

Measurement of iodine species and sulfuric acid using bromide chemical ionization mass spectrometers

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Abstract. Iodine species are important in the marine atmosphere for oxidation and new-particle formation. Understanding iodine chemistry and iodine new-particle formation requires high time resolution, high sensitivity, and simultaneous measurements of many iodine species. Here, we describe the application of bromide chemical ionization mass spectrometers (Br-CIMS) to this task. During ~~the iodine new-particle formation~~oxidation experiments in the Cosmics Leaving OUtdoor Droplets (CLOUD) chamber, we have measured gas-phase iodine species and sulfuric acid using two Br-CIMS, one coupled to a Multi-scheme chemical IONization inlet (Br-MION-CIMS) and the other to a Filter Inlet for Gasses and AEROsols inlet (Br-FIGAERO-CIMS). From offline calibrations and inter-comparisons with other instruments attached to the CLOUD chamber, we have quantified the sensitivities of the Br-MION-CIMS to HOI, I₂, and H₂SO₄ and obtain detection limits of 5.8×10⁶, ~~3.86.3~~3×10⁵, and 2.0×10⁵ molec cm⁻³, respectively, for a 2-min integration time. From binding energy calculations, we estimate the detection limit for HIO₃ to be 1.2×10⁵ molec cm⁻³, based on an assumption of maximum sensitivity. Detection limits in the Br-FIGAERO-CIMS are around one order of magnitude higher than those in the Br-MION-CIMS; for example, the detection limits for HOI and HIO₃ are 3.3×10⁷ and 5.1×10⁶ molec cm⁻³, respectively. Our comparisons of the performance of the MION inlet and the

46 FIGAERO inlet show that bromide chemical ionization mass spectrometers using either atmospheric pressure or
47 reduced pressure interfaces are well-matched to measuring iodine species and sulfuric acid in marine environments.

48

49 **1 Introduction**

50 Reactive iodine species are released into the atmosphere mainly by biological processes in marine environments (i.e.
51 from macro- and micro-algae) (McFiggans et al., 2004), O₃ deposition on the sea surface (Carpenter et al., 2013), as
52 well as from the sea ice (Spolaor et al., 2013) and snowpack in the polar region (Raso et al., 2017). Once emitted,
53 iodine species can modify atmospheric oxidative capacity via a chain of catalytic reactions with O₃ that form iodine
54 oxides, leading to about 20–28 % of O₃ loss in the marine boundary layer (Prados-Roman et al., 2015; Sherwen et al.,
55 2016). Through convection, reactive iodine species can be transported from the lower troposphere to the upper
56 troposphere-lower stratosphere, causing one third of the iodine-induced ozone loss in the upper troposphere-lower
57 stratosphere (Koenig et al., 2020). Another important effect of iodine species is their contribution to atmospheric new-
58 particle formation. O’Dowd et al. (O’Dowd et al., 2002) showed that particles are produced from condensable iodine-
59 containing vapours at a coastal location (Mace Head in Ireland). Recent studies have demonstrated that iodine oxoacids
60 (iodous acid, HIO₂ and iodic acid, ~~iodic acid~~ (HIO₃)) dominates the ~~charged~~-iodine cluster formation processes (He et
61 al., 2021a, 2021b), and drives the bursts of freshly-formed particles in coastal regions (Sipilä et al., 2016). This process
62 thereby may enhance cloud condensation nuclei formation, affecting climate both directly and indirectly (Saiz-Lopez
63 et al., 2012; Simpson et al., 2015).

64

65 Understanding iodine chemistry and iodine driven new-particle formation requires high time resolution, high
66 sensitivity, and simultaneous measurements of iodine species. However, this has been a long-standing challenge due
67 to their low abundance and short atmospheric lifetimes. Previous studies have achieved detection of relatively more
68 abundant molecular iodine (I₂), iodine monoxide (IO), and iodine dioxide (OIO) via optical spectroscopy, such as
69 differential optical absorption (Leigh et al., 2010), cavity ring-down (Bitter et al., 2005), cavity enhanced absorption
70 (Vaughan et al., 2008), laser-induced fluorescence (Dillon et al., 2006), and resonance fluorescence (Gómez Martín
71 et al., 2011). The spectroscopic techniques are invaluable; however, their very specificity limits them to the detection
72 of a few iodine compounds, and they are less sensitive to other iodine species that have congested or broad absorption
73 cross-sections/features, such as hypoiodous acid (HOI) and iodic acid (HIO₃).

74

75 Another commonly used technique is mass spectrometry; it has a fast response time and a low detection limit, but
76 extra calibration efforts are needed for the quantification of the detection sensitivity. For example, photoionization
77 (Gómez Martín et al., 2013) and chemical ionization mass spectrometry (CIMS) have been employed to detect a suite
78 of halogen species. Reagent ions used with CIMS include: SF₅⁺ for HCl and ClONO₂ (Marcy et al., 2004); iodide (I⁻)
79 for atmospheric chlorine and bromine species such as ClONO₂, Cl₂, ClO, BrO, and BrCl (Kercher et al., 2009; Lee et
80 al., 2018; Tham et al., 2016); superoxide (O₂⁻) for molecular iodine (I₂) (Finley and Saltzman, 2008); and both nitrate
81 (NO₃⁻) (Sipilä et al., 2016) and protonated water (H₃O⁺) (Pfeifer et al., 2020) for HIO₃. The nitrate-CIMS and H₃O⁺-

82 CIMS suffer from the limited analyte affinity to the reagent ions. The iodide-CIMS can effectively measure chlorine
83 and bromine species, but it is not suitable to detect iodine species due to the ambiguity in peak identification.

84
85 The bromide ion (Br⁻) exhibits an affinity to a wide spectrum of iodine containing species. Br-CIMS has been
86 routinely used to measure chlorine species (Lawler et al., 2011), HO₂ radicals (Sanchez et al., 2016), organic vapors
87 and sulfuric acid (Rissanen et al., 2019), and nitric acid (Wang et al., 2020). Like chlorine species, iodine species are
88 known to cluster with bromide ions via halogen (or hydrogen) bonds; as such, here we explore using the Br-CIMS to
89 measure gas-phase iodine species and sulfuric acid simultaneously at concentrations relevant to the marine
90 boundary layer. In this study, we demonstrate the detection of various gas-phase inorganic iodine species with the Br-
91 CIMS and explore the effect of relative humidity (RH) on that detection. We then quantify the sensitivities of several
92 gas-phase halogen-iodine species via inter-method calibration, offline calibration, and quantum chemical calculations.
93 Finally, we compare the performance of Br-MION-CIMS and Br-FIGAERO-CIMS and show that both of them are
94 well-suited for iodine species measurement in the atmosphere marine boundary layer.

95

96 **2 Methodology**

97 **2.1 The CLOUD facility**

98 We conducted measurements and instrument inter-comparison at the CERN CLOUD facility, a 26.1 m³
99 electropolished stainless-steel chamber that enables new-particle formation experiments simulating the typical range
100 of tropospheric conditions with scrupulous cleanliness and minimal contamination (Duplissy et al., 2016; Kirkby et
101 al., 2011). The CLOUD chamber is mounted in a thermal housing, capable of keeping temperature constant in a range
102 of -65 °C and +100 °C with ±0.1 °C precision (Dias et al., 2017) and relative humidity commonly between < 0.5%
103 and 80%. Photochemical processes are driven by different light sources, including four 200 W Hamamatsu Hg-Xe
104 lamps with significant spectral irradiance between 250 and 450 nm, and an array of 48 green light LEDs at 528 nm
105 with adjustable optical power up to 153 W. Ion-induced nucleation under different ionization levels is simulated with
106 a combination of electric fields (electrodes at ±30 kV at top and bottom of the chamber) which can be turned on to
107 rapidly scavenge smaller ions, and a high-flux beam of 3.6 GeV pions (π⁺) which enhances ion production when turned
108 on. Mixing is accelerated with magnetically coupled fans mounted at the top and bottom of the chamber. The
109 characteristic gas mixing time in the chamber during experiments is a few minutes. The loss rate of condensable vapors
110 onto the chamber wall is comparable to the condensation sink in pristine boundary layer environments (e.g. 2.2×10⁻³
111 s⁻¹ for H₂SO₄ at 5 °C). To avoid a memory effect between different experiments, the chamber is periodically cleaned
112 by rinsing the walls with ultra-pure water and heating to 100 °C for at least 24 hours, ensuring extremely low
113 contaminant levels of sulfuric acid (< 5×10⁴ cm⁻³) and total organics (< 150 pptv) (Kirkby et al., 2016; Schnitzhofer
114 et al., 2014). The CLOUD gas system is also built to the highest technical standards of cleanliness and performance.

115

116 The dry air supply for the chamber is provided by cryogenic nitrogen (Messer, 99.999 %) and cryogenic oxygen
117 (Messer, 99.999 %) mixed at the atmospheric ratio of 79:21. Ultrapure water vapor, ozone and other trace gases can
118 be precisely added to attain desired mixing ratios at different levels. The total injection rate of the humidified air is

119 fixed at 330 standard liters per minute (slpm) to compensate for the sampling consumption of various instruments.
120 Molecular iodine (I₂) is injected into the chamber by passing a flow of cryogenic nitrogen through a crystalline iodine
121 (I₂, Sigma-Aldrich, 99.999 % purity) reservoir, which is temperature-controlled at 10 °C, to achieve levels of 0.4 to
122 168 pptv in the chamber. The sulfinert-coated injection lines are temperature stabilized to minimize line conditioning
123 effects. High intensity green light emitting diodes (LEDs) are used to photolyze molecular iodine into iodine atoms
124 and initiate the subsequent oxidation reactions in the presence of ozone and water vapor. The 48 green LEDs (light
125 sabre 4, LS4) are mounted in pairs (one upward facing, one downward facing) on a copper cooling bar within a long
126 quartz jacket that protrudes into the chamber in the mid plane. The maximum total optical power output is 153 W,
127 centered on 528 nm. Actinic fluxes are regulated by controlling the number of LEDs used and the set point of
128 individual LEDs. Light fluxes are monitored by two photodiodes and a spectrometer.~~The dry air supply for the~~
129 ~~chamber is provided by cryogenic oxygen (Messer, 99.999 %) and cryogenic nitrogen (Messer, 99.999 %) mixed at~~
130 ~~the atmospheric ratio of 79:21. Ultrapure water vapor, ozone and other trace gases can be precisely added to attain~~
131 ~~mixing ratios at different levels.~~

132 133 **2.2 Br-MION-CIMS**

134 We measured gas-phase iodine species with a bromide chemical ionization atmospheric pressure interface time-of-
135 flight mass spectrometer (Junninen et al., 2010) coupled with a Multi-scheme chemical IONization inlet (Br-MION-
136 CIMS) (Rissanen et al., 2019). The Br-MION inlet consists of an electrically grounded 24 mm inner diameter stainless
137 steel flow tube, attached to an ion source. For the CLOUD measurements, the length of the sampling inlet was ~1.5
138 m and was designed to be in a laminar flow with a fixed total flow rate of ~~20-32~~ standard liters per minute (slpm). An
139 ion filter, operated with positive and negative voltage, was placed at the front of the inlet to filter out any ions in the
140 sample air prior to ion-molecule reaction chamber in the inlet. The reagent ions, bromide (Br⁻) and the bromide-water
141 cluster (H₂O·Br⁻), were produced by feeding 25 standard milliliters per minute (mlpm) of nitrogen (N₂) flow through
142 a saturator containing dibromomethane (CH₂Br₂; > 99.0 %, Tokyo Chemical Industry) into the ion source, where the
143 reagent was ionized by soft X-ray radiation. The resulting ions were then accelerated by a ~~28~~500 V ion accelerator
144 array and focused by a ~~29~~50 V ion deflector into the laminar sampling flow of the inlet via a 5 mm orifice. A small
145 counter flow (~40 mlpm) was applied through the orifice to prevent any mixing of the electrically neutral reagent
146 vapor with the sampling flow. The details of the inlet design, setup, and operation are described ~~in~~ previously (He,
147 2017; Rissanen et al., 2019).

148 149 **2.3 Br-FIGAERO-CIMS**

150 We also measured both the gas- and particle-phase compositions via thermal desorption using a bromide chemical
151 ionization time-of-flight mass spectrometer equipped with a Filter Inlet for Gases and AEROSOLS (Br-FIGAERO-
152 CIMS) (Lopez-Hilfiker et al., 2014). FIGAERO is a manifold inlet for a CIMS with two operating modes. In the
153 sampling mode, gases are directly sampled into a 150 mbar ion-molecule reactor, using coaxial core sampling to
154 minimize their wall losses in the sampling line. The total flow is maintained at 18.0 slpm and the core flow at 4.5 slpm;
155 the CIMS samples at the center of the core flow with a flow rate at ~1.6 slpm. Concurrently, particles are collected on

156 a PTFE filter via a separate dedicated port with a flow rate of 6 slpm. In the desorption mode, the filter is automatically
157 moved into a pure N₂ gas stream flowing into the ion molecule reactor, while the N₂ is progressively heated upstream
158 of the filter to evaporate the particles via temperature programmed desorption. Analytes are then chemically ionized
159 by Br⁻ and extracted into a mass spectrometer.

160
161 We optimize the adduct-ion signals in both the Br-MION-CIMS and Br-FIGAERO-CIMS by tuning the electric field
162 strengths in the first two low-pressure stages of the mass spectrometer as weak as possible to minimize collision
163 induced cluster fragmentation while maintaining sufficient ion transmission. Optimization is achieved by maximizing
164 the ratio of I₂Br⁻/Br⁻ at a constant I₂ concentration. We list relevant instrument specifications and operational
165 conditions in Table S1. It should be noted that these values are specific to our instruments, thus can vary according to
166 instrument parameters and may not be applicable to other instruments.

167 168 **2.4 CE-DOAS**

169 For the quantitative measurement of gas-phase molecular iodine (I₂), we deployed a Cavity Enhanced Differential
170 Optical Absorption Spectroscopy instrument (CE-DOAS) (Meinen et al., 2010). CE-DOAS determines concentrations
171 of trace gases from the strength of differential spectral features in a reference spectrum. The overall accuracy for the
172 I₂ time series is estimated to be 20 %, never better than the detection limit (3-sigma precision), resulting from the
173 uncertainty in cross sections and the stability of the baseline.~~The accuracy of the method is ultimately determined by~~
174 ~~the uncertainty of the respective absorption cross sections.~~ It is thus an absolute method and does not depend on an
175 instrument specific detection efficiency. To maximize the measurement sensitivity towards I₂, we used a setup
176 optimized for the green wavelength range (508-554 nm), where I₂ exhibits strong differential absorption features. The
177 measurement light is provided by a green light emitting diode (LED Engin). Spectral dispersion is established with a
178 Czerny-Turner grating spectrometer (Princeton Instruments Acton 150), resulting in an optical resolution of 0.73 nm
179 full width at half maximum at 546 nm. Intensities are monitored with a CCD detector (Princeton Instruments
180 PIXIS400B) cooled to -70 °C. Highly reflective mirrors (Advanced Thin Films) enhance the 1 m mirror separation to
181 an effective optical path length of 15-23 km. The effective spectral mirror reflectivity was established by comparing
182 light intensity spectra in the presence of N₂ and He (Washenfelder et al., 2008). The abundance of trace gases is then
183 determined by comparing spectra of chamber air relative to reference spectra recorded with ultrapure N₂ without I₂.
184 Chamber air is drawn into the cavity with a constant flow rate of 1 slpm. Variations of the sampling flow did not result
185 in changes in measured I₂ concentrations, indicating that photolysis from the measurement light within the instrument
186 was negligible. The following absorbers were included in the fit: I₂ (Spietz et al., 2006), NO₂ (Vandaele et al., 1998),
187 H₂O (Rothman et al., 2010), O₂-O₂ collision-induced absorption (Thalman and Volkamer, 2013), and a polynomial of
188 sixth order. The setup allowed a 1-minute detection limit of 25 pptv, or 8 pptv for integration times of 10 minutes,
189 respectively. Periodic automated recordings of N₂ reference spectra were recorded to ensure baseline stability. The
190 optical path length at the time of measurement was continuously confirmed for consistency by the measurement of
191 the O₂-O₂ collision-induced absorption and H₂O column in the same analysis window. The overall systematic accuracy

192 for the I₂ time series is estimated to be 20 %, never better than the detection limit, resulting from the uncertainty in
193 cross sections and the stability of the baseline.

195 **2.5 Offline calibration setup**

196 **2.5.1 I₂ permeation device**

197 We used an iodine permeation tube (VICI Metronic) as a source for offline laboratory calibration. The permeation
198 tube was encased within an electronically controlled heating mantle (80 ~ 140 (± 2) °C) to allow for adjustable yet
199 steady iodine permeation rates. The heating device (for holding the permeation tube) was made from a stainless steel
200 tube (½-inch outer diameter, OD) with a length of 25 cm, encased within an electronically controlled heating mantle.
201 The configuration of the permeation device has been described in Tham et al. (Tham et al., 2021). The iodine
202 permeation device was run continuously for at least 72 hours before any calibration experiments to ensure that a
203 complete equilibrium was reached in the system. We then confirmed the robustness of the permeation device by the
204 constant I₂ signal measured with Br-MION-CIMS for over 24 hours.

205
206 To determine the permeation rate of I₂, we trapped iodine in n-hexane at cryogenic temperatures in an all-glass
207 apparatus, following the method described in Chance et al. (Chance et al., 2010). We initially filled the absorption
208 glass vessel with 20 ml n-hexane (99.95 %, Merck), and then weighed it to determine the combined mass. We then
209 immersed the absorption vessel into a wide-necked Dewar vessel, filled with an acetone/dry ice mixture (at -80 ±
210 3 °C). After temperature equilibration, the I₂ molecules, carried by 50 mlpm N₂ flow from the permeation device, were
211 bubbled through the absorption vessel. After a continuous collection for 5 hours, we removed the absorption apparatus
212 from the cooling mixture, and allowed it to warm to room temperature prior to disassembling the setup to prevent any
213 losses of iodine on the tip of the inlet capillary. The absorption vessel was then re-weighed; the mass compared with
214 that prior to absorption was less than 2 %, indicating a negligible loss during the trapping process. The I₂ / n-hexane
215 sample solutions were stored at 4 °C for 14 hours before being subjected to analysis.

216
217 We determined the I₂ concentration of the samples using a UV/Vis spectrophotometer (Shimadzu Model UV2450) at
218 a wavelength of 522 nm. We established a calibration curve via a set of I₂ solutions ranging from 270 to 5300 nmol,
219 diluted with n-hexane from a freshly prepared stock solution (0.5 g L⁻¹). Repetition of the same analysis after 2 and 7
220 days yielded identical results, confirming that the sample solutions were stable at 4 °C. As an alternative analytical
221 approach, we also quantified the I₂ concentration in the sample solutions using an inductively coupled plasma mass
222 spectrometer (ICP-MS, Agilent 7800). Before introducing to the ICP-MS, the sample solutions were treated with
223 NaHSO₃ water solution (0.100 M), accomplishing efficient hexane-to-water extraction and simultaneous reduction of
224 iodine to iodide (Schwehr et al., 2005) (Agilent Clinical Sample Preparation Guide (v3), ref. ISO 17294-2). The ICP-
225 MS results were in good agreement (within 20 % discrepancies) with those from the UV/Vis spectrophotometry.

226
227 We conducted the I₂ trapping and quantification experiments in triplicate with satisfactory reproducibility (standard
228 deviation < 10 %). The calculated iodine permeation rate at 50 mlpm N₂ flow and 140 °C oven temperature is 278 ±

229 12 ng min⁻¹ (mean ± standard deviation). This result was used as the benchmark to estimate temperature-corrected
230 permeation rates according to the formula provided by the permeation tube vendor (VICI Metronic). We checked the
231 validity of the temperature-corrected values by conducting a second iodine absorption experiment in which the iodine
232 permeation tube was kept at 130 °C with an N₂ flow rate of 50 mlpm, and the determined permeation rate agreed
233 within 10 % of the calculated value.

234 **2.5.2 Cl₂ permeation device**

235 We used a commercial chlorine permeation tube (VICI Metronic) as a source for offline calibration. We passed a 20
236 mlpm high-purity nitrogen (99.999 %) flow at room temperature through a 25 cm long stainless-steel tube (½" O.D.)
237 containing the permeation tube. We quantified the permeation rate of Cl₂ following a procedure described in a previous
238 study (Finley and Saltzman, 2008). The output of 20 mlpm flow was bubbled into a buffered aqueous potassium iodide
239 solution (2.0 % KI (m/v), prepared in 1.00 mM aqueous phosphate buffer, pH = 7.0) filled in an all-glass two-stage
240 serial absorption apparatus (stage 1 = 100 ml; stage 2 = 50 ml) for 3 hours and kept at room temperature. The Cl₂
241 oxidized the iodide (I⁻) into iodine (I₂) once contacting with the KI absorption solution, and the I₂ further reacted with
242 the excess KI present in the absorption solution to form I₃⁻, which can be quantified by UV/Vis-spectrophotometry.
243 We analyzed the resulting sample solutions with an UV/Vis spectrophotometer (Shimadzu Model UV-1800) using 1-
244 cm quartz cells at 352 nm, corresponding to the I₃⁻. We detected no I₃⁻ in the second stage absorption solution,
245 indicating that all the chlorine was quantitatively trapped and rapidly converted to I₃⁻ within the first absorption unit.
246 The samples were quantified relative to I₃⁻ standards in the range of 5 to 68 × 10⁻⁶ M, prepared by dilution of a stock
247 obtained by dissolving 174 mg iodine in 200 ml of a solution containing 2 % KI in 1.00 mM aqueous phosphate buffer,
248 pH 7.0. From this calibration curve, we calculated a molar absorptivity of 26,800 L mol⁻¹ cm⁻¹, which is consistent
249 with the values reported in the literature (Finley and Saltzman, 2008; Kazantseva et al., 2002). Samples and standard
250 solutions were re-analyzed after being stored in the dark at room temperature for 24 hours, and the results were within
251 3 % of those obtained with the fresh solutions. We repeated the absorption experiment, and the calculated chlorine
252 permeation rate at room temperature was 764 ± 74 ng Cl₂ min⁻¹ (mean ± standard deviation).

253 **2.5.3 HOI calibrator**

254 We produced a continuous HOI source via the reaction of I₂ and hydroxyl radicals (OH) in a setup similar to the
255 sulfuric acid (H₂SO₄) calibrator (Kürten et al., 2012). The OH was generated by photolyzing H₂O with a mercury (Hg)
256 lamp at 184.9 nm, whose calibrated intensity was used to estimate the OH concentration. We tested the system by
257 removing the I₂ or OH source from the calibrator, upon which HOI production was undetectable, confirming that any
258 single reactant did not produce HOI. A numerical model was constructed to predict the mean HOI concentration
259 entering the Br-MION-CIMS, which is analogous to the model used for H₂SO₄ calibration (Kürten et al., 2012). We
260 only included the formation pathway of I₂ + OH to HOI in the model for simplicity; the other pathway of IO + HO₂
261 was considered minor as its reaction rate is about an order of magnitude slower than that of I₂ + OH. Furthermore, IO
262 is likely at negligible concentration in the calibration device due to the absence of O₃ for IO formation.

266 **2.65 Quantum chemical calculations**

267 We used quantum chemical calculations to estimate the cluster formation enthalpy of halogen containing species and
268 bromide ions. The initial ~~conformational enformer~~ sampling was performed using the Spartan '14 program. The
269 cluster geometry was then optimized using density function theory methods at the ω B97X-D/aug-cc-pVTZ-PP level
270 of theory (Chai and Head-Gordon, 2008; Kendall et al., 1992). Iodine pseudopotential definitions were taken from the
271 EMSL basis set library (Feller, 1996). Calculations were carried out using the Gaussian 09 program (Frisch et al.,
272 2010). An additional coupled-cluster single-point energy correction was carried out on the lowest energy geometry to
273 calculate the final cluster formation enthalpy. The coupled-cluster calculation was performed at the DLPNO-
274 CCSD(T)/def2-QZVPP level using the ORCA program ver. 4.0.0.2 (Neese, 2012; Riplinger and Neese, 2013). In
275 Table 1 we present calculated cluster formation enthalpies based on the optimized geometries.

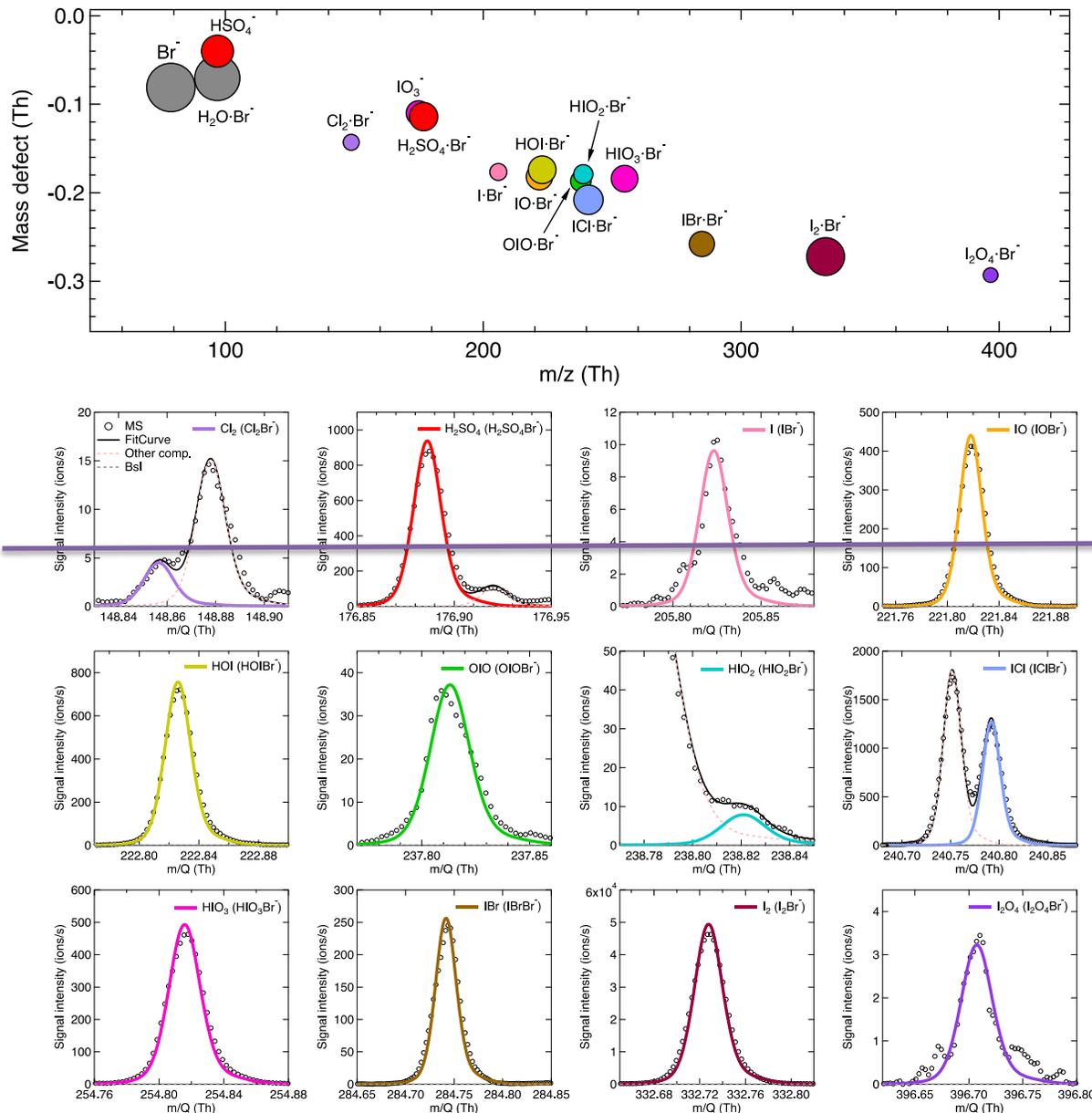
276

277 **3 Results and Discussion**

278 **3.1 Detection of gas-phase inorganic species by Br-MION-CIMS**

279 We show in Fig. 1 the selected inorganic species observed with the Br-MION-CIMS during an iodine ~~new particle~~
280 ~~formation oxidation~~ experiment in the CLOUD chamber. The peak identities are indicated in the labels. Observed
281 species include I_2 and its various oxidation products. There are also a few other halogen-containing inorganic species
282 such as Cl_2 , ICl and IBr , likely coming from the impurities in the iodine source. Non-halogen inorganic species such
283 as H_2SO_4 can also react with bromide ions and are detected. Due to the large negative mass defect of the bromine and
284 iodine atoms, and the high resolution (~ 10000 Th Th^{-1}) of the mass spectrometer, the peaks can be unambiguously
285 distinguished and identified in the mass spectrum. As shown in the lower panel of Fig. 1, most of the iodine-containing
286 species appear as a single peak in the unit mass range, except for $HIO_2 \cdot ^{79}Br^-$ ($m/z = 238.82$), which overlaps with the
287 reagent ion cluster ($^{79}Br_2^{81}Br^-$) ($m/z = 238.75$).

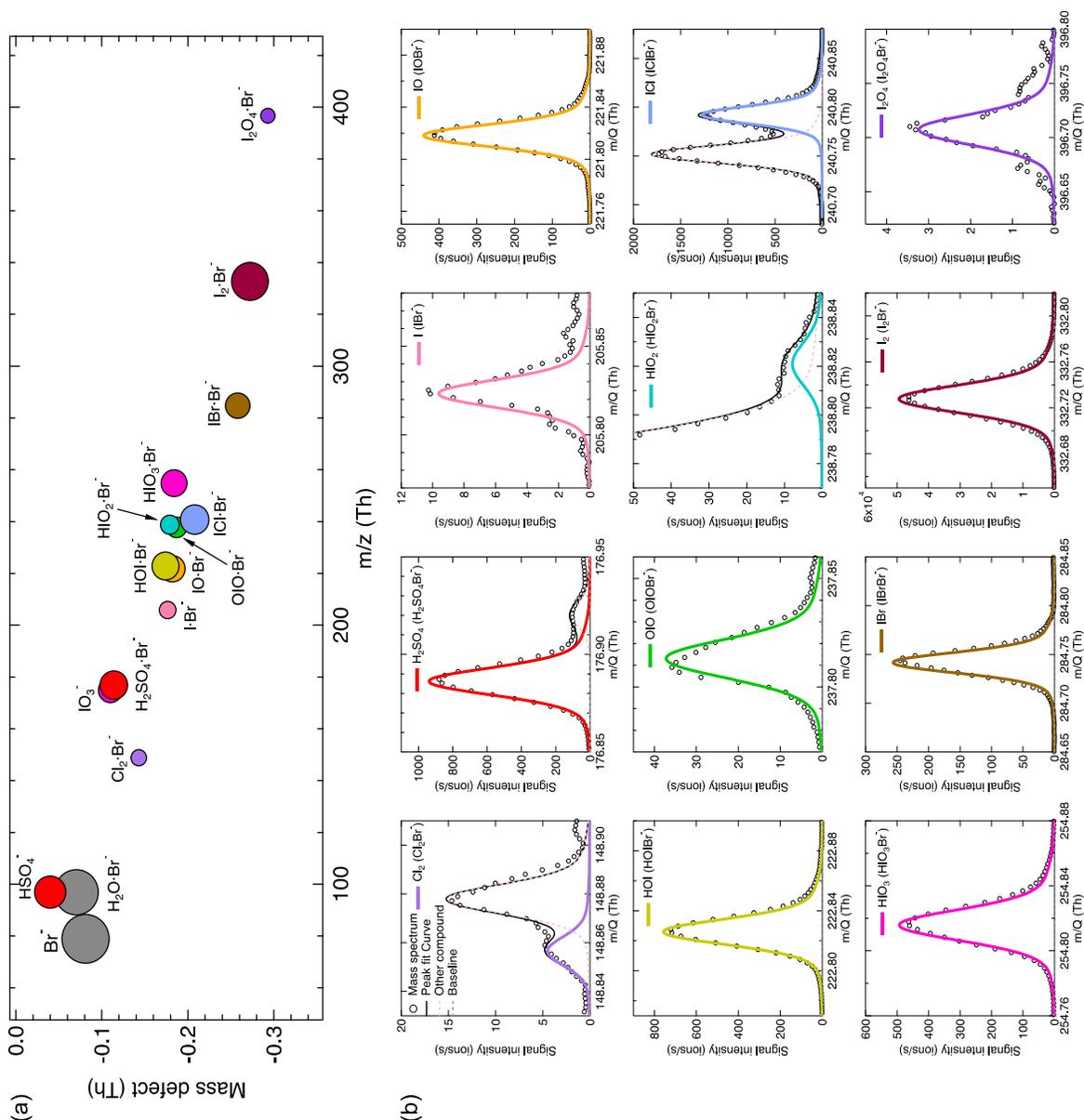
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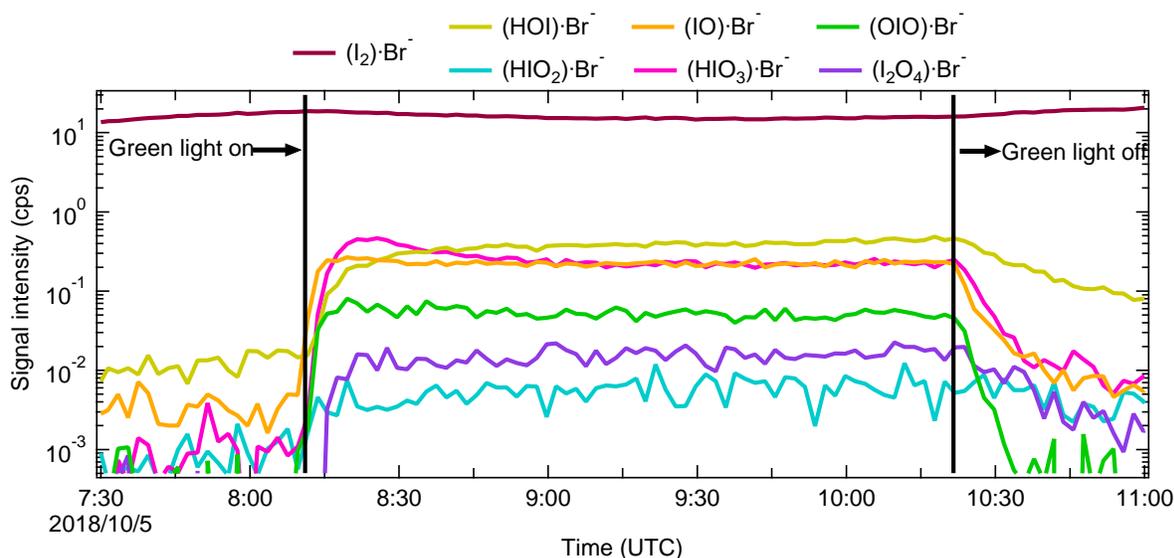
289
 290 **Figure 1.** Gas phase inorganic species measured with the Br MION-CIMS. Mass defect (difference of exact mass to integer mass)
 291 versus m/z of gas phase halogen species and sulfuric acid during an iodine new particle formation experiment at 69 % relative
 292 humidity and -10 °C. Ions shown here are either clustered with or formed via proton transfer to a bromide ion. The area of the
 293 markers is proportional to the logarithm of the signal (counts per second). Shown in the lower panel are the high resolution single
 294 peak fits for species in the mass defect plot in the upper panel.
 295

296 The iodine oxidation new particle formation experiments were conducted under experimental conditions typically
 297 found in the high-latitude marine boundary layer, with a temperature of -10 °C and a relative humidity of 69 %. As
 298 illustrated in Fig. 2, a typical experiment started with illumination of the chamber at constant I_2 (~60 pptv) using the
 299 green light to photolytically produce I atoms. The subsequent reactions of I and ~40 ppbv O_3 led to the formation of
 300 various oxidized iodine species within a few minutes. The most prominent species we observed from these
 301 experiments were IO, HOI and HIO_3 , with lower but significant levels of OIO, HIO_2 , and I_2O_4 . Among these iodine

302 oxides, IO rose the most rapidly; this is consistent with the first-generation production of IO from the I + O₃ reaction.
 303 After a few steps of radical reactions, OIO, HIO₂ and HIO₃ reached steady state almost simultaneously. The only
 304 observed iodine oxide dimer was I₂O₄ in this event, while I₂O₂, I₂O₃ and I₂O₅ were below the detection limit of both
 305 mass spectrometers. A noticeable dip in the HIO₃ traces a few minutes after the onset of the reactions is likely due to
 306 the participation of HIO₃ in new-particle formation, resulting in an extra loss term and a lower steady-state
 307 concentration. When we turned off the green light, the production of I radicals stopped and iodine species decayed
 308 away.



309
 310 **Figure 1.** Gas-phase inorganic species measured with the Br-MION-CIMS. (a) Mass defect (difference of exact mass to integer
 311 mass) versus m/z of gas-phase halogen species and sulfuric acid during an iodine oxidation experiment at 69 % relative humidity
 312 and -10 °C (conducted on 05 October 2018). Ions shown here are either clustered with or formed via proton transfer to a bromide
 313 ion. The area of the markers is proportional to the logarithm of the signal (counts per second). (b) The high-resolution single peak
 314 fits for species in the mass defect plot in the upper panel. Species in both panels are color-coded in the same style.



316
 317 **Figure 2.** Evolution of selected iodine species during a typical run. The experiment was performed at 60 pptv I₂, 40 ppbv O₃, 69 %
 318 relative humidity and -10 °C. The oxidized iodine species start to appear soon after switching on the green light at 08:11, 05 October
 319 2018. The I atom production was halted at 10:21, 05 October 2018 by switching off the green light, and the concentration of
 320 oxidized iodine species decayed away afterwards. All species are color-coded in the same way as in Fig. 1.
 321

322 3.2 Relative humidity dependence

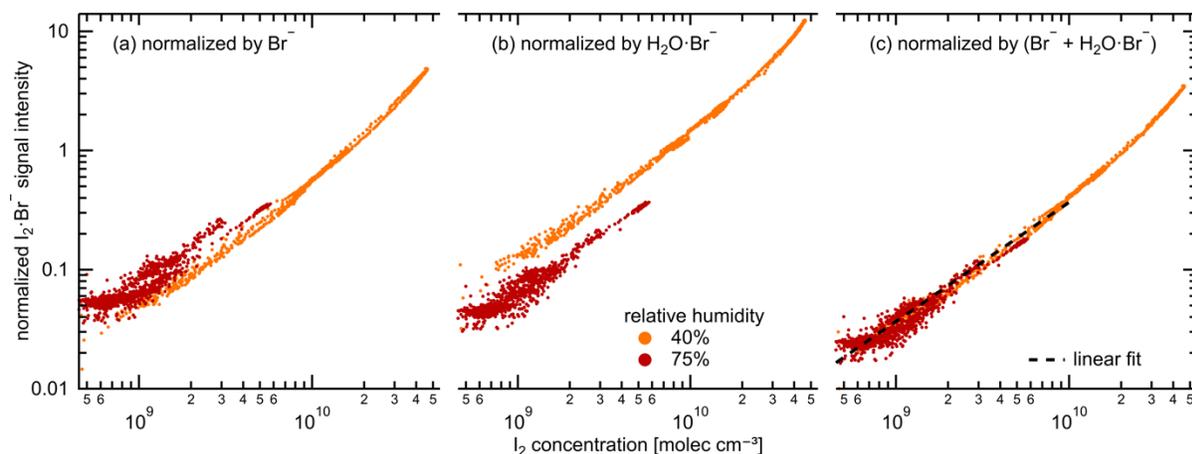
323 Water molecules can cluster with I⁻ to form H₂O·I⁻ in the iodide CIMS. This enhances the instrument sensitivities for
 324 small molecules (i.e. chlorine and bromine inorganics) and reduces them for large molecules (i.e. oxygenated organics)
 325 (Lee et al., 2014). To investigate the role of water concentration in the sensitivity of the Br-MION-CIMS, we varied
 326 the relative humidity (RH) from 40 % to 80 % at a constant temperature of -10 °C. We show in Fig. 3 the correlation
 327 of I₂ time series from the Br-MION-CIMS and the CE-DOAS throughout the experiment.
 328

329 Chemical ionization relies on an ion-molecule reaction to transfer charge from a reagent ion to an analyte, forming
 330 either a product ion or a charged cluster between the analyte and the reagent ion with a rate coefficient k_{IM} . This occurs
 331 in an ion-molecule reactor, with a fixed flow rate and thus reaction time, dt , and ideally under pseudo first-order
 332 conditions where a small fraction of the analyte is ionized and the reagent ion concentration ($[Ion]$) remains constant.
 333 Under these (linear) conditions the fraction of analyte that is ionized is $k_{IM} \times [Ion] \times dt$. However, the primary ion
 334 source strength can vary with time, and so we normalize the analyte signal by reagent ion signal to account for those
 335 small variations in analyte signal.
 336

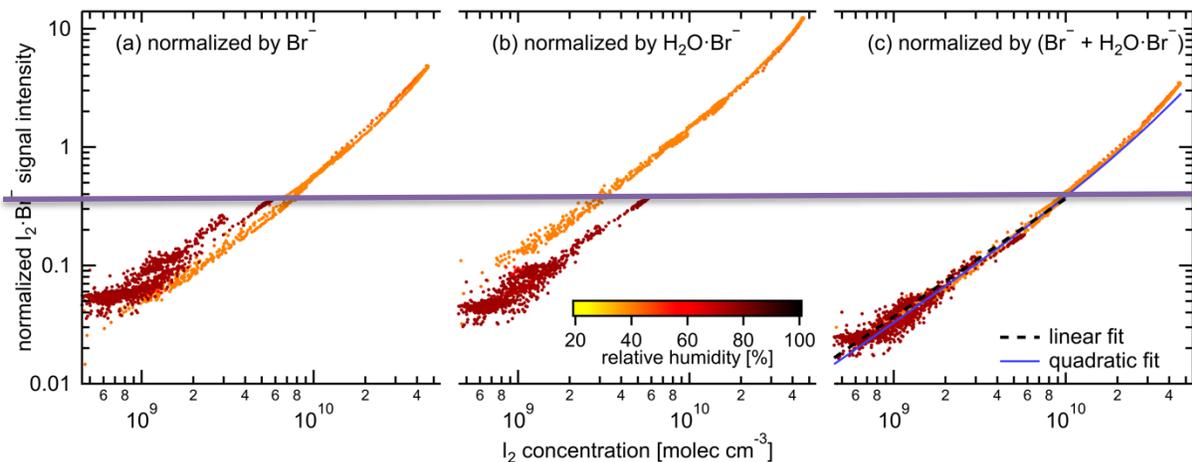
337 During the RH transition, the ratio of the two reagent ions, Br⁻ and H₂O·Br⁻, changed in the Br-MION-CIMS. As
 338 shown in Fig. 3 (a) and (b), using either reagent ion alone for I₂ normalization results in discrepancies in recovered I₂
 339 concentrations at different RH. However, if we use the sum of these two reagent ions (Br⁻ + H₂O·Br⁻) for normalization,
 340 the humidity effect vanishes, as shown in ~~the~~ Fig. 3 (c). Separately, during the I₂ calibration using a permeation tube
 341 (section 3.3.2.1), This suggests that the quantitative detection of I₂ molecules is robust and independent of RH between
 342 20 % and 40 % at 25 °C, as long as a proper normalization method is used for the Br-MION-CIMS. Furthermore, we

343 have also carried out the HOI calibration at 25 °C (section 3.3.2.3) and used the same normalization method as
 344 described in section 3.3.4. During the calibration, we varied the water content in the calibrator to vary OH
 345 concentrations. A good correlation between the modeled HOI concentrations and the measured HOI signals also
 346 indicates that the different H₂O concentrations in the system do not affect the HOI detection (Fig.4(d)). This assertion
 347 may also be applicable to other molecules such as iodine monochloride (ICl) and iodine monobromide (IBr), which
 348 share similar chemical and physical properties with I₂; -but however, further confirmation is needed for other species
 349 such as oxygenated organics.

350



351



352

353 **Figure 3.** Signal normalization methods for the Br-MION-CIMS. Normalized I₂·Br⁻ signal intensity for variable I₂ concentrations,
 354 color coded by relative humidity (orange: 35-45 % relative humidity, red: 70-80 % relative humidity). The charge ions in the ion
 355 source of Br-MION-CIMS are Br⁻ and H₂O·Br⁻ (both ⁷⁹Br and ⁸¹Br). Their abundance depends both on the instrument tuning and
 356 the absolute humidity of the sampled flow. The normalization of the I₂·Br⁻ signal by only Br⁻ (a) or H₂O·Br⁻ (b) does not compensate
 357 for the humidity effect. Using the sum of Br⁻ and H₂O·Br⁻ (c) for normalization yields a tight correlation to the true I₂ as measured
 358 by CE-DOAS, independent of the humidity. The black dashed line and blue solid curve indicates the fitted linear and quadratic
 359 calibration curves, respectively.

360

361 3.3 Quantification of gas-phase inorganic species

362 Inter-method calibrations and offline calibrations were carried out to quantify the sensitivities of the Br-MION-CIMS
 363 to selected calibrants. For the inter-method calibrations, we used the CLOUD chamber as a stable source of I₂ and

364 H_2SO_4 , and inter-compared the signals in the Br-MION-CIMS with the CE-DOAS and a pre-calibrated nitrate-CIMS,
365 respectively. For the offline calibrations, we carried out the experiments separately in a laboratory at the University
366 of Helsinki, using permeation tubes to quantify I_2 and Cl_2 , and a calibrator to quantify HOI.

367 **3.3.1 Inter-method calibrations at CLOUD**

368 **3.3.1.1 I_2 calibration using the CE-DOAS**

369 As shown in Fig. 3, we use the accurate I_2 time series measured with the CE-DOAS to calibrate normalized I_2 signals
370 in the Br-MION-CIMS. The I_2 concentrations used for the calibration span approximately 2 orders of magnitude,
371 reaching up to 4.6×10^{10} molec cm^{-3} . A linear fit, limited to I_2 concentrations smaller than 10^{10} molec cm^{-3} , establishes
372 the calibration factor as follows:

$$373 \quad [\text{I}_2] = 2.7 \times 10^{10} \text{ molec cm}^{-3} \times \text{I}_2 \cdot {}^{79}\text{Br}^- / ({}^{79}\text{Br}^- + \text{H}_2\text{O} \cdot {}^{79}\text{Br}^-)$$

374 For this range of concentrations, which are typically encountered in the atmosphere, the calibrated Br-MION-CIMS
375 time series agrees within error with the CE-DOAS measurement (1-sigma accuracy 20 %, detection limit 25 pptv for
376 1 min data). Deviations between both time series are generally smaller than 10 % (25 and 75 percentile 0.88 and 1.03,
377 respectively). These small differences may be explained by incomplete homogeneity of iodine concentrations in the
378 chamber and the different sampling positions of CE-DOAS and Br-MION-CIMS. As shown in Fig. 3, we use the
379 accurate I_2 time series measured with the CE-DOAS to calibrate normalized I_2 signals in the Br-MION-CIMS. The I_2
380 concentrations used for the calibration span approximately 2 orders of magnitude, reaching up to 4.6×10^{10} molec cm^{-3} .
381 A linear fit, limited to I_2 concentrations smaller than 10^{10} molec cm^{-3} , establishes the calibration factor $[\text{I}_2] = 2.7 \times 10^{10}$
382 molec $\text{cm}^{-3} \times \text{I}_2 \cdot {}^{79}\text{Br}^- / ({}^{79}\text{Br}^- + \text{H}_2\text{O} \cdot {}^{79}\text{Br}^-)$. We also use a quadratic fit to establish the calibration for the entire range of
383 concentrations encountered during whole campaign (solid line in Fig. 3 (c)); two curves agree well. The CE-DOAS I_2
384 detection limit is 6.3×10^8 molec cm^{-3} (25 pptv) for a 1 min integration time, and the total systematic uncertainty is
385 estimated to be 20%. Deviations between both time series are generally smaller than 10% (25 and 75 percentile 0.88
386 and 1.03, respectively). These small differences are consistent with incomplete homogeneity of iodine concentrations
387 in the chamber and the different sampling positions of CE-DOAS and Br-MION-CIMS.

388 **3.3.1.2 H_2SO_4 calibration using a nitrate-CIMS**

389 We derive the H_2SO_4 calibration coefficient for the Br-MION-CIMS using the absolute H_2SO_4 concentrations
390 measured with a pre-calibrated nitrate-CIMS. The calibration protocol of H_2SO_4 in the nitrate-CIMS has been
391 described in detail previously (Kürten et al., 2012). The H_2SO_4 time series used for the inter-method calibration covers
392 a wide concentration range from less than 5.0×10^4 (detection limit of the nitrate-CIMS) to 6.0×10^7 molec cm^{-3} . For
393 Br-MION-CIMS, although both HSO_4^- and $\text{H}_2\text{SO}_4 \cdot \text{Br}^-$ appear as distinct peaks for sulfuric acid, we only use the
394 normalized $\text{H}_2\text{SO}_4 \cdot {}^{79}\text{Br}^-$ for the inter-calibration, as HSO_4^- ($m/z = 96.96$) has substantial interference from the reagent
395 ion $\text{H}_2\text{O} \cdot {}^{79}\text{Br}^-$ ($m/z = 96.93$). We show in Fig. 4 (a) the linear fit:

$$396 \quad [\text{H}_2\text{SO}_4] = 4.1 \times 10^{10} \text{ molec cm}^{-3} \times \text{H}_2\text{SO}_4 \cdot {}^{79}\text{Br}^- / ({}^{79}\text{Br}^- + \text{H}_2\text{O} \cdot {}^{79}\text{Br}^-) - 9.3 \times 10^5$$

399 The H₂SO₄ calibration coefficient is thus 4.1×10¹⁰ molec cm⁻³ per normalized signal (cps cps⁻¹), and the correlation
400 coefficient between the two H₂SO₄ traces is 0.95. The systematic 3-sigma accuracy is +50/-33 % for H₂SO₄ calibration
401 using a nitrate-CIMS; detailed accuracy estimation has been discussed previously (Stolzenburg et al., 2020).
402

403 **3.3.2 Offline laboratory calibrations for I₂, Cl₂, and HOI**

404 **I₂ calibration using a permeation tube**

405 We used an iodine permeation tube (VICI Metronic) as a source for offline calibration. The permeation tube was
406 encased within an electronically controlled heating mantle (80–140 (± 2) °C) to allow for adjustable yet steady iodine
407 permeation rates. The iodine permeation device was run continuously for at least 72 hours before any calibration
408 experiments to ensure that a complete equilibrium was reached in the system. We then confirmed the robustness of
409 the permeation device by the constant I₂ signal measured with Br-MION-CIMS for over 24 hours.

410
411 To determine the permeation rate of I₂, we trapped iodine in n-hexane at cryogenic temperatures in an all-glass
412 apparatus, following the method described in Chance et al. (Chance et al., 2010). We initially filled the absorption
413 glass vessel with 20 ml n-hexane (99.95%, Merck), and then weighed it to determine the combined mass. We then
414 immersed the absorption vessel into a wide-necked Dewar vessel, filled with an acetone/dry ice mixture (at -80 ±
415 3 °C). After temperature equilibration, the I₂ molecules, carried by 50 mlpm N₂ flow from the permeation device, were
416 bubbled through the absorption vessel. After a continuous collection for 5 hours, we removed the absorption apparatus
417 from the cooling mixture, and allowed it to warm to room temperature prior to disassembling the setup to prevent any
418 losses of iodine on the tip of the inlet capillary. The absorption vessel was then re-weighed; the mass compared with
419 that prior to absorption was less than 2%, indicating a negligible loss during the trapping process. The I₂/n-hexane
420 sample solutions were stored at 4 °C for 14 hours before being subjected to analysis.

421
422 We determined the I₂ concentration of the samples using a UV/Vis spectrophotometer (Shimadzu Model UV2450) at
423 a wavelength of 522 nm. We established a calibration curve via a set of I₂ solutions ranging from 270 to 5300 nmol,
424 diluted with n-hexane from a freshly prepared stock solution (0.5 g L⁻¹). Repetition of the same analysis after 2 and 7
425 days yielded identical results, confirming that the sample solutions were stable at 4 °C. As an alternative analytical
426 approach, we also quantified the I₂ concentration in the sample solutions using an inductively coupled plasma mass
427 spectrometer (ICP-MS, Agilent 7800). Before introducing to the ICP-MS, the sample solutions were treated with
428 NaHSO₃-water solution (0.100 M), accomplishing efficient hexane-to-water extraction and simultaneous reduction of
429 iodine to iodide (Schwehr et al., 2005) (Agilent Clinical Sample Preparation Guide (v3), ref. ISO-17294-2). The ICP-
430 MS results were in good agreement with those from the UV/Vis spectrophotometry.

431
432 We conducted the I₂-trapping and quantification experiments in triplicate with satisfactory reproducibility (standard
433 deviation < 10%). The calculated iodine permeation rate at 50 mlpm N₂ flow and 140 °C oven temperature is 278 ±
434 12 ng min⁻¹ (mean ± standard deviation). This result was used as the benchmark to estimate temperature-corrected
435 permeation rates according to the formula provided by the permeation tube vendor (VICI Metronic). We checked the

436 validity of the temperature-corrected values by conducting a second iodine absorption experiment in which the iodine
437 permeation tube was kept at 130 °C with an N₂ flow rate of 50 mlpm, and the determined permeation rate agreed
438 within 10% of the calculated value. ~~We~~For the I₂ calibration, ~~we then~~ diluted the I₂ flow to seven different values and
439 measured the flow with the Br-MION-CIMS. We repeated the calibration five times; we show the data along with a
440 linear fit between the I₂ concentration and normalized I₂ signal in Fig. 4 (ab). The slope of the line gives a calibration
441 coefficient of 6.3×10¹⁰ molec cm⁻³ per normalized signal (cps cps⁻¹), with R² of 0.98 and an overall 1-sigma accuracy
442 uncertainty of ± 45%.

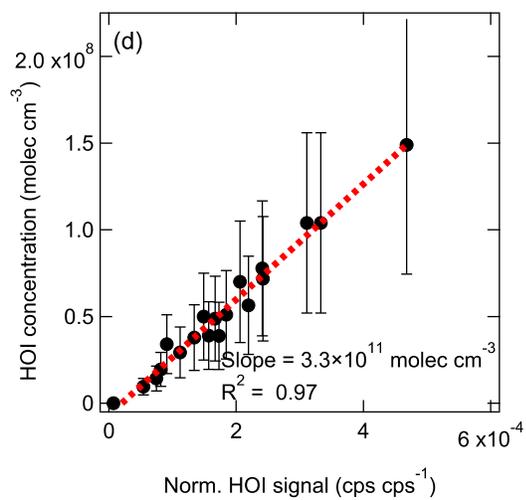
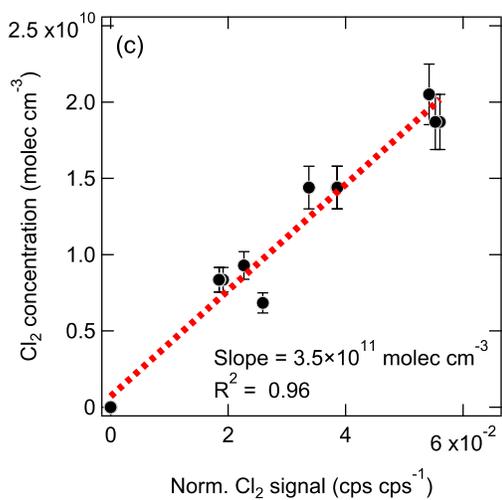
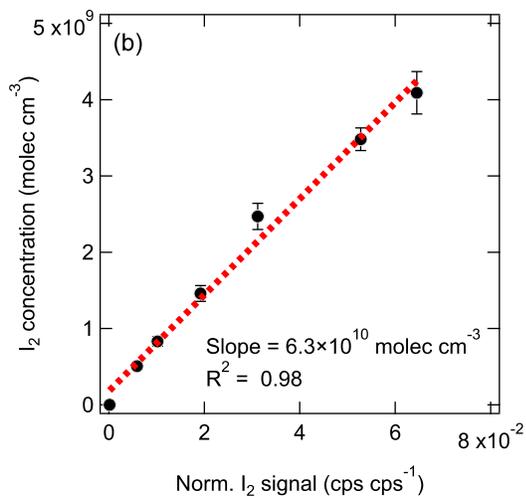
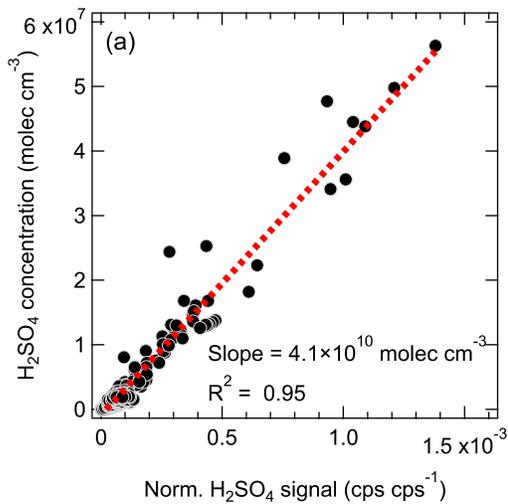
444 3.3.3 Cl₂ calibration using a permeation tube

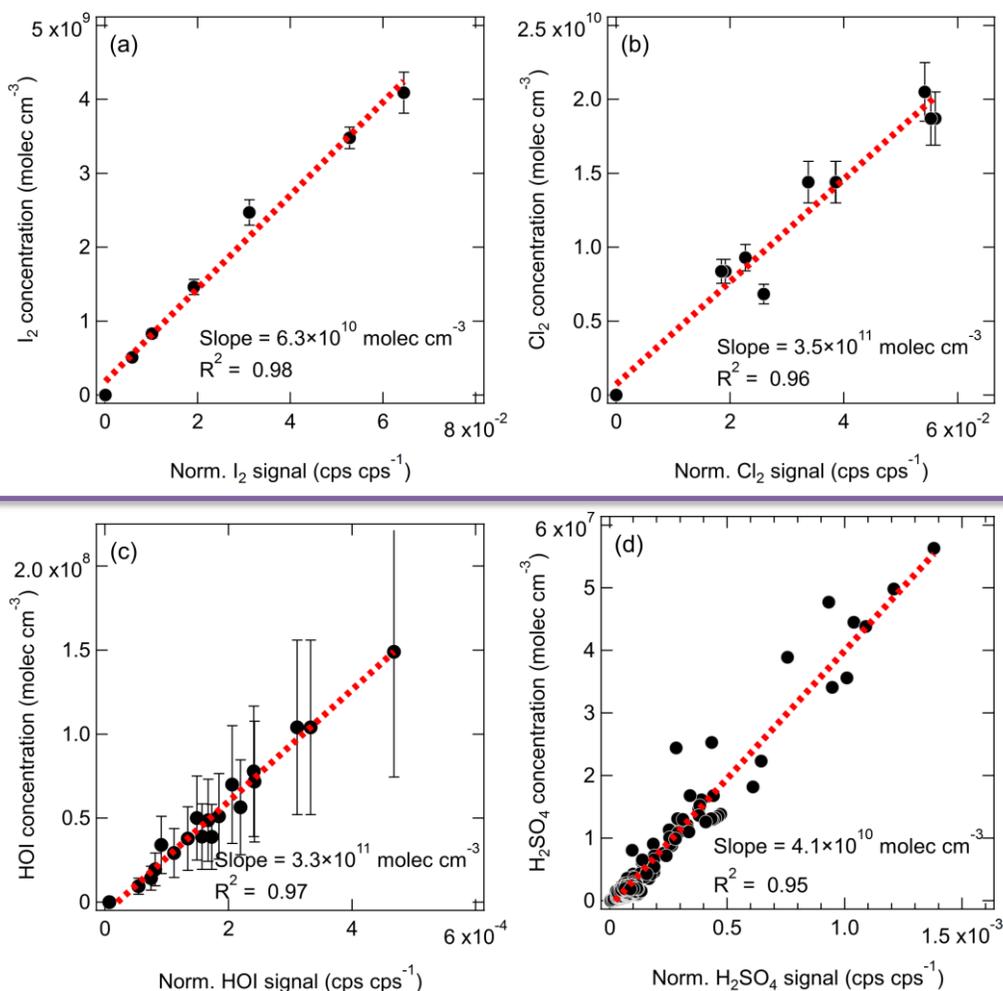
445 ~~We used a commercial chlorine permeation tube (VICI Metronic) as a source for offline calibration. We passed a 20~~
446 ~~mlpm high-purity nitrogen (99.999%) flow at room temperature through a 25-cm long stainless-steel tube (1/2" O.D.)~~
447 ~~containing the permeation tube. We quantified the permeation rate of Cl₂ following a procedure described in a previous~~
448 ~~study (Finley and Saltzman, 2008). The output of 20 mlpm flow was bubbled into a buffered aqueous potassium iodide~~
449 ~~solution (2.0 % KI (m/v), prepared in 1.00 mM aqueous phosphate buffer, pH = 7.0) filled in an all-glass two-stage~~
450 ~~serial absorption apparatus (stage 1 = 100 ml; stage 2 = 50 ml) for 3 hours and kept at room temperature. The Cl₂~~
451 ~~oxidized the iodide (I⁻) into iodine (I₂) once contacting with the KI absorption solution, and the I₂ further reacted with~~
452 ~~the excess KI present in the absorption solution to form I₃⁻, which can be quantified by UV/Vis spectrophotometry.~~
453 ~~We analyzed the resulting sample solutions with an UV/Vis spectrophotometer (Shimadzu Model UV-1800) using 1-~~
454 ~~cm quartz cells at 352 nm, corresponding to the I₃⁻. We detected no I₃⁻ in the second-stage absorption solution,~~
455 ~~indicating that all the chlorine was quantitatively trapped and rapidly converted to I₃⁻ within the first absorption unit.~~
456 ~~The samples were quantified relative to I₃⁻ standards in the range of 5 to 68 × 10⁻⁶ M, prepared by dilution of a stock~~
457 ~~obtained by dissolving 174 mg iodine in 200 ml of a solution containing 2 % KI in 1.00 mM aqueous phosphate buffer,~~
458 ~~pH 7.0. From this calibration curve, we calculated a molar absorptivity of 26,800 L mol⁻¹ cm⁻¹, which is consistent~~
459 ~~with the values reported in the literature (Finley and Saltzman, 2008; Kazantseva et al., 2002). Samples and standard~~
460 ~~solutions were re-analyzed after being stored in the dark at room temperature for 24 hours, and the results were within~~
461 ~~3 % of those obtained with the fresh solutions. We repeated the absorption experiment, and the calculated chlorine~~
462 ~~permeation rate at room temperature was 764 ± 74 ng Cl₂ min⁻¹ (mean ± standard deviation).~~

464 ~~The~~For the Cl₂ calibration, ~~the~~ Cl₂ permeation source was run continuously for 12 hours prior to calibration
465 experiments to ensure complete system equilibrium. A two-stage dilution system similar to the setup of (Gallagher et
466 al., 1997) was set up for diluting the output of the Cl₂ permeation device. The 20 mlpm of N₂ stream emerging from
467 the Cl₂ permeation device (operated at room temperature) was diluted in a stream of 6 slpm of dry N₂. Then, a small
468 fraction of this mixture (50 to 300 mlpm) was further mixed with the total flow of 25 slpm of N₂ (20 slpm dry N₂ + 5
469 slpm humidified N₂) before being sampled by the Br-MION-CIMS. The calibration coefficient for Cl₂ was determined
470 to be 3.5×10¹¹ molec cm⁻³ per normalized signal (cps cps⁻¹) from three separate calibration experiments (Fig 4 (bc)),
471 with an 1-sigma accuracy of 30 %.

3.3.4 HOI calibration using a calibrator

We produced a continuous HOI source via the reaction of I₂ and hydroxyl radical (OH) in a setup similar to the sulfuric acid (H₂SO₄) calibrator (Kürten et al., 2012). The OH was generated by photolyzing H₂O with a mercury (Hg) lamp at 184.9 nm, whose calibrated intensity was used to estimate the OH concentration. We tested the system by removing the I₂ or OH source from the calibrator, upon which HOI production was undetectable, confirming that any single reactant did not produce HOI. A numerical model was constructed to predict the mean HOI concentration entering the Br MION CIMS, which is analogous to the model used for H₂SO₄ calibration (Kürten et al., 2012). We only included the formation pathway of I₂ + OH to HOI in the model for simplicity; the other pathway of IO + HO₂ was considered minor as IO forms at a relatively slow rate via the reaction of I radical and O₃. ~~A Ws for the HOI calibration, we~~ produced a range of HOI concentrations by varying I₂ and OH concentrations in the calibrator. We show in Fig. 4 (ed) the linear correlation between the modeled HOI concentrations and measured HOI signals. The slope of the fit ~~line~~ corresponds to a calibration coefficient of 3.3×10^{11} molec cm⁻³ per normalized signal (cps cps⁻¹), with an overall 1- ~~sigma accuracy uncertainty~~ of $\pm 55\%$. The good correlation ($R^2 = 0.97$) including various H₂O levels also indicates that H₂O concentrations did not affect the HOI detection.





489 **Figure 4.** The normalized signals (cps cps^{-1}) absolute concentrations (molec cm^{-3})-vs. the absolute concentrations (molec cm^{-3})
 490 normalized signals (cps cps^{-1}) measured with the Br-MION-CIMS for (a) H_2SO_4 , (b) Cl_2 , (c) Cl_2HOI , and (d) $HOI-H_2SO_4$.
 491 The red dashed lines are the linear fittings. The overall factor 2 systematic 3-sigma accuracy of +50/-33 % scale uncertainty-on
 492 $[H_2SO_4]$ is not shown here.
 493
 494

495 3.3.36 Connecting sensitivity to binding enthalpy

496 Beyond the species for which we carried out calibrations, there are many more, especially iodine species, that cannot
 497 be directly calibrated due to a lack of authentic standards or generation methods. However, the sensitivity of an iodide-
 498 CIMS towards analytes can be predicted by the cluster binding enthalpy, calculated by relatively simple quantum
 499 chemical methods (Iyer et al., 2016). This holds for the bromideBr-CIMS as well. In the instrument, ion clusters,
 500 formed from reactions between analytes and reagent ions, are guided and focused by ion optics during transmission
 501 to the detector. The electric forces applied to the clusters enhance their collision energies with carrier gas molecules.
 502 If sufficient energy is transferred during the collision, cluster fragmentation may occur, affecting the instrument
 503 sensitivity for the analytes (Passananti et al., 2019). However, clusters with higher binding enthalpy will be more
 504 easily preserved and detected. Analytes that bind to the reagent ions with enthalpies higher than a critical level are
 505 likely detected at maximum sensitivity (kinetic-limited detection) by the instrument. For example, the calculated

506 critical enthalpy is 25 kcal mol⁻¹ for the iodide-CIMS used in Iyer et al. (Iyer et al., 2016) and Lopez-Hilfiker et al.
507 (Lopez-Hilfiker et al., 2016), calculated at the DLPNO-CCSD(T)/def2-QZVPP//PBE/aug-cc-pVTZ-PP level of theory.
508

509 For the bromide chemical ionization, there are two types of fragmentation pathways:

510 1) Reversion to the original form of Br⁻ and analyte



512 2) Proton transfer from the analyte to Br⁻



514 where the X-H is the hydrogen bond donor. An analyte ~~may be expected to~~should be detected at the maximum
515 sensitivity when the dissociation enthalpy for the first pathway is either a) much higher than the critical enthalpy
516 (dissociation of X-H·Br⁻ back to X-H and Br⁻ the reactants does not occur), or b) lower than the critical enthalpy, but
517 much higher than that of the second pathway (dissociation of X-H·Br⁻ to X-H and Br⁻ may occur, but dissociation to
518 X⁻ and HBr back to the reactants would occur, but it is not competitive with the other dissociation channelsis the
519 dominant pathway). Whether the enthalpy for the second pathway is higher than the critical enthalpy does not directly
520 affect the sensitivity, as both X-H·Br⁻ and X⁻ can be measured and counted. The sensitivity toward X-H would be
521 reduced only when the first reversion pathway occurs to a non-negligible extent. Taking H₂SO₄ as an example, the
522 dissociation enthalpies for the first and second pathways are 41.1 and 27.9 kcal mol⁻¹, respectively. If some of the
523 H₂SO₄·Br⁻ dissociate, they preferably become HSO₄⁻ and are detectable by the Br-CIMS. Thus, H₂SO₄ can be detected
524 at the maximum sensitivity. ~~We list the cluster formation enthalpies for a selection of halogen-containing species in~~
525 ~~Table 1 and the corresponding cluster dissociation enthalpies in Table 2.~~

526
527 While we were unable to experimentally establish a correlation between sensitivities and binding enthalpies due to
528 limited quantifiable halogen species, we can predict the tentative critical enthalpy as the binding enthalpy of a species
529 that is likely detected at the maximum sensitivity. We list the cluster formation enthalpies for a selection of halogen
530 containing species in Table 1 and the corresponding cluster dissociation enthalpies in Table 2. Among all the
531 calibration coefficients listed in Table 3, H₂SO₄ and I₂ have the lowest calibration coefficients (highest sensitivities);
532 ~~and their coefficients are almost the same for both the online and offline calibrations, with discrepancies well within~~
533 ~~the systematic uncertainties~~. Thereby, we conclude that both H₂SO₄ and I₂ are detected at the maximum sensitivity,
534 suggesting a critical enthalpy not higher than 33.7 kcal mol⁻¹. We can then infer the sensitivity for other species that
535 are difficult to calibrate by comparing their binding enthalpies to those of the benchmark species. For example, ICl
536 and IBr should have the maximum sensitivity, since the dissociation enthalpies for ICl·Br⁻ and IBr·Br⁻ are both much
537 higher than 33.7 kcal mol⁻¹ (Table 2). Although HIO₃·Br⁻ has a lower dissociation enthalpy than the critical value, the
538 favored dissociation pathway is proton transfer (the second pathway); HIO₃ can thus be considered as a maximum
539 sensitivity species detectable as IO₃⁻ ions after proton transfer. This is consistent with the fact that both HIO₃·Br⁻ and
540 IO₃⁻ are detected in Figure 1, so is the case with H₂SO₄. We thus assume that HIO₃ has a kinetic calibration coefficient
541 of 4.1×10¹⁰ molec cm⁻³ cps cps⁻¹, the value for H₂SO₄. However, the lowest dissociation enthalpies of HOI·Br⁻ and
542 Cl₂·Br⁻ are 26.9 and 22.3 kcal mol⁻¹, respectively, consistent with their higher calibration coefficients of 3.3×10¹¹ and

543 3.5×10^{11} molec cm^{-3} cps cps^{-1} . The dissociation enthalpies for $\text{IO} \cdot \text{Br}^-$, $\text{OIO} \cdot \text{Br}^-$, and $\text{HIO}_2 \cdot \text{Br}^-$ are 24.5, 23.2, and 29.2
544 kcal mol^{-1} , respectively. We would expect that their sensitivities are lower than the maximum sensitivity. Since the
545 dissociation enthalpies for $\text{IO} \cdot \text{Br}^-$ and $\text{OIO} \cdot \text{Br}^-$ are between those of $\text{HOI} \cdot \text{Br}^-$ and $\text{Cl}_2 \cdot \text{Br}^-$, a similar calibration
546 coefficient of around 3.0×10^{11} molec cm^{-3} cps cps^{-1} could may be applied, but direct calibrations are more preferable.
547 We note that when transferring the calibration factor from one species to another, the diffusivity difference should be
548 accounted, since it affects the inlet line loss. This factor is not considered in the cluster enthalpies calculations.

549
550 Further, we estimate the detection limit of the calibrated species. The detection limit is defined as the analyte
551 concentration, corresponding to the sum of the mean signal and three times the standard deviations (3σ) of the
552 background fluctuations during a two-hour background measurement. We derive the detection limit of HOI , HIO_3 , I_2 ,
553 and H_2SO_4 to be 5.8×10^6 , 1.2×10^5 , ~~3.863×10^5~~ , and 2.0×10^5 molec cm^{-3} (or 0.2, 0.005, ~~0.03015~~, and 0.008 pptv),
554 respectively, for a 2-min integration time.

555 **3.4 Comparison between Br-MION-CIMS and Br-FIGAERO-CIMS**

557 While Br-MION-CIMS and Br-FIGAERO-CIMS use the same chemical ionization scheme, their designs differ in the
558 ion-molecule reaction chamber (IMR). MION is an atmospheric pressure (1 bar) drift tube; analyte molecules gain an
559 electric charge in an axial laminar flow. FIGAERO is connected to a cone-shaped IMR chamber operated at a reduced
560 pressure (150 mbar); the sample flow is injected into the inlet via an orifice, necessarily causing turbulence and wall
561 interactions in the IMR region. The atmospheric pressure and reduced pressure IMRs are both widely used for trace
562 gas measurements. We thus compare iodine species measurements from Br-MION-CIMS and Br-FIGAERO-CIMS,
563 to better understand the performance and applicability of the bromide ionization scheme.

564 **3.4.1 Signal trend and detection limit**

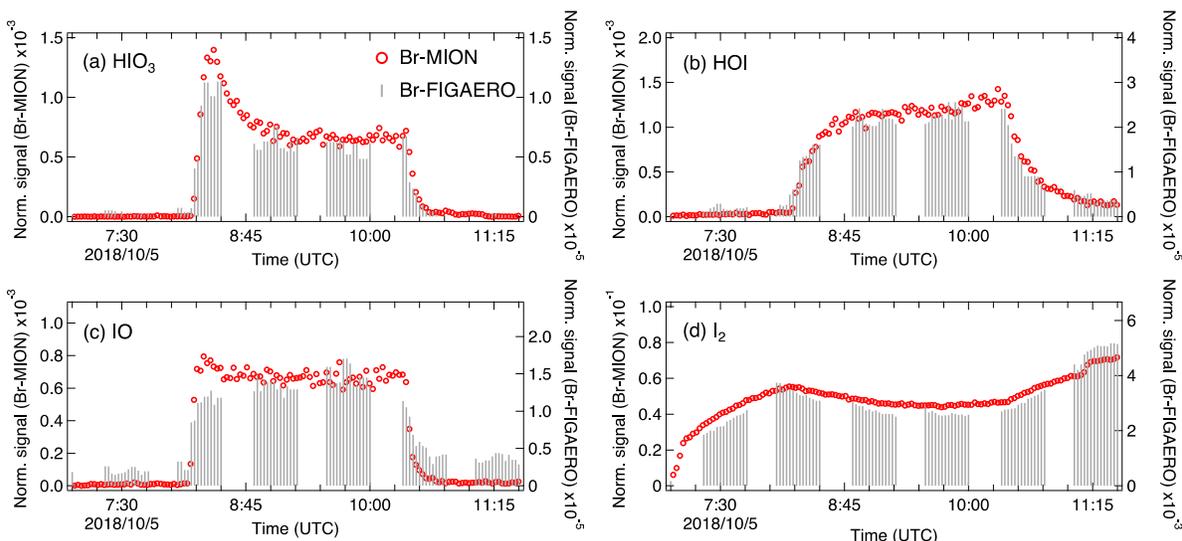
565 We show in Fig. 5 the same ~~iodine oxidation new-particle-formation~~ event as in Fig. 2, to illustrate the time series for
566 $\text{HIO}_3 \cdot \text{Br}^-$, $\text{HOI} \cdot \text{Br}^-$, $\text{IO} \cdot \text{Br}^-$, and $\text{I}_2 \cdot \text{Br}^-$, measured with Br-MION-CIMS (red circles) and Br-FIGAERO-CIMS (grey
567 sticks), respectively. Note that the FIGAERO alternates between gas and particle measurements; here we show only
568 the gas-phase signals. Clear and concurrent signals of HIO_3 , HOI , IO , and I_2 are evident from both the Br-MION-
569 CIMS and Br-FIGAERO-CIMS. Prior to the NPF-iodine oxidation event (08:11), there was no photochemical
570 production and thus virtually no signal of oxidized iodine species in both instruments. The dark reaction of ozone with
571 I_2 did not proceed at a significant rate, due to the low rate coefficient and to low levels of I_2 . Signals detected during
572 this period are considered as the persistent background, coming from electronic noise or other sources such as the
573 ionizer, carrier flows, or long-term “memory” in the case of the Br-FIGAERO-CIMS. Not surprisingly, the Br-MION-
574 CIMS has a near-zero background for all analytes. For HIO_3 (Fig. 5 (a)), the background signal in the Br-FIGAERO-
575 CIMS is also negligible; however, IO shows a substantial persistent background (Fig. 5 (c)) in the Br-FIGAERO-
576 CIMS. ~~After the NPF event (10:21), the photochemical production of oxidized iodine species was terminated and~~
577 ~~vapor concentrations decayed exponentially due to dilution and losses to chamber walls. The I_2 signal increases after~~
578 ~~the event termination because it is no longer photolyzed.~~

580
581 When we initiated the photochemistry, oxidized iodine signals rapidly built up toward an asymptote within timescales
582 of minutes. The HIO₃ signals rose with a time constant of 102 s in the Br-MION-CIMS and 108 s in the Br-FIGAERO-
583 CIMS, both with $t_{app} \approx 300$ s. The fitted time constants of HOI are slightly longer than those of HIO₃, with 120 s in
584 the Br-MION-CIMS and 114 s in the Br-FIGAERO-CIMS. IO signals stabilized the earliest, thus have the fastest time
585 constants of 48 s in the Br-MION-CIMS and 84 s in the Br-FIGAERO-CIMS. The instrumental differences in these
586 timescales are small for HIO₃ and HOI, but larger for IO. When colliding with the IMR surface, HIO₃ condenses
587 irreversibly; it thus makes sense that the Br-MION-CIMS and Br-FIGAERO-CIMS signals show the same time
588 constant timescale for HIO₃. Semi-volatile HOI, however, can return to the gas phase from the walls depending on the
589 surface coverage of HOI and the vapor concentration. Additionally, the heterogeneous reaction of aqueous iodide (I⁻)
590 and ozone (Carpenter et al., 2013) could also contribute to the emission of HOI from the IMR wall in the FIGAERO.
591 As the evaporation flux is typically a function of the amount analyte on the surface, the buffering effect could degrade
592 the instrument time response upon changes in analyte concentration. Here, however, we did not observe a significant
593 memory effect, likely because the HOI concentration was too low to fully saturate the IMR surface or because any
594 HOI evaporation was suppressed due to an enhanced accommodation coefficient of HOI on the metal surface. We
595 expect IO to be prone to loss on the metal surface due to its radical nature.

596
597 After the iodine oxidation event (10:21), the photochemical production of oxidized iodine species was terminated and
598 vapor concentrations decayed exponentially due to dilution and losses to chamber walls. Memory effects could also
599 influence the signal time constant. The dilution loss rate was around $2.1 \times 10^{-4} \text{ s}^{-1}$ (4760 s time constant) for all species
600 in the chamber, determined by the total chamber flow rate and the chamber volume. Wall loss rates, however, vary
601 for species with different diffusion constants. The decay rates of HIO₃ are 400 s for the Br-MION-CIMS and 370 s for
602 the Br-FIGAERO-CIMS, much faster than the dilution loss. For comparison, the time constant for H₂SO₄ vapor loss
603 was 300 s. These time constants are thus consistent with wall loss (around $2.2 \times 10^{-3} \text{ s}^{-1}$). The IO decay time constant
604 is 294 s for the Br-MION-CIMS and 435 s for the Br-FIGAERO-CIMS. The time constant for the Br-MION-CIMS
605 indicates that the decay of IO is also driven by wall loss, so the net flux during this period was thus towards the wall
606 rather than from the wall. Therefore, the difference of IO between instruments may well be attributed to the persistent
607 background from the ionizer of the FIGAERO. The HOI signals have longer decay time constants in both instruments
608 of 909 s for the Br-MION-CIMS and 714 s for the Br-FIGAERO-CIMS; this may reflect a time constant for depletion
609 of HOI adsorbed to the chamber walls. The I₂ signal increases after the event termination because it is no longer
610 photolyzed while its injection continues.

611
612 Applying the calibration coefficients, we convert the Br-MION-CIMS signals to absolute concentrations, and
613 subsequently correlate them with signals measured with Br-FIGAERO-CIMS. We then estimate the tentative detection
614 limits for HOI and HIO₃ in the Br-FIGAERO-CIMS to be 3.3×10^7 and $5.1 \times 10^6 \text{ molec cm}^{-3}$ (versus 5.8×10^6 and 1.2×10^5
615 molec cm^{-3} in the Br-MION-CIMS), respectively, at 3σ of the background signal for a 2-min integration time during
616 a two-hour period; they are in general one order of magnitude higher than those in the Br-MION-CIMS. This is in line

617 with the higher background signals observed in the Br-FIGAERO-CIMS. Note that for both HOI and HIO₃ the
 618 uncertainties introduced from the correlational analysis are negligible compared to the limited accuracy of the
 619 calibration sources (55 % for HOI and +50/-33 % for HIO₃). We are unable to estimate the I₂ detection limit in the Br-
 620 FIGAREO-CIMS, due to a lack of I₂ background measurement; but Br-FIGAERO-CIMS can and did detect I₂ at the
 621 low pptv level with good fidelity.
 622



623 **Figure 5.** Signal comparison of selected iodine species measured with Br-MION-CIMS and (gas-phase) Br-FIGAERO-CIMS,
 624 respectively, during the same iodine oxidation new particle formation experiment shown in Fig. 2.
 625
 626

627 3.4.2 Rise and decay time constants

628 In order to quantitatively compare the performance of the two types of IMRs, we set the initiation (08:11) and
 629 termination (10:21) of NPF event as $t = 0$, and fit the rise and decay rates of HIO₃, HOI and IO in both instruments,
 630 respectively.

631
 632 When we initiated the photochemistry, oxidized iodine species rapidly built up toward an asymptote; we thus fit their
 633 time-series signals individually with a four-parameter sigmoid function using a least-squares fitting algorithm:

$$634 \quad S_i(t) = a + (b - a) / (1 + e^{-(t - t_{app})/c}) \quad (3)$$

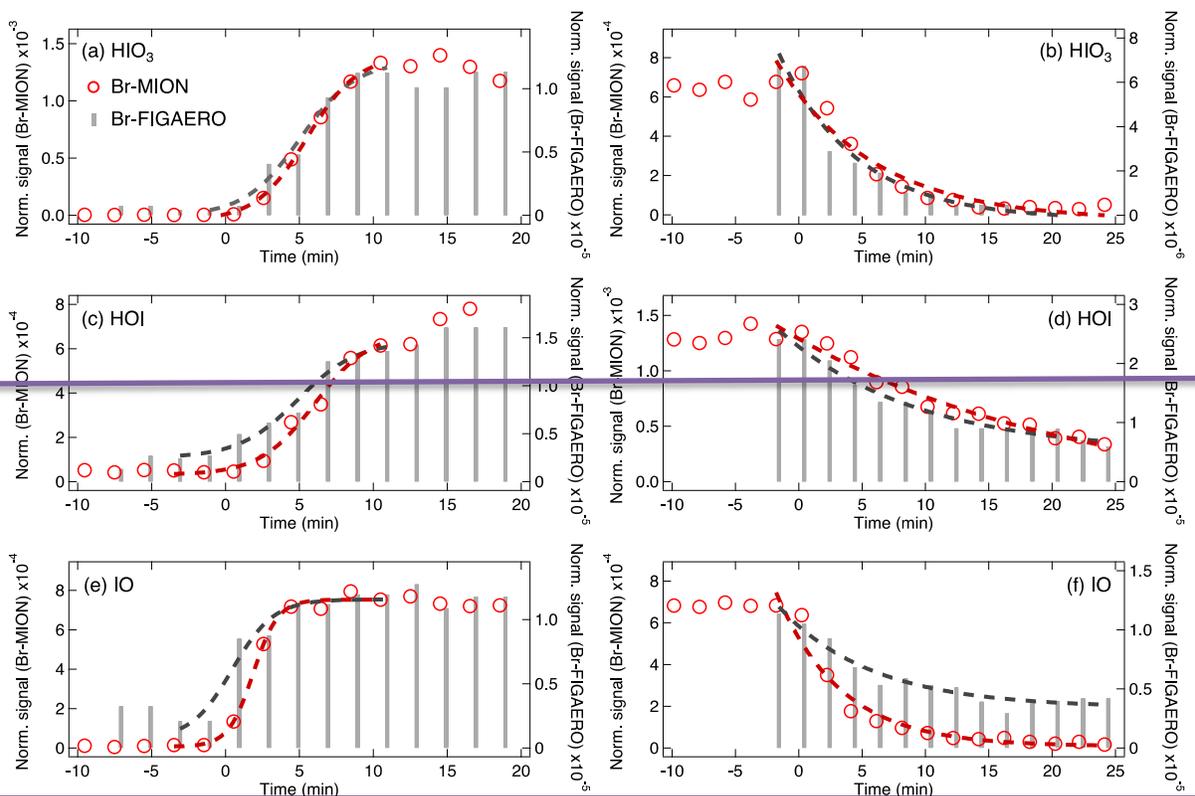
635 where a and b represent the background and asymptotic value of the sigmoid function respectively, c is the exponential
 636 time constant of the signal change, and t_{app} is the time at which the 50% value between plateau and background is
 637 reached (50% appearance time). As we were unable to separate the time scale of chemical reactions from that of
 638 instrument response, we use the time constants to represent the overall response for the purpose of comparing
 639 instrument performance.

640
 641 We show in Fig. 6 (a), (c), and (e) the rise rate fits of HIO₃, HOI, and IO, respectively. HIO₃ signals rose with a time
 642 constant of 102 s in the Br-MION-CIMS and 108 s in the Br-FIGAERO-CIMS, both with $t_{app} = 300$ s. The fitted time
 643 constants of HOI are slightly longer than those of HIO₃, with 120 s in the Br-MION-CIMS and 114 s in the Br-

644 ~~FIGAERO-CIMS. IO signals stabilized the earliest, thus have the fastest time constants of 48 s in the Br-MION-CIMS~~
645 ~~and 84 s in the Br-FIGAERO-CIMS. The instrumental differences are small for HIO₃ and HOI, but larger for IO. When~~
646 ~~colliding with the IMR surface, HIO₃ condenses irreversibly; it thus makes sense that the Br-MION-CIMS and Br-~~
647 ~~FIGAERO-CIMS signals show the same time constant for HIO₃. Semi-volatile HOI, however, can return to the gas~~
648 ~~phase from the walls depending on the surface coverage of HOI and the vapor concentration. Additionally, the~~
649 ~~heterogeneous reaction of aqueous iodide (I⁻) and ozone (Carpenter et al., 2013) could also contribute to the emission~~
650 ~~of HOI from the IMR wall in the FIGAERO. As the evaporation flux is typically a function of the amount analyte on~~
651 ~~the surface, the buffering effect could degrade the instrument time response upon changes in analyte concentration.~~
652 ~~Here, however, we did not observe a significant memory effect, likely because the HOI concentration was too low to~~
653 ~~fully saturate the IMR surface or because any HOI evaporation was suppressed due to an enhanced accommodation~~
654 ~~coefficient of HOI on the metal surface. We expect IO to be prone to loss on the metal surface due to its radical nature.~~

655
656 We also fit the exponential decay time constants of these iodine species to test this interpretation (Fig. 6 (b), (d), and
657 (f)). After photochemistry was terminated at the end of the NPF event, only two sinks drove the vapor concentration
658 decay—dilution and wall loss. Memory effects could also influence the signal time constant. The dilution loss rate
659 was around $2.1 \times 10^{-4} \text{ s}^{-1}$ (4760 s time constant) for all species in the chamber, determined by the total chamber flow
660 rate and the chamber volume. Wall loss rates, however, vary for species with different diffusion constants. The decay
661 rates of HIO₃ are 400 s for the Br-MION-CIMS and 370 s for the Br-FIGAERO-CIMS, much faster than the dilution
662 loss. For comparison, the time constant for H₂SO₄ vapor loss was 300 s. These time constants are thus consistent with
663 wall loss (around $2.2 \times 10^{-3} \text{ s}^{-1}$). The IO decay time constant is 294 s for the Br-MION-CIMS and 435 s for the Br-
664 FIGAERO-CIMS. The time constant for the Br-MION-CIMS indicates that the decay of IO is also driven by wall loss,
665 so the net flux during this period was thus towards the wall rather than from the wall. Therefore, the difference of IO
666 between instruments may well be attributed to the persistent background from the ionizer of the FIGAERO. The HOI
667 signals have longer decay time constants in both instruments of 909 s for the Br-MION-CIMS and 714 s for the Br-
668 FIGAERO-CIMS; this may reflect a time constant for depletion of HOI adsorbed to the chamber walls.

669



670
671 **Figure 6.** Rising ((a), (c), and (e)) and decay ((b), (d), and (f)) rate comparison of selected iodine species measured with Br-MION-
672 CIMS and Br-FIGAERO-CIMS, respectively, during the same iodine new particle formation event in Fig. 2 and 5.
673

674 4 Summary and conclusion

675 We confirm in this study that bromide chemical ionization is a suitable technique for the time-resolved, highly
676 sensitive, and simultaneous measurements of iodine species and sulfuric acid. The Br-MION-CIMS shows constant
677 sensitivity throughout the relative humidity range of 40 to 80 % at -10 °C, as long as the sum of the two reagent ions
678 (Br⁻ + H₂O·Br⁻) is used for signal normalization. This demonstrates the applicability of this technique to field
679 measurements in the ambient marine environment.

680
681 We quantify iodine species and sulfuric acid via offline calibrations (i.e. permeation tube and calibrator) and inter-
682 method calibrations (i.e. CE-DOAS and pre-calibrated nitrate-CIMS); ~~different methods result in consistent~~
683 ~~calibration coefficients.~~ Further, we calculate the binding enthalpies between the calibrated species and reagent ions,
684 which qualitatively agree with the corresponding calibration coefficients. This indicates that the quantum chemical
685 calculations can be employed along with the calibration experiments to determine the sensitivities for unquantifiable
686 species; more work is required to further establish the correlation between calibration coefficients and binding
687 enthalpies.

688
689 Further, using inter-method and offline calibrations, we estimate the detection limits of HOI, HIO₃, I₂, and H₂SO₄ in
690 Br-MION-CIMS being 5.8×10⁶, 1.2×10⁵, 3.86-3×10⁵, and 2.0×10⁵ molec cm⁻³, respectively, for a 2-min integration
691 time during a two-hour period. To our knowledge, the simultaneous measurements of various iodine species and

692 sulfuric acid with low detection limits are unprecedented for online techniques. Detection limits for HOI and HIO₃ in
693 the Br-FIGAERO-CIMS are 3.3×10^7 and 5.1×10^6 molec cm⁻³, which are in general one order of magnitude higher than
694 those in the Br-MION-CIMS. The signal comparison between the two instruments also shows that the Br-CIMS can
695 be coupled to both the atmospheric pressure and the reduced pressure interfaces for iodine species and sulfuric acid
696 measurements in the marine environment.

697

698

699 *Data availability.* Data available on request from the authors.

700

701 *Author Contributions.* M.W., X.-C.H., Y.-J.T and H.F. wrote the manuscript. X.-C.H., Y.-J.T., M.W. and M.Sip.
702 designed the experiments. X.-C.H., Y.-J.T. and J.S. carried out the Br-MION-CIMS measurements, M.W., D.C. and
703 V.H. carried out the Br-FIGAERO-CIMS measurements, and H.F. carried out the CE-DOAS measurements. Y.-J.T.,
704 X.-C.H., H.F., D.C., J.S. and M.Sim. performed the calibrations. S.I., X.-C.H. and T.K. carried out the quantum
705 chemical calculations. M.W. performed the comparison analysis of the Br-FIGAERO-CIMS and Br-MION-CIMS.
706 N.-M.D., T.K., M.R., R.V. and M.Sip. commented on the manuscript. All other co-authors participated in either the
707 development and preparations of the CLOUD facility and the instruments, and/or collecting and analyzing the data.

708

709 *Competing interests.* The authors declare that they have no conflict of interest.

710

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714

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720

721 **References**

722 Bitter, M., Ball, S. M., Povey, I. M. and Jones, R. L.: A broadband cavity ringdown spectrometer for in-situ
723 measurements of atmospheric trace gases, *Atmos. Chem. Phys.*, doi:10.5194/acp-5-2547-2005, 2005.

724 Carpenter, L. J., MacDonald, S. M., Shaw, M. D., Kumar, R., Saunders, R. W., Parthipan, R., Wilson, J. and Plane,
725 J. M. C.: Atmospheric iodine levels influenced by sea surface emissions of inorganic iodine, *Nat. Geosci.*,
726 doi:10.1038/ngeo1687, 2013.

727 Chai, J. Da and Head-Gordon, M.: Long-range corrected hybrid density functionals with damped atom-atom
728 dispersion corrections, *Phys. Chem. Chem. Phys.*, doi:10.1039/b810189b, 2008.

729 Chance, R. J., Shaw, M., Telgmann, L., Baxter, M. and Carpenter, L. J.: A comparison of spectrophotometric and
730 denuder based approaches for the determination of gaseous molecular iodine, *Atmos. Meas. Tech.*, doi:10.5194/amt-
731 3-177-2010, 2010.

732 Dias, A., Ehrhart, S., Vogel, A., Mathot, S., Onnela, A., Almeida, J., Kirkby, J., Williamson, C. and Mumford, S.:
733 Temperature uniformity in the CERN CLOUD chamber, *Atmos. Meas. Tech.*, 10, 5075–5088, 2017.

734 Dillon, T. J., Tucceri, M. E. and Crowley, J. N.: Laser induced fluorescence studies of iodine oxide chemistry: Part
735 II. The reactions of IO with CH₃O₂, CF₃O₂ and O₃, *Phys. Chem. Chem. Phys.*, doi:10.1039/b611116e, 2006.

736 Duplissy, J., Merikanto, J., Franchin, A., Tsagkogeorgas, G., Kangasluoma, J., Wimmer, D., Vuollekoski, H.,

- 737 Schobesberger, S., Lehtipalo, K., Flagan, R. C., Brus, D., Donahue, N. M., Vehkamäki, H., Almeida, J., Amorim,
738 A., Barmet, P., Bianchi, F., Breitenlechner, M., Dunne, E. M., Guida, R., Henschel, H., Junninen, H., Kirkby, J.,
739 Kurten, A., Kupc, A., Maattanen, A., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Praplan, A. P.,
740 Riccobono, F., Rondo, L., Steiner, G., Tome, A., Walther, H., Baltensperger, U., Carslaw, K. S., Dommen, J.,
741 Hansel, A., Petaja, T., Sipila, M., Stratmann, F., Vrtala, A., Wagner, P. E., Worsnop, D. R., Curtius, J. and Kulmala,
742 M.: Effect of ions on sulfuric acid-water binary particle formation: 2. Experimental data and comparison with QC-
743 normalized classical nucleation theory, *J. Geophys. Res.*, 121, 1752–1775, doi:10.1002/2015JD023539, 2016.
- 744 Feller, D.: The role of databases in support of computational chemistry calculations, *J. Comput. Chem.*,
745 doi:10.1002/(SICI)1096-987X(199610)17:13<1571::AID-JCC9>3.0.CO;2-P, 1996.
- 746 Finley, B. D. and Saltzman, E. S.: Observations of Cl₂, Br₂, and I₂ in coastal marine air, *J. Geophys. Res. Atmos.*,
747 doi:10.1029/2008JD010269, 2008.
- 748 Frisch, M. J., Trucks, G. W., Schlegel, H. B., Scuseria, G. E., Robb, M. A., Cheeseman, J. R., Scalmani, G., Barone,
749 V., Mennucci, B., Petersson, G. A., Nakatsuji, H., Caricato, M., Li, X., Hratchian, H. P., Izmaylov, A. F., Bloino, J.,
750 Zheng, G., Sonnenberg, J. L., Hada, M., Ehara, M., Toyota, K., Fukuda, R., Hasegawa, J., Ishida, M., Nakajima, T.,
751 Honda, Y., Kitao, O., Nakai, H., Vreven, T., Montgomery Jr., J. A., Peralta, J. E., Ogliaro, F., Bearpark, M., Heyd, J.
752 J., Brothers, E., Kudin, K. N., Staroverov, V. N., Kobayashi, R., Normand, J., Raghavachari, K., Rendell, A.,
753 Burant, J. C., Iyengar, S. S., Tomasi, J., Cossi, M., Rega, N., Millam, J. M., Klene, M., Knox, J. E., Cross, J. B.,
754 Bakken, V., Adamo, C., Jaramillo, J., Gomperts, R., Stratmann, R. E., Yazyev, O., Austin, A. J., Cammi, R.,
755 Pomelli, C., Ochterski, J. W., Martin, R. L., Morokuma, K., Zakrzewski, V. G., Voth, G. A., Salvador, P.,
756 Dannenberg, J. J., Dapprich, S., Daniels, A. D., Farkas, , Foresman, J. B., Ortiz, J. V., Cioslowski, J. and Fox, D.
757 J.: Gaussian09 Revision D.01, Gaussian Inc. Wallingford CT, Gaussian 09 Revis. C.01, 2010.
- 758 Gallagher, M. S., King, D. B., Whung, P. Y. and Saltzman, E. S.: Performance of the HPLC/fluorescence SO₂
759 detector during the GASIE instrument intercomparison experiment, *J. Geophys. Res. Atmos.*,
760 doi:10.1029/97jd00700, 1997.
- 761 Gómez Martín, J. C., Blahins, J., Gross, U., Ingham, T., Goddard, A., Mahajan, A. S., Ubelis, A. and Saiz-Lopez,
762 A.: In situ detection of atomic and molecular iodine using Resonance and Off-Resonance Fluorescence by Lamp
763 Excitation: ROFLEX, *Atmos. Meas. Tech.*, doi:10.5194/amt-4-29-2011, 2011.
- 764 Gómez Martín, J. C., Gálvez, O., Baeza-Romero, M. T., Ingham, T., Plane, J. M. C. and Blitz, M. A.: On the
765 mechanism of iodine oxide particle formation, *Phys. Chem. Chem. Phys.*, doi:10.1039/c3cp51217g, 2013.
- 766 He, X.-C., Iyer, S., Sipilä, M., Ylisirniö, A., Peltola, M., Kontkanen, J., Baalbaki, R., Simon, M., Kürten, A., Tham,
767 Y. J., Pesonen, J., Ahonen, L. R., Amanatidis, S., Amorim, A., Baccarini, A., Beck, L., Bianchi, F., Brilke, S., Chen,
768 D., Chiu, R., Curtius, J., Dada, L., Dias, A., Dommen, J., Donahue, N. M., Duplissy, J., El Haddad, I., Finkenzeller,
769 H., Fischer, L., Heinritzi, M., Hofbauer, V., Kangasluoma, J., Kim, C., Koenig, T. K., Kubečka, J., Kvashnin, A.,
770 Lamkaddam, H., Lee, C. P., Leiminger, M., Li, Z., Makhmutov, V., Xiao, M., Marten, R., Nie, W., Onnela, A.,
771 Partoll, E., Petäjä, T., Salo, V.-T., Schuchmann, S., Steiner, G., Stolzenburg, D., Stozhkov, Y., Tauber, C., Tomé,
772 A., Väisänen, O., Vazquez-Pufleau, M., Volkamer, R., Wagner, A. C., Wang, M., Wang, Y., Wimmer, D., Winkler,
773 P. M., Worsnop, D. R., Wu, Y., Yan, C., Ye, Q., Lehtinen, K., Nieminen, T., Manninen, H. E., Rissanen, M.,
774 Schobesberger, S., Lehtipalo, K., Baltensperger, U., Hansel, A., Kerminen, V.-M., Flagan, R. C., Kirkby, J., Kurtén,
775 T. and Kulmala, M.: Determination of the collision rate coefficient between charged iodine acid clusters and iodine
776 acid using the appearance time method, *Aerosol Sci. Technol.*, 55(2), 231–242,
777 doi:10.1080/02786826.2020.1839013, 2021a.
- 778 He, X.-C., Tham, Y. J., Dada, L., Wang, M., Finkenzeller, H., Stolzenburg, D., Iyer, S., Simon, M., Kürten, A.,
779 Shen, J. and others: Role of iodine oxoacids in atmospheric aerosol nucleation, *Science* (80-.), 371(6529), 589–
780 595, 2021b.
- 781 He, X.: From the measurement of halogenated species to iodine particle formation, 2017.
- 782 Iyer, S., Lopez-Hilfiker, F., Lee, B. H., Thornton, J. A. and Kurtén, T.: Modeling the Detection of Organic and
783 Inorganic Compounds Using Iodide-Based Chemical Ionization, *J. Phys. Chem. A*, doi:10.1021/acs.jpca.5b09837,
784 2016.
- 785 Junninen, H., Ehn, M., Petaja, T., Luosujärvi, L., Kotiaho, T., Kostianen, R., Rohner, U., Gonin, M., Fuhrer, K.,
786 Kulmala, M., Worsnop, D. R., Petäjä, T., Luosujärvi, L., Kotiaho, T., Kostianen, R., Rohner, U., Gonin, M., Fuhrer,

- 787 K., Kulmala, M. and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion composition,
788 *Atmos. Meas. Tech.*, 3(4), 1039–1053, doi:10.5194/amt-3-1039-2010, 2010.
- 789 Kazantseva, N. N., Ernepesova, A., Khodjamamedov, A., Geldyev, O. A. and Krungal, B. S.: Spectrophotometric
790 analysis of iodide oxidation by chlorine in highly mineralized solutions, in *Analytica Chimica Acta.*, 2002.
- 791 Kendall, R. A., Dunning, T. H. and Harrison, R. J.: Electron affinities of the first-row atoms revisited. Systematic
792 basis sets and wave functions, *J. Chem. Phys.*, doi:10.1063/1.462569, 1992.
- 793 Kercher, J. P., Riedel, T. P. and Thornton, J. A.: Chlorine activation by N₂O₅: Simultaneous, in situ detection of
794 ClNO₂ and N₂O₅ by chemical ionization mass spectrometry, *Atmos. Meas. Tech.*, doi:10.5194/amt-2-193-2009,
795 2009.
- 796 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten,
797 A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim,
798 A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S.,
799 Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J.,
800 Lovejoy, E. R., Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A.,
801 Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A.,
802 Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M.,
803 Carslaw, K. S., Worsnop, D. R., Baltensperger, U. and Kulmala, M.: Role of sulphuric acid, ammonia and galactic
804 cosmic rays in atmospheric aerosol nucleation, *Nature*, 476, 429 [online] Available from:
805 <https://doi.org/10.1038/nature10343>, 2011.
- 806 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan, C.,
807 Almeida, J., Trostl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer, A. K.,
808 Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X. M., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin,
809 A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M.,
810 Kurten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Perakyla, O., Piel,
811 F., Petaja, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A. D. D., Riipinen, I., Rissanen, M. P., Rondo, L.,
812 Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F.,
813 Tome, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M.,
814 Ye, P. L., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M.,
815 Carslaw, K. S., Curtius, J., Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E.,
816 Wimmer, D., Winkler, P. M., Ye, P. L., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R.,
817 Baltensperger, U., Kulmala, M., Carslaw, K. S. and Curtius, J.: Ion-induced nucleation of pure biogenic particles,
818 *Nature*, 533(7604), 521–526, doi:10.1038/nature17953, 2016.
- 819 Koenig, T. K., Baidar, S., Campuzano-Jost, P., Cuevas, C. A., Dix, B., Fernandez, R. P., Guo, H., Hall, S. R.,
820 Kinnison, D., Nault, B. A., Ullmann, K., Jimenez, J. L., Saiz-Lopez, A. and Volkamer, R.: Quantitative detection of
821 iodine in the stratosphere, *Proc. Natl. Acad. Sci. U. S. A.*, doi:10.1073/pnas.1916828117, 2020.
- 822 Kurten, A., Rondo, L., Ehrhart, S. and Curtius, J.: Calibration of a Chemical Ionization Mass Spectrometer for the
823 Measurement of Gaseous Sulfuric Acid, *J. Phys. Chem. A*, 116, 6375–6386, doi:10.1021/jp212123n, 2012.
- 824 Lawler, M. J., Sander, R., Carpenter, L. J., Lee, J. D., Von Glasow, R., Sommariva, R. and Saltzman, E. S.: HOCl
825 and Cl₂ observations in marine air, *Atmos. Chem. Phys.*, doi:10.5194/acp-11-7617-2011, 2011.
- 826 Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurten, T., Worsnop, D. R. and Thornton, J. A.: An iodide-adduct high-
827 resolution time-of-flight chemical-ionization mass spectrometer: application to atmospheric inorganic and organic
828 compounds, *Env. Sci Technol.*, 48(11), 6309–6317, doi:10.1021/es500362a, 2014.
- 829 Lee, B. H., Lopez-Hilfiker, F. D., Veres, P. R., McDuffie, E. E., Fibiger, D. L., Sparks, T. L., Ebben, C. J., Green, J.
830 R., Schroder, J. C., Campuzano-Jost, P., Iyer, S., D’Ambro, E. L., Schobesberger, S., Brown, S. S., Wooldridge, P.
831 J., Cohen, R. C., Fiddler, M. N., Bililign, S., Jimenez, J. L., Kurtén, T., Weinheimer, A. J., Jaegle, L. and Thornton,
832 J. A.: Flight Deployment of a High-Resolution Time-of-Flight Chemical Ionization Mass Spectrometer:
833 Observations of Reactive Halogen and Nitrogen Oxide Species, *J. Geophys. Res. Atmos.*,
834 doi:10.1029/2017JD028082, 2018.
- 835 Leigh, R. J., Ball, S. M., Whitehead, J., Leblanc, C., Shillings, A. J. L., Mahajan, A. S., Oetjen, H., Lee, J. D., Jones,
836 C. E., Dorsey, J. R., Gallagher, M., Jones, R. L., Plane, J. M. C., Potin, P. and McFiggans, G.: Measurements and

837 modelling of molecular iodine emissions, transport and photodestruction in the coastal region around Roscoff,
838 Atmos. Chem. Phys., doi:10.5194/acp-10-11823-2010, 2010.

839 Lopez-Hilfiker, F. D., Mohr, C., Ehn, M., Rubach, F., Kleist, E., Wildt, J., Mentel, T. F., Lutz, A., Hallquist, M.,
840 Worsnop, D. and Thornton, J. A.: A novel method for online analysis of gas and particle composition: description
841 and evaluation of a Filter Inlet for Gases and AEROSols (FIGAERO), Atmos. Meas. Tech., 7(4), 983–1001,
842 doi:10.5194/amt-7-983-2014, 2014.

843 Lopez-Hilfiker, F. D., Iyer, S., Mohr, C., Lee, B. H., D’ambro, E. L., Kurtén, T. and Thornton, J. A.: Constraining
844 the sensitivity of iodide adduct chemical ionization mass spectrometry to multifunctional organic molecules using
845 the collision limit and thermodynamic stability of iodide ion adducts, Atmos. Meas. Tech., doi:10.5194/amt-9-1505-
846 2016, 2016.

847 Marcy, T. P., Fahey, D. W., Gao, R. S., Popp, P. J., Richard, E. C., Thompson, T. L., Rosenlof, K. H., Ray, E. A.,
848 Salawitch, R. J., Atherton, C. S., Bergmann, D. J., Ridley, B. A., Weinheimer, A. J., Loewenstein, M., Weinstock, E.
849 M. and Mahoney, M. J.: Quantifying Stratospheric Ozone in the Upper Troposphere with in Situ Measurements of
850 HCl, Science (80-), doi:10.1126/science.1093418, 2004.

851 McFiggans, G., Coe, H., Burgess, R., Allan, J., Cubison, M., Alfarra, M. R., Saunders, R., Saiz-Lopez, A., Plane, J.
852 M. C., Wevill, D. J., Carpenter, L. J., Rickard, A. R. and Monks, P. S.: Direct evidence for coastal iodine particles
853 from Laminaria macroalgae - Linkage to emissions of molecular iodine, Atmos. Chem. Phys., doi:10.5194/acp-4-
854 701-2004, 2004.

855 Meinen, J., Thieser, J., Platt, U. and Leisner, T.: Technical Note: Using a high finesse optical resonator to provide a
856 long light path for differential optical absorption spectroscopy: CE-DOAS, Atmos. Chem. Phys., doi:10.5194/acp-
857 10-3901-2010, 2010.

858 Neese, F.: The ORCA program system, Wiley Interdiscip. Rev. Comput. Mol. Sci., doi:10.1002/wcms.81, 2012.

859 O’dowd, C. D., Jimenez, J. L., Bahreini, R., Flagan, R. C., Seinfeld, J. H., Hämerl, K., Pirjola, L., Kulmala, M. and
860 Hoffmann, T.: Marine aerosol formation from biogenic iodine emissions, Nature, doi:10.1038/nature00775, 2002.

861 Passananti, M., Zapadinsky, E., Zanca, T., Kangasluoma, J., Myllys, N., Rissanen, M. P., Kurtén, T., Ehn, M.,
862 Attoui, M. and Vehkamäki, H.: How well can we predict cluster fragmentation inside a mass spectrometer?, Chem.
863 Commun., doi:10.1039/c9cc02896j, 2019.

864 Pfeifer, J., Simon, M., Heinritzi, M., Piel, F., Weitz, L., Wang, D., Granzin, M., Müller, T., Bräkling, S., Kirkby, J.,
865 Curtius, J. and Kürten, A.: Measurement of ammonia, amines and iodine compounds using protonated water cluster
866 chemical ionization mass spectrometry, Atmos. Meas. Tech., doi:10.5194/amt-13-2501-2020, 2020.

867 Prados-Roman, C., Cuevas, C. A., Fernandez, R. P., Kinnison, D. E., Lamarque, J. F. and Saiz-Lopez, A.: A
868 negative feedback between anthropogenic ozone pollution and enhanced ocean emissions of iodine, Atmos. Chem.
869 Phys., doi:10.5194/acp-15-2215-2015, 2015.

870 Raso, A. R. W., Custard, K. D., May, N. W., Tanner, D., Newburn, M. K., Walker, L., Moore, R. J., Huey, L. G.,
871 Alexander, L., Shepson, P. B. and Pratt, K. A.: Active molecular iodine photochemistry in the Arctic, Proc. Natl.
872 Acad. Sci. U. S. A., doi:10.1073/pnas.1702803114, 2017.

873 Riplinger, C. and Neese, F.: An efficient and near linear scaling pair natural orbital based local coupled cluster
874 method, J. Chem. Phys., doi:10.1063/1.4773581, 2013.

875 Rissanen, M. P., Mikkilä, J., Iyer, S. and Hakala, J.: Multi-scheme chemical ionization inlet (MION) for fast
876 switching of reagent ion chemistry in atmospheric pressure chemical ionization mass spectrometry (CIMS)
877 applications, Atmos. Meas. Tech., doi:10.5194/amt-12-6635-2019, 2019.

878 Rothman, L. S., Gordon, I. E., Barber, R. J., Dothe, H., Gamache, R. R., Goldman, A., Perevalov, V. I., Tashkun, S.
879 A. and Tennyson, J.: HITEMP, the high-temperature molecular spectroscopic database, J. Quant. Spectrosc. Radiat.
880 Transf., doi:10.1016/j.jqsrt.2010.05.001, 2010.

881 Saiz-Lopez, A., Plane, J. M. C., Baker, A. R., Carpenter, L. J., Von Glasow, R., Gómez Martín, J. C., McFiggans, G.
882 and Saunders, R. W.: Atmospheric chemistry of iodine, Chem. Rev., doi:10.1021/cr200029u, 2012.

883 Sanchez, J., Tanner, D. J., Chen, D., Huey, L. G. and Ng, N. L.: A new technique for the direct detection of HO₂
884 radicals using bromide chemical ionization mass spectrometry (Br-CIMS): Initial characterization, Atmos. Meas.

885 Tech., doi:10.5194/amt-9-3851-2016, 2016.

886 Schnitzhofer, R., Metzger, A., Breitenlechner, M., Jud, W., Heinritzi, M., De Menezes, L.-P., Duplissy, J., Guida,
887 R., Haider, S., Kirkby, J. and others: Characterisation of organic contaminants in the CLOUD chamber at CERN,
888 Atmos. Meas. Tech., 7(7), 2159–2168, 2014.

889 Schwehr, K. A., Santschi, P. H. and Elmore, D.: The dissolved organic iodine species of the isotopic ratio of
890 129I/127I: A novel tool for tracing terrestrial organic carbon in the estuarine surface waters of Galveston Bay,
891 Texas, Limnol. Oceanogr. Methods, doi:10.4319/lom.2005.3.326, 2005.

892 Sherwen, T., Evans, M. J., Carpenter, L. J., Andrews, S. J., Lidster, R. T., Dix, B., Koenig, T. K., Sinreich, R.,
893 Ortega, I., Volkamer, R., Saiz-Lopez, A., Prados-Roman, C., Mahajan, A. S. and Ordóñez, C.: Iodine's impact on
894 tropospheric oxidants: A global model study in GEOS-Chem, Atmos. Chem. Phys., doi:10.5194/acp-16-1161-2016,
895 2016.

896 Simpson, W. R., Brown, S. S., Saiz-Lopez, A., Thornton, J. A. and Von Glasow, R.: Tropospheric Halogen
897 Chemistry: Sources, Cycling, and Impacts, Chem. Rev., doi:10.1021/cr5006638, 2015.

898 Sipilä, M., Sarnela, N., Jokinen, T., Henschel, H., Junninen, H., Kontkanen, J., Richters, S., Kangasluoma, J.,
899 Franchin, A., Peräkylä, O., Rissanen, M. P., Ehn, M., Vehkamäki, H., Kurten, T., Berndt, T., Petäjä, T., Worsnop,
900 D., Ceburnis, D., Kerminen, V. M., Kulmala, M. and O'Dowd, C.: Molecular-scale evidence of aerosol particle
901 formation via sequential addition of HIO₃, Nature, doi:10.1038/nature19314, 2016.

902 Spietz, P., Martín, J. G. and Burrows, J. P.: Effects of column density on i2 spectroscopy and a determination of i2
903 absorption cross section at 500 nm, Atmos. Chem. Phys., doi:10.5194/acp-6-2177-2006, 2006.

904 Spolaor, A., Gabrieli, J., Martma, T., Kohler, J., Björkman, M. B., Isaksson, E., Varin, C., Vallelonga, P., Plane, J.
905 M. C. and Barbante, C.: Sea ice dynamics influence halogen deposition to Svalbard, Cryosphere, doi:10.5194/tc-7-
906 1645-2013, 2013.

907 Stolzenburg, D., Simon, M., Ranjithkumar, A., Kürten, A., Lehtipalo, K., Gordon, H., Ehrhart, S., Finkenzeller, H.,
908 Pichelstorfer, L., Nieminen, T. and others: Enhanced growth rate of atmospheric particles from sulfuric acid, Atmos.
909 Chem. Phys., 20(12), 7359–7372, 2020.

910 Thalman, R. and Volkamer, R.: Temperature dependent absorption cross-sections of O₂-O₂ collision pairs between
911 340 and 630 nm and at atmospherically relevant pressure, Phys. Chem. Chem. Phys., doi:10.1039/c3cp50968k,
912 2013.

913 Tham, Y. J., Wang, Z., Li, Q., Yun, H., Wang, W., Wang, X., Xue, L., Lu, K., Ma, N., Bohn, B., Li, X., Kecorius,
914 S., Größ, J., Shao, M., Wiedensohler, A., Zhang, Y. and Wang, T.: Significant concentrations of nitryl chloride
915 sustained in the morning: Investigations of the causes and impacts on ozone production in a polluted region of
916 northern China, Atmos. Chem. Phys., doi:10.5194/acp-16-14959-2016, 2016.

917 Tham, Y. J., He, X.-C., Li, Q., Cuevas, C. A., Shen, J., Kalliokoski, J., Yan, C., Iyer, S., Lehmusjärvi, T., Jang, S.
918 and others: Direct field evidence of autocatalytic iodine release from atmospheric aerosol, Proc. Natl. Acad. Sci.,
919 118(4), 2021.

920 Vandaele, A. C., Hermans, C., Simon, P. C., Carleer, M., Colin, R., Fally, S., Mérienne, M. F., Jenouvrier, A. and
921 Coquart, B.: Measurements of the NO₂ absorption cross-section from 42 000 cm⁻¹ to 10 000 cm⁻¹ (238-1000 nm) at
922 220 K and 294 K, J. Quant. Spectrosc. Radiat. Transf., doi:10.1016/S0022-4073(97)00168-4, 1998.

923 Vaughan, S., Gherman, T., Ruth, A. A. and Orphal, J.: Incoherent broad-band cavity-enhanced absorption
924 spectroscopy of the marine boundary layer species I₂, IO and OIO, Phys. Chem. Chem. Phys.,
925 doi:10.1039/b802618a, 2008.

926 Wang, M., Kong, W., Marten, R., He, X. C., Chen, D., Pfeifer, J., Heitto, A., Kontkanen, J., Dada, L., Kürten, A.,
927 Yli-Juuti, T., Manninen, H. E., Amanatidis, S., Amorim, A., Baalbaki, R., Baccharini, A., Bell, D. M., Bertozzi, B.,
928 Bräkling, S., Brilke, S., Murillo, L. C., Chiu, R., Chu, B., De Menezes, L. P., Duplissy, J., Finkenzeller, H.,
929 Carracedo, L. G., Granzin, M., Guida, R., Hansel, A., Hofbauer, V., Krechmer, J., Lehtipalo, K., Lamkaddam, H.,
930 Lampimäki, M., Lee, C. P., Makhmutov, V., Marie, G., Mathot, S., Mauldin, R. L., Mentler, B., Müller, T., Onnela,
931 A., Partoll, E., Petäjä, T., Philippov, M., Pospisilova, V., Ranjithkumar, A., Rissanen, M., Rörup, B., Scholz, W.,
932 Shen, J., Simon, M., Sipilä, M., Steiner, G., Stolzenburg, D., Tham, Y. J., Tomé, A., Wagner, A. C., Wang, D. S.,
933 Wang, Y., Weber, S. K., Winkler, P. M., Wlasits, P. J., Wu, Y., Xiao, M., Ye, Q., Zauner-Wieczorek, M., Zhou, X.,

- 934 Volkamer, R., Riipinen, I., Dommen, J., Curtius, J., Baltensperger, U., Kulmala, M., Worsnop, D. R., Kirkby, J.,
935 Seinfeld, J. H., El-Haddad, I., Flagan, R. C. and Donahue, N. M.: Rapid growth of new atmospheric particles by
936 nitric acid and ammonia condensation, *Nature*, doi:10.1038/s41586-020-2270-4, 2020.
- 937 Washenfelder, R. A., Langford, A. O., Fuchs, H. and Brown, S. S.: Measurement of glyoxal using an incoherent
938 broadband cavity enhanced absorption spectrometer, *Atmos. Chem. Phys.*, doi:10.5194/acp-8-7779-2008, 2008.
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941 **Table 1: Cluster formation enthalpies of different species with bromide ions.** The cluster geometries are optimized
 942 at the ω B97X-D/aug-cc-pVTZ-PP level at 298.15 K. The enthalpies are calculated at the DLPNO-CCSD(T)/def2-
 943 QZVPP// ω B97xD/aug-cc-pVTZ-PP level at 298.15 K.

Cluster formation pathway	Formation enthalpies (kcal mol ⁻¹)
$\text{Cl}_2 + \text{Br}^- \rightarrow \text{Cl}_2 \cdot \text{Br}^-$	-22.3
$\text{OIO} + \text{Br}^- \rightarrow \text{OIO} \cdot \text{Br}^-$	-23.2
$\text{IO} + \text{Br}^- \rightarrow \text{IO} \cdot \text{Br}^-$	-24.5
$\text{HIO}_3 + \text{Br}^- \rightarrow \text{HIO}_3 \cdot \text{Br}^-$	-26.6
$\text{HOI} + \text{Br}^- \rightarrow \text{HOI} \cdot \text{Br}^-$	-26.9
$\text{HIO}_2 + \text{Br}^- \rightarrow \text{HIO}_2 \cdot \text{Br}^-$	-29.2
$\text{I}_2 + \text{Br}^- \rightarrow \text{I}_2 \cdot \text{Br}^-$	-33.7
$\text{ICl} + \text{Br}^- \rightarrow \text{ICl} \cdot \text{Br}^-$	-33.8
$\text{IBr} + \text{Br}^- \rightarrow \text{IBr} \cdot \text{Br}^-$	-36.7
$\text{H}_2\text{SO}_4 + \text{Br}^- \rightarrow \text{H}_2\text{SO}_4 \cdot \text{Br}^-$	-41.1
$\text{I}_2\text{O}_4 + \text{Br}^- \rightarrow \text{I}_2\text{O}_4 \cdot \text{Br}^-$	-42.6
$\text{I}_2\text{O}_5 + \text{Br}^- \rightarrow \text{I}_2\text{O}_5 \cdot \text{Br}^-$	-53.2

944
 945 **Table 2: Fragmentation reaction enthalpies of different species with bromide ions.** The cluster geometries are
 946 optimized at the ω B97X-D/aug-cc-pVTZ-PP level at 298.15 K. The enthalpies are calculated at the DLPNO-
 947 CCSD(T)/def2-QZVPP// ω B97xD/aug-cc-pVTZ-PP level at 298.15 K.

Cluster fragmentation pathway	Fragmentation enthalpies (kcal mol ⁻¹)
$\text{Cl}_2 \cdot \text{Br}^- \rightarrow \text{Cl}_2 + \text{Br}^-$	22.3
$\text{Cl}_2 \cdot \text{Br}^- \rightarrow \text{BrCl} + \text{Cl}^-$	22.3
$\text{HIO}_3 \cdot \text{Br}^- \rightarrow \text{HIO}_3 + \text{Br}^-$	26.6
$\text{HIO}_3 \cdot \text{Br}^- \rightarrow \text{IO}_3^- + \text{HBr}$	20.8
$\text{HIO}_3 \cdot \text{Br}^- \rightarrow \text{IO}_2^- + \text{HOBr}$	52.0
$\text{HOI} \cdot \text{Br}^- \rightarrow \text{HOI} + \text{Br}^-$	26.9
$\text{HOI} \cdot \text{Br}^- \rightarrow \text{IO}^- + \text{HBr}$	57.7
$\text{HOI} \cdot \text{Br}^- \rightarrow \text{I}^- + \text{HOBr}$	31.3
$\text{HIO}_2 \cdot \text{Br}^- \rightarrow \text{HIO}_2 + \text{Br}^-$	29.2
$\text{HIO}_2 \cdot \text{Br}^- \rightarrow \text{IO}_2^- + \text{HBr}$	43.8
$\text{HIO}_2 \cdot \text{Br}^- \rightarrow \text{IO}^- + \text{HOBr}$	42.2
$\text{I}_2 \cdot \text{Br}^- \rightarrow \text{I}_2 + \text{Br}^-$	33.7

$I_2 \cdot Br^- \rightarrow IBr + I^-$	33.8
$ICl \cdot Br^- \rightarrow ICl + Br^-$	33.8
$ICl \cdot Br^- \rightarrow IBr + Cl^-$	39.8
$ICl \cdot Br^- \rightarrow BrCl + I^-$	42.0
$IBr \cdot Br^- \rightarrow IBr + Br^-$	36.7
$IBr \cdot Br^- \rightarrow Br_2 + I^-$	39.4
$H_2SO_4 \cdot Br^- \rightarrow H_2SO_4 + Br^-$	41.1
$H_2SO_4 \cdot Br^- \rightarrow HSO_4^- + HBr$	27.9

948

949 **Table 3: Calibration coefficients for selected species.**

Species	Calibration coefficient (molec cm ⁻³ cps cps ⁻¹)	Detection limit (molec cm ⁻³)
^a I ₂	^a 2.7×10 ¹⁰	<u>3.8×10⁵</u>
^a H ₂ SO ₄	<u>4.1×10¹⁰</u>	<u>2.0×10⁵</u>
^b I ₂	^b 6.3×10 ¹⁰	<u>8.8×10⁵</u>
^b Cl ₂	^b 3.5×10 ¹¹	n/a
^b HOI	^b 3.3×10 ¹¹	<u>5.8×10⁶</u>
H₂SO₄	^a 4.1×10¹⁰	
^c HIO ₃	<u>4.1×10¹⁰</u>	<u>1.2×10⁵</u>

950 ^a: Inter-method calibrations

951 ^b: Offline calibrations

952 ^c: Derived from dissociation enthalpies