inhomogeneity limitations. Baumgardner et al., (1989) made a calibration of the FSSP during the airborne Antarctic zone experiment and set further limitations to be applied during the data analysis of this project. A similar study from Baumgardner et al., (1992) was conducted for the FSSP during the airborne arctic stratospheric expedition where an improved forward scattering spectrometer probe, the FSSP-300, was developed and introduced. Wendisch (1998) presented a quantitative comparison of ground based FSSP with a particle volume monitor. He stated that FSSP can be considered as an excellent microphysical sensor in continental, stratiform or cumuliform clouds with mostly small drops, however he noticed some discrepancies in the liquid water content, especially when cloud droplets larger than 25 μm were considered. Gerber et al., (1999) performed and evaluated ground based measurements of liquid water content using also a FSSP and a particle volume monitor (PVM). They observed large discrepancies too and stated that the FSSP overestimate

liquid water content for large cloud droplets due to the inertial concentration effect. Burnet and Brenquier (1999) validated the droplet spectra and the liquid water content using five instruments including the FSSP, the fast FSSP and the CDP. Burnet and Brenquier (2002) investigated further in detail only the FSSP to address the instrumental uncertainties and limitations of them. Coelho et al., (2005) made a detailed discussion for FSSP-100 in low and high droplet concentration measurements with a proposed correction for coincidence effects. Lance et al., (2010) calibrated the CDP and presented a full description of the instrument performance in laboratory and in-flight conditions. Baumgardner et al., (2011) summarized airborne systems for in situ measurements of aerosol particles, clouds and radiation that were currently in use on research aircraft around the world including cloud spectrometer probes. Febvre et al., (2012) highlighted the possible effects of ice crystals in FSSP measurements. Spiegel et al., (2012) made a thoroughly analysis of wind velocity and wind angle impacts at the Junfraujoch comparing the Fog droplet spectrometer (FM-100) to others instruments. FM-100 showed several artifacts at temperatures below zero. One more evaluation regarding cloud ground based measurements which taking into consideration the wind direction was made by Guyot et al., (2015) at the Puy-de-Dôme observatory between seven optical sensors including a FSSP, a fast FSSP, a fog monitor and two CDP probes. They placed one FSSP and the fog monitor at the roof of the observatory and the two CDP probes and one FSSP inside a wind tunnel. Authors showed a good agreement in sizing abilities of the instruments but observed possible discrepancies in number concentration values, fact that also affected the liquid water content values. Lloyed et al. (2015) observed cloud microphysical structures by conducting CAPS, FSSP, CDP-100 and PVM measurements. They mounted all the instruments on a rotator and wing on the terrace rooftop outside the Sphinx Laboratory, (Jungfraujoch, Switzerland). Several developments of the in-situ measurement systems were reviewed and summarized with respect to their strengths, weaknesses, limitations and uncertainties by Baumgardner et al., (2017). The progress in performing in-situ cloud measurements was clearly observed along with developments. Lowenthal et.al., (2019) conducted winter time mixed- phase orographic cloud measurements at the Storm Peak Laboratory (Colorado, USA). They deployed a FSSP-100 on a rotating wind vane (to orient them into the wind).

In this work, we focused on the intercomparison of three cloud spectrometer probes ground setups as they were used during the PaCE 2013. Due to the increased demand for long term continuous ground based in-situ cloud measurements, we tested and evaluated the operation of three FMI owned ground setups to perform continuous ground based measurements in harsh environments. The FMI research station (Sammaltunturi) located in northern Finland is considered as an ideal place to perform in-situ low level cloud measurements, especially during autumn, when the station is usually inside a cloud about 50% of the time. There, along with the FSSP – 100 and the CDP which are classic instruments for in-situ cloud measurements, we were also using the Cloud, Aerosol and Precipitation Spectrometer (CAPS) probe (part of this instrument is the CAS probe) with a purchased inhalation system. CAPS' worldwide unique setup allows us semi-long term (in orders of months) observations when compared to short-term (orders of hours) airborne measurement. The exact set of measurements limitations for each cloud probe that are presented in this work will help us to conduct a detailed further analysis of microphysical cloud properties and their interactions with aerosol during all PaCE campaigns. The previous PaCE campaigns, already resulted in

series of publications on experimental observations and modelling studies (e.g. Komppula et al., 2005 and 2006; Lihavainen et al., 2008 and 2010; Kivekäs et al., 2009; Anttila et al., 2009 and 2012).

A description of the measurement site and the instrumentation as it was installed could be found in section 2.1 and 2.2. Subsequently, in section 2.3, the procedure we followed during data analysis is presented. In section 3, the inter-comparison of the instruments and how they were influenced by the meteorological parameters are presented. Finally, in section 4, we summarized our main results and conclusions in order to set limitations and made recommendations for the data analysis of the three instrument ground based setups during future studies in sub-Arctic environment.

2. Methods

2.1 Measurement site description

Measurements were conducted during autumn 2013, in the Finnish sub-Arctic region at Sammaltunturi station (67°58′N, 24°07′E, 560 m a.s.l.) which is a part of the Pallas Atmosphere – Ecosystem Supersite hosted by the Finnish Meteorological Institute. The station is also part of Global Atmosphere Watch (GAW) programme. Sammaltunturi station is located on the top of a treeless hill. A full detailed description of the site can be found in Hatakka et al., (2003). Autumn was chosen as the best period to run the campaign due to the high chances that the station will be inside a cloud. This allows us to conduct in-situ cloud measurements. All the meteorological parameters were continuously measured by the Vaisala FD12P weather sensor. During our previous Pallas Cloud Experiments (PaCE) clouds microphysical properties and aerosols physico – chemical properties and their interactions were measured and investigated (e.g. Lihavainen et al., 2008; Hyvärinen et al., 2011; Anttila et al., 2012; Collaud Cohen et al., 2013; Jaatinen et al., 2014; Lihavainen et al., 2015; Raatikainen et al., 2015; Gérard et al., 2019). During PaCE2013, our main motivation was to focus on inter-comparison of in-situ cloud properties measured with three different cloud probes, their evaluation and mutual benchmarking regarding PaCE campaigns.

2.2 Cloud instrumentation

The atmospheric in-situ measurements community has identified cloud droplet probes with surface installation as a potential method for continuous cloud in-situ measurements (Wandinger et al., 2018). During PaCE 2013, to perform in-situ measurements of cloud droplets, we used three instruments originally developed for airborne measurements, but tailored for ground-based measurements by manufacturer (DMT, USA). The Cloud, Aerosol and Precipitation Spectrometer probe (CAPS), the Cloud Droplet Probe (CDP) and the Forward Scattering Spectrometer Probe (FSSP-100, hereafter called as FSSP for simplicity). All three of them were installed on the roof top of the measurement site as it is described below in details and share similar measurement technique. Wind tunnel could be considered as the optimal choice to utilize these instruments for ground based setups (e.g Elk Mountain, Baumgardner (1983) and Puy de Dôme, Guyot et al., 2015). There are measurement sites like ours in sub-arctic which do not have this possibility due to both practical and budgetary reasons. However, it was shown that same quality data could be obtained from roof top measurements (Guyot et al., 2015). Ground based measurements

with cloud probes that were originally designed to be used for aircraft were already conducted in several measuring sites without using a wind tunnel (e.g. Jungfraujoch, Lloyed et al., (2015) and Storm Peak, Lowenthal et.al., (2019)). The basic concept is that they use the forward scattering of a laser beam for the detection and sizing of individual particles. Then, using Mie theory (Mie, 1908), they calculate the size of the particle from the intensity of the scattered light.

5   Only data of the Cloud and Aerosol Spectrometer (CAS) probe were used regarding to the CAPS. CAS is one part of the Cloud, Aerosol and Precipitation Spectrometer probe (CAPS, DMT) (0.51-930 µm) which is an instrument that is widely used on airborne measurements for investigating the microphysical properties of clouds (e.g. Baumgardner 2001; Baumgardner et al., 2011). The CAPS probe includes two more instruments, however they are not comparable with the FSSP and the CDP: the Cloud Imaging Probe (CIP) and the Hotwire Liquid Water Content Sensor (Hotwire LWC). CAS measures smaller particles

10   (0.51 µm to 50 µm) and relies on light-scattering. Particles scatter light from an incident laser at a wavelength of 680 nm, a sample area of 0.24 mm^2 and collecting optics guide the light scattered in the 4° to 12° range into a forward-sizing photodetector. The intensity of light is measured and used to infer the particle size. Backscatter optics also measure light intensity in the 168° to 176° range, which allows determination of the real component of a particle's refractive index for spherical particles. The droplets are then classified into 30 size bins. CAS was operational from October 15th until November

15   28th. It was installed and fixed on Sammaltunturi station roof. It was heading to the main wind direction of the station (southwest, ~ 225°). For the instruments' installation we used the manufacturer pylon (height 0.3 m) (same as it is used for CAPS airborne measurements). The whole system was fixed on horizontal metallic circle (D = 0.28 m). The metallic circle was attached to a vertical metallic bar (height 0.3 m) part of a square metallic stand (0.7 m x 0.7 m) (see Fig1a). As a result CAPS had a total height of 0.6 m from the point of the roof it was installed and 4.5 m from the ground. In our setup, CAPS

20   had its own tailored inhalation system, a high flow pump (Baldor – Reliance, USA), which worked continuously. The pump was connected with the CAS probe with 1.14 m long of 0.07 m inner diameter hose. The hose was connected to a triple branch (three parts with I.D. = 0.12) through a 0.12 m to 0.05 m reducer. The triple branch connected the CAS probe through the hose with the high flow pump. The other parts of the branch connected the pump with the CIP through 2 different hoses (L=1.52 m I.D. = 0.12 m). In addition, a stepped CAS inlet (funnel shape reducer I.D. = 0.12 to I.D. = 0.05 m) was attached over the CAS

25   inlet tube (for detailed description please Fig. S1 and Fig. S2 in the Supplement).The probe air speed (PAS) inside CAS was checked daily with a digital thermo anemometer (model 471, Dwyer Inc.) to secure that the flows inside the instrument remained similar. This was done through a small hole near the end of CAS probe outlet and beginning of 0.05 m hose and in a position such that the anemometer inserted into hole was just in the middle of the CAS probe outlet (hose diameter). In cases when PAS changed, data were corrected accordingly. During this campaign PAS values ranged from 17 to 23 ms^{-1}. The

30   calibration of the instrument was done at DMT and also at the Finnish Meteorological Institute before and after the campaign using glass beads and polystyrene latex spheres (PSL) standards.

The Forward Scattering Spectrometer Probe (FSSP, 1.2-47 µm, model SPP-100, DMT), initially manufactured by Particle Measuring Systems (PMS Inc., Boulder CO, USA) is a widely used cloud probe for measuring droplet size distribution (Brenquier 1989). It shares the same measurement principle with the CAS probe and measures the light scattered in the 4° to

12° range with a laser of wavelength 633 nm and a sample area of 0.414 mm^2. Droplets were classified into 40 size bins. During PaCE, the FSSP was installed and placed on a rotating platform, so that the inlet was always heading against the wind direction. The rotating platform was a horizontal metallic base (0.7 x 0.1 x 0.4 m) with a metallic fin fixed at the back of it. This setup ensured that the instrument would follow the wind direction continuously. The rotating platform was placed on a similar squared metallic stand we used also in CAPS setup, but with a higher metallic vertical bar (L = 0.6 m, I.D. = 4 cm). Thus, the instrument had a total height of 0.6 m from the point of the roof it was installed and 5.5 m from the ground. During FSSP installation on the rotating platform, we wanted to prevent the full rotation of the probe which could be dangerous for the cable wiring and safety of the instrument. For this reason, a vertically metallic bar (0.3 m, D. = 0.6 cm) along with two horizontal bars (L = 0.25 m, D = 0.6 cm) were installed (northeast ~ 60°) and they served as a brake (Fig. 1 b). A custom inhalation system with high flow ventilator was employed through the instruments' inlet to secure that the air speed would remain constant (for detailed description please Fig. S3 in the Supplement). In addition, the PAS inside the FSSP tube was checked daily with the digital thermo-anemometer (model 471, Dwyer Inc.). The PAS during the campaign was ~ 10 ms^{-1} which lead to an air speed of ~ 36 ms^{-1} inside the inlet due to necking inside the inlet's mouth from I.D. = 3.8 cm to I.D. = 2.0 cm. This value was used for further data processing. FSSP was installed and operational from September 14th until November 28th 2013. The instrument was calibrated at DMT, USA before and on site after the end of the campaign.

The third instrument that we used was the cloud droplet probe (CDP, 3-50 μm, Droplet Measurement Technologies) (Lance et al., 2010). Similar to the previous instruments it uses the same principle and measures the light scattered in the 4° to 12° range. The laser beam had wavelength of 658 nm and sample area of 0.3 mm^2. The CDP classified droplets into 30 size bins. It was placed next to the FSSP also on a rotating platform to continuously head against the wind direction (Fig. 1 b). The metallic platform covering the instrumental electronics consists of fixed part (0.4 x 0.4 x 0.3 m) at the bottom and on top of that the rotating part (0.4 x 0.4 x 0.1 m) having the probe itself on top of it. The rotating part is equipped with a large fin to keep the inlet towards the wind (for detailed description please Fig. S4 in the Supplement). The instrument had a custom inlet with an external pump to secure a constant PAS (14 ms^{-1}). The CDP was installed and was operational from September 25th 2013 until November 28th 2013. It was calibrated at Finnish Meteorological Institute, Kuopio unit, before the campaign and after the campaign on site using glass beads and PSL standards.

All three instruments were using anti ice heaters as they were deployed by the manufacturer. The external parts of the setup (rotating platforms and inhalation systems) were not using additional heating system. The instruments were installed in a horizontal position and placed close to each other on Sammaltunturi roof. The CDP and FSSP were installed next to each other (approx. 0.5 m) and they had a horizontal distance of ~ 10 m and vertical distance of ~1 meter to the CAS probe. All the probes' parameters are presented in Table 1. During the campaign a routine was consistently followed. The cloud probes functionality was daily checked visually. Ice and snow accumulation could fully or partially block the probes' inlets and affect the flows. In addition, ice and snow could also accumulated in parts of the roof and affect the probes measurements. For this reason, all three cloud probes needed periodical cleaning. When the station was inside a cloud and sub-zero temperatures were observed, the cleaning procedure of the probes during the day was repeated twice or more times per day (if needed).

2.3 Data handling and processing

The presence of a cloud was estimated by the cloud droplet counts measured with all the cloud probes, the visibility and relative humidity measurements at the site. As a cloud event we defined the situation when there were droplet counts (considering cleaned dataset) measured by the cloud probes more than 30 continuous minutes, the horizontal visibility was less than 1000 meters and the relative humidity was ~ 100%.

From each cloud probe we obtained the cloud droplet size distribution. For the intercomparison of the probes we had to eliminate some size bins of the CAS and the FSSP probe in order to use similar size range in each case. The CAS probe, using the PADS software (Droplet Measurement Technologies Manual, 2011), derives the number concentration (N_c), the liquid water content (LWC), the median volume diameter (MVD) and the effective diameter (ED). The same parameters were derived using the following equations, since we have used old software PACS 2.2 (Droplet Measurement Technologies) for data acquisition of the FSSP-100 and the CDP probe:

Number concentration, (N_c, cm^{-3}) was calculated from the division of the total number of sized particles N with the sample volume V_s (cm^{-3})

$$N_c = \frac{N}{V_S},\tag{1}$$

where V_S was defined as

$$V_S = \text{PAS} \times t \times \text{A}\tag{2}$$

where PAS is the probe air speed (ms^{-1}), t is the time of the sampling interval and A is sample area (mm^2) defined as the height of the laser beam (mm) multiplied by the length of the laser beam within the depth of field (DOF, mm). On instrument that records probe time as CAS and FSSP, the sampling interval is calculated by subtracting the previous instance's probe time from that of the current instance. On the CDP the sampling interval is assumed to be the designated sample rate. All three probes were setup to sample at 1s (1Hz).

Liquid water content, (LWC, gm^{-3}) is the mass of liquid water per unit volume of air and it was calculated using the following equation

$$LWC = \sum_i^n LWC_i\tag{3}$$

where

$$LWC_i = c_i 10^{-12} \frac{\pi}{6} m_i^3\tag{4}$$

m_i is the midpoint of its bin and calculated as

$$m_i = \frac{b_i + b_{i+1}}{2}\tag{5}$$

and c_i are the droplets concentrations per bin (m^{-3}).

The factor $\frac{\pi}{6} m_i^3$ in equation indicated that we assume that the particle is a sphere. Another assumption that was made is that water has a density of 1 g cm^{-3}.

Median volume diameter (*MVD*, μm) indicates the droplet diameter which divides the total water volume in the droplet spectrum such that half the water volume is in smaller drops and half is in larger drops and is derived by a linear interpolation with respect to the (i+1) bin as

$$MVD = b_{i^*} + \left(\frac{.5 - cum_{i^*-1}}{pro_{i^*}}\right)(b_{i^*+1} - b_{i^*}) ,$$ (6)

where $pro_{i^*} = \frac{LWC}{LWC_i}$ is the proportion of the spectrum *LWC* that falls in the *i-th* bin and

$cum_{i^*} = pro_1 + \cdots + pro_{i^*}$ is the cumulative proportion of the spectrum *LWC* that falls in the first *i* bins and

i^* is the smallest value of *i* such that $cum_{i^*} > 0.5$.

This interpolation gives an accurate estimation in comparison with the half point of b_{i^*} and b_{i+1}. The second component of the equation scales the amount summed to b_{i^*} according how close b_{i^*} and b_{i+1} each were to 0.5.

Effective diameter (*ED*, μm), which is the ratio of *LWC* to the optical cross sectional area of droplets of a sample droplet spectrum by the following equation:

$$ED = \frac{\sum_{i=1}^{n} p_i r_i^3}{\sum_{i=1}^{n} p_i r_i^2} 2$$ (7)

where *n* is the number of sizing bins, p_i the particle counts for bin *i* and r_i the mean radius in μm of bin *i*.

All our instruments were calibrated in laboratory and /or on site. The calibration of all three instruments was done for size measurements but not for N_c measurements. Also, we should take into account the fact that the instruments faced extreme conditions during the whole campaign, in terms of frequent changes to wind direction and speed and sub-zero temperatures. Those meteorological conditions could possibly lead to unexpected performance.

The sizing accuracy for cloud spectrometers has been estimated as 20% and the concentration accuracy as 16% (Baumgardner, 1983; Dye and Baumgardner, 1984; Baumgardner et al., 2017). The major factors that are usually considered for possible biases in data analysis are coincidence, dead time losses and changing velocity acceptance ratio (*VAR*). Coincidence events occur when more than one droplet is registered by an instrument at the same time resulting in multiple droplets artificially measured as one droplet. Changing velocity acceptance ratio (*VAR*) is a result from the fact that only a part of a laser beam is used to calculate the sampling volume because drops passing the laser beam near the edges are undersized. Lance et al. (2012) showed that at ambient droplet concentrations of 500 cm⁻³, at least 27% undercounting and 20%–30% oversizing bias were observed for CAS. However, during PaCE 2013 droplet number concentrations we observed to reach maximum of 200 cm⁻³ and in majority of cases less than 100 cm⁻³. Due to those low number concentration values we do not take coincidence, dead time losses and *VAR* uncertainties into consideration in this analysis. Liquid water content (*LWC*) as it was derived from the CAS probe has a significant uncertainty of 40% according to Droplet Measurement Technologies Manual. FSSP uncertainties, limitations and corrections have been several times reported in literature (Baumgardner 1984; Baumgardner et al., 1985; Baumgardner and Sportwart, 1990). Febvre et al. (2012) find out that the uncertainty of the FSSP

in derived effective diameter (*ED*) and derived liquid water content (*LWC*) are 3μm and 30% regarding mixed phase clouds. Lance et al. (2012) reported for the CDP probe importance of the coincidence errors even if the number concentrations are as low as 200 cm^{-3} resulting in 25% undercounting error and 30% oversizing error due to coincidence.

During the campaign, measurements were performed with 1 Hz acquisition frequency for all three probes. During the data analysis, minute averages from each cloud probe were calculated when the station was inside a cloud. As we previously highlighted, there were cases when the flows of the cloud probes were blocked. This situation was also visible in raw data. Such cases were cleaned out from all datasets for the final analysis. A typical example case of probe freezing was observed on November 3rd, 2013. The CDP was completely clogged with ice, see Fig. 2, where its size distribution is presented. However, in order to find probe freezing cases a closer look in N_c was carefully done for the whole data set. As an example, Fig. 3 depicts the N_c of each cloud probe during the same day along with the meteorological parameters. There, the sudden decrease in droplet number concentration of the CAS probe from 12:00 and a sudden increase in N_c just before was a clear sign of probes inlet freezing. This behavior in N_c was observed due to the opening of the probe inlet was getting smaller (from the accumulation of ice) which resulted in a raising PAS. During data evaluation we considered the PAS was constant. This led to an underestimation of the PAS which explain this abnormality in the N_c. The same behavior could be also seen for the CDP after 16:00.

3 Results
3.1 Overview of results for PACE 2013

During PaCE 2013, FSSP, CAS, and CDP were installed for the total of 1824, 1080, and 1560 hours respectively, see Table 1 for installation dates. During the campaign, the station was inside a cloud about 664 hours (36.5%). During this period, the temperature ranged from -12.0 to 10.2 °C with averaged temperature -1.9 (std 5.1) °C and the wind speed average was 6.8 (std 2.9) ms^{-1}. The dominant wind direction was ~ 220 ° but there were winds and clouds from all directions. Regarding data coverage, the FSSP and the CAS probe showed good performance with ~ 500 hours (75%) and ~ 220 hours (67%) cleaned data coverage respectively. The CDP probe performed significantly worse with ~ 108 hours (17%) cleaned data coverage. The CDP worse performance was due to its frozen inlet or/and its rotation system during night and encountered frequent operation problems especially at sub-zero temperatures.

3.2 Example cases with all three probes in operation.

Firstly, we present example cases when all three probes were operating. Those cases correspond to different wind directions. This choice was made since we used two different approaches; two probes were installed on a rotating platform (FSSP, CDP) and one probe was installed in a fixed direction (CAS). We provide the time series of the derived N_c, *LWC*, *ED*, *MVD* and the size distributions as they were measured by all three probes along with selected meteorological parameters (temperature, wind speed and wind direction). For each case, the same size range for each probe was adopted. For this reason, we eliminated 12 bins (ranging from 0.51 to 3 μm) of the CAS probe and 2 bins (from 1.2 to 3.5 μm) of the FSSP-100. The

final size ranges for probes comparison were then following: the CAS 3 to 50 μm, the FSSP 3.5 to 47 μm and the CDP 3 to 50 μm.

The above mentioned parameters are presented on November 12th from 15:00 to 18:00, see Fig. 4. The wind speed during this period varied from 6 to 13.5 ms^{-1} and temperature varied from -6.1 to -5.1 °C. Average wind direction was 226.5 (std 6.1)°, which means that all three instruments were pointing to the same direction (Fig. 6e). As we noticed in Fig. 4a, this resulted to a good agreement among all three probes in deriving N_c. The CDP operated without flaws in both its rotation and inhalation system due to cleaning procedure of the instrument done just before this measuring period. In this case, the *LWC* values were in best possible agreement, see Fig. 4b. Derived sizing parameters *ED* and *MVD* are presented in Fig. 4c and d. and both of them had good agreement.

The next example case took place during October, 29th from 12:15 to 15:00, see Fig. 5. The wind speed during this period varied from 2.9 to 8.9 ms^{-1} and temperature varied from -5.4 to 2.9 °C. Average wind direction was 141.2 (std 18.4)°, which indicates that the wind direction was perpendicular to the CAS probe. (Fig. 5e). Here, the CAS probe significantly undercounted N_c (~ factor 5) compared to FSSP and CDP (Fig. 5a). The CDP was also operating with no malfunctions in its rotation and inhalation system. *LWC* as it was derived by the CAS probe was highly affected by a factor of about 7 due to its losses in N_c (Fig. 5b). Furthermore, as we can see in Fig. 5c and d, CAS ability to derive *ED* and *MVD* was not affected by N_c. Especially, when we are comparing *ED* and *MVD* between CAS and FSSP, their difference is less than 20%. However, it was also interesting that even the CDP and the FSSP had a good agreement in droplet counts, they present some differences in the other derived parameters. Investigating in details their size distribution, we found that this was a result of different estimation in sizing. This can be clearly seen in Fig. 5f. In this case there was a shift in CDP sizing when compared with FSSP towards smaller sizes about 5 μm in size range from 15 to 20 μm.

The last example day we present took place on October 25th, from 12:00 to 15:00, see Fig.6. The wind speed varied during this period from 4.4 to 9.9 ms^{-1} and temperature from -5.8 to -5.3 °C. Average wind direction was 85.6 (std 4.8)°, which means that the probes were not facing the same direction (Fig. 6e). This explains why the N_c of CAS was lower than the N_c of FSSP by a factor of 2. The CDP measured the same N_c of cloud droplets as CAS, something that was not really expected (Fig. 6a). This is a typical example case we faced which indicates why the CDP counts were not trusted during sub-zero temperatures. Even if the CDP was not obviously clogged, when observed from the raw data, its rotation system was frozen so it could not follow the wind direction and we were not able to determine where exactly the CDP was pointing. Also, here, we can see that during the period the instruments were not facing the same wind direction, we observed large discrepancies in derived *LWC* (Fig. 6b). In Fig. 6c and 6d we can also see slight discrepancies in derived sizing parameters *ED* and *MVD*. We noticed a slight shift in FSSP sizing towards bigger sizes compared to CAS, about 2.5 μm in size range from 7 to 10 μm and a slight shift in CDP sizing towards smaller sizes, about 2 μm in size range from 5 to 7 μm, as it is depicted in Fig. 6f clearly explains those differences.

3.3 CAS and FSSP counting performance based on the wind direction.

After investigating different example cases, we focused on how the change of the wind direction influenced the droplet counting ability of the instruments in different sectors of the wind rose. In this section, we concentrated only on the inter-comparison of the CAS probe (installed in a fixed direction) with the FSSP (was following the wind direction). The decision to first compare only CAS and FSSP setups was made because their parallel data coverage was the best (~ 243 hours of common cleaned data set). CDP was not used in this section due to only few common data with the CAS. The reason for that was that from the date that the CAS was installed the CDP had operation problems. A detailed analysis regarding the CDP and its behaviour is presented later in section 3.5. To obtain as close as possible size range for both instruments we removed the first ten bins of the CAS and the first bin of the FSSP. As a result, we used the following size ranges; CAS - from 1.25 to 50 µm and FSSP - from 1.2 to 47 µm.

To estimate the possible losses of the cloud droplet counts from each cloud probe we used N_c as a benchmark parameter. Possible CAS sampling losses were investigated by calculating the aspiration efficiency as described in Spiegel et al., 2012. Our expected losses were ~5% for 20 µm, ~20% for 40 µm and 40% for 50 µm cloud droplets. Averaged total N_c values of cleaned data set as they were derived from the CAS probe and the FSSP were 39.8 (std 35.3) cm^{-3} and 44.1(std 26.9) cm^{-3} respectively. We divided the wind rose into 12 parts. This choice was made according to specific factors. First of all, we took into consideration the crucial point orientation of CAS inlet when compares to actual wind direction since the CAS probe was fixed and installed (southwest of the station, ~ 225°). This point helped us to define the areas where the two instruments were performing in wind iso-axial condition. Secondly, we tried to secure that we will have enough amount of data in each part to increase the reliability of our results. Accordingly, the 360 wind rose was divided in to following sectors: 0 to 74, 75 to 94, 95 to 114, 115 to 154, 155 to 184, 185 to 199, 200 to 214, 215 to 235, 236 to 250, 251 to 265, 266 to 295 and 296 to 360°. Fig. 7 shows the ratio of N_c of the CAS to the FSSP probes along with the percentage of observations in each of those sectors of the wind rose and the averaged N_c values from both instruments. There, we can see that each of instruments had different counting performance in each sector. The best counting performance (ratio is close to 1) was found at two sectors (from 200 to 214° and 215 to 235°), where both probes were facing similar direction. On the other hand, when the wind direction was perpendicular (115 to 154°) to CAS fixed direction, the ratio was found lower than 0.4. There the CAS probe undercounted a significant amount of cloud droplets (~ 60%). However, there were also cases where FSSP measured smaller N_c compared to the CAS probe (sectors from 236 to 250° and from 251 to 265°). During those cases, FSSP was not actually freely moving because of the brake set up. Depend on the wind turns, FSSP could be left in wrong orientation for an unknown amount of time. Inside those two sectors the CAS probe measured relatively high N_c (~ 120cm^{-3}) in comparison to the other parts of the wind rose (~ 50 cm^{-3}).

All wind sectors were further investigated to explain the biases in the performance of the two instruments. Firstly, a closer look (see Fig.8) is presented for two sectors (200 - 214, 215 - 235°) where the agreement was found the best according to N_c ratio. For this reason, the wind rose sector from 200 to 235° was adopted as wind iso-axial conditions for the rest of this work. Results indicate that the agreement in both cases was good (R^2 = 0.78 and 0.62 with slope 0.65 and 0.50 respectively, Fig.8a and b) and the maximum difference observed was ~ 30%. When N_c as derived from CAS was more than 80 cm^{-3}, FSSP N_c

was about 25% lower. Temperature and wind speed in range of -11 to -1.4 °C and 1.6 to 13.8 ms^{-1} were also tested for possible biases in wind iso-axial conditions and we found that they did not affected the ability of the probes to derive N_c. A more detailed look of how the two cloud probes measured in wind iso-axial conditions when the station was inside a cloud is presented in Fig. 8c and d, where the averaged number size distribution of the cloud droplets is shown. The CAS probe measured more counts in sizes smaller than 7 µm (~ 3 counts / cm^{-3} more than the FSSP at 1.4 µm and ~ 15 drops / cm^{-3} more at 5 µm, in both cases difference in counts was less than 30%). Also, we can see that the FSSP measured no droplets for sizes larger than 35 µm. Within the size range, meaning from 7 to 20 µm (area which usually represented the peak of the size distribution), the FSSP usually measured higher N_c. This difference could be up to 25% (~ 150 more counts / cm^{-3}). We have noticed also a slight shift in FSSP sizing towards bigger sizes, about 1.5 µm in size range from 7 to 10 µm. Those differences in the counting efficiency of the two instruments explain the slight discrepancies we observed in N_c even when they were measuring in wind iso-axial conditions. In a similar way, all the remained sectors of the wind rose were investigated in detail to reveal more biases (for detailed description please see discussion and Fig. S5, S6 in the Supplement).

From the inter-comparison of the two instruments in each sector of the wind rose, a general benchmarking was created and it is presented in Fig. 9. According to our results we merged some of the wind sectors that we had initially created. As a result, we now have four sectors representing the wind rose; wind iso-axial conditions (from 200 to 235°), perpendicular conditions (from 115 to 154 and 296 to 360°), conditions between iso-axial and perpendicular (from 76 to 114, from 155 to 199 and from 236 to 295°) and the special case where the brake influenced the performance of FSSP (from 1 to 74°). To summarize our results, we should highlight that the best agreement between the two cloud spectrometers was obtained in wind iso-axial conditions (from 200 to 235°, $R^2 = 0.60$) and it covered a cleaned data set of ~ 66 observation hours. The effect of wind direction on the CAS probes ability to measure N_c was immediately noticed when the wind direction was out of the range of the wind iso-axial conditions. The agreement became slightly worst when the spectrometers were facing conditions that wind direction was between iso-axial and perpendicular ($0.46 \leq R^2 \leq 0.50$ for 76 to 114, 155 to 199 and 236 to 295° respectively, ~ 50% of total cleaned data set). The CAS probe performed the worst when the wind direction was perpendicular to the CAS installed direction ($R^2 = 0.32$ and 0.11 for 115 to 154 and 296 to 360° respectively) and represents ~ 40 observation hours. Guyot et al. (2015) performed a similar experiment to investigate the sensitivity of the FSSP to meteorological parameters. Even though we conducted the measurements at different temperatures (in Puy-de-Dôme they sampled clouds only above zero) we found that our results were related. The main reason that caused the discrepancies (mainly in deriving N_c and LWC) to the fixed direction cloud spectrometers ground setups (Pallas – CAPS and Puy-de-Dôme - FSSP) was the wind direction. The strong sensitivity to the wind direction suggested that the cloud spectrometers were sampling anisokinetically in both cases.

3.4 Inter-comparison of CAS and FSSP - derived parameters *LWC*, *ED* and *MVD*.

In this section, we focused on investigating the derived parameters *LWC*, *ED* and *MVD*. First, a comparison of the liquid water content (*LWC)* for the two probes CAS and FSSP is presented. We only present measurements that were performed in wind iso-axial conditions, since the *LWC* was very sensitive to both changes in droplet N_c and changes in shape of the number

size distribution. The discrepancies we observed in droplet N_c in sectors outside the wind iso-axial conditions caused a significant difference in total LWC at least by a factor of 5 or even more. We also noticed differences by factor of 15 especially when the wind direction was perpendicular to CAS fixed direction. Figure 10a shows that the agreement in LWC ($R^2 = 0.34$ and slope 0.53) between CAS and FSSP in iso-axial conditions was found worse than agreement of both probes in N_c. After investigating how different meteorological parameters contribute to apparent biases in more detail, we found that temperature was the main and only factor that affected the instruments ability to derive LWC. Accordingly, we divided our measurements in two temperature data sets. Measurements with temperature range from -11.1 to -4 and from -3.9 to -1.4 °C. Figure 10b presents the agreement for the first set of measurements, temperatures below -4 °C. Excluding the warmer temperature range, we obtained better agreement between the probes ($R^2 = 0.78$ and slope 0.82). On other hand, the second set of temperatures (from -3.9 to -1.4 °C) indicated that the two probes significantly disagreed ($R^2 = 0.02$ and slope 0.07). As we already explained in section 3.3, there was a slight shift in FSSP sizing towards bigger sizes, about 1.5 µm in size range from 7 to 10 µm. However, when applied the correction to FSSP sizing the resulting change in LWC was found marginal (about 0.7%).

Our main conclusion regarding the derived LWC was that the main factor affecting LWC values was the actual difference in the counts in each bin, especially when referring to larger droplets. Taking into account those limitations and biases in deriving LWC our final proposal is to use only LWC values from wind iso-axial conditions. In addition, the critical parameter should be the temperature of the cloud. This suggests that only derived LWC values for temperatures below -4 °C will be considered as acceptable and will be used for further analysis of this data set. However, even when we consider the best agreement the maximum difference in obtained LWC between CAS and FSSP could still be about 40%. In addition, we suggest the deployment of another LWC sensor, e.g. the particle volume monitor (PVM-100, Gerber 1999) during future campaigns in order to obtain another reference LWC values for inter-comparison in wide temperature range. In addition, we are continuously following the development of a new generation of counters designed for ground based in-situ cloud measurements. Thus, it is a matter of future deployment during upcoming PaCE campaigns.

The final step to complete the inter-comparison between the CAS probe and the FSSP was to investigate their ability to derive two sizing parameters, the ED and the MVD. Both of them are significant to identify and evaluate the sizing performance of the cloud spectrometers (e.g. Stephens, 1978; Slingo and Schencker, 1982; Korolev, 1999; Mitchell et al., 2011). The cleaned dataset obtained from the whole wind spectrum plotted in different color scale for wind directions, temperature and wind speed is presented in Fig. 11 a, b and c for ED and in Fig. 11 d, e and f for MVD. The observations when the FSSP did not operate properly due to installation of the brake were excluded from the inter-comparison. The agreement for both sizing parameters was found good ($R^2 = 0.80$, slope 0.79 and $R^2 = 0.78$, slope 0.75 respectively). The best agreement was observed when the wind direction (see Fig.13a, d) was inside the range of iso-axial conditions where all the points were focused along the 1:1 line. When the direction was perpendicular the points were spread wider (maximum observed difference between the two probes was about 20%). Surprisingly, despite the fact that CAS was measuring lower N_c even by a factor of 10 when the wind direction was perpendicular to CAS fixed direction, the derived ED and MVD were not significantly influenced. Both sizing parameters were derived from the measured size distribution as described in section 2.3. We found that even if significant

number of cloud droplets was lost due to inertia, the shape and the position of the peak of the size distribution measured by CAS remained similar. This behaviour was found the same through the whole available cleaned data set (~ 183 hours) with the maximum *ED* and *MVD* of 35 and 30 μm. It has to be pointed out that this behaviour might be exclusive for sub-Arctic conditions with generally small cloud droplets. This fact allows us to use the majority of the data set when investigating those two derived sizing parameters. As a result, it creates a significant and usable data set without need to disqualify data according to particular wind direction. Thus we obtain statistically significant size properties of the cloud droplets at wide range of meteorological conditions. We also investigated the probability that wind speed will affect the sizing parameters (see Fig. 11c, f). When the probes were facing high wind speed, *ED* and *MVD* were slightly influenced (FSSP derived bigger values of *ED* and *MVD* when compared to CAS). On the other hand, while they were facing low wind speeds sizing were again influenced on the opposite way (FSSP derived smaller values when compared to CAS). This could happen due to the isokinetic motion of the particles. The larger particles could not enter the FSSP because the inner diameter necking on the inlet was changing from 3.8 to 2.0 cm. Finally, Fig. 11 b, d indicates that at lower temperatures we observed smaller *ED* and *MVD* values.

According to previous analysis, our main conclusion was that even if there were slight biases and uncertainties the agreement in inter-comparison was considered good as both R^2 and slope were higher than 0.75. As a result, we propose when deriving the sizing parameters, *ED* and *MVD*, all measurements can be used for further research after carefully exclude the FSSP dataset that was obtained from the wind rose sector where the brake influenced its performance.

3.5 Evaluation of the CDP during PaCE 2013.

After comparing and analyzing discrepancies and biases between the CAS and the FSSP cloud probes, we discuss the performance of the CDP cloud probe separately. To evaluate CDP performance we used only FSSP data. We should remind that during the period that both the CDP and the CAS probe were on site (from 15 October to 27 November), the CDP encountered several malfunction and operational problems during icing conditions. As a result, there was a lack of common data between the CDP and the CAS probe.

We used CDP and FSSP data from September 25th to October 14th (~ 70 hours of cleaned data set) since it was the only period that the CDP faced fewer operational problems since average ambient temperatures were mostly above 0 °C. During this inter-comparison, a set of data from the FSSP was removed (0 to 74°, where the FSSP had significant malfunctions due to installation of the brake). For this time period, average temperature at the station was 1.7 (std 1.6) °C and the averaged wind speed was 6.9 (std 3.6) ms^{-1}. In order to compare similar size range for both cloud probes the first 2 bins from the FSSP were removed. This means that the following results depict size range from 3.5 to 47 μm for the FSSP and from 3 to 50 μm for the CDP probe.

As it was already mentioned at the beginning of section 3.2 those two instruments belong to the sub category of the probes that were installed on a rotating platform during PaCE 2013. Figure 12a shows, as it was expected, that the ability of the two instruments to derive N_c was good ($R^2 = 0.84$, slope 1.11). However, there were cases where the difference between them was about 30%. Additionally, we investigated the derived sizing parameters *ED* and *MVD*, see fig.12 b and c. In a range of

temperatures (from -3.9 to 3.8 °C) and wind speeds (0.9 to 19 ms^{-1}) agreement corresponding to the sizing parameters was good (R^2 = 0.82 and 0.79 with slopes 1.23 and 1.25 for *ED* and *MVD* respectively). However, when FSSP derived *ED* and *MVD* for sizes larger than 22.5 µm, we could see a difference that could be even 15 µm smaller in comparison with CDP. This difference was noticed especially when the wind speeds were low. FSSP had similar behavior (section 3.4) when we were comparing CAS and FSSP due to isokinetic motion of the particles.

A significant limitation in derived *LWC* regarding temperature was already discussed above during the comparison of CAS and FSSP. In this case, the temperature ranged from -3.9 to 3.8 °C. This range that was above -4 °C (the temperature point that was set in section 3.4). As a result, the comparison of CDP with FSSP derived *LWC* did not lead to reasonable correlation and no *LWC* data are presented here.

In summary, the CDP was operating well in warm liquid clouds and had good agreement in cloud droplet counts and the sizing derived parameters with FSSP. On the other hand, while we faced sub-zero conditions the CDP operation was considered as problematic. Its probe inlet became often clogged because of supercooled cloud drops accumulation. This happened to both the rotation and inhalation system because its big surfaces were collecting ice and it had a small opening of inhalation system. In conclusion, even that this CDP setup performed well in warm cloud conditions, it is not suitable instrument for semi-long term ground based measurement of clouds in sub-Artic conditions, when we are facing subzero conditions.

4. Conclusions

We conducted ground based in-situ cloud measurements during PaCE 2013 from September 14th until November 28th. We deployed three cloud spectrometers' setups (CAS, FSSP and CDP) on the roof of Sammaltunturi station, located in Finnish sub-Arctic. The obtained data set was analyzed in detail to evaluate the instruments ground based setups' performance and established limitations for future studies. All cloud spectrometers and their setups are owned by FMI and results could be used in campaigns with similar meteorological conditions, sub-Arctic conditions with frequently occurring supercooled clouds. Although there is a possibility we sampled ice particles in some cases, it is expected that the number of supercooled liquid droplets greatly exceed the number of small ice cloud droplets. The CAS was installed and fixed against the main wind direction of the station (~ 225°) and the other two probes (FSSP and CDP) were installed on rotating platforms. Each probe had its own inhalation system. Their ability to measure the size distribution of cloud droplets along with the derived number concentration (N_c), the sizing parameters (*ED* and *MVD*) and the liquid water content (*LWC*) was tested and the above parameters were mutually compared. In this work, CAS and FSSP ground setups were investigated first because their parallel data coverage was the best (~ 243 hours of common cleaned data set). On the other hand, CDP had low common data set with the CAS. The reason for that was that from the date that the CAS was installed we were mainly facing sub-zero temperatures, conditions that proved that were not favorable for this CDP ground setup.

To estimate the droplet counting performance and possible droplets losses, we used number concentration (N_c). Results indicated that when we were deriving N_c, the mutual direction of probe heading and the wind direction were playing the most significant role. From the inter-comparison of the CAS (fixed orientation) against FSSP (rotating platform), it was found that

the CAS probe had the best agreement ($R^2=0.70$) with the FSSP during wind iso-axial conditions (200 to 235°). The CAS probe counting efficiency was strongly dependent on the wind direction, this can be clearly explained by its installation to fixed orientation. When the station was inside warm clouds, both the CDP and the FSSP had good agreement ($R^2= 0.82$) as they were both operating on rotating platforms.

5   The *LWC* was found the most sensitive derived parameter. This is because *LWC* strong dependency both on size and the number of droplets in each size bin. Thus, the wind direction played again the most significant role. For that reason, we strongly recommend that CAS and FSSP derived *LWC* values only from wind iso-axial conditions should be used. Additionally, *LWC* values were found also temperature dependent. For temperatures lower than -4 °C the agreement between the CAS and the FSSP in wind iso-axial conditions was high ($R^2 =0.62$) and that is why temperature -4 °C was adopted as the critical

10   temperature point regarding *LWC* estimation. We excluded all *LWC* values derived from the CDP due to its problematic operation at sub-zero conditions i.e. close to temperature -4 °C , CDP was usually frozen. The CDP probe often accumulated ice in sub-zero condition, both in its rotational platform and inhalation system. This was happening due to presence of supercooled clouds at the station. The big surfaces of the CDPs rotation platform and the inlet with small opening were collecting ice very fast.

15   The analysis of the derived sizing parameters, *ED* and *MVD*, showed good agreement among the three probes during the time they were operating properly. However, our conclusions here concentrate only to CAS and FSSP, because only those two instruments were operating properly in subzero temperatures, temperatures that we usually face during PaCE. The obtained inter-comparison results were surprisingly good even though CAS lost a significant amount of cloud droplets due to its orientation. The wind direction did not significantly affect neither the *ED* values nor the *MVD*, even though large discrepancies

20   (uncertainty ~ 85%) in N_c outside of the wind iso-axial conditions could be found (e.g. when the wind direction was perpendicular to the CAS probe fixed direction, uncertainty for sizing parameters was ~ 20%). The *ED* and *MVD* was not affected because the shape and the peak position of the CAS size droplet distribution did not change significantly. Such behavior held through ~ 183 hours of data set. This result is important as it allows us to use a larger data set without limitations due to wind direction and other meteorological parameters regarding derived *ED* and *MVD*. The small differences (about 2

25   μm) were explained by a closer look in size distribution of each spectrometer and the differences in sizing during operation in real conditions as they were mentioned above.

At the time of PaCE2013, the market did not offer any instrumentation fulfilling our requirements on continuous long-term unattended operation at subzero conditions. As final suggestions regarding performing continuous ground based in-situ cloud measurements in harsh environments, we would like to highlight two major issues. First, the cloud probes should always

30   continuously face the wind direction to minimize the sampling losses. If this is not secured, only the measurements that were conducted in wind iso-axial conditions can be used for further analysis. However, deriving the sizing parameters *ED* and *MVD* for the whole wind direction spectrum is still possible, but must be done with insight and prudence. Secondly, deicing features of the ground setups should be upgraded to make possible their long-term unattended operation in subzero conditions.

Otherwise, the cloud probes need necessary daily or more frequent checkups and cleaning of their inlets. Our final recommendations and our view on the main limitations of each spectrometer ground setup for using and analyzing the obtained data sets during sub-Arctic meteorological conditions with frequently occurring supercooled clouds (including future PaCE campaigns) are summarized and presented in Table 2.

Data availability

The cloud probes and meteorological data used here are available upon request to the corresponding author (Konstantinos.doulgeris@fmi.fi).

Author contribution

KMD, DB wrote the manuscript with contribution from all co-authors. KMD prepared the manuscript and analyzed data from all cloud probes and meteorological data with contribution from DB. DB installed and operated all the instruments during PACE2013. MK and SR provided the CDP ground set up and performed its calibration.

Competing interests

The authors declare that they have no conflict of interest.

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References

Albrecht, B. A.: Aerosols, cloud microphysics, and fractional cloudiness, Science, 245, 1227–1230, 1989.

Anttila, T., Vaattovaara, P., Komppula, M., Hyvärinen, A.-P., Lihavainen, H., Kerminen, V.-M., and Laaksonen, A.: Size-dependent activation of aerosols into cloud droplets at a subarctic background site during the second Pallas Cloud Experiment (2nd PaCE): method development and data evaluation, Atmos. Chem. Phys., 9, 4841-4854, https://doi.org/10.5194/acp-9-4841-2009, 2009.

Anttila, T., Brus, D., Jaatinen, A., Hyvärinen, A.-P., Kivekäs, N., Romakkaniemi, S., Komppula, M., and Lihavainen, H.: Relationships between particles, cloud condensation nuclei and cloud droplet activation during the third Pallas Cloud Experiment, Atmos. Chem. Phys., 12, 11435-11450, https://doi.org/10.5194/acp-12-11435-2012, 2012.

Baumgardner, D.: An analysis and comparison of five water droplet measuring instruments, J. Appl. Meteorol., 22, 891–910, https://doi.org/10.1175/1520-0450(1983)022<0891:AAACOF>2.0.CO;2, 1983.

Baumgardner, D., Strapp, W., and Dye, J. E.: Evaluation of the Forward Scattering Spectrometer Probe – Part II: Corrections for coincidence and dead-time losses, J. Atmos. Oceanic Technol.,2, 626–632, DOI: 10.1175/1520-0426(1985)002<0626:EOTFSS>2.0.CO;2 1985.

Baumgardner, D., Dye, J. E., and Gandrud, B. W.: Calibration of the forward scattering spectrometer probe used on the ER-2 during the Airborne Antarctic Ozone Experiment, Geophys. Res., 94, 16475-16480, doi: 10.1029/JD094iD14p16475, 1989.

Baumgardner, D. and Spowart, M.: Evaluation of the Forward Scattering Spectrometer Probe. Part III: Time Response and Laser Imhomogeneity Limitations, J. Atmos. Oceanic Technol., 7, 666–672, https://doi.org/10.1175/1520-0426(1990)007<0666:EOTFSS>2.0.CO;2 1990.

Baumgardner, D., Dye, J. E., Gandrud, B. W., and Knollenberg, R. G.: Interpretation of Measurements Made by the Forward Scattering Spectrometer Probe (FSSP-300) During the Airborne Arctic Stratospheric Expedition, 97, 8035-8046, 10.1029/91JD02728., 1992.

Baumgardner, D., Jonsson, H., Dawson, W., O'Connor D., and Newton R.: The cloud, aerosol and precipitation spectrometer (CAPS): A new instrument for cloud investigations, Atmos. Res. 59-60, 251-264, doi: 10.1016/S0169-8095(01)00119-3, 2001.

Baumgardner, D., Brenguier, J., Bucholtz, A., Coe, H., DeMott, P., Garrett, T., Gayet, J., Hermann, M., Heymsfield, A., Korolev, A., Kramer, M., Petzold, A., Strapp, W., Pilewskie, P., Taylor, J., ¨ Twohy, C., Wendisch, M., Bachalo, W., and Chuang, P.: Airborne instruments to measure atmospheric aerosol particles, clouds and radiation: A cook's tour of mature and emerging technology, Atmos. Res., 102, 10–29, doi:10.1016/j.atmosres.2011.06.021, 2011.

Baumgardner, D., Abel, S. J., Axisa, D., Cotton, R., Crosier, J., Field, P., Gurganus, C., Heymsfield, A., Korolev, A., Krämer, M., Lawson, P., McFarquhar, G., Ulanowski, Z., and Um, J.: Cloud Ice Properties: In Situ Measurement Challenges, Meteorological Monographs, 58, 9.1–9.23, https://doi.org/10.1175/AMSMONOGRAPHS-D-16-0011.1, 2017.

Boucher, O., D. Randall, P. Artaxo, C. Bretherton, G. Feingold, P. Forster, V.-M. Kerminen, Y. Kondo, H. Liao, U. Lohmann, P. Rasch, S.K. Satheesh, S. Sherwood, B. Stevens, and X.Y. Zhang, 2013: Clouds and aerosols. In *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*. T.F. Stocker, D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Doschung, A. Nauels, Y. Xia, V. Bex, and P.M. Midgley, Eds. Cambridge University Press, pp. 571-657, doi:10.1017/CBO9781107415324.016, 2013.

Brenguier, J.-L., Bachalo, W. D., Chuang, P. Y., Esposito, B. M., Fugal, J., Garrett, T., Gayet, J.-F., Gerber, H., Heymsfield, A., Kokhanovsky, A., Korolev, A., Lawson, R. P., Rogers, D. C., Shaw, R. A., Strapp, W., and Wendisch, M.: In Situ Measurements of Cloud and Precipitation Particles, in Airborne Measurements for Environmental Research, in: Methods and Instruments, edited by: Wendisch, M. and Brenguier, J.-L., Wiley VCH Verlag GmbH & Co. KGaA, Weinheim, Germany, 225– 301, doi:10.1002/9783527653218.ch5, 2013.

Burnet F. and Brenguier, J. L.: Validation of droplet spectra and liquid water content measurements, Phys. Chem. Earth, 24, 249– 254, https://doi.org/10.1016/S1464-1909(98)00046-X, 1999.

Burnet F. and Brenguier, J. L.: Comparison between standard and modified Forward Scattering Spectrometer Probes during the Small Cumulus Microphysics Study, J. Atmos. Oceanic Technol., 19, 1516–1531, https://doi.org/10.1175/1520-0426(2002)019<1516:CBSAMF>2.0.CO;2 ,2002.

Coelho, A., Brenguier, J. L., and Perrin, T.: Droplet spectra measurements with the FSSP-100. Part I: Low droplet concentration measurements. J. Atmos. Oceanic Technol., 22, 1749–1756 https://doi.org/10.1175/JTECH1817.1, 2005.

Coelho, A., Brenguier, J. L., and Perrin, T.: Droplet spectra measurements with the FSSP-100. Part II: Coincidence Effects, J. Atmos. Oceanic Technol., 22, 1749–1756, https://doi.org/10.1175/JTECH1818.1, 2005.

Collaud Coen, M., Andrews, E., Asmi, A., Baltensperger, U., Bukowiecki, N., Day, D., Fiebig, M., Fjaeraa, A. M., Flentje, H., Hyvärinen, A., Jefferson, A., Jennings, S. G., Kouvarakis, G., Lihavainen, H., Lund Myhre, C., Malm, W. C., Mihapopoulos, N., Molenar, J. V., O'Dowd, C., Ogren, J. A., Schichtel, B. A., Sheridan, P., Virkkula, A., Weingartner, E.,

Wandinger U., Apituley A., Blumenstock T., Bukowiecki N., Cammas J.-P., Connolly P., De Mazière M., Dils B., Fiebig M., Freney E., Gallagher M., Godin-Beekmann S., Goloub P., Gysel M., Haeffelin M., Hase F., Hermann M., Herrmann H., Jokinen T., Komppula M., Kubistin D., Langerock B., Lihavainen H., Mihalopoulos N., Laj P., Lund Myhre C., Mahieu E., Mertes ., Möhler O., Mona L., Nicolae D., O'Connor E., Palm M., Pappalardo G., Pazmino A., Petäjä T., Philippin S., Plass-Duelmer C., Pospichal B., Putaud J.-P., Reimann S., Rohrer F., Russchenberg H., Sauvage S., Sellegri K., Steinbrecher R., Stratmann F., Sussmann R., Van Pinxteren D., Van Roozendael M., Vigouroux C., Walden C., Wegene R. and Wiedensohler A. ACTRIS-PPP Deliverable D5.1: Documentation on technical concepts and requirements for ACTRIS Observational Platforms,https://www.actris.eu/Portals/46/Documentation/ACTRIS%20PPP/Deliverables/Public/WP5_D5.1_M18.pdf?ver= 2018-06-28-125343-273 , 2018.

Weller, R., and Laj, P.: Aerosol decadal trends – Part 1: In-situ optical measurements at GAW and IMPROVE stations, Atmos. Chem. Phys., 13, 869-894, https://doi.org/10.5194/acp-13-869-2013, 2013.

Droplet Measurement Technologies Manual: CAPS operator manual, DOC-0066 Revision F, DMT, Boulder, Colorado, USA, 2011

Dye, J. E. and Baumgardner, D.: Evaluation of the forward scattering spectrometer probe, I – Electronic and optical studies, J. Atmos. Ocean. Tech., 1, 329–344, https://doi.org/10.1175/1520-0426(1984)001<0329:EOTFSS>2.0.CO;2 , 1984.

Faber, S., French, J. R., and Jackson, R.: Laboratory and in-flight evaluation of measurement uncertainties from a commercial Cloud Droplet Probe (CDP), Atmos. Meas. Tech., 11, 3645–3659, https://doi.org/10.5194/amt-11-3645-2018, 2018.

Febvre, G., Gayet, J.-F., Shcherbakov, V., Gourbeyre, C., and Jourdan, O.: Some effects of ice crystals on the FSSP measurements in mixed phase clouds, Atmos. Chem. Phys., 12, 8963–8977, doi: 10.5194/acp-12-8963-2012, 2012.

Gerber, H., Frick, G., and Rodi, A.: Ground-based FSSP and PVM measurements of liquid water content. J. Atmos. Oceanic Technol., 16, 1143–1149, doi: 10.1175/1520-0426(1999)016,1143: GBFAPM.2.0.CO;2, 1999.

Gérard, V., Noziere, B., Fine, L., Ferronato, C., Singh, D., K., Frossard, A., Cohen, R.C. Asmi, E., Lihavainen, H.,Kivekäs, N., Aurela, M., Brus, D., Frka, S. and Cvitešić Kušan, A.: Concentrations and Adsorption Isotherms for Amphiphilic Surfactants in PM1 Aerosols from Different Regions of Europe, *Environ. Sci. & Tech.* DOI: 10.1021/acs.est.9b03386, 2019.

Guyot, G., Gourbeyre, C., Febvre, G., Shcherbakov, V., Burnet, F., Dupont, J.-C., Sellegri, K., and Jourdan, O.: Quantitative evaluation of seven optical sensors for cloud microphysical measurements at the Puy-de-Dôme Observatory, France, Atmos. Meas. Tech., 8, 4347-4367, https://doi.org/10.5194/amt-8-4347-2015, 2015.

Hyvärinen A.-P., Kolmonen P., Kerminen V.-M., Virkkula A., Leskinen A., Komppula M., Hatakka J., Burkhart J., Stohl A., Aalto P., Kulmala M., Lehtinen K.E.J., Viisanen Y. and Lihavainen H.: Aerosol black carbon at five background measurement sites over Finland, a gateway to the Arctic. Atmos. Environ. 45, 4042–4050, https://doi.org/10.1016/j.atmosenv.2011.04.026, 2011.

Jaatinen A., Romakkaniemi J., Anttila T., Hyvärinen A.-P., Hao L.Q., Kortelainen A., Miettinen P., Mikkonen S., Smith J.N., Virtanen A. and Laaksonen A.: The third Pallas Cloud Experiment: Consistency between the aerosol hygroscopic growth and CCN activity. Boreal Env. Res. 19, 368–382 http://hdl.handle.net/10138/228774, 2014.

Johnson, B., Turnbull, K., Brown, P., Burgess, R., Dorsey, J., Baran, A. J., Webster, H., Haywood, J., Cotton, R., Ulanowski, Z., Hesse, E., Woolley, A., and Rosenberg, P.: In situ observations of volcanic ash clouds from the FAAM aircraft during the eruption of Eyjafjallajökull in 2010, J. Geophys. Res.-Atmos., 117, D00U24, https://doi.org/10.1029/2011JD016760, 2012.

Jones, H. M., Haywood, J., Marenco, F., O'Sullivan, D., Meyer, J., Thorpe, R., Gallagher, M. W., Krämer, M., Bower, K. N., Rädel, G., Rap, A., Woolley, A., Forster, P., and Coe, H.: A methodology for in-situ and remote sensing of microphysical and radiative properties of contrails as they evolve into cirrus, Atmos. Chem. Phys., 12, 8157–8175, https://doi.org/10.5194/acp12-8157-2012, 2012.

5   Korolev, A.V., Strapp, W.J., and Nevzorov, N. A.: In situ measurements of effective diameter and effective droplet number concentration, Journal of Geophysical Research, 104, 3993-4003, 10.1029/1998JD200071, 1999.

Korolev, A.V., Shashkov, A., and Barker, H.: Calibrations and Performance of the Airborne Cloud Extinction Probe, Journal of Atmospheric and Oceanic Technology, 31, 326-345, 10.1175/JTECH-D-13-00020.1, 2014.

Kivekas, N., Kerminen, V.-M., Raatikainen, T., Vaattovaara, P., and ¨ Lihavainen, H.: Physical and chemical characteristics
10   of the activating particles during the Second Pallas Cloud Experiment (Second PaCE), Boreal. Environ. Res., 14, 515–526, 2009.

Komppula, M., Lihavainen, H., Kerminen, V.-M., Kulmala, M., and Viisanen, Y.: Measurements of cloud droplet activation of aerosol particles at a clean subarctic background site, J. Geophys. Res., 110, D06204, doi: 10.1029/2004JD005200, 2005.

Knollenberg, R. G.: Techniques for Probing Cloud Microstructure, in: Clouds, Their Formation, Optical Properties and
15   Effects, edited by: Hobbs, P. V. and Deepak, A., Academic Press, New York, NY, USA, 15–92, 1981.

Lance, S., Brock, C. A., Rogers, D., and Gordon, J. A.: Water droplet calibration of the Cloud Droplet Probe (CDP) and inflight performance in liquid, ice and mixed-phase clouds during ARCPAC, Atmos. Meas. Tech., 3, 1683–1706, doi:10.5194/amt3-1683-2010, 2010.

Lance, S.: Coincidence Errors in a Cloud Droplet Probe (CDP) and a Cloud and Aerosol Spectrometer (CAS),
20   and the Improved Performance of a Modified CDP, https://doi.org/10.1175/JTECH-D-11-00208.1, 2012.

Lihavainen, H., Kerminen, V.-M., Komppula, M., Hyvarinen, A.- P., Laakia, J., Saarikoski, S., Makkonen, U., Kivekas, N., Hillamo, R., Kulmala, M., and Viisanen, Y.: Measurements of the relation between aerosol properties and microphysics and chemistry of low level liquid water clouds in Northern Finland, Atmos. Chem. Phys., 8, 6925–6938, doi:10.5194/acp-8-6925-2008, 2008.

25   Lihavainen, H., Kerminen, V.-M., and Remer, L. A.: Aerosol-cloud interaction determined by both in situ and satellite data over a northern high-latitude site, Atmos. Chem. Phys., 10, 10987-10995, https://doi.org/10.5194/acp-10-10987-2010, 2010.

Lihavainen H., Hyvärinen A.-P., Asmi E., Hatakka J. and Viisanen Y.: Long term variability of aerosol optical properties in northern Finland. Boreal Env. Res. 20, 526–541, http://hdl.handle.net/10138/228289, 2015.

Lloyd, G., Choularton, T. W., Bower, K. N., Gallagher, M. W., Connolly, P. J., Flynn, M., Farrington, R., Crosier, J.,
30   Schlenczek, O., Fugal, J., and Henneberger, J.: The origins of ice crystals measured in mixed-phase clouds at the high-alpine site Jungfraujoch, Atmos. Chem. Phys., 15, 12953–12969, https://doi.org/10.5194/acp-15-12953-2015, 2015.

Lowenthal, D. H., Hallar, A. G., David, R. O., McCubbin, I. B., Borys, R. D., and Mace, G. G.: Mixed-phase orographic cloud microphysics during StormVEx and IFRACS, Atmos. Chem. Phys., 19, 5387–5401, https://doi.org/10.5194/acp-19-
35   5387-2019, 2019.

Luebke, A. E., Afchine, A., Costa, A., Grooß, J.-U., Meyer, J., Rolf, C., Spelten, N., Avallone, L. M., Baumgardner, D., and Krämer, M.: The origin of midlatitude ice clouds and the resulting influence on their microphysical properties, Atmos. Chem. Phys., 16, 5793–5809, https://doi.org/10.5194/acp-16-5793-2016, 2016.

McFarquhar, G. M., Um, J., Freer, M., Baumgardner, D., Kok, G. L., and Mace, G.: Importance of small ice crystals to cirrus properties: Observations from the Tropical Warm Pool International Cloud Experiment (TWP-ICE), Geophys. Res. Lett., 34, L13803, doi: 10.1029/2007GL029865, 2007.

Mertes, S., Schwarzenböck, A., Laj, P., Wobrock, W., Pichon, J. M., Orsi, G., and Heintzenberg, J.: Changes of cloud microphysical properties during the transition from supercooled to mixed-phase conditions during CIME, Atmos. Res., 58, 267–294, https://doi.org/10.1016/S0169-8095(01)00095-3, 2001

Mitchell, D. L., Lawson, R. P., and Baker, B.: Understanding effective diameter and its application to terrestrial radiation in ice clouds, Atmos. Chem. Phys., 11, 3417-3429, https://doi.org/10.5194/acp-11-3417-2011, 2011.

Nichman, L., Järvinen, E., Dorsey, J., Connolly, P., Duplissy, J., Fuchs, C., Ignatius, K., Sengupta, K., Stratmann, F., Möhler, O., Schnaiter, M., and Gallagher, M.: Intercomparison study and optical asphericity measurements of small ice particles in the CERN CLOUD experiment, Atmos. Meas. Tech., 10, 3231-3248, https://doi.org/10.5194/amt-10-3231-2017, 2017.

Pruppacher, H. R. and Klett, J. D.: Microphysics of Clouds and Precipitation, Kluwer Academic Publishers, Dordrecht, 1997.

Raatikainen, T., Brus, D., Hyvärinen, A.-P., Svensson, J., Asmi, E., and Lihavainen, H.: Black carbon concentrations and mixing state in the Finnish Arctic, Atmos. Chem. Phys., 15, 10057-10070, https://doi.org/10.5194/acp-15-10057-2015, 2015.

Rosenfeld, D. and Ulbrich, C.W.: Cloud Microphysical Properties, Processes, and Rainfall Estimation Opportunities, Meteorological Monographs, 30, 237-237, DOI: 10.1175/0065 9401(2003)030<0237:CMPPAR>2.0.CO;2, 2003.

Small, J. D., P. Y. Chuang, G. Feingold, and H. Jiang: Can aerosol decrease cloud lifetime?, Geophys. Res. Lett., 36, L16806, doi: 10.1029/2009GL038888, 2009.

Smith, H., Connolly, P., Baran, A., Hesse, E., Smedley, A. and Webb, A.: Cloud chamber laboratory investigations into scattering properties of hollow ice particles: Journal of Quantitative Spectroscopy and Radiative Transfer. 157, 106-118 13 p. https://doi.org/10.1016/j.jqsrt.2015.02.015, 2015.

Slingo, A., and Schencker H. M.: On the shortwave radiative properties of stratiform water clouds, Q. J. R. Meteorol. Soc., 108, 407-426, https://doi.org/10.1002/qj.49710845607, 1982.

Stephens, G. L., Radiation profiles in extended water clouds, II, Parameterization scheme, J. Atmos. Sci., 35, 2123-2132, https://doi.org/10.1175/1520-0469(1978)035<2123:RPIEWC>2.0.CO;2, 1978.

Spiegel, J. K., Zieger, P., Bukowiecki, N., Hammer, E., Weingartner, E., and Eugster, W.: Evaluating the capabilities and uncertainties of droplet measurements for the fog droplet spectrometer (FM-100), Atmos. Meas. Tech., 5, 2237–2260, doi:10.5194/amt5-2237-2012, 2012.

Voigt, C., Schumann, U., Minikin, A., Abdelmonem, A., Afchine, A., Borrmann, S., Boettcher, M., Buchholz, B., Bugliaro, L., Costa, A., Curtius, J., Dollner, M., Dörnbrack, A., Dreiling, V., Ebert, V., Ehrlich, A., Fix, A., Forster, L., Frank, F., Fütterer, D., Giez, A., Graf, K., Grooß, J.-U., Groß, S., Heimerl, K., Heinold, B., Hüneke, T., Järvinen, E., Jurkat, T., Kaufmann, S., Kenntner, M., Klingebiel, M., Klimach, T., Kohl, R., Krämer, M., Krisna, T. C., Luebke, A., Mayer, B., Mertes, S., Molleker, S., Petzold, A., Pfeilsticker, K., Port, M., Rapp, M., Reutter, P., Rolf, C., Rose, D., Sauer, D., Schäfler, A., Schlage, R., Schnaiter, M., Schneider, J., Spelten, N., Spichtinger, P., Stock, P., Walser, A., Weigel, R., Weinzierl, B., Wendisch, M., Werner, F., Wernli, H., Wirth, M., Zahn, A., Ziereis, H., and Zöger, M.: ML-CIRRUS: The Airborne Experiment on Natural Cirrus and Contrail Cirrus with the High-Altitude Long-Range Research Aircraft HALO, B. Am. Meteorol. Soc., 98, 271–288, https://doi.org/10.1175/bams-d-15- 00213.1, 2017.

Wagner, R., Bunz, H., Linke, C., Möhler, O., Naumann, K. H., Saathoff, H., Schnaiter, M., and Schurath, U.: Chamber simulations of cloud chemistry: the AIDA Chamber, in: NATO Adv. Res. Work. Environ. Simul. Chambers Appl. to Atmos. Chem. Process. (2004 Zakopane, Poland), edited by: Barnes, I. and Rudzinski, K. J., Springer, Dordrecht, the Netherlands, 2006.

5   Wendisch, M. : A quantitative comparison of ground-based FSSP and PVM measurements, J. Atmos. Oceanic Technol., 15, 887– 900, DOI: 10.1175/1520-0426(1998)015<0887:AQCOGB>2.0.CO;2, 1998.

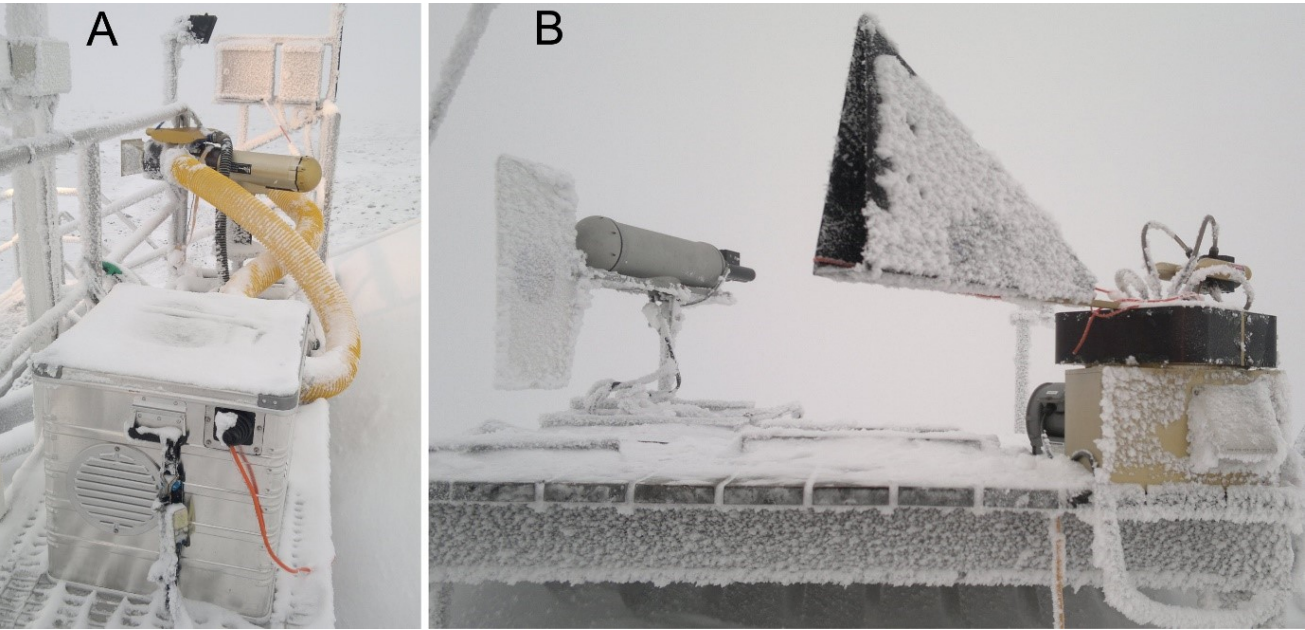
Figures

10   Figure 1. A) The CAPS probe setup and B) the FSSP-100 and the CDP setups as they were installed on Sammaltunturi station during PACE2013.

Table 1. Instruments, wavelengths (nm), sampling area (mm), number of bins, probe air speed (ms^{-1}), size range (μm), time resolution (s), operation starting and ending date are presented.

Instrument	Laser Wavelength (nm)	Sampling area (mm^2)	Number of bins	Probe air speed (m/s)	Size range (μm)	Time resolution (s)	Operation start date	Operation end date
CAPS	680	0.24	30	17-23	0.61 - 50	1	15 Oct	28 Nov
FSSP	633	0.414	40	10	1.2 - 47	1	14 Sept	28 Nov
CDP	658	0.3	30	14	3 - 50	1	25 Sept	28 Nov

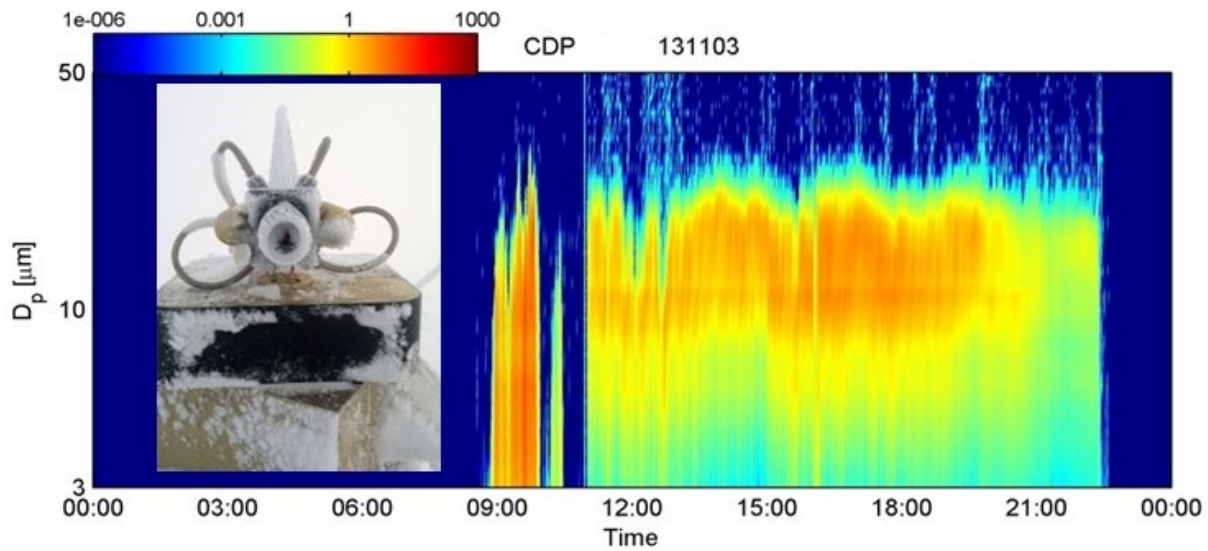



Figure 2. An example case of the CDP probe where its rotational platform and inlet are frozen. The size distribution of the CDP probe at 03.11.2013 is depicted. The instrument was out of order from 00:00 to about 11:30 and CDP cleaning procedure was needed to start operation again.

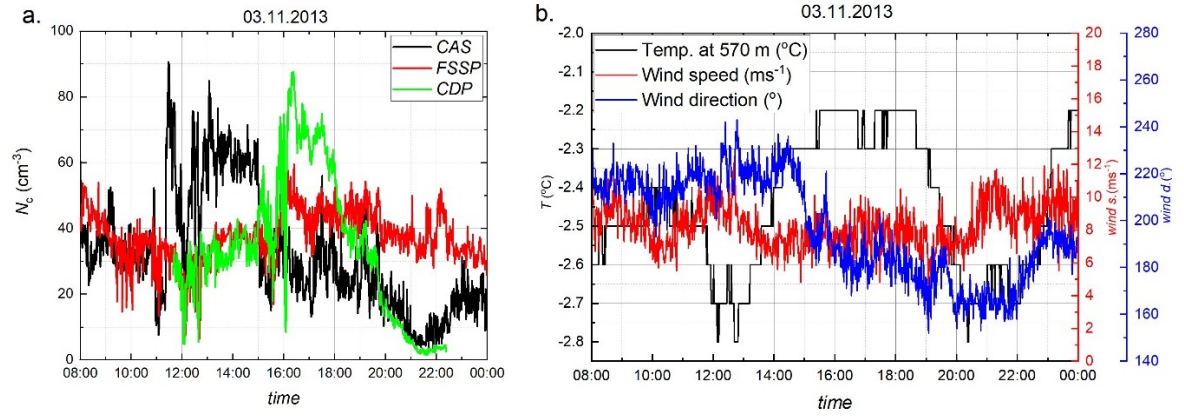



Figure 3. Time series of 1 minute averages during 03.11.2013; (a)N_c of each cloud probe along with (b) temperature, wind speed and wind direction are presented. This is a typical example of the cloud probes accumulating ice. From 12:00 we can see drop in N_c of the CAS. The sudden increase just before was a clear sign of probes inlet freezing. The same behavior could be also seen for the CDP after 16:00. When ice was accumulated, the opening of the probe inlet was getting smaller which resulted in a raising PAS. During deriving N_c to evaluate our data set, we considered the PAS was constant. The underestimation of the PAS explains the abnormality in N_c.

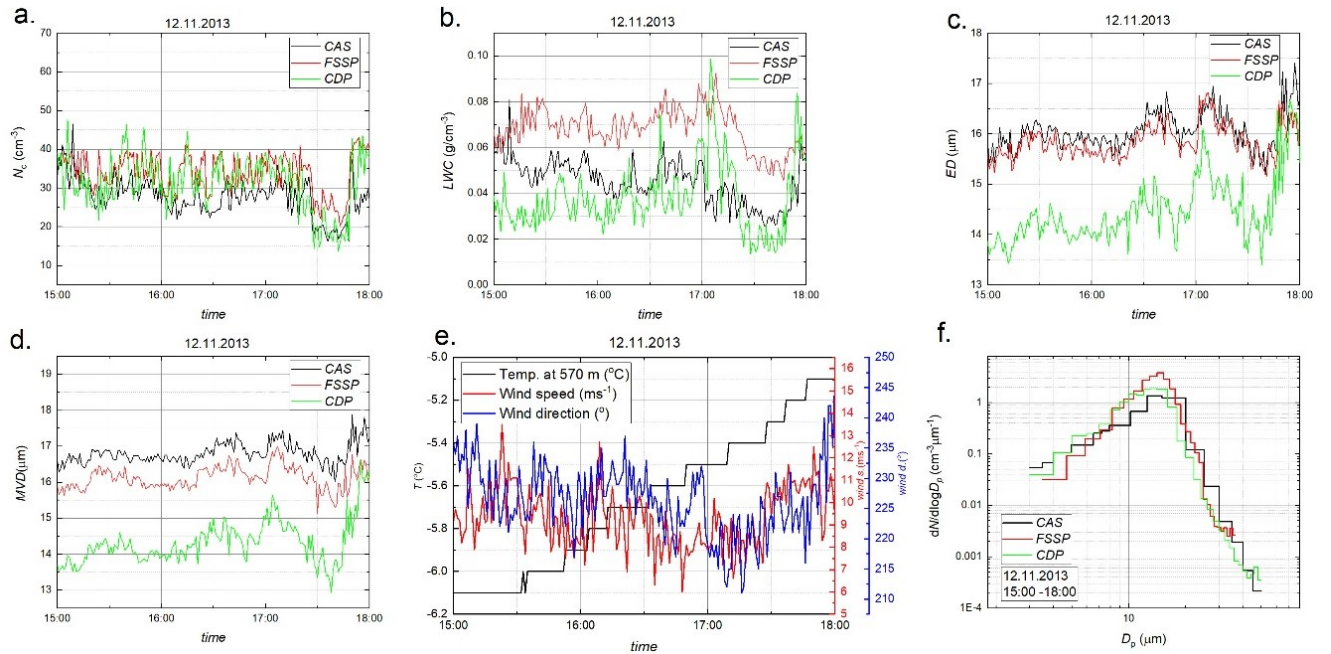



Figure 4. Time series on 12 November case from 15:00 to 18:00, the main parameters as they were derived /measured from all three cloud probes: (a) N_c; (b) LWC; (c) ED; (d) MVD; (e) temperature, wind speed, wind direction and (f) size distribution. All three instruments were pointing to the same direction. This resulted high agreement in N_c for all three probes. In addition, we also noticed good agreement in LWC. The main reason for slight differences in ED and MVD was the different sizing ability of the probes.

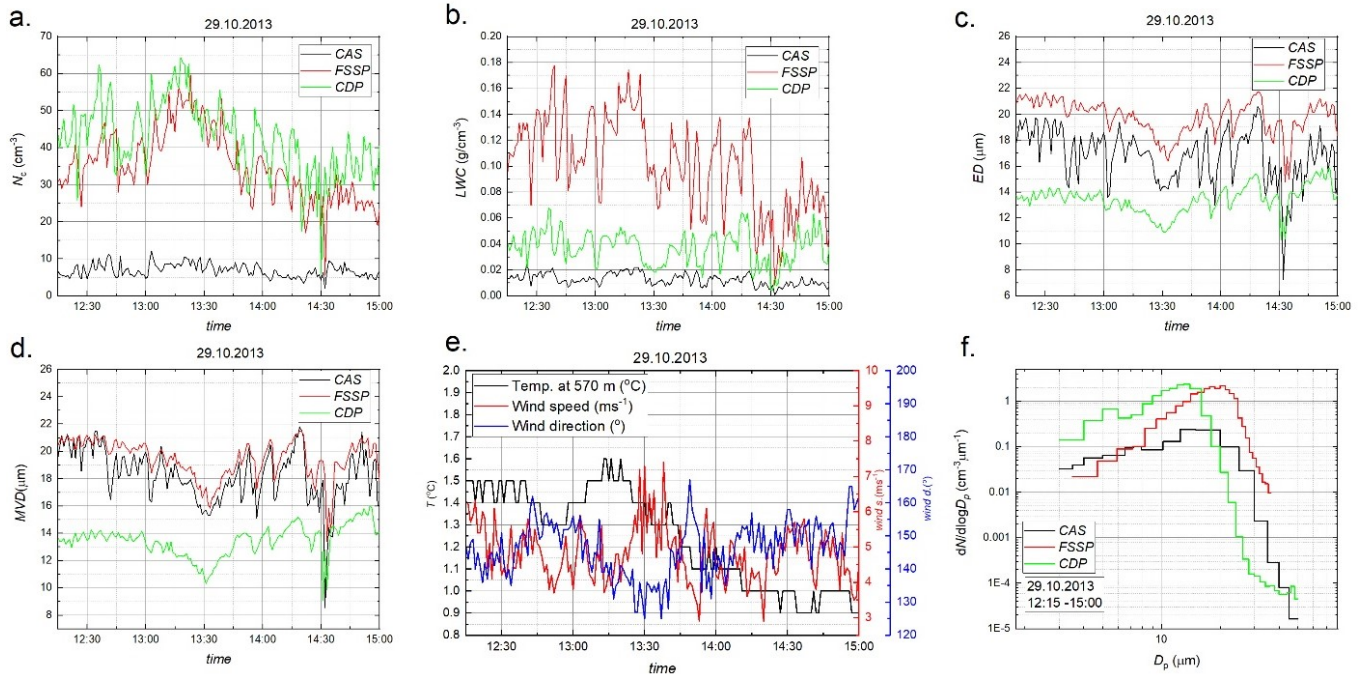



Figure 5. Time series on 29 October from 12:15 to 15:00, the main parameters as they were derived /measured from all three cloud probes: (a) N_c; (b) LWC; (c) ED; (d) MVD; (e) temperature, wind speed, wind direction and (f) size distribution. The wind direction was perpendicular to the CAS probe. This resulted in CAS significantly underestimated N_c and LWC. The main reason for slight differences in ED and MVD was the different sizing ability of the probes.

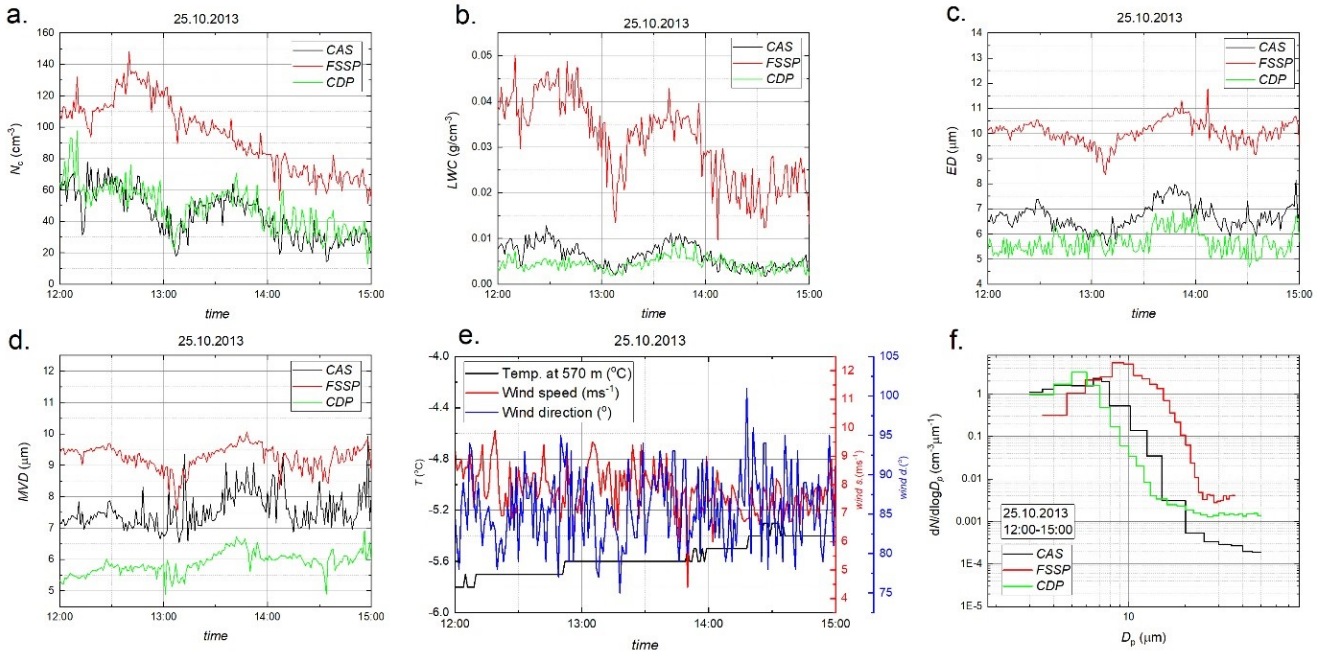



Figure 6. Time series on 25 October from 12:00 to 15:00, the main parameters as they were derived /measured from all three cloud probes: (a) N_c; (b) *LWC*; (c) *ED*; (d) *MVD*; (e) temperature, wind speed, wind direction and (f) size distribution. During this case the probes were not facing the same direction. Agreement in N_c between CAS and CDP indicates that CDP rotation system was frozen and CAS not facing the wind. Large discrepancies were observed in *LWC* and slight discrepancies in *ED* and *MVD*. The main reason for those discrepancies was the different sizing ability of the probes.

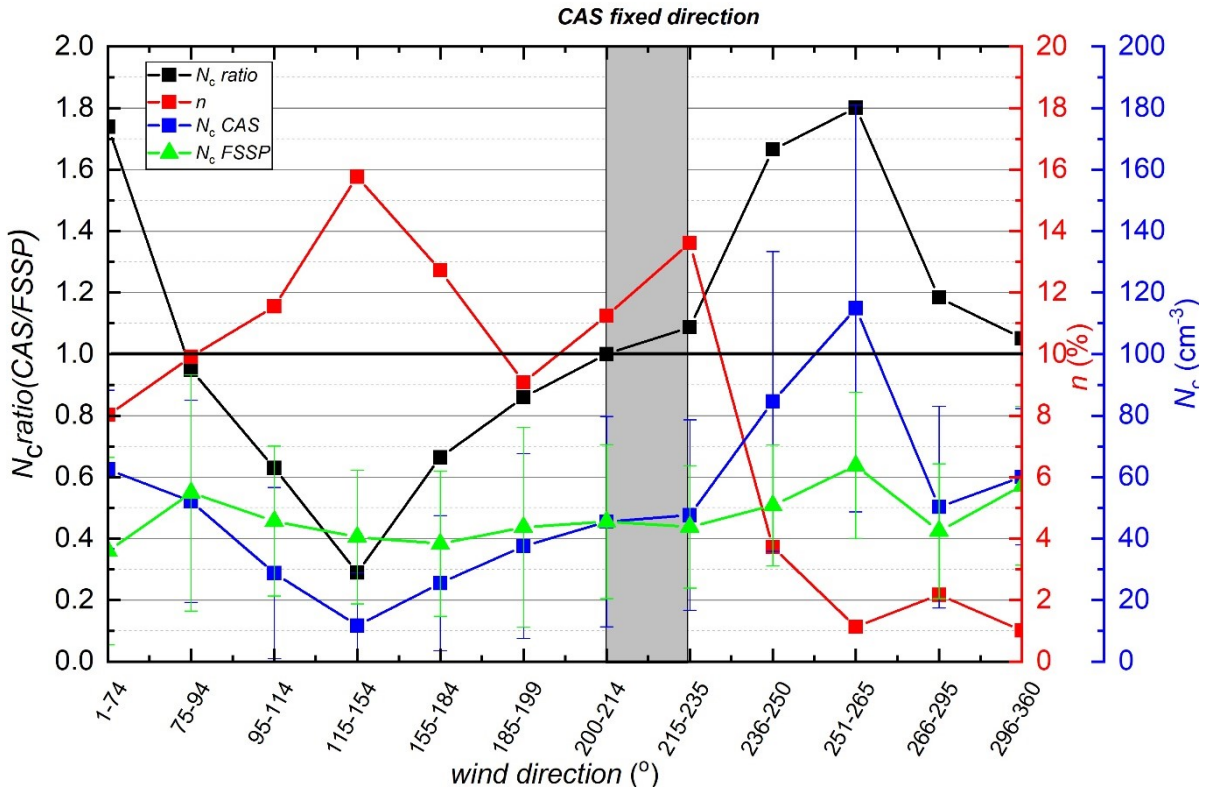



Figure 7. Number concentration ratio ($N_{c\ ratio}$), number of observations (n), and N_c of the CAS probe and the FSSP for each part of the 360° wind rose as it was divided for detailed investigation. The grey rectangle corresponds to wind iso-axial measurements.

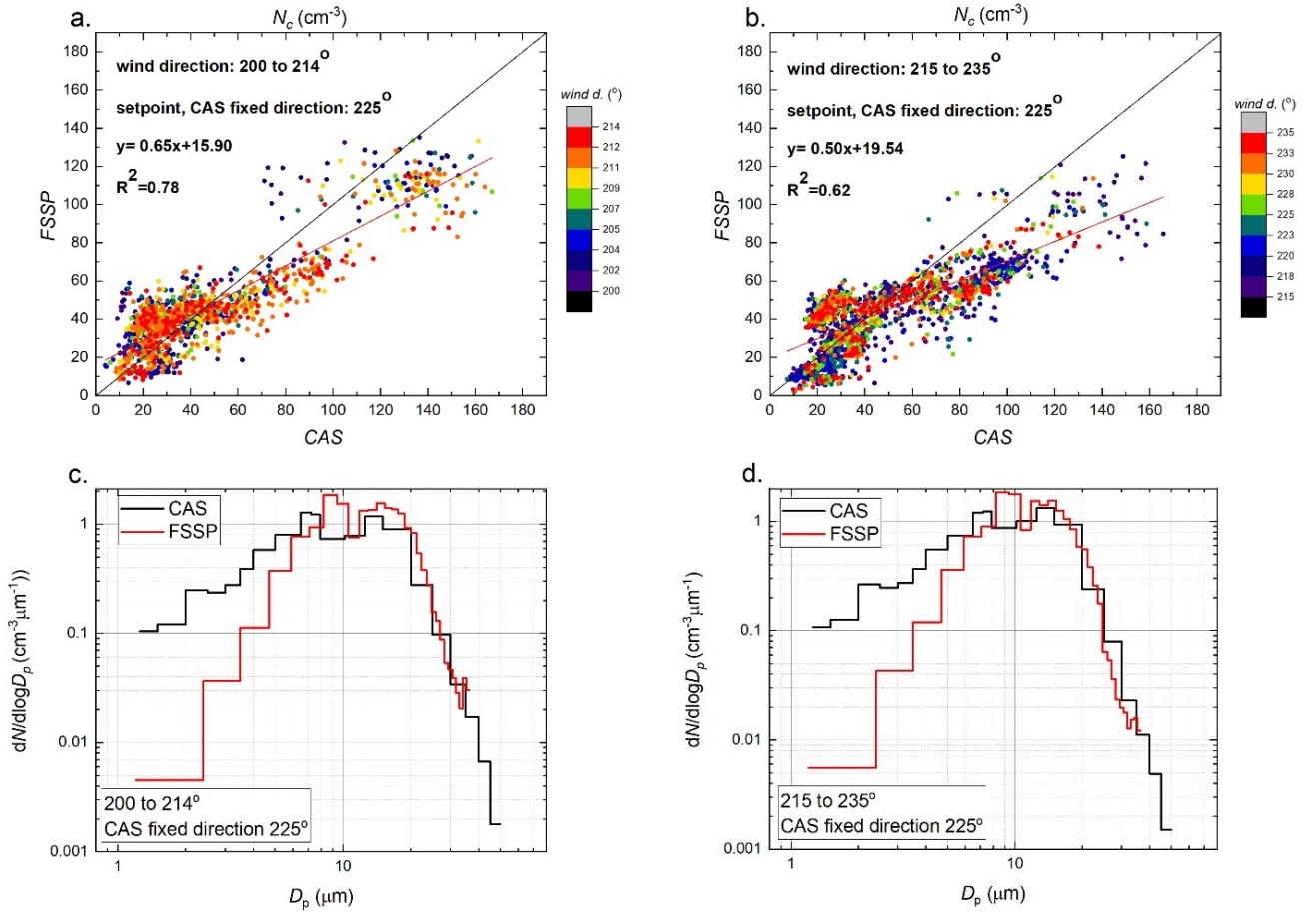



Figure 8. Comparison of number concentration (N_c) as it was derived from the CAS and the FSSP is presented for two sectors of the wind rose during the station was inside a cloud (a) 200 to 214°; (b) 215 to 235°; Color code represent the wind direction, (c) and (d) the size distribution as it was measured from the CAS probe and the FSSP for the same two wind sectors is presented.

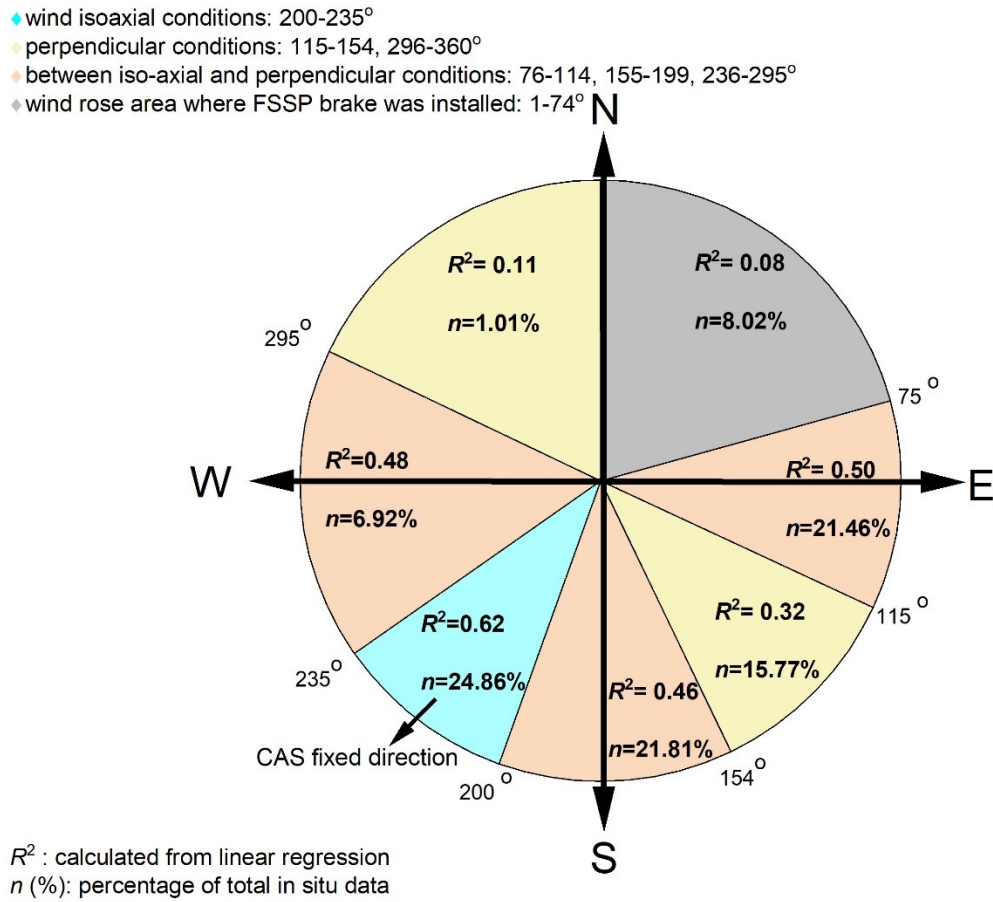



Figure 9. Inter-comparison of number concentration (N_c) between the CAS and the FSSP based on the wind direction. The CAS was set and installed in a fixed direction (southwest, ~ 225°); the FSSP was installed on a rotating platform and following the wind direction. The wind rose was separated to four representative wind direction conditions; wind iso-axial conditions (from 200 to 235°), perpendicular conditions (from 115 to 154 and 296 to 360°), conditions between iso-axial and perpendicular (from 76 to 114, from 155 to 199 and from 236 to 295°) and the special case where the brake influenced the performance of FSSP (from 1 to 74°).

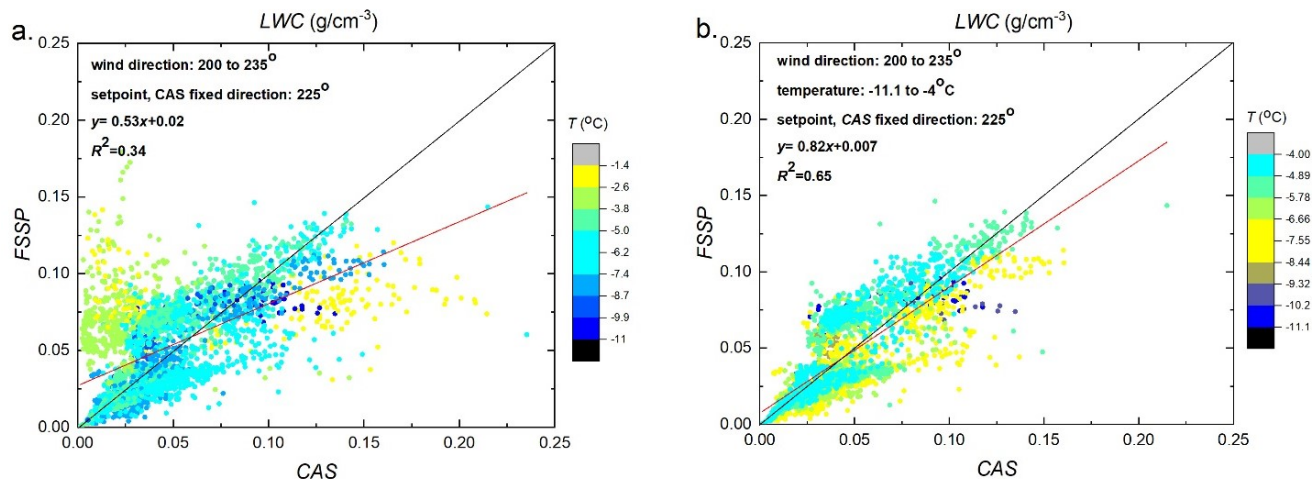



Figure 10. (a) Comparison of *LWC* as it was derived from the CAS and FSSP for wind iso-axial conditions, color code represents full temperature range from -11.1 to -1.4 °C (b) Comparison of *LWC* as it was derived from the CAS and FSSP is presented for the same conditions but only for temperature range from -11.1 to -4 °C.

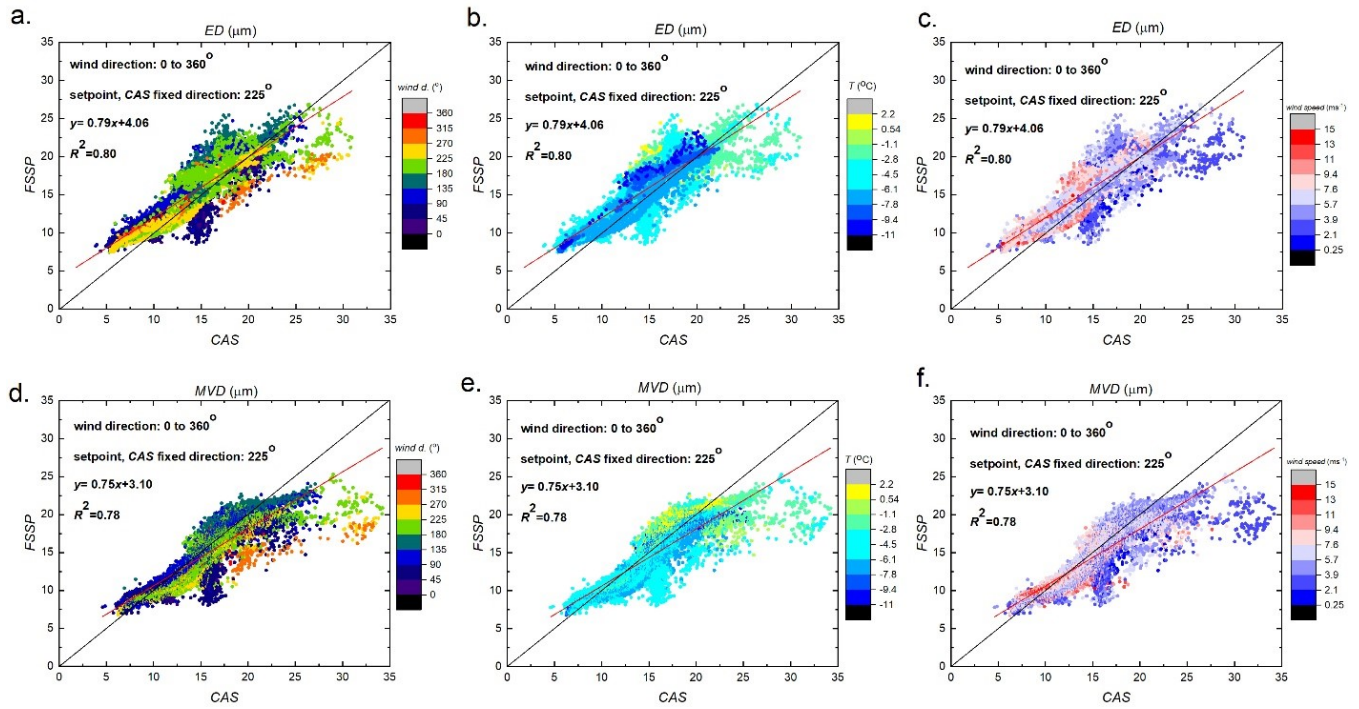



Figure 11. Comparison of *ED* (a) (b) (c) and *MVD* (d) (e) (f) as it was derived from the CAS and FSSP is presented for all wind directions. Color code represents (a) and (d) wind direction; (b) and (e) temperature; (c) and (f) wind speed.

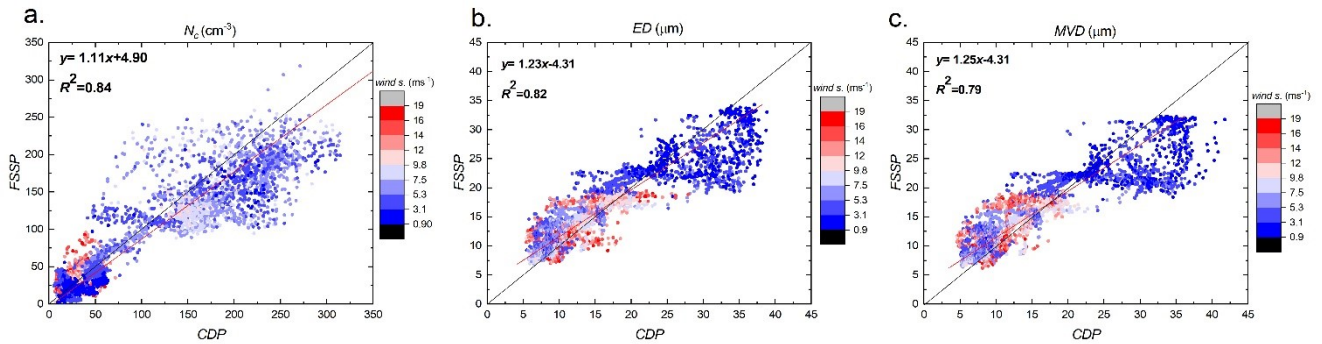



Figure 12. Inter-comparison of (a) N_c; (b) *ED* and (c) *MVD* as it was derived from the CDP and the FSSP is presented for all wind directions. Color code represents the wind speed.

Table 2 presents the final recommendations for data analysis regarding the cloud spectrometers ground based setups for future campaigns in sub-arctic conditions with frequently occurring supercooled clouds.

	N_c	ED, MVD	LWC	Comments
CAS	Only data from wind iso-axial conditions. (± 20° from its fixed direction) should be used	All measurements can be used for further analysis, independent on wind direction in size range of ED and MVD 5 - 35 µm	Only data from wind iso-axial conditions and temperatures below – 4 ºC should be used	Good data coverage (67%), Operating properly both in non-icing and icing conditions, needs daily cleaning
FSSP	Data from all wind sectors will be used except data from wind sector where brake was installed (± 40° brake direction)	All data can be used for further analysis except data from wind sector where brake was installed (± 40° from brake direction)	Only data from wind iso-axial conditions and temperatures below – 4 ºC should be used	The best data coverage (75%), Operating properly both in non-icing and icing conditions, needs daily cleaning
CDP	Usable in warm clouds. Limitation in temperature, operational problems at sub-zero temperatures	All data obtained in non-icing conditions can be used for further analysis	Not usable due to temperature range.	Low data coverage (17%), Operating properly in non-icing conditions, not recommended for sub-zero temperatures