Use of thermal signal for the investigation of near-surface turbulence

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Abstract. Organized motions in the roughness sub-layer were investigated using novel temperature sensing and data science methods. Despite accuracy draw backs, current fibre-optic temperature (DTS) and thermal imaging (TIR) instruments offer frequent, moderately precise and highly localized observations of thermal signal in a domain geometry suitable for micro-meteorological applications near the surface. Horizontal and vertical cross-sections allowed a tomographic investigation of

5 the spanwise and streamwise evolution of organised motion, opening avenues for analysis without assumptions on scale relationships. Variance events on the order of seconds to minutes could be identified, localized, and classified using signal decomposition and machine learning techniques. However, small-scale turbulence patterns at the surface appeared difficult to resolve due to the heterogeneity of the thermal properties of the vegetation canopy, which are not immediately evident with the naked eye. The results highlight a need for physics-aware data sciences techniques that treat scale and shape of temperature

10 structures in combination, rather than as separate features.

1 Introduction

Available methods to determine energy and scalar fluxes from the terrestrial land surface are relatively imprecise due to a multiscale of irregularities in the land surface and the turbulent transport mechanisms (Brutsaert, 1998). The most broadly adopted

15 micro-meteorological methods for quantification of surface exchange fluxes take an integrative approach to observations in the atmospheric surface layer (ASL) and have become a crucial component in global Earth System assessments on climate change mitigation (Foken, 2006). On that account, questions have been asked about additional details contained in such data (e.g., Knauer et al., 2017; Klosterhalfen et al., 2019; Clement and Moncrieff, 2019), but also, more pressingly, what we are missing outside the applicable range of the methodology and assumptions (Mahrt, 2010; Bou-Zeid et al., 2020).

20 The complex relationship between the scale and heterogeneity of a source area and observations of turbulence and temperature (and humidity) in the ASL have been broadly recognized (Garratt, 1980; Schmid, 1994; Raupach and Finnigan, 1995; Mahrt, 1996; Sun et al., 1999; Li et al., 2012; Patton et al., 2016) in decades of work that contributed to the best practise guidelines for the eddy covariance technique (Lee et al., 2005; Aubinet et al., 2012). The atmosphere, the surface and the plant microclimate are in direct exchange in the so-called roughness sublayer, the internal ASL layer below the inertial sublayer

25 (Raupach, 1979; Garratt, 1980; Stull, 1988). Finnigan (2000) provides a detailed review on aspects of turbulence within the RSL, where individual roughness elements modulate the flow in a three-dimensional and inhomogeneous way. The correlation between the statistics of turbulence, air temperature and skin temperature of vegetation has been investigated using single-point observations as well as spatial data (Finnigan, 1979; Paw U et al., 1992; Shaw et al., 1995; Katul et al., 1998), highlighting the presence of organized structures in the RSL in accordance with insights from laboratory experiments on near-wall fluid

30 dynamics (Kline et al., 1967). Coherency of eddies has been observed in profiles within the RSL from vector and scalar data, such as temperature (Antonia et al., 1979; Williams and Hacker, 1992). Many near-surface turbulence studies adopt Taylor's hypothesis (frozen convection) to translate wave numbers to length scales, but it could be questioned if and when the underlying assumptions hold, particularly for non-ideal conditions (stationarity, ergodicity, homogeneity; see, e.g., Higgins et al., 2013).

35 A way to avoid making such assumptions would be to observe turbulence features explicitly in space and time. Dense geometrically distributed observations of air and surface temperature in combination with the wind field have not been explored in great detail outside a wind tunnel, mostly because methods were not available or because the costs would be prohibitive. However, novel opto-electronic technologies may bring a needed change.

The availability of spatial surface temperature observations has improved with advancements in spectral imaging. Thermal

40 infrared imaging (TIR) instruments observe long-wavelength infrared light from which a surface brightness temperature (T_b) can be derived. TIR instruments have proven to be sufficiently sensitive, as demonstrated in recent studies on spatial temperature fluctuations (Christen et al., 2011; Katurji and Zawar-Reza, 2016), planar flow fields (Abram et al., 2013; Inagaki et al., 2013; Burns and Chemel, 2013) and partitioning of energy fluxes (e.g., Kustas and Anderson, 2009). It should be noted that the sensitivity of TIR to measure T_b is affected by thermal signal other than the surface skin temperature; the quantification of T_b

45 outside a controlled laboratory environment is a challenge, in and of itself. Parameters that affect the observation of T_b include external influences between the camera sensor and the object, variation in thermal admittance, thermal reflections, moisture and humidity (Vollmer and Möllmann, 2010; Jones, 1999).

Distributed Temperature Sensing (DTS) provides highly localized temperature observations along optical (silica) fibre (Dakin et al., 1985), making it a tool for spatial environmental sensing (Selker et al., 2006; Tyler et al., 2009). The optical

50 fibre is usually packaged inside a cable for durability and to help minimize attenuation due to (micro-)bending. The cable jacket adds surface and volume, making the measurement more susceptible to the local energy balance terms (Thomas et al., 2012; de Jong et al., 2015). Hence, micro-meteorological studies with DTS have primarily been demonstrated for conditions of nocturnal darkness (Keller et al., 2011; Thomas et al., 2012; Zeeman et al., 2015) or, after adjustments, taking advantage of a differentiating thermal signal governed by wind, evaporation or shading of sunlight (Petrides et al., 2011; Euser et al., 2014;

55 Sayde et al., 2015; Lapo et al., 2020; Schilperoort et al., 2018; van Ramshorst et al., 2020). Fibre-optic cable can be deployed to form a mesh of horizontal and vertical profiles in complex geometries from the ground up.

The goal of this study was to combine DTS and TIR for the investigation of temperature and wind field statistics near a vegetation surface. These novel methods are moderately precise, but constrained in accuracy for observation of air and surface temperature. However, both DTS and TIR promise geometrically dense thermal information that could complement single-

60 point observations using well-known sensing technique, such as ultrasonic anemometry for wind and metal wire resistance for temperature. The combined granularity may allow us to revisit the investigation of scale and structure in the roughness sub-layer in more detail. The examination of multi-dimensional data has taken flight with the availability of novel statistics and machine learning tools, which will also be addressed.

2   Materials and methods

65 The study was conducted at the DE-Fen station, Fendt–Peissenberg, Germany, which is a TERerstrial ENvironment Observatories and ICOS core site (TERENO; Zacharias et al., 2011; Kiese et al., 2018). The site is a permanent grassland situated in a shallow valley bordered to the West by a steep, forested slope towards a plateau at approximately 150 m higher elevation (Figure 1). The Fendter Bach stream and tributaries to the Rott developed a more prominent canyon approximately 2 km to the South-West of the site. Gullies in the valley area flow into the Rott, which drains the area into the Ammer to the North.

70 During the ScaleX campaigns, additional experiments were conducted for the investigation of scale interactions between the atmospheric boundary layer and the surface, as well as validation of measurement techniques (https://scalex.imk-ifu.kit.edu; Wolf et al., 2017; Zeeman et al., 2017; Brosy et al., 2017; Brenner et al., 2018; Mauder and Zeeman, 2018; Zhao et al., 2018; Ibraim et al., 2019; Hald et al., 2019; Zeeman et al., 2019; Kunz et al., 2020). The experiment took place during the ScaleX 2016 campaign in Jun–Aug 2016, but we focus here on reference period 18–22 Jul 2016. These days included a typical summer

75 warm-up phase with mostly clear sky conditions and calm nights, but with limited activity by uncrewed aerial vehicles (UAV) perturbing the air.

2.1   Ultrasonic anemometer (EC) network

A network of wind sensors was deployed, comprising profiles of 3-axis (3D) and 2-axis (2D) ultrasonic anemometers on a guyed lattice mast and four tripods (Figure 1; Table). The EC network was placed approximately 60 m East of the micro-

80 meteorological station of DE-Fen. The sensors were cross-calibrated prior to the experiment (Mauder and Zeeman, 2018). An additional EC station (ICOS; EC with trace gas analysers) was temporarily placed at the center of the sensor network (Figure 1c; Figure 2).

2.2   Distributed temperature sensing (DTS)

The DTS set-up consisted of a high-resolution instrument, custom fibre-optic cable and a calibration set-up. The DTS instru-

85 ment was configured to measure at 1 Hz and 0.15 m range gate intervals up to 1890 m (model Ultima-HS; Silixa Ltd, Hertford-shire, UK). The fibre-optic cable consisted of a bend-optimized optical fibre (Giga-Link communications type; 125 μm with 50 μm core), buffered with aramid fabric (Kevlar) in a white plastic jacket (900 μm outer diameter; AFL Telecommunications, Spartanburg, SC, USA). Both ends of the fibre-optic cable were connected to separate channels on the DTS instrument for measurement in both directions (Hausner et al., 2011). The configuration allowed two profiles along the lattice mast up to 8 m

90 above ground and 316 profiles suspended in the air up to 4.5 m above ground. The latter were spatially configured to form the

sides of a box of $20 \times 20 \times 3.5$ m [L×W×H], or four planes of 80 vertical profiles in 0.25 m intervals, with EC profiles placed at each corner (Figure 1c and Figure 2). The dimensions were chosen as a trade-off between DTS range limitations, instrument spatial resolution and measures to limit systematic shading of direct solar radiation or wind by DTS and EC support structures. The effective thermal sensitivity is dependent on the instrument, the deployed cable geometry and the location along the fibre, which previously was estimated to be 0.5 K or better (Zeeman et al., 2015, see Appendix A3.1).

2.3 Thermal infrared imaging (TIR)

A thermal infrared imaging system was installed inside a weather-proof camera housing at 8.5 m above ground (8–14 μm spectral range, 40 mK thermal sensitivity, $53° \times 38°$ field of view and 382×288 pixels optical resolution; model PI 640, Optris GmbH, Berlin, Germany). The spectroscopic data was recorded at 2 Hz between $00'$ and $53'$ in the hour with automatic self-calibration each 9 s.

2.4 Data integration

Each EC, DTS and TIR record was stored with an accurate time stamp and locations were determined in post-processing. The calibration details are provided in Appendix A. The physical meaning of the EC, DTS and TIR derived temperature measurements must be considered. Until that is fully explained, or until we assure the values match the expectation in atmospheric research, we should discern the nomenclatures from air temperature (T_a) or surface temperature (T_s). DTS can be used to measure temperature of air, but we denote it as "cable" temperature (T_c) below, in order to distinguish it from radiation-shielded and fan-aspirated sensing standards used for air temperature in micro-meteorology. Following similar argumentation, T_b and T_v are differentiated here as surface brightness temperature and temperature from the speed of sound (or acoustic virtual temperature), respectively. The use of the EC-based T_v is common in micro-meteorological studies, once noise and calibration have been handled with care (Mauder and Zeeman, 2018).

3 Results

3.1 General patterns

First, we will briefly cover aspects related to wind and radiation for the study period at the site, as trends or variability in those variables influence air temperature near the surface substantially. The selected days showed an almost cloudless warm-up phase with a diurnal cycle in temperature and wind that is characteristic for locations at this latitude and proximity to the Alps (Figure 3). The different temperature observations show a tooth-shaped pattern with a rapid increase after sunrise and a maximum mid-afternoon, thereafter temperature drops rapidly and continuous to decrease after sundown under radiative cooling (Figure 3d–e). Some short-term variability in surface and air temperature can be explained by transient cloud cover, such as at midnight and at noon on 19 Jul 2016, which changes the radiation balance at the surface by shadowing of sunlight (Figure 3d–e). Sudden temperature shifts also occur in conjunction with a passing storm, such as right after sunrise on 21 Jul

2016 (Figure 3f). This particular rapid change in temperature correlated with increased wind speed, that is atypical for the early morning at this site (Figure 4). Weak wind and absence of clouds allowed unstable conditions to develop after sunrise on these calm cloudless days, as well as after a rain event on the morning of 21 Jul 2016 (Figure 4b–d). High surface energy fluxes during the day correlated with a northern wind sector in mid-morning to mid-afternoon observed near the surface. The

125 wind slows and turns to a S to E wind sector during the evening transition, with notably weak wind close to the surface and continuing through the night (Figure 4a–b; Figure 4e). Please note that this reoccurring diurnal pattern of shifting wind direction is sometimes referred to as 'Alpine pumping'. As a consequence of very weak wind conditions, eddy-covariance flux computations and some turbulence statistics were rarely acceptable in the hours from just before sundown until shortly after sunrise (Figure 4e–g). It is clear that the assessment criteria for quality of flux computations leave plausible values mostly for

130 the unstable case, as expected. In the past, alternative stability classifications have been formulated from bulk Richardson or temperature gradients, which are provided here for comparison (Figure 4g; Appendix B). Using higher resolution temperature gradient information, T_c reveals a vertical deviation in stable classification above 1.5 m and unstable below in the late afternoon, which is confirmed by T_a (Figure 4g). The match foremost highlights agreement in temperature gradient observations between the methods. The stability details are shown here to point out that with sufficient spatial resolution, simple indicators become

135 a means to reveal physical patterns. This will be discussed below.

3.2 Temperature

The air temperature observations agree between the methods (Figure 3a–c). Some dissimilarity between T_a and T_c can be seen at the lowest part of the profile (< 0.5 m) during the day, where the effects of sensor shading and aspiration on T_a can be expected to be most pronounced (Figure 3a–b). A stable temperature profile develops at night in excess of 3 K in the lowest 9 m

140 depth (Figure 3a–c; Figure 5a–b). The unstable daytime temperature gradient is small in comparison, but temperature variance can be significantly larger and increases towards the surface (Figure 5c–f; Figure 5g, positive kurtosis). Furthermore, the shape of the variance distribution of those daytime observations indicate a moderate dominance of short-term hot excursions from the mean (Figure 5g, positive skew). Those are not new findings, but again describe the general conditions at the site.

3.3 Temperature time scale decomposition

145 In addition to aggregating statistics for fixed time intervals, we can decompose the air temperature signal in contributing time scales (Figure 6). The multi-resolution decomposition (MRD) results confirm a variance maximum at the surface during the day, but at night the dominant contribution was found between 0.5 and 2.5 m depth (Figure 6b; details about MRD in Appendix B). Not surprisingly, highest daytime variance was found after the rain storm on 21 Jul 2016 under wet surface conditions (Figure 3f; Figure 5; Figure 6b). During the day, scales of less than 3 minutes contributed most to total variance,

150 whereas periods of 5 minutes or were more prominent at night (Figure 6c). For each of the decomposition levels we can determine a height of maximum contribution to variance. During the day, a peak contribution was observed at the bottom of the profile, also corresponding to higher decomposition levels ($M > 7$, lower frequency signal). During the day, the sunlit surface acts as a slow-response buffer in transfer of energy to the air in closest contact. The height of the maximum variance for smaller

decomposition scales ($M < 6$, higher frequency contribution) was shown to transition from below 0.5 m before noon, to the top of the profile. During the night, the height of the variance maximum for the large and small decomposition levels could be found to coincide, as shown between 1 and 2 m depth on the night of 19 to 20 Jul 2016 (Figure 6d). This suggests an interaction between scales.

The decomposed variance scales for the surface brightness temperature differ in distinct aspects, but generally follow a similar pattern (Figure 7). Similar to T_c, there is a increase in contribution of lower decomposition levels ($M < 7$, higher frequency signal) during the day (Figure 7b). However, the total variance at the surface is very low during the night (Figure 7c). This confirms the results for T_c, and suggests that the stable stratification at night allows the air (a layer) closest to the vegetation to decouple from the atmosphere aloft. Interestingly, spatial patterns do emerge from the T_b variance that correlate to location, also at night when total variance is very low. This is discussed in more detail in Section 3.5.

Please note that the MRD results helped identify noise peaks affecting the whole thermal image in the early morning on most days. This noise signal is likely caused by direct sunlight at very low zenith angles (Figure 7a). Furthermore, the T_b signal shows limited information in the lowest decomposition scales ($M < 2$, high frequency), which is assumed to be a design feature of the camera and likely governed by noise-reduction algorithms that result in a low temporal sensitivity. The T_c results contain more high-frequency signal than the T_b, as is evident from the the lowest decomposition levels and typical for lidar data such as DTS.

3.4 Temperature event classification

The MRD and other decomposition methods, such as wavelet, help decompose variance in time, but do not necessarily discern the shape of the variance events. In fact, the outcomes of MRD differ substantially between a signal with sinewave or ramp shape, which both commonly occur in air temperature time series. For the same event duration, MRD can classify a wave and a ramp shape one decomposition level apart. It is therefore helpful to investigate the shape of variance events in more detail. An indicator for event shapes that likely show a more saw-tooth or ramp-like pattern is positive skewness. The difference between the first half and second half of the event provides addition detail. Such statistical features can be used to identify events, and this is the basis of methods like TED (Kang et al., 2013, 2015, see Appendix B). The original TED classification is based on differences between variance events in Euclidean space, i.e., k-means clustering on principle components analysis of event statistics. Different event types may occur though-out the day and through-out the vertical profile (Figure 8). The classification of some clusters appeared to coincide with the timing and height of dominant variance scales identified and localized by MRD (Figure 6d). This is seem most clearly at night.

The accuracy of event classification statistics dependents on the signal and the ability to separate the noise it includes. Not investigated here is how such statistical event features translate between different sensor types with different noise profiles. Further analysis of lagged spatial correlations would be beneficial for the identification of patterns and the evolution of structures, which the applied algorithms currently do not address.

3.5 Spatial temperature patterns

Based on the scale decomposition and cluster results we focus on a nighttime and daytime period more closely, starting 03:00 UTC 20 Jul 2016 and 10:00 UTC 21 Jul 2016, respectively. An example of night and daytime observations can be viewed in an animation (see Video supplement). During the animated periods, the mean wind is mostly aligned with the DTS measurement set-up, with mean flow from a SSE direction. Both the daytime and nighttime case show intermittent structures. However, animations reveal scale interactions that are not easily identified in the decomposition or classification statistics.

First, hot plumes or thermal streets advect laterally through the observation volume during the day and appear correlated to stream-wise oriented hot and cold streaks on the surface. The plumes maintain a moderately constant spanwise position (Figure 9). Motion of the heat structures is not necessarily aligned with the mean wind as recorded by the EC profiles. Please note that this is in part an issue of apparent (co-variant) alignment of the vertical planes with the flow (Zeeman et al., 2015). Surface heat emits from hot areas in the vegetation, which are mostly sunlit leaves at the top of the canopy. Remnants of tractor tracks are visible parallel to the DTS setup and field track in a WSW to ENE direction (Figure A6; on close inspection also in Figure 2). The vegetation canopy was higher where tractor tires did not disturb the soil and vegetation during the previous harvest. The streaks of hot and cold air appear to move independent of the orientation of those tracks, although hot T_b signal shows up more clearly against the cool background of low vegetation in the tractor tracks, making its visual appearance dependent on the approach use here in which all variance scales are included. It is assumable, that patterns in thermal brightness signal relate to heat and moisture released by evapotranspiration, where variability in moisture contributes to a spatial differentiation in emissivity. Those emissivity values are unknown, hence the T_b signal is not necessarily an accurate measure of T_s. Nevertheless, we can clearly identify structures from motion in the T_b signal during the day. Analysis of decomposed T_c signal are needed to accurately distinguish the heterogeneity in the transmission of heat, related to canopy structure. Second, occasionally, gusts cause divergent motion at the surface, revealed in T_b as a cold patches that are reminiscent of gust landing on flat water, forming so-called capillary waves (see, e.g., Dorman and Mollo-Christensen, 1973). Third, shear-induced wave and wave-breaking patterns were shown at multiple scales in the night-time case, but no clear signal was found in T_b at the surface (Figure 10). The location of the shear (and wave) interfaces coincided with the MRD and TED determined heights for variance maxima discussed above. Most nighttime events correlated with wind direction shifts in the profile. Wind shear events were shown to cause cascades of ripples, which occasionally are only visible in one section of T_c but not elsewhere, suggesting the disturbances could be local.

The computed turbulent kinetic energy (TKE, $\overline{e}$) supports this observation (Figure 11). Although TKE is an order of magnitude less during the night than during the day, peaks that did occur at 3 m were not necessarily registered at all corners of the observation domain, or at others heights in the profile. A peak at 3.0 m at 03:28 UTC on 20 Jul 2016 exemplifies this ((Figure 11b and (Figure 10). The decomposed variance and identified clusters for this event provided further evidence that the wave interaction did not perturb the air closest to the surface (Figure 12). Classified TED events were detected at the same height as the variance maximum from the MRD decomposition, but an exact match to the timing of dominant variance peaks and structures did not occur. Possible explanations were the minimum event window size of approximately 120 s and the effect

220 of high-frequency noise in the T_c time series. Please note that these particular night-time events are characterised by an abrupt negative temperature anomalies.

The T_c and T_b methods resolve an order of magnitude more locations than T_v or T_a data, but can only reveal structures that exhibit substantial temperature differences and spatial (auto-)correlation. Combining the observations of horizontal and vertical planes allows combined investigation of spanwise and streamwise correlations. A decomposition along multiple dimensions

225 (e.g, two-dimensional space), a classification of two-dimensional shape features and the quantification of time-sequential motion are outside the scope of this study.

3.6 Experimental design

There are limitations in the described approach using DTS and TIR data which would require further thought on experimental design. First, the support structures of the fibre-optics and other sensors were placed as far away as practicable with a suspension

230 cable system (Appendix A3.1), but this did not prevent shadows on the surface and profiles at lower zenith angles. The shadow effects are systematic and can be overcome by careful selection of data points. Vertical shadows are less of a concern, as these affect only a small number of vertical profiles at a time, and nighttime data is less affected. Second, a circle-shaped geometry for the fibre-optic cable set-up would make up- and downwind pairs equidistant, potentially simplifying analysis by making the lateral gradients independent of wind direction. Third, in retrospect after analysis, placement of three-dimensional sonic

235 anemometers at lower heights in the profile would have been advantageous. Fourth, parallax issues caused by objects in view of the thermal camera can only be overcome by using multiple cameras from different viewpoints. An uncluttered, regular grid matrix is helpful for time-sequential analysis of motion, for instance, using algorithms developed in computer vision research. Fifth, continuous thermal imaging data would be needed. The hourly interruption of the recording of T_b in this experiment was caused by a meanwhile resolved software issue. Sixth, the addition of a reference surface temperature measurement within

240 view of the TIR camera will be useful for the assessment of noise in the image time-sequence. Seventh, the use of a thinner fibre-optic cable design with a lower thermal inertia could improve the T_c response, although the sensitivity of DTS technique may become a limitation at high recording frequencies (see Appendix A3; 0.2 K white noise at 0.5 Hz). Eight, the TED skill may be improved with the information from the scale decomposition, e.g., by applying the method on a variance subset from which high and low frequency contributions are excluded. Nine, the use of supervised machine learning techniques benefits

245 from the availability of statistical features with a physical meaning. Additional event metrics, particularly those supported by past research on RSL processes, should be explored. Even simplified stability indicators could turn out useful, in favor of the moderately high computational costs of signal decomposition (e.g, MRD and TED) applied to high-resolution sensor networks.

4 Conclusions

Despite accuracy drawbacks, current DTS and TIR applications promise frequent (0.2 Hz or better), moderately precise (0.2 K

250 or better) and highly localized (0.25 m or better) observations of thermal signal in a spatial domain suitable for micro-meteorological applications near the surface (2.5-dimensional or better). The fingerprints of motion and turbulence structure

could be identified in horizontal and vertical planes. Horizontal and vertical cross-sections allow a tomographic investigation of spanwise and streamwise organised motion, in a novel combination with time-sequential thermography. However, small-scale turbulence is difficult to resolve due to the heterogeneous thermal properties of the vegetation surface and sensor precision

255   at short signal integration times. The heterogeneity in thermal activity of the vegetation points at ecophysiological traits that may be contributing factors in the formation and distribution of coherent structures near the surface, which needs to be investigated further. The ability to trace coherent motion in space and time may proof useful for the development of conditional sampling methods that complement the eddy-covariance technique. The comparability of TED results between time series could be improved with a (hybrid-)hierarchical clustering modification, but the modifications only partly resolved limitations

260   of repeatability of the cluster analysis. The identification and classification of events in turbulence time series will benefit from the development of physics-aware machine learning techniques that quantify combinations of scale and shape features, instead of treating those as separate properties. In addition, a supervised approach for event detection using machine learning techniques may lead to new avenues for investigation of turbulence interactions near the surface from increasingly high-resolution observations of atmospheric properties.

265   *Code and data availability.* Data from this study are available with Open Access (DOIs 10.5281/zenodo.4267887, 10.5281/zenodo.4267810 and 10.5281/zenodo.4061242). Software code are available as notebooks for python and R (https://gitlab.imk-ifu.kit.edu/scalex/) or from the author.

Video supplement. The video supplement contains an animation of high-resolution data for a nighttime and daytime hour (DOI: 10.5446/50229).

Appendix A: Data procedures

270   ### A1   Time synchronisation

Sensor data from each EC profile were recorded by a datalogger system with accurate time keeping, comprising a low-power computer (model Beaglebone Black, BeagleBoard.org Foundation, Oakland Twp, MI, USA), four digital communication ports (RS232), a temperature-compensated hardware clock (model ChronoDot, DS3231; Maxim Integrated Products Inc., San Jose, CA, USA) operated by a modified operating system with accurate time synchronisation (Debian Linux; Chrony software

275   module). In addition, an asynchronous network framework was used to record up to four EC sensors simultaneously with low latency (Python 2.7; Twisted software module). Similar systems were used as time reference device in the local computer networks of the DTS and TIR, recording time offset continuously. The time offset between the data logging systems was determined before, during and after the campaign and showed no measurable drift. A GPS unit was used as a backup reference to world time, as access to internet real-time services could not be guaranteed at all times. Please note that time is reported in

280   UTC, but local time was CEST (UTC+02:00; CET summer time).

A2 Georeferencing

The EC and TIR sensors were referenced by measurement tape compared to the central EC location (ICOS station). The same was done for the rings used in the support structure for the DTS fibre-optic cable.

A2.1 Wind direction

285 The orientation of each sonic anemometer was determined using a magnetic compass. However, this appeared imprecise and relative offsets ($< |4°|$) were determined from a period of data with strong wind from a persistent N direction, along the valley axis. The wind vectors in all EC time series were corrected accordingly.

A2.2 Fibre-optic cable

The DTS fibre cable was suspended using white rings with grooves in which the fibre-optic cable was held in place by gravity
290 and light tension (Figure A1). All coils and turns made with the fibre-optic cable throughout the set-up, including the temperature calibration, were made to match the diameter of the grooves of these rings (cut from polypropylene drainage pipe; Ø 0.25 m). The rings were kept in place and aligned with a support cable through the center of the ring. The locations of the straight vertical sections were initially identified in the DTS signal with ice packs at known locations. However, the fibre-optic cable developed in the initial days of deployment (June 2016, prior to the experiment), as the material of the outer jacket
295 stretched under heat and tension until it matched the almost non-elastic Kevlar buffer. Slack in the fibre-optic cables was removed by adding tension to the steel suspension cables using pulleys and by coiling excess cable on the bottom rings. As the position of some sections of cable could still change, e.g., accidentally or by re-adjustment, the location of the straight vertical sections were re-confirmed by detection of peaks in signal attenuation (Stokes signal) that correlate to the bend sections of fibre on the rings, as well as temperature differences at noon (bottom rings are warm, compared to top) and after sundown
300 (bottom rings are cold, compared to top). The latter also helped confirm the direction and height of each vertical profile (Figure A2). The algorithm to detect ring-coiled sections of cable used a combination of (wavelet-)filters on the DTS signals and their derivatives, including minima, maxima and zero-crossing points in $[T'_c(x_i)]$ and $[T''_c(x_i)]$, where x_i is the DTS range-gate interval distance along the fibre. The optical fibre is packaged loosely inside the cable with up to 10 % "overstuff", hence measurement along the cable exterior are only an approximation of location of DTS range gates. A 0.15 m section along the
305 fibre translated to a fibre-optic cable section of approximately 0.125 m.

A2.3 Image projection

To map image pixels to real-world locations, the lens optical distortion was determined, as well as the location, the height, and the roll, pitch and yaw rotation of the sensor (Figure A5). The spatial resolution at the surface varied with distance away from the mast due to camera perspective and parallax. To facilitate computation, the image time series were resampled to a regular
310 0.125 m grid, aligned with the spatial resolution and orientation of the DTS set-up (Figure A6), excluding pixels far outside

the 20×20 m domain or depicting objects. For the translation from a perspective view to an orthogonal map, the surface was assumed to be flat.

A3 Temperature reference

A3.1 DTS cable temperature (T_c)

315 The DTS observations were calibrated using reference temperature measurements. Reference temperature sensors (two PT-1000 probes) are an integral part of the DTS instrument and recorded automatically. The sensors were placed inside two insulated water baths in close proximity to sections of fibre-optic cable. One bath was heated using an aquarium heater (32 °C; model Fluval E-100, Hagen GmbH, Holm, Germany), the other bath was chilled using a laboratory flow-through cooler (4 °C; model Titan 500, AB Aqua Medic GmbH, Bissendorf, Germany), and both baths were continuously mixed using aquarium

320 pumps (10 L min⁻¹ flow rate, directed horizontally at half depth; model PondoCompact 600, PonTec, Hörstel, Germany). At both ends of the fibre-optic cable and in both calibration baths, at least 20 m of cable was placed loosely on a submerged spool made of mesh material, in order to keep the cable away from the walls of the bath and achieve temperature equilibrium with the water. For each measurement, the span and offset values determined for both ends of the cable were linearly interpolated and applied to all intervals in-between the calibration sections (see, e.g., Hausner et al., 2011; Zeeman et al., 2015). Between

325 0.1 and 0.2 K of variance can be attributed to white noise, as shown from an Allan deviation plot on data before calibration (Figure A4; see, e.g., Werle, 2010).

A3.2 TIR brightness temperature (T_b)

The imaging data was converted to a time-series of brightness temperature assuming an emissivity of 0.96. However, a high-frequency noise was found in the time series that affected the whole image. The signal was assumed to be an artifact of small

330 variations in the exposure time of the detector and sensitivity to the interruptions by the automatic self-calibration mechanism. The signal could effectively be removed by subtracting the mean temperature of the whole image (or a selected area) from the value of each pixel. Although this effectively applies a high-pass filter in the time domain, the image data are improved in a meaningful way for the applicability of time-sequential thermography.

A3.3 Aspirated air temperature (T_a)

335 Air temperature measurements were made with resistance temperature devices in fan-aspirated enclosures at four heights along the mast (RTD PT1000; model 41347 and 43502, R.M. Young Company, Traverse City, MI, USA; see Table 1) and recorded at 0.5 Hz (model CR5000, Campbell Scientific Inc, Logan, UT, USA). The temperature sensors were calibrated for span and offset prior to the campaign using a well-mixed water bath (see Appendix A3.1), while increasing the temperature along an approximate ambient range (18–27 °C; PT100 based digital thermometer, model GMH 3710, Greisinger Electronics GmbH,

340 Regenstauf, Germany; Wolz, 2016).

A4 System integration

Some additional measures were needed to integrate all systems. In order to avoid AC power hazards, to limit electrical cross-interference and to accommodate voltage drops on long cables, all outdoor devices were connected to a central 24 V DC power supply and grounding to earth, with subsequent local down-regulation to 5 or 12 V DC. Similarly, network connections were
345 made using fibre-optic links instead of metal wire cables. The DTS system, including a temperature calibration set-up, was operated inside a temperate-regulated housing. To accommodate data storage requirements for continuous operation, all data were automatically transferred on an hourly (DTS, TIR) or daily (EC) basis. Python was used for data acquisition, analysis and figures (Python 3.7; Hoyer and Hamman, 2017; Virtanen et al., 2020; Anaconda, 2020; Plotly Technologies Inc., 2020), with additional analysis produced by R (R Core Team, 2018; Sarkar, 2008; Wickham et al., 2020).

350 ## Appendix B: Definitions

The comparison of computed variables included:

$$\overline{U} = \begin{cases} \overline{u} & \text{after double rotation} \\ \sqrt{\overline{u}^2 + \overline{v}^2} & \text{otherwise} \end{cases} \tag{B1}$$

$$\overline{e} = \tfrac{1}{2}(\overline{u'u'} + \overline{v'v'} + \overline{w'w'}) \tag{B2}$$

$$u_* = \sqrt[4]{\overline{u'w'}^2 + \overline{v'w'}^2} \tag{B3}$$

355 $$R_{\mathrm{B}} = \frac{g}{\overline{T}(z_1)} \frac{\left[\dfrac{\overline{T}(z_1) - \overline{T}(z_2)}{z_1 - z_2}\right]}{\left[\dfrac{\overline{U}(z_1) - \overline{U}(z_2)}{z_1 - z_2}\right]^2} \quad \text{where } [z_1 > z_2] \tag{B4}$$

$$L = -\frac{u_*^3 \; \overline{T}}{\kappa \, g \, (\overline{w'T'})} \tag{B5}$$

$$\zeta = \frac{z}{L} \tag{B6}$$

where $\overline{U}$ is the total horizontal wind velocity, $\overline{e}$ is the turbulent kinetic energy (TKE), u_* the friction velocity and R_{B} the bulk Richardson number, with z the height above ground and g the acceleration due to gravity, L is the Obukhov length with
360 κ the von Karmann constant and ζ a dimensionless scaling parameter (Stull, 1988). The overbars denote averaging in time. Double rotation refers to a method of alignment of the coordinate system into the mean wind (Kaimal and Finnigan, 1994). Stability classes were taken from literature and repeated for completion in Table 2.

A Multi Resolution Decomposition (MRD) approach was used to determine the the weighted contribution of different time scales to the total variance (Howell and Mahrt, 1997; Vickers and Mahrt, 2003). The MRD filter is analogue to a Haar transform
365 in use of a rolling window for discrete decomposition of scales (2^{M} to 2^0 samples) for the computation of (co-)spectral variance. The difference with Haar is a scale-dependent weighting function to resolve the contribution of each scale to the total variance,

simply by cumulatively subtracting the variance of larger scales from subsequent smaller scales. The largest window size (2^M samples) was linearly detrended before MRD computation.

Complementary to decomposition approaches based on scales (MRD, or wavelet), intermittent events can be identified

370 based on characteristics in the turbulence time series. The Turbulence time series Event Detection (TED) approach can identify potential events in temperature time series and apply a subsequent cluster classification based on statistical features of those events (Wang et al., 2006; Kang et al., 2013, 2015, see the R-package TED for code examples). More precisely, the cluster classification is based on a Principal Component Analysis (PCA) on Euclidian distance measures for a set of statistical features, including kurtosis, skew, variance, minima and maxima within the event time window. The method was modified to use a

375 hybrid hierarchical k-means clustering to improve the repeatability of the clustering for the many co-located T_c time series. The presented results are based on PCA on all events and features detected within the T_c array and for a one-day time window. Kang et al. (2015) empirically found a number of six clusters for a day of CASES99 temperature data (Kang et al., 2015; Poulos et al., 2002), which was used accordingly here. Please note that TED does not make assumptions on the duration of events, but a minimal rolling window size of approximately two minutes (90 samples) was applied and events longer than 30

380 minutes were excluded (Fig. A7).

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Table 1. Overview of profile instruments. The 2-axis and 3-axis ultrasonic anemometer are referred to as 2D and 3D, respectively.

Profile	Height (m)	Type	Model	Manufacturer	Serial number
EC & TIR mast	9.0	3D	uSonic-3 Omni	METEK GmbH	106054006
	6.0	3D	1210R3	Gill Instruments Ltd.	595
	9.0	RTD	43347, 43502	R. M. Young Company	15087
	3.0	RTD	43347, 43502	R. M. Young Company	15080
	1.0	RTD	43347, 43502	R. M. Young Company	15084
	0.5	RTD	43347, 43502	R. M. Young Company	15081
EC tripod NW	3.0	3D	CSAT3	Campbell Scientific Inc.	1791
	2.0	2D	WindSonic	Gill Instruments Ltd.	08460064
	1.0	2D	WindSonic	Gill Instruments Ltd.	08460063
EC tripod NE	3.0	3D	CSAT3	Campbell Scientific Inc.	0771-1
	2.0	2D	WindSonic	Gill Instruments Ltd.	08460059
	1.0	2D	WindSonic	Gill Instruments Ltd.	08470015
EC tripod SE	3.0	3D	81000	R. M. Young Company	003149
	2.0	2D	WindSonic	Gill Instruments Ltd.	08460065
	1.0	2D	WindSonic	Gill Instruments Ltd.	08460074
EC tripod SW	3.0	3D	81000	R. M. Young Company	03154
	2.0	2D	WindSonic	Gill Instruments Ltd.	08460066
	1.0	2D	WindSonic	Gill Instruments Ltd.	08470017
EC tripod ICOS	3.0	3D	HS	Gill Instruments Ltd.	152903

Table 2. Stability categories and Pasquill–Gifford classes based on literature, as a function of bulk Richardson number (R_b), Obukhov length (L) and the temperature gradient with height.

Stability	Class	$f(R_\mathrm{b})$	$f(L)_P$	$f(L)_R$	$f(\Delta T/\Delta z)$ [K/100m]
Very Unstable	A	$R_\mathrm{b} < -0.86$	$-3.5 < L < 0$	$-100 < L < 0$	$T < -1.9$
Moderately Unstable	B	$-0.86 < R_\mathrm{b} < -0.37$	$-8 < L < -3.5$	$-200 < L < -100$	$-1.9 < T < -1.7$
Weakly Unstable	C	$-0.37 < R_\mathrm{b} < -0.1$	$-15 < L < -8$	$-500 < L < -200$	$-1.7 < T < -1.5$
Neutral	D	$-0.1 < R_\mathrm{b} < 0.053$	$L > 75 \mid L < -15$	$\mid L \mid > 500$	$-0.5 < T < 1.5$
Weakly Stable	E	$0.053 < R_\mathrm{b} < 0.134$	$35 < L < 75$	$200 < L < 500$	$1.5 < T < 4.0$
Moderately Stable	F	$R_\mathrm{b} > 0.134$	$8 < L < 35$	$50 < L < 200$	$T >= 4.0$
Very Stable	G	N/A	$0 < L < 8$	$0 < L < 50$	N/A
	Pasquill (1961)	Gifford (1976)	Pasquill and Smith (1971)	Gryning et al. (2007)	Woodward (1998)
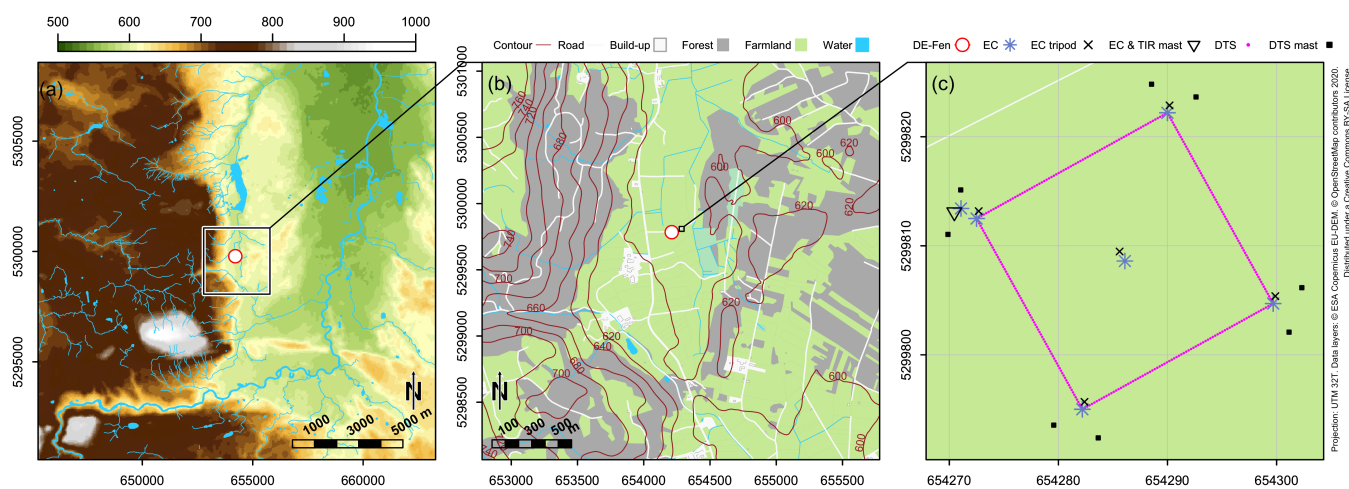
Figure 1. The location of the study site with aspects of (**a**) topography and runoff, (**b**) land use and (**c**) the sensor network set-up. Data layers: © ESA Copernicus EU-DEM. © OpenStreetMap contributors 2020. Distributed under a Creative Commons BY-SA License.
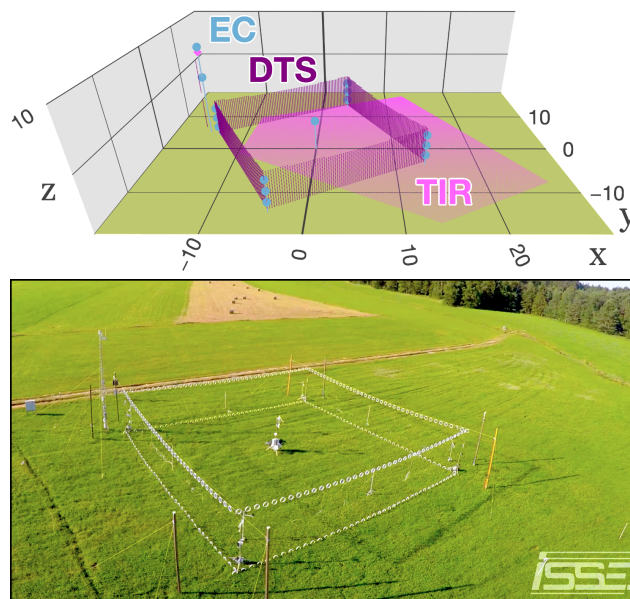
Figure 2. Overview of the set-up in (**top**) schematic perspective and (**bottom**) a bird-eye view from the South. Still image modified from an aerial video by M. Schörner (24 Jul 2016, ISSE/University of Augsburg).
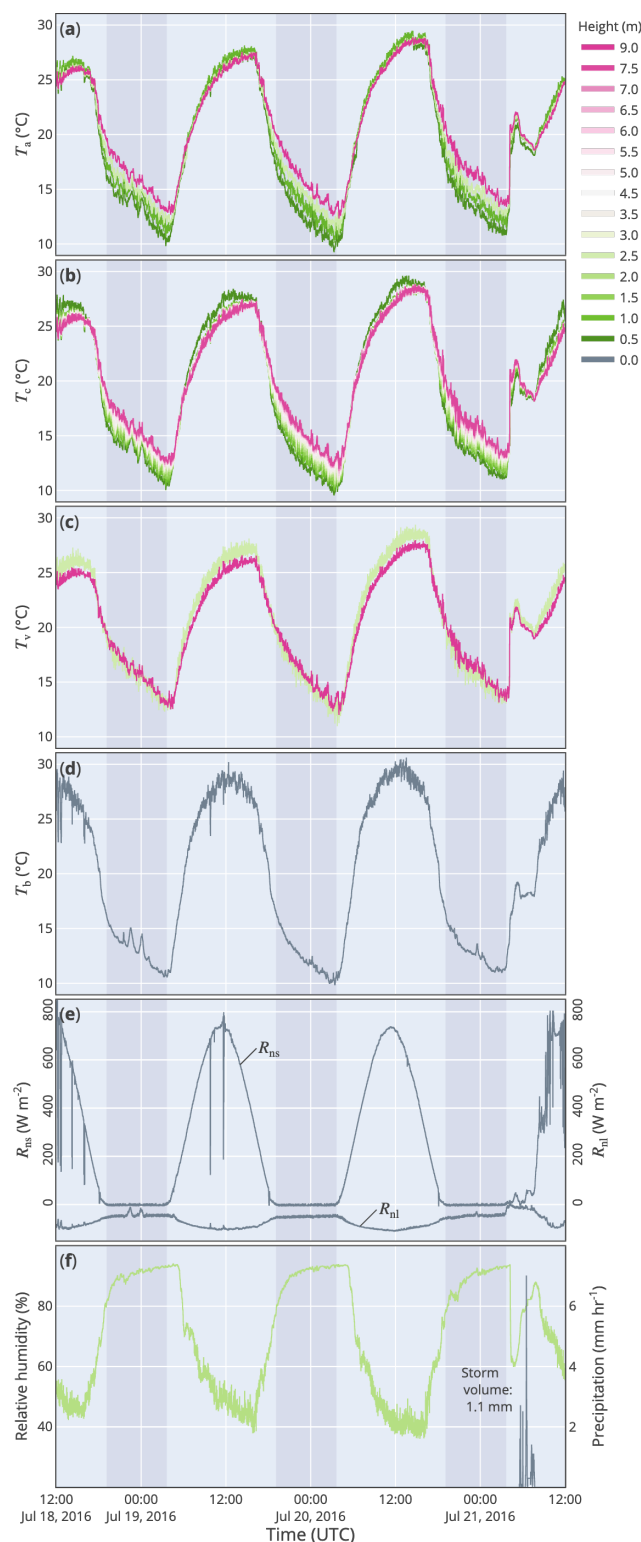
Figure 3. Reference period observations of (**a**) air temperature, (**b**) air temperature from fibre-topic cable (**c**), air temperature from the speed of sound, (**d**) surface brightness temperature from radiance (**e**) net short-wave radiation, R_{ns}, and and net long-wave radiation, R_{nl}, terms, and (**f**) relative humidity and precipitation.
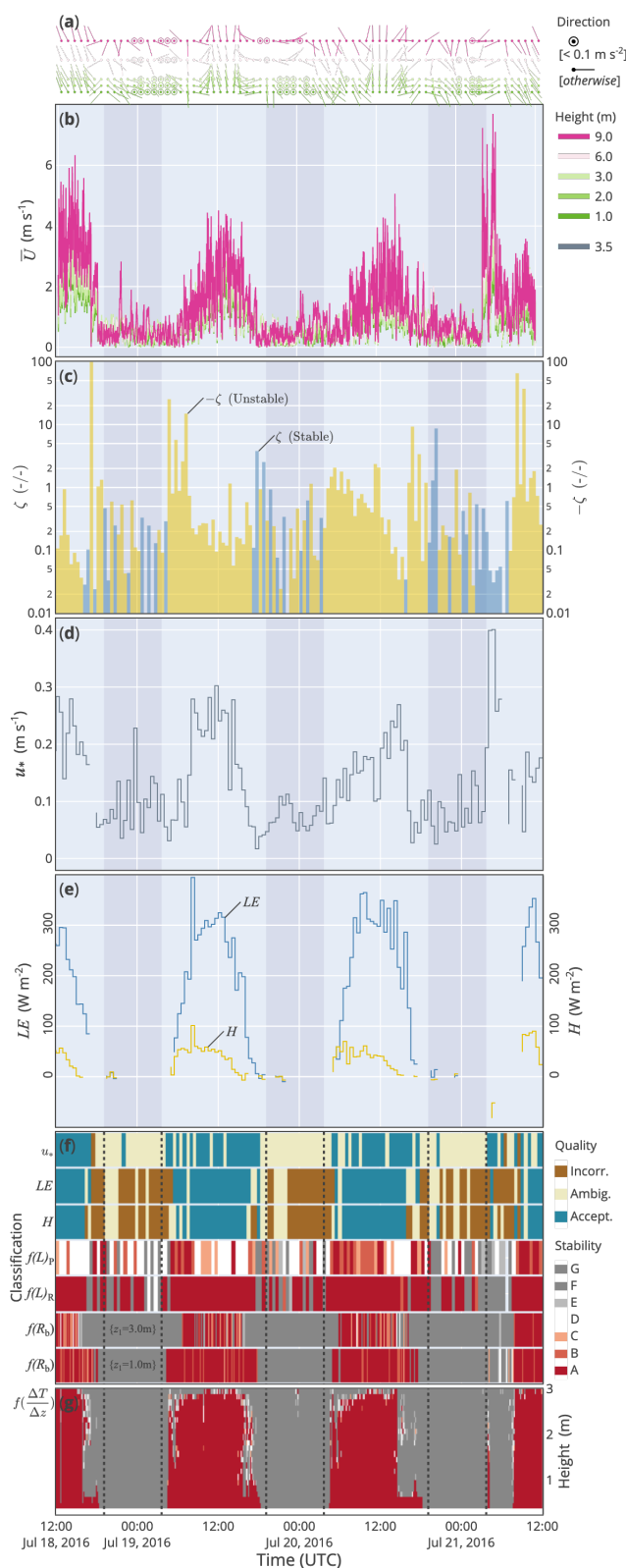
Figure 4. Reference period observations of (**a**) wind direction, (**b**) wind speed, (**c**) the stability length parameter, (**d**) the friction velocity, (**e**) latent heat (LE) and sensible heat (H) flux, (**f**) quality and stability classification and (**g**) stability against height. See text for details.
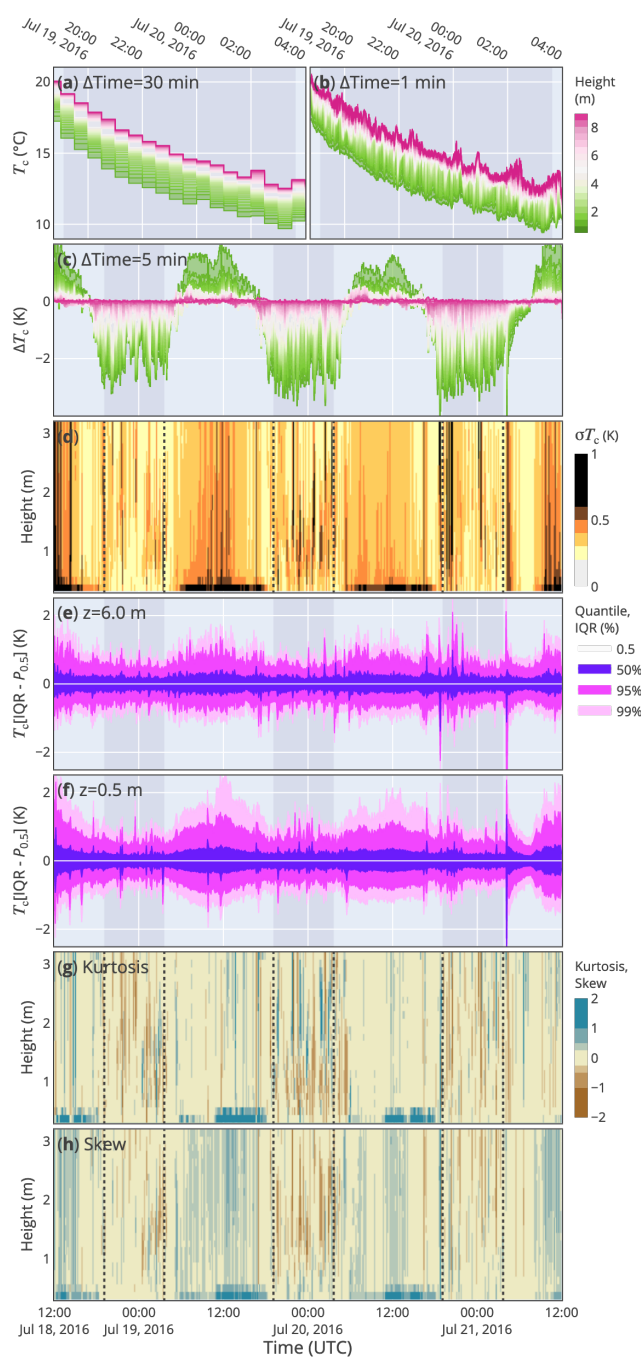
Figure 5. Shown against time are (**a**)(**b**) fibre optic cable temperature (T_c) for a selected night and 30 min and 1 min averaging, respectively, (**c**) the T_c gradient relative to the top of the profile, (**d**) the T_c variance against height, (**e**)(**f**) the inter-quartile range (IQR) relative to the median ($P_0.5$) for 6.0 and 0.5 m above ground, respectively, (g) the kurtosis and skewness of T_c against height, respectively.
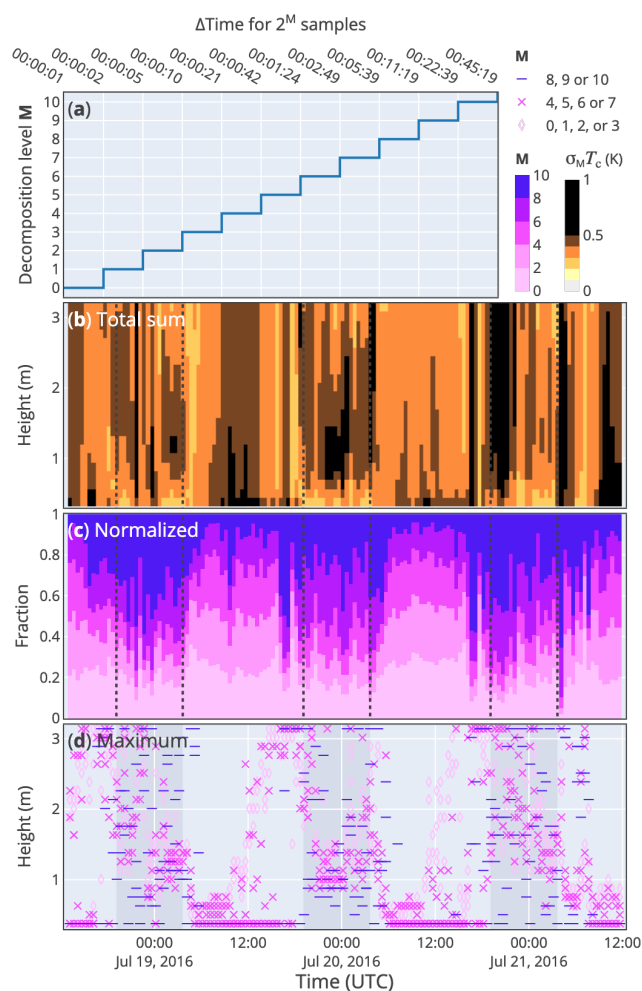
Figure 6. For the Multiresolution Decomposition (MRD) of T_c are shown (**a**) the decomposition levels (M), and against time and height (**b**) the total sum of variance, $\sigma_M T_c$, (**c**) the relative contribution of each level and (**d**) the height of maximum mean variance for three groups of M.
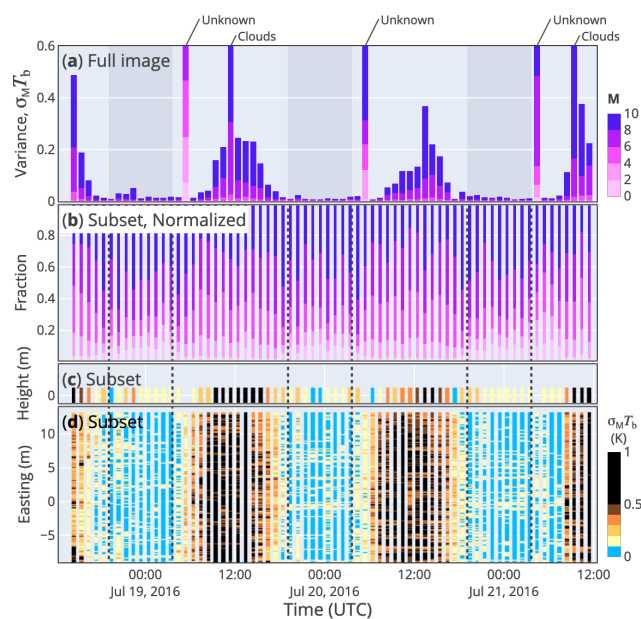
Figure 7. For the Multiresolution Decomposition (MRD) of T_b are shown against time (**a**) the sum of variance for the full thermal image, (**b**) the relative contribution of each decomposition level, (**c**) the total sum of variance, $\sigma_M T_b$, and (**d**) the total sum of variance per grid cell in a subset image area. The subset includes grid cells depicting vegetation within 1 m of a fibre-optic profile, for which the cell location is expressed relative to the center of the domain. See text for details.
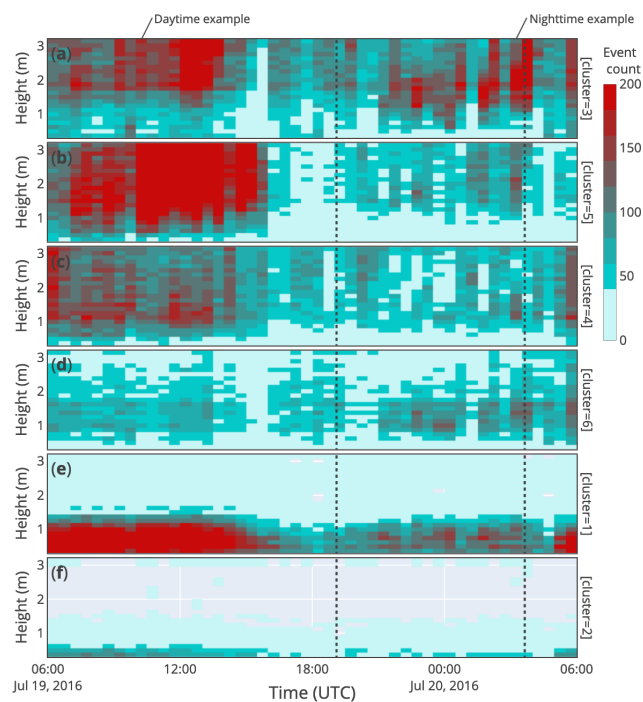
Figure 8. For each TED-determined cluster are shown the number of events per 30-min time interval and height. The clusters are ranked according to frequency, with least frequent in the bottom panel.
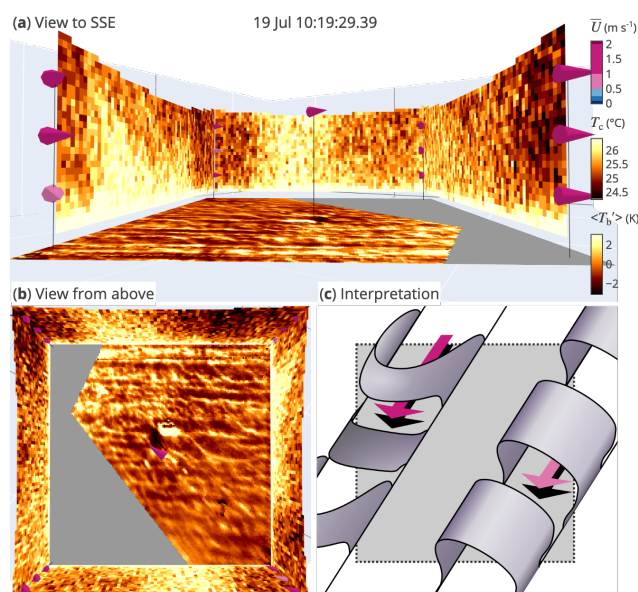
Figure 9. Daytime example of structures and motion in air and surface temperature represented by fibre-optic cable temperature (T_c) and surface brightness temperature (T_b) variance, respectively. Shown are (**a**) a viewpoint from NNW, (**b**) a perspective from above and (**c**) an conceptual interpretation. Observations of wind are shown as cones pointing in the direction of the mean wind ($\overline{U}$).
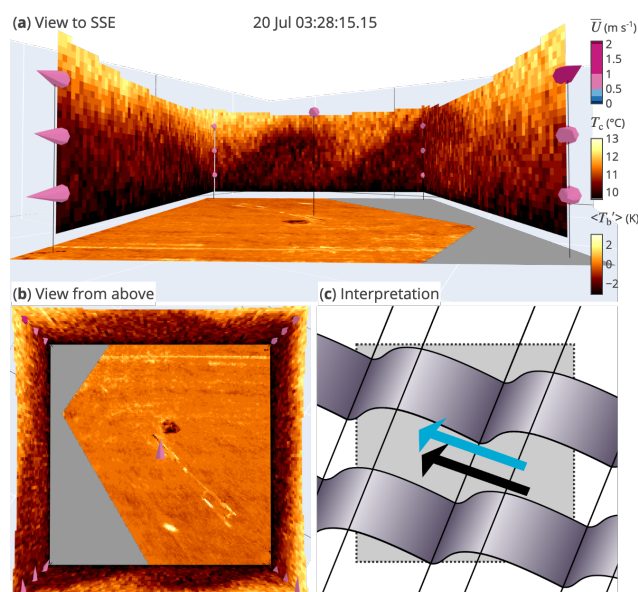
Figure 10. Night-time example of structures and motion in air and surface temperature represented by fibre-optic cable temperature (T_c) and surface brightness temperature (T_b) variance, respectively. Shown are (**a**) a viewpoint from NNW, (**b**) a perspective from above and (**c**) an conceptual interpretation. Observations of wind are shown as cones pointing in the direction of the mean wind ($\overline{U}$).
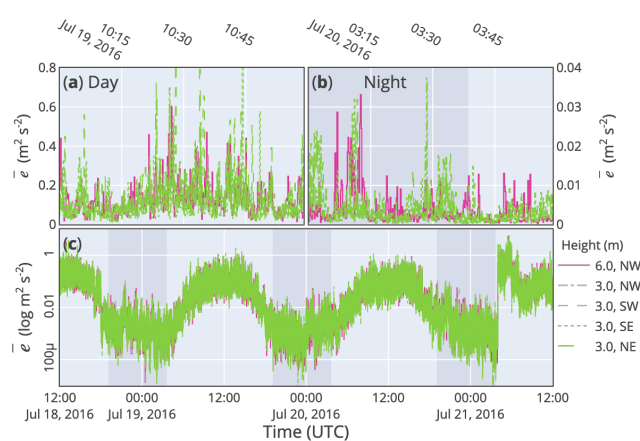
Figure 11. The turblent kinetic energy (TKE; $\overline{e}$) was computed for corner locations at 3.0 or 6.0 m above ground and shown for (**a**) an example daytime hour, (**b**) and example night-time to dusk hour and (c) the study period. Night-time is highlighted with dark areas.
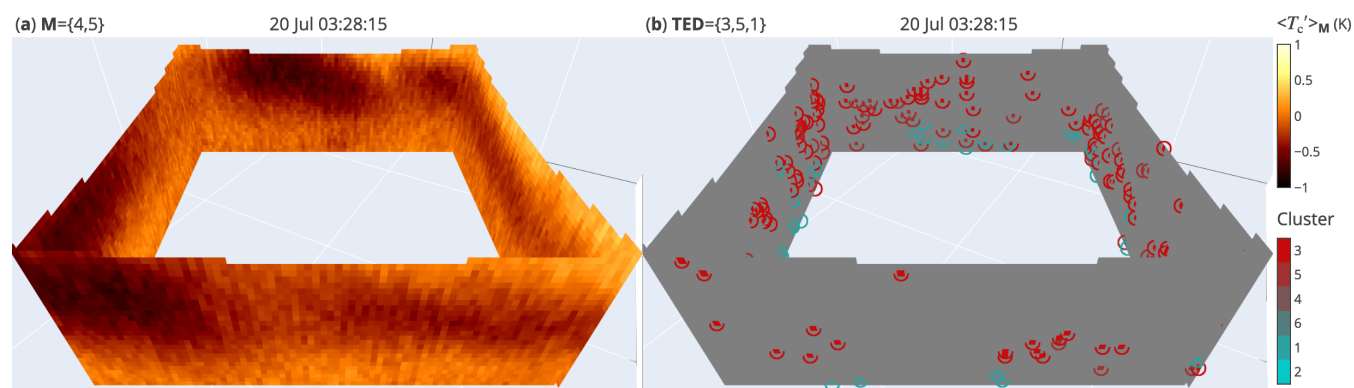
Figure 12. Shown for a night-time instance are (**a**) the variance for selected Multiresolution Decomposition levels (M) of T_c and (**b**) the dominant clusters corresponding to detected events within a 45 s window. See text for details.
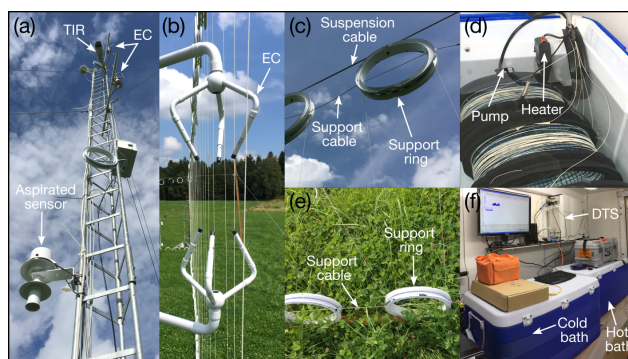
Figure A1. Details of the field set-up showing (**a**) a close-up of the 9 m mast, (**b**) a view along a vertical plane of DTS profiles, (**c**)(**e**) details of the DTS support structure, (**d**) a view inside a calibration bath and (**f**) the DTS instrumentation.
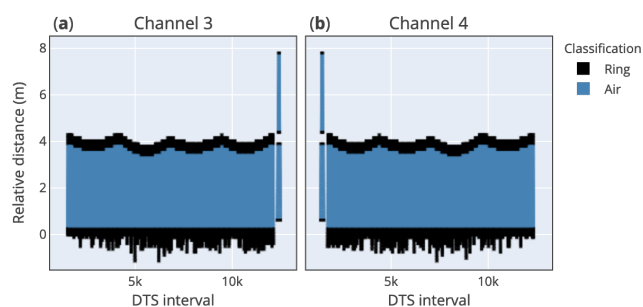
Figure A2. The distance relative to the lower support rings were determined for DTS intervals in (**a**)(**b**) opposite sensing direction. Intervals in contact with the rings were excluded from the analysis, leaving 8286 usable locations.
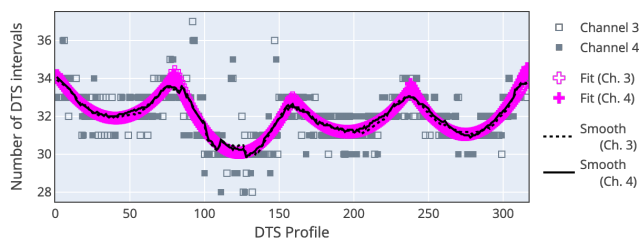
Figure A3. The alignment of DTS profiles was validated using a fit (hyperbole cosine) and smoothing (Savitzky-Golay) function.
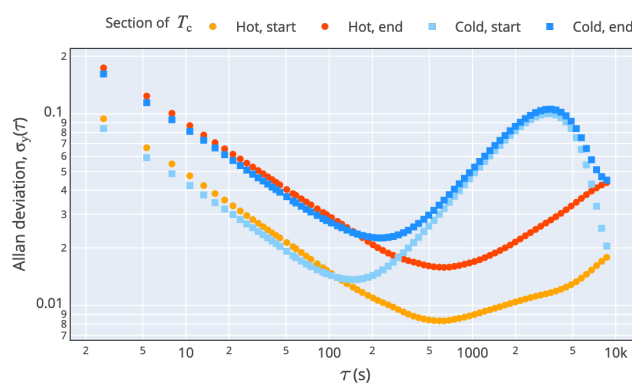
Figure A4. Allan deviation for four DTS range gates in different sections of the fibre-optic cable, in cold and hot baths and at the start and end of the cable, respectively. The −1 slope towards high frequency ($\tau < 100$ s) indicates white noise. Calculations were made before calibration on 24 hours of T_c data.
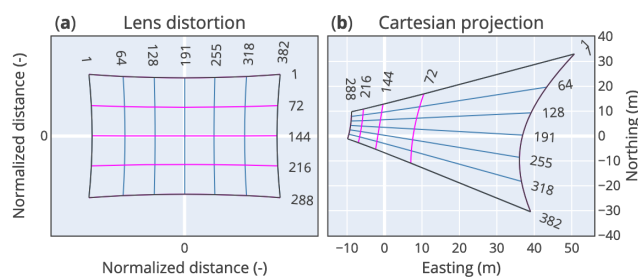
Figure A5. The projection of thermal images to real-world coordinates required (**a**) correction for lens optical distortion and (**b**) project after camera rotation. The numbers refer to pixel count relative to the image top-left corner.
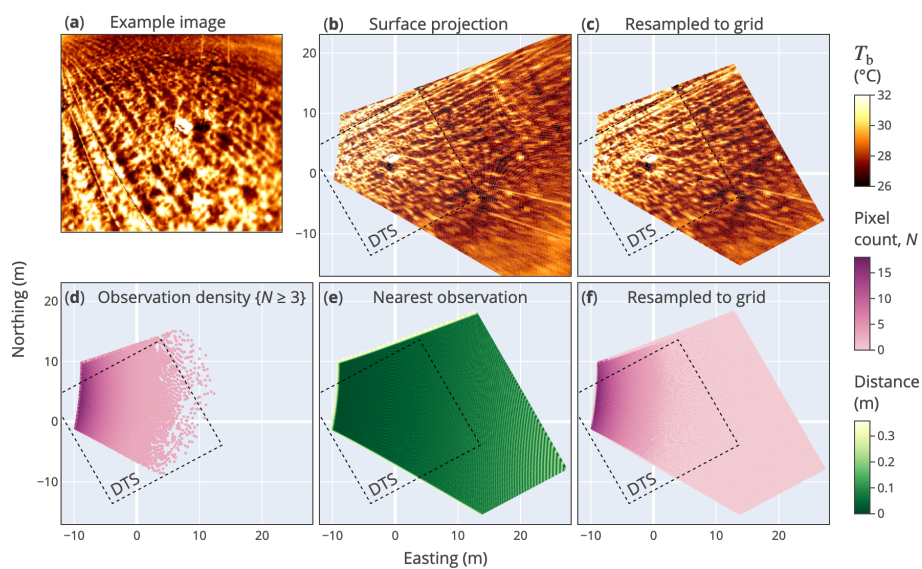
Figure A6. The (**a**)(**b**) thermal image data were (**c**)(**f**) spatially resampled in a regular grid of 0.25×0.25 m, matching the orientation of the fibre-optic array (DTS). At least (**d**) three observations should fall within a grid cell, or else (**e**) the nearest value within a 0.35 m radius was used for computation.
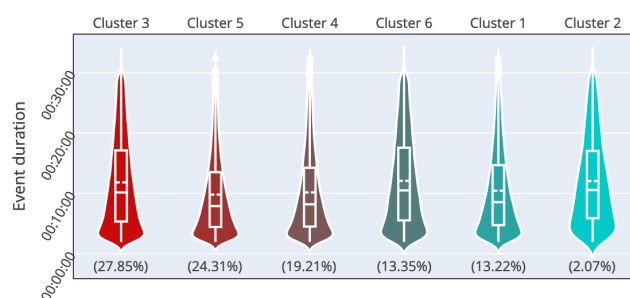
Figure A7. The TED method samples events in temperature time series and applies a k-means clustering on event statistics. Clusters are ranked by fraction of total events, shown as percentage below the violin-box plot distribution of the event duration.