



#### Measurement of iodine species and sulfuric acid using bromide 1 chemical ionization mass spectrometers 2

3

4 Mingyi Wang<sup>1,2\*</sup>, Xu-Cheng He<sup>3\*#</sup>, Henning Finkenzeller<sup>4</sup>, Siddharth Iyer<sup>3</sup>, Dexian Chen<sup>1,5</sup>, Jiali Shen<sup>3</sup>, Mario Simon<sup>6</sup>, Victoria Hofbauer<sup>1,2</sup>, Jasper Kirkby<sup>6,7</sup>, Joachim Curtius<sup>6</sup>, Norbert Maier<sup>8</sup>, 5 Theo Kurtén<sup>3,8</sup>, Douglas R. Worsnop<sup>3,9</sup>, Markku Kulmala<sup>3,10,11,12</sup>, Matti Rissanen<sup>3,13</sup>, Rainer 6 Volkamer<sup>4</sup>, Yee Jun Tham<sup>3#</sup>, Neil M. Donahue<sup>1,2,5,14</sup>, and Mikko Sipilä<sup>3</sup> 7

- 8
- 9 <sup>1</sup>Center for Atmospheric Particle Studies, Carnegie Mellon University, Pittsburgh, PA, 15213, USA
- 10 <sup>2</sup>Department of Chemistry, Carnegie Mellon University, Pittsburgh, PA, 15213, USA
- 11 <sup>3</sup>Institute for Atmospheric and Earth System Research (INAR), University of Helsinki, 00014 Helsinki, Finland
- 12 <sup>4</sup>Department of Chemistry & CIRES, University of Colorado Boulder, Boulder, CO 80309, USA
- 13 <sup>5</sup>Department of Chemical Engineering, Carnegie Mellon University, Pittsburgh, PA, 15213, USA
- 14 <sup>6</sup>Institute for Atmospheric and Environmental Sciences, Goethe University Frankfurt, 60438 Frankfurt am Main, 15 Germany
- 16 <sup>7</sup>CERN, the European Organization for Nuclear Research, CH-1211 Geneve 23, Switzerland
- 17 <sup>8</sup>Department of Chemistry, University of Helsinki, 00014 Helsinki, Finland
- 18 <sup>9</sup>Aerodyne Research, Inc., Billerica, MA, 01821, USA
- 19 <sup>10</sup>Helsinki Institute of Physics, P.O. Box 64 (Gustaf Hallstromin katu 2), FI-00014 University of Helsinki, Finland
- 20 <sup>11</sup>Joint International Research Laboratory of Atmospheric and Earth System Sciences, Nanjing University, Nanjing, 21 China
- 22 <sup>12</sup>Aerosol and Haze Laboratory, Beijing Advanced Innovation Center for Soft Matter Science and Engineering, Beijing 23 University of Chemical Technology, Beijing, China
- 24 25 <sup>13</sup>Aerosol Physics Laboratory, Physics Unit, Faculty of Engineering and Natural Sciences, Tampere University,
- Tampere, Finland
- 26 27 <sup>14</sup>Department of Engineering and Public Policy, Carnegie Mellon University, Pittsburgh, PA, 15213, USA
- \*These authors contributed equally to this work
- 28
- 29 <sup>#</sup>Correspondence to:
- 30 Xu-Cheng He (xucheng.he@helsinki.fi) and Yee Jun Tham (yee.tham@helsinki.fi).
- 31
- 32 Abstract. Iodine species are important in the marine atmosphere for oxidation and new-particle formation.
- 33 Understanding iodine chemistry and iodine new-particle formation requires high time resolution, high sensitivity, and
- 34 simultaneous measurements of many iodine species. Here, we describe the application of bromide chemical ionization
- 35 mass spectrometers (Br-CIMS) to this task. During iodine new-particle formation experiments in the Cosmics Leaving
- 36 OUtdoor Droplets (CLOUD) chamber, we have measured gas-phase iodine species and sulfuric acid using two Br-
- 37 CIMS, one coupled to a Multi-scheme chemical IONization inlet (Br-MION-CIMS) and the other to a Filter Inlet for
- 38 Gasses and AEROsols inlet (Br-FIGAERO-CIMS). From offline calibrations and inter-comparisons with other
- 39 instruments attached to the CLOUD chamber, we have quantified the sensitivities of the Br-MION-CIMS to HOI, I2,
- 40 and H<sub>2</sub>SO<sub>4</sub> and obtain detection limits of 5.8×10<sup>6</sup>, 6.3×10<sup>5</sup>, and 2.0×10<sup>5</sup> molec cm<sup>-3</sup>, respectively, for a 2-min
- 41 integration time. From binding energy calculations, we estimate the detection limit for HIO<sub>3</sub> to be  $1.2 \times 10^5$  molec cm<sup>-</sup>
- 42 <sup>3</sup>, based on an assumption of maximum sensitivity. Detection limits in the Br-FIGAERO-CIMS are around one order
- 43 of magnitude higher than those in the Br-MION-CIMS; for example, the detection limits for HOI and HIO<sub>3</sub> are  $3.3 \times 10^7$
- 44 and 5.1×106 molec cm<sup>-3</sup>, respectively. Our comparisons of the performance of the MION inlet and the FIGAERO inlet
  - 1





- 45 show that bromide chemical ionization mass spectrometers using either atmospheric pressure or reduced pressure
- 46 interfaces are well-matched to measuring iodine species and sulfuric acid in marine environments.
- 47

# 48 1 Introduction

49 Reactive iodine species are released into the atmosphere mainly by biological processes in marine environments (i.e. 50 from macro- and micro-algae) (McFiggans et al., 2004), O<sub>3</sub> deposition on the sea surface (Carpenter et al., 2013), as 51 well as from the sea ice (Spolaor et al., 2013) and snowpack in the polar region (Raso et al., 2017). Once emitted, 52 iodine species can modify atmospheric oxidative capacity via a chain of catalytic reactions with O<sub>3</sub> that form iodine 53 oxides, leading to about 20-28 % of O<sub>3</sub> loss in the marine boundary layer (Prados-Roman et al., 2015; Sherwen et al., 54 2016). Through convection, reactive iodine species can be transported from the lower troposphere to the upper 55 troposphere-lower stratosphere, causing one third of the iodine-induced ozone loss in the upper troposphere-lower 56 stratosphere (Koenig et al., 2020). Another important effect of iodine species is their contribution to atmospheric new-57 particle formation. O'Dowd et al., (O'Dowd et al., 2002) showed that particles are produced from condensable iodine-58 containing vapours at a costal location (Mace Head in Ireland). Recent studies have demonstrated that iodic acid (HIO3) 59 dominates the charged iodine cluster formation, and drives the bursts of freshly-formed particles in coastal regions 60 (He et al., 2020; Sipilä et al., 2016). This process thereby may enhance cloud condensation nuclei formation, affecting 61 climate both directly and indirectly (Saiz-Lopez et al., 2012; Simpson et al., 2015).

62

63 Understanding iodine chemistry and iodine driven new-particle formation requires high time resolution, high 64 sensitivity, and simultaneous measurements of iodine species. However, this has been a long-standing challenge due 65 to their low abundance and short atmospheric lifetimes. Previous studies have achieved detection of relatively more 66 abundant molecular iodine (I2), iodine monoxide (IO), and iodine dioxide (OIO) via optical spectroscopy, such as 67 differential optical absorption (Leigh et al., 2010), cavity ring-down (Bitter et al., 2005), cavity enhanced absorption 68 (Vaughan et al., 2008), laser-induced fluorescence (Dillon et al., 2006), and resonance fluorescence (Gómez Martín 69 et al., 2011). The spectroscopic techniques are invaluable; however, their very specificity limits them to the detection 70 of a few iodine compounds, and they are less sensitive to other iodine species that have congested or broad absorption 71 cross sections such as hypoiodous acid (HOI) and iodic acid (HIO<sub>3</sub>).

72

73 Another commonly used technique is mass spectrometry; it has a fast response time and a low detection limit, but 74 extra calibration efforts are needed for the quantification of the detection sensitivity. For example, photoionization 75 (Gómez Martín et al., 2013) and chemical ionization mass spectrometry (CIMS) have been employed to detect a suite 76 of halogen species. Reagent ions used with CIMS include: SF5<sup>-</sup> for HCl and ClONO<sub>2</sub> (Marcy et al., 2004); iodide (I<sup>-</sup>) 77 for atmospheric chlorine and bromine species such as ClNO2, Cl2, ClO, BrO, and BrCl (Kercher et al., 2009; Lee et 78 al., 2018; Tham et al., 2016); superoxide ( $O_2^{-}$ ) for molecular iodine ( $I_2$ ) (Finley and Saltzman, 2008); and both nitrate 79 (NO<sub>3</sub><sup>-</sup>) (Sipilä et al., 2016) and protonated water (H<sub>3</sub>O<sup>+</sup>) (Pfeifer et al., 2020) for HIO<sub>3</sub>. The nitrate-CIMS and H<sub>3</sub>O<sup>+</sup>-80 CIMS suffer from the limited analyte affinity to the reagent ions. The iodide-CIMS can effectively measure chlorine 81 and bromine species, but it is not suitable to detect iodine species due to the ambiguity in peak identification.





#### 82

83 Bromide ion (Br) exhibits an affinity to a wide spectrum of iodine containing species. Br-CIMS has been routinely 84 used to measure chlorine species (Lawler et al., 2011), HO<sub>2</sub> radicals (Sanchez et al., 2016), organic vapors and sulfuric 85 acid (Rissanen et al., 2019), and nitric acid (Wang et al., 2020). Like chlorine species, iodine species are known to 86 cluster with bromide ions via halogen (or hydrogen) bonds; as such, here we explore using the Br-CIMS to measure 87 gas-phase iodine species and sulfuric acid simultaneously at concentrations relavant to the marine boundary layer. In 88 this study, we demonstrate the detection of various gas-phase inorganic iodine species with the Br-CIMS and explore 89 the effect of relative humidity (RH) on that detection. We then quantify the sensitivities of several gas-phase halogen 90 species via inter-method calibration, offline calibration, and quantum chemical calculations. Finally, we compare the 91 performance of Br-MION-CIMS and Br-FIGAERO-CIMS and show that both of them are well-suited for iodine

- 92 species measurement in the atmosphere.
- 93

#### 94 2 Methodology

# 95 2.1 The CLOUD facility

96 We conducted measurements and instrument inter-comparison at the CERN CLOUD facility, a 26.1 m<sup>3</sup> 97 electropolished stainless-steel chamber that enables new-particle formation experiments simulating the typical range 98 of tropospheric conditions with scrupulous cleanliness and minimal contamination (Duplissy et al., 2016; Kirkby et 99 al., 2011). The CLOUD chamber is mounted in a thermal housing, capable of keeping temperature constant in a range 100 of -65 °C and +100 °C with  $\pm 0.1$  °C precision (Dias et al., 2017) and relative humidity commonly between < 0.5% 101 and 80%. Photochemical processes are driven by different light sources, including four 200 W Hamamatsu Hg-Xe 102 lamps with significant spectral irradiance between 250 and 450 nm, and an array of 48 green light LEDs at 528 nm 103 with adjustable optical power up to 153 W. Ion-induced nucleation under different ionization levels is simulated with 104 a combination of electric fields (electrodes at ±30 kV at top and bottom of the chamber) which can be turned on to 105 rapidly scavenge smaller ions, and a high-flux beam of 3.6 GeV pions ( $\pi^+$ ) which enhances ion production when turned 106 on. Mixing is accelerated with magnetically coupled fans mounted at the top and bottom of the chamber. The 107 characteristic gas mixing time in the chamber during experiments is a few minutes. The loss rate of condensable vapors 108 onto the chamber wall is comparable to the condensation sink in pristine boundary layer environments (e.g.  $2.2 \times 10^{-3}$ 109 s<sup>-1</sup> for H<sub>2</sub>SO<sub>4</sub> at 5 °C). To avoid a memory effect between different experiments, the chamber is periodically cleaned 110 by rinsing the walls with ultra-pure water and heating to 100 °C for at least 24 hours, ensuring extremely low 111 contaminant levels of sulfuric acid ( $< 5 \times 10^4$  cm<sup>-3</sup>) and total organics (< 150 ppty) (Kirkby et al., 2016; Schnitzhofer 112 et al., 2014). The CLOUD gas system is also built to the highest technical standards of cleanliness and performance. 113 The dry air supply for the chamber is provided by cryogenic oxygen (Messer, 99.999 %) and cryogenic nitrogen 114 (Messer, 99.999 %) mixed at the atmospheric ratio of 79:21. Ultrapure water vapor, ozone and other trace gases can 115 be precisely added to attain mixing ratios at different levels. 116

#### 117 2.2 Br-MION-CIMS





118 We measured gas-phase iodine species with a bromide chemical ionization atmospheric pressure interface time-of-119 flight mass spectrometer (Junninen et al., 2010) coupled with a Multi-scheme chemical IONization inlet (Br-MION-120 CIMS) (Rissanen et al., 2019). The Br-MION inlet consists of an electrically grounded 24 mm inner diameter stainless 121 steel flow tube, attached to an ion source. For the CLOUD measurements, the length of the sampling inlet was ~1.5 122 m and was designed to be in a laminar flow with a fixed total flow rate of 20 standard liters per minute (slpm). An ion 123 filter, operated with positive and negative voltage, was placed at the front of the inlet to filter out any ions in the 124 sample air prior to ion-molecule reaction chamber in the inlet. The reagent ions, bromide (Br) and the bromide-water 125 cluster (H<sub>2</sub>O·Br<sup>-</sup>), were produced by feeding 25 standard milliliters per minute (mlpm) of nitrogen (N<sub>2</sub>) flow through 126 a saturator containing dibromomethane (CH<sub>2</sub>Br<sub>2</sub>; > 99.0 %, Tokyo Chemical Industry) into the ion source, where the 127 reagent was ionized by soft X-ray radiation. The resulting ions were then accelerated by a 2500 V ion accelerator array 128 and focused by a 250 V ion deflector into the laminar sampling flow of the inlet via a 5 mm orifice. A small counter 129 flow (~40 mlpm) was applied through the orifice to prevent any mixing of the electrically neutral reagent vapor with 130 the sampling flow. The details of the inlet design, setup, and operation are described in Rissanen et al., 2019.

131

### 132 2.3 Br-FIGAERO-CIMS

133 We also measured both the gas- and particle-phase compositions via thermal desorption using a bromide chemical 134 ionization time-of-flight mass spectrometer equipped with a Filter Inlet for Gases and AEROsols (Br-FIGAERO-135 CIMS) (Lopez-Hilfiker et al., 2014). FIGAERO is a manifold inlet for a CIMS with two operating modes. In the 136 sampling mode, gases are directly sampled into a 150 mbar ion-molecule reactor, using coaxial core sampling to 137 minimize their wall losses in the sampling line. The total flow is maintained at 18.0 slpm and the core flow at 4.5 slpm; 138 the CIMS samples at the center of the core flow with a flow rate at ~1.6 slpm. Concurrently, particles are collected on 139 a PTFE filter via a separate dedicated port with a flow rate of 6 slpm. In the desorption mode, the filter is automatically 140 moved into a pure N2 gas stream flowing into the ion molecule reactor, while the N2 is progressively heated upstream 141 of the filter to evaporate the particles via temperature programmed desorption. Analytes are then chemically ionized 142 by Br and extracted into a mass spectrometer.

143

### 144 2.4 CE-DOAS

145 For the quantitative measurement of gas-phase molecular iodine (I2), we deployed a Cavity Enhanced Differential 146 Optical Absorption Spectroscopy instrument (CE-DOAS) (Meinen et al., 2010). CE-DOAS determines concentrations 147 of trace gases from the strength of differential spectral features in a reference spectrum. The accuracy of the method 148 is ultimately determined by the uncertainty of the respective absorption cross sections. It is thus an absolute method 149 and does not depend on an instrument specific detection efficiency. To maximize the measurement sensitivity towards 150 I<sub>2</sub>, we used a setup optimized for the green wavelength range (508-554 nm), where I<sub>2</sub> exhibits strong differential 151 absorption features. The measurement light is provided by a green light emitting diode (LED Engin). Spectral 152 dispersion is established with a Czerny-Turner grating spectrometer (Princeton Instruments Acton 150), resulting in 153 an optical resolution of 0.73 nm full width at half maximum at 546 nm. Intensities are monitored with a CCD detector 154 (charge-coupled device, Princeton Instruments PIXIS400B) cooled to -70 °C. Highly reflective mirrors (Advanced





- 155 Thin Films) enhance the 1 m mirror separation to an effective optical path length of 15-23 km. The effective spectral 156 mirror reflectivity was established by comparing light intensity spectra in the presence of N<sub>2</sub> and He (Washenfelder 157 et al., 2008). The abundance of trace gases is then determined by comparing spectra of chamber air relative to reference 158 spectra recorded with ultrapure N<sub>2</sub> without I<sub>2</sub>. Chamber air is drawn into the cavity with a constant flow rate of 1 slpm. 159 Variations of the sampling flow did not result in changes in measured I2 concentrations, indicating that photolysis 160 from the measurement light within the instrument was negligible. The following absorbers were included in the fit: I2 161 (Spietz et al., 2006), NO<sub>2</sub> (Vandaele et al., 1998), H<sub>2</sub>O (Rothman et al., 2010), O<sub>2</sub>-O<sub>2</sub> collision-induced absorption 162 (Thalman and Volkamer, 2013), and a polynomial of sixth order. The setup allowed a 1-minute detection limit of 25 163 pptv, or 8 pptv for integration times of 10 minutes, respectively. Periodic automated recordings of N<sub>2</sub> reference spectra 164 were recorded to ensure baseline stability. The optical path length at the time of measurement was continuously 165 confirmed for consistency by the measurement of the O2-O2 collision-induced absorption and H2O column in the same 166 analysis window. The overall systematic accuracy for the  $I_2$  time series is estimated to be 20 %, never better than the 167 detection limit, resulting from the uncertainty in cross sections and the stability of the baseline.
- 168

### 169 2.5 Quantum chemical calculations

170 We used quantum chemical calculations to estimate the cluster formation enthalpy of halogen containing species and 171 bromide ions. The initial conformer sampling was performed using the Spartan '14 program. The cluster geometry 172 was then optimized using density function theory methods at the @B97X-D/aug-cc-pVTZ-PP level of theory (Chai 173 and Head-Gordon, 2008; Kendall et al., 1992). Iodine pseudopotential definitions were taken from the EMSL basis 174 set library (Feller, 1996). Calculations were carried out using the Gaussian 09 program (Frisch et al., 2010). An 175 additional coupled-cluster single-point energy correction was carried out on the lowest energy geometry to calculate 176 the final cluster formation enthalpy. The coupled-cluster calculation was performed at the DLPNO-CCSD(T)/def2-177 QZVPP level using the ORCA program ver. 4.0.0.2 (Neese, 2012; Riplinger and Neese, 2013). In Table 1 we present 178 calculated cluster formation enthalpies based on the optimized geometries.

179

#### 180 3 Results and Discussion

#### 181 3.1 Detection of gas-phase inorganic species by Br-MION-CIMS

182 We show in Fig. 1 the selected inorganic species observed with the Br-MION-CIMS during an iodine new-particle 183 formation experiment in the CLOUD chamber. The peak identities are indicated in the labels. Observed species 184 include I2 and its various oxidation products. There are also a few other halogen-containing inorganic species such as 185 Cl<sub>2</sub>, ICl and IBr, likely coming from the impurities in the iodine source. Non-halogen inorganic species such as H<sub>2</sub>SO<sub>4</sub> 186 can also react with bromide ion and are detected. Due to the large negative mass defect of the bromine and iodine 187 atoms, and the high resolution (~10000 Th Th<sup>-1</sup>) of the mass spectrometer, the peaks can be unambiguously 188 distinguished and identified in the mass spectrum. As shown in the lower panel of Fig. 1, most of the iodine-containing 189 species appear as a single peak in the unit mass range, except for  $HIO_2$ .<sup>79</sup>Br (m/z = 238.82), which overlaps with the 190 reagent ion cluster ( $^{79}Br_2^{81}Br$ )<sup>-</sup> (m/z = 238.75). 191





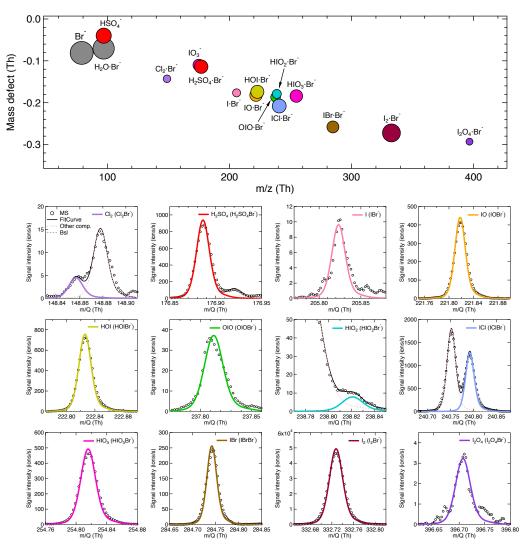




Figure 1. Gas-phase inorganic species measured with the Br-MION-CIMS. Mass defect (difference of exact mass to integer mass) versus m/z of gas-phase halogen species and sulfuric acid during an iodine new-particle formation experiment at 69 % relative humidity and -10 °C. Ions shown here are either clustered with or formed via proton transfer to a bromide ion. The area of the markers is proportional to the logarithm of the signal (counts per second). Shown in the lower panel are the high-resolution single peak fits for species in the mass defect plot in the upper panel.

198

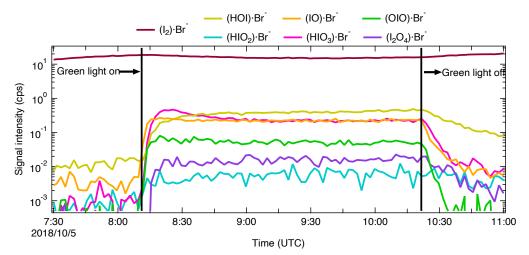
The iodine new-particle formation experiments were conducted under experimental conditions typically found in the high-latitude marine boundary layer, with a temperature of -10 °C and a relative humidity of 69 %. As illustrated in Fig. 2, a typical experiment started with illumination of the chamber at constant  $I_2$  (~60 pptv) using the green light to photolytically produce I atoms. The subsequent reactions of I and ~ 40 ppbv O<sub>3</sub> led to the formation of various oxidized iodine species within a few minutes. The most prominent species we observed from these experiments were IO, HOI and HIO<sub>3</sub>, with lower but significant levels of OIO, HIO<sub>2</sub>, and  $I_2O_4$ . Among these iodine oxides, IO rose the most





rapidly; this is consistent with the first-generation production of IO from the I + O<sub>3</sub> reaction. After a few steps of radical reactions, OIO, HIO<sub>2</sub> and HIO<sub>3</sub> reached steady state almost simultaneously. The only observed iodine oxide dimer was I<sub>2</sub>O<sub>4</sub> in this event, while I<sub>2</sub>O<sub>2</sub>, I<sub>2</sub>O<sub>3</sub> and I<sub>2</sub>O<sub>5</sub> were below the detection limit of both mass spectrometers. A noticeable dip in the HIO<sub>3</sub> traces a few minutes after the onset of the reactions is likely due to the participation of HIO<sub>3</sub> in new-particle formation, resulting in an extra loss term and a lower steady-state concentration. When we turned off the green light, the production of I radicals stopped and iodine species decayed.

211



212

Figure 2. Evolution of selected iodine species during a typical run. The experiment was performed at 60 pptv I<sub>2</sub>, 40 ppbv O<sub>3</sub>, 69 %
 relative humidity and -10 °C. The oxidized iodine species start to appear soon after switching on the green light at 08:11, 05 October
 2018. The I atom production was halted at 10:21, 05 October 2018 by switching off the green light, and the concentration of oxidized iodine species decayed away afterwards. All species are color-coded in the same way as in Fig. 1.

217

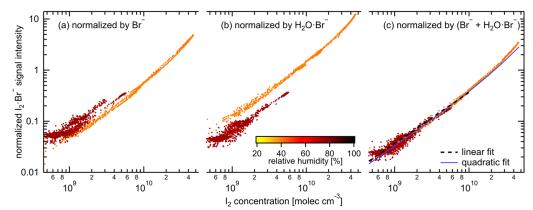
# 218 3.2 Relative humidity dependence

219 Water molecules can cluster with I to form  $H_2O I$  in the iodide CIMS. This enhances the instrument sensitivities for 220 small molecules (i.e. inorganics) and reduces them for large molecules (i.e. organics) (Lee et al., 2014). To investigate 221 the role of water concentration in the sensitivity of the Br-MION-CIMS, we varied the relative humidity (RH) from 222 40 % to 80 % at a constant temperature of -10 °C. We show in Fig. 3 the correlation of I<sub>2</sub> time series from the Br-223 MION-CIMS and the CE-DOAS throughout the experiment. During the RH transition, the ratio of the two reagent 224 ions, Br and H2O Br, changed in the Br-MION-CIMS. As shown in Fig. 3 (a) and (b), using either reagent ion alone 225 for I<sub>2</sub> normalization results in discrepancies in recovered I<sub>2</sub> concentrations at different RH. However, if we use the 226 sum of these two reagent ions ( $Br + H_2O \cdot Br$ ) for normalization, the humidity effect vanishes, as shown in the Fig. 3 227 (c). This suggests that the quantitative detection of I<sub>2</sub> molecules is robust and independent of RH, as long as a proper 228 normalization method is used for the Br-MION-CIMS. Furthermore, we have also carried out the HOI calibration and 229 used the same normalization method as described in section 3.3.4. During the calibration, we varied the water content 230 in the calibrator to vary OH concentrations. A good correlation between the modeled HOI concentrations and the 231 measured HOI signals also indicates that the different H<sub>2</sub>O concentrations in the system do not affect the HOI detection. 232 This assertion may also be applicable to other molecules, but further confirmation is needed.





233



l<sub>2</sub> concentration [molec cm<sup>-5</sup>]
 Figure 3. Signal normalization methods for the Br-MION-CIMS. The charger ions in the ion source of Br-MION-CIMS are Br-and H<sub>2</sub>O·Br<sup>-</sup> (both <sup>79</sup>Br and <sup>81</sup>Br). Their abundance depends both on the instrument tuning and the absolute humidity of the sampled flow. The normalization of the l<sub>2</sub>·Br<sup>-</sup> signal by only Br<sup>-</sup> (a) or H<sub>2</sub>O·Br<sup>-</sup> (b) does not compensate for the humidity effect. Using the sum of Br<sup>-</sup> and H<sub>2</sub>O·Br<sup>-</sup> (c) for normalization yields a tight correlation to the true I<sub>2</sub> as measured by CE-DOAS, independent of the humidity. The black dashed line and blue solid curve indicate the fitted linear and quadratic calibration curves, respectively.

#### 241 3.3 Quantification of gas-phase inorganic species

# 242 3.3.1 I<sub>2</sub> calibration using the CE-DOAS

243 As shown in Fig. 3, we use the accurate I<sub>2</sub> time series measured with the CE-DOAS to calibrate normalized I<sub>2</sub> signals 244 in the Br-MION-CIMS. The I<sub>2</sub> concentrations used for the calibration span approximately 2 orders of magnitude, 245 reaching up to 4.6×10<sup>10</sup> molec cm<sup>-3</sup>. A linear fit, limited to I<sub>2</sub> concentrations smaller than 10<sup>10</sup> molec cm<sup>-3</sup>, establishes 246 the calibration factor  $[I_2] = 2.7 \times 10^{10}$  molec cm<sup>-3</sup> ×  $I_2$ .<sup>79</sup>Br'/(<sup>79</sup>Br' + H<sub>2</sub>O.<sup>79</sup>Br'). We also use a quadratic fit to establish 247 the calibration for the entire range of concentrations encountered during whole campaign (solid line in Fig. 3 (c)); two 248 curves agree well. The CE-DOAS I<sub>2</sub> detection limit is  $6.3 \times 10^8$  molec cm<sup>-3</sup> (25 pptv) for a 1 min integration time, and 249 the total systematic uncertainty is estimated to be 20%. Deviations between both time series are generally smaller than 250 10% (25 and 75 percentile 0.88 and 1.03, respectively). These small differences are consistent with incomplete 251 homogeneity of iodine concentrations in the chamber and the different sampling positions of CE-DOAS and Br-252 MION-CIMS.

253

### 254 3.3.2 I<sub>2</sub> calibration using a permeation tube

We used an iodine permeation tube (VICI Metronic) as a source for offline calibration. The permeation tube was encased within an electronically controlled heating mantle ( $80 \sim 140 (\pm 2)$  °C) to allow for adjustable yet steady iodine permeation rates. The iodine permeation device was run continuously for at least 72 hours before any calibration experiments to ensure that a complete equilibrium was reached in the system. We then confirmed the robustness of the permeation device by the constant I<sub>2</sub> signal measured with Br-MION-CIMS for over 24 hours.

260

To determine the permeation rate of I<sub>2</sub>, we trapped iodine in n-hexane at cryogenic temperatures in an all-glass apparatus, following the method described in Chance et al. (Chance et al., 2010). We initially filled the absorption



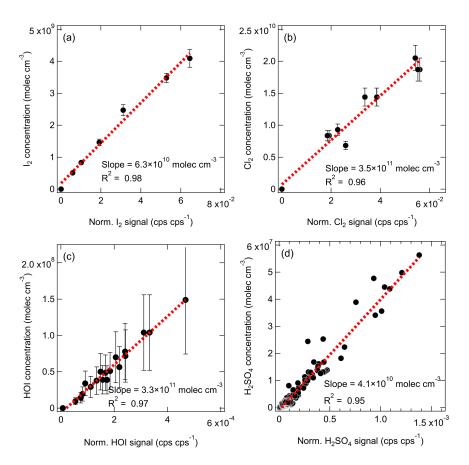


263 glass vessel with 20 ml n-hexane (99.95%, Merck), and then weighed it to determine the combined mass. We then 264 immersed the absorption vessel into a wide-necked Dewar vessel, filled with an acetone/dry ice mixture (at  $-80 \pm$ 265 3 °C). After temperature equilibration, the I<sub>2</sub> molecules, carried by 50 mlpm N<sub>2</sub> flow from the permeation device, were 266 bubbled through the absorption vessel. After a continuous collection for 5 hours, we removed the absorption apparatus 267 from the cooling mixture, and allowed it to warm to room temperature prior to disassembling the setup to prevent any 268 losses of iodine on the tip of the inlet capillary. The absorption vessel was then re-weighed; the mass compared with 269 that prior to absorption was less than 2%, indicating a negligible loss during the trapping process. The  $I_2$  / n-hexane 270 sample solutions were stored at 4 °C for 14 hours before being subjected to analysis. 271 272 We determined the I<sub>2</sub> concentration of the samples using a UV/Vis spectrophotometer (Shimadzu Model UV2450) at 273 a wavelength of 522 nm. We established a calibration curve via a set of I<sub>2</sub> solutions ranging from 270 to 5300 nmol, diluted with n-hexane from a freshly prepared stock solution (0.5 g L<sup>-1</sup>). Repetition of the same analysis after 2 and 7 274 275 days yielded identical results, confirming that the sample solutions were stable at 4 °C. As an alternative analytical 276 approach, we also quantified the I2 concentration in the sample solutions using an inductively coupled plasma mass 277 spectrometer (ICP-MS, Agilent 7800). Before introducing to the ICP-MS, the sample solutions were treated with 278 NaHSO3 water solution (0.100 M), accomplishing efficient hexane-to-water extraction and simultaneous reduction of 279 iodine to iodide (Schwehr et al., 2005) (Agilent Clinical Sample Preparation Guide (v3), ref. ISO 17294-2). The ICP-280 MS results were in good agreement with those from the UV/Vis spectrophotometry. 281 282 We conducted the I<sub>2</sub> trapping and quantification experiments in triplicate with satisfactory reproducibility (standard 283 deviation < 10%). The calculated iodine permeation rate at 50 mlpm N<sub>2</sub> flow and 140 °C oven temperature is 278  $\pm$ 284 12 ng min<sup>-1</sup> (mean  $\pm$  standard deviation). This result was used as the benchmark to estimate temperature-corrected 285 permeation rates according to the formula provided by the permeation tube vendor (VICI Metronic). We checked the 286 validity of the temperature-corrected values by conducting a second iodine absorption experiment in which the iodine 287 permeation tube was kept at 130 °C with an N<sub>2</sub> flow rate of 50 mlpm, and the determined permeation rate agreed 288 within 10% of the calculated value. We then diluted the I2 flow to seven different values and measured the flow with 289 the Br-MION-CIMS. We repeated the calibration five times; we show the data along with a linear fit between the I<sub>2</sub> 290 concentration and normalized I<sub>2</sub> signal in Fig. 4 (a). The slope of the line gives a calibration coefficient of  $6.3 \times 10^{10}$ 291 molec cm<sup>-3</sup> per normalized signal (cps cps<sup>-1</sup>), with R<sup>2</sup> of 0.98 and an overall uncertainty of  $\pm 45\%$ .

292







293

Figure 4. The absolute concentrations (molec cm<sup>-3</sup>) vs. the normalized signals (cps cps<sup>-1</sup>) measured with the Br-MION-CIMS for
 (a) I<sub>2</sub>, (b) Cl<sub>2</sub>, (c) HOI, and (d) H<sub>2</sub>SO<sub>4</sub>. The red dashed lines are the linear fittings. The overall factor 2 systematic scale uncertainty
 on [H<sub>2</sub>SO<sub>4</sub>] is not shown here.

297

### 298 3.3.3 Cl<sub>2</sub> calibration using a permeation tube

299 We used a commercial chlorine permeation tube (VICI Metronic) as a source for offline calibration. We passed a 20 300 mlpm high-purity nitrogen (99.999%) flow at room temperature through a 25 cm long stainless-steel tube (1/2" O.D.) 301 containing the permeation tube. We quantified the permeation rate of Cl2 following a procedure described in a previous 302 study (Finley and Saltzman, 2008). The output of 20 mlpm flow was bubbled into a buffered aqueous potassium iodide 303 solution (2.0 % KI (m/v), prepared in 1.00 mM aqueous phosphate buffer, pH = 7.0) filled in an all-glass two-stage 304 serial absorption apparatus (stage 1 = 100 ml; stage 2 = 50 ml) for 3 hours and kept at room temperature. The Cl<sub>2</sub> 305 oxidized the iodide (I) into iodine ( $I_2$ ) once contacting with the KI absorption solution, and the  $I_2$  further reacted with 306 the excess KI present in the absorption solution to form I3, which can be quantified by UV/Vis-spectrophotometry. 307 We analyzed the resulting sample solutions with an UV/Vis spectrophotometer (Shimadzu Model UV-1800) using 1-308 cm quartz cells at 352 nm, corresponding to the  $I_3$ . We detected no  $I_3$  in the second stage absorption solution, 309 indicating that all the chlorine was quantitatively trapped and rapidly converted to I3- within the first absorption unit.





- The samples were quantified relative to I<sub>3</sub><sup>-</sup> standards in the range of 5 to  $68 \times 10^{-6}$  M, prepared by dilution of a stock obtained by dissolving 174 mg iodine in 200 ml of a solution containing 2 % KI in 1.00 mM aqueous phosphate buffer, pH 7.0. From this calibration curve, we calculated a molar absorptivity of 26,800 L mol<sup>-1</sup> cm<sup>-1</sup>, which is consistent with the values reported in the literature (Finley and Saltzman, 2008; Kazantseva et al., 2002). Samples and standard solutions were re-analyzed after being stored in the dark at room temperature for 24 hours, and the results were within 3 % of those obtained with the fresh solutions. We repeated the absorption experiment, and the calculated chlorine permeation rate at room temperature was  $764 \pm 74$  ng Cl<sub>2</sub> min<sup>-1</sup> (mean ± standard deviation).
- 317

The Cl<sub>2</sub> permeation source was run continuously for 12 hours prior to calibration experiments to ensure complete system equilibrium. A two-stage dilution system similar to the setup of (Gallagher et al., 1997) was set up for diluting the output of the Cl<sub>2</sub> permeation device. The 20 mlpm of N<sub>2</sub> stream emerging from the Cl<sub>2</sub> permeation device (operated at room temperature) was diluted in a stream of 6 slpm of dry N<sub>2</sub>. Then, a small fraction of this mixture (50 to 300 mlpm) was further mixed with the total flow of 25 slpm of N<sub>2</sub> (20 slpm dry N<sub>2</sub> + 5 slpm humidified N<sub>2</sub>) before being sampled by the Br-MION-CIMS. The calibration coefficient for Cl<sub>2</sub> was determined to be  $3.5 \times 10^{11}$  molec cm<sup>-3</sup> per normalized signal (cps cps<sup>-1</sup>) from three separate calibration experiments (Fig 4 (b)), with an accuracy of 30 %.

325

#### 326 3.3.4 HOI calibration using a calibrator

327 We produced a continuous HOI source via the reaction of I2 and hydroxyl radical (OH) in a setup similar to the sulfuric 328 acid (H<sub>2</sub>SO<sub>4</sub>) calibrator (Kürten et al., 2012). The OH was generated by photolyzing H<sub>2</sub>O with a mercury (Hg) lamp 329 at 184.9 nm, whose calibrated intensity was used to estimate the OH concentration. We tested the system by removing 330 the I<sub>2</sub> or OH source from the calibrator, upon which HOI production was undetectable, confirming that any single 331 reactant did not produce HOI. A numerical model was constructed to predict the mean HOI concentration entering the 332 Br-MION-CIMS, which is analogous to the model used for H<sub>2</sub>SO<sub>4</sub> calibration (Kürten et al., 2012). We only included 333 the formation pathway of  $I_2$  + OH to HOI in the model for simplicity; the other pathway of IO + HO<sub>2</sub> was considered 334 minor as IO forms at a relatively slow rate via the reaction of I radical and O<sub>3</sub>. We produced a range of HOI 335 concentrations by varying I<sub>2</sub> and OH concentrations in the calibrator. We show in Fig. 4 (c) the linear correlation 336 between the modeled HOI concentrations and measured HOI signals. The slope of the fit line corresponds to a 337 calibration coefficient of  $3.3 \times 10^{11}$  molec cm<sup>-3</sup> per normalized signal (cps cps<sup>-1</sup>), with an overall uncertainty of  $\pm 55\%$ . The good correlation (R<sup>2</sup> = 0.97) including various H<sub>2</sub>O levels also indicates that H<sub>2</sub>O concentrations did not affect 338 339 the HOI detection.

340

### 341 3.3.5 H<sub>2</sub>SO<sub>4</sub> calibration using a nitrate-CIMS

We derive the H<sub>2</sub>SO<sub>4</sub> calibration coefficient for the Br-MION-CIMS using the absolute H<sub>2</sub>SO<sub>4</sub> concentrations
measured with a pre-calibrated nitrate-CIMS. The calibration protocol of H<sub>2</sub>SO<sub>4</sub> in the nitrate-CIMS has been
described in detail previously (Kürten et al., 2012). The H<sub>2</sub>SO<sub>4</sub> time series used for the inter-method calibration covers
a wide concentration range from less than 5.0×10<sup>4</sup> (detection limit of the nitrate-CIMS) to 6.0×10<sup>7</sup> molec cm<sup>-3</sup>. For
Br-MION-CIMS, although both HSO<sub>4</sub><sup>-</sup> and H<sub>2</sub>SO<sub>4</sub>·Br<sup>-</sup> appear as distinct peaks for sulfuric acid, we only use the





347	normalized $H_2SO_4$ . <sup>79</sup> Br <sup>-</sup> for the inter-calibration, as HSO <sub>4</sub> <sup>-</sup> (m/z = 96.96) has substantial interference from the reagent			
348	ion H <sub>2</sub> O <sup>-79</sup> Br <sup>-</sup> (m/z = 96.93). We show in Fig. 4 (d) the linear fit ([H <sub>2</sub> SO <sub>4</sub> ] = $4.1 \times 10^{10}$ molec cm <sup>-3</sup> × H <sub>2</sub> SO <sub>4</sub> · <sup>79</sup> Br <sup>-</sup> /( <sup>79</sup> Br <sup>-</sup> )			
349 250	+ $H_2O^{.79}Br$ ) – $9.3 \times 10^5$ ) between the two $H_2SO_4$ traces with a correlation coefficient of 0.95. The calculated $H_2SO_4$			
350	calibration coefficient is $4.1 \times 10^{10}$ molec cm <sup>-3</sup> per normalized signal (cps cps <sup>-1</sup> ).			
351				
352	3.3.6 Connecting sensitivity to binding enthalpy			
353	Beyond the species for which we carried out calibrations, there are many more, especially iodine species, that cannot			
354	be directly calibrated due to a lack of authentic standards or generation methods. However, the sensitivity of an iodide-			
355	CIMS towards analytes can be predicted by the cluster binding enthalpy, calculated by relatively simple quantum			
356	chemical methods (Iyer et al., 2016). This holds for the bromide-CIMS as well. In the instrument, ion clusters, formed			
357	from reactions between analytes and reagent ions, are guided and focused by ion optics during transmission to the			
358	detector. The electric forces applied to the clusters enhance their collision energies with carrier gas molecules. If			
359	sufficient energy is transferred during the collision, cluster fragmentation may occur, affecting the instrument			
360	sensitivity for the analytes (Passananti et al., 2019). However, clusters with higher binding enthalpy will be more			
361	easily preserved and detected. Analytes that bind to the reagent ions with enthalpies higher than a critical level are			
362	likely detected at maximum sensitivity (kinetic-limited detection) by the instrument. For example, the calculated			
363	critical enthalpy is 25 kcal mol <sup>-1</sup> for the iodide-CIMS used in Iyer et al. (Iyer et al., 2016) and Lopez-Hilfiker et al.			
364	(Lopez-Hilfiker et al., 2016), calculated at the DLPNO-CCSD(T)/def2-QZVPP//PBE/aug-cc-pVTZ-PP level of theory.			
365				
365 366	For the bromide chemical ionization, there are two types of fragmentation pathways:			
366 367	1) Reversion to the original form of Br <sup>-</sup> and analyte			
366 367 368	1) Reversion to the original form of Br <sup>-</sup> and analyte $X-H \cdot Br^- \rightarrow X-H + Br^-$ (1)			
366 367 368 369	<ol> <li>Reversion to the original form of Br<sup>-</sup> and analyte X-H·Br<sup>-</sup> → X-H + Br<sup>-</sup> (1)</li> <li>Proton transfer from the analyte to Br<sup>-</sup></li> </ol>			
366 367 368 369 370	1) Reversion to the original form of Br and analyte $X-H \cdot Br' \rightarrow X-H + Br'$ (1)2) Proton transfer from the analyte to Br $X-H \cdot Br' \rightarrow X' + HBr$ (2)			
366 367 368 369 370 371	1) Reversion to the original form of Br <sup>-</sup> and analyte $X-H \cdot Br^- \rightarrow X-H + Br^-$ (1)         2) Proton transfer from the analyte to Br <sup>-</sup> $X-H \cdot Br^- \rightarrow X^- + HBr$ (2)         where the X-H is the hydrogen bond donor. An analyte may be expected to be detected at the maximum sensitivity			
366 367 368 369 370 371 372	1) Reversion to the original form of Br and analyte $X-H \cdot Br' \rightarrow X-H + Br'$ (1)         2) Proton transfer from the analyte to Br' $X-H \cdot Br' \rightarrow X' + HBr$ (2)         where the X-H is the hydrogen bond donor. An analyte may be expected to be detected at the maximum sensitivity when the dissociation enthalpy for the first pathway is either a) much higher than the critical enthalpy (dissociation			
366 367 368 369 370 371 372 373	<ol> <li>Reversion to the original form of Br<sup>-</sup> and analyte</li></ol>			
366 367 368 369 370 371 372 373 373	1) Reversion to the original form of Br <sup>-</sup> and analyte $X-H \cdot Br^- \rightarrow X-H + Br^-$ (1) 2) Proton transfer from the analyte to Br <sup>-</sup> $X-H \cdot Br^- \rightarrow X^- + HBr$ (2) where the X-H is the hydrogen bond donor. An analyte may be expected to be detected at the maximum sensitivity when the dissociation enthalpy for the first pathway is either a) much higher than the critical enthalpy (dissociation back to the reactants does not occur), or b) lower than the critical enthalpy, but much higher than that of the second pathway (dissociation back to the reactants would occur, but it is not competitive with the other dissociation channel).			
366 367 368 369 370 371 372 373 374 375	<ol> <li>Reversion to the original form of Br<sup>-</sup> and analyte</li></ol>			
366 367 368 369 370 371 372 373 374 375 376	1) Reversion to the original form of Br <sup>-</sup> and analyte $X-H \cdot Br^- \rightarrow X-H + Br^-$ (1) 2) Proton transfer from the analyte to Br <sup>-</sup> $X-H \cdot Br^- \rightarrow X^- + HBr$ (2) where the X-H is the hydrogen bond donor. An analyte may be expected to be detected at the maximum sensitivity when the dissociation enthalpy for the first pathway is either a) much higher than the critical enthalpy (dissociation back to the reactants does not occur), or b) lower than the critical enthalpy, but much higher than that of the second pathway (dissociation back to the reactants would occur, but it is not competitive with the other dissociation channel). Whether the enthalpy for the second pathway is higher than the critical enthalpy does not directly affect the sensitivity, as both X-H ·Br <sup>-</sup> and X <sup>-</sup> can be measured and counted. Taking H <sub>2</sub> SO <sub>4</sub> as an example, the dissociation enthalpies for			
366 367 368 369 370 371 372 373 374 375 376 377	1) Reversion to the original form of Br <sup>-</sup> and analyte $X-H\cdot Br^- \rightarrow X-H + Br^-$ (1) 2) Proton transfer from the analyte to Br <sup>-</sup> $X-H\cdot Br^- \rightarrow X^- + HBr$ (2) where the X-H is the hydrogen bond donor. An analyte may be expected to be detected at the maximum sensitivity when the dissociation enthalpy for the first pathway is either a) much higher than the critical enthalpy (dissociation back to the reactants does not occur), or b) lower than the critical enthalpy, but much higher than that of the second pathway (dissociation back to the reactants would occur, but it is not competitive with the other dissociation channel). Whether the enthalpy for the second pathway is higher than the critical enthalpy does not directly affect the sensitivity, as both X-H·Br and X <sup>-</sup> can be measured and counted. Taking H <sub>2</sub> SO <sub>4</sub> as an example, the dissociation enthalpies for the first and second pathways are 41.1 and 27.9 kcal mol <sup>-1</sup> , respectively. If some of the H <sub>2</sub> SO <sub>4</sub> ·Br dissociate, they			
366 367 368 369 370 371 372 373 374 375 376 377 378	1) Reversion to the original form of Br and analyte $X-H \cdot Br \rightarrow X-H + Br$ (1) 2) Proton transfer from the analyte to Br $X-H \cdot Br \rightarrow X^- + HBr$ (2) where the X-H is the hydrogen bond donor. An analyte may be expected to be detected at the maximum sensitivity when the dissociation enthalpy for the first pathway is either a) much higher than the critical enthalpy (dissociation back to the reactants does not occur), or b) lower than the critical enthalpy, but much higher than that of the second pathway (dissociation back to the reactants would occur, but it is not competitive with the other dissociation channel). Whether the enthalpy for the second pathway is higher than the critical enthalpy does not directly affect the sensitivity, as both X-H ·Br and X <sup>-</sup> can be measured and counted. Taking H <sub>2</sub> SO <sub>4</sub> as an example, the dissociation enthalpies for the first and second pathways are 41.1 and 27.9 kcal mol <sup>-1</sup> , respectively. If some of the H <sub>2</sub> SO <sub>4</sub> ·Br dissociate, they preferably become HSO <sub>4</sub> <sup>-</sup> and are detectable by the Br-CIMS. We list the cluster formation enthalpies for a selection			
366 367 368 369 370 371 372 373 374 375 376 377 378 379	1) Reversion to the original form of Br <sup>-</sup> and analyte $X-H\cdot Br^- \rightarrow X-H + Br^-$ (1) 2) Proton transfer from the analyte to Br <sup>-</sup> $X-H\cdot Br^- \rightarrow X^- + HBr$ (2) where the X-H is the hydrogen bond donor. An analyte may be expected to be detected at the maximum sensitivity when the dissociation enthalpy for the first pathway is either a) much higher than the critical enthalpy (dissociation back to the reactants does not occur), or b) lower than the critical enthalpy, but much higher than that of the second pathway (dissociation back to the reactants would occur, but it is not competitive with the other dissociation channel). Whether the enthalpy for the second pathway is higher than the critical enthalpy does not directly affect the sensitivity, as both X-H·Br and X <sup>-</sup> can be measured and counted. Taking H <sub>2</sub> SO <sub>4</sub> as an example, the dissociation enthalpies for the first and second pathways are 41.1 and 27.9 kcal mol <sup>-1</sup> , respectively. If some of the H <sub>2</sub> SO <sub>4</sub> ·Br dissociate, they			
366 367 368 369 370 371 372 373 374 375 376 377 378 379 380	<ol> <li>Reversion to the original form of Br and analyte</li></ol>			
366 367 368 369 370 371 372 373 374 375 376 377 378 379 380 381	1) Reversion to the original form of Br and analyte $X-H\cdot Br' \rightarrow X-H + Br'$ (1) 2) Proton transfer from the analyte to Br $X-H\cdot Br' \rightarrow X' + HBr$ (2) where the X-H is the hydrogen bond donor. An analyte may be expected to be detected at the maximum sensitivity when the dissociation enthalpy for the first pathway is either a) much higher than the critical enthalpy (dissociation back to the reactants does not occur), or b) lower than the critical enthalpy, but much higher than that of the second pathway (dissociation back to the reactants would occur, but it is not competitive with the other dissociation channel). Whether the enthalpy for the second pathway is higher than the critical enthalpy does not directly affect the sensitivity, as both X-H·Br' and X' can be measured and counted. Taking H <sub>2</sub> SO <sub>4</sub> as an example, the dissociation enthalpies for the first and second pathways are 41.1 and 27.9 kcal mol <sup>-1</sup> , respectively. If some of the H <sub>2</sub> SO <sub>4</sub> ·Br' dissociate, they preferably become HSO <sub>4</sub> ' and are detectable by the Br-CIMS. We list the cluster formation enthalpies for a selection of halogen containing species in Table 1 and the corresponding cluster dissociation enthalpies in Table 2.			
366 367 368 369 370 371 372 373 374 375 376 377 378 379 380	<ol> <li>Reversion to the original form of Br and analyte</li></ol>			





- 384 I<sub>2</sub> have the lowest calibration coefficients (highest sensitivities); and their coefficients are almost the same for both 385 the online and offline calibrations, with discrepancies well within the systematic uncertainties. Thereby, we conclude 386 that both H<sub>2</sub>SO<sub>4</sub> and I<sub>2</sub> are detected at the maximum sensitivity, suggesting a critical enthalpy not higher than 33.7 387 kcal mol<sup>-1</sup>. We can then infer the sensitivity for other species that are difficult to calibrate by comparing their binding 388 enthalpies to those of the benchmark species. For example, ICl and IBr should have the maximum sensitivity, since 389 the dissociation enthalpies for ICl Br and IBr Br are both much higher than 33.7 kcal mol<sup>-1</sup> (Table 2). Although 390 HIO3 Br has a lower dissociation enthalpy than the critical value, the favored dissociation pathway is proton transfer 391 (the second pathway); HIO<sub>3</sub> can thus be considered as a maximum sensitivity species detectable as IO<sub>3</sub><sup>-</sup> ions after 392 proton transfer. This is consistent with the fact that both HIO<sub>3</sub> Br and IO<sub>3</sub> are detected in Figure 1, so is the case with 393 H<sub>2</sub>SO<sub>4</sub>. We thus assume that HIO<sub>3</sub> has a kinetic calibration coefficient of  $3.8 \times 10^{10}$  molec cm<sup>-3</sup> cps cps<sup>-1</sup>, the value for 394 H<sub>2</sub>SO<sub>4</sub>. However, the lowest dissociation enthalpies of HOI Br and Cl<sub>2</sub> Br are 26.9 and 22.3 kcal mol<sup>-1</sup>, respectively, 395 consistent with their higher calibration coefficients of  $3.3 \times 10^{11}$  and  $3.5 \times 10^{11}$  molec cm<sup>-3</sup> cps cps<sup>-1</sup>. The dissociation 396 enthalpies for IO·Br, OIO·Br, and HIO<sub>2</sub>·Br are 24.5, 23.2, and 29.2 kcal mol<sup>-1</sup>, respectively. We would expect that 397 their sensitives are lower than the maximum sensitivity. Since the dissociation enthalpies for IO Br and OIO Br are 398 between those of HOI Br and Cl2 Br, a similar calibration coefficient of around 3.0×10<sup>11</sup> molec cm<sup>-3</sup> cps cps<sup>-1</sup> could 399 be applied.
- 400

Further, we estimate the detection limit of the calibrated species. The detection limit is defined as the analyte concentration, corresponding to the sum of the mean signal and three times the standard deviations  $(3\sigma)$  of the background fluctuations during a two-hour background measurement. We derive the detection limit of HOI, HIO<sub>3</sub>, I<sub>2</sub>, and H<sub>2</sub>SO<sub>4</sub> to be  $5.8 \times 10^6$ ,  $1.2 \times 10^5$ ,  $6.3 \times 10^5$ , and  $2.0 \times 10^5$  molec cm<sup>-3</sup> (or 0.2, 0.005, 0.03, and 0.008 pptv), respectively, for a 2-min integration time.

406

### 407 3.4 Comparison between Br-MION-CIMS and Br-FIGAERO-CIMS

While Br-MION-CIMS and Br-FIGAERO-CIMS use the same chemical ionization scheme, their designs differ in the ion-molecule reaction chamber (IMR). MION is an atmospheric pressure (1 bar) drift tube; analyte molecules gain an electric charge in an axial laminar flow. FIGAERO is connected to a cone-shaped IMR chamber operated at a reduced pressure (150 mbar); the sample flow is injected into the inlet via an orifice, necessarily causing turbulence and wall interactions in the IMR region. The atmospheric pressure and reduced pressure IMRs are both widely used for trace gas measurements. We thus compare iodine species measurements from Br-MION-CIMS and Br-FIGAERO-CIMS, to better understand the performance and applicability of bromide ionization scheme.

415

#### 416 3.4.1 Signal trend and detection limit

417 We show in Fig. 5 the same new-particle formation event as in Fig. 2, to illustrate the time series for HIO<sub>3</sub> Br, HOI Br,

418 IO·Br<sup>-</sup>, and I<sub>2</sub>·Br<sup>-</sup>, measured with Br-MION-CIMS (red circles) and Br-FIGAERO-CIMS (grey sticks), respectively.

419 Note that the FIGAERO alternates between gas and particle measurements; here we show only the gas-phase signals.

420 Clear and concurrent signals of HIO<sub>3</sub>, HOI, IO, and I<sub>2</sub> are evident from both the Br-MION-CIMS and Br-FIGAERO-

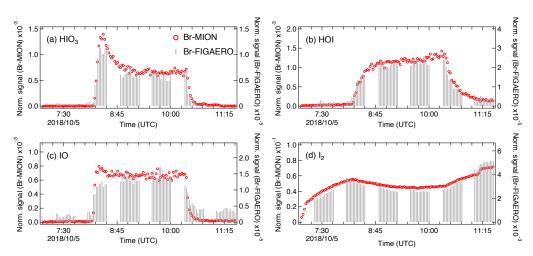




421 CIMS. Prior to the NPF event (08:11), there was no photochemical production and thus virtually no signal of oxidized 422 iodine species in both instruments. The dark reaction of ozone with I2 did not proceed at a significant rate, due to the 423 low rate coefficient and low levels of I<sub>2</sub>. Signals detected during this period are considered as the persistent background, 424 coming from electronic noise or other sources such as the ionizer, carrier flows, or long-term "memory" in the case 425 of the Br-FIGAERO-CIMS. Not surprisingly, the Br-MION-CIMS has a near-zero background for all analytes. For 426 HIO<sub>3</sub> (Fig. 5 (a)), the background signal in the Br-FIGAERO-CIMS is also negligible; however, IO shows a substantial 427 persistent background (Fig. 5 (c)) in the Br-FIGAERO-CIMS. After the NPF event (10:21), the photochemical 428 production of oxidized iodine species was terminated and vapor concentrations decayed exponentially due to dilution 429 and losses to chamber walls. The I<sub>2</sub> signal increases after the event termination because it is no longer photolyzed. 430 431 Applying the calibration coefficients, we convert the Br-MION-CIMS signals to absolute concentrations, and 432 subsequently correlate them with signals measured with Br-FIGAERO-CIMS. We then estimate the tentative detection 433 limits for HOI and HIO3 in the Br-FIGAERO-CIMS to be 3.3×107 and 5.1×106 molec cm<sup>-3</sup> (versus 5.8×106 and 1.2×105 434 molec cm<sup>3</sup> in the Br-MION-CIMS), respectively, at  $3\sigma$  of the background signal for a 2-min integration time during

435 a two-hour period; they are in general one order of magnitude higher than those in the Br-MION-CIMS. This is in line

- 436 with the higher background signals observed in the Br-FIGAERO-CIMS. We are unable to estimate the I<sub>2</sub> detection
- 437 limit in the Br-FIGAREO-CIMS, due to a lack of I<sub>2</sub> background measurement; but Br-FIGAERO-CIMS can and did
- 438 detect  $I_2$  at the low pptv level with good fidelity.
- 439



- 440
- 441 442

Figure 5. Signal comparison of selected iodine species measured with Br-MION-CIMS and (gas-phase) Br-FIGAERO-CIMS, respectively, during the same iodine new-particle formation experiment shown in Fig. 2.

443

# 444 3.4.2 Rise and decay time constants

445 In order to quantitatively compare the performance of the two types of IMRs, we set the initiation (08:11) and 446 termination (10:21) of NPF event as t = 0, and fit the rise and decay rates of HIO<sub>3</sub>, HOI and IO in both instruments, 447 respectively.





449	When we initiated the photochemistry, oxidized iodine species rapidly built up toward an asymptote; we thus fit their
450	time-series signals individually with a four-parameter sigmoid function using a least-squares fitting algorithm:

451

448

# $S_i(t) = a + (b - a) / (1 + e^{(-(t - t_{app})/c)})$ (3)

where *a* and *b* represent the background and asymptotic value of the sigmoid function respectively, *c* is the exponential time constant of the signal change, and  $t_{app}$  is the time at which the 50 % value between plateau and background is reached (50 % appearance time). As we were unable to separate the time scale of chemical reactions from that of instrument response, we use the time constants to represent the overall response for the purpose of comparing instrument performance.

457

458 We show in Fig. 6 (a), (c), and (e) the rise rate fits of HIO<sub>3</sub>, HOI, and IO, respectively. HIO<sub>3</sub> signals rose with a time 459 constant of 102 s in the Br-MION-CIMS and 108 s in the Br-FIGAERO-CIMS, both with  $t_{app} = \sim 300$  s. The fitted time 460 constants of HOI are slightly longer than those of HIO<sub>3</sub>, with 120 s in the Br-MION-CIMS and 114 s in the Br-461 FIGAERO-CIMS. IO signals stabilized the earliest, thus have the fastest time constants of 48 s in the Br-MION-CIMS 462 and 84 s in the Br-FIGAERO-CIMS. The instrumental differences are small for HIO3 and HOI, but larger for IO. When 463 colliding with the IMR surface, HIO3 condenses irreversibly; it thus makes sense that the Br-MION-CIMS and Br-464 FIGAERO-CIMS signals show the same time constant for HIO3. Semi-volatile HOI, however, can return to the gas 465 phase from the walls depending on the surface coverage of HOI and the vapor concentration. Additionally, the 466 heterogeneous reaction of aqueous iodide (I<sup>-</sup>) and ozone (Carpenter et al., 2013) could also contribute to the emission 467 of HOI from the IMR wall in the FIGAERO. As the evaporation flux is typically a function of the amount analyte on 468 the surface, the buffering effect could degrade the instrument time response upon changes in analyte concentration. 469 Here, however, we did not observe a significant memory effect, likely because the HOI concentration was too low to 470 fully saturate the IMR surface or because any HOI evaporation was suppressed due to an enhanced accommodation 471 coefficient of HOI on the metal surface. We expect IO to be prone to loss on the metal surface due to its radical nature.

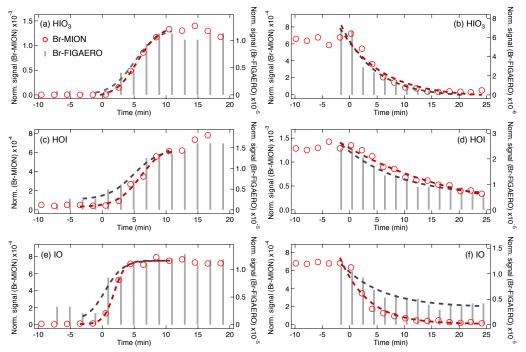
472

473 We also fit the exponential decay time constants of these iodine species to test this interpretation (Fig. 6 (b), (d), and 474 (f)). After photochemistry was terminated at the end of the NPF event, only two sinks drove the vapor concentration 475 decay - dilution and wall loss. Memory effects could also influence the signal time constant. The dilution loss rate 476 was around  $2.1 \times 10^{-4}$  s<sup>-1</sup> (4760 s time constant) for all species in the chamber, determined by the total chamber flow 477 rate and the chamber volume. Wall loss rates, however, vary for species with different diffusion constants. The decay 478 rates of HIO3 are 400 s for the Br-MION-CIMS and 370 s for the Br-FIGAERO-CIMS, much faster than the dilution 479 loss. For comparison, the time constant for H2SO4 vapor loss was 300 s. These time constants are thus consistent with 480 wall loss (around 2.2×10<sup>-3</sup> s<sup>-1</sup>). The IO decay time constant is 294 s for the Br-MION-CIMS and 435 s for the Br-481 FIGAERO-CIMS. The time constant for the Br-MION-CIMS indicates that the decay of IO is also driven by wall loss, 482 so the net flux during this period was thus towards the wall rather than from the wall. Therefore, the difference of IO 483 between instruments may well be attributed to the persistent background from the ionizer of the FIGAERO. The HOI





- 484 signals have longer decay time constants in both instruments of 909 s for the Br-MION-CIMS and 714 s for the Br-
- 485 FIGAERO-CIMS; this may reflect a time constant for depletion of HOI adsorbed to the chamber walls.
- 486



# 487

Figure 6. Rising ((a), (c), and (e)) and decay ((b), (d), and (f)) rate comparison of selected iodine species measured with Br-MION-CIMS and Br-FIGAERO-CIMS, respectively, during the same iodine new-particle formation event in Fig. 2 and 5.

490

# 491 4 Summary and conclusion

492 We confirm in this study that bromide chemical ionization is a suitable technique for the time-resolved, highly 493 sensitive, and simultaneous measurements of iodine species and sulfuric acid. The Br-MION-CIMS shows constant 494 sensitivity throughout the relative humidity range of 40 to 80 % at -10 °C, as long as the sum of the two reagent ions 495 (Br<sup>-</sup> + H<sub>2</sub>O·Br<sup>-</sup>) is used for signal normalization. This demonstrates the applicability of this technique to field 496 measurements in the ambient marine environment.

497

We quantify iodine species and sulfuric acid via offline calibrations (i.e. permeation tube and calibrator) and intermethod calibrations (i.e. CE-DOAS and pre-calibrated nitrate-CIMS); different methods result in consistent calibration coefficients. Further, we calculate the binding enthalpies between the calibrated species and reagent ions, which qualitatively agree with the corresponding calibration coefficients. This indicates that the quantum chemical calculations can be employed along with the calibration experiments to determine the sensitivities for unquantifiable species; more work is required to further establish the correlation between calibration coefficients and binding enthalpies.

505





506	Further, using inter-method and offline calibrations, we estimate the detection limits of HOI, HIO <sub>3</sub> , I <sub>2</sub> , and H <sub>2</sub> SO <sub>4</sub> in
507	Br-MION-CIMS being 5.8×10 <sup>6</sup> , 1.2×10 <sup>5</sup> , 6.3×10 <sup>5</sup> , and 2.0×10 <sup>5</sup> molec cm <sup>-3</sup> , respectively, for a 2-min integration time
508	during a two-hour period. To our knowledge, the simultaneous measurements of various iodine species and sulfuric
509	acid with low detection limits are unprecedented for online techniques. Detection limits for HOI and HIO <sub>3</sub> in the Br-
510	FIGAERO-CIMS are 3.3×10 <sup>7</sup> and 5.1×10 <sup>6</sup> molec cm <sup>-3</sup> , which are in general one order of magnitude higher than those
511	in the Br-MION-CIMS. The signal comparison between the two instruments also shows that the Br-CIMS can be
512	coupled to both the atmospheric pressure and the reduced pressure interfaces for iodine species and sulfuric acid
513	measurements in the marine environment.
514	
515	
516	Data availability. Data available on request from the authors.
517	
518 519	Author Contributions. M.W., XC.H., YJ.T and H.F. wrote the manuscript. XC.H., YJ.T., M.W. and M.Sip. designed the experiments. XC.H., YJ.T. and J.S. carried out the Br-MION-CIMS measurements, M.W., D.C. and
520	V.H. carried out the Br-FIGAERO-CIMS measurements, and H.F. carried out the CE-DOAS measurements. YJ.T.,
521	XC.H., H.F., D.C., J.S. and M.Sim. performed the calibrations. S.I., XC.H. and T.K. carried out the quantum
522	chemical calculations. M.W. performed the comparison analysis of the Br-FIGAERO-CIMS and Br-MION-CIMS.
523	NM.D., T.K., M.R., R.V. and M.Sip. commented on the manuscript. All other co-authors participated in either the
524 525	development and preparations of the CLOUD facility and the instruments, and/or collecting and analyzing the data.
525 526	Competing interests. The authors declare that they have no conflict of interest.
527	
528	Acknowledgement. We thank the European Organization for Nuclear Research (CERN) for supporting CLOUD with
529	important technical and financial resources and for providing a particle beam from the CERN Proton Synchrotron.
530	We also thank Juhani Virkanen and Heini Ali-Kovero for providing assistance in the laboratory analytical experiments.
531 532	Financial support. This research has received funding from the US National Science Foundation (AGS-1531284,
533	AGS-1801574 and AGS-1801280), Academy of Finland (projects: 296628, 328290, Centre of Excellence 1118615)
534	and the European Research Council (ERC) under the European Union's Horizon 2020 research and innovation
535	programme (GASPARCON, grant agreement no. 714621). The FIGAERO-CIMS was supported by an MRI grant for
536 537	the US NSF AGS-1531284 as well as the Wallace Research Foundation.
538	References
539	Bitter, M., Ball, S. M., Povey, I. M. and Jones, R. L.: A broadband cavity ringdown spectrometer for in-situ
540	measurements of atmospheric trace gases, Atmos. Chem. Phys., doi:10.5194/acp-5-2547-2005, 2005.
541 542	Carpenter, L. J., MacDonald, S. M., Shaw, M. D., Kumar, R., Saunders, R. W., Parthipan, R., Wilson, J. and Plane, J. M. C.: Atmospheric iodine levels influenced by sea surface emissions of inorganic iodine, Nat. Geosci.,
543	doi:10.1038/ngeo1687, 2013.
544 545	Chai, J. Da and Head-Gordon, M.: Long-range corrected hybrid density functionals with damped atom-atom dispersion corrections, Phys. Chem. Chem. Phys., doi:10.1039/b810189b, 2008.
546 547 548	Chance, R. J., Shaw, M., Telgmann, L., Baxter, M. and Carpenter, L. J.: A comparison of spectrophotometric and denuder based approaches for the determination of gaseous molecular iodine, Atmos. Meas. Tech., doi:10.5194/amt-3-177-2010, 2010.
549	Dias, A., Ehrhart, S., Vogel, A., Mathot, S., Onnela, A., Almeida, J., Kirkby, J., Williamson, C. and Mumford, S.:





- 550 Temperature uniformity in the CERN CLOUD chamber, Atmos. Meas. Tech., 10, 5075–5088, 2017.
- 551 Dillon, T. J., Tucceri, M. E. and Crowley, J. N.: Laser induced fluorescence studies of iodine oxide chemistry: Part
- 552 II. The reactions of IO with CH3O2, CF3O2 and O3, Phys. Chem. Chem. Phys., doi:10.1039/b611116e, 2006.
- 553 Duplissy, J., Merikanto, J., Franchin, A., Tsagkogeorgas, G., Kangasluoma, J., Wimmer, D., Vuollekoski, H.,
- 554 Schobesberger, S., Lehtipalo, K., Flagan, R. C., Brus, D., Donahue, N. M., Vehkamaki, H., Almeida, J., Amorim,
- 555 A., Barmet, P., Bianchi, F., Breitenlechner, M., Dunne, E. M., Guida, R., Henschel, H., Junninen, H., Kirkby, J.,
- 556 Kurten, A., Kupc, A., Maattanen, A., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Praplan, A. P.,
- 557 Riccobono, F., Rondo, L., Steiner, G., Tome, A., Walther, H., Baltensperger, U., Carslaw, K. S., Dommen, J.,
- 558 Hansel, A., Petaja, T., Sipila, M., Stratmann, F., Vrtala, A., Wagner, P. E., Worsnop, D. R., Curtius, J. and Kulmala,
- 559 M.: Effect of ions on sulfuric acid-water binary particle formation: 2. Experimental data and comparison with QC-
- 560 normalized classical nucleation theory, J. Geophys. Res., 121, 1752-1775, doi:10.1002/2015JD023539, 2016.
- 561 Feller, D.: The role of databases in support of computational chemistry calculations, J. Comput. Chem.,
- 562 doi:10.1002/(SICI)1096-987X(199610)17:13<1571::AID-JCC9>3.0.CO;2-P, 1996.
- 563 Finley, B. D. and Saltzman, E. S.: Observations of Cl2, Br2, and I2 in coastal marine air, J. Geophys. Res. Atmos., 564 doi:10.1029/2008JD010269, 2008.
- 565 Frisch, M. J., Trucks, G. W., Schlegel, H. B., Scuseria, G. E., Robb, M. A., Cheeseman, J. R., Scalmani, G., Barone,
- 566 V., Mennucci, B., Petersson, G. A., Nakatsuji, H., Caricato, M., Li, X., Hratchian, H. P., Izmaylov, A. F., Bloino, J.,
- 567 Zheng, G., Sonnenberg, J. L., Hada, M., Ehara, M., Toyota, K., Fukuda, R., Hasegawa, J., Ishida, M., Nakajima, T.,
- 568 Honda, Y., Kitao, O., Nakai, H., Vreven, T., Montgomery Jr., J. A., Peralta, J. E., Ogliaro, F., Bearpark, M., Heyd, J.
- 569 J., Brothers, E., Kudin, K. N., Staroverov, V. N., Kobayashi, R., Normand, J., Raghavachari, K., Rendell, A.,
- 570 Burant, J. C., Iyengar, S. S., Tomasi, J., Cossi, M., Rega, N., Millam, J. M., Klene, M., Knox, J. E., Cross, J. B.,
- 571 Bakken, V., Adamo, C., Jaramillo, J., Gomperts, R., Stratmann, R. E., Yazyev, O., Austin, A. J., Cammi, R.,
- 572 Pomelli, C., Ochterski, J. W., Martin, R. L., Morokuma, K., Zakrzewski, V. G., Voth, G. A., Salvador, P.,
- 573 Dannenberg, J. J., Dapprich, S., Daniels, A. D., Farkas, ,, Foresman, J. B., Ortiz, J. V, Cioslowski, J. and Fox, D.
- 574 J.: Gaussian09 Revision D.01, Gaussian Inc. Wallingford CT, Gaussian 09 Revis. C.01, 2010.
- 575 Gallagher, M. S., King, D. B., Whung, P. Y. and Saltzman, E. S.: Performance of the HPLC/fluorescence SO2
- 576 detector during the GASIE instrument intercomparison experiment, J. Geophys. Res. Atmos.,
- 577 doi:10.1029/97jd00700, 1997.
- 578 Gómez Martín, J. C., Blahins, J., Gross, U., Ingham, T., Goddard, A., Mahajan, A. S., Ubelis, A. and Saiz-Lopez, 579 A.: In situ detection of atomic and molecular iodine using Resonance and Off-Resonance Fluorescence by Lamp 580 Excitation: ROFLEX, Atmos. Meas. Tech., doi:10.5194/amt-4-29-2011, 2011.
- 581 Gómez Martín, J. C., Gálvez, O., Baeza-Romero, M. T., Ingham, T., Plane, J. M. C. and Blitz, M. A.: On the
- 582 mechanism of iodine oxide particle formation, Phys. Chem. Chem. Phys., doi:10.1039/c3cp51217g, 2013.
- 583 He, X.-C., Iyer, S., Sipilä, M., Ylisirniö, A., Peltola, M., Kontkanen, J., Baalbaki, R., Simon, M., Kürten, A., Tham,
- 584 Y. J. and others: Determination of the Collision Rate Coefficient between Charged Iodic Acid Clusters and Iodic 585 Acid using the Appearance Time Method, Aerosol Sci. Technol., 1–17, 2020.
- 586 Iyer, S., Lopez-Hilfiker, F., Lee, B. H., Thornton, J. A. and Kurtén, T.: Modeling the Detection of Organic and 587 Inorganic Compounds Using Iodide-Based Chemical Ionization, J. Phys. Chem. A, doi:10.1021/acs.jpca.5b09837,
- 588 2016.
- 589 Junninen, H., Ehn, M., Petaja, T., Luosujarvi, L., Kotiaho, T., Kostiainen, R., Rohner, U., Gonin, M., Fuhrer, K.,
- 590 Kulmala, M., Worsnop, D. R., Petäjä, T., Luosujärvi, L., Kotiaho, T., Kostiainen, R., Rohner, U., Gonin, M., Fuhrer,
- 591 K., Kulmala, M. and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion composition,
- 592 Atmos. Meas. Tech., 3(4), 1039–1053, doi:10.5194/amt-3-1039-2010, 2010.
- 593 Kazantseva, N. N., Ernepesova, A., Khodjamamedov, A., Geldyev, O. A. and Krumgalz, B. S.: Spectrophotometric 594 analysis of iodide oxidation by chlorine in highly mineralized solutions, in Analytica Chimica Acta., 2002.
- 595 Kendall, R. A., Dunning, T. H. and Harrison, R. J.: Electron affinities of the first-row atoms revisited. Systematic 596 basis sets and wave functions, J. Chem. Phys., doi:10.1063/1.462569, 1992.
- 597 Kercher, J. P., Riedel, T. P. and Thornton, J. A.: Chlorine activation by N2O5: Simultaneous, in situ detection of 598
- CINO2 and N2O5 by chemical ionization mass spectrometry, Atmos. Meas. Tech., doi:10.5194/amt-2-193-2009,





599 2009.

- 600 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten,
- 601 A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim,
- 602 A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S.,
- 603 Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J.,
- 604 Lovejoy, E. R., Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A.,
- 605 Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A.,
- 606 Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M.,
- 607 Carslaw, K. S., Worsnop, D. R., Baltensperger, U. and Kulmala, M.: Role of sulphuric acid, ammonia and galactic
- 608 cosmic rays in atmospheric aerosol nucleation, Nature, 476, 429 [online] Available from:
- 609 https://doi.org/10.1038/nature10343, 2011.
- 610 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan, C.,
- 611 Almeida, J., Trostl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer, A. K.,
- 612 Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X. M., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin,
- 613 A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M.,
- 614 Kurten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Perakyla, O., Piel,
- 615 F., Petaja, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A. D. D., Riipinen, I., Rissanen, M. P., Rondo, L.,
- 616 Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipila, M., Steiner, G., Stozhkov, Y., Stratmann, F.,
- 617 Tome, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M.,
- 618 Ye, P. L., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M.,
- 619 Carslaw, K. S., Curtius, J., Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E.,
- 620 Wimmer, D., Winkler, P. M., Ye, P. L., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R.,
- 621 Baltensperger, U., Kulmala, M., Carslaw, K. S. and Curtius, J.: Ion-induced nucleation of pure biogenic particles, 622 Nature, 533(7604), 521-526, doi:10.1038/nature17953, 2016.
- 623 Koenig, T. K., Baidar, S., Campuzano-Jost, P., Cuevas, C. A., Dix, B., Fernandez, R. P., Guo, H., Hall, S. R.,
- 624 Kinnison, D., Nault, B. A., Ullmann, K., Jimenez, J. L., Saiz-Lopez, A. and Volkamer, R.: Quantitative detection of 625 iodine in the stratosphere, Proc. Natl. Acad. Sci. U. S. A., doi:10.1073/pnas.1916828117, 2020.
- 626 Kurten, A., Rondo, L., Ehrhart, S. and Curtius, J.: Calibration of a Chemical Ionization Mass Spectrometer for the 627 Measurement of Gaseous Sulfuric Acid, J. Phys. Chem. A, 116, 6375-6386, doi:Doi 10.1021/Jp212123n, 2012.
- 628 Lawler, M. J., Sander, R., Carpenter, L. J., Lee, J. D., Von Glasow, R., Sommariva, R. and Saltzman, E. S.: HOCI 629 and Cl2 observations in marine air, Atmos. Chem. Phys., doi:10.5194/acp-11-7617-2011, 2011.
- 630 Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurten, T., Worsnop, D. R. and Thornton, J. A.: An iodide-adduct high-
- 631 resolution time-of-flight chemical-ionization mass spectrometer: application to atmospheric inorganic and organic 632 compounds, Env. Sci Technol, 48(11), 6309-6317, doi:10.1021/es500362a, 2014.
- 633 Lee, B. H., Lopez-Hilfiker, F. D., Veres, P. R., McDuffie, E. E., Fibiger, D. L., Sparks, T. L., Ebben, C. J., Green, J. 634
- R., Schroder, J. C., Campuzano-Jost, P., Iyer, S., D'Ambro, E. L., Schobesberger, S., Brown, S. S., Wooldridge, P.
- 635 J., Cohen, R. C., Fiddler, M. N., Bililign, S., Jimenez, J. L., Kurtén, T., Weinheimer, A. J., Jaegle, L. and Thornton,
- 636 J. A.: Flight Deployment of a High-Resolution Time-of-Flight Chemical Ionization Mass Spectrometer:
- 637 Observations of Reactive Halogen and Nitrogen Oxide Species, J. Geophys. Res. Atmos.,
- 638 doi:10.1029/2017JD028082, 2018.
- 639 Leigh, R. J., Ball, S. M., Whitehead, J., Leblanc, C., Shillings, A. J. L., Mahajan, A. S., Oetjen, H., Lee, J. D., Jones, 640 C. E., Dorsey, J. R., Gallagher, M., Jones, R. L., Plane, J. M. C., Potin, P. and McFiggans, G.: Measurements and
- 641 modelling of molecular iodine emissions, transport and photodestruction in the coastal region around Roscoff,
- 642 Atmos. Chem. Phys., doi:10.5194/acp-10-11823-2010, 2010.
- 643 Lopez-Hilfiker, F. D., Mohr, C., Ehn, M., Rubach, F., Kleist, E., Wildt, J., Mentel, T. F., Lutz, A., Hallquist, M.,
- 644 Worsnop, D. and Thornton, J. A.: A novel method for online analysis of gas and particle composition: description
- 645 and evaluation of a Filter Inlet for Gases and AEROsols (FIGAERO), Atmos. Meas. Tech., 7(4), 983-1001, 646 doi:10.5194/amt-7-983-2014, 2014.
- 647 Lopez-Hilfiker, F. D., Iyer, S., Mohr, C., Lee, B. H., D'ambro, E. L., Kurtén, T. and Thornton, J. A.: Constraining
- 648 the sensitivity of iodide adduct chemical ionization mass spectrometry to multifunctional organic molecules using
- 649 the collision limit and thermodynamic stability of iodide ion adducts, Atmos. Meas. Tech., doi:10.5194/amt-9-1505-





#### **650** 2016, 2016.

- 651 Marcy, T. P., Fahey, D. W., Gao, R. S., Popp, P. J., Richard, E. C., Thompson, T. L., Rosenlof, K. H., Ray, E. A.,
- 652 Salawitch, R. J., Atherton, C. S., Bergmann, D. J., Ridley, B. A., Weinheimer, A. J., Loewenstein, M., Weinstock, E.
- 653 M. and Mahoney, M. J.: Quantifying Stratospheric Ozone in the Upper Troposphere with in Situ Measurements of
- 654 HCl, Science (80-. )., doi:10.1126/science.1093418, 2004.
- 655 McFiggans, G., Coe, H., Burgess, R., Allan, J., Cubison, M., Alfarra, M. R., Saunders, R., Saiz-Lopez, A., Plane, J.
- 656 M. C., Wevill, D. J., Carpenter, L. J., Rickard, A. R. and Monks, P. S.: Direct evidence for coastal iodine particles
- from Laminaria macroalgae Linkage to emissions of molecular iodine, Atmos. Chem. Phys., doi:10.5194/acp-4 701-2004, 2004.
- 659 Meinen, J., Thieser, J., Platt, U. and Leisner, T.: Technical Note: Using a high finesse optical resonator to provide a
- long light path for differential optical absorption spectroscopy: CE-DOAS, Atmos. Chem. Phys., doi:10.5194/acp-10-3901-2010, 2010.
- 662 Neese, F.: The ORCA program system, Wiley Interdiscip. Rev. Comput. Mol. Sci., doi:10.1002/wcms.81, 2012.
- O'dowd, C. D., Jimenez, J. L., Bahreini, R., Flagan, R. C., Seinfeld, J. H., Hämerl, K., Pirjola, L., Kulmala, M. and
   Hoffmann, T.: Marine aerosol formation from biogenic iodine emissions, Nature, doi:10.1038/nature00775, 2002.
- Passananti, M., Zapadinsky, E., Zanca, T., Kangasluoma, J., Myllys, N., Rissanen, M. P., Kurtén, T., Ehn, M.,
- Attoui, M. and Vehkamäki, H.: How well can we predict cluster fragmentation inside a mass spectrometer?, Chem.
  Commun., doi:10.1039/c9cc02896j, 2019.
- Pfeifer, J., Simon, M., Heinritzi, M., Piel, F., Weitz, L., Wang, D., Granzin, M., Müller, T., Bräkling, S., Kirkby, J.,
   Curtius, J. and Kürten, A.: Measurement of ammonia, amines and iodine compounds using protonated water cluster
- 670 chemical ionization mass spectrometry, Atmos. Meas. Tech., doi:10.5194/amt-13-2501-2020, 2020.
- Prados-Roman, C., Cuevas, C. A., Fernandez, R. P., Kinnison, D. E., Lamarque, J. F. and Saiz-Lopez, A.: A
  negative feedback between anthropogenic ozone pollution and enhanced ocean emissions of iodine, Atmos. Chem.
  Phys., doi:10.5194/acp-15-2215-2015, 2015.
- 674 Raso, A. R. W., Custard, K. D., May, N. W., Tanner, D., Newburn, M. K., Walker, L., Moore, R. J., Huey, L. G.,
- Alexander, L., Shepson, P. B. and Pratt, K. A.: Active molecular iodine photochemistry in the Arctic, Proc. Natl.
  Acad. Sci. U. S. A., doi:10.1073/pnas.1702803114, 2017.
- Riplinger, C. and Neese, F.: An efficient and near linear scaling pair natural orbital based local coupled cluster
   method, J. Chem. Phys., doi:10.1063/1.4773581, 2013.
- 679 Rissanen, M. P., Mikkilä, J., Iyer, S. and Hakala, J.: Multi-scheme chemical ionization inlet (MION) for fast
- switching of reagent ion chemistry in atmospheric pressure chemical ionization mass spectrometry (CIMS)
   applications, Atmos. Meas. Tech., doi:10.5194/amt-12-6635-2019, 2019.
- Rothman, L. S., Gordon, I. E., Barber, R. J., Dothe, H., Gamache, R. R., Goldman, A., Perevalov, V. I., Tashkun, S.
  A. and Tennyson, J.: HITEMP, the high-temperature molecular spectroscopic database, J. Quant. Spectrosc. Radiat.
  Transf., doi:10.1016/j.jgsrt.2010.05.001, 2010.
- Saiz-Lopez, A., Plane, J. M. C., Baker, A. R., Carpenter, L. J., Von Glasow, R., Gómez Martín, J. C., McFiggans, G. and Saunders, R. W.: Atmospheric chemistry of iodine, Chem. Rev., doi:10.1021/cr200029u, 2012.
- 687 Sanchez, J., Tanner, D. J., Chen, D., Huey, L. G. and Ng, N. L.: A new technique for the direct detection of HO2
- radicals using bromide chemical ionization mass spectrometry (Br-CIMS): Initial characterization, Atmos. Meas.
   Tech., doi:10.5194/amt-9-3851-2016, 2016.
- 690 Schnitzhofer, R., Metzger, A., Breitenlechner, M., Jud, W., Heinritzi, M., De Menezes, L.-P., Duplissy, J., Guida,
- R., Haider, S., Kirkby, J. and others: Characterisation of organic contaminants in the CLOUD chamber at CERN,
   Atmos. Meas. Tech., 7(7), 2159–2168, 2014.
- 693 Schwehr, K. A., Santschi, P. H. and Elmore, D.: The dissolved organic iodine species of the isotopic ratio of
- 694 129I/127I: A novel tool for tracing terrestrial organic carbon in the estuarine surface waters of Galveston Bay,
   695 Texas, Limnol. Oceanogr. Methods, doi:10.4319/lom.2005.3.326, 2005.
- 696 Sherwen, T., Evans, M. J., Carpenter, L. J., Andrews, S. J., Lidster, R. T., Dix, B., Koenig, T. K., Sinreich, R.,





- 697 Ortega, I., Volkamer, R., Saiz-Lopez, A., Prados-Roman, C., Mahajan, A. S. and Ordóñez, C.: Iodine's impact on 698 tropospheric oxidants: A global model study in GEOS-Chem, Atmos. Chem. Phys., doi:10.5194/acp-16-1161-2016, 699 2016.
- 700 Simpson, W. R., Brown, S. S., Saiz-Lopez, A., Thornton, J. A. and Von Glasow, R.: Tropospheric Halogen
- 701 Chemistry: Sources, Cycling, and Impacts, Chem. Rev., doi:10.1021/cr5006638, 2015.
- 702 Sipilä, M., Sarnela, N., Jokinen, T., Henschel, H., Junninen, H., Kontkanen, J., Richters, S., Kangasluoma, J.,
- 703 Franchin, A., Peräkylä, O., Rissanen, M. P., Ehn, M., Vehkamäki, H., Kurten, T., Berndt, T., Petäjä, T., Worsnop, 704 D., Ceburnis, D., Kerminen, V. M., Kulmala, M. and O'Dowd, C.: Molecular-scale evidence of aerosol particle
- 705 formation via sequential addition of HIO3, Nature, doi:10.1038/nature19314, 2016.
- 706 Spietz, P., Martín, J. G. and Burrows, J. P.: Effects of column density on i2 spectroscopy and a determination of i2 707 absorption cross section at 500 nm, Atmos. Chem. Phys., doi:10.5194/acp-6-2177-2006, 2006.
- 708 Spolaor, A., Gabrieli, J., Martma, T., Kohler, J., Björkman, M. B., Isaksson, E., Varin, C., Vallelonga, P., Plane, J.
- 709 M. C. and Barbante, C.: Sea ice dynamics influence halogen deposition to Svalbard, Cryosphere, doi:10.5194/tc-7-
- 710 1645-2013, 2013.
- 711 Thalman, R. and Volkamer, R.: Temperature dependent absorption cross-sections of O2-O 2 collision pairs between 712 340 and 630 nm and at atmospherically relevant pressure, Phys. Chem. Chem. Phys., doi:10.1039/c3cp50968k, 713 2013.
- 714 Tham, Y. J., Wang, Z., Li, Q., Yun, H., Wang, W., Wang, X., Xue, L., Lu, K., Ma, N., Bohn, B., Li, X., Kecorius,
- 715 S., Größ, J., Shao, M., Wiedensohler, A., Zhang, Y. and Wang, T.: Significant concentrations of nitryl chloride
- 716 sustained in the morning: Investigations of the causes and impacts on ozone production in a polluted region of
- 717 northern China, Atmos. Chem. Phys., doi:10.5194/acp-16-14959-2016, 2016.
- 718 Vandaele, A. C., Hermans, C., Simon, P. C., Carleer, M., Colin, R., Fally, S., Mérienne, M. F., Jenouvrier, A. and 719 Coquart, B.: Measurements of the NO2 absorption cross-section from 42 000 cm-1 to 10 000 cm-1 (238-1000 nm) at 720 220 K and 294 K, J. Quant. Spectrosc. Radiat. Transf., doi:10.1016/S0022-4073(97)00168-4, 1998.
- 721 Vaughan, S., Gherman, T., Ruth, A. A. and Orphal, J.: Incoherent broad-band cavity-enhanced absorption
- 722 spectroscopy of the marine boundary layer species I2, IO and OIO, Phys. Chem. Chem. Phys.,
- 723 doi:10.1039/b802618a, 2008.
- 724 Wang, M., Kong, W., Marten, R., He, X. C., Chen, D., Pfeifer, J., Heitto, A., Kontkanen, J., Dada, L., Kürten, A., Yli-Juuti, T., Manninen, H. E., Amanatidis, S., Amorim, A., Baalbaki, R., Baccarini, A., Bell, D. M., Bertozzi, B.,
- 725
- 726 Bräkling, S., Brilke, S., Murillo, L. C., Chiu, R., Chu, B., De Menezes, L. P., Duplissy, J., Finkenzeller, H.,
- 727 Carracedo, L. G., Granzin, M., Guida, R., Hansel, A., Hofbauer, V., Krechmer, J., Lehtipalo, K., Lamkaddam, H.,
- 728 Lampimäki, M., Lee, C. P., Makhmutov, V., Marie, G., Mathot, S., Mauldin, R. L., Mentler, B., Müller, T., Onnela,
- 729 A., Partoll, E., Petäjä, T., Philippov, M., Pospisilova, V., Ranjithkumar, A., Rissanen, M., Rörup, B., Scholz, W.,
- 730 Shen, J., Simon, M., Sipilä, M., Steiner, G., Stolzenburg, D., Tham, Y. J., Tomé, A., Wagner, A. C., Wang, D. S.,
- 731 Wang, Y., Weber, S. K., Winkler, P. M., Wlasits, P. J., Wu, Y., Xiao, M., Ye, Q., Zauner-Wieczorek, M., Zhou, X.,
- 732 Volkamer, R., Riipinen, I., Dommen, J., Curtius, J., Baltensperger, U., Kulmala, M., Worsnop, D. R., Kirkby, J., 733 Seinfeld, J. H., El-Haddad, I., Flagan, R. C. and Donahue, N. M.: Rapid growth of new atmospheric particles by
- 734 nitric acid and ammonia condensation, Nature, doi:10.1038/s41586-020-2270-4, 2020.
- 735 Washenfelder, R. A., Langford, A. O., Fuchs, H. and Brown, S. S.: Measurement of glyoxal using an incoherent 736 broadband cavity enhanced absorption spectrometer, Atmos. Chem. Phys., doi:10.5194/acp-8-7779-2008, 2008.
- 737
- 738





- 739 Table 1: Cluster formation enthalpies of different species with bromide ion. The cluster geometries are optimized
- 740 at the ωB97X-D/aug-cc-pVTZ-PP level at 298.15 K. The enthalpies are calculated at the DLPNO-CCSD(T)/def2-
- 741 QZVPP//\0B97xD/aug-cc-pVTZ-PP level at 298.15 K.

Cluster formation pathway	Formation
	enthalpies
	(kcal mol <sup>-1</sup> )
$Cl_2 + Br \rightarrow Cl_2 \cdot Br \rightarrow$	-22.3
$OIO + Br^- \rightarrow OIO \cdot Br^-$	-23.2
$IO + Br \rightarrow IO \cdot Br \rightarrow$	-24.5
$\mathrm{HIO}_3 + \mathrm{Br}^{\text{-}} \rightarrow \mathrm{HIO}_3 {\cdot} \mathrm{Br}^{\text{-}}$	-26.6
$\mathrm{HOI} + \mathrm{Br}^{\text{-}} \rightarrow \mathrm{HOI} \cdot \mathrm{Br}^{\text{-}}$	-26.9
$\mathrm{HIO}_2 + \mathrm{Br}^{\text{-}} \rightarrow \mathrm{HIO}_2 \cdot \mathrm{Br}^{\text{-}}$	-29.2
$I_2 + Br^{\text{-}} \rightarrow I_2 \cdot Br^{\text{-}}$	-33.7
$ICl + Br \rightarrow ICl \cdot Br$	-33.8
$\mathrm{IBr} + \mathrm{Br} \to \mathrm{IBr} \cdot \mathrm{Br}$	-36.7
$H_2SO_4 + Br^- \rightarrow H_2SO_4 \cdot Br^-$	-41.1
$I_2O_4 + Br^- \rightarrow I_2O_4 \cdot Br^-$	-42.6
$I_2O_5 + Br^- \rightarrow I_2O_5 \cdot Br^-$	-53.2

742

743 Table 2: Fragmentation reaction enthalpies of different species with bromide ion. The cluster geometries are

744 optimized at the ωB97X-D/aug-cc-pVTZ-PP level at 298.15 K. The enthalpies are calculated at the DLPNO-

745 CCSD(T)/def2-QZVPP//\0B97xD/aug-cc-pVTZ-PP level at 298.15 K.

Fragmentation
enthalpies
(kcal mol <sup>-1</sup> )
22.3
22.3
26.6
20.8
52.0
26.9
57.7
31.3
29.2
43.8
42.2
33.7





$I_2{\cdot}Br^{\text{-}} \to IBr + I^{\text{-}}$	33.8
$\mathrm{ICl} \ \mathrm{Br} \to \mathrm{ICl}{+} \ \mathrm{Br}$	33.8
$\mathrm{ICl} \cdot \mathrm{Br} \to \mathrm{IBr} + \mathrm{Cl}$	39.8
$ICl \cdot Br \to BrCl + I$	42.0
$\mathrm{IBr} \; \mathrm{Br} \to \mathrm{IBr} + \mathrm{Br}$	36.7
$\mathrm{IBr} \cdot \mathrm{Br}^{\text{-}} \to \mathrm{Br}_2 + \mathrm{I}^{\text{-}}$	39.4
$H_2SO_4{\cdot}Br^{-} \rightarrow H_2SO_4 + Br^{-}$	41.1
$\mathrm{H_2SO_4}{\cdot}\mathrm{Br}^{\scriptscriptstyle -} \to \mathrm{HSO_4}^{\scriptscriptstyle -} + \mathrm{HBr}$	27.9

746

# 747 Table 3: Calibration coefficients for selected species.

Calibration	coefficient
(molec cm <sup>-3</sup> c	ps cps <sup>-1</sup> )
<sup>a</sup> 2.7×10 <sup>10</sup>	
<sup>b</sup> 6.3×10 <sup>10</sup>	
${}^{b}3.5{\times}10^{11}$	
<sup>b</sup> 3.3×10 <sup>11</sup>	
$^{a}4.1 \times 10^{10}$	
	(molec cm <sup>-3</sup> c <sup>a</sup> 2.7×10 <sup>10</sup> <sup>b</sup> 6.3×10 <sup>10</sup> <sup>b</sup> 3.5×10 <sup>11</sup> <sup>b</sup> 3.3×10 <sup>11</sup>

748 a: inter-method calibrations

b: offline calibrations