



1 **Estimation of sulfuric acid concentrations using ambient ion**
2 **composition and concentration data obtained by ion mass**
3 **spectrometry measurements (APi-TOF)**

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15
16 **Abstract**

17 Sulfuric acid (H₂SO₄, SA) is the key compound in atmospheric new particle formation. Therefore, it is crucial to
18 observe its concentration with sensitive instrumentation, such as chemical ionisation inlets coupled to
19 Atmospheric Pressure interface Time-of-Flight mass spectrometers (CI-APi-TOF). However, there are
20 environmental conditions and physical reasons when chemical ionisation cannot be used, for example in certain
21 remote places or flight measurements with limitations regarding chemicals. In these cases, it is important to
22 estimate the SA concentration based on ambient ion composition and concentration measurements that are
23 achieved by APi-TOF alone. Here we derive a theoretical expression to estimate SA concentration and validate it
24 with accurate CI-APi-TOF observations. The developed estimate works very well during daytime and with SA
25 concentrations above 2·10⁶ cm⁻³.

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28 **1 Introduction**

29 Sulfuric acid (H₂SO₄, SA) is the key compound in atmospheric new particle formation (e.g. Weber et al., 1995,
30 1996; Birmili et al., 2003; Kulmala et al., 2004; Kuang et al., 2008; Kerminen et al., 2010; Wang et al., 2011;
31 Kulmala et al., 2014; Yao et al., 2018; Cai et al., 2021), therefore it is crucial to have accurate observations of its
32 concentration. However, ambient concentrations of H₂SO₄ are low, commonly less than a part per trillion by
33 volume (~2·10⁷ molecules cm⁻³), making it challenging to detect this concentration. During the recent years there
34 has been clear steps towards the reliable detection of H₂SO₄ in the atmosphere, particularly via the development
35 of a Chemical Ionisation Atmospheric Pressure interface Time-of-Flight mass spectrometer (CI-APi-TOF,
36 Jokinen et al., 2012), using nitric acid as a reagent ion.

37



38 During the past decade or so, Atmospheric Pressure interface Time-of-Flight mass spectrometers (APi-TOF,
39 Junninen et al., 2010) have been deployed in several measurement campaigns where the use of a CI inlet was
40 either not possible or desired. In these instances, the APi-TOF only observed the composition and concentration
41 of ambient ions. The APi-TOF is capable of directly sampling and detecting naturally charged gas-phase ions,
42 including molecular clusters, and is often being used to detect clustering processes as a first step of new particle
43 formation on a molecular basis (e.g. Schobesberger et al., 2013; Jokinen et al., 2018; Beck et al., 2021). While a
44 CI-APi-TOF at best has a limit of detection around $\sim 10^4$ molecules cm^{-3} (\sim ppq level), the APi-TOF can detect
45 1% of the ambient ion concentration. With an average ion concentration of ~ 1000 cm^{-3} per polarity (Hirsikko et
46 al., 2011), the APi-TOF is measuring 10 ion $\text{cm}^{-3}\text{s}^{-1}$ with a limit of detection of ~ 0.01 counts per second, hence
47 0.1 ions cm^{-3} . This corresponds to approximately pps level ($100 \cdot 10^{-21}$), showing that the limit of detection of APi-
48 TOF in comparison to CI-APi-TOF is lower by five orders of magnitudes.

49
50 A detailed description of this instrument can be found in Junninen et al., (2010). Since concentrations of neutral
51 clusters are below the detection limit of CI-APi-TOF in many atmospheric conditions and environments, using
52 the APi-TOF is currently the only way to directly detect atmospheric clustering. Therefore, if we can estimate
53 H_2SO_4 concentration particularly during initial steps of new particle formation, based on the same dataset, we can
54 readily get better insight into the process itself.

55
56 Since there are only limited long term observations of H_2SO_4 concentrations, several proxies on this concentration
57 have also been developed (e.g. Petäjä et al., 2009; Mikkonen et al., 2011; Lu et al., 2019; Dada et al., 2020). These
58 proxies attempt to approximate the ambient H_2SO_4 concentrations using more readily measured quantities, in
59 particular the sulfur dioxide concentration, (UV) radiation intensity and pre-existing particle number size
60 distribution that can be used to calculate the condensation sink for gas-phase H_2SO_4 . In case we are able to estimate
61 the H_2SO_4 concentration from its influence on the ion mass spectra obtained by APi-TOF, that estimate may be
62 more accurate than the use of H_2SO_4 proxies.

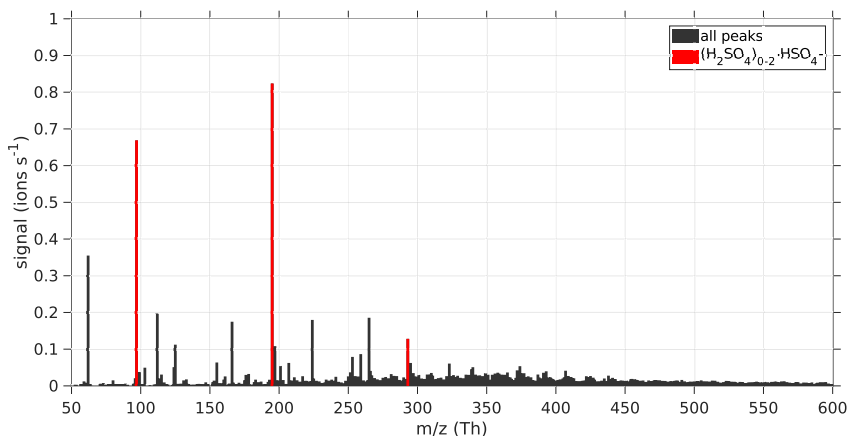
63
64 Motivated by the reasoning outlined above, we derive here an expression to estimate H_2SO_4 concentration based
65 primarily on APi-TOF observations and validate it.

66 67 68 **2 Theoretical estimation of sulfuric acid concentration with bisulphate ion and H_2SO_4 clusters**

69 Ambient ion mass spectra commonly contain clear evidence of gas-phase H_2SO_4 , predominantly in the form of
70 bisulphate ion (HSO_4^-) and its adducts involving H_2SO_4 , forming so-called dimers ($\text{H}_2\text{SO}_4 \cdot \text{HSO}_4^-$) as well as
71 larger clusters (Ehn et al., 2010). These are due to the efficient scavenging of negative charge by ambient H_2SO_4
72 via proton donation, and due to the high stability of the sulfuric acid-bisulphate ion clusters, in particular for the
73 dimer (Ortega et al., 2014). In order to estimate the sulfuric acid concentration (H_2SO_4) using the measured
74 naturally charged ions (see Fig. 1), we theoretically explain that concentration by following the bisulphate ion
75 HSO_4^- , herein denoted $\text{SA}_{\text{monomer}}$, the dimer cluster $\text{H}_2\text{SO}_4 \cdot \text{HSO}_4^-$ (SA_{dimer}) and trimer cluster $(\text{H}_2\text{SO}_4)_2 \cdot \text{HSO}_4^-$
76 ($\text{SA}_{\text{trimer}}$). Any other H_2SO_4 -containing ion clusters, in particular those larger than the $\text{SA}_{\text{trimer}}$, typically occur at
77 much smaller concentrations and will be neglected.



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81 **Figure 1** Mass spectrum from 50 to 600 Th measured with the Api-TOF on 24 May 2017 during the time period 08:00 – 18:00
 82 (local time). The bisulphate ion HSO_4^- and the H_2SO_4 clusters containing it were used for the estimation of H_2SO_4
 83 concentration and are coloured in red.

84
 85

86 If we assume that the concentration of $\text{SA}_{\text{monomer}}$ depends generally on its production rate (P_1) and that its loss is
 87 by condensation onto aerosol particles (condensation sink, CS) and its loss to the SA_{dimer} , we get the following
 88 equation for this concentration:

89

$$\frac{d[\text{SA}_{\text{monomer}}]}{dt} = P_1 - \text{CS} \cdot [\text{SA}_{\text{monomer}}] - P_2, \quad (1)$$

90

91 where $P_2 = k_1 \times [\text{SA}_{\text{monomer}}] \times [\text{H}_2\text{SO}_4]$ is the dimer production rate due to $\text{SA}_{\text{monomer}}\text{-H}_2\text{SO}_4$ collisions, and the
 92 collision rate k_1 is assumed to be constant.

93

94 For the dimer concentration we consider the production P_2 , the loss due to CS, and the clustering of the SA_{dimer}
 95 with H_2SO_4 with a rate constant k_2 :

96

$$\frac{d[\text{SA}_{\text{dimer}}]}{dt} = P_2 - \text{CS} \cdot [\text{SA}_{\text{dimer}}] - k_2 \cdot [\text{SA}_{\text{dimer}}] \cdot [\text{H}_2\text{SO}_4], \quad (2)$$

97

98 And with substituting P_2 , eq. 2 for SA_{dimer} changes to:

99



$$\frac{d[SA_{dimer}]}{dt} = k_1 \cdot [SA_{monomer}] \cdot [H_2SO_4] - CS \cdot [SA_{dimer}] - k_2 \cdot [SA_{dimer}] \cdot [H_2SO_4]. \quad (3)$$

100

101 Finally, to produce SA_{trimer} we consider the collision of the SA_{dimer} with H_2SO_4 and the loss to the CS. For the sake
102 of completeness, we would additionally have to consider the loss of $SA_{trimers}$ to form the tetramer $(H_2SO_4)_3 \cdot HSO_4^-$
103 , however this additional term is rather small and will therefore be neglected in this derivation. Therefore, we get
104 the simplified equation for SA_{trimer} :

105

$$\frac{d[SA_{trimer}]}{dt} = k_2 \cdot [SA_{dimer}] \cdot [H_2SO_4] - CS \cdot [SA_{trimer}]. \quad (4)$$

106

107 For simplification, we consider a pseudo-steady state condition for both dimers and trimers by setting the left-
108 hand side of eqs. 3 and 4 to be zero, which is justified when the dimer and trimer concentrations change at rates
109 smaller than their overall production and loss rates. Thereby, from eq. (3) we obtain:

110

$$k_1 \cdot [SA_{monomer}] \cdot [H_2SO_4] = CS \cdot [SA_{dimer}] + k_2 \cdot [SA_{dimer}] \cdot [H_2SO_4] \quad (5)$$

111

112 and from eq. (4) we obtain:

113

$$k_2 \cdot [SA_{dimer}] \cdot [H_2SO_4] = CS \cdot [SA_{trimer}]. \quad (6)$$

114

115 If we now deploy equation (6) in equation (5) and solve for H_2SO_4 , the result is:

116

$$k_1 \cdot [SA_{monomer}] \cdot [H_2SO_4] = CS \cdot [SA_{dimer}] + CS \cdot [SA_{trimer}], \quad (7)$$

117

118 Besides the steady-state assumption, it should be noted that in deriving eq. 8 monomers, dimers and trimers were
119 assumed to have the same loss rate (CS) onto pre-existing aerosol particles. This causes an additional, yet minor,
120 uncertainty in estimated H_2SO_4 concentrations, as such loss rates are dependent on the size/mass of the clusters
121 (e.g. Lehtinen et al., 2007).

122

123 From equation 8 we also see that the concentration of H_2SO_4 is proportional to relative concentrations of sulfuric
124 acid monomers, dimers and trimers clustered with the bisulphate ion:

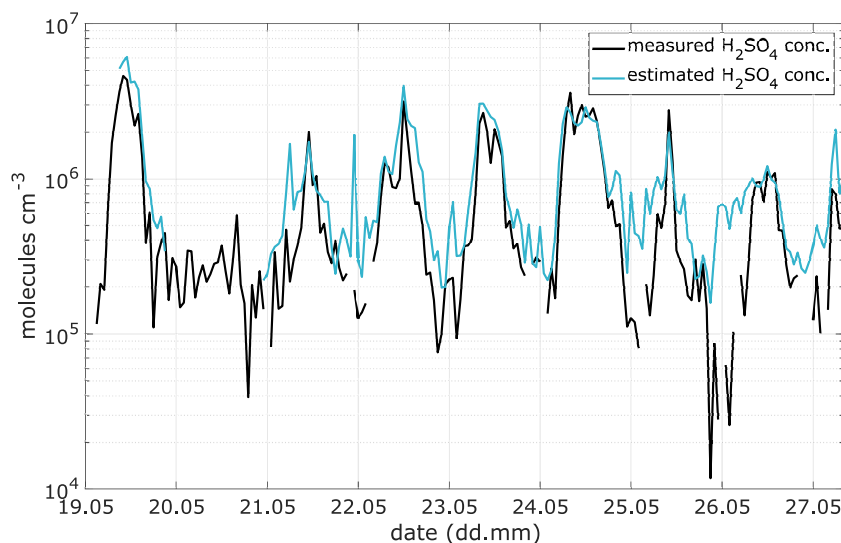
125

$$[H_2SO_4] \sim \frac{[SA_{dimer}] + [SA_{trimer}]}{[SA_{monomer}]} \quad (9)$$

126



127 To estimate the H_2SO_4 concentration with the ion mode API-TOF, we can therefore use this theoretical approach,
128 in particular Eq. 8. For the collision rate of H_2SO_4 with HSO_4^- we use $k_1 = 2 \cdot 10^{-9} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ as in [Loyejoy](#)
129 [et al., \(2004\)](#). The value of CS is calculated based on [Kulmala et al., \(2012\)](#). Even if the CS was unknown due,
130 for example, to the lack of particle measurements, the daytime variability of the H_2SO_4 concentration could still
131 be estimated only by using the relation of the H_2SO_4 -containing cluster with HSO_4^- , as it is proportional to the
132 H_2SO_4 concentration (see eq. 9).
133



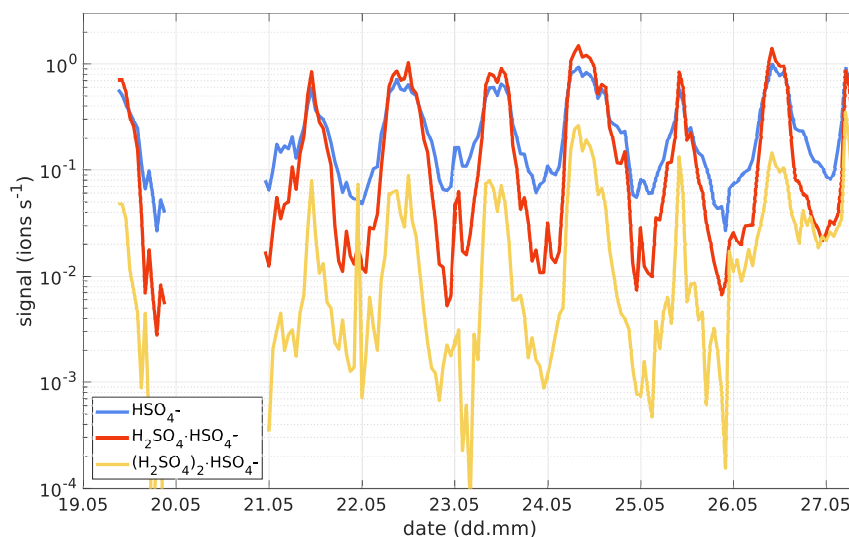
134
135 **Figure 2** Time series of measured H_2SO_4 concentration from the CI-API-TOF (black) and estimated H_2SO_4 concentration from
136 the API-TOF (blue) between 19 and 28 May 2017. The concentration is given in molecules cm^{-3} .
137

138

139 3 Validation

140 We tested the expression derived above using a dataset collected during inter-comparison measurements at the
141 SMEAR II station in Hyytiälä, Finland ([Hari and Kulmala, 2005](#)). In Fig. 2 we show the time series of estimated
142 and observed H_2SO_4 concentrations, based on API-TOF measurements together with Eq. 8 and CI-API-TOF
143 measurements, respectively. The estimated H_2SO_4 concentration agrees well with the measured one during most
144 of the daytime. During night-time, however, the estimated H_2SO_4 concentration seems to be considerably higher
145 than the measured one. This could be because the SA_{dimer} concentration is considerably lower than the $\text{SA}_{\text{monomer}}$
146 (see Fig. 3).
147

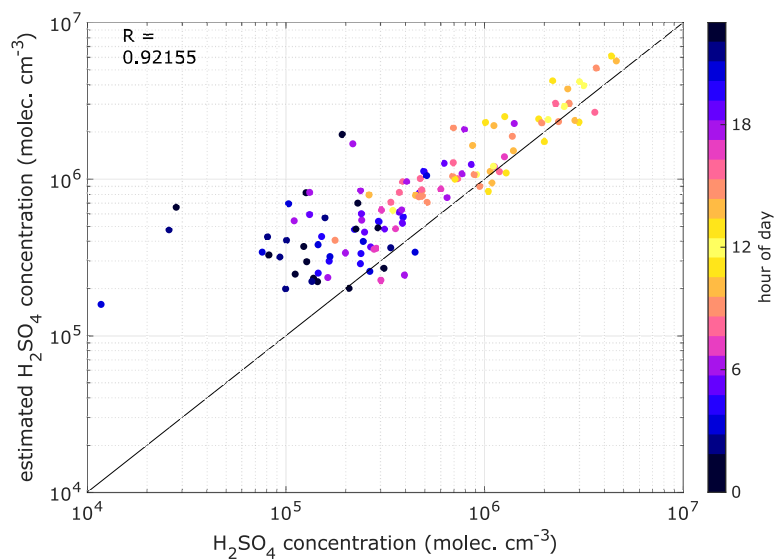
148 The scatter plot in Fig. 4 shows that the estimated H_2SO_4 concentrations agree well with the observed one when
149 the H_2SO_4 concentrations are larger than $2 \times 10^6 \text{ cm}^{-3}$, demonstrating that our method works particularly well
150 during conditions that favour the formation of H_2SO_4 -containing clusters.



151

152 **Figure 3** Time series of the bisulphate ion (HSO_4^- , $\text{SA}_{\text{monomer}}$), H_2SO_4 clustered with bisulphate ($\text{H}_2\text{SO}_4\cdot\text{HSO}_4^-$, SA_{dimer}) and
153 two H_2SO_4 molecules clustered with the bisulphate ion ($(\text{H}_2\text{SO}_4)_2\cdot\text{HSO}_4^-$, $\text{SA}_{\text{trimer}}$) between 19 and 28 May 2017. The
154 concentration is given in ions s^{-1} as measured by the APi-TOF.

155



156

157 **Figure 4** Measured H_2SO_4 concentration using a CI-APi-TOF (x-axis) versus estimated H_2SO_4 concentration based on APi-
158 TOF results (y-axis). The colour is indicating the hour of the day and the black line is the 1:1 ratio. Between 08:00 and 16:00
159 local time, the concentrations are agreeing well. The shown data contains the time period from 19 to 28 May 2017. The overall
160 correlation coefficient (Pearson) is 0.92.

161

162



163 **4 Conclusions**

164 Here we derived a theoretical expression to estimate H₂SO₄ concentrations based on APi-TOF measurements of
165 ambient ions. The estimation agrees well with the measured concentration during daytime and with high enough
166 H₂SO₄ concentrations, indicating that the estimation is able to give accurate enough H₂SO₄ concentrations during
167 most of the time when active clustering of sulfuric acid is taking as the initial step(s) of atmospheric new particle
168 formation.

169
170 The APi-TOF's "ion mode", i.e. direct ion sampling without chemical ionisation, remains a crucial tool in many
171 field deployments and laboratory studies, since it is extremely sensitive and allows for observing atmospheric
172 clustering molecule by molecule, which in most cases is impossible when relying on chemical ionization.
173 Therefore, having available a reliable estimate of H₂SO₄ concentration allows us to utilise the APi-TOF ion mode
174 even more effectively.

175

176 **Data availability**

177 The data can be accessed via Zenodo (not yet published, the correct link and doi will follow).

178

179 **Author contribution**

180 LJB, SS, VMK and MK designed the study. LJB performed the measurements. SS and LJB derived the equations.
181 LJB processed and analysed the data and performed the data visualisation. MK and VMK supervised the process.
182 All authors commented and edited the paper.

183

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193 References

- 194 Beck, L. J., Sarnela, N., Junninen, H., Hoppe, C. J. M., Garmash, O., Bianchi, F., Riva, M., Rose, C.,
195 Peräkylä, O., Wimmer, D., Kausiala, O., Jokinen, T., Ahonen, L., Mikkilä, J., Hakala, J., He, X.-C.,
196 Kontkanen, J., Wolf, K. K. E., Cappelletti, D., Mazzola, M., Traversi, R., Petroselli, C., Viola, A. P., Vitale,
197 V., Lange, R., Massling, A., Nøjgaard, J. K., Krejci, R., Karlsson, L., Zieger, P., Jang, S., Lee, K., Vakkari,
198 V., Lampilahti, J., Thakur, R. C., Leino, K., Kangasluoma, J., Duplissy, E.-M., Siivola, E., Marbouti, M.,
199 Tham, Y. J., Saiz-Lopez, A., Petäjä, T., Ehn, M., Worsnop, D. R., Skov, H., Kulmala, M., Kerminen, V.-M.,
200 and Sipilä, M.: Differing Mechanisms of New Particle Formation at Two Arctic Sites, *Geophysical Research*
201 *Letters*, 48, e2020GL091334, <https://doi.org/10.1029/2020GL091334>, 2021.
- 202 Birmili, W., Berresheim, H., Plass-Dülmer, C., Elste, T., Gilge, S., Wiedensohler, A., and Uhrner, U.: The
203 Hohenpeissenberg aerosol formation experiment (HAFEX): a long-term study including size-resolved
204 aerosol, H₂SO₄, OH, and monoterpenes measurements, 3, 361–376, <https://doi.org/10.5194/acp-3-361-2003>,
205 2003.
- 206 Cai, R., Yan, C., Yang, D., Yin, R., Lu, Y., Deng, C., Fu, Y., Ruan, J., Li, X., Kontkanen, J., Zhang, Q.,
207 Kangasluoma, J., Ma, Y., Hao, J., Worsnop, D. R., Bianchi, F., Paasonen, P., Kerminen, V.-M., Liu, Y.,
208 Wang, L., Zheng, J., Kulmala, M., and Jiang, J.: Sulfuric acid–amine nucleation in urban Beijing, 21, 2457–
209 2468, <https://doi.org/10.5194/acp-21-2457-2021>, 2021.
- 210 Dada, L., Yliviikka, I., Baalbaki, R., Li, C., Guo, Y., Yan, C., Yao, L., Sarnela, N., Jokinen, T., Daellenbach,
211 K. R., Yin, R., Deng, C., Chu, B., Nieminen, T., Wang, Y., Lin, Z., Thakur, R. C., Kontkanen, J.,
212 Stolzenburg, D., Sipilä, M., Hussein, T., Paasonen, P., Bianchi, F., Salma, I., Weidinger, T., Pikridas, M.,
213 Sciare, J., Jiang, J., Liu, Y., Petäjä, T., Kerminen, V.-M., and Kulmala, M.: Sources and sinks driving sulfuric
214 acid concentrations in contrasting environments: implications on proxy calculations, 20, 11747–11766,
215 <https://doi.org/10.5194/acp-20-11747-2020>, 2020.
- 216 Ehn, M., Junninen, H., Petäjä, T., Kurtén, T., Kerminen, V.-M., Schobesberger, S., Manninen, H. E., Ortega,
217 I. K., Vehkamäki, H., Kulmala, M., and Worsnop, D. R.: Composition and temporal behavior of ambient
218 ions in the boreal forest, 10, 8513–8530, <https://doi.org/10.5194/acp-10-8513-2010>, 2010.
- 219 Hirsikko, A., Nieminen, T., Gagné, S., Lehtipalo, K., Manninen, H. E., Ehn, M., Hörrak, U., Kerminen, V.-
220 M., Laakso, L., McMurry, P. H., Mirme, A., Mirme, S., Petäjä, T., Tammet, H., Vakkari, V., Vana, M., and
221 Kulmala, M.: Atmospheric ions and nucleation: a review of observations, 11, 767–798,
222 <https://doi.org/10.5194/acp-11-767-2011>, 2011.
- 223 Jokinen, T., Sipilä, M., Kontkanen, J., Vakkari, V., Tisler, P., Duplissy, E.-M., Junninen, H., Kangasluoma,
224 J., Manninen, H. E., Petäjä, T., Kulmala, M., Worsnop, D. R., Kirkby, J., Virkkula, A., and Kerminen, V.-
225 M.: Ion-induced sulfuric acid–ammonia nucleation drives particle formation in coastal Antarctica, *Sci Adv*,
226 4, eaat9744, <https://doi.org/10.1126/sciadv.aat9744>, 2018.
- 227 Junninen, H., Ehn, M., Petäjä, T., Luosujärvi, L., Kotiaho, T., Kostianen, R., Rohner, U., Gonin, M., Fuhrer,
228 K., Kulmala, M., and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion
229 composition, 3, 1039–1053, <https://doi.org/10.5194/amt-3-1039-2010>, 2010.
- 230 Kerminen, V.-M., Petäjä, T., Manninen, H. E., Paasonen, P., Nieminen, T., Sipilä, M., Junninen, H., Ehn,
231 M., Gagné, S., Laakso, L., Riipinen, I., Vehkamäki, H., Kurten, T., Ortega, I. K., Dal Maso, M., Brus, D.,
232 Hyvärinen, A., Lihavainen, H., Leppä, J., Lehtinen, K. E. J., Mirme, A., Mirme, S., Hörrak, U., Berndt, T.,
233 Stratmann, F., Birmili, W., Wiedensohler, A., Metzger, A., Dommen, J., Baltensperger, U., Kiendler-Scharr,
234 A., Mentel, T. F., Wildt, J., Winkler, P. M., Wagner, P. E., Petzold, A., Minikin, A., Plass-Dülmer, C.,
235 Pöschl, U., Laaksonen, A., and Kulmala, M.: Atmospheric nucleation: highlights of the EUCAARI project
236 and future directions, 10, 10829–10848, <https://doi.org/10.5194/acp-10-10829-2010>, 2010.
- 237 Kuang, C., McMurry, P. H., McCormick, A. V., and Eisele, F. L.: Dependence of nucleation rates on sulfuric
238 acid vapor concentration in diverse atmospheric locations, 113, <https://doi.org/10.1029/2007JD009253>,
239 2008.



- 240 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V.-M., Birmili, W., and
241 McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a review of observations,
242 *Journal of Aerosol Science*, 35, 143–176, <https://doi.org/10.1016/j.jaerosci.2003.10.003>, 2004.
- 243 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M., Aalto, P.
244 P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and Kerminen, V.-M.:
245 Measurement of the nucleation of atmospheric aerosol particles, 7, 1651–1667,
246 <https://doi.org/10.1038/nprot.2012.091>, 2012.
- 247 Kulmala, M., Petäjä, T., Ehn, M., Thornton, J., Sipilä, M., Worsnop, D. R., and Kerminen, V.-M.: Chemistry
248 of Atmospheric Nucleation: On the Recent Advances on Precursor Characterization and Atmospheric
249 Cluster Composition in Connection with Atmospheric New Particle Formation, 65, 21–37,
250 <https://doi.org/10.1146/annurev-physchem-040412-110014>, 2014.
- 251 Lehtinen, K. E. J., Dal Maso, M., Kulmala, M., and Kerminen, V.-M.: Estimating nucleation rates from
252 apparent particle formation rates and vice versa: Revised formulation of the Kerminen–Kulmala equation,
253 *Journal of Aerosol Science*, 38, 988–994, <https://doi.org/10.1016/j.jaerosci.2007.06.009>, 2007.
- 254 Lovejoy, E. R., Curtius, J., and Froyd, K. D.: Atmospheric ion-induced nucleation of sulfuric acid and water,
255 109, <https://doi.org/10.1029/2003JD004460>, 2004.
- 256 Lu, Y., Yan, C., Fu, Y., Chen, Y., Liu, Y., Yang, G., Wang, Y., Bianchi, F., Chu, B., Zhou, Y., Yin, R.,
257 Baalbaki, R., Garmash, O., Deng, C., Wang, W., Liu, Y., Petäjä, T., Kerminen, V.-M., Jiang, J., Kulmala,
258 M., and Wang, L.: A proxy for atmospheric daytime gaseous sulfuric acid concentration in urban Beijing,
259 19, 1971–1983, <https://doi.org/10.5194/acp-19-1971-2019>, 2019.
- 260 Mikkonen, S., Romakkaniemi, S., Smith, J. N., Korhonen, H., Petäjä, T., Plass-Duelmer, C., Boy, M.,
261 McMurry, P. H., Lehtinen, K. E. J., Joutsensaari, J., Hamed, A., Mauldin III, R. L., Birmili, W., Spindler,
262 G., Arnold, F., Kulmala, M., and Laaksonen, A.: A statistical proxy for sulphuric acid concentration, 11,
263 11319–11334, <https://doi.org/10.5194/acp-11-11319-2011>, 2011.
- 264 Ortega, I. K., Olenius, T., Kupiainen-Määttä, O., Loukonen, V., Kurtén, T., and Vehkamäki, H.: Electrical
265 charging changes the composition of sulfuric acid–ammonia/dimethylamine clusters, 14, 7995–8007,
266 <https://doi.org/10.5194/acp-14-7995-2014>, 2014.
- 267 Petäjä, T., Mauldin, I. I. I., Kosciuch, E., McGrath, J., Nieminen, T., Paasonen, P., Boy, M., Adamov, A.,
268 Kotiaho, T., and Kulmala, M.: Sulfuric acid and OH concentrations in a boreal forest site, 9, 7435–7448,
269 <https://doi.org/10.5194/acp-9-7435-2009>, 2009.
- 270 Schobesberger, S., Junninen, H., Bianchi, F., Lönn, G., Ehn, M., Lehtipalo, K., Dommen, J., Ehrhart, S.,
271 Ortega, I. K., Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy, J., Almeida, J., Amorim,
272 A., Breitenlechner, M., Downard, A. J., Dunne, E. M., Flagan, R. C., Kajos, M., Keskinen, H., Kirkby, J.,
273 Kupc, A., Kürten, A., Kurtén, T., Laaksonen, A., Mathot, S., Onnela, A., Praplan, A. P., Rondo, L., Santos,
274 F. D., Schallhart, S., Schnitzhofer, R., Sipilä, M., Tomé, A., Tsagkogeorgas, G., Vehkamäki, H., Wimmer,
275 D., Baltensperger, U., Carslaw, K. S., Curtius, J., Hansel, A., Petäjä, T., Kulmala, M., Donahue, N. M., and
276 Worsnop, D. R.: Molecular understanding of atmospheric particle formation from sulfuric acid and large
277 oxidized organic molecules, *PNAS*, 110, 17223–17228, <https://doi.org/10.1073/pnas.1306973110>, 2013.
- 278 Wang, Z. B., Hu, M., Yue, D. L., Zheng, J., Zhang, R. Y., Wiedensohler, A., Wu, Z. J., Nieminen, T., and
279 Boy, M.: Evaluation on the role of sulfuric acid in the mechanisms of new particle formation for Beijing
280 case, 11, 12663–12671, <https://doi.org/10.5194/acp-11-12663-2011>, 2011.
- 281 Weber, R. J., McMurry, P. H., Eisele, F. L., and Tanner, D. J.: Measurement of Expected Nucleation
282 Precursor Species and 3–500-nm Diameter Particles at Mauna Loa Observatory, Hawaii, 52, 2242–2257,
283 [https://doi.org/10.1175/1520-0469\(1995\)052<2242:MOENPS>2.0.CO;2](https://doi.org/10.1175/1520-0469(1995)052<2242:MOENPS>2.0.CO;2), 1995.



- 284 Weber, R. J., Marti, J. J., McMurry, P. H., Eisele, F. L., Tanner, D. J., and Jefferson, A.: Measured
285 Atmospheric New Particle Formation Rates: Implications for Nucleation Mechanisms, 151, 53–64,
286 <https://doi.org/10.1080/00986449608936541>, 1996.
- 287 Yao, L., Garmash, O., Bianchi, F., Zheng, J., Yan, C., Kontkanen, J., Junninen, H., Mazon, S. B., Ehn, M.,
288 Paasonen, P., Sipilä, M., Wang, M., Wang, X., Xiao, S., Chen, H., Lu, Y., Zhang, B., Wang, D., Fu, Q.,
289 Geng, F., Li, L., Wang, H., Qiao, L., Yang, X., Chen, J., Kerminen, V.-M., Petäjä, T., Worsnop, D. R.,
290 Kulmala, M., and Wang, L.: Atmospheric new particle formation from sulfuric acid and amines in a Chinese
291 megacity, 361, 278–281, <https://doi.org/10.1126/science.aao4839>, 2018.
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