Highly time-resolved characterization of carbonaceous aerosols using a two-wavelength
 Sunset thermal-optical carbon analyzer

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22 Abstract

Carbonaceous aerosols have great influence on the air quality, human health and climate 23 change. Except for organic carbon (OC) and elemental carbon (EC), brown carbon (BrC) mainly 24 originates from biomass burning, as a group of OC with strong absorption from the visible to near-25 26 ultraviolet wavelengths and makes a considerable contribution to global warming. Large amounts of studies have reported long-term observation of OC and EC concentrations throughout the word, 27 but studies of BrC based on long-term observations are rather limited. In this study, we established 28 a two-wavelength method (658 nm and 405 nm) applied in the Sunset thermal-optical carbon 29 analyzer. Based on a one-year observation, we firstly investigated the characteristics, 30 meteorological impact and transport process of OC and EC. Since BrC absorbs light at 405 nm 31

more effectively than 658 nm, we defined the enhanced concentrations (dEC = $EC_{405 nm}$ - $EC_{658 nm}$) 32 and gave the possibility to provide an indicator of BrC. The receptor model and MODIS fire 33 information were used to identify the presence of BrC aerosols. Our results showed that the 34 carbonaceous aerosol concentrations were highest in winter and lowest in summer. Traffic 35 emission was an important source of carbonaceous aerosols in Nanjing. Receptor model results 36 showed that strong local emissions were found for OC and EC; however, dEC was significantly 37 affected by regional or long-range transport. The dEC/OC and OC/EC ratios showed similar 38 diurnal patterns and the dEC/OC increased when the OC/EC ratios increased, indicating strong 39 secondary sources or biomass burning contributions to dEC. Two biomass burning events both in 40 summer and winter were analyzed and the results showed that the dEC concentrations were 41 obviously higher in biomass burning days; however, no similar levels of the OC and EC 42 43 concentrations were found both in biomass burning days and normal days in summer, suggesting that biomass burning emission made a great contribution to dEC and the sources of OC and EC 44 were more complicated. Large number of open fire counts from the northwest and southwest areas 45 of the study site were observed in winter and significantly contributed to OC, EC and dEC. In 46 47 addition, the near-by Yangtze River Delta area was one of the main potential source areas of dEC, suggesting that anthropogenic emissions could also be important sources of dEC. The results 48 49 proved that dEC can be an indicator of BrC in biomass burning days. Our modified twowavelength instrument provided more information than the traditional single-wavelength thermal-50 51 optical carbon analyzer and gave a new idea about the measurement of BrC; the application of dEC data needs to be further investigated. 52

531. Introduction

Carbonaceous aerosols including organic carbon (OC) and elemental carbon (EC), which 54 55 have significant influence on the global radiative transfer, human health and atmospheric visibility, 56 have been the focus of research in the atmospheric environment field for many years (Lelieveld et al., 2015; Wu and Yu, 2016; Wang et al., 2018; Zhang et al., 2017; Liu et al., 2019; Zhang et al., 57 2019). EC mainly originates from fossil fuel and biomass combustion and is estimated to be the 58 second largest warming factor behind CO₂ contributing to climate change (Liu et al., 2015; Zhang 59 and Kang, 2019; Cao and Zhang, 2015). OC originates both from primary emissions and gas-to-60 particle conversion as secondary organic carbon (SOC) and can scatter the solar radiation which 61 causes negative forcing globally (Zhou et al., 2014; Huang et al., 2014). 62

In the recent decades, brown carbon (BrC), as a kind of light-absorbing organic carbon which 63 can absorb light especially from near-UV to visible wavelength, has caused global concern due to 64 its positive climate effect (Andreae and Gelencsér, 2006; Zhang et al., 2020). BrC is mainly 65 emitted from anthropogenic and biogenic emissions (Zhang et al., 2011). Previous studies have 66 proved that biomass burning and biofuel combustion are the most important sources of primary 67 BrC (Saleh et al., 2014; Wu et al., 2020; Lei et al., 2018). Recent researches reported that in 68 developing countries such as China and India, the contribution of fossil fuel combustion to BrC 69 cannot be ignored (Satish et al., 2017; Yan et al., 2017; Kirillova et al., 2014). Secondary BrC is 70 mainly produced by heterogeneous photo-oxidation reactions or aqueous reactions of 71 anthropogenic and biogenic precursors (Zhang et al., 2020; Li et al., 2020; Zhang et al., 2011). 72 However, due to the lack of understanding of BrC at the molecular level and in situ BrC data, there 73 74 are still large uncertainties in the estimates of the distribution and the magnitude of the BrC climate effect in both remote sensing and modeling (Arola et al., 2011; Feng et al., 2013). 75

The thermal-optical analysis (TOA) method is one of the most widely used quantitative 76 methods for OC and EC taking use of the difference between the thermal-optical properties of OC 77 78 and EC (Birch and Cary, 1996; Chow et al., 2004). OC and EC will be volatilized at different heating protocol. The reflectance/transmittance of one laser source (near-infrared wavelength) 79 80 through the sample filter are continuously monitored and the return of the reflectance/transmittance to its initial value on the thermograph is taken as the split point between 81 82 OC and EC. This way, the formation of pyrolyzed carbon which can also absorb the light and make the sample darker, is corrected. This method has been widely used in studies employing the NIOSH 83 protocol or IMPROVE A protocol (Ji et al., 2016; Chow et al., 2007). However, the thermal-84 optical approach assumed that EC is the only light-absorbing species, the presence of BrC, which 85 86 is part of OC but also a light-absorbing component, shifts this separation towards EC, resulting in 87 overestimated EC values and underestimated OC values (Chen et al., 2015; Birch and Cary, 1996). Sandradewi et al. (2008) pointed out that light absorption measurements at different wavelength 88 by the aethalometer can be used to quantify the contributions of wood combustion and traffic 89 emissions to aerosols since wood smoke contains organic compounds which enhance the light 90 absorption in the ultraviolet wavelength. But traffic emissions produce more black carbon (BC), 91 which dominates the light absorption in the near-infrared wavelength. They took use of 92 aethalometer data measured at 470 nm and 950 nm to quantify the BC distinction between wood 93

burning and traffic emission. With the similar principle, Wang et al. (2011) used a two-wavelength 94 Aethalometer (370 and 880 nm) to identify the presence of residential wood combustion (RWC) 95 particles which was closely associated with BrC. Organic components of wood smoke particles 96 absorb light at 370 nm more effectively than 880 nm in two-wavelength aethalometer 97 measurements. They believed that the enhanced absorption (Delta-C=BC_{370nm}-BC_{880nm}) can serve 98 as an indicator of RWC particles. This method was further used by Wang et al. (2012a; 2012b). 99 Chen et al. (2015) used a modified seven-wavelength thermal-optical transmittance/thermal-100 optical reflectance (TOT/TOR) instrument (Thermal Spectral Analysis - TSA) allowing the 101 determination of the OC-EC split at different wavelengths and light absorption measurements to 102 be made with wavelength-specific loading corrections, providing additional information including 103 the optical properties of black carbon (BC) and BrC from the IR to UV parts of the solar spectrum 104 105 and their contributions. Massabò et al. (2016) further corrected the OC/EC split point using the Multi-Wavelength Absorbance Analyzer (MWAA) which provides the aerosol absorbance values 106 107 at five wavelengths from IR to UV together with a Sunset OC/EC analyzer to achieve the BrC concentration. With a set of samples collected during wintertime in the Ligurian Apennines in 108 109 Italy, clear correlations were found between the BrC and levoglucosan mass concentration. A further step of BrC quantification taking use of TSA was reported by Chow et al. (2018), further 110 111 proving that the use of seven wavelengths in thermal-optical carbon analysis allows contributions from biomass burning and secondary organic aerosols to be estimated. Their results clearly 112 113 demonstrated the role of BrC in the thermal-optical analysis. However, these techniques focus on the light absorption measurement of BrC and are still limited reported in previous researches, 114 though they provide quartz-fiber filter samples that are currently being characterized for OC and 115 EC by thermal-optical analysis. These methods mentioned above still cannot achieve the 116 117 observation of long-term real-time BrC mass concentrations.

Since the establishment of the TOT method by the Sunset Laboratory, the Sunset OC/EC instrument, as part of the Chemical Speciation Network (CSN), with over 100 monitors across the United States over 15 years, offering long-term measurement of OC and EC concentrations, has been widely used in the United States and throughout the world providing important in-situ data of OC and EC (U.S.EPA, 2019; Birch and Cary, 1996). This instrument had been designed with a tuned diode laser (red 660 nm) to correct the formation of pyrolyzed carbon. In this study, we modified the Sunset instrument to a two-wavelength (658 nm and 405 nm) Sunset carbon analyzer

by adding one more violet diode laser at λ =405 nm. The violet diode laser together with the red 125 diode laser focus through the sample chamber, then the laser beam passes through the filter to 126 correct for the pyrolysis-induced error. Previous work reported by Chen et al. (2015) as mentioned 127 above was integrating the optical instrument like the aethalometer to the traditional OC/EC 128 analyzer; in this way, they provided the light absorption contributions of BC and BrC. The 129 enhanced carbon analyzer provided new insight into more accurate OC and EC measurements. 130 Their work was conducted in offline mode; based on their work, our instrument can get the real-131 time OC and EC mass concentrations both at 658 nm and 405 nm. BrC particles absorb light at 132 405 nm more effectively than 658 nm in the two-wavelength Sunset carbon measurements. We 133 define dEC=EC_{405 nm}-EC_{658 nm} and hope it can be an indicator of BrC aerosols so that we can divide 134 real-time BrC mass concentration measurement from the two-wavelength measurement. 135

Nanjing, as one of the largest cities in the Yangzi River Delta region, represents a heavy industry
area with a dense population. In addition, due to its topography, Nanjing is very sensitive to
regional transport of air masses from its surrounding areas. OC, EC and dEC were measured from
June 2015 to July 2016 at Nanjing University of Information Science and Technology (NUIST).
Based on the abundant data, together with MODIS fire information, we can analyze the temporal
variation, transport processes and sources of carbonaceous aerosols in North Nanjing and evaluate
the biomass burning impact on dEC, which can be the scientific basis of pollution control policy.

1432. Methods

144 **2.1 Study site**

In this study, the sampling site is located at Nanjing University of Information Science and 145 Technology (NUIST) in the North Suburb of Nanjing (32°207'N, 118°717'E). The study site 146 is surrounded by housing and industrial areas. Many chemical enterprises, for example, Yangzi 147 148 Petrochemical, Nanjing Chemical Industry and Nanjing Iron and Steel Group are located at the 149 northeast of the study region, which produces exhaust with large amounts of aerosol particles. The study site is adjacent to a heavily trafficked road (Ningliu Road) located near the site, 150 approximately 600 m to the east. Therefore, this region has intense human activities, industrial 151 emissions and heavy traffic flow. 152

153 **2.2 Two-wavelength TOT measurement**

Hourly concentrations of OC and EC in PM_{2.5} were sampled and measured by a semicontinuous carbon analyzer (Model-4, Sunset Lab, USA). Air samples were collected continuously with a sample flow of ~ 8 L/min through a PM_{2.5} cyclone. The collection time was set at 45 min for each cycle. The airstream passed through a parallel plate organic denuder to reduce the effect of volatile organic compounds and finally deposited on a quartz filter with a diameter of ~ 17 mm.

After a sample was collected, OC and EC were determined using the TOT method by applying 159 a slightly modified NIOSH 5040 protocol. The details of the heating setup are shown in Table S1. 160 Figure 1 shows the structure and operational principle of the instrument. Briefly, it consists of two-161 stages: the oven is first purged with helium and the oven temperature increased in a stepped ramp 162 to 840°C, OC is volatilized in this stage. Then the oven temperature is kept at 840°C for a while 163 and goes down to 550°C. In the second stage, EC is volatilized in a second temperature ramp to 164 850°C while purging the oven with a mixture containing 2% oxygen and 98% helium. The 165 pyrolysis products are converted to carbon dioxide (CO₂) which is quantified using a self-166

167 contained nondispersive infrared (NDIR) system.

Also, in this study, we used a two-diode lasers (658 nm and 405 nm) equipped Sunset analyzer; 168 thus mass concentrations of OC and EC at different wavelengths can be measured with the 2-lasers 169 system. The split point between OC and EC is detected automatically by the RTCalc731 software 170 171 provided by Sunset Lab. The principle is the same as for the traditional Sunset carbon analyzer (Birch and Cary, 1996). An example thermogram of sample analysis using the two-wavelength 172 Sunset semi-continuous carbon analyzer is shown in Fig. 2. During the sample analysis, the laser 173 beam at 658 nm and 405 nm are both sent through the filter and the transmitted light signal is 174 monitored to correct the undesired formation of pyrolyzed carbon (PyrC) and then to determine 175 176 the split point of OC and EC at both wavelengths. BrC aerosols absorb light at 405 nm more significantly than 658 nm in the 2-lasers system. Due to the strong absorption of BrC at the near-177 ultraviolet wavelength, the enhanced absorption at 405 nm can serve as an indicator of BrC 178 aerosols (Liu et al., 2015). We define dEC data as the difference of EC concentrations at two 179 180 wavelengths (dEC=EC_{405nm}-EC_{658nm}) to identify the presence of BrC aerosols. Our study provides a one-year measurement of dEC mass concentrations. Besides, OC and EC represent the OC and 181 EC concentrations at 658 nm in this paper without a special explanation. 182

183 At the end of each analysis, a fixed volume of an internal standard containing 5% methane 184 and 95% helium is injected and thus a known carbon mass can be derived. The external sucrose 185 standard ($4.207 \ \mu g \ \mu L^{-1}$) calibration was conducted every week to insure repeatable quantification. 186 Calibration with an instrument blank was conducted every day. Both detection limit for OC and

EC of the instrument was 0.5 μ g m⁻³. We also did the measurements of OC and EC in PM_{2.5} filter 187 samples using the same method followed by the NIOSH protocol. All data were corrected to blank 188 measurement before comparison. Figure S1 shows the correlations between the real-time OC, EC 189 concentrations and sampling OC, EC concentrations at the same time. The results showed that the 190 online and offline data during the corresponding periods had good correlations with R² of 0.8 for 191 OC, R² of 0.4 for EC and R² of 0.8 for TC. In order to evaluate the impact of PyrC, we calculated 192 the PyrC at 658 nm fraction of dEC and the average PyrC/dEC was 4.4%, indicating the little 193 influence of PyrC. 194

195 **2.3 Test of the new dEC data**

To evaluate the new dEC data, parallel BC concentrations were measured with a sevenwavelength Aethalometer with dEC concentrations in December, 2019. Radiation attenuation of an aerosol deposition on a filter (ATN $_{\lambda}$) is determined by the Beer-Lambert law:

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$$ATN_{\lambda} = ln \frac{I_{0,\lambda}}{I_{\lambda}}$$
 (1)

Where $I_{0, \lambda}$ and I_{λ} are the measured wavelength-specific laser reflectance signals. ATN_{λ} is used to calculate the attenuation coefficient with Eq. (2):

$$202 \qquad b_{ATN} = \frac{A}{V} \tag{2}$$

Where A is the filter area and V is the sampled air volume. Then a simplified two-component model is used to calculate the contribution of light attenuation to both BC and BrC (Chow et al., 2018; Chen et al., 2015; Sandradewi et al., 2008; Hareley et al., 2008):

206
$$b_{ATN}(\lambda) = q_{BC} \times \lambda^{-AAE_{BC}} + q_{BrC} \times \lambda^{-AAE_{BrC}}$$
 (3)

Where q_{BC} and q_{BrC} are fitting coefficients, AAE is the absorption Ångström exponent which 207 represents the wavelength-dependent characteristics of light absorption capability of aerosols. The 208 AAE of BC was assumed to be 1. Fitting coefficients in Eq. (3) were obtained for potential AAE_{BrC} 209 between 1 and 8 by least squares linear regression and the AAE_{BrC} leading to the overall best fit in 210 terms of r^2 is selected as the effective AAE_{BrC}. Using these fitting coefficients, the b_{ATN} due to BC 211 and BrC are calculated at each wavelength. Figure S2 shows that the fitted b_{ATN} at 405 nm are 212 within ± 5 % of the measured values for $b_{ATN} > 0.01$. Figure 3 shows the relationship between the 213 b_{ATN} due to BrC at 405 nm and the dEC. Good correlation between them is found with R square 214 215 of 0.64, indicating that dEC was associated with BrC.

216 **2.4 Sampling**

217 2.4.1 Real-time PM_{2.5} observation

The real-time $PM_{2.5}$ concentrations were measured through the Tapered Element Oscillating Microbalance (TEOM) method (TEOM1405-DF, Thermo Scientific, America) from August, 2015 to July, 2016. The resolution of the measured data was 6 min. The instrumental operation maintenance, data assurance and quality control were performed according to the Chinese Ministry of Environmental Protection Standards for PM_{10} and $PM_{2.5}$ which was named "HJ 653-2013" (Zhang and Cao, 2015b).

224 2.4.2 Sample collections

PM_{2.5} in the atmosphere was collected on 8*10 inch prebaked quartz fiber filters (QFF, PALL, America) by a high volume air sampler (KC-1000, Qingdao, China) at a flow rate of 999 L min⁻¹ in four months: 4 June to 18 June, 6 October to 2 November and 10 December to 31 December in 2015, 10 May to 31 May in 2016. Sampling started and ended at around 8:00 and 20:00 every day; each sample was collected for 12 hours. A total of 148 samples were collected including four field blanks in the four seasons using 10 min exposure to ambient air without active sampling.

All QFFs were pre-baked at 450 °C for 6 h before sampling to remove residual carbon. Before and after sampling, all QFFs were weighed with an electronic balance (Sartorius, 0.1 mg, Germany). After weighting, the filters were wrapped in aluminum foils, packed in air-tight polyethylene bags and stored at -20°C until further analysis. All procedures during handling of filters were strictly quality controlled to avoid any possible contamination.

236 **2.5 Identification of potential regional sources**

The Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSLPIT4.8) model, provided 237 by the National Oceanic and Atmospheric Administration (NOAA), was used to investigate the air 238 mass origins of carbonaceous aerosols. The 48-hour back trajectories at Nanjing (32.2°N, 118.7°E) 239 240 were calculated every hour (Draxler and Hess, 1998; Rolph et al., 2017; Cohen et al., 2015). In 241 order to evaluate the behavior of the air mass circulation in the planetary boundary layer (PBL), the trajectories at 500 m corresponding to the upper-middle height of the PBL were calculated, 242 representing well-mixed convective boundary layer for regional transport investigation (Xu and 243 Akhtar, 2010). The National Center for Environmental Prediction Global Data Assimilation 244 System (NCEP GDAS) data obtained from NOAA with a spatial resolution of 1° ×1° and 24 levels 245 of the vertical resolution were used as meteorological data input to the model. The Potential Source 246 Contribution Function (PSCF) model was usually applied to localize the potential sources of 247

- pollutants. The details about the setup of the model can be found in Bao et al. (2017).
- 249 **3. Results and discussion**
- 250 **3.1** Characteristics of carbonaceous aerosols
- 251 3.1.1 Concentrations of carbonaceous aerosols

The statistics for the PM_{2.5}, OC, EC and dEC mass concentrations at the NUIST site are 252 summarized in Table 1. The hourly OC concentrations ranged from 0.5 to 45.8 µg m⁻³ (average of 253 $8.9 \pm 5.5 \ \mu g \ m^{-3}$), and the EC concentrations ranged from 0.0 to 17.6 $\mu g \ m^{-3}$ (average of 3.1 ± 2.0 254 μ g m⁻³). The results are comparable to those reported by Chen et al. (2017) in the Xianlin Campus 255 of Nanjing University (5.7 µg m⁻³ for OC and 3.2 µg m⁻³ for EC), which site was located in the 256 southeast suburb of Nanjing and close to the G25 highway and were also affected by traffic sources. 257 The higher OC concentrations in this study are probably due to the around chemical enterprise 258 259 emissions. The average contributions of OC and EC to the total measured PM_{2.5} mass was 12.8% and 4.3%, respectively, suggesting that carbonaceous fraction made an important contribution to 260 fine particulate matter. The average dEC mass concentration was 0.8 µg m⁻³ contributing 10.0% to 261 OC, 22.3% to EC and 1.3% to the PM_{2.5} concentrations with maximum concentration of 8.1 μ g m⁻ 262 ³ contributing 48.2% to OC, 97.8% to EC and 17.6% to total PM_{2.5} concentrations. This 263 information can be further applied in the PMF analysis to evaluate the sources of the carbonaceous 264 265 aerosols (Zhu et al., 2014; Sahu et al., 2011; Yan et al., 2019).

Compared with carbonaceous aerosol levels in other cities (Table S2), the OC and EC 266 267 concentrations in Nanjing were generally lower than those observed in urban sites such as Beijing and Shanghai and inland cities like Chengdu and Chongqing which are affected by the basin terrain 268 characteristics with static wind and unfavorable diffusion conditions, but higher than those 269 observed in the southern coastal cities such as Guangzhou, which is a megacity in China. It could 270 271 be explained since the site in Guangzhou was a rural site. In general, the level of carbonaceous 272 aerosol concentrations in China is higher than that in developed countries in the United States and Europe and lower than that in developing countries like India, though the sampling period in India 273 was from late autumn to winter and the much higher concentrations in India indicated the heavy 274 pollution level. The average OC/EC ratio in this study was 3.6, which is lower than most of those 275 reported in other studies, indicating the important impact of vehicle emissions at our study site. 276

Figure 4 shows the mass fractions of hourly carbonaceous aerosols and OC/EC ratios at different PM_{2.5} concentration intervals during the study period. During that period, 84.2% of the

PM_{2.5} samples exceeded the daily averaged Chinese national ambient air quality standard (NAAQS) 279 of 35.0 μ g m⁻³ for the first grade and 40.1% of the total samples exceeded the NAAQS of 75.0 μ g 280 m⁻³ for the second grade, reflecting heavy aerosol pollution in the study area. Generally, the 281 fractions of carbonaceous components decreased with increasing PM2.5 pollution level. A larger 282 mass fraction (about 32.3%) of carbonaceous aerosols in PM2.5 was found for relatively lower 283 $PM_{2.5}$ levels (0–20 µg m⁻³) compared to high $PM_{2.5}$ levels (300–500 µg m⁻³) with a carbonaceous 284 aerosol mass fraction of 5.2%. The results indicate that other components like secondary inorganic 285 aerosol (SIA) contribute more significantly to heavy haze events in Nanjing, which was also found 286 in other cities in the Yangtze River Delta area (Yang et al., 2011; Zhang and Zhang, 2019). The 287 contribution of dEC to OC decreased with the increase of PM2.5 concentrations between 0-200 µg 288 m⁻³, and then increased with the increase of PM_{2.5} concentrations between 200-500 μ g m⁻³. The 289 dEC contributed most significantly to OC of 14.3% for PM_{2.5} concentrations below 20 µg m⁻³. A 290 similar trend was found for the OC/EC ratios which showed a sharp increase along with enhanced 291 $PM_{2.5}$ level above 150 µg m⁻³. Previous studies have reported that high OC/EC ratios were related 292 to SOC formation or biomass burning emissions whereas low OC/EC ratios were related to vehicle 293 294 exhaust (Wang et al., 2015). We divided the dEC/OC at different intervals of OC/EC ratios and found that the dEC/OC increased when the OC/EC ratios increased in the four seasons, indicating 295 296 strong secondary sources or biomass burning contributions to dEC during heavy pollution periods (Fig. S3). 297

298 3.1.2 Seasonal variations of carbonaceous aerosols

As shown in Fig. 5, the OC, EC, dEC concentrations and dEC/OC ratios showed similar 299 variations with highest in winter and lowest in summer. The average OC and EC concentration in 300 winter was ~1.4 times and 1.5 times higher than that in summer and the average dEC 301 302 concentrations and dEC/OC in winter were approximately 1.4 and 1.6 times higher than those in 303 summer (Table 1). High dEC/OC was found in January and February in winter, indicating strong influence of anthropogenic sources on dEC, such as coal combustion. In addition, we found strong 304 biomass burning activities in February, which significantly contributed to the high concentrations 305 of dEC in February; more details can be found in section 3.3. The seasonality of carbonaceous 306 species in PM_{2.5} was strongly influenced by seasonal variations in emission intensities and 307 meteorological parameters. Table S3 summarizes the meteorological parameters in the four 308 seasons during the study period. The high carbonaceous aerosol concentrations in winter were 309

mainly a result of relatively stable atmospheric conditions with low temperature, relative humidity and boundary layer on one hand, and on the other hand, increasing emissions from fossil-fuel combustion for heating from the chemical enterprises nearby. In summer, the higher boundary layer resulted in the dispersion of aerosols in the atmosphere, and the higher temperature promoted the partitioning of semi-volatile organic compounds (SVOCs) into the gaseous phase (Yang et al., 2011). In addition, large precipitation in summer (586 mm in total) favored the wet scavenging processes of aerosols.

The average OC/EC ratios in spring, summer, autumn and winter were 3.9, 4.0, 2.8 and 3.4, 317 respectively (Table 1). The OC/EC ratio could give some information about primary and secondary 318 organic carbon (Turpin and Huntzicker, 1995; Lim and Turpin, 2002). In summer, strong 319 convective activities in the atmospheric boundary layer and solar radiation, high temperature and 320 321 plenty of moisture in the atmosphere were favorable for the formation of SOC. On the other hand, the high OC/EC ratios in June in this study were also strongly related to biomass burning which 322 323 will be discussed in section 3.3. The lower ratios of OC to EC in autumn and winter indicate strong primary sources in these two seasons. It should be noted that the OC/EC ratios were a rough 324 325 indicator to estimate the primary and secondary organic carbon; further analysis of the formation of SOC needs to be conducted in the future (Pio et al., 2011; Wu and Yu, 2016). 326

327 3.1.3 Diurnal variation of carbonaceous aerosols

The diurnal pattern of carbonaceous aerosols can be affected by both meteorological 328 329 parameters and sources (Ji et al., 2016). Figure 6 depicts the diurnal variation of OC, EC, dEC, dEC/OC and OC/EC ratios during the study period. Clear diurnal variations were observed in OC 330 and EC. Both the OC and EC concentrations kept high levels at night and low levels in the daytime, 331 indicating the strong influence of the atmospheric boundary layer on air quality in northern 332 333 Nanjing. The peak occurred in the morning both for OC and EC, indicating the significant impact 334 of traffic sources on the OC and EC concentrations. The dEC/OC and OC/EC ratios showed similar trends in the daytime with gradually increase from morning till afternoon, indicating the 335 importance of the contribution of secondary sources to dEC. Similar though not so obvious diurnal 336 variations were found in dEC. It should be noted that the vehicle emissions and the boundary layer 337 height had no significant effect on the diurnal variation of dEC/OC, suggesting there were no 338 significant local sources of dEC. There was a small peak in dEC/OC at 3:00 am, which might be 339 related to the aqueous secondary organic aerosol formation during nighttime (Sullivan et al., 2016). 340

The relative humidity (RH) and Temperature (T) dependent distributions of OC, EC mass 341 concentrations and dEC/OC and OC/EC throughout the study period are shown in Fig. 7. High 342 dEC/OC (>30 %) can be found in three areas, first shown in the right area with relatively high T 343 at 25-40 °C and RH at 40-60 %, which were usually found in the summer afternoon which was 344 closely related to the strong formation of SOC. This distribution was also seen in OC/EC. The 345 second area is in the upper region with RH over 80 % and T at 10-20 °C and the third area appears 346 for RH below 30 % and T at about 10 °C, corresponding to nighttime and winter afternoon. In 347 general, dEC had no strong dependence on the RH and T distribution, indicating the complex 348 formation mechanism of dEC. OC and EC show similar distributions with the highest mass loading 349 (OC: > 20 μ g m⁻³; EC: > 8 μ g m⁻³) at relatively high RH at 60-80 % which usually occurred at 350 night with relatively low boundary layer height, leading to the accumulation of aerosols. However, 351 the corresponding OC/EC ratios were low, suggesting the importance of primary sources to OC 352 and EC in northern Nanjing, which will be verified in the wind rose of OC and EC (Fig. 8). 353

354 **3.2 Air mass transport**

355 3.2.1 Windrose of carbonaceous aerosols

356 To investigate the influence of air mass transport to the study site, the wind rose of OC, EC and dEC/OC using hourly data in the four seasons is shown in Fig. 8 (Carslaw and Ropkins, 2012). 357 Two points should be noted. First, high OC and EC mass concentrations were found near the field 358 site (indicated by wind speed (WS) $< 1 \text{ m s}^{-1}$), suggesting that local and primary emissions (e.g., 359 360 industrial and vehicle emissions) were stable and important sources contributing to atmospheric OC and EC mass concentrations in northern Nanjing. The OC mass concentrations from the 361 southwest increased with the increase of WS in summer, indicating that the sources of OC are 362 complicated in summer including secondary reaction during long-range or regional transport. 363 364 Second, compared with OC and EC, dEC showed no significant local sources. The dEC/OC increased with increasing WS and the highest dEC/OC were found for WS over 3 m s⁻¹. Long-365 range or regional transport were highly likely the main sources contributing to the dEC mass 366 concentrations. 367

368 3.2.2 The potential source areas of carbonaceous aerosols

The possible source contributions were evaluated using the PSCF model and the PSCF maps are shown in Fig. 9 (Petit et al., 2017). The areas with high PSCF values were highly likely the potential pollution source areas. As shown in Fig. 9, the PSCF results further proved the strong

regional transport contribution to dEC and local contributions to OC and EC. In spring, the 372 potential source areas of OC and EC were mainly from the southwest of Nanjing; however, the 373 potential source areas of dEC were from the east of Nanjing, indicating obvious different sources 374 between OC, EC and dEC. In summer, local areas were the main source areas of EC and the near-375 by Yangtze River Delta City Group from the southeast of Nanjing including developed cities like 376 Shanghai were the main sources areas of OC and dEC. The anthropogenic emissions from these 377 areas might be important sources of OC and dEC. Besides, both the potential source areas of dEC 378 379 and EC were in the northwest of Nanjing in summer, suggesting strong primary sources of dEC from this area which were very likely associated with biomass burning, more details are given in 380 section 3.3. In autumn, local sources from the study site were strongest for OC and EC. However, 381 dEC mainly originated from regional transport from the northwest and southeast areas of Nanjing. 382 383 Biomass burning has been proved to be an important source of air pollutants in the Yangtze River Delta (YRD) area, especially in the wheat harvest seasons (e.g., June and October) (Cheng et al., 384 385 2014; Zhang and Cao, 2015a). In addition, the YRD area is the most economically developed region in China and has lots of industrial cities, which means that industrial emissions and 386 387 anthropogenic sources contributed to high carbonaceous aerosol pollution levels. In winter, dEC was mainly from long-range transport from northern cities and regional transport from the 388 389 southwest areas of Nanjing while both long-range transport and local sources were found in OC and EC concentrations. 390

391 **3.3** The characteristics of carbonaceous aerosols during biomass burning periods

The biomass burning emission has been proved to be an important source of BrC on a global 392 scale; it is consistently observed in large-scale forest fire events (Laskin et al., 2015). Based on the 393 Fire Information for Resource Management System (FIRMS) derived from the Moderate 394 395 Resolution Imaging Spectroradiometer (MODIS), we found that the fire points amounted to 2028, 396 1773 and 967 on 11 Jun 2015, 7 February 2016 and 2 March 2016, respectively, in the areas around our study site, suggesting there were strong biomass burning events on these days (Fig. S4). To 397 further investigate the biomass burning impact on dEC, we analyzed the temporal trends of 398 carbonaceous aerosols from 4 June 2015 to 19 June 2015 and 7 February 2016 to 3 March 2016, 399 respectively. Combining the observed aerosol concentrations and fire information, we divided the 400 periods into normal days and biomass burning days. It should be noted that the biomass burning 401 days are not determined based only on fire points. We also considered the 48-h backward 402

trajectories and open biomass burning areas. For example, we found lots of fire points from 11 403 June 2015 to 12 June 2015 and from 7 February 2016 to 10 February 2016, respectively, and the 404 48-h back trajectories passed over these biomass burning areas (Fig. S5b, c). However, although 405 there were large amounts of fire points in the northwest of Nanjing from 8 June 2015 to 9 June 406 2015, the backward trajectories showed that the air masses during the periods came from the 407 southeast areas where no open fire points were found (Fig. S5a). In contrast, there were only a few 408 fire points found near the study site from 26 February 2016 to 27 February 2016, the 48-h backward 409 trajectories showed the air masses came exactly from the area (Fig. S5d). 410

As shown in Fig. 10 and Fig. 11, we found that dEC concentrations, dEC/OC and OC/EC 411 ratios showed peaks during each biomass burning period which were not that obvious in OC and 412 EC concentrations, suggesting the unique biomass burning impact on dEC and the sources of OC 413 414 and EC were more complicated. It should be noted that there were peaks of dEC on 9 June 2015 and 13 February 2016, which were not biomass burning days, suggesting that biomass burning 415 was not the only source of dEC. As mentioned in sections 3.1 and 3.2, anthropogenic emissions 416 could be the sources of dEC and the secondary sources cannot be ignored, either. Summarized in 417 418 Table 2 are the average and standard deviation values of OC, EC, OC/EC, dEC and dEC/OC during biomass burning and normal days. The OC/EC, dEC concentrations and dEC/OC were obviously 419 420 higher in biomass burning days than in normal days, but similar levels of the OC and EC concentrations were found both in biomass burning days and normal days in summer, suggesting 421 422 the great contribution of biomass burning emissions to dEC and there were other sources of OC and EC in summer. All the carbonaceous aerosols were higher in biomass burning days in winter; 423 in addition, the locations of open fire counts were mainly in the northwest and southwest area of 424 the study site (Fig. S5c, d), which were the potential source areas of OC, EC and dEC in winter as 425 426 discussed in section 3.2.2, indicating strong contributions of biomass burning emissions to all the carbonaceous aerosols in winter. 427

428 4. Conclusions

In this study, the characteristics and sources of carbonaceous aerosols in North Nanjing were investigated and we introduced a two-wavelength method by modifying the Sunset carbon analyzer. We incorporated a new diode laser at λ =405 nm in the instrument, making it possible to detect the laser beam passing through the filter at both wavelength λ =658 nm and λ =405 nm, so that we can obtain the dEC concentrations. Our study illustrated the feasibility of using dEC to characterize

the BrC aerosols, providing a new idea about the measurement of BrC. The results showed that 434 high (low) OC, EC and dEC concentrations were found in winter (summer), indicating the 435 significant impact of the increase of various emission sources in winter and wet scavenging by rain 436 in summer. Similar diurnal cycles for OC and EC concentrations were found with high at night 437 and low in daytime, strongly affected by the boundary layer. Traffic emissions were found to have 438 a significant influence on the concentrations of OC and EC. Similar trends were found in the 439 diurnal cycle of dEC/OC and OC/EC and the dEC/OC increased when the OC/EC ratio increased, 440 indicating strong secondary sources or biomass burning impact on dEC. The wind rose and 441 receptor model results showed that strong local emissions were found for OC and EC; however, 442 dEC was significantly affected by regional or long-range transport. The near-by YRD area was one 443 of the main potential source areas of dEC, suggesting that anthropogenic emissions could be the 444 445 sources of dEC. Together with the back trajectories analysis and MODIS fire information, we analyzed two biomass burning events both in summer and winter. The results showed that the 446 sources of OC and EC were more complicated than those of dEC in summer. Biomass burning 447 emission made a great contribution to dEC concentrations in summer. A large number of open fire 448 449 counts from the northwest and southwest areas of the study site was observed; these fires significantly contributed to the carbonaceous aerosol pollution. 450

451 Our modified two-wavelength instrument provided more information than the traditional single-wavelength thermal-optical carbon analyzer. The results proved that dEC can be an 452 453 indicator of BrC in biomass burning days. It should be noted that the sources of dEC were complicated and the anthropogenic emissions and secondary formations of dEC aerosols could not 454 be ignored; further chemical analysis needs to be conducted in the future. The evaluation of SOC 455 formation and the relationship between dEC and SOC can be conducted. In addition, more 456 457 chemical analysis such as the analysis for ions, organic matter or sugars in PM_{2.5} can be made; 458 thus we can get some information of the tracers of different sources and more accurate and quantitative source apportionment can be done (Bhattaraia et al., 2019; Wu et al., 2018; 2019). We 459 also hope that the dEC data can be further applied in more research. 460

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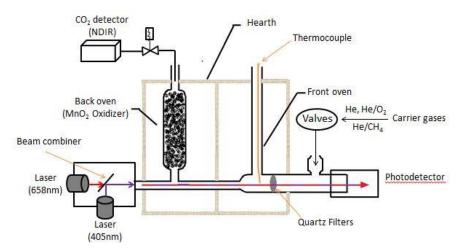
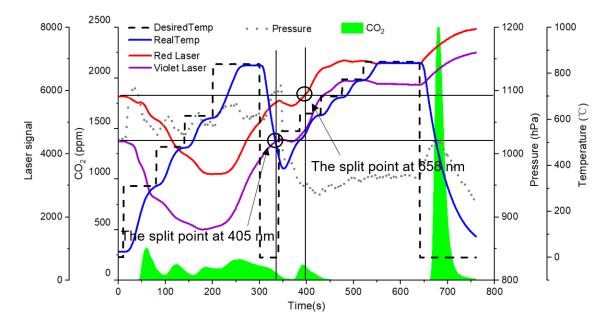


Figure 1. Principle and structure of the Sunset semi-continuous carbon analyzer.



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Figure 2. Example thermogram of sample analysis using the two-wavelength Sunset semicontinuous carbon analyzer.

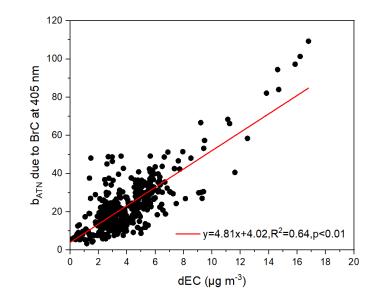
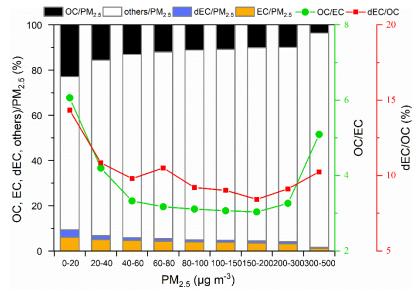




Figure 3. Relationship between the b_{ATN} due to BrC at 405 nm and the dEC concentrations.



 $\begin{array}{rl} & \mathsf{PM}_{2.5} \ (\mu g \ m^{-3}) \\ \textbf{682} & \textbf{Figure 4. Carbonaceous species fractions of } \mathsf{PM}_{2.5} \ \text{and } \mathsf{OC/EC} \ ratios \ at \ different \ \mathsf{PM}_{2.5} \\ \end{array}$

concentration intervals at NUIST from June 2015 to August 2016.

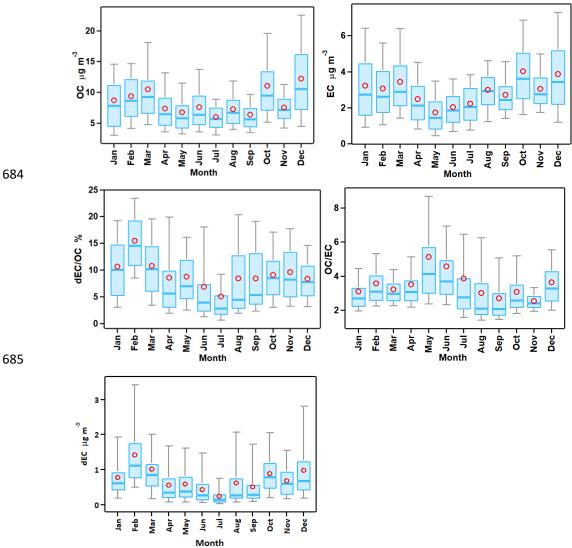


Figure 5. Monthly variations of OC, EC, dEC, dEC/OC and OC/EC ratios at NUIST from June 687 2015 to August 2016. The boundary of the box indicates the 25% and 75% percentile, 688 respectively. The lower and upper whiskers indicate the 10% and 90% percentile, respectively. 689 The red circle within the box marks the average while the line within the box marks the median. 690

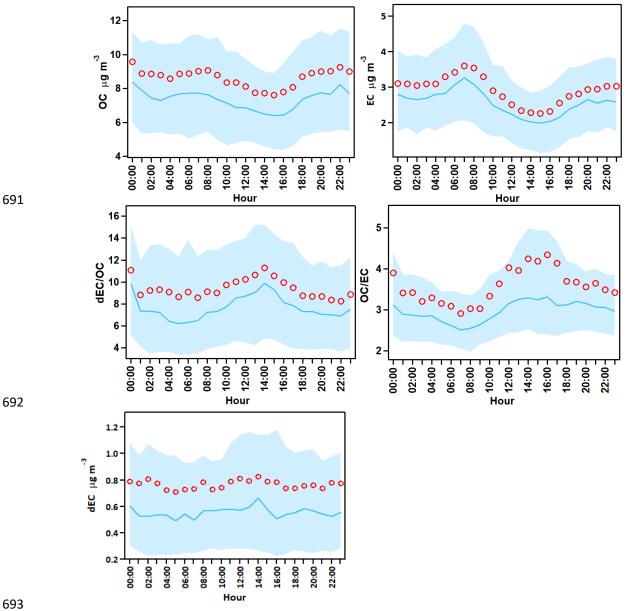


Figure 6. Diurnal variations of OC, EC, dEC concentrations, dEC/OC and OC/EC ratios during
the study period. The boundary of the shaded area indicates the 25% and 75% percentile,
respectively. The red circle marks the average while the blue line marks the median.

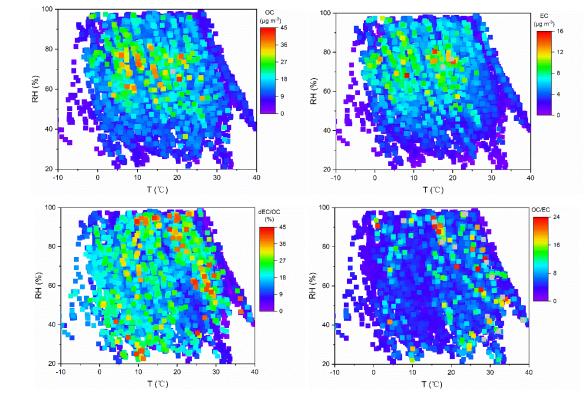


Figure 7. RH/T dependence of OC, EC, dEC/OC and OC/EC ratios during the study period.

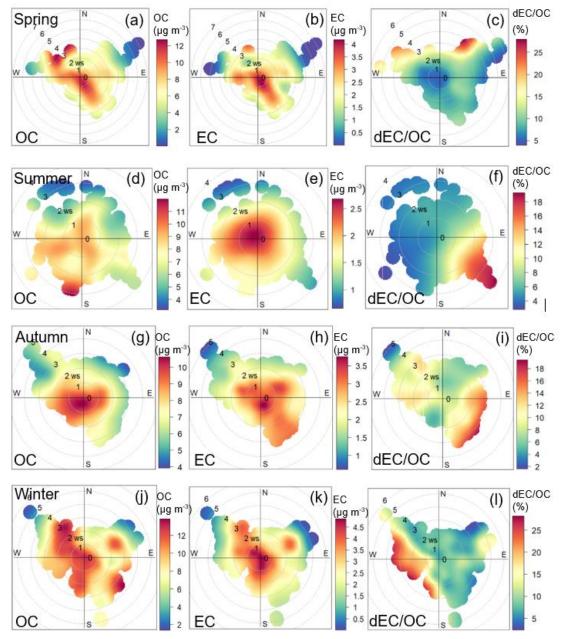
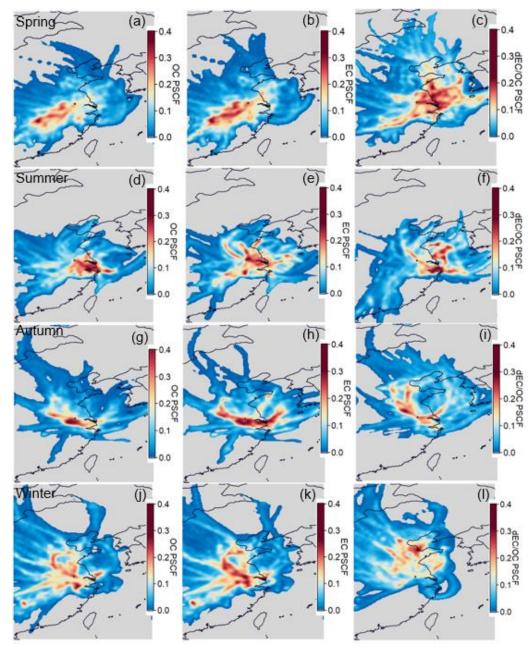


Figure 8. Wind rose of OC, EC and dEC/OC in spring ((a), (b), (c)), summer ((d), (e), (f)), autumn

^{702 (}g), (h), (i)) and winter ((j), (k), (l)).





703 704 Figure 9. PSCF map for OC, EC and dEC/OC in spring ((a), (b), (c)), summer ((d), (e), (f)), autumn

705 (g), (h), (i)) and winter ((j), (k), (l)).

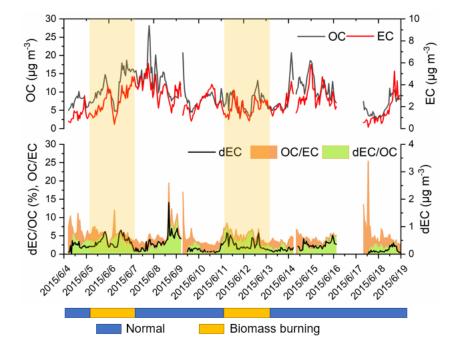
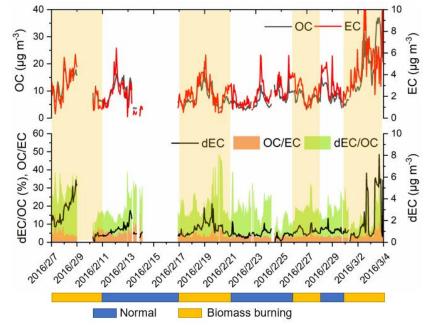


Figure 10. Time series of OC, EC, dEC/OC, dEC and OC/ EC from 4 June 2015 to 19 June 2015.

The period was divided into normal days (blue bars) and biomass burning days (yellow bars). The
yellow shadow represents the biomass burning periods.



710
711 Figure 11. Time series of OC, EC, dEC/OC, dEC and OC/ EC from 7 February 2016 to 3 March

712 2016. The period was divided into normal days (blue bars) and biomass burning days (yellow bars).

713 The yellow shadow represents the biomass burning periods.

	Annual					Spring	Summer	Autumn	winter
N=5113	Average	Standard rage Median Min Ma Deviation		Max	Average	Average	Average	Average	
PM _{2.5} (µg m ⁻³)	77.2	48.6	65.0	2.5	458.1	72.1	47.9	70.5	91.8
OC (µg m ⁻³)	8.9	5.5	7.5	0.5	45.8	8.4	7.2	8.4	10.2
EC (µg m ⁻³)	3.1	2.0	2.6	0.0	17.6	2.6	2.3	3.3	3.4
OC/EC	3.5	2.4	2.9	1.0	29.3	3.9	4.0	2.8	3.4
$dEC (\mu g m^{-3})$	0.8	0.8	0.6	0.0	8.1	0.8	0.5	0.7	1.1
dEC/OC (%)	10.0	7.2	8.6	0.0	48.2	9.5	6.9	9.0	11.3
dEC/EC (%)	22.3	16.7	18.5	0.1	97.8	24.5	18.2	18.7	25.9
OC/PM _{2.5} (%)	12.8	5.6	11.6	0.7	66.2	13.2	14.4	14.1	11.1
EC/PM _{2.5} (%)	4.3	2.3	3.9	0.0	33.2	3.9	4.7	5.8	3.7
dEC/PM _{2.5} (%)	1.3	1.2	0.9	0.0	17.6	1.4	1.3	1.2	1.3

Table 1. Statistical summary on the $PM_{2.5}$ and carbon species concentrations.

		OC	EC	OC/EC	dEC	dEC/OC
		(µg m ⁻³)	(µg m ⁻³)	UC/EC	(µg m ⁻³)	(%)
June 4 th to 19 th	Normal days	9.5 ± 4.5	2.6±1.3	4.3±2.3	0.2 ± 0.1	2.5 ± 1.3
	Biomass burning days	9.0 ± 3.6	2.0 ± 0.9	4.8±1.6	0.4 ± 0.2	4.6 ± 1.4
February 7 th to March 3 rd	Normal days	7.5 ± 3.3	2.5 ± 1.2	3.3±1.3	0.8 ± 0.3	12.7 ± 5.6
	Biomass burning days	11.2±7.2	3.1±1.9	4.0±1.8	1.7±1.4	15.4±7.8

Table 2. Statistics of OC, EC, OC/EC, dEC and dEC/OC during biomass burning days and normal
days. The values represent average±standard deviation.