

Investigation of structural changes of atmospheric aerosol samples during two thermal-optical measurement procedures (EUSAAR2, NIOSH870)

Theresa Haller¹, Eva Sommer¹, Thomas Steinkogler³, Christian Rentenberger¹, Anna Wonaschuetz¹, Anne Kasper-Giebl³, Hinrich Grothe², Regina Hitzenberger¹

5 ¹University of Vienna, Faculty of Physics, Vienna, 1090, Austria

²TU Wien, Institute of Materials Chemistry, Vienna, 1060, Austria

³TU Wien, Institute of Chemical Technologies and Analytics, Vienna, 1060, Austria

Correspondence to: theresa.haller@univie.ac.at

Abstract

10 Thermal-optical measurement techniques are widely used for the monitoring of carbonaceous aerosols. Although results of different thermal-optical measurement techniques are comparable for total carbon, they can vary widely for values of elemental carbon especially in the presence of brown carbon. Charring of organic material during the inert heating phase of thermal-optical measurements has been found to be a major confounder, but no literature about investigations of structural changes during this process in atmospheric aerosols is available. In a recent study we investigated these structural changes
15 for combustion aerosol standard soot (CAST). Now we apply this approach to selected atmospheric aerosol filter samples and a subset of eight washed filter samples with low loadings of water soluble carbon (WSOC). To investigate structural changes, Raman spectra were obtained for samples heated to the corresponding temperature levels and gas atmospheres of the EUSAAR2 and NIOSH870 protocols. The temperature levels where changes in the Raman spectra occurred (i.e. changes in structure) varied for different samples. For the washed samples with low WSOC loadings and absence of other water
20 soluble aerosol components such as inorganic salts, changes in structural ordering and darkening of the samples were not observed. We were able to show for the first time that the darkening of a sample (measured in terms of transmission laser signal) is not necessarily caused by an increase of structural ordering in the sample. Possible transformations at lower temperatures could include a formation of non-graphitic light absorbing intermediate organic carbon, a release of C-H groups or a decrease of carbonyl groups.

25 1 Introduction

Carbonaceous matter is an important component of atmospheric aerosols because of its adverse impacts on climate and human health (e.g. Bond and Bergstrom, 2006; Highwood et al., 2006). It spans a rather continuous range from highly light absorbing graphitic-like material to non-absorbing organic molecules (e.g. Pöschl, 2005) and is usually classified optically into Light Absorbing Carbon (LAC) covering Brown Carbon (BrC) and Black Carbon (BC) or thermo-chemically into

30 Organic Carbon (OC), Elemental Carbon (EC) and inorganic or Carbonate Carbon (CC) (Petzold et al., 2013). However, these definitions are tied to measurement methods rather than to the actual internal structure of the material which makes reliable measurements challenging. Using thermal measurement techniques OC is defined as the thermally unstable carbonaceous material which evolves in the absence of oxygen at temperatures below 550°C – 700°C (Bond and Bergstrom, 2006; Chow et al., 2004). EC is defined as the fraction which does not evolve below 550°C - 700°C in the absence of oxygen
35 and combusts in the presence of oxygen at temperatures above 600°C (Andreae and Gelencsér, 2006).

Results for EC and OC measured with different thermal-optical techniques suffer from biases by up to 44% for EC, while the values of total carbon (TC=EC+OC) are comparable within 5%-15% (e.g. Reisinger et al. 2008; Cavalli et al., 2010; Yu et al., 2002; Cheng et al., 2012; Hitzenberger et al., 2006; Panteliadis et al., 2015; Müller et al., 2004; Venkatachari et al., 2006; Watson et al., 2005; ten Brink et al., 2004). The largest discrepancies were found for samples obtained under wintry
40 conditions and in the presence of biomass smoke (Reisinger et al., 2008; Wonaschuetz et al., 2009, Cheng et al., 2012), when large amounts of BrC can be found (Lukács et al., 2007). The optical absorptivity of BrC is markedly higher in the shortwave range of the visible spectrum than in the longwave range. BrC is more thermally refractory than other organic matter and can therefore be seen as a substance with properties intermediate between those of elemental and organic carbon (e.g. Pöschl et al. 2003).

45 In thermal-optical measurement techniques a filter sample is heated stepwise first in an inert atmosphere and then in an oxidizing atmosphere and the amount of carbon leaving the filter is measured. Different temperature protocols vary in height and duration of the temperature steps and particularly in the maximum temperature of the inert phase (550°C - 900°C, Cavalli et al., 2010). A potential darkening of the sample during the inert phase, commonly explained by pyrolysis and/or charring of the sample producing “Pyrolyzed Carbon” (PC), is monitored with a transmission and/or reflection laser signal.

50 The transmission/reflection laser signal is used to correct for PC: The amount of carbon detected before this signal reaches its initial value (i.e. the split point) is assigned to OC and carbon detected after this point is assigned to EC. However, this optical correction is based on two essential assumptions which are not generally fulfilled: firstly PC does not evolve totally before EC, and secondly the specific light absorption coefficients of PC and EC are not equal (Yu et al., 2002; Chow et al., 2004; Subramanian et al., 2006; Han et al., 2007). Moreover, the specific light absorption coefficient of PC can change
55 during the heating procedure (Yu et al., 2002). Because of these uncertainties the optical correction can lead to an over- or underestimation of EC or OC (Yu et al., 2002; Cheng et al., 2012).

Extensive efforts were undertaken to refine the original thermal-optical measurement procedure (Birch and Cary, 1996) in order to achieve a more accurate setting of the split point and to reduce the formation of PC (e.g. Cavalli et al., 2010; Chow et al., 2004; Cheng et al., 2012). EUSAAR2 was found to produce less PC than other protocols during analyses of ambient
60 European aerosols (Cavalli et al., 2010) and was subsequently selected as the European standard method (Brown et al., 2017).

Despite all efforts, pyrolysis is still a problematic factor in thermal-optical analyses depending on sample composition. Water soluble organic carbon (WSOC) was found to be responsible for a large fraction of charring (13%-66%; Yu et al.,

2002). Several inorganic constituents can influence charring in complex and sometimes contradictory ways: Ammonium
65 bisulfate enhances the charring of starch and cellulose but reduces the charring of levoglucosan (Yu et al. 2002). Metal salts
can generally enhance the charring of OC at least in diesel soot and ambient aerosols (Wang et al., 2010). As a consequence
of this interplay of several complex chemical reactions there is no unequivocal effect of non-carbonaceous aerosol
constituents on the charring behavior (Wang et al., 2010; Yu et al., 2002; Bladt et al., 2014) and a simple prediction of the
charring behavior of an atmospheric aerosol sample is currently impossible.

70 Most of the studies of pyrolysis during thermal-optical measurement procedures only analyze the laser signals (e.g. Cavalli
et al., 2010; Cheng et al., 2012; Yu et al., 2002) or the thermal properties of PC (Yu et al., 2002). The terms “charring” and
“pyrolysis” are often used interchangeably to describe processes leading to sample darkening which is implicitly assumed to
be caused by a formation of a more structured phase of carbon. Underlying structural changes are not explained. Le et al.
75 (2019) investigated structural changes of laboratory generated aerosols (produced in a CAST (combustion aerosol standard)
soot generator) during heat treatment in N₂ and give detailed information about evolution and restructuring of carbonaceous
species. However, the gas atmospheres as well as the heating rates differ from the conditions during thermal-optical
analyses.

To our knowledge no study to date has addressed the structural changes of an atmospheric aerosol sample during a thermal
optical measurement procedure by analyzing graphitization or structural ordering. In a previous study (Haller et al., 2019) we
80 investigated the increase of structural ordering of two widely different combustion aerosol standard soot types (produced by
propane combustion in a CAST burner) during a thermal-optical heating procedure (NIOSH870). Using Raman
spectroscopy, we found that structural ordering increased at 870°C when a BrC rich (and BC poor) sample was heated in a
He atmosphere, while a BC rich (and BrC poor) sample did not further change its already high degree of graphitization. This
approach is now applied to the much more complex situation of atmospheric aerosols. For 21 atmospheric aerosol samples
85 selected from a two-year set of daily filter samples the changes of structural ordering during the inert phases of the
NIOSH870 (Birch and Cary, 1996) and EUSAAR2 (Cavalli et al., 2010) protocols were investigated. From a subset of
filters, WSOC and other soluble aerosol material was removed by washing. The protocols were chosen because of the
different production of PC reported by Cavalli et al. (2010). A dual-optics thermal-optical analyzer (Sunset Instruments Inc.)
90 was used to produce heated samples and to determine EC and OC. Samples heated to each temperature plateau of the inert
phases of both temperature protocols were analyzed with Raman spectroscopy to investigate the degree of intermolecular
ordering in the material. Several properties of the samples (e.g. ionic composition, EC/OC, BC/BrC) were compared with the
progress of structural changes derived from the Raman spectra.

2 Experimental

2.1 Sampling

Atmospheric aerosol was sampled on quartz fiber filters (Pall Tissuquartz 2500 QAT-UP, 47mm) at the rooftop lab of the 100 physics building of the University of Vienna (35m above ground) located about 1.5 km from the city center. Although the building is located in an area highly impacted by traffic, the aerosol at the rooftop lab can be seen as an urban background, since the sampling inlet faces the enclosed source-free courtyard of the building. The aerosol was sampled with an automatic 105 sequential filter sampler (SEQ47/50, Sven Leckel) equipped with a PM2.5 inlet and set to a flow rate of 2.3 m³/h. Sampling time was set to 23 h from 12:00 to 11:00 (UTC+1) of the next day. The loaded filters were stored in the filter sampler for 110 about one to two weeks in a closed magazine before they were transferred to the lab and stored at -22°C in petri dishes closed with PARAFILM®. Before sampling the filters had been heated at 450°C for an hour to remove volatile organic substances and stored for at least 24h in a water-vapor-saturated atmosphere to prevent re-adsorption of OCs (Jankowski et al., 2008) and to ensure low OC blanks (blank values typically around 0.2 µgC/cm²). The samples were collected in 2014 and 2015 for nearly every day. BC and BrC values were obtained from these filters using the Integrating Sphere method (Sommer, 2020).

2.2 Integrating sphere method

Light absorbing carbon (LAC) (i.e. BrC and BC) was analyzed for all filter samples with an extension (Wonaschuetz et al., 115 2009) of the original integrating-sphere technique (an extensive description is given by Hitzenberger and Tohno, 2001). Circular filter punches with diameters of 10 mm were immersed in a mixture of 10% isopropanol, 40% H₂O and 50% acetone in polyethylene (PE) vials. The immersion in this mixture reduces enhanced absorption caused by possible non-absorbing coatings of the particles: Soluble material is removed, and the effect of insoluble coatings is reduced due to the 120 similar refractive indices of the liquid and the coatings (Hitzenberger and Tohno, 2001). The PE vials were placed in the center of a 6 inch integrating sphere coated internally with a highly diffusely reflective material (Spectraflect™). The sample was illuminated with a halogen light source equipped with a diffusor and three interference filters (450, 550 and 650 nm) and the wavelength-dependent light signal was recorded with a photodiode.

The BC and BrC content in the sample was then calculated in an iterative procedure described by Wonaschuetz et al. (2009), comparing the wavelength-dependent light signals with calibration curves obtained with a proxy for BC (Elftex 124, Cabot Corp.) and a proxy for BrC (humic acid sodium salt, Acros Organics, 20 no. 68131044). The values for BrC should 125 therefore be understood as “humic acid sodium salt equivalent” and refer to the total mass of BrC including other atoms (humic acid sodium salt contains 47% carbon by mass). Detection limits are 1 µg per sample for BC and 10 µg per sample for BrC.

2.3 Sample selection and preparation

From the total set of 660 filter samples 21 samples were chosen for further analysis. As the focus of this study is on the restructuring of carbonaceous material with special emphasis on the presence of BrC, winter samples with high BrC loadings

130 and high BrC/LAC ratios and summer samples with BrC loadings below detection limit were selected. The seven samples from summer 2015 had LAC loadings between about 3 and 5 $\mu\text{g}/\text{cm}^2$ with BrC/LAC = 0. The fourteen samples from winter 2014/2015 had LAC loadings between about 8 and 27 $\mu\text{g}/\text{cm}^2$ and BrC/LAC ratios from about 0.3 to 0.7 which were the highest BrC/LAC ratios in this period. TC values varied from 7 to 34 $\mu\text{g}/\text{cm}^2$ for summer samples and from 13 to 71 $\mu\text{g}/\text{cm}^2$ for winter samples.

135 To investigate the effect of WSOC, eight randomly selected filters (out of the 21 samples) were cut in half and one piece of each filter was immersed in 80 ml of MilliQ water (18 M Ω cm, Millipore) in 100 ml bottles and shaken mechanically in a shaking water bath (GFL Technology, type GFL 1086) for about 20 minutes at room temperature with 100 rpm to remove most of WSOC as well as other soluble aerosol material (e.g. Gao et al., 2003; Kirchstetter and Novakov, 2004). The washed filters were removed from the bottles, transferred into petri dishes and dried in a desiccator for 3-4 days. The dry filters were
140 then stored at -22°C in closed petri dishes. In the following discussions, aliquots of the untreated (i.e. unheated and unwashed) filters are referred to as “original” samples.

2.4 Thermal-optical measurements and preparation of heated samples

The EC and OC content of the samples was measured with a dual-optics thermal-optical instrument (Sunset Instruments Inc.). In this instrument the sample is heated stepwise first in He, then in an oxidizing atmosphere (2% O₂ in He). The

145 duration times of the steps and the temperatures and atmospheres for two different heating protocols (EUSAAR2 and NIOSH870) are listed in Table 1. The evolving carbon during each temperature step is tracked by a flame ionization detector (FID). The FID signal is proportional to the evolving carbon content and is calibrated by introducing a known amount of methane at the end of each filter measurement. The darkening of the sample during the heating procedure in the inert phase is monitored with a laser signal (in transmission and reflection with wavelength: 635 nm) which is used afterwards to correct
150 the results for pyrolysis: A split point is defined when the laser signal reaches its initial value and the carbon evolving prior to the split point is assigned to OC and the carbon evolving afterwards to EC. The precision of the instrument is 6% for OC as well as for EC and the lower detection limit is 0.1 $\mu\text{gC}/\text{cm}^2$ for OC and EC, respectively.

Twelve out of the 21 atmospheric aerosol samples were analyzed with both EUSAAR2 and NIOSH870 protocols (Table 1), the other eight samples only with EUSAAR2, as parts of these filters were washed to remove water soluble material
155 including WSOC (see below) and were therefore not available for extra analysis with another protocol. For both protocols heated samples were prepared corresponding to the temperature levels of the inert phase and the first oxidizing temperature level of EUSAAR2: Aliquots of the filters were heated in the instrument following the NIOSH870 and EUSAAR2 temperature steps and atmospheres, respectively, before the procedure was interrupted at the desired temperature level. All

samples were cooled down in He to below 75°C, removed from the oven and stored in a closed petri dish at laboratory

160 temperature.

Carrier Gas	NIOSH870				EUSAAR2			
	duration [s]	T [°C]	code	Raman	duration [s]	T [°C]	code	Raman
He	80	310	NI310	✓	120	200	EU200	✓
He	80	475	NI475	✓	150	300	EU300	✓
He	80	615	NI615	✓	180	450	EU450	✓
He	110	870	NI870	✓	180	650	EU650	✓
He	45	550			30	-		
He+O ₂	45	550			120	500	EU500	✓
He+O ₂	625	45			120	550		
He+O ₂	700	45			70	700		
He+O ₂	775	45			80	850		
He+O ₂	850	45			-	-		

Table 1: Duration times, temperatures and gas atmospheres for the two different thermal-optical measurement protocols (NIOSH870, EUSAAR2). The “Raman” column indicates that heated samples were analyzed with Raman spectroscopy for the respective temperature (see below). The codes are used when we refer to samples heated to a certain temperature – symbolized by the numeric index – following NIOSH870 (NI) or EUSAAR2 (EU).

165 **2.5 Raman Spectroscopy**

For the investigation of the soot structure a confocal Raman microscope (Horiba Jobin-Yvon LabRAM 800HR) was used.

The Raman microscope was equipped with a 632.8 nm HeNe-laser (maximum output <20 mW), a 532 nm frequency-doubled NdYAG DPSS laser (Oxxius LMX-532, maximum output <52 mW) and a CCD camera (Peltier cooled to -60°C).

The laser beam was focused onto the sample with a 20x magnification objective lens (CF Plan, 20x/0.35, WD 20.5 mm,

170 Nikon) and a 100x long working distance (LWD) objective lens (MPlanN, 100x/0.75, WD 0.21 mm, Olympus). The spectra were calibrated with the silicon peak at 520.07 cm⁻¹. Generally Raman signals were obtained for 10 spots for each (unheated and heated) filter sample within an area of (1x2) mm². We found that this area was large enough to characterize a whole sample. Only for some samples with too low loadings five spots were focused manually. The acquisition times were chosen between 0.005 and 3 seconds and 30 – 3000 single spectra were accumulated for each position (slit 400 µm, confocal hole

175 500 µm). A laser power between 25% and 100% was used. Generally laser power was lowered for highly absorbing samples to prevent thermal destruction of the material, and for samples containing fluorescent material to reduce the interfering fluorescence background. The particular settings (acquisition time, number of accumulations, laser power) were chosen individually to obtain the best signal to noise ratio for each sample, since the fluorescence background intensity was different for the samples and for different excitation wavelengths.

180 Due to instrument availability all samples were measured using the HeNe (red) laser except the heated and unheated samples of the washed filters (see below) and their unwashed counterparts which were measured using the NdYAG (green) laser. For samples heated to 500°C in O₂+He (EUSAAR2) and 870°C in He (NIOSH870), which were measured using the red laser, a 100x LWD objective lens and a smaller confocal hole (3 μ m) were used to focus on single particles since the red laser in combination with the 20x objective lens was not sufficient to obtain spectra for the small amounts of material left on the
185 filters at these heating stages. The laser power was decreased to 1% and accumulations of 400 single spectra and an acquisition time of 0.5 sec were required. Because of these restrictions, Raman spectra were obtained for only five different spots for these samples. Mean spectra and their standard deviations were calculated for the spectra of all samples using the 5 to 10 spectra obtained from different measurement positions.

Typical Raman spectra of soot and related carbonaceous material show two overlapping peaks at about 1350 cm⁻¹ (D-peak or 190 defect peak) and 1580 cm⁻¹ (G-peak or graphitic peak) (e.g. Fig. 1). Several authors (e.g. Sadezky et al., 2005; Le et al., 2019; Zickler et al., 2006 and references there) suggest that these peaks arise from five or more bands attributed to different vibrational modes in the soot structure. Detailed descriptions are given by e.g. Rosen and Novakov, 1978; Sadezky et al., 195 2005; Ferrari and Robertson, 2000; Schmid et al., 2011; Knauer et al., 2009; Ivleva et al., 2007. The ratio of the D and G peak intensities (D/G) is an indication of the degree of structural ordering in the carbonaceous material (Tuinstra and Koenig, 1970; Ferrari and Robertson, 2000). For small crystallite sizes $L_a < 2$ nm, as can be found in combustion soot particles (see e.g. Haller et al., 2019), D/G is proportional to the probability to find 6-fold aromatic rings in the cluster, which means that an increasing D/G ratio indicates increasing crystallite sizes (Ferrari and Robertson, 2000; Zickler et al., 2006) and therefore an increase of structural ordering. Ferrari and Robertson (2000) propose proportionality between D/G and L_a^2 with a wavelength dependent factor. Because of this wavelength dependence, we discuss here only samples where all spectra (original, washed, heated, unheated) were obtained with the same excitation wavelength. A more detailed description of the interpretation of Raman spectra of soot can be found in the studies by Ferrari and Robertson (2000) or Haller et al. (2019). Several other authors use this interpretation of the D/G ratio for Raman spectra of combustion aerosol samples (e.g. Commodo et al., 2016; Ess et al., 2016) and use a higher D/G ratio as an indication of a higher degree of structural ordering within the material.

205 2.7 Electron Microscopy

As crystallite sizes are important for the interpretation of the D/G ratio in the Raman spectra, the crystallite sizes of three samples (16.11.2014, 09.06.2015, 10.11.2015) were exemplarily determined from electron diffraction patterns obtained with an electron microscope. A 200 kV transmission electron microscope (TEM) Philips CM200 equipped with a GatanTM Orius
210 CCD camera was used to obtain electron diffraction patterns. Intensity profiles of the diffraction patterns were obtained by both azimuthal integration along rings and background correction using PASAD-tools (Gammer et al., 2010). By fitting pseudo-Voigt functions to the intensity profiles taken from at least 6 different locations on each sample, the average full

width at half maximum of the first intense diffraction peak was determined and used to calculate the mean crystallite size with the Scherrer equation (Fultz et al., 2001).

2.8 Ion chromatography

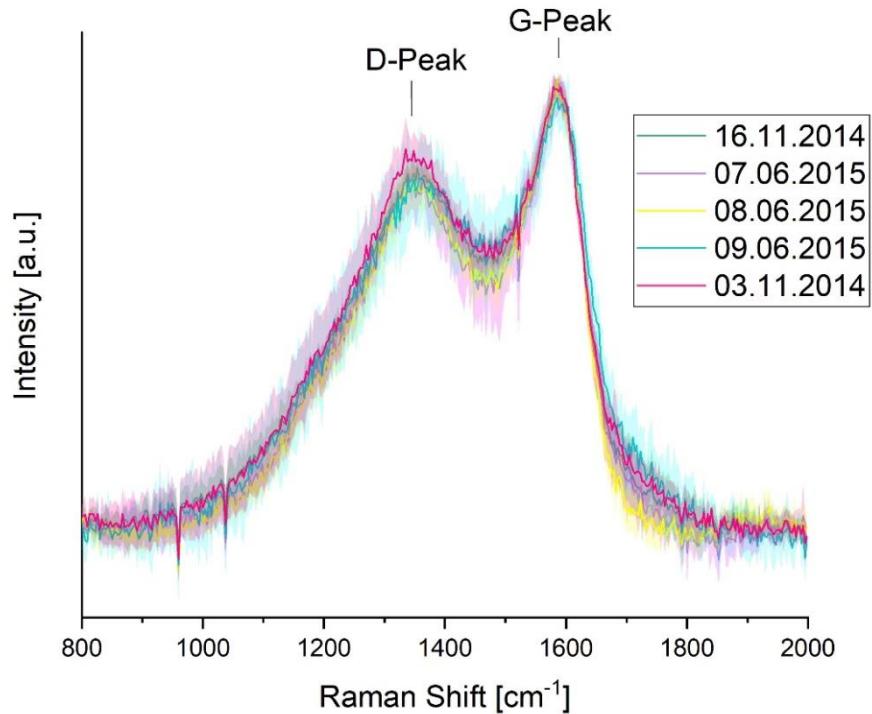
215 Aliquots of the filters (areas between 78 and 101 mm² depending on filter loadings) were extracted in 2 mL of ultrapure water (MilliQ, 18 MΩcm, Millipore). After ultrasonic agitation for 20 minutes and centrifugation for 5 minutes at 13400 rpm, the extracts were analyzed by suppressed ion chromatography. For the determination of the anions (Cl⁻, NO₂⁻, NO₃⁻ and SO₄²⁻) a Dionex Aquion system (Thermo Fisher) equipped with CS16 and CG16 columns, a CERS 500 self-regenerating suppressor and a DS6 heated conductivity detector was used. The cations (Na⁺, NH₄⁺, Mg²⁺, Ca²⁺ and K⁺) were 220 determined with a Dionex ICS-1100 system (Thermo Fisher) equipped with AS22 and AG22 columns, an AERS 500 self-regenerating suppressor and a DS6 heated conductivity detector. The calibration was performed with external standards, prepared from Merck stock solutions with 1000 ppm. The limit of detection for Cl⁻ is 0.005 mg/L, for all other analytes 0.01 mg/L.

3 Results

225 3.1 Properties of the original (unwashed and unheated) samples

The Raman spectra of all samples (e.g. Fig. 1, 2, 3) show two distinct peaks around 1350 cm⁻¹ (D-peak or defect peak) and 1580 cm⁻¹ (G-peak or graphitic peak) which is typical for soot. The Raman spectra of the original samples are surprisingly similar (Fig. 1), although the samples differ in composition. This could be explained by an appreciably higher contribution of graphitic like material to the Raman signal in comparison to the contribution of possibly coexisting organic or brown carbon.

230



235 **Figure 1: Raman spectra for a representative subset of original (unheated, unwashed) samples. Error bars (standard deviation of the mean spectra) are displayed as shaded areas. The spectra are similar among themselves within their error bars especially in the peak heights. For a better illustration, only five spectra were selected. All other spectra of original samples are comparable within their error bars.**

Sommer (2020) found for the whole set of our samples that the fraction of BC in PM2.5 (measured with the integrating sphere method) was relatively constant over the whole year. Traffic is one of the major sources of BC in Vienna in summer

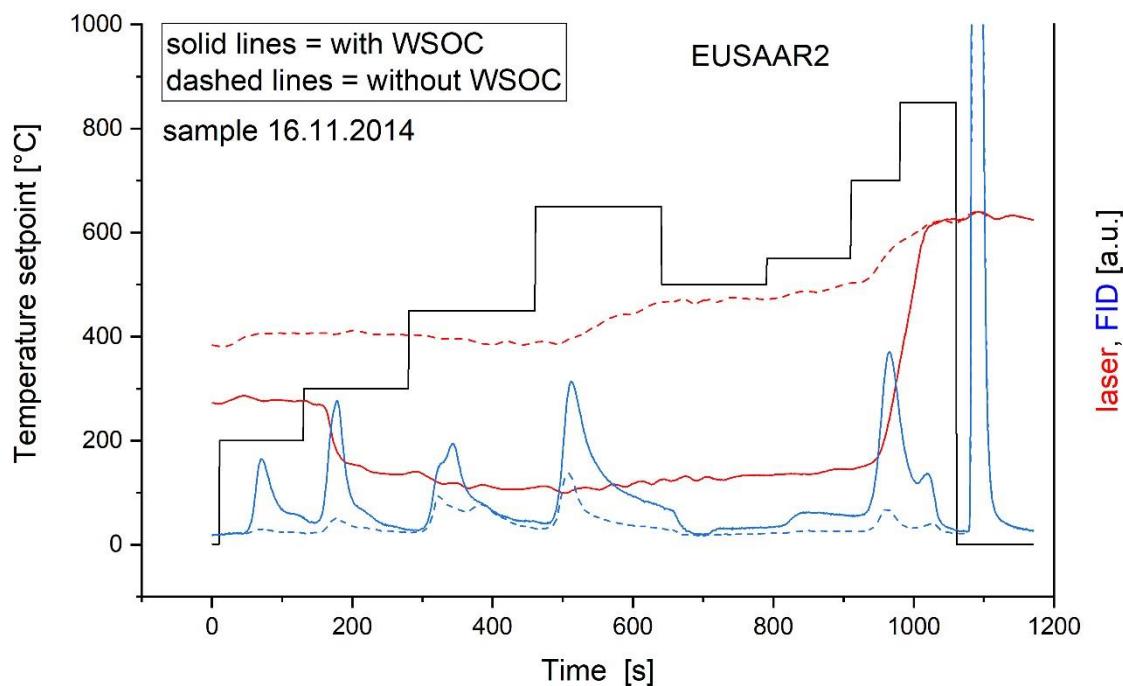
240 as well as in winter. In winter space heating contributes only a small amount to BC because district heating and natural gas are used in most households in Vienna (about 45% each in 2017/18, Statistik Austria, online) followed by electricity (about 6% in 2017/18, Statistik Austria, online). All of these are relatively clean systems regarding particle production. On the other hand wood combustion is used by less than 1% of the households in Vienna, and oil or coal combustion are negligible (Statistik Austria, online). BrC emission factors of wood combustion can be nearly 10 times the emission factors for BC (e.g.

245 Sun et al., 2020). Our winter samples have less than 70% BrC in LAC, therefore also for the samples with the highest amount of BrC, nearly 80% of winter time BC derives from traffic. In 2015 about 57% of all passenger cars and nearly 100% of heavy duty vehicles in Vienna had diesel engines. Diesel fuel accounted for about 80% to total fuel sales in 2015 in Austria (WKO, Mineralölindustrie, online). Traffic is therefore a relatively constant and homogeneous source of carbonaceous aerosol, especially Diesel soot. This could explain the similarity of the Raman spectra of all original samples.

250 Crystallite sizes of the three samples analyzed by TEM are well below 2 nm (mean crystallite size of all three samples: 1.13 ± 0.1 nm). These sizes match with the crystallite sizes of two CAST soots in our recent publication (recalculated: 1.2 ± 0.1 nm for a BC rich and 0.8 ± 0.1 nm for a BrC rich sample; Haller et al., 2019). We therefore assume that the crystallite sizes of the other atmospheric aerosol samples in this study also should be below 2 nm.

3.2 Structural changes of the washed samples

255 Washed and unwashed filter samples were analyzed with EUSAAR2 and Raman spectra were obtained for the non-heated and heated washed and unwashed samples. Figure 2 shows two typical thermograms one of an unwashed and one of a washed part of the same filter. While the transmission laser signal of the unwashed sample decreases during the inert phase, the signal of the washed sample is nearly constant until it starts to increase during the 650°C step in He. The initial laser signal for the washed sample is higher than the one for the original sample indicating that part of the light absorbing material
260 was removed during the washing procedure. Since the samples were shaken in this process also substances which are not soluble in water (e.g. BC) could have been removed mechanically. This was confirmed by integrating sphere measurements: part of BC (about 60%) is absent and BrC is below detection limit also for samples with a high amount of BrC in the unwashed counterpart.



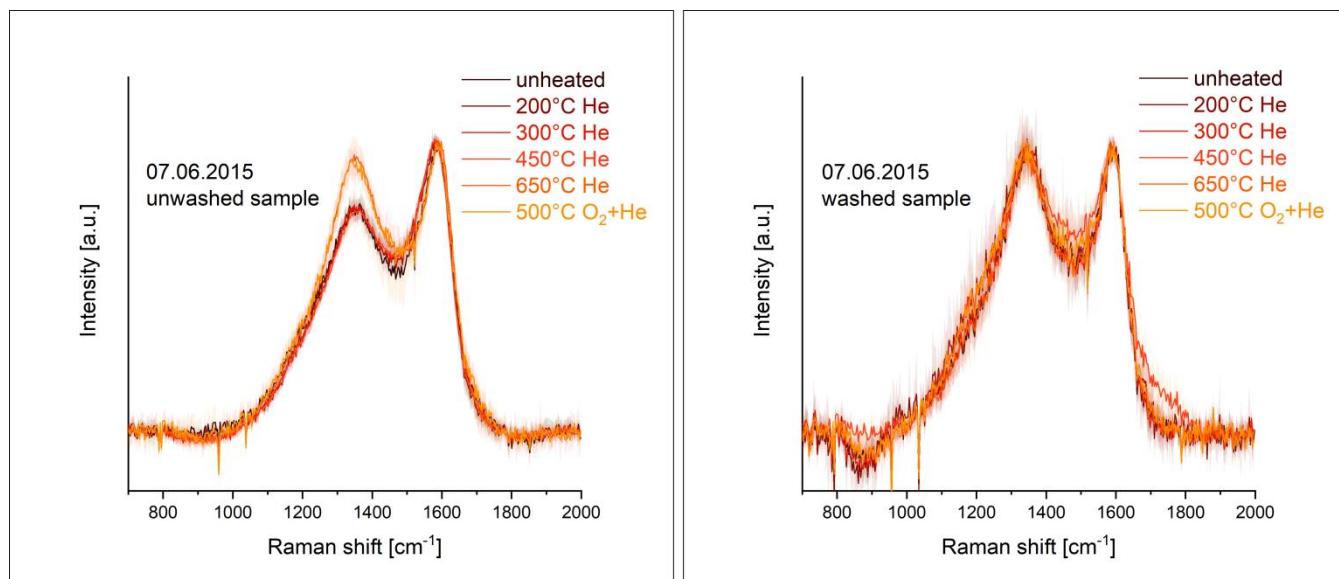
265 **Figure 2: Typical thermogram of a sample with (solid lines) and without (dashed lines) WSOC. The transmission laser signal decreases during the first three temperature levels for the unwashed sample (red solid line) but stays nearly constant for the**

washed sample (red dashed line). The FID signals are lower for all temperature levels for the washed sample (blue dashed line) than for the unwashed sample (blue solid line). The black solid line represents the ideal progress of the sample temperature.

The differences in the laser signals between original and washed samples hold for all of the 8 investigated samples (Supplement S1). These findings are in good agreement with Yu et al. (2002) who found that WSOC accounted for a large contribution (13-66%) to the charring during NIOSH-like heating procedures.

The D/G ratios of the Raman spectra of most unwashed samples increase during the heating procedure above 450°C (Fig 3a and section 3.3.2) indicating an increase of structural ordering according to Ferrari and Robertson (2000). On the other hand the Raman spectra of the washed samples do not change noticeably during the heating process even when their unwashed counterparts have changing Raman signals (Fig. 3) indicating that no increase of structural ordering occurs during heating of the washed samples. This is in good agreement with the findings of Yu et al. (2002) who based their interpretation of charring of WSOC on thermal and optical properties of their heated samples. We can confirm this now by showing that the washed samples indeed do not change their structure (i.e. do not build graphitic like material) during heating. In contrast to cases, where the laser signal decreases despite constant D/G ratios (as described in 3.3.1) it seems to be possible to conclude from an unchanged laser signal to a lack of structural ordering of the material.

The laser signals start to increase during the last temperature step in He (650°C) also for the washed samples. The Raman spectra of these samples show that structural ordering does not increase during the heating procedure and the relatively constant (transmission/reflectance) laser signals during the first three temperature levels suggest that no noticeable pyrolysis occurs. Therefore we assume that some of the light absorbing material which was originally on the filter evolves in the absence of oxygen at 650°C. This could be caused by a premature evolution of EC due to oxygenation with oxygen supplied by mineral oxides or catalysis by metal or other inorganic salts (Chow et al., 2001; Wang et al., 2010).

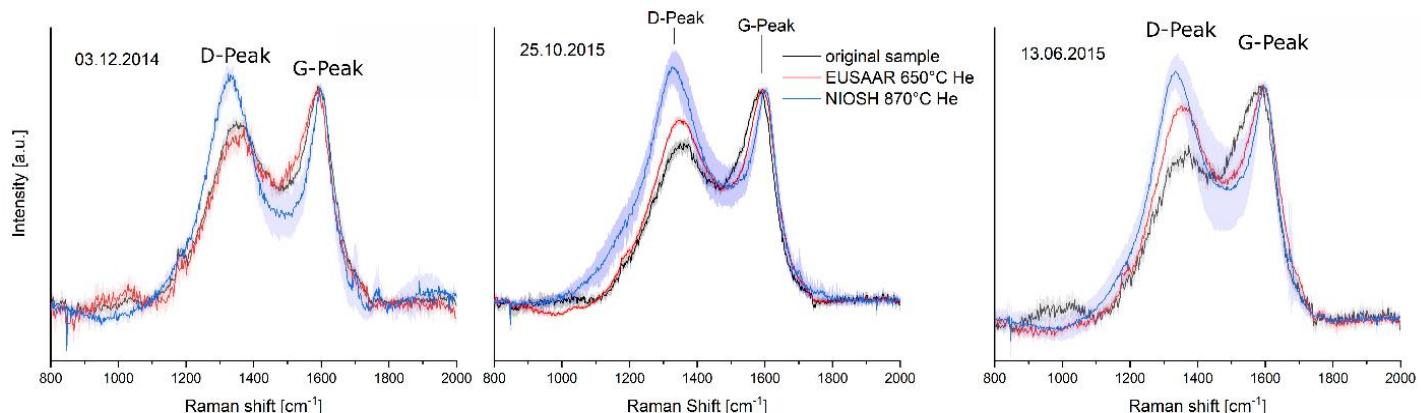


290 **Figure 3: Progress of the Raman spectra during the heating procedure (EUSAAR2) of the washed (left) and unwashed (right) part of a representative atmospheric aerosol sample (07.06.2015). Error bars are shown as shaded areas. The Raman spectra of the unwashed sample change at 650°C in He (late and fast change; see categorization below) while the spectra of the washed sample do not change noticeably during heating following EUSAAR2**

3.3 Analysis of the unwashed (heated and original) samples

295 3.3.1 Comparison of structural changes EUSAAR2 vs. NIOSH870 after the inert phase

A comparison between NIOSH870 and EUSAAR2 protocols was performed for a subset of 6 unwashed samples (5 winter, 1 summer). The D/G ratio increased during the course of the inert phase for both protocols (EUSAAR2 and NIOSH870). However, for three samples (13.06.2015, 03.12.2014, 25.10.2015) out of six, the D/G ratio at the maximum He temperature is significantly larger for NIOSH870 than for EUSAAR2 (Fig 4). For sample 25.10.2015 also D/G of EU500, (i.e. the first 300 temperature step in He+O₂) remains lower than the D/G of NI870. For the other three samples (08.01.2015, 17.02.2015, 09.01.2015), relative D-peak intensities are similar for NI870 and EU650 as well as for EU500.



305 **Figure 4: Comparison of Raman spectra of EU650 (red lines) and NI870 (blue lines) for the three samples where D/G of NI870 is larger than D/G of EU650 (from left to right: 03.12.2014, 25.10.2015 and 13.06.2015). Error bars are shown as shaded areas. The D/G ratio increases for NIOSH870 as well as for EUSAAR2 compared to the spectrum of the original sample (black lines) but the relative D-peak of NI870 is visibly higher than the D-peak of EU650 for all three samples.**

The larger increase of D/G for NI870 for some samples can be interpreted as a stronger increase of structural ordering, when these samples are heated according to the NIOSH870 protocol compared to EUSAAR2.

310 These differences could be caused by a stronger graphitization during the NIOSH870 protocol than during EUSAAR2 at least in three cases out of six. This would mean that PC developed during the heating with different protocols might also have different optical properties which would influence the decrease of the transmission / reflection laser signal during the thermal-optical measurement procedure. This would be in good agreement with Yu et al. (2002), who found that light absorption coefficients of PC formed during a thermal-optical heating procedure are not identical over the whole heating

315 procedure. It is plausible that this might be true also for two protocols with different maximum temperatures in He. The higher D/G ratio for a higher heating temperature is also in accordance with Le et al. (2019) who found increasing D/G ratios between 600°C and 800°C for their OC rich samples heated in N₂.

As suggested by Chow et al. (2001) part of EC could leave the filter already during the 870°C temperature level in helium because of oxygenation or catalysis caused by mineral or other oxides at temperatures >700°C. This premature evolution of 320 relatively structured material could also affect the D/G ratio during the highest inert temperature step in NIOSH870.

Therefore we assume that the higher D/G ratio of NI870 for some samples could be a combination of stronger graphitization caused by the higher temperature and premature oxidation of EC during the 870°C level in NIOSH870.

3.3.2 Structural changes during EUSAAR2

For all of the 21 (unwashed) samples Raman spectra were taken for unheated samples and for samples heated according to 325 the EUSAAR2 protocol. In most cases (17 out of 21) the D/G ratio increases during the heating process indicating an increase of structural ordering. However, distinct changes of the Raman spectra occur at different temperature levels during the heating procedure in the inert phase (450°C – 650°C) or the oxidizing phase (500°C). Moreover in some cases the change happens only at one temperature level or extends over two temperature levels. Therefore the samples were classified into six 330 categories (subsequently referred to as “Raman categories”) depending on the temperature levels where significant changes in the Raman signals occur: The significance was investigated by performing unpaired, two-sided t-tests for all spectra compared to the spectra of the next higher and lower temperature level, respectively, using a significance level α of 5%. Some slight decreases in the D/G peak ratio between EU200 and EU300 are visible but they are not significant except for one sample out of 21 (13.06.2015, $p=0.035$). A possible explanation of this decrease is given by Le et al. (2019) who 335 observed a similar behavior during heating in N₂ and gave as a reason the evaporation of OC volatile components and cyclization of molecular structures. Since this decreasing D/G ratio is not significant in most of our samples we focus here only on increasing D/G ratios. The six categories were defined as follows: (Sample codes are explained in Table 1 and examples for the categories are shown in Fig. 5 a-e.)

340 *Early structural change:* (four samples) D/G ratio of the Raman spectrum obtained for EU450 is significantly higher according to a t-test with $\alpha=5\%$ than D/G ratio of the spectrum for EU300. And D/G ratio of EU650 is significantly higher than D/G of EU450. The D/G ratios do not increase significantly between unheated sample, EU200 and EU300 and between EU650 and EU500 (Fig. 5, a).

345 *Late structural change:* (two samples) D/G ratio of the Raman spectrum of EU650 is significantly higher than D/G of the Raman spectrum of EU450 and D/G of the Raman spectrum of EU500 is significantly higher than D/G of EU650. The D/G ratio does not increase significantly between the unheated sample, EU200, EU300, EU450 and EU650 (Fig. 5, b).

350 *No change:* (four samples) All mean spectra lie within the error bars of the other spectra and D/G ratios do not change significantly (Fig. 5, c).

355 *Fast and early structural change:* (one sample) D/G ratio of the Raman spectrum of EU450 is significantly higher than D/G of EU300 and D/G does not decrease significantly between EU200 and EU300 as well as between EU450, EU650 and EU500 (Fig. 5, d).

360 *Fast and late structural change:* (six samples) D/G ratio of the Raman spectrum of EU650 is significantly higher than D/G of EU450 and D/G does not increase significantly between EU200, EU300 and EU450 as well as between EU650 and EU500 (Fig. 5, e).

365 *Undefined:* (four samples) Changes of the Raman spectra cannot be assigned unambiguously to one of the other five categories as a consequence of too strong noise signals.

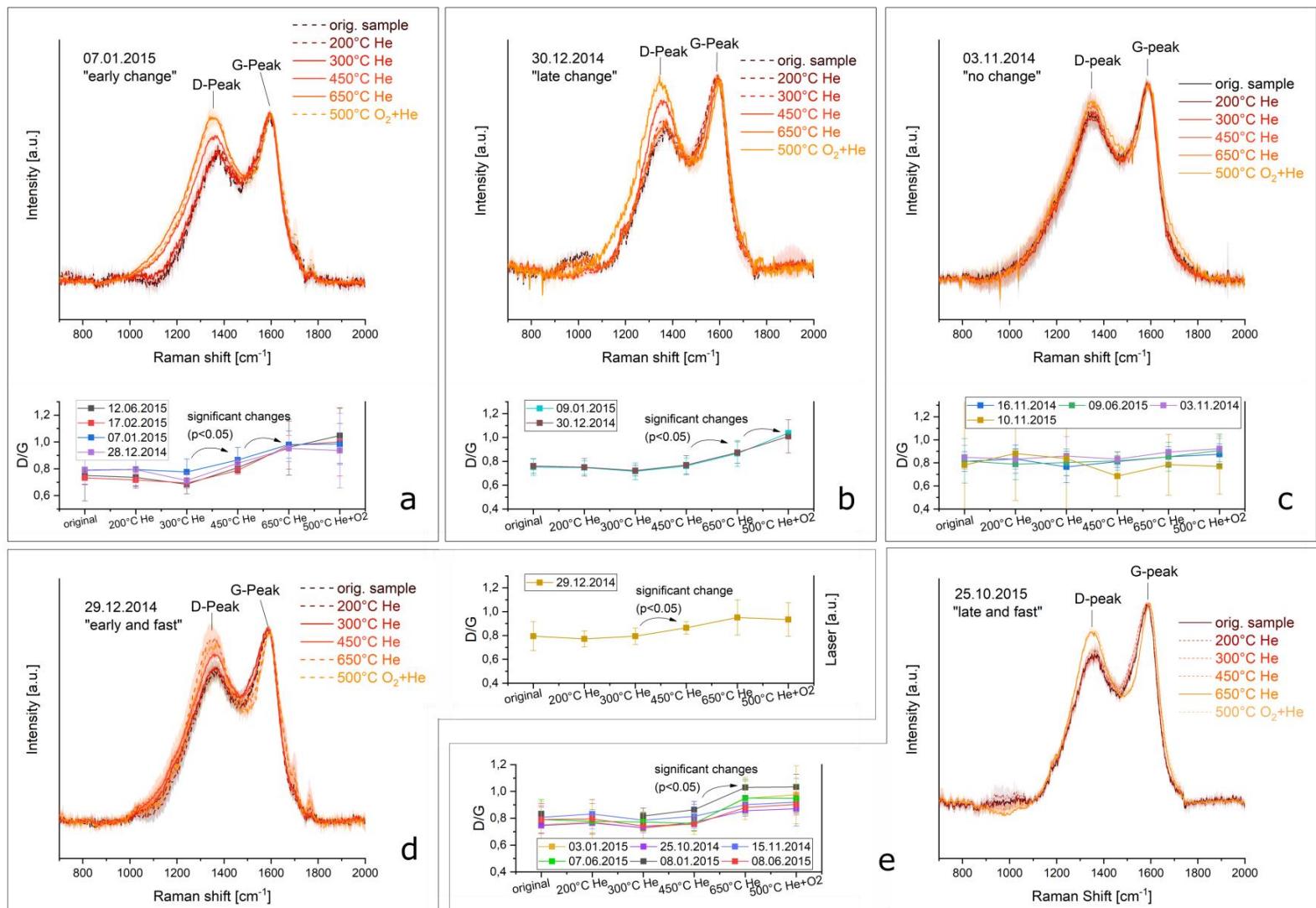


Figure 5 a-e: Examples of Raman spectra for heated and unheated filter samples assigned to the categories "early structural change" (a), "late structural change" (b), "no structural change" (c), "early and fast structural change" (d) and "late and fast structural change" (e). Error bars are shown as shaded areas. The changes of D/G ratio with heating temperature are shown for all samples assigned to the respective category in an additional image. Arrows indicate the temperature levels where significant changes were identified according to a t-test ($\alpha=5\%$).

The changes of peak ratios in dependence of the temperature level are shown in Figure 5 for all samples of each category as
370 well as examples for the changes in the Raman spectra. It is remarkable that the spectra of the heated samples differ and show different progresses over the heating procedure (Fig. 5) although the Raman spectra of the original (unheated) samples are comparable. Therefore we conclude that the differences in the progress of restructuring might not be caused by the degree of structural ordering of the original samples or at least cannot be simply predicted by their Raman spectra.

3.3.3 Structural changes and laser signals (EUSAAR2)

375 During the thermal-optical analysis procedure with EUSAAR2 as well as with NIOSH870 decreasing (transmission and reflection) laser signals were observed during the inert phase for all unwashed samples. The signals started to increase again during the last temperature level in He for both protocols. This increase was also described by other authors (e.g. Chow et al., 2004; Yu et al., 2002).

For a more detailed analysis of the progress of the laser signals we focus on EUSAAR2 here. During the heating following
380 EUSAAR2, the reflection laser signals decrease in more or less pronounced steps at each temperature level in the He phase and no distinct differences between the samples and therefore no relationship between the changes of the Raman spectra and the reflection laser signals are visible. For the transmission signals two different types of decrease can be observed: For some samples the signal decreases with a pronounced step during the 300°C temperature level in He. For other samples the transmission laser signal shows a relatively steady decrease before it reaches its minimal value. However, no relation
385 between the slope of the decrease and the Raman categories is visible. The diagrams can be found in the supplementary material (S2 and S3).

Figure 6 shows a comparison of the D/G ratio and the value of the transmission laser signal at the end of each temperature level in EUSAAR2 sorted by the Raman categories. It is obvious that the decrease of the transmission signal is not related to the increase of the D/G ratio. Samples with similar behavior of the Raman signals can have different curves of the
390 transmission laser signals and vice versa. It is also noteworthy that the strongest decreases of the transmission laser signals do not occur at the temperature levels where changes in the Raman spectra are observable (450°C, 650°C). We found similar effects in our earlier study (Haller et al., 2019) where we compared UV-VIS spectra of heated (NIOSH870 protocol) CAST soot samples with the changes of the respective Raman spectra: The UV-VIS spectra started to change at lower temperatures while the Raman spectra changed only at the highest temperature in He (870°C).

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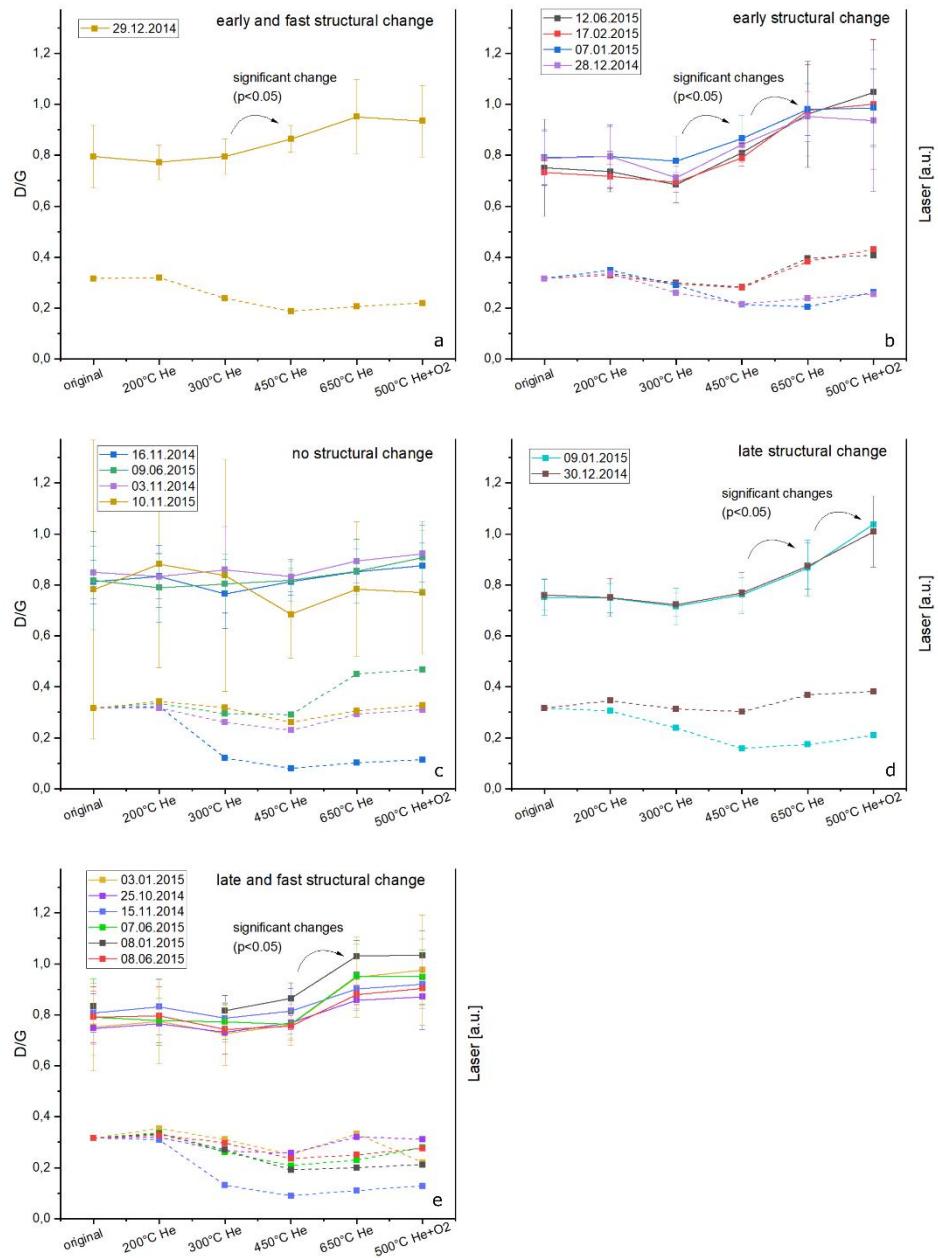
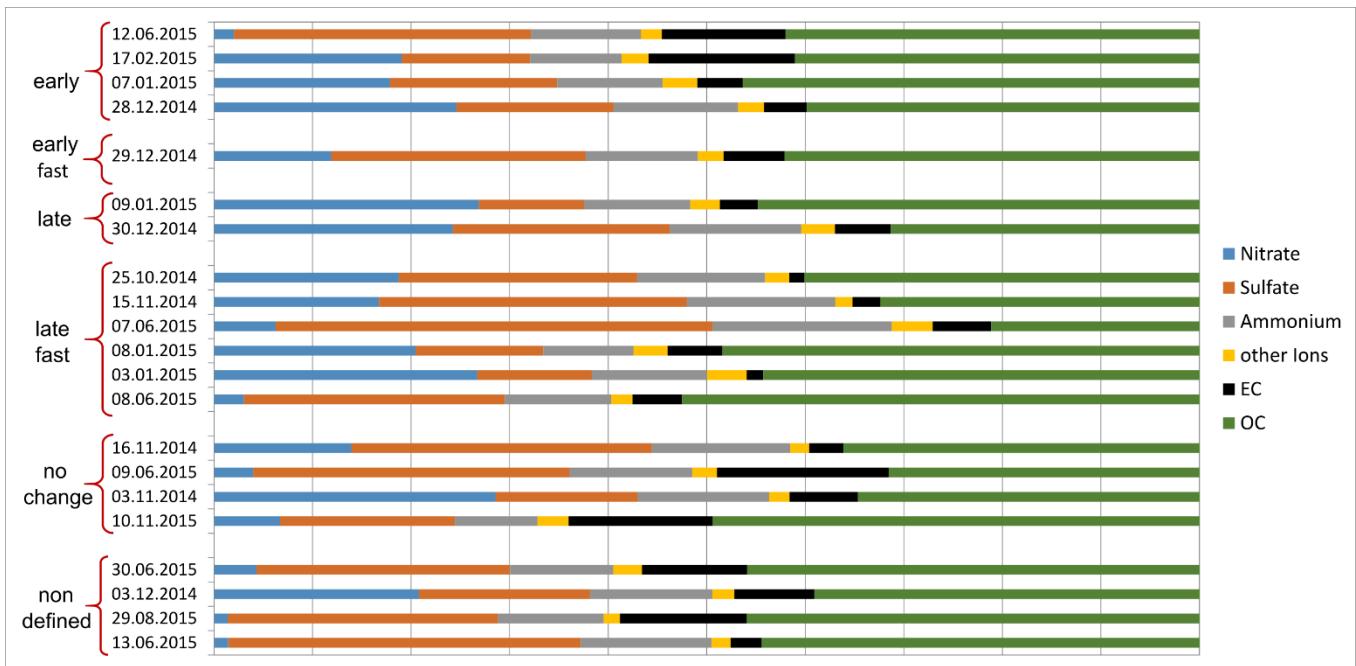


Figure 6 a-e: Comparison of the changes of D/G peak ratios (solid lines, left axis) and transmission laser signals (dashed lines, right axis) at the end of each temperature level sorted by Raman categories. Standard deviations of D/G are calculated from the standard deviations of the Raman spectra. All laser signals are normalized to the signal at the end of the measurement procedure (i.e. signal of the clean filter). The curves for the different filters are shifted at the y-axis so that all curves start at the same point (original sample). It is visible that changes in D/G and changes in laser signal are not related.

This lack of correlation between laser signal and D/G ratio is most striking for the four samples with no change in the Raman spectra of the heated samples (Fig. 6c). Although the Raman spectra do not indicate an increase of structural ordering, the (transmission, reflection) laser signals decrease during the heating procedure in He. Therefore we conclude that the 405 formation or increase of graphitic like structures cannot be the only reason for the darkening (i.e. stronger absorption of the red laser signal). Our data give evidence that non-graphitic material with smaller optical band gaps and therefore higher absorbance at 635 nm than in the unheated material might be formed. This corresponds also to the findings of Yu et al. (2002) who assume that light absorbing intermediate OC products are formed during the inert phase of the thermal-optical 410 measurement procedures they used (a NIOSH method and other experimental temperature protocols). Le et al. (2019) performed a detailed analysis of several heated CAST soots and suggest a removal of C-H “out-of-plane” bonds around 350-400°C, a decomposition of oxygen containing surface groups (C-O, C=O) or a decrease of cross-linkages between polyaromatic units and carbon chains. All of these transformations could influence the optical properties without an increase of graphitic like structures.

3.3.4 Possible relations of structural changes with other parameters

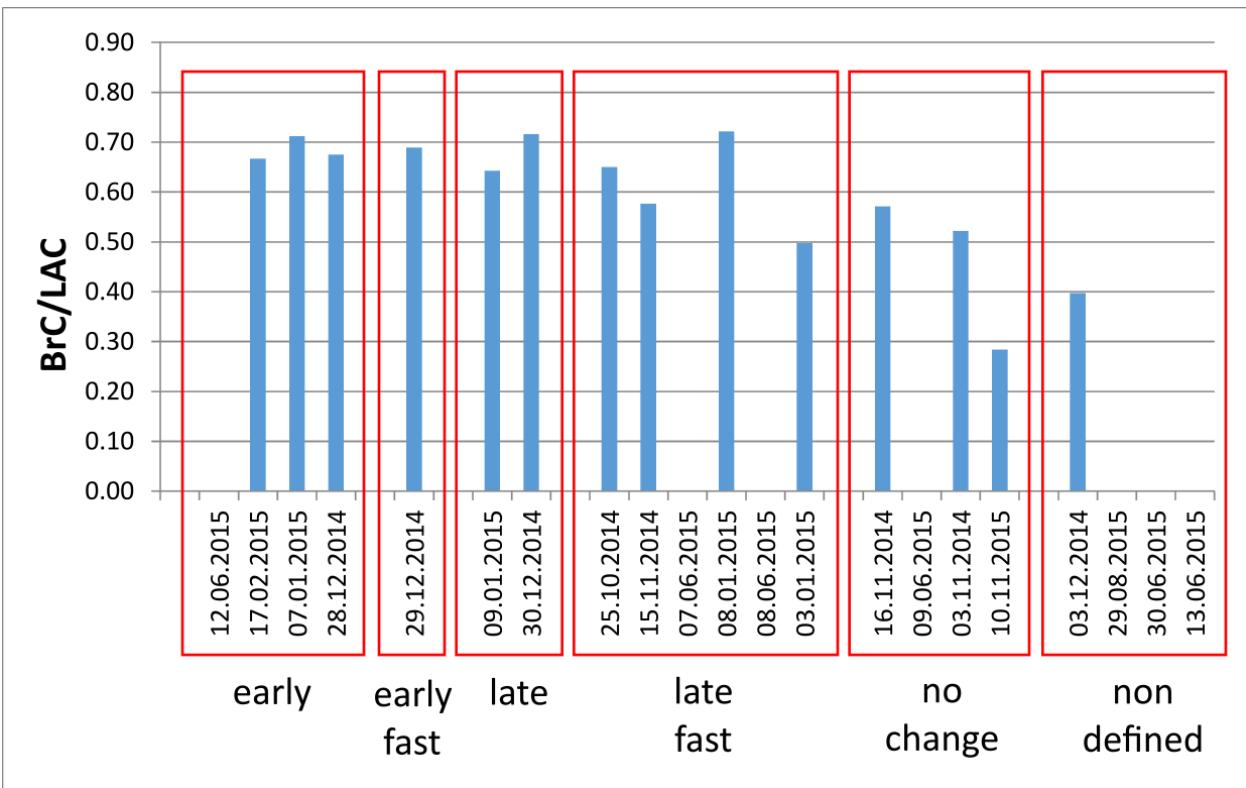
415 Generally the charring behavior of aerosol samples or the reliability of thermal-optical measurements is often reported to be influenced by the composition of the aerosol (e.g. Wang et al., 2010; Yu et al., 2002; Reisinger et al., 2008; Wonaschuetz et al., 2009). Since our samples show differing charring behaviors, we investigated possible relations to the aerosol composition using the unwashed samples. We compared the progress of structural changes (expressed by the assignment to the 420 Raman categories) to the chemical composition of the non-carbonaceous matter and to the composition of the carbonaceous matter (LAC/BC, EC/OC). The ionic composition of the samples sorted by their Raman categories is shown in Fig. 7.



425 **Figure 7: Fractions of analyzed ions, elemental carbon (EC) and organic carbon (OC) sorted by the categories representing the behavior of the Raman spectra during the thermal optical heating procedure. The contribution of minor ions to the category “other ions” can be found in the supplement.**

No relationship between the ionic components and the progress of restructuring of the material is visible. This could be explained by the study by Yu et al. (2002) who showed that one and the same compound can enhance or reduce charring depending on the chemistry of the organic substances present in the sample. Although they analyzed only three standards for 430 organic carbon, it is possible that similar counteracting influences could be found also for other organic substances which could suppress a possible relationship between charring and the inorganic material within a sample.

The progress of structural changes is also unrelated to the BrC/LAC ratio (Fig. 8). Samples with BrC below detection limit are found in nearly all the Raman categories. Especially the samples in the category “no change” during heating do not have appreciably lower BrC/LAC ratios than others. This is quite interesting, because previous studies in the Austrian aerosol 435 (e.g. Reisinger et al., 2008; Wonaschuetz et al., 2009) found larger discrepancies between EC values obtained with different thermal-optical methods for atmospheric aerosol samples with high contributions of BrC to LAC, up to about 0.9 (Wonaschuetz et al., 2009) and 0.75 (Reisinger et al., 2008). These values are comparable to the BrC/LAC of about 0.7 in our study. At the time of the earlier studies it was not possible to determine whether these discrepancies were due to a higher tendency of BrC to char (and therefore influence the EC/OC split) or from other BrC-related processes. Our findings suggest 440 that the relative amount of BrC has no influence on the increase of structural ordering during the heating procedure at least for the samples analyzed here. The discrepancies between different thermal-optical methods for BrC-rich samples might therefore have other reasons than a restructuring of BrC in terms of increased structural ordering.



445 Figure 8: Comparison of BrC/LAC with the categories representing the behavior of the Raman signal during the heating procedure of the thermal-optical analysis. For samples with BrC/LAC=0 (i.e. no bar in the diagram) BrC was below detection limit (<12.7 $\mu\text{g}/\text{cm}^2$).

4 Conclusion

450 The aim of this study was to investigate structural changes of carbonaceous matter during thermal-optical analyses. Two different thermal-optical measurement protocols with different maximum temperatures in the He phase (EUSAAR2, NIOSH870) were used to show the progress of restructuring of carbonaceous material during heating. Raman spectra of unheated and heated samples were taken for 21 atmospheric aerosol samples. A subset of these samples was washed with water to remove WSOC and other water soluble constituents. The transmission laser signal patterns during the EUSAAR2 protocol and the progress of structural changes were compared for the washed samples and their unwashed counterparts.

455 We found that the laser signals did not decrease during the heating procedure of the washed samples but decreased when the unwashed samples were heated. This is in good agreement with Yu et al. (2002) who observed that water soluble material contributed most to charring. The D/G peak ratios of the Raman spectra of the washed samples did not change appreciably during heating in He following the EUSAAR2 protocol. We therefore infer that the relatively constant transmission laser signal can indeed be interpreted as a lack of graphitization.

460 A comparison between EUSAAR2 and NIOSH870 was performed for six unwashed samples. EUSAAR2 is used in Europe as a standard method to monitor EC and OC (Brown et al., 2017). One of the arguments for the introduction of EUSAAR2 was that less PC is formed during EUSAAR2 in comparison with other (NIOSH-like) temperature protocols (Cavalli et al., 2010). However, this argument was solely based on the decrease of the transmission laser signal. Structural changes of the material were not investigated. We found now that the D/G peak ratios of the Raman spectra can be higher for samples
465 heated to the maximum He temperature of NIOSH870 compared to samples heated to the maximum He temperature of EUSAAR2 at least for some samples, which means that PC formed during NIOSH870 might have a higher degree of structural ordering in these cases.

When comparing the Raman spectra for 21 samples heated following the EUSAAR2 protocol we found a rather complex behaviour of structural changes. Since changes in the Raman signals (i.e. the ratio of the peak intensities) occur at different
470 temperature levels, we defined six “Raman categories” of samples depending on the temperature levels, where significant changes in the Raman spectra occur and how fast they proceed. We tried to find similarities for samples with similar progress of structural changes (i.e. the attribution to one of the six Raman categories) and chemical composition (BC, BrC, major ions), but did not find any relationship. This is in good agreement with Wang et al. (2010) and Yu et al. (2002) who found that non-carbonaceous components have conflicting influences on the charring behavior during thermal-optical
475 analyses.

We also investigated the progress of the thermal-optical transmission laser signals in relation to the Raman categories corresponding to the temperature levels where visible structural changes occur. It was obvious that the decrease of the transmission signal did not occur at the same temperature levels where structural changes in terms of increased ordering were found. This lack of a relation is most prominent for samples where the D/G peak ratios of the Raman spectra do not
480 change during the whole heating procedure in He. Although the Raman spectra do not indicate an increase of structural ordering, the transmission (and reflection) laser signals still decrease during the He phase. Moreover no noticeable differences can be seen between the transmission laser signal patterns of samples where no increase of ordering occurs in comparison with samples where the Raman spectra change during heating.

We therefore conclude that some of the processes that lead to a darkening of the sample (i.e. more absorption at 635 nm) do
485 not affect the relative peak heights of the Raman spectra which means that structural ordering does not increase (similar findings as well as their interpretation were described for CAST soot samples in our previous publication; Haller et al., 2019). Processes suggested by various authors include e. g. the separation of oxygen and hydrogen at temperatures above 250°C in He (Petzold et al., 2013; Chow et al., 2004, Le et al., 2019), the decomposition of oxygen containing surface groups (C-O, C=O) or a decrease of cross-linkages between polycyclic aromatic units and carbon chains (Le et al., 2019). All of
490 these transformations could affect the optical properties of the material without affecting the D/G peak ratio of the Raman spectra. These findings are in good agreement with Yu et al. (2002) who describe a possible formation of light absorbing intermediate (non-graphitic) OC products during the inert phase of thermal-optical analyses.

In the literature the darkening of a sample during thermal-optical analysis is often explained by “charring” or “pyrolysis” although exact descriptions of the underlying structural changes are not given. Based on our findings, it is important to point
495 out that what is called “charring” or “pyrolysis” is not necessarily an increase of structural ordering. Particularly the group of samples where the D/G peak ratio does not change over the whole heating procedure gives evidence that the darkening of the sample cannot be understood as building of graphitic like structures in all cases. We therefore suggest that the widely used terms “charring” and “pyrolysis” should be used carefully when the darkening of a sample during thermal-optical measurement procedures is interpreted.

500 In summary we can say that restructuring of carbonaceous matter during thermal-optical analysis is very complex and not easily predictable for atmospheric aerosol samples. At least for the washed samples the constant transmission / reflection signal indicates an absence of graphitization, but for all samples a decrease of the transmission or reflection laser signal during the heating procedures is not a reliable indicator for an increase of structural ordering since other processes can also lead to a darkening of the material.

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Data availability. Data can be accessed by contacting the corresponding author.

Author contributions. TH performed the conceptualization of the experimental setup, most measurements (except TEM, BrC/BC, ion chromatography), the data evaluation and interpretation and prepared the manuscript. ES performed the
510 BrC/BC measurements. TS performed the ion chromatography measurements and wrote the text in the measurement section. CR performed the TEM measurements and wrote the TEM section. AW supervised the BrC/BC measurements. AKG contributed to discussions and provided input for the ion chromatography and to the manuscript in general. HG contributed to discussions and provided input for the Raman measurements. RH performed the conceptualization and supervision of the study, contributed to discussions, and provided extensive input to the text.

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Competing interests: The authors declare that they have no conflict of interest.

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525 **References**

Anderson, J. O., Thundiyil, J. G., Stolbach, A.: Clearing the Air: A Review of the Effects of Particulate Matter Air Pollution on Human Health, *J. Med. Toxicol.*, **8** (2), 166–175, DOI: 10.1007/s13181-011-0203-1, 2012.

Andreae, M. O., Gelencsér, A.: Black carbon or brown carbon? The nature of light-absorbing carbonaceous aerosols, *Atmos. Chem. Phys.*, **6** (10), 3131–3148, DOI: 10.5194/acp-6-3131-2006, 2006.

530 Birch, M. E., Cary, R. A.: Elemental Carbon-Based Method for Monitoring Occupational Exposures to Particulate Diesel Exhaust, *Aerosol Science and Technology*, **25** (3), 221–241, DOI: 10.1080/02786829608965393, 1996.

Bladt, H., Ivleva, N. P., Niessner, R.: Internally mixed multicomponent soot: Impact of different salts on soot structure and thermo-chemical properties, *Journal of Aerosol Science*, **70**, 26–35, DOI: 10.1016/j.jaerosci.2013.11.007, 2014.

535 Bond, T. C., Bergstrom, R. W.: Light Absorption by Carbonaceous Particles: An Investigative Review, *Aerosol Science and Technology*, **40** (1), 27–67, DOI: 10.1080/02786820500421521, 2006.

Brown, R. J. C., Beccaceci, S., Butterfield, D. M., Quincey, P. G., Harris, P. M., Maggos, T., Panteliadis, P., John, A., Jedynska, A., Kuhlbusch, T. A. J., Putaud, J.-P., Karanasiou, A.: Standardisation of a European measurement method for organic carbon and elemental carbon in ambient air: results of the field trial campaign and the determination of a measurement uncertainty and working range, *Environ. Sci.: Processes Impacts*, **19** (10), 1249–1259, DOI: 10.1039/C7EM00261K, 2017.

540 Cavalli, F., Viana, M., Yttri, K. E., Genberg, J., Putaud, J.-P.: Toward a standardised thermal-optical protocol for measuring atmospheric organic and elemental carbon: the EUSAAR protocol, *Atmos. Meas. Tech.*, **3** (1), 79–89, DOI: 10.5194/amt-3-79-2010, 2010.

Cheng, Y., Duan, F.-k., He, K.-b., Du, Z.-y., Zheng, M., Ma, Y.-l.: Intercomparison of thermal-optical method with different temperature protocols: Implications from source samples and solvent extraction, *Atmospheric Environment*, **61**, 453–462, DOI: 545 10.1016/j.atmosenv.2012.07.066, 2012.

Chow, J. C., Watson, J. G., Crow, D., Lowenthal, D. H., Merrifield, T.: Comparison of IMPROVE and NIOSH Carbon Measurements, *Aerosol Science and Technology*, **34** (1), 23–34, DOI: 10.1080/027868201300081923, 2001.

550 Chow, J. C., Watson, J. G., Chen, L.-W. A., Arnott, W. P., Moosmüller, H., Fung, K.: Equivalence of Elemental Carbon by Thermal/Optical Reflectance and Transmittance with Different Temperature Protocols, *Environ. Sci. Technol.*, **38** (16), 4414–4422, DOI: 10.1021/es034936u, 2004.

Commodo, M., Falco, G. de, Larciprete, R., D'Anna, A., Minutolo, P.: On the hydrophilic/hydrophobic character of carbonaceous nanoparticles formed in laminar premixed flames, *Experimental Thermal and Fluid Science*, **73**, 56–63, DOI: 10.1016/j.expthermflusci.2015.09.005, 2016.

Draxler, R. R.: HYSPLIT4 User's Guide. NOAA Tech. Memo. ERL ARL-230, NOAA Air Resources Laboratory, Silver Spring, MD, 555 1999.

Draxler, R. R., Hess, G. D.: Description of the HYSPLIT4 modeling system: NOAA Tech. Memo. ERL ARL-224, NOAA Air Resources Laboratory, Silver Spring, MD, 1997.

Draxler, R.R. and Hess, G.D.: An Overview of the HYSPLIT_4 Modelling System for Trajectories, Dispersion on and Deposition, *Australian Meteorological Magazine*, **47**, 295–308, 1998.

560 EAA: Air quality in Europe - 2017 report, EAA report, Copenhagen, 2017.

Ess, M. N., Ferry, D., Kireeva, E. D., Niessner, R., Ouf, F.-X., Ivleva, N. P.: In situ Raman microspectroscopic analysis of soot samples with different organic carbon content: Structural changes during heating, *Carbon*, **105**, 572–585, DOI: 10.1016/j.carbon.2016.04.056, 2016.

565 Ferrari, A. C., Robertson, J.: Interpretation of Raman spectra of disordered and amorphous carbon, *Phys. Rev. B*, **61** (20), 14095–14107, DOI: 10.1103/PhysRevB.61.14095, 2000.

Fultz, B. and Howe, J. M.: Transmission electron microscopy and diffractometry of materials, *Physics and astronomy online library*, Springer, Berlin, 2001.

570 Gammer, C., Mangler, C., Rentenberger, C., Kärnthal, H. P.: Quantitative local profile analysis of nanomaterials by electron diffraction, *Scripta Materialia*, **63** (3), 312–315, DOI: 10.1016/j.scriptamat.2010.04.019, 2010.

Gao, S., Hegg, D. A., Hobbs, P. V., Kirchstetter, T. W., Magi, B. I., Sadilek, M.: Water-soluble organic components in aerosols associated with savanna fires in southern Africa: Identification, evolution, and distribution, *J. Geophys. Res.*, **108** (D13), n/a-n/a, DOI: 10.1029/2002JD002324, 2003.

Haller, T., Rentenberger, C., Meyer, J. C., Felgitsch, L., Grothe, H., Hitzenberger, R.: Structural changes of CAST soot during a 575 thermal-optical measurement protocol, *Atmos. Meas. Tech.*, **12** (7), 3503–3519, DOI: 10.5194/amt-12-3503-2019, 2019.

Han, Y., Cao, J., Chow, J. C., Watson, J. G., An, Z., Jin, Z., Fung, K., Liu, S.: Evaluation of the thermal/optical reflectance method for discrimination between char- and soot-EC, *Chemosphere*, **69** (4), 569–574, DOI: 10.1016/j.chemosphere.2007.03.024, 2007.

Häusler, T., Gebhardt, P., Iglesias, D., Rameshan, C., Marchesan, S., Eder, D., Grothe, H.: Ice Nucleation Activity of Graphene and Graphene Oxides, *J. Phys. Chem. C*, **122** (15), 8182–8190, DOI: 10.1021/acs.jpcc.7b10675, 2018.

580 Highwood, E. J., Kinnersley, R. P.: When smoke gets in our eyes: The multiple impacts of atmospheric black carbon on climate, air quality and health, *Environment International*, **32** (4), 560–566, DOI: 10.1016/j.envint.2005.12.003, 2006.

Hitzenberger, R., Petzold, A., Bauer, H., Ctyroky, P., Pouresmaeil, P., Laskus, L., Puxbaum, H.: Intercomparison of Thermal and 585 Optical Measurement Methods for Elemental Carbon and Black Carbon at an Urban Location, *Environ. Sci. Technol.*, **40** (20), 6377–6383, DOI: 10.1021/es051228v, 2006.

590 Hitzenberger, R., Tohno, S.: Comparison of black carbon (BC) aerosols in two urban areas – concentrations and size distributions, *Atmospheric Environment*, **35** (12), 2153–2167, DOI: 10.1016/S1352-2310(00)00480-5, 2001.

Ikhenazene, R., Pirim, C., Noble, J. A., Irimie, C., Carpentier, Y., Ortega, I. K., Ouf, F.-X., Focsa, C., Chazallon, B.: Ice Nucleation Activities of Carbon-Bearing Materials in Deposition Mode: From Graphite to Airplane Soot Surrogates, *J. Phys. Chem. C*, **124** (1), 489–503, DOI: 10.1021/acs.jpcc.9b08715, 2020.

Ivleva, N. P., Messerer, A., Yang, X., Niessner, R., Pöschl, U.: Raman Microspectroscopic Analysis of Changes in the Chemical 595 Structure and Reactivity of Soot in a Diesel Exhaust Aftertreatment Model System, *Environ. Sci. Technol.*, **41** (10), 3702–3707, DOI: 10.1021/es0612448, 2007.

Jankowski, N., Schmidl, C., Marr, I. L., Bauer, H., Puxbaum, H.: Comparison of methods for the quantification of carbonate carbon in atmospheric PM10 aerosol samples, *Atmospheric Environment*, **42** (34), 8055–8064, DOI: 595 10.1016/j.atmosenv.2008.06.012, 2008.

Kirchstetter, T. W., Novakov, T., Hobbs, P. V.: Evidence that the spectral dependence of light absorption by aerosols is affected by organic carbon, *J. Geophys. Res.*, **109** (D21), n/a-n/a, DOI: 10.1029/2004JD004999, 2004.

Knauer, M., Schuster, M. E., Su, D., Schlögl, R., Niessner, R., Ivleva, N. P.: Soot structure and reactivity analysis by Raman 600 microspectroscopy, temperature-programmed oxidation, and high-resolution transmission electron microscopy, *The journal of physical chemistry. A*, **113** (50), 13871–13880, DOI: 10.1021/jp905639d, 2009.

Le, K. C., Pino, T., van Pham, T., Henriksson, J., Török, S., Bengtsson, P.-E.: Raman spectroscopy of mini-CAST soot with various fractions of organic compounds: Structural characterization during heating treatment from 25 °C to 1000 °C, *Combustion and Flame*, **209**, 291–302, DOI: 10.1016/j.combustflame.2019.07.037, 2019.

605 Lukács, H., Gelencsér, A., Hammer, S., Puxbaum, H., Pio, C., Legrand, M., Kasper-Giebl, A., Handler, M., Limbeck, A., Simpson, D., Preunkert, S.: Seasonal trends and possible sources of brown carbon based on 2-year aerosol measurements at six sites in Europe, *J. Geophys. Res.*, 112 (D23), 1052, DOI: 10.1029/2006JD008151, 2007.

Müller, K., Spindler, G., Maenhaut, W., Hitzenberger, R., Wiprecht, W., Baltensperger, U., Brink, H. ten: INTERCOMP2000, a campaign to assess the comparability of methods in use in Europe for measuring aerosol composition, *Atmospheric Environment*, 38 (38), 6459–6466, DOI: 10.1016/j.atmosenv.2004.08.031, 2004.

610 Panteliadis, P., Hafkenscheid, T., Cary, B., Diapouli, E., Fischer, A., Favez, O., Quincey, P., Viana, M., Hitzenberger, R., Vecchi, R., Saraga, D., Sciai, J., Jaffrezo, J. L., John, A., Schwarz, J., Giannoni, M., Novak, J., Karanasiou, A., Fermo, P., Maenhaut, W.: ECOC comparison exercise with identical thermal protocols after temperature offset correction – instrument diagnostics by in-depth evaluation of operational parameters, *Atmos. Meas. Tech.*, 8 (2), 779–792, DOI: 10.5194/amt-8-779-2015, 2015.

Partanen, A.-I., Landry, J.-S., Matthews, H. D.: Climate and health implications of future aerosol emission scenarios, *Environ. Res. Lett.*, 13 (2), 24028, DOI: 10.1088/1748-9326/aaa511, 2018.

Petzold, A., Ogren, J. A., Fiebig, M., Laj, P., Li, S.-M., Baltensperger, U., Holzer-Popp, T., Kinne, S., Pappalardo, G., Sugimoto, N., Wehrli, C., Wiedensohler, A., Zhang, X.-Y.: Recommendations for reporting "black carbon" measurements, *Atmos. Chem. Phys.*, 13 (16), 8365–8379, DOI: 10.5194/acp-13-8365-2013, 2013.

Pierce, J. R., Chen, K., Adams, P. J.: Contribution of primary carbonaceous aerosol to cloud condensation nuclei: processes and uncertainties evaluated with a global aerosol microphysics model, *Atmos. Chem. Phys.*, 7 (20), 5447–5466, DOI: 10.5194/acp-7-5447-2007, 2007.

Pöschl, U.: Atmospheric Aerosols: Composition, Transformation, Climate and Health Effects, *Angew. Chem. Int. Ed.*, 44 (46), 7520–7540, DOI: 10.1002/anie.200501122, 2005.

625 Prenni, A. J., DeMott, P. J., Sullivan, A. P., Sullivan, R. C., Kreidenweis, S. M., Rogers, D. C.: Biomass burning as a potential source for atmospheric ice nuclei: Western wildfires and prescribed burns, *Geophys. Res. Lett.*, 39 (11), n/a-n/a, DOI: 10.1029/2012GL051915, 2012.

Reisinger, P., Wonaschütz, A., Hitzenberger, R., Petzold, A., Bauer, H., Jankowski, N., Puxbaum, H., Chi, X., Maenhaut, W.: Intercomparison of Measurement Techniques for Black or Elemental Carbon Under Urban Background Conditions in Wintertime: Influence of Biomass Combustion, *Environ. Sci. Technol.*, 42 (3), 884–889, DOI: 10.1021/es0715041, 2008.

630 Rosen, H., Novakov, T.: Identification of primary particulate carbon and sulfate species by Raman spectroscopy, *Atmospheric Environment* (1967), 12 (4), 923–927, DOI: 10.1016/0004-6981(78)90031-8, 1978.

Sadezky, A., Muckenhuber, H., Grothe, H., Niessner, R., Pöschl, U.: Raman microspectroscopy of soot and related carbonaceous materials: Spectral analysis and structural information, *Carbon*, 43 (8), 1731–1742, DOI: 10.1016/j.carbon.2005.02.018, 2005.

Schmid, J., Grob, B., Niessner, R., Ivleva, N. P.: Multiwavelength Raman Microspectroscopy for Rapid Prediction of Soot Oxidation Reactivity, *Anal. Chem.*, 83 (4), 1173–1179, DOI: 10.1021/ac102939w, 2011.

Sommer, E.: Long Term Trends of Black and Brown Carbon Concentrations in the Urban Aerosol of Vienna, Master Thesis (in preparation), University of Vienna, Vienna, 2020.

Spracklen, D. V., Carslaw, K. S., Pöschl, U., Rap, A., Forster, P. M.: Global cloud condensation nuclei influenced by carbonaceous combustion aerosol, *Atmos. Chem. Phys.*, 11 (17), 9067–9087, DOI: 10.5194/acp-11-9067-2011, 2011.

640 Statistik Austria:
https://www.statistik.at/web_de/statistiken/energie_umwelt_innovation_mobilitaet/energie_und_umwelt/energie/energieeinsatz_der_haushalte/index.html, last access 18.02.2021.

Sun, J., Zhang, Y., Zhi, G., Hitzenberger, R., Jin, W., Chen, Y., Wang, L., Tian, C., Li, Z., Chen, R., Xiao, W., Cheng, Y., Yang, W., Yao, L., Cao, Y., Huang, D., Qiu, Y., Xu, J., Xia, X., Yang, X., Zhang, X., Zong, Z., Song, Y., Wu, C.: Brown carbon's emission factors and

645 optical characteristics in household biomass burning: developing a novel algorithm for estimating the contribution of brown carbon, *Atmos. Chem. Phys.*, 21 (4), 2329–2341, DOI: 10.5194/acp-21-2329-2021, 2021.

Subramanian, R., Khlystov, A. Y., Robinson, A. L.: Effect of Peak Inert-Mode Temperature on Elemental Carbon Measured Using Thermal-Optical Analysis, *Aerosol Science and Technology*, 40 (10), 763–780, DOI: 10.1080/02786820600714403, 2006.

ten Brink, H., Maenhaut, W., Hitzenberger, R., Gnauk, T., Spindler, G., Even, A., Chi, X., Bauer, H., Puxbaum, H., Putaud, J.-P.,
650 Tursic, J., Berner, A.: INTERCOMP2000: the comparability of methods in use in Europe for measuring the carbon content of aerosol, *Atmospheric Environment*, 38 (38), 6507–6519, DOI: 10.1016/j.atmosenv.2004.08.027, 2004.

Tuinstra, F., Koenig, J. L.: Raman Spectrum of Graphite, *The Journal of Chemical Physics*, 53 (3), 1126–1130, DOI: 10.1063/1.1674108, 1970.

Venkatachari, P., Zhou, L., Hopke, P. K., Schwab, J. J., Demerjian, K. L., Weimer, S., Hogrefe, O., Felton, D., Rattigan, O.: An
655 Intercomparison of Measurement Methods for Carbonaceous Aerosol in the Ambient Air in New York City, *Aerosol Science and Technology*, 40 (10), 788–795, DOI: 10.1080/02786820500380289, 2006.

Wang, Y., Chung, A., Paulson, S. E.: The effect of metal salts on quantification of elemental and organic carbon in diesel exhaust particles using thermal-optical evolved gas analysis, *Atmos. Chem. Phys. Discuss.*, 10 (7), 16941–16968, DOI: 10.5194/acpd-10-16941-2010, 2010.

660 Watson, J. G., Chow, J. C., Chen, L.-W. A.: Summary of Organic and Elemental Carbon/Black Carbon Analysis Methods and Intercomparisons, *Aerosol Air Qual. Res.*, 5 (1), 65–102, DOI: 10.4209/aaqr.2005.06.0006, 2005.

WKO, Mineralölindustrie: <https://www.wko.at/branchen/industrie/mineraloelindustrie/statistik.html>, last access 19.02.2020.

Wonaschütz, A., Hitzenberger, R., Bauer, H., Pouresmaeil, P., Klatzer, B., Caseiro, A., Puxbaum, H.: Application of the Integrating Sphere Method to Separate the Contributions of Brown and Black Carbon in Atmospheric Aerosols, *Environ. Sci. Technol.*, 43
665 (4), 1141–1146, DOI: 10.1021/es8008503, 2009.

Yttri, K. E., Aas, W., Bjerke, A., Cape, J. N., Cavalli, F., Ceburnis, D., Dye, C., Emblico, L., Facchini, M. C., Forster, C., Hanssen, J. E., Hansson, H. C., Jennings, S. G., Maenhaut, W., Putaud, J. P., Tørseth, K.: Elemental and organic carbon in PM₁₀: a one year measurement campaign within the European Monitoring and Evaluation Programme EMEP, *Atmos. Chem. Phys.*, 7 (22), 5711–5725, DOI: 10.5194/acp-7-5711-2007, 2007.

670 Yu, J. Z., Xu, J., Yang, H.: Charring Characteristics of Atmospheric Organic Particulate Matter in Thermal Analysis, *Environ. Sci. Technol.*, 36 (4), 754–761, DOI: 10.1021/es015540q, 2002.

Zickler, G. A., Smarsly, B., Gierlinger, N., Peterlik, H., Paris, O.: A reconsideration of the relationship between the crystallite size La of carbons determined by X-ray diffraction and Raman spectroscopy, *Carbon*, 44 (15), 3239–3246, DOI: 10.1016/j.carbon.2006.06.029, 2006.

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