



A new lidar design for operational atmospheric wind and cloud / aerosol survey from space

Didier Bruneau and Jacques Pelon

5 LATMOS/IPSL, Sorbonne Université, UVSQ, CNRS, Paris, France

Correspondence to: J. Pelon (jacques.pelon@latmos.ipsl.fr)

Abstract. Global wind profile measurement has for long been a first priority for numerical weather prediction. The demonstration from ground-based observations that a double-edge Fabry-Perot interferometer could be efficiently used for deriving wind profiles from the molecular scattered signal in a very large atmospheric vertical domain has led to the choice of the direct detection technique in space and the selection of the Atmospheric Dynamic Mission (ADM) Aeolus by ESA in 1999. ADM-Aeolus was successfully launched in 2018, after the technical issues raised for the lidar development have been solved, providing first global wind profiles from space in the whole troposphere. Simulated and real time assimilation of the projected horizontal wind information were able to confirm the expected improvements in forecast score, validating the concept of a wind profiler using a fixed line-of-sight lidar from space.

15 The question is raised here about consolidating results gained from ADM-Aeolus mission with a potential operational follow-on instrument. Maintaining the configuration of the instrument as close as possible to the one achieved (UV emission lidar with a single slanted line-of-sight) we revisit the concept of the receiver by replacing the arrangement of the Fizeau and Fabry-Perot interferometers with a unique Quadri-channel Mach-Zehnder (QMZ) interferometer which relaxes the system operational constraints and extends the observation capabilities to recover the radiative properties of clouds. This ability is meeting first and second profiling priorities of the meteorological forecasting community on atmospheric dynamics and radiation.

We discuss the optimization of the key parameters that may preside to the selection of a high performance system. The selected optical path difference (3.2 cm) of the QMZ leads to a very compact design allowing the realization of a high quality interferometer and offering a large field-angle acceptance. Performance simulation of horizontal wind speed measurements with different backscatter profiles shows results in agreement with the targeted ADM-Aeolus random errors, using an optimal 45° line-of-sight angle. The Doppler measurement is, from principle, unbiased by the atmospheric conditions (temperature, pressure, particle scattering) and only weakly affected by the instrument calibration errors. The study of the random systematic errors arising from the uncertainties in the instrumental calibration and in the modelled atmospheric parameters used for the backscatter analysis shows a limited impact under realistic conditions. The particle



30 backscatter coefficients can be retrieved with uncertainties better than a few percent in the boundary layer and in semi-transparent clouds. Extinction coefficients and depolarization ratio can be derived accordingly.

1 Introduction

Direct wind profiles in the meteorological atmosphere (0–25 km) are lacking over the oceans and in the tropics as indirect retrievals from temperature sounding are of no help in this region due to the lack of geostrophic equilibrium. First priority in
35 global atmospheric observations from space for weather forecasting was set on wind profiling to get better information on atmospheric circulation (WMO, 1996, 2012).

Aerosols are good tracers of atmospheric dynamics as their reduced Brownian motion allows for high accuracy spectral measurements of the small Doppler shifts induced in lidar backscattered light. This is why the choice of heterodyne lidar was first targeted for space missions (LAWS Instrument Panel report, 1987). The feasibility of such heterodyne system operating
40 on particle scattering was discussed for long in the community, but an important drawback was that the vertical profiling extent was limited by the very low value of the backscattering coefficient of the upper tropospheric particles, especially in the infrared. This leads to a very constraining lidar design in terms of mass and power needs for a space operating system.

Lidar backscattered signal from molecules, though much more spectrally broadened due to their high speed thermal motion at molecular scale, was however shown to be very effective for deriving upper tropospheric and stratospheric winds, as
45 shown from pioneering work performed at Service d'Aéronomie -now Laboratoire ATmosphères, Milieux et Observations Spatiales -LATMOS- (Chanin et al., 1989, Garnier et al., 1992). Specific interferometric techniques involving multi Fabry-Perot in differential detection, also known as double-edge technique, were implemented to analyze Doppler shift due to mean atmospheric motions. A few drawbacks are inherent to this approach, such as to require a narrow field of view and a particulate scattering correction in the analysis to reduce biases (Garnier et al., 1990). For the Aeolus mission, a technique
50 combining two interferometers in cascade, one matched to the broad molecular spectrum (Rayleigh channel), the other matched to the narrow aerosol spectrum (Mie channel) was chosen to be implemented in the space-borne Atmospheric LAser Doppler INstrument (ALADIN) providing high signal to noise over the whole atmosphere. Lidar wavelength operation was selected in the ultraviolet (UV) for space observations because of eye-safety issues, and because it maximized molecular return. Though first space observational tests in the UV revealed their difficulty (McCormick et al., 1993), the
55 candidate Atmospheric Dynamics Mission was accepted as the first Earth Explorer mission by ESA in 1999. An airborne demonstrator was developed which showed that needed performances could be achieved (Reitebuch et al, 2009). The satellite launch occurred in 2018, after a rather long delay induced by the large number of problems to be overcome. Though recent data analyses (Martin et al, 2020) show that the horizontal line-of-sight (HLOS) wind measurements exhibit seasonal and orbital dependent biases slightly larger than the missions requirements, the measurements are considered so good that
60 the European Centre for Medium-Range Weather Forecasts is now using them in their forecasts (https://www.esa.int/Applications/Observing_the_Earth/Aeolus/Aeolus_winds_now_in_daily_weather_forecasts). The



success of the Aeolus mission is then validating the concept of space based wind measurements using high spectral Doppler analysis on backscattered molecular signal.

Other priorities in atmospheric observation have been identified on the retrieval of aerosol and cloud radiative parameters as well as precipitations, which are addressing climate studies as more recently emphasized (National academies, 2018). This need has stimulated the development of the High Spectral Resolution Lidar (HSRL) technique which offers the capacity of retrieving the particle extinction and backscatter coefficients without using problematic assumptions. Such systems have been implemented onboard aircraft for preparing and validating observations from space missions. The first airborne HSRL systems were developed at DLR in Europe [Esselborn et al., 2008] and NASA in the USA [Hair et al., 2008], based on Iodine Cell absorption technique. The new Nasa/LaRC HSRL-2 system now also operate at 355 nm using a Michelson interferometric technique that was used for multispectral aerosol characterization (Müller et al., 2014). At the same time the lidar developed in the frame of the French CNES-CNRS project LEANDRE (Etude des Aérosols des Nuages et du Rayonnement) was upgraded and implemented as a New Generation of HSRL design for aircraft operation. In contrast with the DLR and NASA systems which were not designed to allow wind measurements, the French multi-wavelength system LEANDRE-NG (LNG) is operating as a HSRL at 355 nm for aerosol/cloud with a Doppler capacity allowing for wind measurements. It is based on the use of a single Mach-Zehnder Interferometer with four detection channels in phase quadrature (Bruneau and Pelon, 2003, hereafter BP03). This HSR Doppler (HSRD) Lidar design was chosen to better meet objectives of combined radar-lidar atmospheric observations (Delanoë et al., 2012). HSRD-LNG has been successfully validated and involved in several field experiments including the Aeolus CAL/VAL and EarthCARE preparation (Bruneau et al., 2015, Schäfler et al., 2018, Cazenave et al., 2019).

These technical developments offer new observational capabilities to the NASA-CNES CALIPSO mission (Winker et al., 2010) launched in 2006 and still in operation for the measurement of aerosol and cloud properties from space. The successes of the AEOLUS and of the CALIPSO missions have proved the potential of lidar sounding from space and its importance for meteorological and climatological applications. Many technical challenges for UV operation within AEOLUS have now been overcome by Airbus Defense and Space and ESA that pave the way for the EarthCARE mission focused on clouds and aerosols. This mission, developed in collaboration with JAXA, is to be launched by ESA in 2022. The EarthCARE payload, embarking a UV HSRL (ATLID) and a new cloud radar should take over CALIPSO lidar and CloudSat radar instruments designed for radiation budget analysis through the retrieval of cloud and aerosol properties (Stephens et al., 2018).

Taking advantage of observed performance and measurement capabilities from airborne measurements, based on the HSRD-LNG system, we study in this paper a MZI-based system for a future operational space mission. The proposed design is aiming at a performance-improved operation of a UV spaceborne lidar addressing atmospheric dynamics as a first goal. Such a design may also contribute to radiation budget analysis after the EarthCARE mission. The paper is organized as follows: in a first section, we recall the choices that are to be made to meet wind profiling objectives, and in the second one, we develop the lidar design and present its main characteristics, discussing its advantages. The performance assessment is then presented and discussed in a comparative way in the last section.



2. Preliminary choices

As previously introduced, the requirement of acquiring high accuracy wind profiling over a broad vertical atmospheric domain is ruled by the selection of efficient Doppler analysis on the molecular backscattered signal. Only molecular scattering can provide a high signal-to-noise ratio (SNR) in lidar detection in the upper troposphere and stratosphere at the global scale. Such a capacity has proven to be of high importance to Numerical Weather Prediction (LePichon et al., 2015). As intensity of the Rayleigh scattering cross-section is inversely proportional to the fourth power of the emission wavelength, the use of a UV system is preferred to optimize lidar profiling.

The ADM-Aeolus main requirements are based on this approach to perform accurate wind measurements in clear air over the whole troposphere (ESA publication SP-1233 (4), 1999, ESA Aeolus MRD, 2016). It is based on newly developed narrow line UV emission from solid state laser sources, and takes advantage of the direct detection of high molecular scattering signals at 355 nm through a large diameter telescope and a cascade spectral interferometric discriminator complemented by large efficiency detectors (<https://earth.esa.int/web/eoportal/satellite-missions/a/adm-aeolus>, ESA SP-1311, 2008). The signal detection design for ALADIN is based on the use of the Accumulation Charge Coupled Devices (A-CCD) which offer a very high efficiency (above 75 %) and a low noise which allow near quantum noise limited detection in the UV. The ALADIN design is thus meeting most of the needs for wind measurement, but the interferometer configuration is also leading to operation and performance constraints proper to the choice made.

The first wind measurements by lidar on molecular scattering have been performed at the Observatoire de Haute Provence (OHP) in the upper-atmosphere with a double-edge (DE) technique using a Fabry-Perot (FP) interferometer (Garnier et al., 1992). This photometric differential technique uses two spectral channels precisely positioned on each side of the molecular spectrum. Temperature and pressure are significantly varying on the vertical in the atmosphere and the dependence of the molecular spectral broadening with temperature and pressure needs to be accounted for in the analysis process (Dabas et al., 2008, Zhai et al., 2020). Particulate scattering can furthermore be contributing to within a few percent and even more in the troposphere, and this has been shown to be a constraint for DE-FP techniques, requiring performing independent particulate scattering ratio measurements for bias correction (Souprayen et al, 1999a, 199b, Witchas et al., 2020). In addition to the DE-FP of the Rayleigh channel dedicated to the analysis of the molecular return, ALADIN uses a Fizeau interferometer for the Mie channel dedicated to the analysis of high particulate returns. One important question in the optical design of the interferometers is to match the aperture of the telescope without degrading overall performance in case of misalignment. Both DE-FP and Fizeau interferometers have a small angular acceptance. The optical adaptation of the interferometers to the large telescope aperture is resulting in a very small field-of-view and implies a high accuracy in the pointing of the laser emission to maintain co-alignment. The solution used in ALADIN is to take advantage of the large telescope magnification for both the emission and the reception to reduce alignment sensitivity. The emission beam is sent through the telescope using a polarization by-pass for separating emission and reception paths (<https://earth.esa.int/web/eoportal/satellite-missions/a/adm-aeolus>). A drawback of this optical design is that only the backscattered light co-polarized with the emission



can be detected. This is leading to significant losses in the detection of particulate backscattered signal in the case of dust
130 and ice crystals that are inducing large depolarization. It also introduces an important limitation on the retrieval of the
aerosol and cloud parameters using the analysis of the Fizeau interferometer A-CCD channels (Flamant et al., 2008).

Two-wave interferometers such as Michelson (MI) or Mach-Zehnder (MZ) interferometers are a second group of techniques
that can be used in Doppler lidar measurements. MI and MZ interferometers in photometric or fringe imaging modes have
already been proposed in HSRL systems for wind measurements or scattering analysis (Cézard et al., 2009, Liu et al., 2012,
135 Cheng et al., 2015, Bruneau et al., 2015). These interferometers can include a field-compensation design (Bouchareine and
Connes, 1963) which allows a large incident field angle and facilitates their accommodation with a large beam étendue
system (i.e. large telescope and relatively wide field angle). MI and MZ interferometers can be designed to operate with two
or four detection channels (see BP03). Dual channel techniques are very similar to the DE-FP and present the same
limitations due to their sensitivity to temperature, pressure and aerosol scattering. The quadri-channel detection technique
140 (BP03) which provides four signals in phase quadrature, allows independent determinations of interference modulation and
phase, with the wind measurements being derived from the interference phase. Both particulate and molecular returns
produce the same interference phase and therefore the same wind measurement. Particulate backscatter thus improves the
SNR and the wind speed accuracy without introducing bias. It can also appear convenient to use the linear fringes
interferometric pattern imaged on a CCD array as its position depends linearly on the Doppler shift. It has been shown that
145 fringe imaging and quadri-channel techniques provide the same theoretical measurement precision (Bruneau, 2002, BP03). It
must be also considered that the accommodation of linear fringes on a square image zone with a circular aperture receiver
necessitates a truncation of the aperture that would decrease the signal and hence the measurement precision. We therefore
rule out the fringe imaging technique from our choices.

In addition, the Quad-channel techniques offer another advantage, as they do not require any specific spectral positioning of
150 the laser emission relative to the interferometer, provided that a reference signal characterizing the frequency at the emission
is acquired. Appropriate frequency stability is just required during the signal accumulation needed for a single measurement.
A large spectral drift of the laser source and the interferometer is tolerated over a longer timescale. Off-axis MI
interferometer associated with quadri-channel detection and QMZ present the same advantages in terms of measurement
robustness: insensitivity on molecular spectral shape and particulate backscatter, tolerance with regard to spectral drifts and
155 misalignments (thanks to a field compensated design). Although the MZ optical arrangement is slightly less simple than the
off-axis MI's, the latter brings the difficulty that one output port is necessarily very close to the input one and would hamper
the positioning of the detectors. The QMZ technique thus appears as a near-optimal choice in this group, and we made this
choice for optimizing the new receiver design.

Besides, an experimental comparison has been performed between the Dual-channel MZ (DMZ) and the DE-FP techniques
160 at the OHP (Bruneau et al., 2004). It was shown that, in the same conditions of measurement, the DMZ statistical error is
lower than that of the DE-FP by a factor of 1.4 thanks to a higher optical efficiency. As the Quadri-detection MZ (QMZ)
produces a statistical error larger than that of the DMZ by the same factor (Bruneau, 2001), the QMZ and the DE-FP are



equivalent in terms of statistical error. It is to be further noted that an optimized and original optical design is used in ALADIN: the reflected signal from one edge of the DE-FP interferometer is reinjected to the other edge. This scheme improves the DE-FP optical efficiency as compared to the common method that splits the incident signal on two halves of the FP aperture, as done in the OHP setting. It leads to a theoretical advantage to the ALADIN concept in terms of measurement precision, as compared to a Quadri-detection technique, by a factor that depends on the efficiency of the double-pass optical setting but cannot be larger than 1.4 (case of a perfect efficiency).

Comparison of advantages and disadvantages guiding the choice of interferometric design for wind measurement on the molecular scattering is summarized in Table 1. In contrast with ALADIN, which utilizes a second interferometer (Mie channel) for regions of the atmosphere where the aerosol load is significant, and taking advantage of the QMZ insensitivity to particulate backscatter (in terms of bias) we chose to use a unique interferometer to probe all the layers of the atmosphere. This choice allows a simplified receiver design considered to ensure good robustness and reliability. In the conducted analysis, we will rely on numerical simulations and on results obtained from QMZ operation with the airborne HSR-LNG to define realistic parameter for the interferometer and derive the performance assessment.

An additional advantage of the QMZ technique is that it also allows measurements of the scattering ratio (derived from the contrast of the interference). The particulate signal produces an interference contrast near unity, significantly higher than that produced by the molecular signal. The measurement of the interference contrast given by the total atmospheric signal then leads directly to the particulate backscatter ratio and from there to a precise quantification of aerosol backscattering. We will come back on this advantage for the analysis of aerosol and cloud properties.

Table 1: Summary of advantages (+) and disadvantages (-) of measurement techniques for direct detection wind profiling on molecular return.

Doppler measurement technique	DE-FP ALADIN design	QMZ Proposed design
Sensitivity to molecular spectral shape	Yes (-)	No (+)
Sensitivity to particulate scattering	Yes (-)	No (+)
Sensitivity to lidar alignment	High (-)	Small (+)
Spectral stability requirement	High (-)	Low (+)
Statistical error for an optimized design	σ_0 (+)	$\sigma_0 \leq \sigma < 1.4\sigma_0$ (-)



185 3. Lidar design

3.1. General architecture and characteristics

As mentioned earlier, the goal for the new design is to simplify and ruggedize the existing one as much as possible, keeping as many solutions based on Aeolus and ATLID designs as needed, and accounting for recommendations made by the Aeolus Science Advisory Group (ASAG, 2020). The emitter is thus based on the nominal Nd:YAG laser in a redundant design with frequency converters to 355 nm, as developed for Aeolus. As the field-compensated MZI acceptance angle is much wider than that of the DE-FP, a very small emission divergence and a mono-static emitter-receiver is no longer required. So, in contrast to ALADIN, but similarly to the ATLID lidar for EarthCARE (Hélière, 2012), a bi-static design is chosen. The schematic of the lidar optical design is presented in Fig. 1. The lasers can be coupled to a small telescope (typically about 15 cm in aperture) that assures an emitted beam divergence of 50 μ rd for a 99 % relative encircled energy. The emitters and the expanders can be mounted on the same plate, and mirrors can be tilted to maintain alignment with the receiving optics. The emitted energy will conservatively be kept similar to the actual ALADIN laser source (e.g. 65 mJ @355nm). We however will extrapolate, with little risk, the possibility to extend the laser repetition rate to 100 Hz. A short laser pulse duration of 5-7 ns (corresponding to a high energy extraction from a Q-Switched oscillator and leading to a high frequency conversion efficiency) is allowed to keep an appropriate spectral linewidth (less than 150 MHz). We will also assume that pulse linewidth and the spectral jitter or drift allows maintaining the average spectral width within 200 MHz during an accumulation of 50 shots as observed in spaceborne Aeolus observations (Reitebuch et al., 2019a).

The receiver telescope with a 1.5 m aperture is coupled to a silica-silica optical fiber through an assembly of front optics which includes a 0.1 nm-bandwidth background filter. The signal fiber, with a core diameter of 340 μ m (and a numerical aperture of 0.22) is used as the field stop and defines a reception field angle of 100 μ rd which allows a good margin for the lidar emission-reception alignment. Note also that the receiver is not sensitive to the polarization state of the collected light and hence to the depolarization by particles scattering. As discussed in the literature, it could be useful to measure the particulate depolarization to better identify cloud and aerosol types as emphasized by the Aeolus SAG (ASAG, 2020) -see also for example Burton et al, 2012-. This can be handled inserting a polarization splitter between the telescope and the interferometer in the front optics system (see Fig. 1). The interferometric analysis is done in this case on the parallel polarized signal, and the perpendicularly polarized signal is measured in an additional channel as done in the DHRS-LNG instrument (Bruneau et al., 2015). However we will not discuss this polarization analysis in the basic lidar design, keeping it as an option.

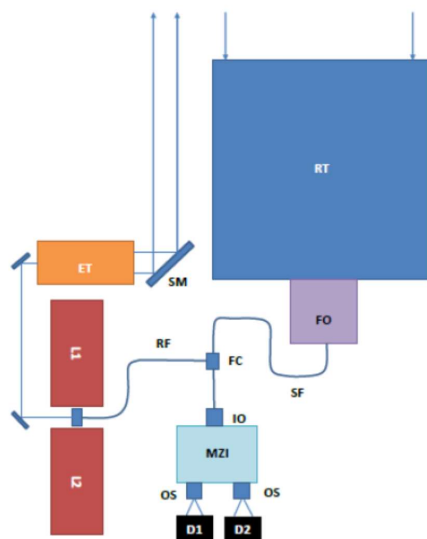


Fig. 1: Diagram of the lidar system. L1, L2 lasers (nominal and redundant), ET emission telescope, SM steering mirror, RT reception telescope, FO front optics (including background filter) SF signal fiber, RF reference fiber, FC fiber coupler and mode scrambler, IO input optics, MZI Mach-Zehnder interferometer, OS output separation optics, D1, D2 A-CCD detectors.

An optical mode scrambler is inserted on the fiber path in order to obtain a uniform illumination distribution at the interferometer input. This arrangement ensures that the interferometer response is not biased by emission/reception misalignment. The fiber and the scrambler ensure the complete depolarization of the signal before it arrives on the MZI (even when including a polarization splitter in the front optics). The mode scrambler also allows the injection of a small amount of the emitted pulse used as the reference signal and transported by a second optical fiber. This was implemented and successfully tested on the airborne LNG system. The output of the fiber is then collimated by a 15-mm-focal-length lens at the MZI input port.

The main characteristics of the proposed lidar design are reported in Table 2 and compared to the nominal (Paffrath, 2006; <https://earth.esa.int/web/eoportal/satellite-missions/a/adm-aeolus>) and actual (Reitebuch et al., 2019a) ALADIN parameters (see also ASAG, 2020). This table also includes parameters discussed in next sub-sections.



Table 2: Main characteristics of the lidar design used for the analysis of performance. Parameters are compared to Aeolus ones as a reference actual in space values reported by Reitebuch et al., 2019).

System Parameters	Aeolus nominal and actual values	Our Study
Satellite altitude	395-425 km	400 km
Satellite speed	7.2 km/s	7.2 km/s
Line-of-sight angle	35°	45°
Horizontal resolution	50 km (planned) 90 km (in space)	50 km
Vertical resolution	0.5/1/2 km	0.5 km (0.25 km optional in the PBL)
Wavelength	355 nm	355 nm
Emitted energy (laser source)	130 mJ (planned) 65 mJ (in space)	65 mJ
Repetition rate	100 Hz (planned) 50 Hz (in space)	100 Hz
Emitter linewidth	30 MHz	200 MHz
Emitter optical transmission	0.66 (monostatic)	0.9 (bistatic)
Telescope aperture	1.5 m	1.5 m
Telescope field of view	0.02 mrd	0.1 mrd
Receiver optical transmission	0.42 (polarization by-pass)	0.5 (optical fiber)
Blocking filter bandwidth	1 nm	0.1 nm
Doppler sensor	2 interferometers: Double-Edge Fabry-Perot and Fizeau	1 interferometer: Quadri-channel Mach-Zehnder
A-CCD background	6 p-e /pixel / 50 shots	6 p-e /pixel / 50 shots

235 3.2. QMZ Design

The field-compensated QMZ design has been detailed in several papers [Bruneau, 2001, BP03, Smith and Chu, 2016]. Making the choice of a unique interferometer, the optical path difference (OPD) Δ between the two arms of the interferometer is the key parameter for optimizing the wind speed retrieval over the whole troposphere using both the molecular and the particulate backscattering. In contrast with BP03 that focused mainly on the HRS separation of molecular
 240 and particulate backscattering, we give here priority to the optimization with regard to wind measurement and consider



backscattering as a by-product (though of great interest). We will then consider performance obtained on the analysis of aerosol and cloud properties for this wind performance driven design.

The principles of the lidar interferometric system are presented in Fig. 2a. The optical signal incident on the interferometer is previously depolarized by passing through a multimode optical fiber. A quarter-wave plate is inserted in one arm of the interferometer and polarizers at its output separate the signal in four interference components in phase quadrature.

The interferometer is built in the same way as for the HSR-LNG (Bruneau et al., 2015) with an assembly of prisms and plates (Fig 2b). For the selected MZ optical path difference (see next section) the total prism assembly stands within a volume of approximately 1.5 cm x 2 cm x 4 cm. The different elements can be optically contacted resulting in a single block interferometer. The interferometer adjustment is obtained by construction and is unalterable.

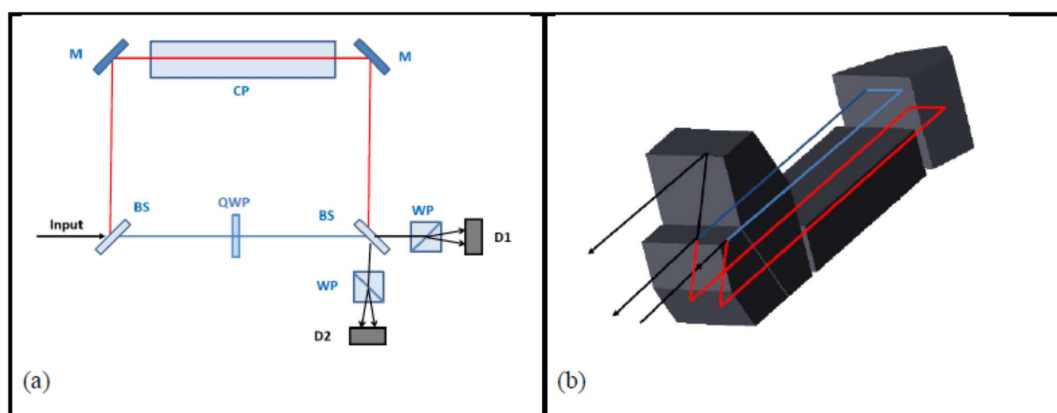


Fig. 2. (a) QMZ schematics: BS beam splitter, M mirror, CP field compensation plate, QWP quarter-wave plate, WP Wollaston polarizer, D1,D2 A-CCD detectors; (b) MZ actual prism arrangement(QWP,WP and detectors not shown here).

As mentioned above, an alternative design would consist in replacing the MZI by an off-axis Michelson interferometer but still in quadri-channel detection. In this case, the quarter-wave plate should be replaced by a $\frac{1}{8}$ -wave plate used in double path. One difficulty of this design is the accommodation of one detector close to the input port of the interferometer. An additional fiber between interferometer and detector could be a solution. The alternative choice of a quadri-detection off-axis Michelson interferometer would lead to the same equations and results. We however keep considering in this paper the QMZ as the nominal design, as we refer to proven optical solutions from our airborne system.

The beam issued from each of the two MZI output ports is focused on an A-CCD detector by means of a 6.5-mm-focal-length lens. A quartz Wollaston polarizer, positioned just ahead of the focusing lens, separates the two polarizations issued from one MZI output port with an angular separation angle of 25 mrd. This way the fiber output is imaged on the detector as



two spots of 150 μm in diameter separated by 160 μm . Thanks to the mode scrambler these images are well defined disk of
265 homogeneous illuminations. They fill a 8 by 8 pixels area of the A-CCD image zone in the same way as the Rayleigh
channel of ALADIN. The two A-CCD provide the four channel signals (Eq. 1) of the QMZ technique.

The average signal delivered by each channel of the A-CCD outputs, for $i = 1$ to 4, can be written

$$S_i = \frac{S_{\text{tot}}}{4} a_i \left[1 + M_i M_{\text{atm}} \sin \left(\varphi + (i - 1) \frac{\pi}{2} \right) \right] + S_{b_i} \quad (1)$$

with S_{tot} the total signal, and a_i , M_i the relative photometric sensitivity and instrumental interference modulation of channel i ,
270 respectively. M_{atm} is the interference modulation given by the atmospheric backscattered signal and φ the interference phase.
 S_{b_i} is the background signal due to solar detected light and A-CCD intrinsic noise.

The interference phase, from which the wind speed is derived, and the interference modulation, which leads to the scattering
ratio, can be determined independently from these four signals. Actually, the line-of-sight (LOS) wind speed is calculated
from the phase difference between the atmospheric signal and a reference signal obtained on a highly attenuated pick-up of
275 the emission.

The equations which give the wind and backscatter products, as well as the respective random errors are detailed in
Appendix A. These equations are the basis of numerical simulations that have been performed to analyze measurement
performance of the selected design. As the statistical error on wind is depending on the OPD and interference modulation
(see Appendix A), we will first examine the optimal determination of these two parameters more directly linked to the
280 interferometer design and then discuss measurement SNR depending on the lidar system and mission parameters in Sect. 4,
as derived from the numerical simulations performed.

3.3 MZI optical path difference

As seen from Eq. A16 in Appendix, the statistical error on wind measurement for a given SNR is inversely proportional to
the OPD and to the interference modulation M_{atm} . The interference modulation M_{atm} is in counterpart a fast decreasing
285 function of the OPD for a given signal spectral width. There is then a compromise to achieve between the frequency
discrimination (large OPD) and the interference modulation (small OPD). It has been shown (Bruneau, 2000) that, for
measurements in a pure molecular atmosphere at $\lambda = 355$ nm, the optimum OPD is 3.2 cm. As we want to optimize
measurement in conditions of clean air (scattering ratio smaller than 1.01), the OPD should be close to this value. The
optimal OPD is thus quite smaller than the one proposed in BP03 (10 cm) that aimed at a compromise between wind and
290 scattering ratio measurements. A complete parametric study is presented in Sect. 4. Such a small-ODP MZI, can be easily
integrated as a solid state device derived from the design used for the airborne system, as reported in Fig. 2b (Bruneau et al.,
2015).

The modulation produced by the molecular return M_{mol} is dominated by the Doppler spectral width of the molecules
Brownian movement and varies from 0.49 to 0.59 depending on the atmospheric temperature with an average value of 0.56.
295 Note that at high atmospheric pressure, near the surface, Brillouin molecular scattering causes a relative broadening of this



spectrum by approximately 1.7 %. This effect causes itself a small decrease in the molecular modulation M_{mol} from 0.49 to 0.48 (see Appendix C).

As the spectral broadening caused by the particle Brownian motion is small compared to the emitted linewidth, the modulation produced by the particle return M_{par} is very close to that given by the emitted linewidth. For an emission linewidth of 200 MHz, including the broadening by frequency shift or jitter during the measurement integration duration, and a 3.2 cm OPD, M_{par} is still as high as 0.95, according to equations given in BP03.

Note again that, as the QMZ technique allows the determination of wind speed (through the interference phase) and the scattering ratio (through the interference modulation) independently, the superposition of a narrow particle spectrum to the wide molecular spectrum does not bring any measurement bias. The increase in interference modulation caused by an additional particle scattering will indeed improve the wind measurement accuracy.

3.4 MZI optical quality

The interferometer optical quality is obviously an important aspect of the design as it determines the intrinsic modulation factors M_i (Eq. 1). These modulation factors can be slightly different from one channel to the other but for clarity we consider here that they are all equal to M_0 . The equations of Appendix A show that the random errors (both on wind speed and scattering ratio) are, at first order, inversely proportional to this intrinsic modulation factor M_0 .

We can also consider that M_0 is the product of two factors: M_W modulation degradation caused by imperfections of optical surfacing and non-homogeneities of the optical index resulting in a global wave-front error (WFE) and M_R the modulation degradation caused by the imbalance of the reflection-transmission factors of the beam-splitters.

It has been shown that the modulation degradation of a Michelson interferometer is a function of the RMS value of the WFE (σ_{WFE}), almost independently of the shape of this distortion (Liu et al., 2012). This result is obviously also valid for a MZI. Assuming a Gaussian distribution of the WFE we obtain the modulation factor M_W as

$$M_W = \frac{1}{\sigma_{\text{WFE}}\sqrt{2\pi}} \int \exp\left(-\frac{\delta^2}{2\sigma_{\text{WFE}}^2}\right) \cos\left(\frac{2\pi\delta}{\lambda}\right) d\delta \quad (2)$$

We can also admit that σ_{WFE} is proportional to the area of the useful interferometer aperture. It is therefore essential to evaluate the useful aperture of the MZI which can withstand the large beam étendue of the system. The calculation, detailed in Appendix B, leads for the chosen lidar parameters to a useful aperture $D_{\text{MZI}} = 6.2$ mm.

Using Eq. (2) we calculate M_W as a function of the MZI aperture for different values of the σ_{WFE} per square cm and an operation wavelength of 355 nm (Fig. 3a). For comparison, the intrinsic instrumental contrast of the HSRD-LNG interferometer is 0.65 for a useful aperture of 20 mm (and a 20 cm OPD) that corresponds to a σ_{WFE} less than 20 nm/cm². For the same optical quality $\sigma_{\text{WFE}} = 20$ nm/cm² (which can be improved) a modulation factor $M_W > 0.99$ is achievable for the proposed MZI with a useful aperture of 6.2 mm (see Fig. 3a).



In the case of a non-homogeneous temperature distribution in the optical elements, and more particularly a transverse temperature gradient in the compensation plate, leading to an angular tilt of one wavefront plane, we computed that the impact is an additional degradation of about 0.5 % in M_w for a thermal gradient of 0.1 K per cm. This shows a relative acceptance of temperature changes, not requiring a high standard in temperature control of the interferometer.

Note also that, because of the field compensation, the variation of OPD is only dependent on the fourth power of the source angle (the second power dependence is canceled). The related RMS-WFE can be expressed by $\sigma_{WFE} = \frac{(n-1)^2}{n^2(n^2-1)} \frac{D_{tel}^2 \alpha_{tel}^2}{640\Delta}$. With the system parameters presented above, the source full angle brings a negligible contribution of less than 0.1 nm to the RMS-WFE.

The second intrinsic modulation factor M_R stems from the fact that it is not possible for a pure (non-absorbing) dielectric coating to ensure a perfect balance between the reflectivity R and the transmission T for both polarizations, at an incidence angle of several tens of degrees. The consequent imbalance of the beam-splitter reflectivity causes a modulation factor $M_{Rmin} = \frac{2RT}{(R^2+T^2)}$ on two of the four channels of the QMZ, the two others channels being unaffected. The average reflectivity modulation factor is then $M_R = \frac{1}{2(R^2+T^2)}$. Fig 3b presents the decrease of instrument intrinsic modulation M_R as a function of the reflectivity-transmission difference $R-T$. For comparison, in the HSRD-LNG interferometer the incidence angle on the beam-splitters is 40° and a reflection-transmission imbalance of 0.1 has been achieved leading to $M_R = 0.99$.

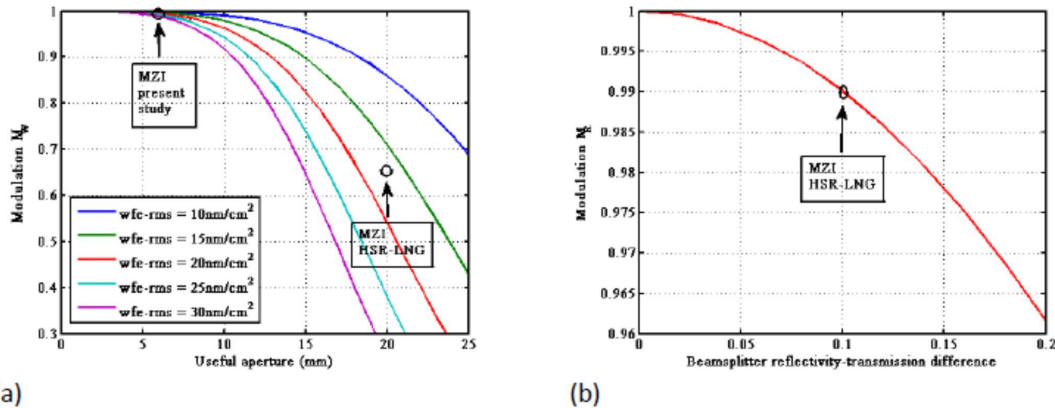


Fig. 3 (a) Instrumental modulation factor M_w as a function of MZI useful aperture for different RMS-WFE. The dots represent the MZI of the present study and the actual MZI of HSR-LNG. (b) Instrumental modulation factor M_R as a function of the beam splitter $R-T$ difference.



We can then conclude that an intrinsic instrument modulation $M_0 = 0.98$ is achievable for a MZI of small OPD (3.2 cm) associated to a space-borne lidar with a receiver aperture – field-angle product as large as 0.15 m.mrd.

Note also that the optical quality requirements on the MZI are less stringent than for a FPI, which necessitates a larger aperture, because of low field angle acceptance, and a high surfacing quality to avoid losses in transmission and finesse.

3.5 Calibration

Prior to processing actual atmospheric signals, it is necessary to perform a calibration of the MZI for determining the transmission and modulation factors a_i , M_i parameters of the MZI defined in Eq. (1). The sensitivity parameters a_i can be determined by recording the four QMZ signals in absence of interference either by masking sequentially one arm of the interferometer and the other, and adding the signals of the two sequences, or by using the laser in multimode operation (without injection seeding) to get a flat interferometric response. The calibration of the instrumental modulations M_i is then obtained by recording a long sequence of reference signals with a slowly varying OPD phase (with temperature for instance). The recorded signals are fitted by a least-square method on the four modeled signals (Eq. A2) by adjusting the M_i parameters. The impact of calibration errors on the measurements is discussed in Sect. 4.6.

4. Performance assessment

4.1 Model description

Performance modeling starts by computing the total signal S_{tot} (Eq.(1)) expressed in number of photo-electrons (p-e) for a single shot, is based on the canonical lidar equation:

$$S_{tot}(R) = \frac{\eta E A T_{inst}}{h\nu} \int_{r=R-\delta R}^{r=R+\delta R/2} \frac{T_{atm}^2(r) \beta(r)}{r^2} dr \quad (3)$$

where R (m) is the range from the instrument, δR (m) is the range gate, η is the detector quantum efficiency, E (J) is the emitted energy, A (m²) is the receiver aperture area, T_{inst} is the total optical efficiency product of emitter and receiver efficiencies, $h\nu$ is emitted the photon energy, T_{atm} is the atmospheric transmission from instrument to range r for a line-of-sight (LOS) nadir angle θ , β (m⁻¹sr⁻¹) is the total (molecular and particulate) backscatter coefficient.

The atmospheric backscattering and transmission are computed from the ESA reference model of atmosphere (RMA) including a measurement statistics of aerosol backscattering (Vaughan et al., 1998). A cirrus cloud between 9.5 and 10.5 km altitude with an optical thickness of 0.1 (backscattering coefficient equal to 4.10⁻⁶m⁻¹sr⁻¹) can also be introduced. The RMA comprises data of different aerosol backscatter, cloud backscatter, extinction, background radiance, and ground albedo. Extinction and backscatter coefficients corresponding to the Lower Quartile (LQ), Median (MD), Higher Quartile (HQ) and Median profiles with an additional cirrus cloud (MD+CIR) are given in Fig.4.

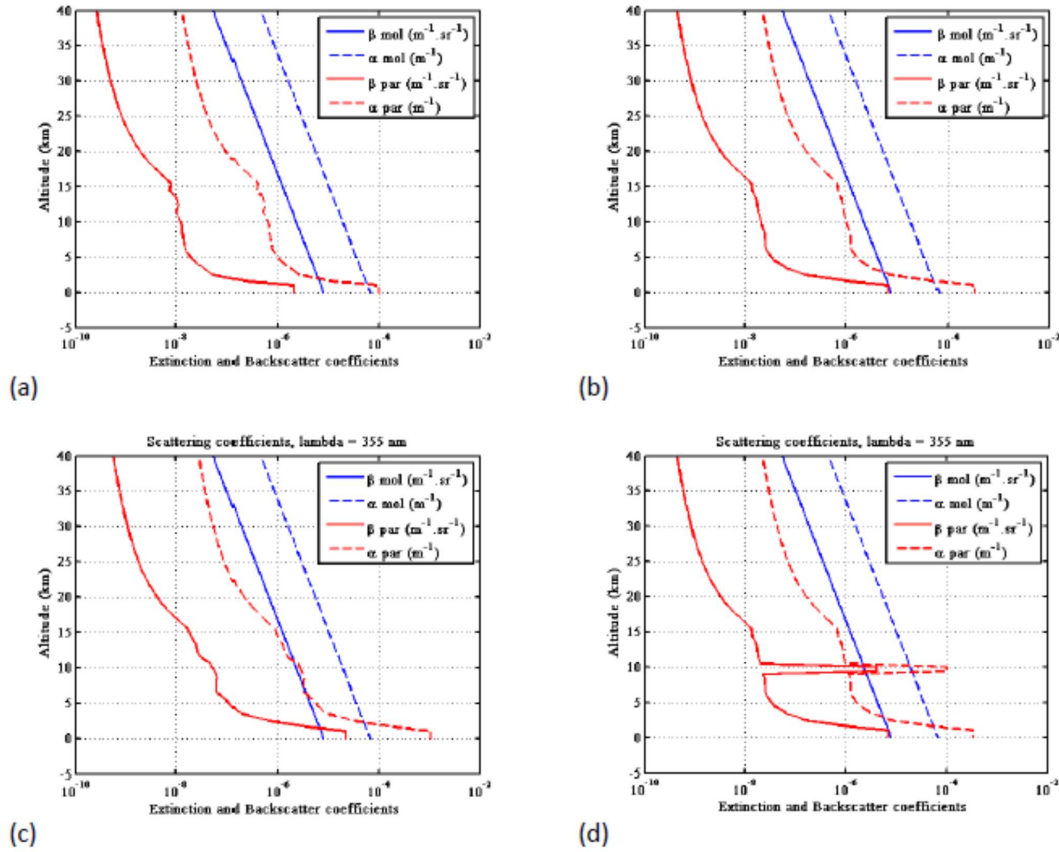


Fig 4. Backscatter and extinction coefficients vertical profiles used in the performance model. (a) RMA Lower Quartile, (b) RMA Median, c: Higher Quartile, d: Median plus cirrus.

A background signal S_{bs} caused by the sun illumination and the intrinsic reading noise S_{bd} due to the A-CCD detector are added on all the range gates so that the total background noise is written as

$$S_b = S_{bs} + S_{bd} = \eta A \Omega T_{rec} I_s(\lambda) \delta \lambda \cos(\theta_s) \frac{alb}{\pi} T_{atm}(\theta_s, z_s) T_{atm}(\theta, z_s) \frac{2\delta}{c} + S_{bd} \quad (4)$$

where Ω (sr) is the receiver field solid angle, T_{rec} is the receiver efficiency, I_s (Wm^{-2}) is the sun irradiance at top of atmosphere at wavelength λ , $\delta \lambda$ is the bandwidth of the blocking filter, θ_s is the sun elevation angle, alb is the albedo (surface or cloud) and z_s is the altitude of the scattering layer. For these computations we take $alb=0.3$ for the ground and add



385 alb=0.015 for the thin cirrus. Assuming a dawn-dusk sun-synchronous polar orbit, similarly to Aeolus, we take $\theta_s = 80^\circ$. The blocking filter has a bandwidth $\delta\lambda = 0.1$ nm.

An important advantage of the A-CCD detectors is their high quantum efficiency ($\eta = 0.85$) allied to a low reading noise. The signal is pre-accumulated on 50 shots on the A-CCD (e.g. 0.5s corresponding to an along track horizontal sampling resolution of 3.6 km, similar to Aeolus) before being read by the analog-to-digital converter that add a noise equivalent to a
390 background signal S_{bd} of 6 p-e per detector pixel (Paffrath, 2006).

The signals are summed on 14 elementary samplings corresponding to a total of $N_s = 700$ shots, for an observation horizontal resolution of 50 km and the resulting total SNR is

$$SNR = \frac{S_{tot}\sqrt{N_s}}{\sqrt{S_{tot}+S_b}} \quad (5)$$

The recorded total signals and SNR (i.e. calculated on the sum of the four detection channels) vertical profiles used in the
395 performance model are presented in Fig. 5 for the median (MD) RMA backscatter profile with the cirrus cloud and a 45° LOS angle. Except a large increase in the cirrus cloud and at ground, the SNR is slowly varying over the whole troposphere. The standard deviation on the projected horizontal (HLOS) wind speed is calculated from Eq. A16 as

$$\sigma(V_{HLOS}) = \frac{\sigma(V_{LOS})}{\sin(\theta)} \frac{R_E+z}{R_E+z_{sat}} \quad (6)$$

where R_E is the Earth radius and z_{sat} the satellite altitude.

400

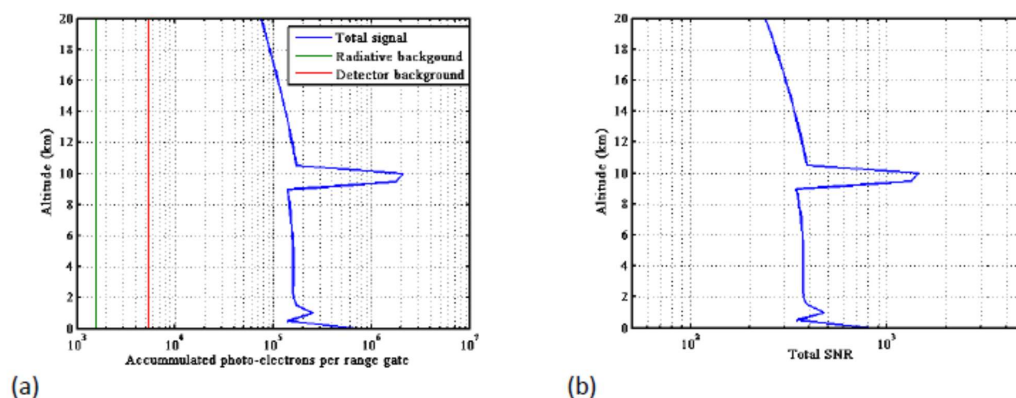


Fig 5 (a) Number of accumulated photo-electrons per range gate (b) SNR for RMA MD backscatter profile plus cirrus at 45° LOS angle.



405 The MZI intrinsic modulation is calculated assuming a σ_{WFE} of 20 nm/cm², a R-T imbalance of 10 % and an emitted spectral width and jitter of 200 MHz RMS. The modulation on molecular return M_{mol} is computed with the standard model of temperature at mid latitudes.

The background and the reference signals are assumed to be recorded with a SNR much higher than that of the atmospheric signal, so that they add a negligible variance to the measurements when subtracted to the total signal. This can be done on
410 each acquisition sequence to avoid possible bias due to hot pixels, as observed and corrected with A-CCDs in space (Weiler et al., 2019).

4.2 OPD and LOS angle optimizations

We first come back on the OPD optimization to detail results previously given on average. The standard deviation of the statistical error on the retrieved horizontal wind speed V_{HLOS} as given by Eq. (6) is presented in Fig. 6 for the median, higher
415 quartile and lower quartile aerosol backscatter distributions and for OPDs varied from 2.2 to 4.2 cm. The LOS angle is set to 45° in these computations.

Fig. 6d thus shows that the optimum OPD of 3.2 cm is effective for all atmospheric models corresponding to scattering ratios smaller than 2 as it is the case outside the boundary layer.

We now look at the pointing angle. The standard deviation of the statistical error on horizontal wind speed is presented in
420 Fig. 7 for the median, higher quartile and lower quartile aerosol backscatter distributions and for LOS angle varied from 35° to 55°. Results are analyzed (as for OPD) looking to HLOS wind error averaged between 0.5 and 15 km, and on vertical profiles. The MZI OPD is set to 3.2 cm on these computations.

One can see from Fig 7d that the optimal 3.2 cm OPD and 40°- 47° LOS angles are well matched to aerosols distributions from lower to higher quartiles of RMA aerosol distribution. The errors are noticeably degraded in the troposphere below
425 10km in altitude for angles smaller than 40° or larger than 50°. For all the aerosol distributions, a LOS angle close to 45° appears to give better results. We thus chose 45° as the optimal value for conducting final analysis performance on HLOS wind retrievals. The standard deviation of the horizontal wind speed statistical error averaged between 0.5 and 15 km in altitude is slightly less than 2 ms⁻¹ in the troposphere for the three distributions.

4.3 Comparison with Aeolus requirements

430 Fig 8a shows the HLOS wind speed error profiles for the selected parameters and the aerosols distributions LQ, MD, HQ and MD + cirrus for the parameters listed in Table 2, corresponding to the previously discussed choices and optimization. Fig. 8b shows the same results but for a variable vertical resolution according to Aeolus requirements. The modelled statistical errors for the RMA-MD aerosol distribution presented in the previous section are summarized in Table 3 in comparison with Aeolus requirements. One can see that in the free troposphere the Aeolus requirements are fulfilled with a good safety
435 margin. For a constant vertical resolution of 500 m, results obtained in the atmospheric boundary layer are better than in the clean atmosphere, due to a larger scattering ratio which allows to compensate the decrease in transmission, till the aerosol



attenuation becomes too high. The errors are nevertheless slightly higher than the requirement (1 ms^{-1}) in the PBL with values around 1.5 ms^{-1} on average.

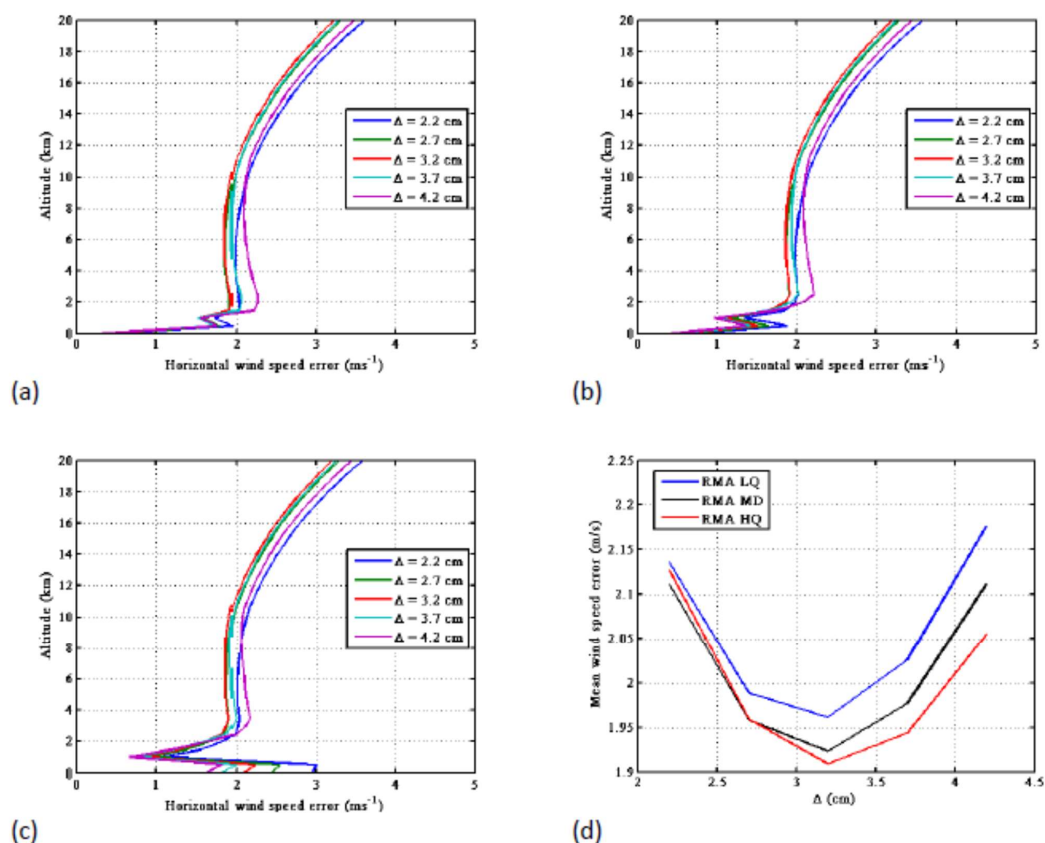


Fig. 6 (a, b, c) Horizontal wind speed (HLOS) statistical error as a function of altitude for OPD for Lower Quartile, Median, Higher Quartile RMA aerosols distribution respectively. (d) wind speed error averaged between 0.5 and 15 km in altitude as a function of OPD.

In presence of a cirrus cloud the wind speed accuracy is improved to a fraction of ms^{-1} in the due to the increased signal return and interference modulation M_a (Eq.A3 and A12) but is degraded in clear air below the cloud layer by 0.5 ms^{-1} due to



the decreased atmospheric transmission (by about 20 % on the 2-way transmission). The effect of a cirrus cloud on wind and backscatter measurement accuracies at different levels is discussed in Sect. 4.5 as a function of the cloud optical thickness.

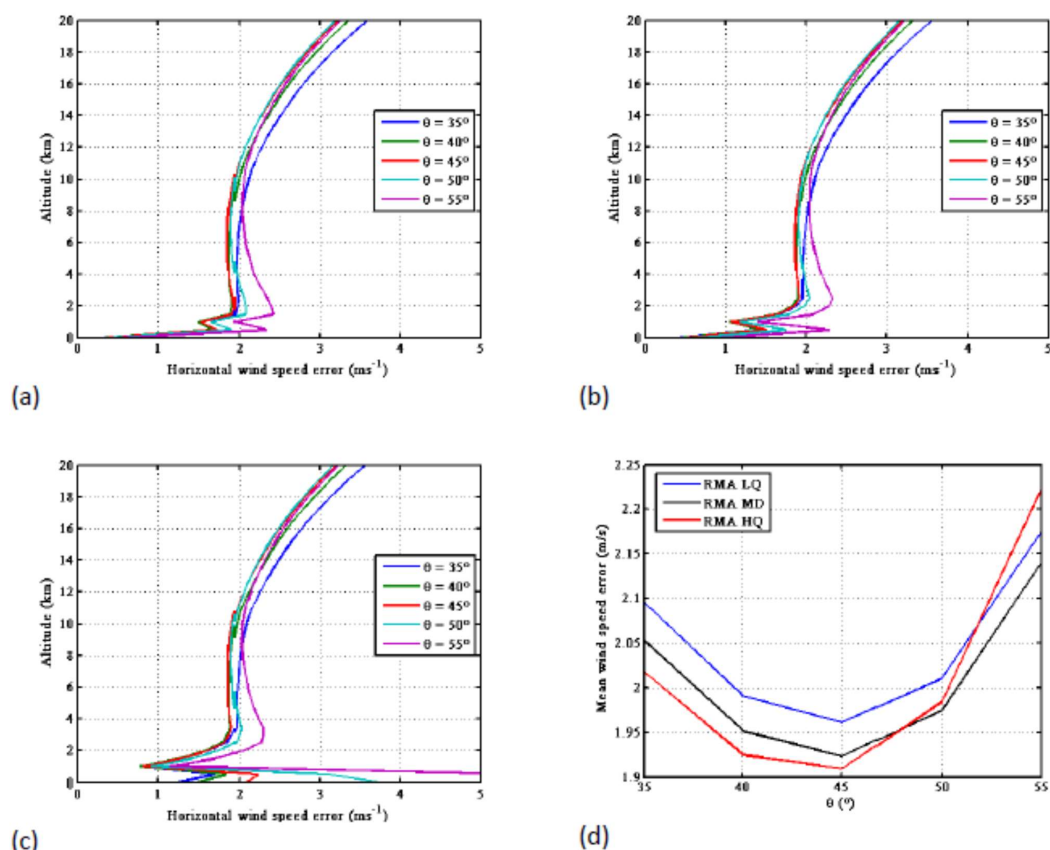
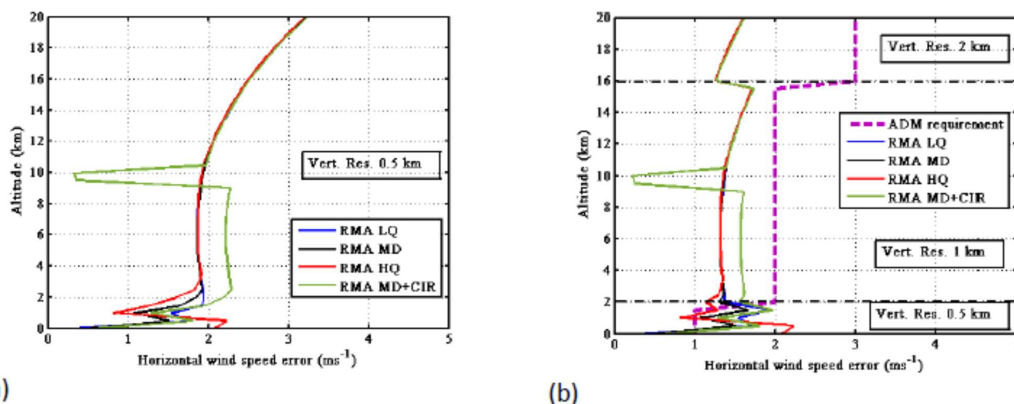


Fig. 7 (a, b, c) HLOS wind speed statistical error as a function of altitude for LOS angles between 35° and 55° , for Lower Quartile, Median, Higher Quartile RMA aerosols distribution respectively. (d) HLOS wind speed error averaged between 0.5 and 15 km in altitude as a function of LOS angle.



455 **Figure 8:** Vertical profile of the horizontal wind speed error (std deviation) for different backscatter profiles: LQ, MD, HQ and MD + cirrus for the optimal parameters $\Delta = 3.2$ cm and $\theta = 45^\circ$; (a) constant vertical resolution of 0.5 km; (b) variable vertical resolution (0.5 to 2 km) and requirements according to Aeolus.

Table 3. HLOS wind statistical error of the proposed QMZ design compared to the original Aeolus requirements.

Altitude (km)	Vertical resolution (m)	Aeolus horizontal wind speed accuracy requirements (ms^{-1})	QMZ averaged horizontal wind speed statistical error modelled for RMA MD (ms^{-1})
0-2	500	1	1.5
2-16	1000	2	1.4
16-20	2000	3	1.9

460

4.4 Backscatter measurements

The choice of a small OPD is a strong advantage for wind measurements in clear air but is not optimized for retrieving particulate backscatter (the OPD should be larger, as discussed in BP03). The QMZ technique with the selected parameters can nevertheless retrieve backscattering coefficients and determine the lidar scattering ratio R_β (see Appendix A) with a relatively good accuracy. The derivation of the particulate backscattering requires the knowledge of the molecular backscattering β_{mol} and the calculation of the molecular and particle interference modulations M_{mol} and M_{par} respectively.



These calculations bring necessarily additional errors in the backscatter retrieval process. They are summarized later in this section, and detailed in Appendix C.

Let us first look to noise induced errors, assuming no other random error source. Fig. 9 shows the relative standard deviation of the statistical error on R_β and on β_p (according to Eqs. A9 and A10 in Appendix A) for the LQ, MD, HQ and MD with cirrus backscatter profiles.

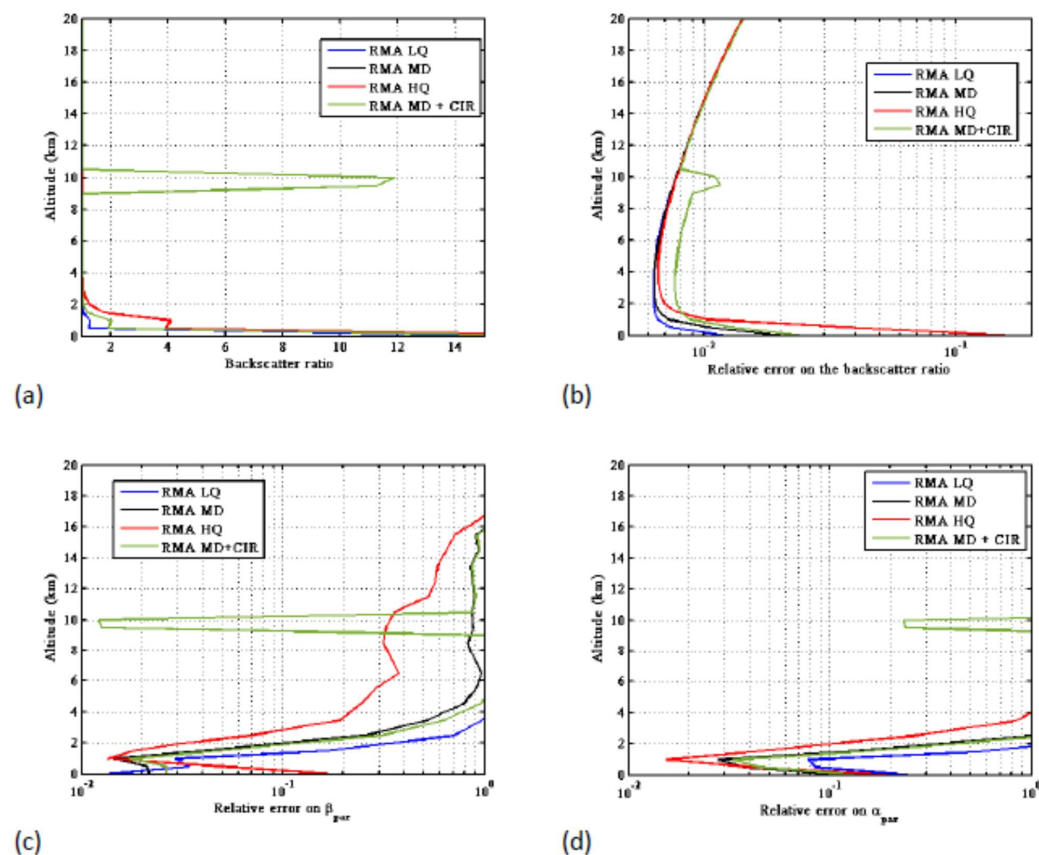


Fig. 9 Actual backscatter ratio (a) Relative statistical error (standard deviation) on the backscatter ratio measurement (b) on the backscattering coefficient (c) and on the extinction coefficient (d) for LQ, MD, HQ and MD plus cirrus backscatter profiles.



The backscatter ratio can be retrieved with a relative statistical error of approximately 1 % over the troposphere except above the tropical tropopause and in the first kilometer above the surface where the error can increase up to 10 % due to the strong atmospheric attenuation for the HQ aerosol distribution. One can argue that in this case the scattering ratio corresponds to a more polluted situation ($R_b \sim 3$), where such an accuracy may be more acceptable than bias for assimilation purpose (Ma et al., 2019).

The error on α_{par} is somehow more important than for β_{par} due to the simple linear signal derivative procedure used in this analysis. More suitable and more efficient methods can be applied by distinguishing and processing homogenous aerosol layers one by one. Nevertheless, as seen in Fig. 9c, the relative error remains smaller than 10 % in the boundary layer, using this derivative method. Other approaches can be used to reduce this error, taking advantage of the simultaneous retrieval of backscattering and transmission extending previous analyses (Young, 1995; Young et al., 2009) possibly using variational methods. This is however beyond the scope of this paper.

The errors due to the use of meteorological re-analyses for the retrieval of the particulate backscatter are discussed in Appendix C. These errors, often qualified as random-systematic, are inherent to the modelling of the molecular scattering which depends on the temperature profile using an a priori guess and include random errors. They are for a large part common to all the retrieval processes whatever the measurement method. Indeed, as shown in Fig C2, assuming a temperature error of 2K, the relative errors induced on β_{par} and α_{par} are of the order of 1 % and 10 % respectively, that is, of the same order as the instrumental random errors.

4.5 Measurements in presence of semi-transparent clouds

The evolution of the measurement performance is modelled for the RMA-MD scattering profile with the addition of a semi-transparent (cirrus) cloud whose optical thickness is varied from 0 to 0.5. Fig.10 shows the SNR, errors in HLOS wind, backscatter and extinction coefficients at different levels. The measurements above the cloud are almost not affected, despite the increase of the background noise due to the increasing cloud albedo. Performance in the PBL below the cloud is degraded by the lower atmospheric transmission while performance in the cloud is improved due to the higher return signal.

4.6 Instrument calibration dependent errors

An error in the determination of the instrumental sensitivities a_i and instrumental contrasts M_i , propagates to the calculation of complex Q (Eq. A4) and then to the LOS velocity and scattering ratio measurements. These calibration errors apply equally on the reference complex signal Q_r but with a slightly different impact due to a different modulation factor and to the Doppler phase shift. The consecutive measurement errors are interference-phase dependent with a negligible averaged value but with a perceptible amplitude.

The maximal value of the LOS velocity error and the relative error on the scattering ratio measurement are presented on Fig 11, for a 10^{-3} relative error in the sensitivity calibration of a single QMZ channel, as a function of the scattering ratio and for LOS wind speeds extending from 0 to 100 ms^{-1} .



510 The sensitivity parameters a_i and modulation parameters M_i are determined by the calibration sequence presented in Sect. 3.5. As tested on HSR-LNG QMZ, fits of the recorded signals with the model described by Eq. (1) are obtained with a correlation better than 99.5 % (Bruneau et al., 2015). This allows to determine the a_i and M_i parameters with a high accuracy.

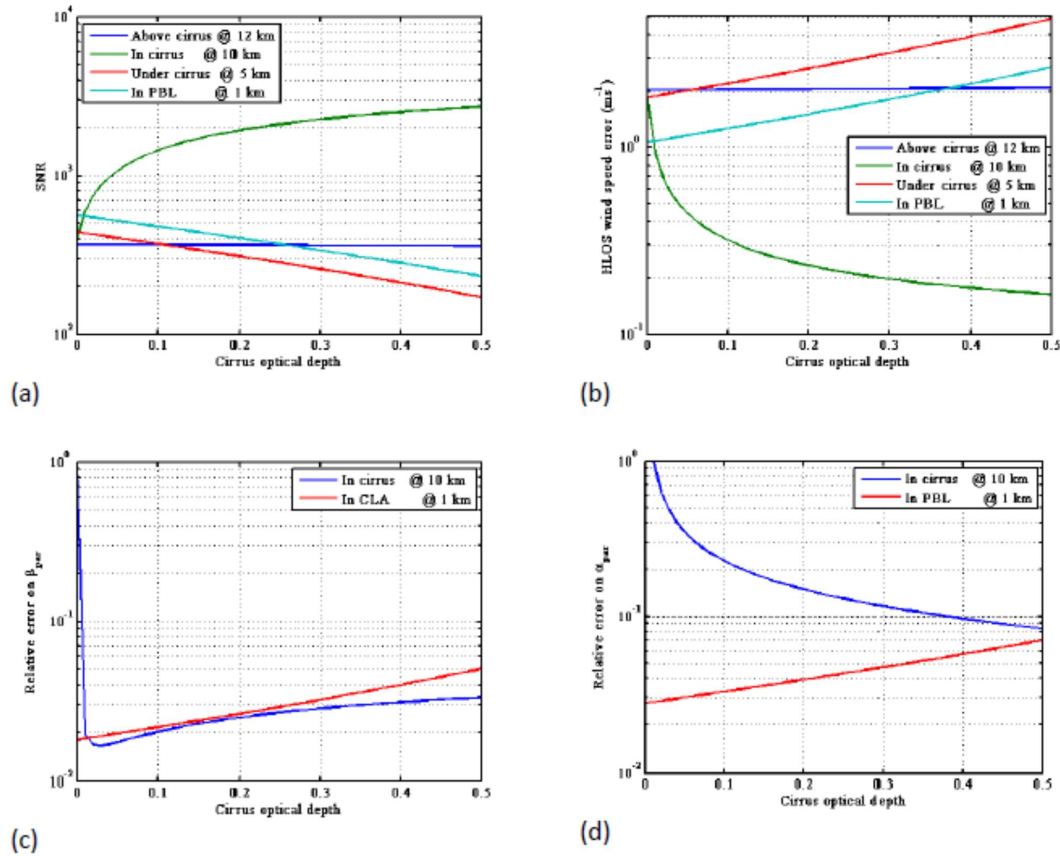


Fig 10. (a) SNR, (b)HLOS wind speed error, (c) β_{par} relative error and (d) α_{par} relative error as a function of the cirrus vertical optical depth for different measurement altitudes.

5. Discussion

From parameters reported in Table 2, and results presented in the previous section, the performance in the determination of the HLOS wind speed can be optimized using a single compact QMZ. This MZI design offers a high value alternative to the whole cascade FP-DE and Fizeau interferometric system implemented in the present Aeolus design. A better overall



performance is expected to be achieved as based on realistic parameters derived from airborne operation with a minimized risk and increased design compactness and reliability. As compared to the 35° LOS angle of Aeolus, a higher 45° slant viewing allows to increase the LOS horizontal projection and also to get the same vertical resolution with a larger range gate and hence collected signal, slightly increasing performance, despite a lower atmospheric transmission.

One can see from Table 3 that the performance is better than the requirement in the lower stratosphere and free troposphere (even keeping a 1000 m and 500 m resolution, respectively, see Fig. 8), and slightly worse in the boundary layer for a varying vertical resolution. This does not appear critical as meteorological stations are providing measurements at the surface as well as spaceborne scatterometers over ocean. We believe that the achievable performance is still acceptable and does not justify the introduction of a second interferometer optimized for particulate scattering, which would significantly complicate the instrument.

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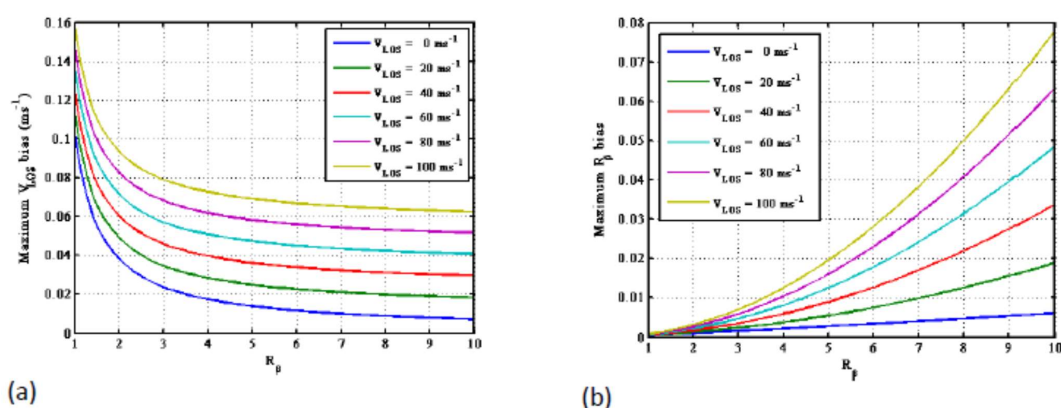


Fig. 11. Maximum measurement bias for a 10^{-3} relative calibration error in the sensitivity a_i of one QMZ channel, as a function of the scattering ratio (a) LOS velocity error, (b) scattering ratio relative error for different LOS velocities (V_{LOS}).

535

Aeolus has been studied to be operated with a 130 mJ / 100 Hz emitter to meet the requirements but the operational system only allowed using a 65 mJ / 50 Hz emitter. Considering equations of errors with SNR depending on the square root of the emitted power (product of energy and repetition rate) at high SNR, and the reduction in the horizontal resolution (90 km instead of 50 km) ALADIN in space should produce a statistical error increased by a factor 2.7 as compared to its original dimensioning. The actual standard deviation of the Aeolus measurements achieved is about 4 ms^{-1} with the Rayleigh channel in the free troposphere and about 2 to 3 ms^{-1} with the Mie channel in the lower two kilometers, both for a 90 km horizontal resolution (Reitebuch et al., 2019, Geiss et al., 2019, Witchas et al., 2020, Martin et al, 2020). In the troposphere the actual Aeolus performance appears to be mostly degraded by transmission losses by a factor 1.5 to 2 with respect to simulations (Reitebuch et al., 2019). This is mostly due to the narrow field of view requiring high pointing accuracy due to the low DE-

540



545 FP acceptance angle. The field compensated QMZ offers more flexibility and its actual performance are expected to be closer to the simulations.

As presented in Sect. 4.6, an instrument calibration error in the sensitivity a_i on one channel by 10^{-3} propagates to wind measurements with a maximum LOS velocity error of 0.15 ms^{-1} (0.2 ms^{-1} in HLOS) depending on the interference phase (the bias averaged on all phases is negligible). This can appear relatively sensitive, but in fact the calibration process given in
550 Sect. 3.5 has been validated with the HSRD-LNG airborne system and indeed allows such a calibration accuracy. Actually, the average 0.12 ms^{-1} LOS wind speed bias observed with this system on a long series of ground returns (Bruneau et al, 2015) shows that the residual phase dependent bias is very limited. To monitor system health and avoid degradation of performance with time due to ageing of optical components, the calibration operation is nevertheless to be done regularly. Note also that the high sensitivity of the wind speed measurement with regard to the instrument calibration is inherent to the
555 spectral discrimination based on molecular scattering because of its large spectral distribution. For comparison, the characteristics of the DE-FP used at the Observatoire de Haute Provence (Souprayen et al, 1999b) lead to a bias of 0.38 ms^{-1} in the LOS velocity measurement for a 10^{-3} sensitivity error on one channel. The actual ALADIN HLOS wind velocity measurement bias has been estimated to be around 1.5 ms^{-1} with a latitude and orbit phase dependence (Martin et al, 2020).

The wind measurement accuracy in presence of a cirrus cloud of variable optical depth, as frequently observed in the
560 atmosphere, is not degraded so much for low cloud optical depths (Fig 10b). HLOS wind speed errors are maintained below 3 m.s^{-1} and 5 m.s^{-1} at altitudes of 1 km (top of the atmospheric boundary layer) and 5 km respectively in presence of cirrus with optical depths up to 0.5. Measurement errors in cirrus clouds remain low, allowing higher horizontal resolutions, which is interesting as they are close to the upper tropospheric jet.

Unlike usual HSRL technique (as that to be used in ATLID), the QMZ does not attempt to separate molecular and particulate
565 signals but delivers directly, in addition to the total signal S_{tot} , the atmospheric interference modulation M_{atm} from which the backscatter ratio R_{β} is derived. The performance of the QMZ in terms of backscatter measurements are thus expressed slightly differently than from those given by the EarthCare instrument. A separation of the molecular and particulate components of the signal is however possible by computing Eq. A11. The retrieval of extinction can then be done in a very similar way as for conventional HSRL. We nevertheless can see (Fig. 9b) that the backscatter ratio random error of 1 % in
570 cirrus clouds with scattering ratios larger than 5 (with 500 m vertical and 50 km horizontal resolutions), as it is the case for most semi-transparent cirrus clouds, compares favorably with the radiometric errors of 50 % and 15 % for particulate and molecular signals respectively required for EarthCare (for 100 m / 300 m vertical and 10 km horizontal resolutions - Hélière et al, 2012). Starting from the backscatter ratio (as well as from separated molecular and particulate signals) it is further needed to use an atmospheric model to derive the particulate backscatter and extinction coefficients coefficient (Eq. A12,
575 A13). An error of about 2K in the temperature leads to an error smaller than 2 % for β_{par} and 10 % for α_{par} in the PBL. These errors account for the molecular density, as for conventional HSRL, and for the M_{mol} temperature dependence which can be seen as the equivalent of the temperature dependence of the transmission of the molecular channel in conventional technique. The QMZ performance simulations have been conducted with a vertical resolution derived from the Aeolus requirements



and A-CCD characteristics taken from the ALADIN design. New A-CCDs with smaller time gates as those planned for
580 ATLID would result in a higher vertical resolution compatible with the EarthCare requirements. Nevertheless, with the
proposed configuration aiming, as a first priority, to the wind measurements the performance is first requesting a low reading
noise per range gate as compared to the useful signal, and reducing this value only allows a better range resolution for wind
and backscatter measurements with an accuracy pertinent for the survey of dynamics and cloud parameters at the global
scale
585 In addition, the QMZ measurement is performed with a reference taken from the emitted pulse and does not require any
frequency locking of the laser or the interferometer. The instrument is also tolerant to emitter frequency linewidth, jitter (200
MHz RMS during half a second, included in the performance simulations) and unaffected by long-term frequency drifts. The
unfolded measurement range is $\pm 832 \text{ ms}^{-1}$ accepting the maximal earth rotational speed of 463 ms^{-1} without the need of a
latitude dependent LOS correction.
590 The design includes a relatively wide field angle ($100 \text{ } \mu\text{rd}$, about 5 times larger than that of ALADIN, similarly to ATLID
and CALIOP instruments) which relaxes the requirements on the emitter-receiver alignment. This wider aperture angle
allows also an increase in the emission cone, and thus a reduction in the energy density at the surface, potentially also
allowing the emission of other wavelengths in eye safety conditions. The solar background is however kept at the same level
due a more severe interference filtering.

595 6. Conclusion

The Aeolus mission has emphasized new capabilities of lidar in space for wind measurements. The discussion of the Aeolus
follow-on programme allows to revisit some of the choices made that have revealed technical constraints that limit
achievable performance. In this paper we explore possible solutions to relax some of the constraints identified, keeping
initial performance objectives identified. A new spaceborne lidar design for the horizontal wind speed measurement using a
600 single QMZ interferometer technique is proposed for consideration for operational missions following Aeolus and EarthCare,
taking advantage of advances made for them both. The changes in the payload are minimized as the lidar system would use
the same telescope aperture as Aeolus and a comparable 355-nm emitter. The instrument is optimized for wind
measurements over the whole troposphere as for Aeolus. It also authorizes the retrieval of cloud and aerosol optical
properties with a satisfying accuracy as compared to the EarthCare mission. It further allows the implementation of a
605 depolarization channel without degrading overall performance. A single Quadri-channel Mach-Zehnder interferometer
(QMZ) is selected to replace the cascade double-edge Fabry Perot (DE-FP) and Fizeau interferometers. The design of the
MZI can provide robustness and compactness, and can be developed with molecular adherence as for Aeolus interferometers.
The MZI offers a large acceptance angle reducing mechanical and thermal constraints. These choices allow to overcome the
limitations induced by the need of a narrow field of view for the DE-FP and the use of a high accuracy alignment control
610 and/or a drastic reduction in thermo-mechanical constraints. The MZI design also relaxes constraints on the telescope



focusing. There is no bias induced by aerosol scattering on Doppler shift, which also represents a strong advantage for the needed bias corrections in the assimilation process. For the retrieval of aerosol and cloud backscattering properties, the necessary knowledge of the molecular spectrum (including Brillouin scattering) can be derived from meteorological analyzes to elaborate products with the requested accuracy.

- 615 The performance study shows that a statistical error on the horizontal wind speed better than 2 ms^{-1} can be achieved, from the boundary layer up to the tropopause for a vertical resolution of 500 m and a 50 km horizontal resolution. Accordingly, the measurement bias can be as small as 0.2 m.s^{-1} . In addition to wind speed, the instrument can retrieve the backscatter ratio with a relative random error of 1 to 2 % over most of the troposphere, well matched to high cloud study and radiatively significant aerosol load. The performance analysis on the QMZ interferometer is supported by measurements performed in
 620 the frame of the UV HSRD-LNG airborne lidar developed and operated by LATMOS.

Appendix A. Signal processing and noise dependent measurement errors

Signal processing

- In this appendix we refer to BP03 to recall expressions of errors of atmospheric wind speed and attenuated backscattering coefficients using a QMZ interferometer. Let us start from the optical lidar signal S_{atm} (in number of photons) at the QMZ
 625 input as

$$S_{\text{atm}}(R) = S_{\text{mol}}(R) + S_{\text{par}}(R) = \frac{EAT_{\text{inst}}}{h\nu} \int_{r=R-\delta R/2}^{r=R+\delta R/2} \frac{(\beta_{\text{mol}}(r) + \beta_{\text{par}}(r))}{r^2} T_{\text{atm}}^2(r) dr \quad (A1)$$

where E is the emitted energy, A the telescope area, T_{inst} the instrumental (emission and reception) transmission, $h\nu$ is the emission photon energy, R is the range, $T_{\text{atm}}(r) = \exp[-\int_0^r (\alpha_{\text{mol}}(r') + \alpha_{\text{par}}(r')) dr']$ and the mol and par subscripts are related to the molecular and particulate scattering respectively.

- 630 The signal delivered by each channel of the A-CCD outputs (in photoelectrons), for $i = 1$ to 4, can be written

$$S_i = \frac{S_{\text{tot}}}{4} a_i \left[1 + M_i M_{\text{atm}} \sin \left(\varphi + (i-1) \frac{\pi}{2} \right) \right] + S_{b,i} \quad (A2)$$

a_i is the relative photometric efficiency of channel i with $\sum_{i=1}^4 a_i = 4$, M_i is the intrinsic modulation of channel i . $S_{b,i}$ is the background signal due to solar detected light and A-CCD intrinsic noise

As all the photons incident on the MZI arrive on the detectors, we have:

$$635 \quad S_{\text{tot}} = \frac{4}{\sum_{i=1}^4 M_i^{-1}} \sum_{i=1}^4 \frac{(S_i - S_{b,i})}{a_i M_i} = \eta S_{\text{atm}} \quad (A3)$$

where η is the mean quantum efficiency of the detectors.

φ is the interference phase and M_{atm} is the atmospheric effective interference modulation given by the molecular and particulate atmospheric backscattered signals as



$$M_{atm} = \frac{M_{par}(R_{\beta}-1)+M_{mol}}{R_{\beta}} \quad (A4)$$

- 640 where the lidar backscattering ratio R_{β} is defined by the ratio of the total to the molecular backscattering as $R_{\beta} = \frac{\beta_{mol}+\beta_{par}}{\beta_{mol}}$. Depending on the backscattering ratio, the atmospheric modulation coefficient M_{atm} varies between M_{mol} and M_{par} , the interference modulations given by pure molecular and particulate signals respectively. After subtraction of the background, the four signals are combined two-by-two in order to produce a complex signal Q (with in-phase and quadrature components):

$$645 \quad Q = Q_2 + iQ_1 \quad \text{with} \quad Q_1 = \frac{a_3(S_1-S_{b1})-a_1(S_3-S_{b3})}{a_3M_3(S_1-S_{b1})+a_1M_1(S_3-S_{b3})}, \quad Q_2 = \frac{a_4(S_2-S_{b2})-a_2(S_4-S_{b4})}{a_4M_4(S_2-S_{b2})+a_2M_2(S_4-S_{b4})} \quad (A5)$$

The interference phase φ , is obtained by the argument of the complex signal Q :

$$\varphi = \arg(Q) \quad (A6)$$

Subtracting the reference phase φ_r , obtained by the same way on a highly attenuated pick-up of the laser emission one can obtain the line-of-sight (LOS) particles velocity V_{LOS} with:

$$650 \quad V_{LOS} = \frac{c\lambda}{4\pi\Delta}(\varphi - \varphi_r) \quad (A7)$$

where Δ is the MZI optical path difference, λ the operating wavelength and c the light celerity both in vacuum. Using this differential approach V_{LOS} can be obtained on the whole measurement range $\pm \frac{c\lambda}{4\Delta}$ without the need to lock the emitting frequency with the interference phase.

The atmospheric modulation is obtained by the modulus of Q divided by the modulus of the reference signal:

$$655 \quad M_{atm} = \frac{|Q|}{|Q_r|} \quad (A8)$$

One can see, from Eq. A3 that, once M_{mol} and M_{par} are determined, M_{atm} is giving access to the scattering ratio R_{β} .

$$R_{\beta} = \frac{M_{par}-M_{mol}}{M_{par}-M_{atm}} \quad (A9)$$

Separated molecular and particulate signals can be obtained (with the same instrumental constant) using:

$$S_{mol} = \frac{M_{par}-M_{atm}}{M_{par}-M_{mol}} \frac{S_{tot}}{\eta} \quad ; \quad S_{par} = \frac{M_{atm}-M_{mol}}{M_{par}-M_{mol}} \frac{S_{tot}}{\eta} \quad (A10)$$

- 660 but this step is not necessary for the retrieval of the particulate backscatter and extinction coefficients obtained as follows:

$$\beta_{par} = \beta_{mol}(R_{\beta} - 1) = \beta_{mol} \frac{M_{atm}-M_{mol}}{M_{par}-M_{atm}} \quad (A11)$$

The total particulate optical depth over the vertical column can be derived from the total signal and the scattering ratio



$$\tau_{par} = \frac{1}{2} \ln \left(\frac{R_{\beta} \cdot \beta_{mol}}{r^2 s_{tot}} \right) - \tau_{mol} \quad (A12)$$

One way to determine the particulate extinction coefficient is to derive equation A12 giving the particulate optical depth with altitude z as

$$\alpha_{par} = \frac{1}{2} \frac{d}{dz} \left(\ln \left(\frac{R_{\beta} \cdot \beta_{mol}}{r^2 s_{tot}} \right) \right) - \alpha_{mol} \quad (A13)$$

which allows to get rid of calibration. This method is sensitive to noise, and other ways to derive extinction can provide better results. We will however use this conservative approach for the sake of simplicity.

Preliminary evaluation of M_{par} and M_{mol}

Assuming Gaussian spectral profiles, the two interference modulations M_{mol} and M_{par} can be expressed as a function of the optical path difference Δ and the 1/e linewidth γ (expressed in wavenumber) of the related spectral functions of atmospheric scattering convolved by the laser emitted width, as given in BP03, so that

$$M_s = \exp[-(\pi \gamma_s \Delta)^2] \quad \text{where } s \text{ stands for mol or par} \quad (A14)$$

For the particle scattering, γ_{par} is defined as a first approximation by the laser source linewidth γ_{las} . Assuming γ_{las} on the order of $3.10^{-3} \text{ cm}^{-1}$ (or 100 MHz, corresponding to the transform limit of a 3ns pulse), we see that we obtain $\gamma_{par} \Delta \approx 10^{-2}$ and $M_{par} \approx 1$ for an OPD value of 3 cm. We thus remain in the case of a high contrast for the particulate signal, with some margin on the laser linewidth.

For the molecular scattering, γ_{mol} is dominated by the thermal molecular velocity:

$$\gamma_{mol} = \frac{2}{\lambda c} \sqrt{\frac{2k}{m}} \quad (A15)$$

and is about $7.10^{-2} \text{ cm}^{-1}$, which implies that $\gamma_{mol} \cdot \Delta \approx 0.2$ and $M_{mol} \approx 0.6$ for an OPD of 3 cm.

Noise dependent statistical error

We give here the random error depending on the detection noise. Assuming, for simplicity, that all the relative sensitivities a_i are equal to 1 and that all the instrumental modulations M_i are equal to M_0 , the standard deviation of V_{LOS} is given in BP03 as

$$\sigma(V_{LOS}) = \frac{c\lambda}{4\pi\Delta SNR M_0 M_{atm}} \sqrt{1 - \frac{1}{2} F_B M_0^2 M_{atm}^2 \sin^2(2\varphi)} \quad (A16)$$

where SNR is the signal-to-noise ratio and F_B is a correlation coefficient between the four detection channels given by $F_B = \frac{s_{tot} - s_b}{s_{tot} + s_b}$ where s_{tot} is the number of signal photo-electrons and s_b is the total number of photo-electrons of the radiative and



detection background (both summed with the ponderation given in Eq. A3). We also assume here that the background can be measured over a long duration and subtracted for any measurement pixel with a negligible impact on bias and SNR. This assumes that the background noise is taken with a much higher SNR than the atmospheric signal. For accurate measurements, the SNR needs to be high so we have $S_{\text{tot}} \gg S_b$ and $F_B \approx 1$ leading to a minimum square-root factor. If, as we propose, the laser frequency is not locked to the interferometer, the phase can take any value between 0 and 2π . For the performance assessment of the instrument, we thus average the error on φ and obtain:

$$\sigma(V_{\text{LOS}}) = \frac{c\lambda}{4\pi\Delta} \frac{\sqrt{2}}{\text{SNR}} \frac{1}{M_0 M_{\text{atm}}} \sqrt{1 - \frac{M_0^2 M_{\text{atm}}^2}{4}} \quad (\text{A17})$$

The factor $D_c = \frac{1}{M_0 M_{\text{atm}}} \sqrt{1 - \frac{M_0^2 M_{\text{atm}}^2}{4}}$ can be seen as a degradation factor on the wind error due to the contrast degradation. For high scattering ratios we have $M_{\text{atm}} \approx M_{\text{par}} \approx 1$, and assuming a perfect MZI, we have $M_0 \approx 1$, so that $D_c = 1$. In clear air, for which we have $M_{\text{atm}} \approx M_{\text{mol}} \approx 0.6$, the degradation factor is $D_c = 1.6$. This is intrinsic to the QMZ technique, for which error is reduced in presence of particle scattering.

The relative standard deviation of the statistical error on R_β is linked to the error on modulus of Q (Eq. A8). After averaging over φ , it can be expressed as (BP03)

$$\frac{\sigma(R_\beta)}{R_\beta} = \frac{\sqrt{2}}{\text{SNR}} \frac{R_\beta}{M_0 (M_{\text{par}} - M_{\text{mol}})} \sqrt{1 - \frac{3}{4} F_B M_0^2 M_{\text{atm}}^2} \quad (\text{A18})$$

We can then derive the error on β_{par} :

$$\sigma(\beta_{\text{par}}) = \frac{\sqrt{2}}{\text{SNR}} \frac{R_\beta (\beta_{\text{par}} + \beta_{\text{mol}})}{M_0 (M_{\text{par}} - M_{\text{mol}})} \sqrt{1 - \frac{3}{4} F_B M_0^2 M_{\text{atm}}^2} \quad (\text{A19})$$

The error on the particulate extinction coefficient is given (after averaging over φ) by (BP03):

$$\sigma(\alpha_{\text{par}}) = \frac{1}{\sqrt{2}\delta r} \frac{1}{\text{SNR}} \sqrt{1 + \frac{2R_\beta^2}{M_0^2 (M_{\text{par}} - M_{\text{mol}})^2} \left(1 + \frac{3}{4} F_B M_0^2 M_{\text{atm}}^2\right) - \frac{R_\beta M_{\text{atm}}}{M_{\text{par}} - M_{\text{mol}}} (1 + F_B)} \quad (\text{A20})$$

Appendix B. MZI aperture

As the beam étendue is maintained over its propagation through the instrument, the product of the useful interferometer aperture D_{MZI} by the apparent source field angle α is equal to the product of the receiving telescope aperture D_{tel} by its field-of-view α_{tel} . Besides, it can be shown that the minimal aperture D_L for a beam propagating on a distance L with a specified D_{tel} by α_{tel} product is $D_L = 2\sqrt{D_{\text{tel}}\alpha_{\text{tel}}L}$. As, for field compensation, a plate of thickness e and optical index n must be inserted in the longer arm of the interferometer with $e = \frac{n}{n^2 - 1} \Delta$, the minimal optical propagation length in the MZI (both



arms are equals) is $L_{min} = \frac{\Delta}{n^2 - 1}$. We can therefore consider that a realistic MZI can be built with an internal propagation length of 2Δ . The resulting useful diameter of the MZI is then:

$$D_{MZI} = 2\sqrt{2D_{tel}\alpha_{tel}\Delta} \quad (B1)$$

- 715 Then, for the instrumental parameters of this study, $D_{tel} = 1.5$ m, $\alpha_{tel} = 0.1$ mrd and $OPD = 3.2$ cm, the useful MZI aperture is $D_{MZI} = 6.2$ mm.

Appendix C. Atmosphere-modeling dependent errors on particulate scattering retrieval

- As already pointed out, the wind speed measurement does not require any atmospheric modeling (unlike DE-FP operation) since the interference phase determination is independent of the actual interference modulation M_{atm} . Nevertheless, the
- 720 particle backscatter and extinction retrieval (Eq. A8 –A10) requires the knowledge of β_{mol} , M_{mol} and M_{par} obtained with some atmosphere modeling and assumptions. Wind speed variation or turbulence in the probed atmosphere can induce a variation in M_{par} that is not taken into account in the simulation. With a Doppler shift of 5.6 MHz per ms^{-1} the broadening of the particle backscattered spectrum is very small as compared to an emitted linewidth of 100-200 MHz and the incidence on M_{par} is however negligible.
- 725 The calculation of M_{mol} is not as straightforward as that given by Eq. A14 since it must include the effect of Brillouin scattering. For an operating wavelength of 355 nm the Brillouin doublet separation is much smaller than the thermal linewidth. The resulting spectrum is a broadening and distortion of the Gaussian profile. A correction can be derived from a model (Tenti et al., 1974) or measurements (Witschas et al., 2008). The effect is roughly proportional to the pressure with a maximum relative broadening of 1.7 % near ground. This broadening causes a relative decrease of M_{mol} of $6 \cdot 10^{-3}$ as
- 730 compared as calculated from the Gaussian thermal line (Fig C1). The Brillouin effect must be included in the calculation of M_{mol} but the uncertainty on the actual atmospheric pressure leads to a second-order variation of M_{mol} that can be neglected. This constraint is less critical here for molecular backscattering quantification than for wind measurements using a DE-FP device (Dabas et al., 2008).

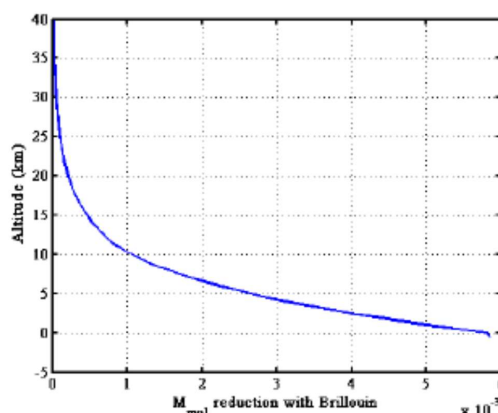


Fig. C1 Reduction of M_{mol} caused by the Brillouin effect

In contrast, the dependence of M_{mol} on temperature, given in Fig C2a for a standard profile, must be taken into account. γ_{mol} varies as the square root of temperature (see Eq A15), which leads to a variation of about 20 % over the atmospheric column. The knowledge of atmospheric temperature to within 2K allows reducing uncertainty on M_{mol} to less than 0.5 %.

740 The other source of error on the particle backscatter coefficient comes from the modeling of the molecular backscatter coefficients which is proportional to the molecular density and then depends on the temperature vertical profile. The sensitivity of the relative error with regard to temperature is presented on Fig. C2b.

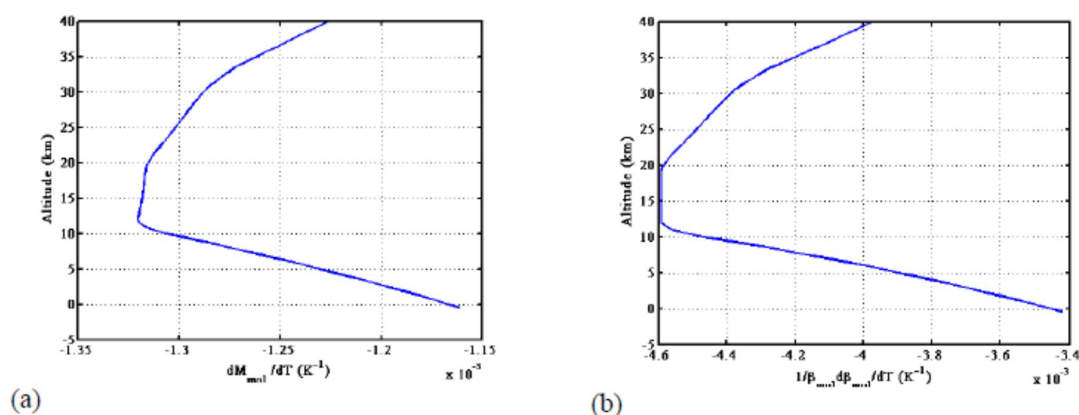


Fig C2 Temperature modeling sensitivity for (a) M_{mol} and (b) β_{mol}



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Assuming random and independent errors on these parameters, the relative error on the retrieval of the particle backscattering coefficient caused by atmospheric analysis errors is given by:

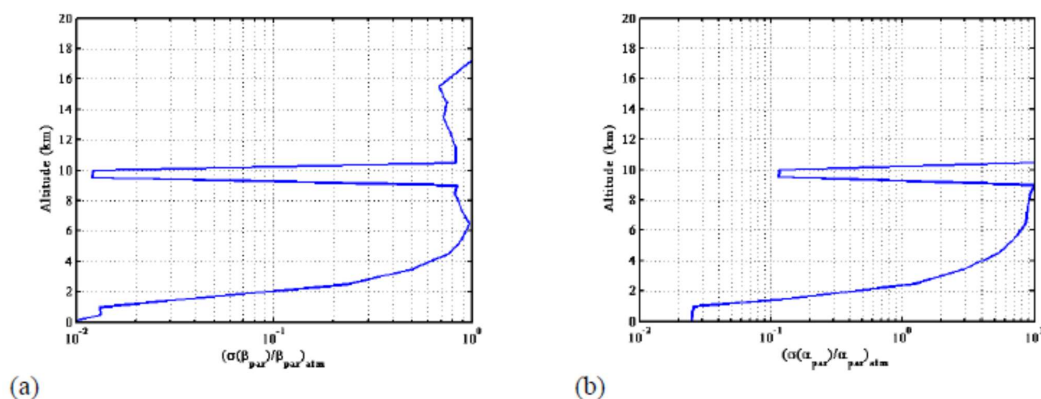
$$\left[\frac{\sigma(\beta_{\text{par}})}{\beta_{\text{par}}} \right]_{\text{atm}} = \left[\left(\frac{\sigma(\beta_{\text{mol}})}{\beta_{\text{mol}}} \right)^2 + \left(\frac{\sigma(M_{\text{mol}})}{M_{\text{par}} - M_{\text{mol}}} \frac{R_{\beta}}{R_{\beta} - 1} \right)^2 \right]^{1/2} \quad (C1)$$

750 In clear air the uncertainty on M_{mol} dominates the total error but in regions where the aerosol load is significant ($R_{\beta} \geq 2$) the contributions of the uncertainty on β_{mol} and M_{mol} are of the same order of magnitude. Figure C3a shows the relative error on β_{par} given by the uncertainties on the atmosphere temperature modeling $\sigma(T) = 2\text{K}$, computed for the MD+CIR aerosol profile and a standard temperature profile. This error must be quadratically added to the error caused by detection noise and presented in Sect. 4.4. In the regions where the aerosol load is significant, the atmosphere-modeling error is slightly lower than the SNR dependent error and the total relative standard deviation remains on the order of a few percent.

755 Considering Eq. A13 and assuming that the temperature errors between two range gates are independent, we obtain for the error on the particle extinction coefficient caused by the atmosphere modelling uncertainties:

$$\left[\sigma(\alpha_{\text{par}}) \right]_{\text{atm}} = \left[\left(\frac{\cos(\theta)}{\sqrt{2}\Delta z} \frac{\sigma(\beta_{\text{mol}})}{\beta_{\text{mol}}} \right)^2 + \left(\frac{\cos(\theta)}{\sqrt{2}\Delta z} \frac{\sigma(M_{\text{mol}})}{M_{\text{par}} - M_{\text{mol}}} \right)^2 + (\sigma(\alpha_{\text{mol}}))^2 \right]^{1/2} \quad (C2)$$

where Δz is the vertical resolution and θ the LOS nadir angle.



760 Fig. C3 (a) Relative standard deviation on β_{par} , (b) standard deviation on α_{par} for the MD+CIR caused by uncertainties of 2K in the atmospheric analysis.



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Authors contributions : JP is at the origin of the paper and has established the main issues. DB defined the instrumental design and established the performance modeling. Both authors participated in the drafting.

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