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# Fragment ion-functional group relationships in organic aerosols using aerosol mass spectrometry and mid-infrared spectroscopy

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#### Abstract.

Aerosol mass spectrometry (AMS) and mid-infrared spectroscopy (MIR) are two analytical methods for characterizing the chemical composition of OMorganic matter (OM). While AMS provides high-temporal-resolution bulk measurements, the extensive fragmentation during the electron impact (EI)-ionization makes the characterization of OM components limited. The analysis of aerosols collected on PTFE filters using MIR, on the other hand, provides functional group (FG)-information with reduced sample alteration but results in a relatively low temporal resolution. In this work, we compared and combined MIR and AMS measurements for several environmental chamber experiments of combustion-related aerosols to achieve a better understanding of the AMS spectra and the OM chemical evolution by with aging. Fresh emissions of wood and coal burning were injected into an environmental simulation chamber and aged with hydroxyl and nitrate radicals. A high-resolution time-of-flight (HR-TOF) AMS measured the bulk chemical composition of fine PMOM. Fine aerosols were also sampled on PTFE filters before and after aging for the offline MIR analysis. After comparing AMS and MIR bulk measurements, we used multivariate statistics to identify the influential functional groups contributing to AMS OM mass functional groups associated the most with the AMS OM for different aerosol sources and aging processes oxidants. We also identified the key mass fragments resulting from fragment ions resulting from molecules containing each functional group for the complex OM generated from biomass and fossil fuel combustion. Finally, we developed a statistical model that enables the estimation of the high-time-resolution functional group composition of OM using collocated AMS and MIR measurements. Using this approach, AMS spectra can be used to interpolate the functional group measurements by MIR, allowing using this approach. The latter allows us to better understand the evolution of OM during the aging process.

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#### 1 Introduction

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Particulate matter (PM) impacts visibility, climate, and human health (Hallquist et al., 2009). Organic matter (OM), which accounts for an important fraction of total fine atmospheric PM mass, has profound effects on aerosol-related phenomena (Turpin and Lim, 2001; Russell, 2003; Shiraiwa et al., 2017). Characterizing the organic fraction is necessary to reduce the uncertainties associated with the impact of PM and can in turn affect the policies related to climate change and air quality management (Zhang et al., 2011; Turpin and Lim, 2001). However, OM chemical composition and Nonetheless, the chemical composition of OM and its formation mechanisms have not yet been fully understood due to their complexity.

Different analytical and computational techniques exist for the determination of the chemical composition of organic aerosols (OAs). Modeling all important SOA-related reaction reactions and species is not feasible for large-scale three-dimensional models (Jathar et al., 2015) and simpler models often do not simple models do not always reproduce the measured concentrations and dynamics of SOA evolution in polluted regions (Volkamer et al., 2006) of OM (Volkamer et al., 2006; Theodoritsi et al., 2020). Among the analytical techniques, aerosol mass spectrometry (AMS) and mid-infrared (MIR) spectroscopy are able to provide bulk chemical information for most of the OM mass (Hallquist et al., 2009).

AMS provides information about the chemical composition of OM and its temporal variations in terms of ensemble mass spectra acquired over short time intervals (Zhang et al., 2011). Aerodyne Research aerosol mass spectrometer (used in this work and referred to as "AMS") is the most widely used thermal desorption-based mass spectrometers thermal-desorption-based mass spectrometer in aerosol research. AMS is capable of quantifying non-refractory species (e.g., sulfate, nitrate, ammonium, chloride, and OM) by thermal vaporization (typically at 600 °C) and electron impact ionization (EI; typically at 70 eV) (Canagaratna et al., 2007). In spite of the valuable information that AMS provides, the AMS OM fragment ions are not molecule-specific and AMS spectra are difficult to interpret due to the extensive fragmentation of molecules with the highenergy electron impact (EI) ionization and flash vaporization. This limits the level of molecular details that can be extracted from the AMS mass spectra (Canagaratna et al., 2007; Kumar et al., 2018; Faber et al., 2017; Chhabra et al., 2011a). Organic acrosol OM components can also undergo oxidation, dehydration, and/or decarboxylation reactions inside the AMS ionization chamber (Canagaratna et al., 2015a). In addition, uncertainties regarding the relative ionization efficiency (Xu et al., 2018), fragmentation tables (Aiken et al., 2008), the gas-phase interference (Canagaratna et al., 2015a), and the collection efficiency (Frossard et al., 2014) have been reported. There are soft ionization methods, such as electrospray ionization (ESI) and chemical ionization (CI) that minimize the analyte fragmentation at the expense of the variable ionization efficiency and the signal-tonoise ratio, and quantifying bulk OM composition (Lopez-Hilfiker et al., 2019; Nozière et al., 2015; Iyer et al., 2016; Zahardis et al., 2011).

MIR spectroscopy, which is commonly performed off-line on polytetrafluoroethylene (PTFE) filters (Takahama et al., 2013; Ruthenburg (Maria et al., 2002; Takahama et al., 2013; Ruthenburg et al., 2014), is used as a complementary method to AMS in this work. This non-destructive method gives direct functional group (FG) information; information, provides measurements consistent with commonly used instruments in monitoring networks (Boris et al., 2019); and , and it is capable of differentiating be-

tween the composition of different oxidized OMs. For example, Liu et al. (2012) observed very similar AMS mass spectra for several SOA samples, while the complementary MIR spectra indicated clear chemical differences for these aerosols. In addition to FG functional group quantification, MIR spectroscopy has been recently used to quantify biomass burning markers in the atmospheric aerosols (e.g., levoglucosan and lignin; ?) (levoglucosan and lignin-like compounds; Yazdani et al., 2021b). However, MIR spectroscopy on filters has low temporal resolution compared to on-line online instruments such as AMS (Faber et al., 2017; ?) (Faber et al., 2017; Yazdani et al., 2021b). Moreover, the volatilization of volatile organic compounds from PTFE filters during or after sampling can affect the sampling period can affect the OM mass and composition (Subramanian et al., 2004). Uncertainties regarding variable absorptivities by different organic molecules the variable absorptivities of different organic molecules (Hastings et al., 1952), peak overlaps (Pavia et al., 2008), scattering, and light scattering by the filter membrane, and the PTFE interference (Takahama et al., 2013) have also been reported for this technique.

Past studies compared AMS and MIR OM, O:C, and positive matrix factorization (PMF) factors in field campaigns (Gilardoni et al., 2009; Russell et al., 2009b; Frossard et al., 2011; Liu et al., 2011; Corrigan et al., 2013; Frossard et al., 2014) reporting and reported reasonable agreement between the instruments despite the aforementioned uncertainties. Two controlled laboratory studies tried to understand the relationship between fragment ions and functional groups (Faber et al., 2017; Russell et al., 2009a) using univariate correlations. In this work, we compare and combine the AMS and MIR measurements for the organic aerosols OM of moderate to high complexity from biomass burning and coal combustion emissions, two major sources of atmospheric OM, in an environmental simulation chamber. We apply additional uni- and multivariate techniques to further interpret the relationships between more than 300 AMS fragment ions and 4 MIR FGsfunctional groups, and provide a method to predict the high-time-resolution evolution of FGs using AMS functional groups using only AMS spectra.

#### 2 Methods

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In the following sections, the experimental set-up (Sect. 2.1), on-line and off-line measurement techniques (Sects. 2.2 and 2.3) are described. Thereafter, statistical methods used for combining AMS and MIR measurements are explained (Sects. 2.4–2.6). The experimental set-up, procedure, and data used in this work are the same as those reported by **?**Yazdani et al. (2021b).

#### 2.1 Laboratory experimental set-up and procedure

Briefly, we conducted four wood burning (WB) experiments with beech wood logs and five coal combustion (CC) experiments with bituminous coal using ordinary modern stoves (Bruns et al., 2015). The emissions were diluted and then injected into a 6 m<sup>3</sup> Teflon bag environmental chamber of at the Paul Scherrer Institute (PSI) in Villigen, Switzerland. The injections were continued until the concentration of PM and OM measured by the seanning mobility particle sizer (SMPS) and high-resolution time-of-flight (HR-ToF) AMS reached atmospherically-relevant values. The emissions were held in the chamber for 30 minutes after injection to improve mixing. Thereafter, primary emissions were chemically aged using the hydroxyl or nitrate radical in order to simulate daytime and nighttime aging mechanisms, respectively. For the diurnal aging simulations, the OH radical was produced by the photolysis of HONO and the OH exposures reached (2–3)×10<sup>7</sup> molec cm<sup>-3</sup>h (measured using butanol-d9;

Barmet et al., 2012) corresponding to 20–30 hours of aging in the atmosphere. For the nocturnal aging experiments, the NO<sub>3</sub> radical was produced by a single injection of O<sub>3</sub> and NO<sub>2</sub> in the chamber. The nitrate radical concentration was estimated to be  $(1.5-2.5)\times10^7$  molec cm<sup>-3</sup> for the first hour of aging process based on the phenol concentration decay in the gas phase. There are in total four experiment categories (two different fuels and oxidants indicated by WB\_OH, WB\_NO<sub>3</sub>, CC\_OH, and CC\_NO<sub>3</sub>) with one to three similar experiments in each category.

#### 2.2 Online AMS PM measurements

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Non-refractory particle composition was measured with a HR-ToF HR-TOF AMS operating in V mode (mass resolution  $\Delta m/m = 2100$  for m/z 200; DeCarlo et al., 2006) with a 2.5 µm inlet aerodynamiclens throughout the experiment. The raw signal was postprocessed in Igor Pro 6.3 (Wave Metrics) using SQUIRREL 1.57 and PIKA 1.15Z routines. Elemental ratios of OM were estimated following the approach of Canagaratna et al. (2015a) (Fig. S1). The AMS OM concentrations reported in this work are not corrected for the chamber wall losses and the measured nitrate is assumed to be inorganic for ease of comparison with MIR.

#### 2.3 Offline MIR PM measurements

Two 47 mm Teflon filters (Pall corporation) were used for each experiment to sample the primary PM after its injection into the chamber and the aged PM after approximately three to four hours of aging. The filter holder was placed downstream of a sharp-cut-off cyclone and a silica gel denuder and the flow rate through the filter was maintained at 8 L min<sup>-1</sup>. We used a similar naming convention for the filters to that of ?Yazdani et al. (2021b). Filters were immediately stored in filter petri dishes at 253 K after sampling and before MIR analysis to minimize volatilization and chemical reactions. The PTFE filters were analyzed using a Bruker Vertex 80 Fourier transform infrared (FT-IR) instrument equipped with an α deuterated lanthanum alanine doped triglycine sulfate (DLaTGS) detector, at a resolution of 4 cm<sup>-1</sup>. The spectra were averaged over 64 scans.

#### 2.3.1 MIR Spectral spectral postprocessing

The MIR spectra were baseline-corrected to eliminate the contribution of light scattering by filter membrane and particles as well as absorption by graphitic carbon (Parks et al., 2021). We used a smoothing spline smoothing splines described by Kuzmiakova et al. (2016). After baseline correction, the scaled and baseline-corrected spectrum of a blank filter was subtracted from the baseline-corrected sample spectra-blank subtraction was performed to minimized the interference of PTFE C-F bands. After baseline correction and blank subtraction, the Yazdani et al. (2021b). The multiple peak-fitting algorithm described by Takahama et al. (2013) was applied on the spectra to obtain FG to obtain functional group abundances of alcohol (aCOH), carboxylic acid (COOH), alkane (aCH), non-acid carbonyl (naCO) (Supplement Fig. S1). After obtaining FG functional group abundances, the O:C, H:C, and OM:OC ratios (Fig. S1) were calculated assuming 0.5 C atom for each of aCH and aCOH bonds (Chhabra et al., 2011b; Russell, 2003; Maria et al., 2002) (Maria et al., 2002; Russell, 2003; Reggente et al., 2019b).

### 2.4 Identifying influential MIR absorbances for AMS OM

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The AMS OM estimates and the MIR spectra are combined statistically to identify the functional groups that influence the former are the best predictors of AMS OM mass concentrationthe most. The results of this method, which are affected neither by. This method is not affected by either uncertainties of MIR peak fitting nor by absorption coefficients, identify important predictors of OM estimated by more routinely applied methods prior to peak fitting. This technique or absorption coefficients and can be applied even when absorption coefficient data are not available for all FGs, and help decide which FGs functional groups. It also helps decide which functional groups are needed to be included in the MIR peak fitting. The averaged AMS OM concentrations over the filter sampling periods were regressed against the corresponding MIR spectra using partial least squares regression (PLSR). Thereafter, the influential absorbances in the MIR spectra for the organic OM concentration were determined based on the variable importance in projection (VIP) scores of the PLSR model method(Fig. S2). This procedure was applied separately for the primary and aged aerosols of each source to highlight the compositional differences. By When regressing AMS OM concentration concentrations against MIR absorbances, we seek a solution of the following linear equation for coefficients a:

$$150 \quad y = Xa + e, \tag{1}$$

where  $\mathbf{X}$   $(n \times p)$  is the MIR spectra matrix with n samples and p independent variables (wavenumbers),  $\mathbf{y}$   $(n \times 1)$  is the vector of the response variable, the AMS OM concentration, and  $\mathbf{e}$  is the vector of residuals. In this work, the univariate partial least squares regression (Wold et al., 1983) is used. The univariate PLSR projects  $\mathbf{X}$  onto  $\mathbf{P}$   $(p \times h)$  (h is the number of latent variables) basis with orthogonal scores  $\mathbf{T}$   $(n \times h)$ , while maximizing the covariance between scores and the response variable,  $\mathbf{y}$ . In Eq. (3),  $\mathbf{b}$  and  $\mathbf{f}$  indicate the regression coefficients and the vector of residuals, respectively.

$$\mathbf{X} = \mathbf{T}\mathbf{P}^{\top} + \mathbf{E},\tag{2}$$

$$y = \mathbf{T}b + f. \tag{3}$$

After solving the regression equation using PLSR for different number of latent variable (LVs), a A repeated 10-fold cross validation was applied to find the optimal number of latent variables and avoid under/overfitting(LV) for the PLSR model. Examining loadings and coefficients directly can be informative about the important absorbances. For instance, the first weight vector,  $w_1$ , can be a good estimate of important bands but it is limited to the cases that signal is not dominated by other factors rather than the analyte, such as inorganics, and filter absorption (Haaland and Thomas, 1988). In this work, we used a more general method, VIP scores (Wold et al., 1993), to identify the important absorption bands. This metric is a root mean square of loading weights of all h latent variables used in the model weighted by their fraction of the captured response (Chong and Jun, 2005; Takahama et al., 2016). The VIP score of the jth wavenumber is calculated by considering all h latent variables in the model as shown in Eq. (4). Since the average of squared VIP scores is equal to one, generally, the wavenumbers with VIP score greater than one are considered influential due to higher-than-average contribution to estimating the response variable. The influential functional groups are those associated the most with the AMS OM or are the functional groups that are the best

predictors of AMS OM. In Eq. (4),  $t_k$  and  $w_k$  represent the kth columns of T, score matrix, and W ( $p \times h$ ), weight matrix, respectively and the. The relationship between T and W is described by Eq. (5) (Helland, 1988).

$$VIP_{j} = \sqrt{p \frac{\sum_{k=1}^{h} SS(b_{k} \boldsymbol{t}_{k}) (\boldsymbol{w}_{jk} / \|\boldsymbol{w}_{k}\|)^{2}}{\sum_{k=1}^{h} SS(b_{k} \boldsymbol{t}_{k})}},$$
where  $SS(b_{k} \boldsymbol{t}_{k}) = b_{k}^{2} \boldsymbol{t}_{k}^{\top} \boldsymbol{t}_{k}.$ 

$$(4)$$

$$\mathbf{T} = \mathbf{X}\mathbf{W}(\mathbf{P}^{\top}\mathbf{W})^{-1}.$$

# 2.5 Identifying FG-ion functional group-ion fragment relationships

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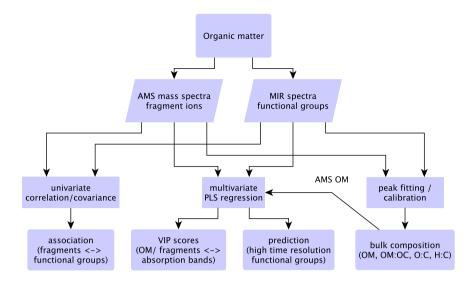
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Covariance and correlation coefficients were used to understand the connection between fragment ions (up to m/z 202 for which the signal-to-noise ratios are ratio is still significant) and FGs functional groups (Fig. S3a). We used normalized functional group abundances by the MIR total OM and normalized fragment ion concentrations by the AMS total OM (averaged over filter sampling periods) and calculated covariances and correlations between 4 FGs functional groups and more than 300 fragment ions. The major difference between the fragment-FG correlation and covariance is that the former is more informative about the fragments with low concentrations while the latter highlights the fragments with higher concentrations. Data normalization was performed to avoid correlations introduced by the changes in the total OM mass concentration (e.g., due to SOA condensation) as oxygenated fragments are highly correlated before normalization (Fig. S5). In addition, negative Negative correlations (anti-correlations) and covariances were omitted as they do not show the production of fragments by FGs from molecules containing the functional group of interest. For example, often the aCH relative abundance decreases with aging as the relative concentrations of oxygenated FGs functional groups such as COOH and oxygenated fragment ions such as CO<sub>2</sub>+ increase, leading to a significant anti-correlation between the aCH FG functional group and the CO<sub>2</sub>+ fragment. Russell et al. (2009a) and Faber et al. (2017) have already applied the univariate fragment-FG (correlation) analysis for different sources using unit-mass-resolution and HR AMS data, respectively. However, their analysis has been limited to only a few light small fragment ions.

Univariate methods can be difficult to interpret when ion fragments are produced by associated with multiple functional groups. Therefore, in In addition to the univariate methods, the VIP scores method was used to highlight the influential spectral regions and FGs functional groups for major fragment ions with high concentrations ( $CO_2^+$ ,  $CHO^+$ ,  $C_2H_3O^+$ ,  $C_3H_5^+$ ) and for two biomass-burning-related fragment ions ( $C_2H_4O_2^+$  for levoglucosan and  $C_8H_9O_2^+$  for lignin). This multivariate approach  $\frac{1}{2}$  which is not affected by the MIR peak fitting uncertainties, is is similar to identifying the influential MIR spectral regions for the AMS OM as discussed in Sect. 2.4 except that the concentrations of individual fragment ions are regressed against the MIR spectra (Fig. S3b).

#### 2.6 Interpolating FG functional group abundances using AMS mass spectra



**Figure 1.** Statistical relations and strategy for comparison of MIR and AMS measurements. The correlation/covariance analysis is performed between AMS mass fragments and MIR functional group abundances from peak fitting. The PLS regression PLSR is performed using the AMS total OM or individual fragment concentrations as the independent variable and the MIR absorbance spectra as the dependent variables.

In order to estimate the high-time-resolution FG composition of OM, FG Functional group abundances for all filters (normalized by the MIR total OM mass concentration), which were calculated from peak fitting, were regressed against the AMS spectra (normalized by the AMS total OM mass concentration and averaged over the filter sampling periods) using PLSR. A repeated 10-fold cross validation was applied to indicate the optimal number of latent variables. These models were then used to interpolate (high-time-resolution) FG compositions using them functional group compositions using the AMS spectra and to investigate the evolution OM during the course of oxidation when only AMS measurements existed (Fig. S4). Thereafter, the high-resolution The contribution of oxygenated functional groups to the bulk O:C ratios separated by FG contribution ratio was calculated from their high-time-resolution abundances (O:C =  $O_{COOH}$ :  $C_{total} + O_{naCO}$ :  $C_{total} + O_{aCOH}$ :  $C_{total}$ ) were calculated from the high-resolution resolution FG compositions following the same approach of Sect. 2.3.1., where  $C_{total} = C_{aCOH} + C_{aCOH} + C_{aCOOH} + C_{naCO}$ .

#### 3 Results and discussions

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In the following subsections, bulk OM parameters from AMS and MIR are combined and compared in Sect. 3.1. There after, fragment ion-FG relationships are investigated in Sect. 3.2. Finally, PLSR models are developed to predict FG-functional group composition of OM using the AMS mass spectra (Sect. 3.3). Our approach for combining and comparing the AMS and MIR measurements is demonstrated in the diagram of Fig. 1.

#### 3.1 Combination and comparison of OM measurements

Influential spectral regions of the MIR spectra and their corresponding FGs functional groups are determined for the AMS OM using VIP scores (Sect. 3.1.1). Thereafter, the OM mass concentration, OM:OC, O:C, and H:C ratios calculated using peak fitting to MIR spectra are compared to the average values from AMS for the primary and aged aerosols (Sects. 3.1.2, 3.1.3, and 3.1.4)

#### 3.1.1 Influential group frequencies for total AMS OM

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VIP scores of the PLSR models regressing MIR absorbances against AMS OM mass concentration highlight certain FGs to be important regarding the OM mass for each fuel burned and aerosol age. As can be seen from Fig. 2, carbonyl CO. aCOH have the highest VIP scores (greater than one) for primary WB wood burning aerosols, highlighting their importance in the primary WB wood burning OM. The high VIP scores of aCOH is consistent with the fact that aCOH it is a major part of wood constituents. Although  $\nu(CH_2)$  and  $\nu(CH_3)$  (stretching vibrational modes) at 2800–3000 cm<sup>-1</sup> do not have high VIP scores for primary WB-wood burning aerosols, the VIP scores for  $\delta(CH_2)$  at 1470 cm<sup>-1</sup> (bending vibrations) are high, 225 suggesting the importance of aCH. The PLSR model probably use the information from the bending peak due to the lower overlap with other bands in the primary WB spectra. The peak peak around 1600 cm<sup>-1</sup> which has a greater-than-one VIP score for primary WB-wood burning is the result of several overlapping peaks attributed to the organonitrates, aromatic ring  $\nu$ (C=C), amine  $\delta$ (N-H), amide  $\delta$ (N-H), and carboxylate  $\nu$ (C=O) (Pavia et al., 2008). These overlapping absorbances make 230 peak assignment in this region uncertain and complex. However this This peak is accompanied by the lignin-related  $\nu(C=C)$ vibrations at 1515 cm<sup>-1</sup> (?)(Yazdani et al., 2021b), suggesting the abundance of lignin-like products in primary WB-wood burning OM as also proposed by Bertrand et al. (2018). For the aged WB-wood burning aerosols, VIP scores are the highest for the broad carboxylic  $\nu(OH)$  absorbances at 2400–3400 cm<sup>-1</sup> and the carbonyl  $\nu(CO)$  (acid carbonyl) at 1700 cm<sup>-1</sup>, suggesting carboxylic acids to be important contributors to the OM mass after SOA formation (?). In contrast to the primary WB aerosols, 235 aCOH (Yazdani et al., 2021b). The aCOH group does not have high VIP scores for the aged WB acrosol (?) wood burning aerosol (Yazdani et al., 2021b).

For the primary  $\[CC]$  coal combustion aerosols,  $\nu(CH_2)$  has the highest VIP scores, suggesting the abundance of hydrocarbons likely from volatile compounds of coal. The aromatic ring  $\nu(C=C)$  peak at 1600 cm<sup>-1</sup>, however, has relatively lower VIP scores, implying that the aromatic rings do not constitute the majority of primary  $\[CC]$  OM although coal is mainly composed of highly substituted aromatics. The ammonium  $\nu(N-H)$  peaks at 3200–3400 cm<sup>-1</sup> have negative coefficients and high VIP scores, implying that the PLSR model compensates ammonium interference with organics by assigning negative coefficients to the region. coal combustion OM. For the aged  $\[CC]$  coal combustion aerosols, which are mostly composed of  $\[CC]$  SOA, the VIP scores of  $\nu(CH_2)$  are considerably lower. By contrast, carbonyl  $\nu(CO)$  and aCOH regions bands (observed on the shoulder of  $\nu(N-H)$  peaks) have the highest VIP scores, suggesting that  $\[CC]$  the SOA is mostly composed of carbonyls and alcohols. The out-of plane aromatic CH band,  $\gamma(CH)$ , at 750 cm<sup>-1</sup>, in spite of being visible, does not have high VIP scores in spite of their positive coefficient, suggesting that aromatic CH (rCH) is not a major part-constituent of the aged  $\[CC]$  Caerosols. The

RONO<sub>2</sub> absorption region at 1630 cm<sup>-1</sup> does not have high VIP scores although this region is very prominent in the aerosols aged with the nitrate radical (?). This is because the AMS OM concentrations used in this study do not consider the majority organonitrate mass as both  $NO_2^+$  and  $NO^+$  are attributed to inorganic nitrate. As a result, the organonitrate abundances do not affect the regression models.

In general, the important FGs highlighted by the VIP scores method functional groups associated the most with the AMS are the same ones targeted in past studies of atmospheric aerosols (e.g., Ruthenburg et al., 2014; Russell et al., 2009b) and are consistent with our knowledge of POAs and SOAs related to combustion sources (e.g., Bertrand et al., 2018, 2017; ?) (e.g., Bertrand et al., 2017, 2018; Li et al., 2020; Yazdani et al., 2021b). In addition, they provide insight into the fraction of the combusted fuel that is important for OM formation (e. g., hydrocarbons in CC). chemical difference between the unburned fuel and the POA and SOA.

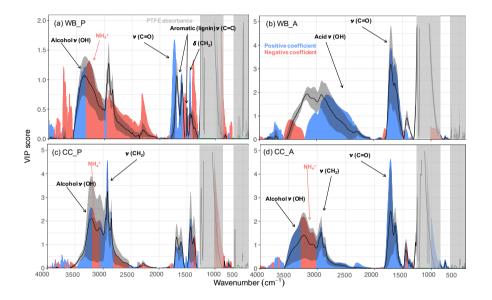


Figure 2. VIP scores of the MIR absorbances regressed against the AMS OM concentration (averaged over the filter sampling periods) for (a) primary wood burning, WB\_P (ba), aged wood burning, WB\_A (eb), primary coal combustion, and CC\_P (dc), and aged coal combustion aerosols, CC\_A (d). Blue/red regions correspond to wavenumbers with positive/negative regression coefficients in the PLSR models, respectively. Solid curves show the average spectrum (± one standard deviation shown by the shaded bands) for each category. Important FGs and their locations functional groups are indicated for each category. The PTFE C—F absorption interference regions are masked by gray rectangles.

#### 3.1.2 AMS and MIR OM mass concentrations

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It was shown in the last section that four FGsfunctional groups, aCH, COOH, aCOH, and naCO non-acid carbonyl (naCO) are the most influential functional groups regarding OM mass. The abundances of the mentioned FGs functional groups were estimated using peak fitting to the MIR spectra. The aromatic C=C group (rC=C) was not quantified due to the in-

terference with other functional groups and the lack of absorption coefficient data. The peak-fitting results show that the OM concentration estimates from AMS and MIR are highly correlated (Fig. 3a,  $R^2 = 0.92$ ). The slope of the MIR OM concentration versus that of AMS (not corrected for collection efficiency) is 1.3. This slope is within the previously reported range (Gilardoni et al., 2009; Russell et al., 2009a, b; Liu et al., 2011) considering the collection efficiency of AMS (?Kumar et al., 2018; Canagaratna et al., 2007) (Yazdani et al., 2021b; Kumar et al., 2018; Canagaratna et al., 2007) and the aerosol volatilization artifacts from PTFE filters (Ruggeri, 2017; Subramanian et al., 2004). ? Yazdani et al. (2021b) reported the AMS the collection efficiency to range between 0.7 and 1.1 for the same experiments. The OM concentrations estimated by both methods indicate the significant enhancement with aging even without particle and vapor wall loss consideration (on average 2.4 and 2.7 times by AMS and MIR, respectively). The enhancement ratios are in the range of values that were previously reported for SOA formation from logwood stoves (Bertrand et al., 2017; Tiitta et al., 2016; Grieshop et al., 2009; Heringa et al., 2011; Hennigan et al., 2010). Using different absorption coefficient values for MIR FGs-functional groups (discussed by Reggente et al., 2019a) has little effect on the correlation coefficient.

#### 3.1.3 AMS and MIR OM:OC ratios

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The OM:OC ratios calculated from the AMS mass spectra were averaged over the filter sampling periods and compared to those measured from peak fitting to MIR spectra. The OM:OC estimates of these two methods agree very well ( $R^2 = 0.82$ and slope = 0.99; Fig. 3b) with an average difference of less than  $0.15 \stackrel{?}{(?)}$  (Yazdani et al., 2021b). The correlation coefficient of OM:OC ratios is also found to be insensitive to the choice of absorption coefficients reported by Reggente et al. (2019a) for MIR spectroscopy. The fact that both methods capture similar OM:OC and mass concentration trends, suggests that a similar fraction of OM is monitored by both and the uncertainties associated with each method is less than variations due to fuel sources and aging processes. The primary CC coal combustion aerosols are estimated to have the lowest OM:OC ratios (1.35–1.5), which is justified by their strong hydrocarbon (aCH) signatures (Fig. 2c). The primary WB-wood burning samples have slightly higher OM:OC ratios (1.6–1.7 from AMS) primarily due to a relatively higher concentration of aCOH (Fig. 2a). Both instruments estimate that the aged aerosols of the two sources, regardless of the aging method, have higher OM:OC than the primary ones (Fig. 3b). The aged WB-wood burning aerosols have the highest OM:OC ranging from 1.9 to 2.1 (from AMS), with high concentrations of COOH(from MIR). The aged CC-coal combustion aerosols have lower average OM:OC ratios compared to the aged WB-wood burning aerosols, ranging from 1.6 to 1.8. For both emission sources, AMS and MIR show that aerosols aged with the hydroxyl radical have higher OM:OC ratios than those aged with the nitrate radical (Fig. 3b). Attributing the total AMS nitrate to organics to estimate an upper bound for OM:OC, makes this difference less prominent. However, the nitrate radical only reacts efficiently with certain precursors compared to the hydroxyl radical, resulting in different SOA composition that is reflected in both AMS and MIR measurements.

#### 3.1.4 AMS and MIR van Krevelen trajectories

The local slope of the aging trajectory in the van Krevelen diagram is informative about the changes in the functionality of OM (Heald et al., 2010; Ziemann and Atkinson, 2012; Chhabra et al., 2011a), which is also directly measured with MIR

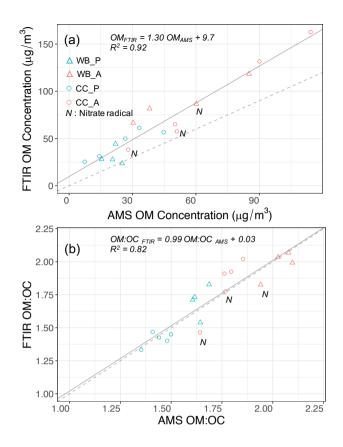


Figure 3. (a) Scatter plot comparing OM concentration (a) and OM:OC (b) estimates by AMS and MIR for primary (P) and aged (A) aerosols of wood burning (WB) and coal combustion (CC) the environmental chamber experiments. (b) Scatter plot comparing OM:OC estimates from AMS and MIR for primary and aged aerosols in each of the environmental chamber experiments. The fitted and 1:1 lines are solid and dashed, respectively.

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spectroscopy. Figure 4 shows the van Krevelen diagram of the WB and CC OM in different experiments. In the WB. In the wood burning experiments, AMS oxidation trajectories vary between a straight line and a convex (L-shaped) curve - (Fig. 4). In the first WB wood burning experiment with the hydroxyl radical (WB\_OH\_1), AMS aging trajectory is almost a straight line, implying a monotonic change of FGs functional groups during the course of aging (Fig. 4a). In the second experiment (WB\_OH\_2), however, similar to the WB\_OH\_1 experiment) the trajectory is convex with a reduced slope toward the end of aging, implying an increase in the abundance of FGs resulting functional groups that result in a low trajectory slope, e.g. acids (Fig. 4b). This is supported by the high concentration of the COOH group observed in the MIR spectra of the corresponding aged aerosols (?) (Yazdani et al., 2021b) and will be investigated further in Sect. 3.3. The WB wood burning experiment with the nitrate radical (WB\_NO3\_1) has a slope close to zero with a lower final O:C probably due to the exclusion of the organonitrate group and organonitrates and the different SOA formation reactions of the nitrate radical. The relatively low-small decrease in H:C with aging is supported by a relatively lower decrease of athe prominent aCH absorptions in the MIR

spectra of WB-wood burning aerosols that are aged with the nitrate radical (?)(Yazdani et al., 2021b). The modest decrease in H:C with aging is observed to be a characteristic of aging with the nitrate radical regardless of the emission source (Fig. 4d, h, and i), suggesting a more effective H atom abstraction by OH.

The starting points of the WB-wood burning oxidation trajectories (from AMS) have H:C ratios in the range of 1.6–1.8 and O:C ratios in the range of 0.3–0.4. The ending points have H:C ratios in the range 1.4–1.6 and O:C ratios in the range of 0.6–0.7. The observed values are close to that of OA-OM measured by Chhabra et al. (2011a), the OA-OM emissions of logwood combustions by Tiitta et al. (2016), and ambient OA-OM reported by Heald et al. (2010). The average O:C and H:C trends calculated from MIR spectroscopy are generally consistent with that of AMS, showing a decline in H:C ratio and increase in O:C ratio. However, there is an offset in the absolute values; in general, H:C is estimated to be approximately 0.2 higher by MIR spectroscopy both for the primary and aged WB aerosols. wood burning aerosols. The positive H:C offset for the wood burning aerosols might be due to the uncharacterized carbon from aromatic rings of lignin pyrolysis products, leading in to an overestimation of H:C.

For the CC coal combustion experiments, AMS oxidation trajectories usually start at H:C ratios around 1.7–1.9 (higher than that of WB supported by the strong aCH signature; Fig. 2ewood burning) and O:C ratios around 0.20–0.25 and end usually at H:C and O:C ratios around 1.5–1.7 and 0.35–0.55, respectively. The high H:C ratios before aging support the low amount of aromatics to aliphatic CH observed from are consistent with the high ratio of aliphatic CH to aromatic carbon observed in MIR spectra.. In most of the experiments, a positive curvature in trajectory is observed (Fig. 4e-i), implying a change in the type of FGs-functional groups produced during the course of aging. The average slopes are close to –1 (from AMS) in the majority of experiments and are slightly higher than those of the WB-wood burning experiments. The average oxidation slopes that are estimated from MIR spectroscopy are also higher for the CC coal combustion experiments compared to WB-wood burning. This observation is supported by the formation of SOAs with higher naCO abundances for CC (?)non-acid carbonyl abundances for coal combustion (Yazdani et al., 2021b), resulting in higher trajectory slopes. However, MIR generally estimates higher O:C (by 0.05–0.1) and lower H:C ratios (by approximately 0.2) for the aged CC coal combustion aerosols compared to AMS.

The observed deviations between the discrepancies between the measurements of the two instruments might stem from the low OM mass collected on the filtersthat, which increases the baseline correction and peak fitting uncertainties in MIR analysis, in addition to the existence of FGs. The existence of functional groups that are not considered in the peak fitting algorithm (e.g., ethers, PAHs, rC=C and rCH), and the assumption about the fractional carbon associated with each functional group might also play a role. Sampling biases of semi-volatile compounds on PTFE filters (Subramanian et al., 2004) and the uncertainties of AMS elemental ratio calibrations (Canagaratna et al., 2015b; Aiken et al., 2008) can also affect the results.

## 3.2 AMS fragment ion-MIR FG functional group relationships

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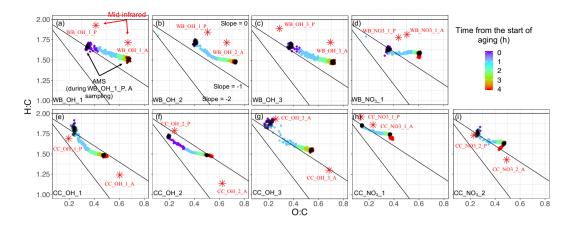
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In Sects. 3.2.1 and 3.2.2, the connection between the AMS fragment ions and MIR FGs functional groups is investigated using different statistical methods (covariance, correlation, and VIP scores). The combined summary of these analyses is shown in Table 1.



**Figure 4.** Comparison of AMS van Krevelen (H:C vs O:C) aging trajectories (color circles) for wood burning (WB) and coal combustion (CC) experiments with MIR estimates for aerosols collected on PTFE filters before and after aging (red stars). Black circles in AMS trajectories correspond to the filter sampling periods for the primary and aged aerosols. The filter names are the same as **?**Yazdani et al. (2021b). P: primary, A: aged. There are one to three similar experiments in each category (WB\_OH, WB\_NO<sub>3</sub>, CC\_OH, and CC\_NO<sub>3</sub>).

#### 3.2.1 Correlation and covariance analyses (univariate)

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The aCH group has high covariance with  $C_xH_{2x-1}$  and  $C_xH_{2x+1}$  fragments ( $C_3H_5^+$ ,  $C_3H_7^+$ ,  $C_4H_9^+$ ,  $C_4H_9^+$ , and  $C_5H_9^+$ ; Fig. 5). The highest correlations are between the aCH group and  $C_3H_5^+$ ,  $C_3H_7^+$ ,  $C_5H_7^+$ , and  $C_6H_9^+$  (Fig. 6). The relationship of heavier-larger fragments such as  $C_7H_{13}^+$  and  $C_8H_{15}^+$  with aCH is more prominent in the correlation analysis. These fragments are especially abundant in the primary CC coal combustion aerosols, suggesting these aerosols is composed of relatively are composed of longer chain hydrocarbons relative to wood burning aerosols that even after fragmentation, produce relatively heavy-large fragments. This observation is also supported by the MIR spectra of these samples, which possess sharp  $CH_2$  and weak  $CH_3$  peaks (??)(Yazdani et al., 2021a, b). Faber et al. (2017) have previously shown the relation between  $C_4H_7^+$  and aCH. The m/z 57 signal in the unit-mass-resolution mass spectra (which also includes  $C_4H_9^+$  signalincludes  $C_4H_9^+$ ) has been proposed to be a tracer of unburned fuel emissions (Schneider et al., 2006). However, its correlation coefficient with aCH has been shown to be quite variable and sometimes negative (Russell et al., 2009a). This discrepancy partly stems from the contribution of  $C_3H_5O^+$  to m/z 57 (Faber et al., 2017) and partly from the fact that molecules with different chain-lengths, degrees of branching, and heteroatoms produce different and source-dependent fragmentation patterns for  $C_xH_y^+$ . In addition, the existence of several highly correlated ion fragments with aCH in this study suggests the superiority of a multi-variate approach to obtain information about aCH from the AMS mass spectraths aCH group.

The COOH group has the highest covariance with  $CO_2^+$ ,  $CO^+$ ,  $C_2H_3O^+$ , and  $CHO^+$  . Since the  $CO^+$  concentration ( $CO^+$  is estimated from that of  $CO_2^+$ , the former fragment it not investigated separately  $CO_2^+$ ). The highest correlations are , on the other hand, with  $C_2H_3O_2^+$  and several heavier larger fragments with multiple oxygen atoms such as  $C_7H_5O_4^+$ , which are abundant in the aged  $\overline{WB}$ -wood burning aerosols. The high covariance with the  $CO_2^+$  fragment is supported by the fact

**Table 1.** Summary of important fragment ions for each functional group based on the analysis method. Important fragments are shown in blue.

FGs	Covariance	Correlation	VIP scores	Multivariate regression CO <sub>2</sub> +, CHO+, C <sub>2</sub> H <sub>3</sub> O+
Alkane (aCH)	$C_x H_{2x\pm 1}^+$ (e.g., $C_3 H_7^+$ , $C_4 H_9^+$ )	$C_xH_Y^+$ (e.g., $C_3H_5^+$ , $C_3H_7^+$ , $C_6H_9^+$ , $C_8H_9^+$ , heavy large fragments: $C_8H_{17}^+$ )	$C_3H_5^+$	-
Alcohol (aCOH)	$CHO^{+}, CH_{3}O^{+},$ $C_{2}H_{4}O_{2}^{+}$	$C_2H_5O^+, C_3H_6O_2^+, C_5H_6^+$ (phenol)	$CHO^{+}, C_{2}H_{4}O_{2}^{+},$ $C_{8}H_{9}O_{2}^{+}$	$\text{CHO}^+, \text{C}_2\text{H}_3\text{O}^+$
Acid (COOH)	$CO_2^+, CO^+, C_2H_3O^+,$ $CHO^+$	$C_x H_y O_{>1}^+$ (e.g., $C_2 H_3 O_2^+$ , $C_7 H_5 O_4^+$ )	CO <sub>2</sub> <sup>+</sup>	CO <sub>2</sub> <sup>+</sup> , CHO <sup>+</sup>
Non-acid carbonyl (naCO)	$CO_2^+$ , $CO^+$ , $C_2H_3O^+$	$C_x H_y O_1^+ (e.g., C_7 H_4 O^+)$	$C_2H_3O^+$	$C_2H_3O^+$
Organonitra (RONO <sub>2</sub> )	ate –	$C_2H_3O^+, C_4H_9NO_3^+,  C_6H_{10}NO_2^+$	$CHO^{+}, C_{2}H_{3}O^{+}$	-

that  $CO_2^+$  is produced from the fragmentation of mono- and dicarboxylic acids (Duplissy et al., 2011; Zhang et al., 2005). The heavier larger fragments with multiple oxygen atoms ( $C_xH_yO_{z>1}^+$ ) are also indicative of the COOH FG functional group as also reported by Lambe et al. (2012) and might be source-specific<del>as, for.</del> For example,  $C_7H_5O_4^+$  is only detected in the WB aerosols. In the WB wood burning aerosols and can be a potential wood burning SOA marker. Sun et al. (2010) reported observed the  $C_7H_5O_4^+$  fragment in the AMS spectra of syringol SOA. In the wood burning aerosols, the COOH group is correlated significantly ( $r \sim 0.96$ ) with CHO<sub>2</sub>+, the fragment known to be produced from the  $\alpha$ -cleavage of carboxylic acids (Pavia et al., 2008). The  $C_2H_4O_2^+$  fragment is also known to be produced from acids having  $\gamma$  hydrogen through McLafferty rearrangement (Pavia et al., 2008) and its concentration increases with the extensive aging for the WB and CC aerosols (?) wood burning and coal combustion aerosols (Yazdani et al., 2021b). However, the strong interference of levoglucosan fragmentation, abundant in WB emissions, with  $C_2H_4O_2^+$  makes significant contribution of levoglucosan to  $C_2H_4O_2^+$  in wood burning aerosols makes the investigation of COOH- $C_2H_4O_2^+$  relation difficult. With the help of MIR, which does not suffer from the same interference, samples with negligible levoglucosan concentrations were separated. For these samples, a fairly strong correlation ( $r \sim 0.82$ ) between COOH and  $C_2H_4O_2^+$  was observed.

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The aCOH group covaries most significantly the most with CHO<sup>+</sup>, CH<sub>3</sub>O<sup>+</sup>, C<sub>3</sub>H<sub>5</sub>O<sup>+</sup> (which contributes to m/z 57 in the unit-mass-resolution spectra), C<sub>2</sub>H<sub>3</sub>O<sup>+</sup>, and C<sub>2</sub>H<sub>4</sub>O<sub>2</sub><sup>+</sup>. The CHO<sup>+</sup> fragment has been often interpreted as the tracer of esters, polyols and compounds with polyfunctional groups without the carboxylic COOH (Canagaratna et al., 2015a). This fragment  $\frac{1}{2}$  however, is also known to be produced by aldehydes although but the aldehyde C-H band is not observed in the MIR spectra of the samples under study. Faber et al. (2017) also showed that the signal ratio of C<sub>2</sub>H<sub>3</sub>O<sup>+</sup> to C<sub>4</sub>H<sub>7</sub><sup>+</sup> is linearly correlated with the molar ratio of aCOH to aCH. The connection of C<sub>2</sub>H<sub>3</sub>O<sup>+</sup> with alcohols, however, should be treated with caution as carbonyls can also produce the same mass fragment. The C<sub>2</sub>H<sub>4</sub>O<sub>2</sub><sup>+</sup> fragment appears to be important for aCOH and has been previously shown to be related to anhydrous sugars in the biomass burning smoke (Schneider et al., 2006). The CH<sub>3</sub>O<sup>+</sup> fragment is produced from the  $\alpha$  cleavage of alcohols (Pavia et al., 2008). The highest correlations in this work are  $\frac{1}{2}$  however, between the aCOH group and C<sub>2</sub>H<sub>5</sub>O<sup>+</sup> and C<sub>3</sub>H<sub>6</sub>O<sub>2</sub><sup>+</sup> fragments and some other fragments such as C<sub>5</sub>H<sub>6</sub><sup>+</sup>. The C<sub>2</sub>H<sub>5</sub>O<sup>+</sup> fragment is also known to be produced from the  $\alpha$  cleavage of alcohols (Pavia et al., 2008). The C<sub>5</sub>H<sub>6</sub><sup>+</sup> fragment, correlated to a lower extent with aCOH, can be produced by phenol fragmentation after loosing CO (Pavia et al., 2008), which is also abundant in WB wood burning emissions (Bruns et al., 2017).

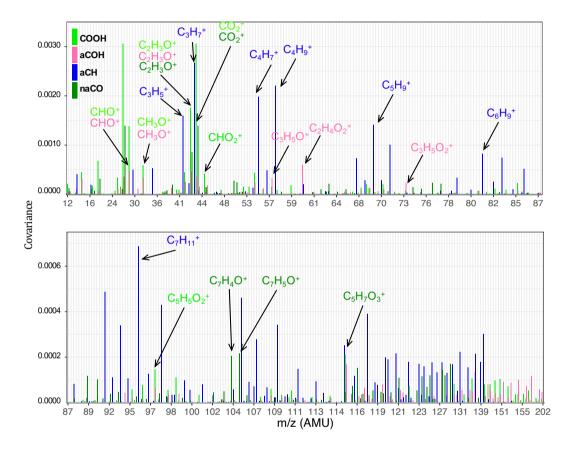
The naCO group covaries most significantly non-acid carbonyl group (naCO) covaries the most with CO<sup>+</sup>, C<sub>2</sub>H<sub>3</sub>O<sup>+</sup>, and CO<sub>2</sub><sup>+</sup>. Contrary to COOH, CHO<sup>+</sup> appears to have a low covariance with naCO. The C<sub>2</sub>H<sub>3</sub>O<sup>+</sup> fragment is known to be produced by aliphatic ketones and aldehydes (Pavia et al., 2008; Eadon et al., 1971). As discussed by **?**Yazdani et al. (2021b), the naCO in the CC coal combustion samples are believed to be mostly ketone based on their C=O frequency. The naCO group is highly correlated with some  $C_xH_y^+$  fragments (e.g.,  $C_4H_3^+$ ,  $C_6H_4^+$ ) and some single-oxygen fragments (e.g.,  $C_5HO^+$ ,  $C_7H_4O^+$ , and  $C_7H_5O^+$ ). The  $C_7H_5O^+$  fragment is known to be produced by aromatic ketones (Pavia et al., 2008) and  $C_xH_yO_1^+$  has been attributed to carbonyls (Lambe et al., 2012). The  $C_2H_3O^+$ :CO<sub>2</sub><sup>+</sup> ratio is relatively higher in samples aged with the nitrate radical or samples that have considerable amounts of the naCO group (Fig. S5), suggesting that  $C_2H_3O^+$  is produced mainly by molecules possessing naCO or SOAs-SOA species formed with the nitrate radical.

There are mid-infrared signatures attributed to levoglucosan and lignin-lignin-like compounds that are prominent in the primary WB-wood burning aerosols and diminish with aging. These signatures are important markers of biomass burning and have been used to identify atmospheric smoke-impacted samples (?). Correlation analysis of these features. The correlation analysis for these signatures with the AMS ion fragments (Fig. S6) show suggest that the  $C_8H_9O^+$  fragment is related to ligninlignin-like compounds. In fact, one might attribute the m/z 121 fragment to two peaks  $C_7H_5O_2^+$  and  $C_8H_9O^+$  for hydroxyphenyl (H) lignin and  $C_4H_9O_2^+$  of guaiacyl (G) lignin, respectively (Li et al., 2012; Tolbert and Ragauskas, 2017). The  $C_2H_4O_2^+$  and  $C_3H_5O_2^+$  fragments have high correlation correlations with MIR levoglucosan signatures. The fragment at m/z 102, in spite of having generally a low concentration, (Fig. S6) has the highest correlation with the levoglucosan concentration measured from the MIR spectra. This fragment might be used alternatively in case the interference of other compounds (e.g., acids) for the lighter smaller fragments related to levoglucosan is substantial.

In this work, organonitrates were not quantified. However, the The MIR peak attributed to RONO<sub>2</sub> has high correlation coefficients with  $C_2H_3O^+$  and several other oxygenated fragments such as  $C_4H_7O^+$ . Nitrogenated fragments containing the nitrate and nitro groups such as  $C_4H_9NO_3^+$  and  $C_6H_{10}NO_2^+$  also appear to have moderate correlation coefficients (approximately 0.6)

with the RONO<sub>2</sub> peak in the MIR spectrum (Fig. S7)<del>although the</del>. The quantification of nitrogenated fragments is, however, known to be complicated in the V mode.

To summarize, the high correlation coefficients of several fragments with each FG suggest that FG functional group suggest that functional group information is retained to a good degree in the AMS spectra. We also found that multiple FGs functional groups are correlated with each of the major oxygenated fragments, (CO<sup>+</sup>, C<sub>2</sub>H<sub>3</sub>O<sup>+</sup>, CO<sub>2</sub><sup>+</sup>, and CHO<sup>+</sup>). As a result, a multivariate approach should be taken to infer FG functional group abundances from the AMS spectra.

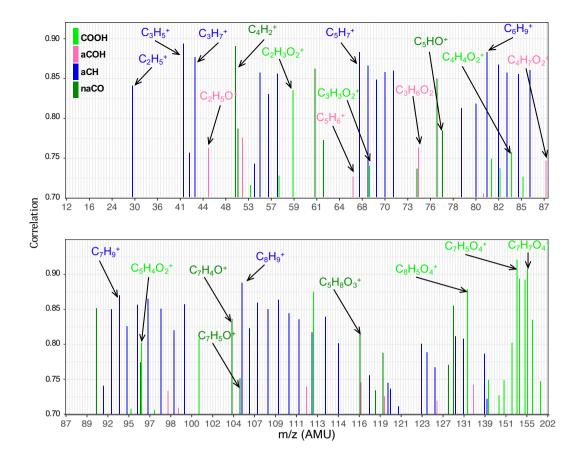


**Figure 5.** Bar plots showing positive covariances of normalized AMS fragment ion concentrations and normalized FG functional group abundances. Only positive covariances are shown.

#### 3.2.2 VIP scores (multivariate)

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As can be seen from Fig. 7, the  $\mathrm{CO_2}^+$  fragment has the highest VIP scores for the carbonyl  $\nu(\mathrm{CO})$  and broad acid  $\nu(\mathrm{OH})$  peaks from 2400 to 3400 cm<sup>-1</sup>. This is consistent with previous studies (e.g., Zhang et al., 2005) and our univariate analyses (Sect. 3.2.1). On the other hand, the  $\nu(\mathrm{CH})$  region (2800 to  $\frac{3200}{3000}\,\mathrm{cm}^{-1}$ ), interfering with the broad acid OH stretching band, has high VIP scores with negative regression coefficient, showing that aCH relative concentration is anti-correlated with  $\mathrm{CO_2}^+$ .



**Figure 6.** Bar plots showing the Pearson correlation coefficients of normalized individual fragment ion concentrations and normalized FG functional group abundances. Only values > 0.7 are shown.

Although some interference for the  $CO_2^+$  fragment is expected from gas-phase  $CO_2$  in the AMS spectra, our results show that the compensation method appears to eliminate this interference this interference is eliminated effectively.

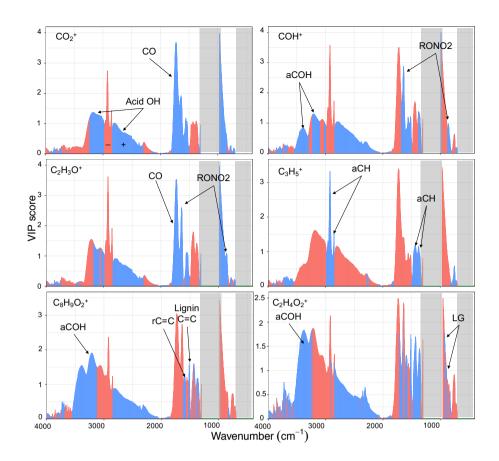
The COH<sup>+</sup> has the highest VIP scores for the RONO<sub>2</sub> peaks and the broad alcohol  $\nu(OH)$  at 3400 cm<sup>-1</sup>. These results suggest that alcohols and the SOA species produced during the aging with the nitrate radical (that can also be alcohols) are mostly responsible for producing this mass fragment. Although the interference form the gas-phase <sup>15</sup>N <sup>14</sup>N can be significant for CHO<sup>+</sup>, our results show that CHO<sup>+</sup> appears to be meaningfully indicative of alcohols after compensation for the the subtraction of the gas-phase interference.

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The  $C_2H_3O^+$  fragment has the highest VIP scores for the carbonyl CO (likely from non-acid carbonyl) and the RONO<sub>2</sub> peaks and also to a lesser extent for the broad acid  $\nu(OH)$  peak. This observation suggests that  $C_2H_3O^+$  is mostly mainly produced by fragmentation of carbonyls and SOA species formed by aging with the nitrate radical and to a lesser extent carboxylic acids.

The  $C_3H_5^+$  fragment was chosen for the VIP scores analysis due to having high concentrations for both  $\frac{CC}{C}$  and  $\frac{C$ 



**Figure 7.** VIP scores of MIR absorbances regressed against AMS fragment ion concentrations (averaged over the filter sampling periods). Blue and red scores correspond to wavenumbers with positive and negative regression coefficients in the <u>PLS PLSR</u> models, respectively. Important <u>FGs</u>-functional groups for each mass fragment are indicated and PTFE absorption regions are masked by gray rectangles.

 $\nu$ (CH) (2800–3000 cm<sup>-1</sup>) and  $\delta$ (CH) (1300–1500 cm<sup>-1</sup>) peaks, showing that this fragment is directly related to aCH for both sources. This result has been expected but also highlights the fact that  $C_xH_y^+$  fragments should be chosen wisely based on the aerosol source to provide useful information about the aCH group. For example, in this study, the  $C_3H_5^+$  fragment appears to be more correlated with the aCH group than the commonly used fragments,  $C_4H_9^+$ .

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Regarding the biomass burning markers, the  $C_8H_9O_2^+$  fragmentthat, which is proposed to be produced by the fragmentation of lignin molecules (Li et al., 2012; Tolbert and Ragauskas, 2017), has the highest VIP scores with positive coefficients in the aromatic  $\nu(C=C)$  (1515 and 1600 cm<sup>-1</sup>) and aCOH regions, suggesting. This observation suggests the connection of this fragment with aromatic compounds having a that have lignin-like substitution substitutions that generates the sharp peak at 1515 cm<sup>-1</sup> and aCOH groups. The  $C_2H_4O_2^+$  fragmentthat, which is proposed to be produced predominantly by fragmentation of levelucosan molecules has the highest VIP scores with positive coefficients in the aCOH region (3400 cm<sup>-1</sup>), suggesting the abundance of aCOH in molecules producing this fragment. In addition, high VIP scores with

positive regression coefficients is observed in the 850–1000 cm<sup>-1</sup> region, which. This region was previously proposed to be related to contain levoglucosan fingerprint absorbances (?)(Yazdani et al., 2021b).

We also performed a simple multivariate linear regression between the oxygenated FGs functional groups (aCOH, naCO, and COOH) and major fragments (CO<sub>2</sub><sup>+</sup>, CHO<sup>+</sup> and C<sub>2</sub>H<sub>3</sub>O<sup>+</sup>; Fig S8) for the experiments in which the hydroxyl radical was used exident. As shown in Fig. S8, regressing CO<sub>2</sub><sup>+</sup> against COOH, aCOH, and naCO results in the highest regression coefficient for the COOH group. In a similar regression for CHO<sup>+</sup>, the relative contribution of aCOH increases (Fig. S8). However, a high regression coefficient for COOH is still observed. The regression for C<sub>2</sub>H<sub>3</sub>O<sup>+</sup> highlights a relatively greater contribution of naCO (Fig. S8). However, as for CHO<sup>+</sup>, a high regression coefficient for COOH is also observed. As summarized in Table 1, different statistical methods suggest that the major fragments are usually produced more by a certain oxygenated FG functional group, while interference from other FGs functional groups might also be significant. This motivates the use of multivariate methods for predicting FG functional group abundances using fragment ion concentrations in the following section.

#### 3.3 MIR FG functional group interpolation using AMS mass spectra

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We showed in previous sections that AMS and MIR measurements are consistent. We also found that FG-functional group information is maintained in the AMS mass spectra, which motivated the use of multivariate methods to access this information. For this purpose, normalized AMS spectra were regressed against normalized FG-functional group compositions from MIR peak fitting. The fit quality of the developed models is reasonable (Figs. S9 and S10) with their  $R^2$  ranging from 0.71 to 0.94. These models use mass fragments to predict the FG-functional group compositions that were found to be important in previous sections. We used the developed PLSR models to interpolate the functional group composition of WB and CC wood burning and coal combustion OM between the filter sampling periods (primary and aged) using the AMS spectra (Fig. 8). These models are especially helpful as AMS has a considerably better time resolution and can be used to investigate the FG detailed functional group evolution of OM during the course of aging.

The interpolated functional group compositions (Figs. 8, S11, and S12) show different FG functional group compositions and trends during the course of oxidation for WB and CC wood burning and coal combustion aerosols. This is predominantly seen in the fraction of oxygenated functional groups that emerge with aging. For the WB-wood burning experiments, the aCH relative abundance falls steeply as aging with the hydroxyl radical starts (Fig. 8a). This is also observed for aCOH. On the other hand the COOH relative abundance increases significantly as soon as the aging starts and levels off after two hours of aging.

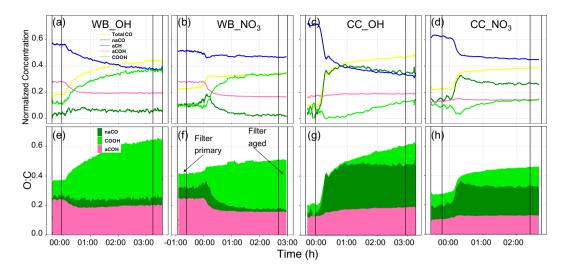
The relative abundance of naCO, however, does not change significantly compared to other FGs for the WB functional groups for the wood burning experiment (Fig. 8a). There are slight differences between different experiments of WB-wood burning aging with the hydroxyl radical (also observed in their van Krevelen trajectory in Fig. 4a-c). For instance, the relatively linear trajectory of the WB\_OH\_1 experiment (Fig. 8a) is concurrent with monotonic FG-functional group composition change (Fig. S11a). On the other hand, the curved van Krevelen trajectories of WB\_OH\_2 and WB\_OH\_3 (Fig. S12b-c) correspond to the consumption of naCO produced after the start of aging and the gradual increase of the COOH relative abundance (Fig. S12b-c). The different SOA species formed by oxidation with hydroxyl and nitrate radicals is also reflected in the evolution of OM FG-functional group composition. When aging with the nitrate radical, the decrease in the relative abundance of aCH

is much less prominent after the start of aging compared to aging with the hydroxyl radical (Fig. 8b) although the OM mass enhancement is comparable between the two (Fig. S11a-d). This observation suggests the formation of different SOA species with higher relative abundance of aCH when the nitrate radical is used. This is also supported by the horizontal trajectory in the Van krevelen diagram (no decrease in H:C) (Fig. 4d). No clear difference in the composition of oxygenated FGs functional groups (except organonitrate) is observed between aging with the hydroxyl and nitrate radicals. However, when the nitrate radical is used, the O:C ratio increases to lower levels and reaches a plateau faster (Fig. 8e–f; also true for the CC coal combustion OM). This observation is consistent with the fact that the nitrate radical is produced with a single injection of ozone but the hydroxyl radical is generated continuously throughout the aging. It is observed that most prominent changes in the FG functional group composition for both oxidants occur in the first hour of aging when the OM mass changes the most (Fig S11), while and only small changes are observed toward the end of aging. Looking at the absolute abundances of functional groups, we observe that the mass concentrations of all FGs functional groups including aCH and aCOH increase during the course of aging (Fig S11) and it is the different rate of their rates increase that changes their relative abundance as shown in Fig. 8.

We observe for the CC For the coal combustion experiments a different composition of FGs emerging functional groups emerges after the start of aging that also evolves differently as aging continues. Like for the WB-wood burning experiments, the aCH relative abundance decreases drastically with aging (Fig. 8c), while its absolute concentration increases only slightly with aging (Fig. S11). The decrease in the aCH relative abundance is, however, less prominent when the nitrate radical is used (Fig. 8d) as also supported by the lower decrease in the H:C shown the the Van krevelen van Krevelen plots (Fig. 4h-i). Unlike the WB-wood burning experiments, the relative abundance of the aCOH group increases slightly with aging when using both exidents (Fig. 8c-d). The relative abundances of naCO and COOH show more complex behaviors. The relative abundance of naCO increases sharply and naCO becomes the major FG-functional group with the start of aging for both exident but decreases slightly after continued aging (Fig. 8c-d). The relative abundance of COOH decreases initially (Fig. 8c), however, after about one hour into the aging process (earlier with the nitrate radical), when there is no more significant OM enhancement, the COOH relative abundance starts to increase gradually. This observation is consistent with the ripening phenomenon (Wang et al., 2018) in which the composition of the SOA keeps changing and becomes more exidized, while the change in the OM mass is minimal. This phenomenon is also observed in the L-shaped exidation trajectories of CC coal combustion OM in the van Krevelen plot of Figs. 4e-i for both exidants.

To summarize, the interpolated FG functional group compositions are supported by the van Krevelen trajectories, but provide insight insights into the oxidation pathways that cannot be independently obtained from the van Krevelen plots (e.g., several combinations of FGs functional groups can produce similar slopes). For two CC coal combustion experiments, negative concentrations of COOH are predicted (Fig S12e–f). These unphysical values are believed to resulted due to uncertainties of PLSR models and quantification uncertainties for the COOH group from the MIR peak fitting. The predicted trends, however, are still informative.

#### 4 Concluding remarks



**Figure 8.** Time series of normalized concentration of functional groups interpolated using AMS mass spectra (a–d) and time series of O:C ratios calculated from the interpolated functional groups (e–h). An example for each source (CC and WB) and oxidant (OH and NO<sub>3</sub>) has been shown. The time zero indicates the start of aging (UV lights turned on or oxidant injected). The horizontal lines indicate the periods of filter sampling.

#### The four MIR FGs highlighted by the VIP scores method (i.e.,

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We combined statistically collocated AMS and MIR measurements in an environmental simulation chamber and found that AMS OM is associated the most with aCH, aCOH, COOH, and non-acid carbonyl (naCO) explain the functional groups for aerosols generated from the combustion of wood and coal. The OM mass, OM:OC, H:C, and O:C estimated from the abundances of the mentioned functional groups were in good agreement with those of AMS for the WB and CC aerosols. By using univariate and multivariate methods, we found that several light and heavy (generally source-specific from AMS measurements and showed clear variations across fuel types and oxidants. These functional groups are those that were used in previous studies of atmospheric aerosols using MIR (e.g., Maria et al., 2003; Russell et al., 2009c; Reggente et al., 2019a).

Previous studies of functional group-ion fragment relationships were limited to small fragments and did not consider marker signatures in the mid-infrared spectra (Russell et al., 2009a; Faber et al., 2017). We performed a univariate (correlation and covariance) analysis on the four mentioned functional groups and more than 300 fragment ions up to m/z 212 in the AMS spectra and found several small (low m/z) and large (m/z > 100) AMS fragment ions are to be informative about the FG composition of OM. Our analysis indicates that the fragmentation of OM functional group composition of POA and SOA from the combustion sources. For example,  $C_7H_6O_4^+$ , which was only detected in wood burning SOA and was highly correlated with the COOH functional group, might be used as a potential marker of biomass burning SOA in the atmosphere. The peaks in the FTIR spectra that were believed to be associated with biomass burning markers (levoglucosan and lignin-like compounds) were also found to be highly correlated with the fragments related these markers.

Our multivariate (VIP scores) analysis indicated that when OMs with different proportions of oxygenated FGs (COOH, functional groups (i.e., COOH, aCOH, and naCO) in AMS produces are fragmented in the AMS, they produce different

proportions of the major oxygenated fragments (i.e.,  $CO_2^+$ ,  $CHO^+$ ,  $C_2H_3O^+$ ). A multivariate method was used to For examples,

C<sub>2</sub>H<sub>3</sub>O<sup>+</sup> was found to be associated the most with non-acid carbonyl and SOA species formed with nitrate radical. However, each of these small oxygenated fragments does not represent only a single functional group and contribution from other

functional groups is expected.

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Finally, we developed a method to extract the high-time-resolution FG-functional group information from the AMS spectra

to better understand the evolution of OM composition with the OM composition during the course of aging. The results of this

method provide insights into the oxidation pathways that cannot be independently obtained from the van Krevelen plot. The

results of this method, which can be easily implemented in other environmental chamber studies, suggest the formation of mod-

erately oxygenated FGs functional groups (e.g., naCOnon-acid carbonyl) soon after the start of aging and their transformation

to the more oxygenated FG, COOH, after the transformation of moderately oxygenated functional groups to more oxygenated

functional groups (e.g., acid) with continued aging.

Code availability. TEXT

Data availability. TEXT

Code and data availability. TEXT

Sample availability. TEXT

Author contributions. IEH and ST and AY conceived of the project and manuscript. AB and IEH performed the chamber experiments. AB provided AMS spectra. ND prepared and assembled the filter sampling set-up and took their FT-IR spectra. AMD provided atomized compounds and ambient sample spectra. AY wrote the code for data analysis and post processing, performed the data analysis, and wrote the manuscript. ST edited the manuscript and provided regular feedback on the analysis. IEH, ASHP, AB, AMD and ND provided input on the analysis and further editing of the manuscript. ST and IEH provided overall supervision of the project.

Competing interests. We declare that no competing interests are present

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#### References

565

- Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper,
  D., Sun, Y., Zhang, Q., Trimborn, A., Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot, A. S. H., Dommen, J., Duplissy, J., Metzger, A., Baltensperger, U., and Jimenez, J. L.: O/C and OM/OC Ratios of Primary, Secondary, and Ambient Organic Aerosols with High-Resolution Time-of-Flight Aerosol Mass Spectrometry, Environ. Sci. Technol., 42, 4478–4485, https://doi.org/10.1021/es703009q, 2008.
- Barmet, P., Dommen, J., DeCarlo, P. F., Tritscher, T., Praplan, A. P., Platt, S. M., Prévôt, A. S. H., Donahue, N. M., and Baltensperger, U.: OH
  Clock Determination by Proton Transfer Reaction Mass Spectrometry at an Environmental Chamber, Atmos. Meas. Tech., 5, 647–656, https://doi.org/10.5194/amt-5-647-2012, 2012.
  - Bertrand, A., Stefenelli, G., Bruns, E. A., Pieber, S. M., Temime-Roussel, B., Slowik, J. G., Prévôt, A. S. H., Wortham, H., El Haddad, I., and Marchand, N.: Primary Emissions and Secondary Aerosol Production Potential from Woodstoves for Residential Heating: Influence of the Stove Technology and Combustion Efficiency, Atmos. Environ., 169, 65–79, https://doi.org/10.1016/j.atmosenv.2017.09.005, 2017.
- Bertrand, A., Stefenelli, G., Jen, C. N., Pieber, S. M., Bruns, E. A., Ni, H., Temime-Roussel, B., Slowik, J. G., Goldstein, A. H., Haddad, I. E., Baltensperger, U., Prévôt, A. S. H., Wortham, H., and Marchand, N.: Evolution of the Chemical Fingerprint of Biomass Burning Organic Aerosol during Aging, Atmos. Chem. Phys., 18, 7607–7624, https://doi.org/10.5194/acp-18-7607-2018, 2018.
  - Boris, A. J., Takahama, S., Weakley, A. T., Debus, B. M., Fredrickson, C. D., Esparza-Sanchez, M., Burki, C., Reggente, M., Shaw, S. L., Edgerton, E. S., and Dillner, A. M.: Quantifying Organic Matter and Functional Groups in Particulate Matter Filter Samples from the Southeastern United States Part 1: Methods, Atmos. Meas. Tech., 12, 5391–5415, https://doi.org/10.5194/amt-12-5391-2019, 2019.
  - Bruns, E., Krapf, M., Orasche, J., Huang, Y., Zimmermann, R., Drinovec, L., Močnik, G., El-Haddad, I., G. Slowik, J., Dommen, J., Baltensperger, U., and Prevot, A.: Characterization of Primary and Secondary Wood Combustion Products Generated under Different Burner Loads, Atmos. Chem. Phys., 15, 2825–2841, https://doi.org/10.5194/acp-15-2825-2015, 2015.
- Bruns, E. A., Slowik, J. G., Haddad, I. E., Kilic, D., Klein, F., Dommen, J., Temime-Roussel, B., Marchand, N., Baltensperger, U., and
  Prévôt, A. S. H.: Characterization of Gas-Phase Organics Using Proton Transfer Reaction Time-of-Flight Mass Spectrometry: Fresh and
  Aged Residential Wood Combustion Emissions, Atmos. Chem. Phys., 17, 705–720, https://doi.org/10.5194/acp-17-705-2017, 2017.
  - Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Chemical and Microphysical Characterization of Ambient Aerosols with the Aerodyne Aerosol Mass Spectrometer, Mass Spectrom. Rev., 26, 185–222, https://doi.org/10.1002/mas.20115, 2007.
  - Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne, J. T., and Worsnop, D. R.: Elemental Ratio Measurements of Organic Compounds Using Aerosol Mass Spectrometry: Characterization, Improved Calibration, and Implications, Atmos. Chem. Phys., 15, 253–272, https://doi.org/10.5194/acp-15-253-2015, 2015a.
- Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne, J. T., and Worsnop, D. R.: Elemental Ratio Measurements of Organic Compounds Using Aerosol Mass Spectrometry: Characterization, Improved Calibration, and Implications, Atmos. Chem. Phys., 15, 253–272, https://doi.org/10.5194/acp-15-253-2015, 2015b.

- Chhabra, P. S., Ng, N. L., Canagaratna, M. R., Corrigan, A. L., Russell, L. M., Worsnop, D. R., Flagan, R. C., and Seinfeld, J. H.: Elemental

  Composition and Oxidation of Chamber Organic Aerosol, Atmos. Chem. Phys., 11, 8827–8845, https://doi.org/10.5194/acp-11-8827-2011, 2011a.
  - Chhabra, P. S., Ng, N. L., Canagaratna, M. R., Corrigan, A. L., Russell, L. M., Worsnop, D. R., Flagan, R. C., and Seinfeld, J. H.: Supplemental Material for Elemental Composition and Oxidation of Chamber Organic Aerosol, Atmos. Chem. Phys., p. 16, https://doi.org/10.5194/acp-11-8827-2011, 2011b.
- 590 Chong, I.-G. and Jun, C.-H.: Performance of Some Variable Selection Methods When Multicollinearity Is Present, Chemom. Intell. Lab. Syst., 78, 103–112, https://doi.org/10.1016/j.chemolab.2004.12.011, 2005.

- Corrigan, A. L., Russell, L. M., Takahama, S., Äijälä, M., Ehn, M., Junninen, H., Rinne, J., Petäjä, T., Kulmala, M., Vogel, A. L., Hoffmann, T., Ebben, C. J., Geiger, F. M., Chhabra, P., Seinfeld, J. H., Worsnop, D. R., Song, W., Auld, J., and Williams, J.: Biogenic and Biomass Burning Organic Aerosol in a Boreal Forest at Hyytiälä, Finland, during HUMPPA-COPEC 2010, Atmos. Chem. Phys., 13, 12 233–12 256, https://doi.org/10.5194/acp-13-12233-2013, 2013.
- DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.: Field-Deployable, High-Resolution, Time-of-Flight Aerosol Mass Spectrometer, Anal. Chem., 78, 8281–8289, https://doi.org/10.1021/ac061249n, 2006.
- Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R., Metzger, A., Barmpadimos, I., Prevot, A. S. H., Weingartner, E., Tritscher, T.,
   Gysel, M., Aiken, A. C., Jimenez, J. L., Canagaratna, M. R., Worsnop, D. R., Collins, D. R., Tomlinson, J., and Baltensperger, U.: Relating Hygroscopicity and Composition of Organic Aerosol Particulate Matter, Atmos. Chem. Phys., 11, 1155–1165, https://doi.org/10.5194/acp-11-1155-2011, 2011.
- Eadon, G., Djerassi, C., Beynon, J. H., and Caprioli, R. M.: The Fragmentation of Aliphatic Ketones in the Mass Spectrometer: A Detailed Study of Nonan-4-One Using Ion Kinetic Energy Spectroscopy, Int. J. Mass Spectrom., 5, 917–933, https://doi.org/10.1002/oms.1210050803, 1971.
  - Faber, P., Drewnick, F., Bierl, R., and Borrmann, S.: Complementary Online Aerosol Mass Spectrometry and Offline FT-IR Spectroscopy Measurements: Prospects and Challenges for the Analysis of Anthropogenic Aerosol Particle Emissions, Atmos. Environ., 166, 92–98, https://doi.org/10.1016/j.atmosenv.2017.07.014, 2017.
- Frossard, A. A., Shaw, P. M., Russell, L. M., Kroll, J. H., Canagaratna, M. R., Worsnop, D. R., Quinn, P. K., and Bates, T. S.: Springtime

  Arctic Haze Contributions of Submicron Organic Particles from European and Asian Combustion Sources, J. Geophys. Res.-Atmos., 116,
  D05 205, https://doi.org/10.1029/2010JD015178, 2011.
  - Frossard, A. A., Russell, L. M., Massoli, P., Bates, T. S., and Quinn, P. K.: Side-by-Side Comparison of Four Techniques Explains the Apparent Differences in the Organic Composition of Generated and Ambient Marine Aerosol Particles, Aerosol Sci. Tech., 48, v–x, https://doi.org/10.1080/02786826.2013.879979, 2014.
- 615 Gilardoni, S., Liu, S., Takahama, S., Russell, L. M., Allan, J. D., Steinbrecher, R., Jimenez, J. L., De Carlo, P. F., Dunlea, E. J., and Baumgardner, D.: Characterization of Organic Ambient Aerosol during MIRAGE 2006 on Three Platforms, Atmos. Chem. Phys., 9, 5417–5432, https://doi.org/10.5194/acp-9-5417-2009, 2009.
- Grieshop, A. P., Donahue, N. M., and Robinson, A. L.: Laboratory Investigation of Photochemical Oxidation of Organic Aerosol from Wood Fires 2: Analysis of Aerosol Mass Spectrometer Data, Atmos. Chem. Phys., 9, 2227–2240, https://doi.org/10.5194/acp-9-2227-2009, 2009.

- Haaland, D. M. and Thomas, E. V.: Partial Least-Squares Methods for Spectral Analyses. 1. Relation to Other Quantitative Calibration Methods and the Extraction of Qualitative Information, Analytical Chemistry, 60, 1193–1202, https://doi.org/10.1021/ac00162a020, 1988.
- Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maen-
- haut, W., McFiggans, G., Mentel, T. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The Formation, Properties and Impact of Secondary Organic Aerosol: Current and Emerging Issues, Atmos. Chem. Phys., 9, 5155–5236, https://doi.org/10.5194/acp-9-5155-2009, 2009.
  - Hastings, S. H., Watson, A. T., Williams, R. B., and Anderson, J. A.: Determination of Hydrocarbon Functional Groups by Infrared Spectroscopy, Anal. Chem., 24, 612–618, https://doi.org/10.1021/ac60064a002, 1952.
- 630 Heald, C. L., Kroll, J. H., Jimenez, J. L., Docherty, K. S., DeCarlo, P. F., Aiken, A. C., Chen, Q., Martin, S. T., Farmer, D. K., and Artaxo, P.: A Simplified Description of the Evolution of Organic Aerosol Composition in the Atmosphere, Geophys. Res. Lett., 37, https://doi.org/10.1029/2010GL042737, 2010.
  - Helland, I. S.: On the Structure of Partial Least Squares Regression, Communications in Statistics Simulation and Computation, 17, 581–607, https://doi.org/10.1080/03610918808812681, 1988.
- Hennigan, C. J., Sullivan, A. P., Collett, J. L., and Robinson, A. L.: Levoglucosan Stability in Biomass Burning Particles Exposed to Hydroxyl Radicals, Geophys. Res. Lett., 37, L09 806, https://doi.org/10.1029/2010GL043088, 2010.

- Heringa, M. F., DeCarlo, P. F., Chirico, R., Tritscher, T., Dommen, J., Weingartner, E., Richter, R., Wehrle, G., Prévôt, A. S. H., and Baltensperger, U.: Investigations of Primary and Secondary Particulate Matter of Different Wood Combustion Appliances with a High-Resolution Time-of-Flight Aerosol Mass Spectrometer, Atmos. Chem. Phys., 11, 5945–5957, https://doi.org/10.5194/acp-11-5945-2011, 2011.
- Iyer, S., Lopez-Hilfiker, F., Lee, B. H., Thornton, J. A., and Kurtén, T.: Modeling the Detection of Organic and Inorganic Compounds Using Iodide-Based Chemical Ionization, J. Phys. Chem. A, 120, 576–587, https://doi.org/10.1021/acs.jpca.5b09837, 2016.
- Jathar, S. H., Cappa, C. D., Wexler, A. S., Seinfeld, J. H., and Kleeman, M. J.: Multi-Generational Oxidation Model to Simulate Secondary Organic Aerosol in a 3-D Air Quality Model, Geosci. Model Dev., 8, 2553–2567, https://doi.org/10.5194/gmd-8-2553-2015, 2015.
- Kumar, N. K., Corbin, J. C., Bruns, E. A., Massabó, D., Slowik, J. G., Drinovec, L., Močnik, G., Prati, P., Vlachou, A., Baltensperger, U., Gysel, M., El-Haddad, I., and Prévôt, A. S. H.: Production of Particulate Brown Carbon during Atmospheric Aging of Residential Wood-Burning Emissions, Atmos. Chem. Phys., 18, 17 843–17 861, https://doi.org/10.5194/acp-18-17843-2018, 2018.
  - Kuzmiakova, A., Dillner, A. M., and Takahama, S.: An Automated Baseline Correction Protocol for Infrared Spectra of Atmospheric Aerosols Collected on Polytetrafluoroethylene (Teflon) Filters, Atmos. Meas. Tech., 9, 2615–2631, https://doi.org/10.5194/amt-9-2615-2016, 2016.
- 650 Lambe, A. T., Onasch, T. B., Croasdale, D. R., Wright, J. P., Martin, A. T., Franklin, J. P., Massoli, P., Kroll, J. H., Canagaratna, M. R., Brune, W. H., Worsnop, D. R., and Davidovits, P.: Transitions from Functionalization to Fragmentation Reactions of Laboratory Secondary Organic Aerosol (SOA) Generated from the OH Oxidation of Alkane Precursors, Environ. Sci. Technol., 46, 5430–5437, https://doi.org/10.1021/es300274t, 2012.
- Li, J., Li, J., Wang, G., Zhang, T., Dai, W., Ho, K. F., Wang, Q., Shao, Y., Wu, C., and Li, L.: Molecular Characteristics of Organic Compositions in Fresh and Aged Biomass Burning Aerosols, Science of The Total Environment, 741, 140 247, https://doi.org/10.1016/j.scitotenv.2020.140247, 2020.

- Li, Y. J., Yeung, J. W. T., Leung, T. P. I., Lau, A. P. S., and Chan, C. K.: Characterization of Organic Particles from Incense Burning Using an Aerodyne High-Resolution Time-of-Flight Aerosol Mass Spectrometer, Aerosol Sci. Tech., 46, 654–665, https://doi.org/10.1080/02786826.2011.653017, 2012.
- 660 Liu, S., Day, D. A., Shields, J. E., and Russell, L. M.: Ozone-Driven Daytime Formation of Secondary Organic Aerosol Containing Car-boxylic Acid Groups and Alkane Groups, Atmos. Chem. Phys., 11, 8321–8341, https://doi.org/10.5194/acp-11-8321-2011, 2011.

670

- Liu, S., Ahlm, L., Day, D. A., Russell, L. M., Zhao, Y., Gentner, D. R., Weber, R. J., Goldstein, A. H., Jaoui, M., Offenberg, J. H., Kleindienst, T. E., Rubitschun, C., Surratt, J. D., Sheesley, R. J., and Scheller, S.: Secondary Organic Aerosol Formation from Fossil Fuel Sources Contribute Majority of Summertime Organic Mass at Bakersfield, J. Geophys. Res.-Atmos., 117, https://doi.org/10.1029/2012JD018170, 2012.
- Lopez-Hilfiker, F. D., Pospisilova, V., Huang, W., Kalberer, M., Mohr, C., Stefenelli, G., Thornton, J. A., Baltensperger, U., Prevot, A. S. H., and Slowik, J. G.: An Extractive Electrospray Ionization Time-of-Flight Mass Spectrometer (EESI-TOF) for Online Measurement of Atmospheric Aerosol Particles, Atmos. Meas. Tech., 12, 4867–4886, https://doi.org/10.5194/amt-12-4867-2019, 2019.
- Maria, S. F., Russell, L. M., Turpin, B. J., and Porcja, R. J.: FTIR Measurements of Functional Groups and Organic Mass in Aerosol Samples over the Caribbean, Atmos. Environ., 36, 5185–5196, https://doi.org/10.1016/S1352-2310(02)00654-4, 2002.
- Maria, S. F., Russell, L. M., Turpin, B. J., Porcja, R. J., Campos, T. L., Weber, R. J., and Huebert, B. J.: Source Signatures of Carbon Monoxide and Organic Functional Groups in Asian Pacific Regional Aerosol Characterization Experiment (ACE-Asia) Submicron Aerosol Types, J. Geophys. Res.-Atmos., 108, 8637, https://doi.org/10.1029/2003JD003703, 2003.
- Nozière, B., Kalberer, M., Claeys, M., Allan, J., D'Anna, B., Decesari, S., Finessi, E., Glasius, M., Grgić, I., Hamilton, J. F., Hoffmann, T.,

  Iinuma, Y., Jaoui, M., Kahnt, A., Kampf, C. J., Kourtchev, I., Maenhaut, W., Marsden, N., Saarikoski, S., Schnelle-Kreis, J., Surratt, J. D.,

  Szidat, S., Szmigielski, R., and Wisthaler, A.: The Molecular Identification of Organic Compounds in the Atmosphere: State of the Art
  and Challenges, Chem. Rev., 115, 3919–3983, https://doi.org/10.1021/cr5003485, 2015.
  - Parks, D. A., Griffiths, P. R., Weakley, A. T., and Miller, A. L.: Quantifying Elemental and Organic Carbon in Diesel Particulate Matter by Mid-Infrared Spectrometry, Aerosol Sci. Tech., 0, 1–14, https://doi.org/10.1080/02786826.2021.1917764, 2021.
- Pavia, D. L., Lampman, G. M., Kriz, G. S., and Vyvyan, J. A.: Introduction to Spectroscopy, Brooks Cole, Belmont, CA, fourth edn., 2008.
  Reggente, M., Dillner, A. M., and Takahama, S.: Analysis of Functional Groups in Atmospheric Aerosols by Infrared Spectroscopy:
  Systematic Intercomparison of Calibration Methods for US Measurement Network Samples, Atmos. Meas. Tech., 12, 2287–2312, https://doi.org/10.5194/amt-12-2287-2019, 2019a.
  - Reggente, M., Höhn, R., and Takahama, S.: An Open Platform for Aerosol InfraRed Spectroscopy Analysis AIRSpec, Atmos. Meas. Tech., 12, 2313–2329, https://doi.org/10.5194/amt-12-2313-2019, 2019b.
  - Ruggeri, G.: On the Functional Group Composition of Organic Aerosol, Ph.D. thesis, Ecole Polytechnique Federale de Lausanne Lausanne (EPFL), https://doi.org/10.5075/epfl-thesis-7578, 2017.
  - Russell, L. M.: Aerosol Organic-Mass-to-Organic-Carbon Ratio Measurements, Environ. Sci. Technol., 37, 2982–2987, https://doi.org/10.1021/es026123w, 2003.
- Russell, L. M., Bahadur, R., Hawkins, L. N., Allan, J., Baumgardner, D., Quinn, P. K., and Bates, T. S.: Organic Aerosol Characterization by Complementary Measurements of Chemical Bonds and Molecular Fragments, Atmos. Environ., 43, 6100–6105, https://doi.org/10.1016/j.atmosenv.2009.09.036, 2009a.

- Russell, L. M., Takahama, S., Liu, S., Hawkins, L. N., Covert, D. S., Quinn, P. K., and Bates, T. S.: Oxygenated Fraction and Mass of Organic Aerosol from Direct Emission and Atmospheric Processing Measured on the R/V Ronald Brown during TEXAQS/GoMACCS 2006, J. Geophys. Res.-Atmos., 114, D00F05, https://doi.org/10.1029/2008JD011275, 2009b.
  - Russell, L. M., Takahama S., Liu S., Hawkins L. N., Covert D. S., Quinn P. K., and Bates T. S.: Oxygenated Fraction and Mass of Organic Aerosol from Direct Emission and Atmospheric Processing Measured on the R/V Ronald Brown during TEXAQS/GoMACCS 2006, J. Geophys. Res.-Atmos., 114, D00F05, https://doi.org/10.1029/2008JD011275, 2009c.
- Ruthenburg, T. C., Perlin, P. C., Liu, V., McDade, C. E., and Dillner, A. M.: Determination of Organic Matter and Organic Matter to Organic 700 Carbon Ratios by Infrared Spectroscopy with Application to Selected Sites in the IMPROVE Network, Atmos. Environ., 86, 47–57, https://doi.org/10.1016/j.atmosenv.2013.12.034, 2014.
  - Schneider, J., Weimer, S., Drewnick, F., Borrmann, S., Helas, G., Gwaze, P., Schmid, O., Andreae, M. O., and Kirchner, U.: Mass Spectrometric Analysis and Aerodynamic Properties of Various Types of Combustion-Related Aerosol Particles, Int. J. Mass Spectrom., 258, 37–49, https://doi.org/10.1016/j.ijms.2006.07.008, 2006.
- Shiraiwa, M., Ueda, K., Pozzer, A., Lammel, G., Kampf, C. J., Fushimi, A., Enami, S., Arangio, A. M., Fröhlich-Nowoisky, J., Fujitani, Y., Furuyama, A., Lakey, P. S. J., Lelieveld, J., Lucas, K., Morino, Y., Pöschl, U., Takahama, S., Takami, A., Tong, H., Weber, B., Yoshino, A., and Sato, K.: Aerosol Health Effects from Molecular to Global Scales, Environ. Sci. Technol., 51, 13545–13567, https://doi.org/10.1021/acs.est.7b04417, 2017.
- Subramanian, R., Khlystov, A. Y., Cabada, J. C., and Robinson, A. L.: Positive and Negative Artifacts in Particulate Organic Carbon Measurements with Denuded and Undenuded Sampler Configurations Special Issue of Aerosol Science and Technology on Findings from the Fine Particulate Matter Supersites Program, Aerosol Sci. Tech., 38, 27–48, https://doi.org/10.1080/02786820390229354, 2004.
  - Sun, Y. L., Zhang, Q., Anastasio, C., and Sun, J.: Insights into Secondary Organic Aerosol Formed via Aqueous-Phase Reactions of Phenolic Compounds Based on High Resolution Mass Spectrometry, Atmos. Chem. Phys., 10, 4809–4822, https://doi.org/10.5194/acp-10-4809-2010, 2010.
- Takahama, S., Johnson, A., and Russell, L. M.: Quantification of Carboxylic and Carbonyl Functional Groups in Organic Aerosol Infrared Absorbance Spectra, Aerosol Sci. Tech., 47, 310–325, https://doi.org/10.1080/02786826.2012.752065, 2013.
  - Takahama, S., Ruggeri, G., and Dillner, A. M.: Analysis of Functional Groups in Atmospheric Aerosols by Infrared Spectroscopy: Sparse Methods for Statistical Selection of Relevant Absorption Bands, Atmos. Meas. Tech., 9, 3429–3454, https://doi.org/10.5194/amt-9-3429-2016, 2016.
- Theodoritsi, G. N., Ciarelli, G., and Pandis, S. N.: Simulation of the Evolution of Biomass Burning Organic Aerosol with Different Volatility

  Basis Set Schemes in PMCAMx-SRv1.0, Geoscientific Model Development Discussions, pp. 1–33, https://doi.org/10.5194/gmd-2020-295, 2020.
  - Tiitta, P., Leskinen, A., Hao, L., Yli-Pirilä, P., Kortelainen, M., Grigonyte, J., Tissari, J., Lamberg, H., Hartikainen, A., Kuuspalo, K., Kortelainen, A.-M., Virtanen, A., Lehtinen, K. E. J., Komppula, M., Pieber, S., Prévôt, A. S. H., Onasch, T. B., Worsnop, D. R., Czech, H.,
- Zimmermann, R., Jokiniemi, J., and Sippula, O.: Transformation of Logwood Combustion Emissions in a Smog Chamber: Formation of Secondary Organic Aerosol and Changes in the Primary Organic Aerosol upon Daytime and Nighttime Aging, Atmos. Chem. Phys., 16, 13 251–13 269, https://doi.org/10.5194/acp-16-13251-2016, 2016.
  - Tolbert, A. and Ragauskas, A. J.: Advances in Understanding the Surface Chemistry of Lignocellulosic Biomass via Time-of-Flight Secondary Ion Mass Spectrometry, Energy Sci. Eng., 5, 5–20, https://doi.org/10.1002/ese3.144, 2017.

- 730 Turpin, B. J. and Lim, H.-J.: Species Contributions to PM2.5 Mass Concentrations: Revisiting Common Assumptions for Estimating Organic Mass, Aerosol Sci. Tech., 35, 602–610, https://doi.org/10.1080/02786820119445, 2001.
  - Volkamer, R., Jimenez, J. L., Martini, F. S., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L. T., Worsnop, D. R., and Molina, M. J.: Secondary Organic Aerosol Formation from Anthropogenic Air Pollution: Rapid and Higher than Expected, Geophys. Res. Lett., 33, https://doi.org/10.1029/2006GL026899, 2006.
- Wang, N., Kostenidou, E., Donahue, N. M., and Pandis, S. N.: Multi-Generation Chemical Aging of Alpha-Pinene Ozonolysis Products by Reactions with OH, Atmos. Chem. Phys., 18, 3589–3601, https://doi.org/10.5194/acp-18-3589-2018, 2018.
  - Wold, S., Martens, H., and Wold, H.: The Multivariate Calibration Problem in Chemistry Solved by the PLS Method, in: Matrix Pencils, edited by Kågström, B. and Ruhe, A., Lect. Notes Math., pp. 286–293, Springer Berlin Heidelberg, 1983.
- Wold, S., Johansson, E., and Cocchi, M.: 3D QSAR in Drug Design: Theory, Methods and Applications, ESCOM, Leiden, Holland, pp. 523–550, 1993.
  - Xu, W., Lambe, A., Silva, P., Hu, W., Onasch, T., Williams, L., Croteau, P., Zhang, X., Renbaum-Wolff, L., Fortner, E., Jimenez, J. L., Jayne, J., Worsnop, D., and Canagaratna, M.: Laboratory Evaluation of Species-Dependent Relative Ionization Efficiencies in the Aerodyne Aerosol Mass Spectrometer, Aerosol Sci. Tech., 52, 626–641, https://doi.org/10.1080/02786826.2018.1439570, 2018.
- Yazdani, A., Dillner, A. M., and Takahama, S.: Estimating Mean Molecular Weight, Carbon Number, and OM/OC with MidInfrared Spectroscopy in Organic Particulate Matter Samples from a Monitoring Network, Atmos. Meas. Tech., 14, 4805–4827, https://doi.org/10.5194/amt-14-4805-2021, 2021a.
  - Yazdani, A., Dudani, N., Takahama, S., Bertrand, A., Prévôt, A. S. H., El Haddad, I., and Dillner, A. M.: Characterization of Primary and Aged Wood Burning and Coal Combustion Organic Aerosols in Environmental Chamber and Its Implications for Atmospheric Aerosols, Atmos. Chem. Phys., 21, 10273–10293, https://doi.org/10.5194/acp-21-10273-2021, 2021b.
- 750 Zahardis, J., Geddes, S., and Petrucci, G. A.: Improved Understanding of Atmospheric Organic Aerosols via Innovations in Soft Ionization Aerosol Mass Spectrometry, Anal. Chem., 83, 2409–2415, https://doi.org/10.1021/ac102737k, 2011.
  - Zhang, Q., Worsnop, D. R., Canagaratna, M. R., and Jimenez, J. L.: Hydrocarbon-like and Oxygenated Organic Aerosols in Pittsburgh: Insights into Sources and Processes of Organic Aerosols, Atmos. Chem. Phys., 5, 3289–3311, https://doi.org/10.5194/acp-5-3289-2005, 2005.
- Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., and Sun, Y.: Understanding Atmospheric Organic Aerosols via Factor Analysis of Aerosol Mass Spectrometry: A Review, Anal. Bioanal. Chem., 401, 3045–3067, https://doi.org/10.1007/s00216-011-5355-y, 2011.
  - Ziemann, P. J. and Atkinson, R.: Kinetics, Products, and Mechanisms of Secondary Organic Aerosol Formation, Chem. Soc. Rev., 41, 6582–6605, https://doi.org/10.1039/C2CS35122F, 2012.