https://doi.org/10.5194/amt-2021-263
Preprint. Discussion started: 6 September 2021

Atmospheric Measurement Techniques
Discussions
Open Access

A Systematic Re-evaluation of Methods for Quantification of Bulk Particle-phase Organic Nitrates Using Real-time Aerosol Mass Spectrometry

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Abstract

Organic nitrate ($RONO_2$) formation in the atmosphere represents a sink of NO_x ($NO_x = NO + NO_2$) and termination of the NO_x/HO_x ($HO_x = HO_2 + OH$) ozone formation and radical propagation cycles, can act as a NO_x reservoir transporting reactive nitrogen, and contributes to secondary organic aerosol formation. While some fraction of $RONO_2$ is thought to reside in the particle phase, particle-phase organic nitrates ($pRONO_2$) are infrequently measured and thus poorly understood. There is an increasing prevalence of aerosol mass spectrometer (AMS) instruments, which have shown promise for determining quantitative total organic nitrate functional group contribution to aerosols. A simple approach that relies on the relative intensities of NO^+ and NO_2^+ ions in the AMS spectrum, the calibrated NO_x^+ ratio for NH_4NO_3, and the inferred ratio for $pRONO_2$ has been proposed as a way to apportion the total nitrate signal to NH_4NO_3 and $pRONO_2$. This method is increasingly being applied to field and laboratory data. However, the methods applied have been largely inconsistent and poorly characterized, and therefore, a detailed evaluation is timely. Here, we compile an extensive survey of NO_x^+ ratios measured for various $pRONO_2$ compounds and mixtures from multiple AMS instruments, groups, and laboratory and field measurements. We show that, in the absence of $pRONO_2$ standards, the $pRONO_2$ NO_x^+ ratio can be estimated using a ratio referenced to the calibrated NH_4NO_3 ratio, a so-called "Ratio-of-Ratios" method (*RoR*=2.75±0.41). We systematically explore the basis for quantifying $pRONO_2$ (and NH_4NO_3) with the *RoR* method using ground and aircraft field measurements conducted over a large range of conditions. The method is compared to another AMS method (positive matrix factorization, PMF) and other $pRONO_2$ and related (e.g., total gas + particle $RONO_2$) measurements, generally showing good agreement / correlation. A broad survey of ground and aircraft AMS measurements shows a pervasive trend of higher fractional contribution of $pRONO_2$ to total nitrate with lower total nitrate concentrations, which generally corresponds to shifts from urban-influenced to rural/remote regions. Compared to ground campaigns, observations from all aircraft campaigns showed substantially lower $pRONO_2$ contributions at mid ranges of total nitrate (0.01-0.1 up to 2-5 $\mu g\ m^{-3}$), suggesting that the balance of effects controlling NH_4NO_3 and $pRONO_2$ formation and lifetimes — such as higher humidity, lower temperatures, greater dilution, different sources, higher particle acidity, and $pRONO_2$ hydrolysis (possibly accelerated by particle acidity) — favors lower $pRONO_2$ contributions for those environments and altitudes sampled.

1 Introduction

Organic nitrate ($RONO_2$) formation in the atmosphere, through oxidation of VOCs (volatile organic compounds) in the presence of NO_x ($NO_x = NO + NO_2$), represents a sink of NO_x and termination of the catalytic NO_x/HO_x ($HO_x = OH + HO_2$) ozone formation and radical propagation cycles, can act as a NO_x reservoir transporting (or removing) reactive nitrogen, and contribute to secondary organic aerosol formation (Zare et al., 2018 and references therein). Particle-phase organic nitrates ($pRONO_2$) have been shown to contribute substantial mass to organic aerosol (OA) (Ng et al., 2017 and references therein), can provide insight into the chemistry controlling SOA formation (e.g., Pye et al., 2015; Xu et al., 2015b; Lee et al., 2016; Ng et al., 2017), may constitute a semivolatile component of OA and dynamically partition between the gas- and particle-phases (e.g., Fry et al., 2013; Rollins et al., 2013; Pye et al., 2015), and represent a loss mechanism for $RONO_2$ or reactive nitrogen oxides (e.g., via hydrolysis or deposition) (Fisher et al., 2016; Lee et al., 2016; Zare et al., 2018). However, $pRONO_2$ have infrequently been measured in ambient air until recently and thus are still poorly understood (Ng et al., 2017).

The recent emergence of a variety of online and offline methods of both speciated and bulk $pRONO_2$ and their applications to ambient aerosol measurements are summarized in Ng et al. (2017). Instrumentation and methods include: (online bulk) aerosol mass spectrometry (AMS; (Jayne et al., 2000)) and its monitoring versions (known as Aerosol Chemical Speciation Monitors, ACSM; (Ng et al., 2011; Fröhlich et al., 2013)); thermal dissociation - laser induced fluorescence (TD-LIF; (Day et al., 2002)); (online speciated) filter inlet for gases and aerosols (FIGAERO) - chemical ionization mass spectrometry (CIMS) (Lopez-Hilfiker et al., 2014); (offline speciated) high-pressure liquid chromatography - mass spectrometry (HPLC/MS) often with electrospray ionization (ESI) (Surratt et al., 2006); (offline bulk) Fourier Transform InfraRed (FTIR) spectroscopy (Maria et al., 2002). While speciated methods can provide more detailed source or mechanistic information, they are slow and, to date, none (online nor offline) has demonstrated quantitative measurement of the bulk of $pRONO_2$ for ambient measurements. Therefore, bulk measurements provide useful constraints on the budgets, formation and loss rates of gas- and aerosol-phase $RONO_2$ in the atmosphere; and fast online methods are essential when ambient concentrations are rapidly changing, especially for aircraft sampling.

For most field applications of the AMS, typically aerosol nitrate concentrations have been reported as a single total (organic plus inorganic) concentration, due to the fact that nearly all of the signal of the nitrate functional group for any nitrate type (or nitrite) is measured at a couple of common ion peaks (NO^+ and NO_2^+ in high-resolution (HR) instruments or *m/z* 30 and *m/z* 46 in unit mass resolution (UMR) instruments) (Farmer et al., 2010). Early on in the application of the AMS, an implicit assumption was often made that ammonium nitrate (NH_4NO_3) typically dominated aerosol nitrate, based on early urban studies that showed semivolatile behavior consistent with NH_4NO_3 (e.g., Jimenez et al., 2003; Hogrefe et al., 2004; Zhang et al., 2004). However, a few early reports on field measurements using UMR AMS (Allan et al., 2004b, 2006) showed that the *m/z* 46 - to - *m/z* 30 ratio (hereinafter “46/30 ratio”) was too low to be associated with only NH_4NO_3, suggesting substantial contributions from mineral nitrates ($NaNO_3$, $Ca(NO_3)_2$), $pRONO_2$, or possibly other reduced organo-nitrogen, or organic ion interferences. In a study focusing on cluster analysis of ambient (UMR) AMS spectra, Marcolli et al. (2006) also reported 46/30 ratios substantially smaller than NH_4NO_3 and found several spectra cluster categories with dominant *m/z* 30 peaks (but not *m/z* 46) and suggested that these signals may be associated with organic nitrates. Similarly, Alfarra et al. (2006) reported 46/30 ratios from chamber-generated SOA (photooxidation of trimethyl benzene and α-pinene) ~2–4 times lower than NH_4NO_3, which they attributed to $pRONO_2$ or nitro-compounds. A few years later, reports from chamber studies where $pRONO_2$-rich SOA was generated (β-pinene or isoprene +NO_3 radicals), using an HR-AMS, showed NO_2^+/NO^+ ratios (hereafter "NO_x^+ ratio") ~2–4 times lower than pure NH_4NO_3 (Fry et al., 2009; Rollins et al., 2009).

Subsequently, broader surveys of the fragmentation patterns of aerosol nitrates (and nitrites) in the AMS were reported, including consistently low NO_x^+ ratios for $pRONO_2$ (Bruns et al., 2010; Farmer et al., 2010). Farmer et al. (2010) evaluated the fragmentation patterns of single-component $pRONO_2$ isolated from SOA, and showed that ~95% the nitrogen-containing signal was observed as NO_x^+ ions with the balance as HNO_3^+ and very little signal at $C_xH_yO_zN^+$ ions. Farmer et al. evaluated several methods for constraining $pRONO_2$ contribution to AMS nitrate signal including using: 1) NO_x^+ ratios, 2) HNO_3^+ ions, 3) $C_xH_yO_zN^+$ ions, 4) "ammonium balance", and 5) AMS total nitrate comparison to inorganic nitrate-only measured with another instrument (typically ion chromatography-based). For the urban dataset evaluated in that study, all methods appeared to be associated with relatively large uncertainties. Bruns et al. (2010) reported NO_x^+ ratios for SOA formed from several monoterpenes and isoprene (with NO_3 radicals) as well as $NaNO_3$ and $NaNO_2$ (with the sodium salts showing greatly reduced NO_x^+ ratios). Other studies have used the ammonium balance (hereafter NH_{4_Bal}) of AMS data, or comparisons to other instruments to estimate $pRONO_2$ content (Aiken et al., 2009; Zaveri et al., 2010; Docherty et al., 2011; Häkkinen et al., 2012; Xu et al., 2015a); however, in most cases, uncertainties were large or not assessed. Since the Farmer et al. study, several other laboratory studies reported NO_x^+ ratios for $pRONO_2$-containing SOA, which are summarized in Sect. 3. Additionally, a number of analyses of field studies have used the NO_x^+ ratios (or its 46/30 UMR equivalent) to support qualitative or semi-quantitative statements about the presence (or low contribution) of $pRONO_2$ (Setyan et al., 2012; Brown et al., 2013; Xu et al., 2016; Schneider et al., 2017; Bottenus et al., 2018) or to quantify $pRONO_2$ (Fry et al., 2013, 2018; Ayres et al., 2015; Kostenidou et al., 2015; Xu et al., 2015a, 2021; Fisher et al., 2016; Kiendler-Scharr et al., 2016; Lee et al., 2016, 2019; Nault et al., 2016; Zhou et al., 2016; Zhu et al., 2016, 2021; Florou et al., 2017; Palm et al., 2017; Brito et al., 2018; de Sá et al., 2018, 2019; Reyes-Villegas et al., 2018; Schulz et al., 2018; Avery et al., 2019; Dai et al., 2019; Huang et al., 2019a, 2019b; Yu et al., 2019; Chen et al., 2020, 2021). Yu et al. (2019) also used the particle size dependence of the 46/30 ratio to investigate particle size and temporal (diurnal and seasonal) trends in $pRONO_2$. Other studies have used positive matrix factorization (PMF) of AMS spectra including both the OA and NO_x^+ signals to quantify $pRONO_2$ (Sun et al., 2012; Hao et al., 2014; Xu et al., 2015a; Zhang et al., 2016; Kortelainen et al., 2017; Yu et al., 2019; Zhu et al., 2021). Recently, Xu et al., (2021) demonstrated another method, using AMS thermal denuder measurements.Thus there is promising use of AMS measurements for quantifying bulk $pRONO_2$ functional group contribution to ambient aerosols (and in addition, providing higher quality NH_4NO_3 concentrations). However, the methods have not been standardized and uncertainties of the different methods have not been well-characterized, and were reported to be large by at least some studies. Together with the increasing prevalence of AMS (and ACSM) field measurements, a detailed evaluation of $pRONO_2$ quantification methods is timely. Here we explore the application of the AMS NO_x^+ ratio method to separate and quantify inorganic and organic nitrate and discuss the methods in detail, as well as comparison to other methods, and some scientific applications. In addition to drawing from available literature whenever possible, new analyses for several field and laboratory datasets are used extensively throughout this manuscript to explore and support findings. Descriptions of those datasets and data processing methods can be found in Supp. Info. Sect. S1 (including Fig. S1).

2 Previous use and methods for pRONO$_2$ quantification using AMS NO$_x^+$ ratios

An equation for quantitative apportionment of the AMS nitrate signal into pRONO$_2$ and NH$_4$NO$_3$ using the NO$_x^+$ ratio was first presented by Farmer et al. (2010) (equation 1 from Farmer et al., and derived in their supporting information, here substituting different notation for some terms for consistency with this manuscript):

$$f_{pRONO_2} = \frac{(R_{ambient} - R_{NH_4NO_3})(1 + R_{pRONO_2})}{(R_{pRONO_2} - R_{NH_4NO_3})(1 + R_{ambient})} \qquad (1)$$

where f_{pRONO2} is the fraction of total AMS nitrate (hereafter pNO$_3$) that is pRONO$_2$, and R_{NH4NO3}, R_{pRONO2}, and R_{ambient} are the NO$_x^+$ ratios (NO$_2^+$/NO$^+$) for pure NH$_4$NO$_3$, pure pRONO$_2$, and the ambient aerosol nitrate mixture measured, respectively. Note that here we use the NO$_2^+$/NO$^+$ ratio for all terms, while Farmer et al. and some others have used NO$^+$/NO$_2^+$. This formulation is preferred since NO$_2^+$ tends to be lower than NO$^+$ for all nitrates, and thus using NO$_2^+$/NO$^+$ avoids ratios trending toward infinity as detection limits are approached. This usage has been applied in several publications, such as Fry et al. (2013) and Kiendler-Scharr et al. (2016), as presented in equations 11 and 1 in those papers, respectively. The equation is identical regardless of the inversion of the NO$_x^+$ ratio. That can be shown by simply swapping all the instances of NO and NO$_2$ in the definitions and derivation shown in Farmer et al. or by substituting 1/R$_x$ for each ratio term in Eq. 1 above, multiplying all parenthetical terms by $R_{\mathrm{ambient}}R_{\mathrm{NH4NO3}}R_{\mathrm{pRONO2}}$, factoring out the same term in the numerator and denominator then canceling, and finally multiplying the first parenthetical terms in the numerator and denominator by -1. While typically R_{NH4NO3} is measured frequently as pure NH$_4$NO$_3$ is periodically sampled by the AMS as a primary calibrant for sensitivity (Canagaratna et al., 2007), regular calibration using pRONO$_2$ is generally not practical. Moreover, it is not immediately clear that all pRONO$_2$ produce the same R_{pRONO2} in the AMS. Values reported in the literature for R_{NH4NO3} and R_{pRONO2} both appear to have a substantial range (factor of ~3) and generally R_{pRONO2} is 2–4 times lower than R_{NH4NO3} (see Sects. 1 and 3).

Several studies have applied Eq. 1 to quantify pRONO$_2$ and NH$_4$NO$_3$, using different assumptions regarding R_{pRONO2}. Farmer et al. (2010) applied their measurements of R_{pRONO2} from their lab study to estimate an upper limit of 50% for the pRONO$_2$ contribution to pNO$_3$ for the urban SOAR campaign, substantially higher than with other methods they applied. They considered that method to be a high upper limit, due to the possible influence of non-refractory nitrates. However, we note that the R_{pRONO2} used in that early study was nearly a factor of two different than we suggest in this study, in the direction favorable to higher pRONO$_2$ fractions. For calculation of pRONO$_2$ for the BEACHON-RoMBAS campaign, Fry et al. (2013) assert that R_{NH4NO3} and R_{pRONO2} likely co-vary for an instrument and therefore define the term "ratio-of-ratios" (hereafter $RoR = R_{\mathrm{NH4NO3}}/R_{\mathrm{pRONO2}}$) in order to estimate R_{pRONO2} from in-field R_{NH4NO3} measurements and literature reports of R_{pRONO2} and R_{NH4NO3}. The *RoR* value applied by Fry et al. (2013) was 2.25, based on the Farmer et al. (2010) average. On the other hand, in an analysis of pRONO$_2$ contribution to OA throughout Europe, Kiendler-Scharr et al. (2016) applied a fixed R_{pRONO2} of 0.1 based on literature reports of R_{pRONO2}, and the argument that it was the minimum ratio observed in the ambient datasets examined (noting that "such low ratios of NO$_2^+$/NO$^+$ were also detected in some data sets where R_{NH4NO3} was reported high"). Those authors state that their approach represents a lower limit of pRONO$_2$. Similarly, Brito et al. (2018), Schulz et al. (2018), Huang et al. (2019a, 2019b), and Avery et al. (2019), applied a fixed R_{pRONO2} of 0.1 (citing Kiendler-Sharr et al. (2016)) for aircraft measurements in West Africa, aircraft measurements in the Amazon, rural forest and urban sites in Germany, and seasonal variations of indoor/outdoor air, respectively. The same method has been applied to laboratory studies of biomass burning aging (Tiitta et al., 2016), composition from photooxidation of terpenes (Zhao et al., 2018; Pullinen et al., 2020), and the composition, optical properties, and aging of particles from a wide variety of biomass burning fuel sources (Cappa et al., 2020; McClure et al., 2020). However, in the latter study, the organic component is classified as "organonitrogen", assuming it includes contributions from both organic nitrate and nitro-organic (i.e. nitroaromatics) functional groups (and assumed to have the same NO_x^+ ratio).

In a regional and seasonal survey of pRONO$_2$ in the SE US, Xu et al. (2015a) used the *RoR* concept. They estimated lower (2.2) and upper (4.4) limits for *RoR* (or R_{pRONO2} = 0.1-0.2 for their corresponding R_{NH4NO3}) from literature reports of SOA formed from isoprene+NO$_3$ radicals (Bruns et al., 2010) and β-pinene+NO$_3$ radicals (Fry et al., 2009; Bruns et al., 2010; Boyd et al., 2015), respectively. The rationale for their approach is that, for their region of study, those two BVOC may represent major contributions to the mixture of pRONO$_2$, and that the literature suggests there may be some source/composition dependence of R_{pRONO2}. For the same region, Chen et al. (2020) used bounds of R_{pRONO2} (0.1-0.2), based on similar logic, however not derived from a *RoR* calculation (however equivalent to a *RoR* of 1.7-3.3). In a study of pRONO$_2$ and SOA formation from Alberta oil sands extraction emissions from ground and aircraft measurements, Lee et al. (2019) used the same bounds of R_{pRONO2} (0.1-0.2), also not derived from a *RoR* calculation and citing Xu et al (2015a) and Farmer et al. (2010) (equivalent to a *RoR* of 1.4-2.9 and 1.5-3.0 for the two datasets). The same methods as Xu et al. (2015a) were used (applying the same range of *RoR*), for measurements conducted in Houston, TX (Dai et al., 2019) and the North China Plain (Xu et al., 2021). However Xu et al. (2021) adjusted the R_{NH4NO3} to match the highest NO_2^+/NO^+ ratios observed, since it was substantially higher than the calibration R_{NH4NO3} (assuming for those periods, nitrate was purely NH_4NO_3). Thus, those five studies report their concentrations and inorganic/organic nitrate split accordingly, and report lower and upper bounds; however, Lee et al., (2019) largely focused on results for the upper limit pRONO$_2$ concentrations for the scientific analysis (with equivalent *RoRs*: 1.4/1.5). Zhou et al. (2016), Zhu et al. (2016), and Yu et al. (2019) applied the *RoR* concept, citing a range of 2–4 from the literature, and thus reported estimated lower/upper limit averages for contribution of pRONO$_2$ to pNO$_3$ in New York City (summer, 67%/95%), a background site in China (spring, 15/22%), and an urban site in China (during spring, 13%/21%; summer, 41%/64%; autumn, 16%/25%), respectively. Similarly Zhu et al., (2021) applied the *RoR* concept, citing a range of 1.4–4.0 from the literature reporting upper(12%)/lower(7.8%) bounds for contribution of pRONO$_2$ to pNO$_3$ at a rural site in the North China Plains during summer. Kostenidou et al. (2015), on the other hand, estimated the R_{pRONO2} as the minimum $R_{ambient}$ observed in ambient data during the campaigns, resulting in effective *RoR*s of 5.6 and 12 for the two campaigns investigated. The same method is used by Reyes-Villegas et al. (2018) (using 46/30, and resulting in an effective *RoR* of 5) and Florou et al. (2017) (resulting in high effective *RoR*s of 14 and 15 for the two campaigns investigated). Other field studies have followed the methods of Fry et al. (2013) (but using a few different fixed values for the *RoR*) using HR data (Ayres et al., 2015; Fisher et al., 2016;

Lee et al., 2016; Palm et al., 2017; de Sá et al., 2018, 2019; Nault et al., 2018; Chen et al., 2021) or UMR data (Fry et al., 2018; Schulz et al., 2018).

3 Survey of NO_x^+ ratios for particle-phase nitrates

Given the numerous applications of NO_x^+ ratios to separate $pRONO_2$ and NH_4NO_3 in AMS measurements, yet many variations in methods and the numerical values used within each method, we have conducted a systematic survey of literature values and trends of NO_x^+ ratios for different nitrates. Such data compilation is aimed at evaluating the evidence that supports using a fixed *RoR* to estimate R_{pRONO2} from the calibration R_{NH4NO3} and to investigate the variability in R_{pRONO2} produced from different sources. Figure 1 shows a compilation of *RoR* values for $pRONO_2$ derived for chamber-generated SOA, isolated compounds (from chamber SOA or standards), and ambient measurements (using instrument comparisons or PMF separation). Figure 1 also shows the *RoR* for the same data as a histogram and average, as well as the correlations of the $pRONO_2$ vs NH_4NO_3 (inverse) NO_x^+ ratios. Details of the values used to compute the ratios and uncertainties, data sources, and any additional calculations for the information included in Fig. 1, are provided in Table S1.

The correlation between the R_{pRONO2} and R_{NH4NO3} is fairly strong (R^2=0.54), considering the variety of data sources and substantial measurement uncertainties. It provides strong evidence that, to first order, the *RoR* method is consistent and supported by various methods, species/mixtures, instruments and operating conditions. The slopes of the linear regression constrained to a zero intercept using an ODR fit (2.66±0.11; assuming both variables contribute comparable uncertainty) is equivalent to an overall *RoR* and is similar to the average of the individual *RoR* datapoints (mean±standard error: 2.75±0.11). Highlighted in the scatterplot in Fig. 1 are a couple of pairs of datapoints that are averages from several experiments conducted in our laboratory with two different AMS during two different years, with substantially different measured calibration R_{NH4NO3} while sampling the same chamber SOA (see S1.2). The trends in those points are similar to the overall trend and provide an example of the validity of the *RoR* method when only differences in instrument / operating conditions are present. Fig. S2 shows a complementary histogram to that in Fig. 1 for the R_{pRONO2}, without normalizing to R_{NH4NO3}. Compared to the normalized values shown in Fig. 1 (i.e., *RoR*s), a factor of two larger relative variability is apparent, with a relative standard deviation of 49% compared to 25%. Also of note is that the average value is 0.21±0.10, twice as high as used in several literature studies. Finally, Fig. S3 shows a complementary plot to the scatter plot in Fig. 1, with the inverse NO_x^+ ratios and axes swapped, which emphasizes different data and outliers, and yields similar but slightly higher (<10%), *RoR* slopes and the same degree of correlation. While the representation in Fig. S3 uses the inverse NO_x^+ ratio of that used throughout this manuscript, it places the R_{NH4NO3} on the x-axis, and thus a non-ODR fit may be appropriate under the assumption that most uncertainty is contributed by the $pRONO_2$ ratios. The ODR and non-ODR fits (2.83±0.12, 2.66±0.12, respectively) bracket the simple average value (2.75).

The compilation shown in Fig. 1 allows for consideration of dependencies of the *RoR* on species/mixtures or methods. Generally, the *RoR*s cluster around 1.5–4 for most studies. The variability within duplicated VOC-oxidant pairs (e.g., β-pinene+NO_3 SOA), similar compound classes (e.g., monoterpenes, isoprene, aromatics, long-chain alkanes or alkenes), or measurement methods (SOA mixtures, isolated compounds, ambient measurements) is similar to the variability between such groupings. Therefore, given the data currently available, there does not appear to be any strong evidence to support any general chemical-dependence of the $pRONO_2$ *RoR*. While such a dependence may in fact exist, evaluation likely would require comparison of several organic nitrate molecules and/or mixtures systematically with the same instrumentation, operation conditions, and analysis methods, together with duplication by different instruments.

Therefore, for applications and further evaluation described in this manuscript, we use the average and variability of the *RoR* determined from data highlighted in Fig. 1: 2.75 (mean) and standard deviation (±0.70, 25%) or standard error (±0.11, 4.0%). The 25th/50th/75th percentiles are 2.12, 2.73, 3.12 (interquartile range / median +14%/-22%). Given the approximate symmetry for the limited statistics available, we treat the variability and uncertainty of the *RoR* as approximately a normal distribution. The standard deviation should be considered an upper limit of the uncertainty of the applicable *RoR* and corresponds to the assumption that the variability in reported values is primarily attributable to true differences in ratios for different types of $pRONO_2$. The lack of clear differences among different sources suggests that some of the variability may instead be instrument/operator related, and that the std. error may be a more relevant characterization of the uncertainty. Complex mixtures of $pRONO_2$ in the atmosphere would likely represent an ensemble of those ratios, and thus result in values closer to the average. In fact, for the limited (7) examples of ambient-derived *RoR*s, the average is similar and the variability somewhat smaller (2.99±0.51, ±17%) compared to the overall survey data. The standard error of the overall survey can be considered a measure of the uncertainty under the assumption that the *RoR* is invariable with source/type and the R_{NH4NO3} for an instrument is a perfect predictor of R_{pRONO2}. A separate manuscript will include further discussions on the *RoR* uncertainty and applications to estimation of the overall nitrate apportionment and concentrations uncertainties.

We recommend the use of the average *RoR* value computed here for future separations of $pRONO_2$ and NH_4NO_3 in ambient aerosol with AMS until there is additional information available to support a different or more complex formulation. On the other hand, where additional constraints on the expected $pRONO_2$ ratio response may be available, a more specific value may be applied. For example, Takeuchi and Ng (2019) measured *RoR*s during dry chamber experiments for different SOA types where only $pRONO_2$ nitrate was generated, and then used those system-specific *RoR*s to separate $pRONO_2$ and NH_4NO_3 during wet experiments where substantial NH_4NO_3 was also formed. We note that in a recent study, Xu et al., (2021) inferred a substantial variability in R_{pRONO2} for ambient measurements on diurnal timescales and with varying pollution levels; however, that relied on comparison of the NO_x^+ ratio method to a newly-proposed method using thermal denuder profiles, which they acknowledge has several potentially large uncertainties or biases that were not quantified.

It is important to emphasize that under strong influence of particle-phase *nitrites* or semi/non-refractory nitrates (e.g., $NaNO_3$, $Ca(NO_3)_2$), quantitative separation of nitrate types may be hindered or simply not feasible (Schroder et al., 2018). As a few studies have reported, nitrites and mineral nitrates produce substantially lower NO_2^+/NO^+ ratios (thus higher *RoR*) in the AMS. For example, *RoR*s of ~10–60 for $NaNO_3$ (Alfarra, 2004; Bruns et al., 2010; Hu et al., 2017b), 17 for $Ca(NO_3)_2$ (Alfarra, 2004), 3.9 for $Mg(NO_3)_2$ (Alfarra, 2004), 9.7 for KNO_3 (Drewnick et al., 2015), and ~300 for $NaNO_2$ (Alfarra, 2004) have been previously reported. We report additional measurements from our laboratory for $NaNO_3$, KNO_3, and KNO_2 showing similarly high values. Table S2 provides additional details and Fig. S4 shows a graphical representation and comparison to $pRONO_2$ for literature reports and our new data. Consequently, even if the expected ratios of other compounds was accurately known,

https://doi.org/10.5194/amt-2021-263
Preprint. Discussion started: 6 September 2021

apportioning the different nitrates or nitrites using a formulation like Eq. 1 would be under-constrained, as there would be more unknowns than equations. Therefore, care must be taken to screen for measurements that may be substantially influenced by such interferences (e.g., seasalt, dust). Additionally, during a recent aircraft campaign focused on biomass burning, we conducted regular calibrations with 4-nitrocatechol, a nitroaromatic (Pagonis et al., 2021). The *RoR* was relatively similar to $pRONO_2$ at 3.35 ± 0.81 (1σ, standard deviation) (Table S2, Figs. S4, S5).

4 Evaluation of calibration R_{NH4NO3} and *RoR* using ambient data

A survey of NO_x^+ ratios for multiple field studies is explored here in order to assess the framework of using measured calibration R_{NH4NO3} and a *RoR* to apportion NH_4NO_3 and $pRONO_2$ concentrations. See Sect. S1.1 and Table S3 for details and a summary of all field campaigns for which data is used within this manuscript. Figure 2 shows frequency distributions of $R_{ambient}$ for ambient aerosol from two aircraft-based remote continental (SEAC^{4}RS, DC3) and two ground-based forest campaigns (SOAS, BEACHON-RoMBAS). The data is shown as the calibration R_{NH4NO3} divided by $R_{ambient}$, so that all data is comparable. For all campaigns, the large majority of the data fall between the R_{NH4NO3} (1 on Fig. 2, indicating all NH_4NO_3) and the *RoR*-determined R_{pRONO2} (2.75 on Fig. 2, indicating all $pRONO_2$). The small fraction of data outside that range may be due to a combination of instrument noise, drifts in the instrument NO_x^+ ratio response not captured by periodic calibrations, and/or the inability of the fixed *RoR* to perfectly capture the R_{pRONO2} response. However, these results show that under a large range of chemical conditions and instrument R_{NH4NO3} (spanning a factor of 2.4 for these campaign averages), the data are generally consistent with the *RoR* apportionment model. Figure S6 shows the same distributions as Fig. 2, except as simple frequency distributions, rather than weighted by mass concentration as in Fig. 2. The broadening and shift to the right for simple frequency distributions (compared to those weighted by mass concentration), reflect the typical trend that $pRONO_2$ tends to constitute higher fractions of pNO_3 when pNO_3 is lower. Distributions are similar for other campaigns (not shown in Figs. 2, S6), as can be inferred from Figs. 5 and S9, which are discussed below.

The effects of estimating R_{pRONO2} using time-variant vs constant R_{NH4NO3} is explored in Fig. S7. For the SEAC^{4}RS campaign, the flight-to-flight calibration R_{NH4NO3} were highly variable due to some instrument instability (range: 0.40–1.49, mean±stdev: 0.80±0.31; Figs. S8, S9e), compared to the very stable ratios measured during the other campaigns (see Fig. 2 caption). Therefore, two histograms are shown overlaid in Fig. S7, one normalized to flight-dependent calibration R_{NH4NO3} and the other normalized to the campaign-averaged R_{NH4NO3}. For the standard frequency distributions (Fig. S7a), there is substantial narrowing when using the flight-dependent ratios, indicating that application of the time-variant ratios provides better constraints on the instrument response to the NH_4NO_3 — $pRONO_2$ mixture. Conversely, normalizing to arbitrary R_{NH4NO3} would be expected to broaden the distribution. The most prominent differences for the mass concentration-weighted distributions (Fig. S7b) are largely due to data with high NH_4NO_3 concentrations where the measured $R_{ambient}$ were beyond the campaign-averaged R_{NH4NO3} (resulting in a substantial fraction of the distribution <1). There is also subtle broadening toward the $pRONO_2$ portion of the distribution. These comparisons support that using the variable calibration R_{NH4NO3} better represents ambient NH_4NO_3 ratios (left side of plots) and tying R_{pRONO2} to R_{NH4NO3} (i.e. using the *RoR* method, rather than fixed R_{pRONO2}) better represents $pRONO_2$ ratios (right side of plot).

https://doi.org/10.5194/amt-2021-263
Preprint. Discussion started: 6 September 2021

Additional support for the practice of using the measured calibration R_{NH4NO3} and anchoring the R_{pRONO2} to those calibrations with a fixed *RoR* can be drawn from the R_{ambient} vs pNO_3 plots shown in Fig. S9a/b. Five studies shown in those figures had relatively constant (within each campaign), but differing (among campaigns; factor of 3.2 range), calibration R_{NH4NO3} (SOAR, MILAGRO, SOAS, BEACHON-RoMBAS, KORUS-AQ; 0.47, 0.84, 0.44, 0.30, 0.97, respectively). However, as pNO_3 increases for the urban-influenced studies (SOAR, MILAGRO, SOAS, KORUS-AQ) or for the oxidation flow reactor (OFR) measurements during SOAS (Fig. S8c), R_{ambient} tends to approximately converge at the calibration R_{NH4NO3}. This suggests that NH_4NO_3 in mixed ambient aerosol is well-represented by offline-calibrations for a range of conditions and instruments. Additionally, the corresponding average ratios at the lowest pNO_3 concentration (same 5 studies in Fig. S9a/b) converge at a similar range of ratios (0.26, 0.52, 0.15, 0.10, 0.40, respectively; range of 4.0). If assuming that the low-pNO_3 observed R_{ambient} approximate pure $pRONO_2$ ratios, a relatively narrower range is computed for an inferred *RoR* (1.6–3.0, factor of 1.9; 2.36±0.63), which is also similar to expected *RoR*s (albeit low possibly due to urban ground studies never sampling pure $pRONO_2$).

Further evidence supporting the use of calibration R_{NH4NO3} and the *RoR* using ambient data is presented in Sect. S2 using campaign datasets where the calibration R_{NH4NO3} showed large variability (DAURE, $SEAC^4RS$ campaigns). Exploration of the NO_x^+ ratios vs pNO_3 relationships showed similar relationships to those discussed above for campaigns where R_{NH4NO3} was constant or changed little, but with the curves shifting with the measured R_{NH4NO3}. Similar values of *RoR* to those presented in the literature survey in Sect. 3 were also inferred from the $SEAC^4RS$ dataset. Finally, both datasets were used to evaluate biases when using a fixed value of R_{pRONO2} vs estimation of a dynamic value using the *RoR* method. Additional evidence from ambient measurements supporting use of calibration R_{NH4NO3} and the *RoR* is presented in Sect. 5.2 where applications of PMF separation are discussed.

5 Demonstrations of *RoR* apportionment and comparisons to other measurements/methods

5.1 $pRONO_2$ - NH_4NO_3 separation compared to total (gas+particle) $RONO_2$ (Tot-$RONO_2$)

Figure 3 shows time series of AMS $pRONO_2$ and NH_4NO_3 concentrations for a $SEAC^4RS$ flight (RF16) in the Southeast US. The nitrate components were apportioned according to Eqs. 2/3 and a *RoR* of 2.75. "Total $RONO_2$" (gas+particle; hereafter Tot-$RONO_2$) concentrations, as measured by thermal dissociation - laser induced fluorescence (TD-LIF) (Day et al., 2002; Perring et al., 2009), are shown for comparison. A wide range of sources were sampled including (and indicated by) biogenic (monoterpenes and/or isoprene and photochemical products such as IEPOX, MVK), anthropogenic (e.g., NO_x, NO_y, aromatics), biomass burning (e.g., acetonitrile and f_{60}, an AMS tracer (Cubison et al., 2011)), likely agricultural, as well as mixtures of these sources or relatively clean free tropospheric air. Flight tracks are shown in Fig. S10 and approximate periods and corresponding source influences are listed in the caption. A large and variable range of pNO_3 was observed (<10 ng m^{-3} or <4 ppt up to ~5 µg m^{-3} or ~1800 ppt) and ranging from $pRONO_2$-dominated to NH_4NO_3-dominated. The $pRONO_2$ and Tot-$RONO_2$ tracked remarkably closely. NH_4NO_3 concentrations exhibited more plume-like behavior, rapidly increasing and decreasing, often while both $pRONO_2$ and Tot-$RONO_2$ remained relatively constant or in some cases showed moderate and similar increases. Overall, $pRONO_2$ was correlated with Tot-$RONO_2$ (R^2=0.49 for all data, R^2=0.69 for data with f_{pRONO2} >0.3) with a regression slope of 0.029 (0.033), indicating that on average ~3% of $RONO_2$ was in the particle phase (Fig. 3, bottom left). NH_4NO_3 showed little overall relationship

https://doi.org/10.5194/amt-2021-263
Preprint. Discussion started: 6 September 2021

to Tot-RONO$_2$ beyond the trend that at higher altitudes, well above the boundary layer and outside of plumes, both concentrations tended to be low (Fig. 3, top and bottom right). Note that the reference R_{NH4NO3} in the particle nitrate apportionment here (in Eqs. 2/3) was 0.70 which was based on the measured calibration R_{NH4NO3} and PMF results (see Sect. 5.2 just below). Measured R_{NH4NO3} during calibrations in days bracketing this flight were 0.96 (2 days before) and 0.71 (1 day after). PMF results support a value of 0.70 (see Sect. 5.2.2), which was used here since it was similar to the nearest calibrations and provides an additional constraint on the otherwise variable calibration R_{NH4NO3} characteristic of this campaign (see Sect. 4). Using a higher R_{NH4NO3} increases the pRONO$_2$ vs Tot-RONO$_2$ slope in Fig 4 (bottom left) and can improve the correlation a bit (mainly by moving the low values at low f_{pRONO2} toward the regression line).

Taken together, these observations indicate that the AMS nitrate apportionment method effectively separated pRONO$_3$ and NH$_4$NO$_3$ over a large range of concentrations, relative contributions, and source influences. However, it is clear that there are limitations when the f_{pRONO2} is very low (see Sect. 5.2). It would not be surprising if the pRONO$_2$ and Tot-RONO$_2$ showed large variability in relative ratios for different sources and locations, since: 1) pRONO$_2$ is only a small subset of Tot-RONO$_2$ and 2) changes in chemical composition and ambient conditions (e.g., OA concentration, temperature) could have large impacts on gas-particle partitioning. However, in this case those effects do not appear to be large factors (or fortuitously cancel out), which in part may be due to relatively similar temperatures and OA concentrations combined with regionally consistent biogenic chemical sources of RONO$_2$ compounds. Regardless of the exact reasons for the relatively invariant partitioning, it provides an excellent test case, since it would be very unlikely that the strong temporal/spatial correlation would be observed if there were major artifacts in either or both the AMS and TD-LIF methods.

There were no measurements of inorganic nitrate onboard the aircraft with fast enough time resolution to compare with the rapidly changing NH$_4$NO$_3$ concentrations calculated from the AMS. Therefore, as a rough indicator of possible changes in the NH$_4$ related to NH$_4$NO$_3$, "Excess NH$_4$" was calculated as the AMS-measured NH$_4$ - 1.2 x SO$_4$ (as molar concentrations). A molar ratio of 1.2 was roughly consistent with the observed ratio when no indications of NH$_4$NO$_3$ were present (NH$_4$=1.2 x SO$_4$) and substantial concentrations of SO$_4$ were present, as shown in Fig. S11. That ratio represents a mixture of (NH$_4$)$_2$SO$_4$ and ammonium bisulfate or an ammonium balance (NH_{4_Bal}) of ~0.7 (NH_{4_Bal} = molar ratio of NH$_4$/(NO$_3$+2SO$_4$)). During periods of elevated NH$_4$NO$_3$ concentrations, the measured NH$_4$NO$_3$ tracked the estimated "Excess NH$_4$" very closely with roughly half the concentration (Fig. S11). As suggested by some negative "Excess NH$_4$" values and the factor of two between NH$_4$NO$_3$ and "Excess NH$_4$", the assumption of constant NH$_4$/SO$_4$ ratios based on composition in the absence of NH$_4$NO$_3$ is not always valid (and not surprising) and clearly a more sophisticated thermodynamic model would be required to accurately predict NH$_4$NO$_3$ concentrations. Nonetheless, the similar features suggest the assignment of NH$_4$NO$_3$ is consistent with variations in the other AMS-measured inorganic compounds. The factor of two suggests that ~half of the "Excess NH$_4$" was associated with sulfate and half with nitrate. During this flight, with the exception of the large biomass burning plume, the elevated NH$_4$NO$_3$ concentrations were observed when the aircraft flew at altitudes of ~2000–4000 m and never during the low-altitude (~300–400 m) legs (S20 bottom left/middle). This effect may have been due to the substantially cooler temperatures (0–15°C vs 25–30°C) at those altitudes, favoring partitioning to the particle-phase, since there did not appear to be any clear relationship between NH$_4$NO$_3$ and gas-phase HNO$_3$ (Fig. S11, bottom right). Increases in available NH_3 gas (not measured) could also be a factor (and consistent with both more sulfate- and nitrate-associated ammonium).

Another example for a different flight (RF18) during the SEAC^{4}RS aircraft campaign is shown in Fig. S12, and was also selected due to large relative and absolute variability in calculated $pRONO_2$ and NH_4NO_3 concentrations and diverse source types sampled (see Fig. S13 for flight track and description). Similarly, the $pRONO_2$ and Tot-$RONO_2$ track remarkably well during periods when NH_4NO_3 concentrations are low or elevated and variable, and there is little correlation between NH_4NO_3 and Tot-$RONO_2$. Overall, $pRONO_2$ was correlated with Tot-$RONO_2$ (R^2=0.51 for all data, R^2=0.71 for data with f_{pRONO2}>0.3) with a regression slope of 0.050 (0.068), indicating that on average ~5–7% of $RONO_2$ was in the particle phase (Fig. S12a, bottom left). The measured NH_4NO_3 tracked the estimated "Excess NH_4" reasonably well and showing similar sharp features (and roughly half the concentration; Fig. S12b, top). In contrast to RF16 discussed above, for RF18 most of the elevated NH_4NO_3 was observed in the warm boundary layer and often coincident with elevated $pRONO_2$ (Fig. S12a,b).

5.2 Positive Matrix Factorization separation of AMS nitrate

5.2.1 Prior studies using PMF for $pRONO_2$ separation

For the vast majority of analyses of AMS data using PMF, only traditional OA ions have been included in the input data matrices. Ions typically associated with nitrate, sulfate, ammonium, and chloride have generally been excluded, with the mindset that they are already separated as unambiguous inorganic species using the standard AMS analyses. However, since organic molecules (e.g., organic nitrates, organosulfates, reduced organic nitrogen) can in fact produce some of the same ions as those inorganic species, inclusion with the OA ions in PMF analysis may allow for separation of inorganic and organic components, as well help identify associations with more well-established source factors.

A few studies have reported results for using PMF of ambient AMS spectra including both the OA and NO_x^+ signals to quantify $pRONO_2$ (and sometimes NH_4NO_3), with mixed results (Sun et al., 2012; Hao et al., 2014; Xu et al., 2015a, 2021; Zhang et al., 2016; Kortelainen et al., 2017; Yu et al., 2019; Zhu et al., 2021). Additionally, a couple other studies have reported results where NO_x^+ ions or calculated $pRONO_2$ (using the NO_x^+ ratio method) are included in PMF analysis, while not explicitly apportioning the inorganic-organic nitrate directly with the PMF results in the laboratory (Tiitta et al., 2016) and field (Kim et al., 2018; Reyes-Villegas et al., 2018). Details and discussions of those studies are presented in Sect. S3 and key results are summarized in Table S4, as related to the PMF analyses.

5.2.2 New results for PMF separation of $pRONO_2$ and comparison to ***RoR*** method

We conducted PMF on the combined OA and NO_x^+ ion time series for the same two flights from the SEAC^{4}RS campaign (as discussed above in Sect. 5.1; RF16, RF18) to test PMF separation of nitrates and the information it can provide, explore strategies, and compare to the *RoR* method. Details and an extended discussion of that analysis is documented in Sect. S4 and key results are summarized in Table S4 alongside previous published analyses. A brief summary is provided here.

As discussed in Sect. 5.1, those two flights included sampling of a wide range of source types and concentrations. PMF was conducted initially on 1-s data; however, although robust overall factors were separated, results suggested that the S/N was not adequate to apportion the NO_x^+ ions to secondary factors at ratios that reflected $pRONO_2$ ratios. Therefore, all analyses discussed here are from 1-min

https://doi.org/10.5194/amt-2021-263
Preprint. Discussion started: 6 September 2021

measurements (which were more effective). Several strategies were used to explore the separation of OA, nitrate, and the NO_x^+ ratios (in separate and combined factors), including: number of factors, varying FPEAK, upweighting and downweighting NO_x^+ ions, bootstrapping, seeding, constraining NO_x^+ ratios, and removing large biomass burning plumes. For both flights, five factors were robustly separated: NH_4NO_3, BBOA (biomass burning OA), IEPOX-SOA (IEPOX-derived SOA), LO-OOA (less-oxidized oxygenated OA), and MO-OOA (more-oxidized OOA) (Figs. S14–S28). See the Glossary and Sects. S3/S4 for more details on factor types. Generally, the best separations with the most information were for FPEAK at or near 0, using standard NO_x^+ ion S/N (no downweighting/upweighting), not constraining NO_x^+ ratios, not removing any plume data, and using bootstrapping to extract averages and assess uncertainty/robustness.

The NH_4NO_3 factors and the BBOA factors had very similar NO_x^+ ratios that were consistent with calibration R_{NH4NO3}, with little variability across the 100 bootstrapping runs (Figs. S17, S25). While the apportionment of nitrate between the NH_4NO_3 and BBOA factors was very consistent across bootstrapping runs, changes in FPEAK had large effects on that relative apportionment as well as the amount of OA ions in the NH_4NO_3 factor spectrum. For the OOA/SOA factors (IEPOX-SOA, LO-OOA, and MO-OOA) the NO_x^+ ratios for LO-OOA and the combination of all three factors were consistent with expected $pRONO_2$ NO_x^+ ratios using the *RoR* (Figs. S17, S25). Across bootstrapping runs, there was modest variability for those ratios (Figs. S17, S25), including some solutions where the LO-OOA had only NO^+ (but not for the combined OAA/SOA factor). The averages and standard deviations of the NO_x^+ ratios for the combined OOA/SOA factor are included in the survey of $pRONO_2$ *RoR*s (Fig. 1, Table S1). For calculation of NH_4NO_3 and $pRONO_2$ concentrations, the nitrate contributions from the NH_4NO_3 and BBOA factors were summed as were the three OOA/SOA factors, respectively. The majority of the $pRONO_2$ was contributed by the LO-OOA factor, followed by MO-OOA and then IEPOX-SOA (Figs. S18, S27). The variability in the factor spectra NO_x^+ ratios and nitrate concentration apportionment across bootstrapping tended to follow the same trend (higher variability for factors with lower $pRONO_2$ contribution; e.g., Figs. S17, S18a, S25, S27). Additionally, substantial trends were observed between factor spectra NO_x^+ ratios and the amount of nitrate apportioned to that factor for some OOA/SOA factors. Bootstrapping and exploration of FPEAK was useful to investigate those dependencies.

Comparisons of NH_4NO_3 and $pRONO_2$ concentrations using the *RoR* and PMF methods are shown for each flight in Figs. 4 and S12a as time series and scatter plots. For both flights there is very good agreement (near unity slope, 0.99–1.04, and R^2>0.99) between methods for NH_4NO_3, certainly in part due to the dominance of NH_4NO_3 during higher concentrations periods. There is reasonable agreement for $pRONO_2$ (slopes of 0.86–1.50, R^2 of 0.51–0.65 depending of the flight and fitting method; and improved to slopes of 1.04–1.42, R^2 of 0.68–0.84 for f_{pRONO2}>0.3) but with notable differences. $pRONO_2$ concentrations tended to be noisier for the *RoR* method compared to the PMF method when nitrate was dominated by NH_4NO_3 or when pNO_3 was very low. This may be due to the additional S/N and constraints that the inclusion of the other OA ions provide, as well as the sensitivity (for both precision and accuracy) of apportionment for the *RoR* method when ratios approach the R_{NH4NO3} limit. On the other hand, the PMF method may dampen some real variability due to the fact that the factor spectra are fixed and cannot chemically evolve in the PMF model. In order to assess the true accuracy of either method, an independent and reliable determination of $pRONO_2$ would be required. Finally, the comparison between the PMF-determined $pRONO_2$ and the TD-LIF Tot-$RONO_2$ showed substantially-improved correlation (compared to using the *RoR* method) for one of the two flights (Fig. 4 vs 3).

5.2.3 Summary of PMF method for nitrate separation

The results from our investigation of PMF and analyses described in the literature summarized above highlight some general aspects, as well as some potential advantages and disadvantages of using PMF to apportion nitrate between organic and inorganic. One major potential advantage is that with PMF, the nitrates can be immediately associated with different source factors. On the other hand, the NO_x^+ ratio method can be used first and then correlations of nitrates with OA-only factors can be explored and even apportioned. PMF may provide additional resolving power and S/N by inclusion of associated OA ions, potentially more precisely separating nitrate concentrations, especially when either $pRONO_2$ or NH_4NO_3 dominate the nitrate. Also, prior knowledge of the NO_x^+ ratio for NH_4NO_3 (or $pRONO_2$) may not be necessary if the ratios are robustly resolved with PMF. Additionally, the NO_x^+ ratios resolved for PMF factors is a product for exploring ratios for ambient aerosol response, and validating application of offline calibration R_{NH4NO3} and *RoR*s derived largely from laboratory studies. PMF may also be useful in separating other species that produce NO_x^+ ions (e.g. nitrites, nitro-organics, mineral nitrates), from just NH_4NO_3 and $pRONO_2$, when they are present and have a unique NO_x^+ ratio.

Some potential drawbacks or cautionary aspects are as follows. Since the PMF model requires fixed profile spectra, this means that nitrate-to-OA ratios are fixed for each factor. Therefore, if this ratio is in fact substantially variable over the period/space of analysis, for example driven by processes such as $pRONO_2$ hydrolysis or gas-particle partitioning, substantial biases or uncertainties in nitrate apportionment can be introduced. While consideration of additional factors could help mitigate such effects, PMF is not designed to concisely separate profiles that are a continuum. Sometimes factors with clear NH_4NO_3 or $pRONO_2$ NO_x^+ ratio signatures are not resolved. We suspect that datasets where neither type of nitrate is dominant for some periods may be more susceptible to that issue; however, those issues may sometimes be resolvable with more extensive investigation with available PMF exploration tools (e.g., seeding, bootstrapping, FPEAK, constraining a NH_4NO_3 factor from offline calibrations). Otherwise, apportioning nitrate using results with profile spectra that do not have clear nitrate signatures may introduce large uncertainties which are difficult to estimate. Variable NO_x^+ ratios due to instrument drifts or changes (e.g., vaporizer bias voltage drifts or tuning) may lead to uncertainty in nitrate apportionment since PMF computes fixed factor spectra. In practice, for using the NO_x^+ ratio method this is not problematic, as long as regular offline NH_4NO_3 calibrations were performed. For PMF, separating the dataset into periods where the NO_x^+ ratio was stable/constant and performing PMF separately for each period is one option to mitigate instrument drift issues; however, this can be very laborious if the dataset requires separate analysis of multiple periods. Another option may be to apply the “rolling method” recently made available with ME-2/SoFi, where a sub-window is moved across the PMF input along the time coordinate, allowing factor profiles to vary with each sub-window shift (Canonaco et al., 2021). Theoretically, offline calibration ratios of NH_4NO_3 may not be necessary for such application, although they would be preferable to have for validation.

A few other notable trends and observations are as follows. PMF-resolved $pRONO_2$ often tends to have the largest contribution from (and association with) LO-OOA/SV-OOA, followed by MO-OOA/LV-OOA, especially for biogenically-influenced locations. That is consistent with $pRONO_2$

forming in fresh SOA (i.e. LO-OOA/SV-OOA) and being partly lost as the OA ages and/or MO-OOA/LV-OOA consisting of a mix of aged OA, some of which was not associated with $pRONO_2$. Nitrate associated with BBOA tends to be dominantly NH_4NO_3; however, primary and secondary $pRONO_2$ associated with BBOA emission has been reported. When NH_4NO_3 factors are resolved, they tend to contain substantial contributions (~20–80%) of OA (non-NO_x^+) ions. Generally, those non-NO_x^+ contributions seem to be higher for strongly biogenically-influenced measurements and less so during cooler wintertime periods when NH_4NO_3 comprises a larger fraction of nitrates. Our experience through exploration of various approaches (e.g., upweighting the NO_x^+ ions, increasingly positive FPEAK, increasing number of factors) suggests that efforts at "cleaning" the NH_4NO_3 factor tends to be ineffective and/or lead to degradation of the overall PMF solutions. Since the OA contained in the NH_4NO_3 tends to not be a large overall fraction of the OA, this does not appear to be a major issue. Finally, evidence suggests that inclusion of NO_x^+ ions in PMF does not tend to have much influence on overall OA-dominated factors (factor spectra nor concentration time series), which is not surprising given that their overall contribution to the S/N among the many OA ions is fairly small. Consequently, there does not appear to be any drawbacks or complications associated with also including NO_x^+ ions when running PMF on AMS data.

Overall, PMF appears to be a useful tool for apportioning nitrates and investigating their associations with sources. The case for quantitative apportionment of nitrate with PMF is strongly bolstered when the NO_x^+ ratios resolved for both the NH_4NO_3 factor and separate or combined $pRONO_2$-associated factors are similar to NH_4NO_3 calibration and expected $pRONO_2$ NO_x^+ ratios. When those criteria are not met, using the NO_x^+ ratio method may be better, as it is likely less prone to such biases or ambiguities, and uncertainties can be better defined.

5.3 Comparison of $pRONO_2$ quantification with AMS and other instruments in the lab and field

Several studies have reported quantitative comparisons of $pRONO_2$ concentrations, as measured by AMS vs other instrumental methods (alternate AMS-based methods, FTIR, TD-(LIF/CRDS/CAPS), and FIGAERO-CIMS). Section S5 provides details and discussions and Table S5 presents a summary of key aspects of those comparisons. Overall, those comparisons show good agreement in most cases (1:1 within known uncertainties) and substantial differences in a few cases (factors up to 2–4). In some of the cases where substantial differences were observed, possible explanations were discussed and sometimes explored. There do not appear to be any consistent reasons for the differences. In some of the field comparisons and all of the laboratory experiments, the nitrate sampled was dominated by (or exclusively) $pRONO_2$, and thus largely serve as a test of $pRONO_2$ quantification (general calibration/quantification factors, RIE, collection efficiency, etc.). Consequently, taken together the evidence available does not support use of an RIE for $pRONO_2$ quantification with AMS that is significantly different from that measured for (and regularly calibrated with) NH_4NO_3. In order to narrow the uncertainties in $pRONO_2$ quantification (in the field and laboratory), controlled laboratory-based intercomparisons of total and speciated organic nitrates using AMS and other methods are needed.

6 Physical basis for NO_x^+ ratios observed for nitrate types and variability among instruments

As Farmer et al. (2010) points out, it is probable that a large fraction of $RONO_2$ molecules thermally decompose to RO and NO_2 at the AMS vaporizer after which NO_2 gas is ionized. For example, the TD-LIF technique (and CRDS/CAPS equivalent methods) rely on quantitative thermal dissociation of

$RONO_2$ to NO_2 in the gas phase, which occurs at ~350 °C in ~50 ms at near ambient pressures (Day et al., 2002). The timescale of evaporation/decomposition/ionization/detection for the AMS are on order tens of µs (Drewnick et al., 2015; Jimenez et al., 2016); however, at 600°C the dissociation rate coefficient for $pRONO_2$ is ~4 orders of magnitude larger (compared to 350 °C). That said, it is not clear what the pressures or temperatures of the gases are in the evaporation plume. Nevertheless, Farmer et al. note that thermal decomposition of $pRONO_2$ to NO_2 in the AMS would be consistent with the higher NO^+/NO_2^+ ratios observed for $pRONO_2$ than NH_4NO_3. Their reasoning is that reported ratios of NO_2 gas ionization (3.0) are substantially higher than those reported for HNO_3 (0.5) gas as well as their measurements of particle-phase NH_4NO_3. Using the simplest assumption that only NO_2 (from $RONO_2$ thermal decomposition) and HNO_3 (from NH_4NO_3 evaporation) are ionized would yield a *RoR* of 6, which is double that observed. Moreover, fixed values would be expected for $pRONO_2$ and NH_4NO_3 rather than the observed range of ~4. Clearly, the behavior is more complicated than this simple model. Given that mass discrimination (ion transmission or detector efficiency differences) for the *m/z* range of the NO^+ and NO_2^+ ions is expected to be minor for the AMS (Hu et al., 2017b), the values and variability in NO_x^+ ratios likely originate in the vaporizer and/or ionizer region. As discussed in Hu et al. (2017b), the values and range of NO_x^+ ratios observed for NH_4NO_3 (combined with other observations) are consistent with EI from a combination of HNO_3, NO_2, and NO gases that are formed through thermal decomposition. They show the greatly-enhanced importance of such neutral gas-phase decomposition for measurements where a "capture vaporizer" is substituted for the standard AMS vaporizer. The capture vaporizer has a different geometry (optimized for limiting particle bounce) that results in longer gas-phase residence time near the hot vaporizer surfaces. Consequently, an order of magnitude lower NO_2^+/NO^+ ratio is observed for NH_4NO_3 (0.04–0.07), likely due to a shift in ionization toward primarily NO gas. Similar thermal decomposition processes would be expected for $RONO_2$. However, thermal decomposition to RO and NO_2 may occur much faster and always to near completion, given the thermal instability of the O-NO_2 bond and near absence of $C_xH_yO_zN^+$ fragments in AMS $pRONO_2$ spectra (Farmer et al., 2010). Hu et al. (2017a) report a large reduction in the NO_2^+/NO^+ ratios for $pRONO_2$ when using the capture vaporizer compared to the standard vaporizer (with a $pRONO_2$ ratio ten times lower than for NH_4NO_3 with the capture vaporizer).

As shown in Drewnick et al. (2015) and Jimenez et al. (2016), single-particle detection timescales for ions when sampling NH_4NO_3 show a range of a factor of two (and ~25 µs differences), primarily with NO^+ being longer than NO_2^+ and NH_x^+ ions. Those observations are interpreted as evidence for additional processes occurring at longer timescales than flash vaporization at the nominal temperature such as vaporization at lower effective temperatures, slower vaporization or thermal decomposition, and adsorption/desorption from ionizer surfaces. They also showed that the signal-particle detection timescales were insensitive to vaporizer temperatures above 300°C. On the other hand, Hu et al. (2017b) showed a small dependence of the R_{NH4NO3} on vaporizer temperature decreasing by 25% from 200°C to 800°C, consistent with more thermal decomposition to NO_2 and NO gases. Other studies have reported no dependence of NO_x^+ ratios on vaporizer temperature (~200–600°C) for $pRONO_2$-containing chamber SOA (Fry et al., 2009) or ambient (mixed nitrate) aerosol (Docherty et al., 2015). Overall, these observations point toward the timescales of interaction, and effects of spatial distribution of competing processes, playing a more important role in affecting observed ion ratios, rather than vaporizer temperature. In part, this relative insensitivity to vaporizer temperature may be because the physical process of particle vaporization occurs at lower temperature than the nominal vaporizer temperature due to evaporative cooling (Saleh et al., 2017). Another observation that Hu et al. reported for using the capture vaporizer was that the vaporization timescales (based on UMR PToF distributions) for NO^+ was much longer than for NO_2^+ for NH_4NO_3, but the reverse for $pRONO_2$. Such apparent spatiotemporal differences in thermal decomposition and ionization could potentially be used as another method for differentiating nitrates. However, low S/N of NO_2^+, differences in sizes and broader distributions for ambient aerosol nitrates, and the possibility that some of the differences Hu et al. observed were from $CH_2O_x^+$, may seriously limit such approach and would require further evaluation (using HR-PToF).

A few other evaluations of R_{NH4NO3}, described in Hu et al., (2017b) (using the standard vaporizer), showed dependencies of NO_x^+ ratios of only <20% including varying the location on which particles impact the vaporizer (by horizontally translating the aerodynamic lens position) and varying the vaporizer bias voltage over ranges expected for typical AMS operation. On the other hand, varying the vaporizer bias voltage over a wider range, such as slightly beyond the settings where the aerosol signal peaks and where the gaseous "airbeam" signal peaks, can result in nearly a factor of two shift in the R_{NH4NO3} (Fig. S29). This behavior reflects the ability of the vaporizer bias voltage tuning to preferentially sample ions produced in different regions of the ionizer. It has also been shown for the signals of other ions, such as CO_2^+ (Jayne et al., 2015). While proper tuning of the AMS vaporizer bias voltage typically aims at optimizing the aerosol signal, that may not always be performed by AMS operators and likely in some cases the airbeam signal may be optimized instead (which can be different than the particle signal peak as in Fig. S29, although not always). Therefore, variability in this tuning parameter may explain a substantial fraction of the range in NH_4NO_3 (and possibly $pRONO_2$) NO_x^+ ratios shown in Fig. 1. Another effect that appears to be able to substantially alter the NO_x^+ ratios is related to exposure to high concentrations of OA for extended periods, possibly coating the vaporizer (and is possibly related to the "Pieber Effect" where nitrate aerosol produces CO_2^+ signal from interactions at the vaporizer surface), and will be discussed in a future publication. Taking all the evidence available at present, the range in NO_x^+ ratios for NH_4NO_3 and $pRONO_2$ among instruments, settings, and operating conditions appears to be driven by changes in the amount of chemical decomposition and the overlap of those products with the ionizing electron beam. This aspect highlights the importance of periodic measurement of the NO_x^+ ratios with a standard (i.e., NH_4NO_3), especially after making significant instrument changes, when quantifying $pRONO_2$ and NH_4NO_3 with the AMS.

7 Multisite survey of inorganic/organic fractionation

An overview of the inorganic vs organic nitrate apportionment for all of the campaigns discussed in this manuscript is shown in Fig. 5. The campaigns span: late-winter to summer across the northern hemisphere and wet/dry seasons near the equator; from ground level to the upper troposphere; and urban to remote locations. Overall, the f_{pRONO2} shows an inverse relationship with the pNO_3, approaching 100% at low pNO_3, primarily at rural/remote locations. At high pNO_3 and strongly urban-influenced locations, the nitrate is dominantly NH_4NO_3. However, urban and urban-influenced locations can often exceed 50% contributions from $pRONO_2$, when pNO_3 is lower (<1–2 μg m^{-3}). At the urban ground sites (MILAGRO, SOAR), the modulation of the variability in pNO_3 tended to be driven by large increases in NH_4NO_3 from photochemical production of HNO_3 during morning to early afternoon, followed by evaporation at higher temperatures during afternoon driving concentrations to minima that were generally sustained through nighttime (Aiken et al., 2009; Docherty et al., 2011). At the rural/remote sites, nitrate is nearly always dominated by $pRONO_2$ and with low concentrations. At the mid-latitude sites (BEACHON, SOAS), a large contribution to the variability in concentrations was attributed to nighttime production of $pRONO_2$ from BVOC (Fry et al., 2013; Xu et al., 2015b). For the Amazon studies, substantial variability was observed on sub-day and synoptic timescales, especially during the lower-concentration wet season measurements, with episodic elevated inorganic contributions (de Sá et al., 2018, 2019). Thus, variability may have largely been driven by transport changes and large-scale regional processes; however, the factors controlling particle-phase nitrate for those studies have not been thoroughly explored. For DAURE, an urban-downwind site with high pNO_3, consistent diurnal patterns were not observed, and pNO_3 variability was likely dominantly driven by variability in transport (Minguillon´ et al., 2011; Zhang and Jimenez, 2021).

The aircraft campaigns span the entire range of the urban and rural/remote sites combined, since they include urban and biomass burning sampling, as well as rural/remote and free tropospheric sampling. However, there are notable differences among them and compared to ground-based studies. A major difference is the shift toward lower f_{pRONO2} or pNO_3 in the intermediate ranges by factors of ~2 or ~10, respectively. The large divergence as pNO_3 decreases from ~2 to ~0.2 µg m^{-3} coincides with the range where the aircraft measurements show NH_{4_Bal} transitions from balanced ($NH_{4\_Bal} \sim 1$) to a modest deficit in ammonium ($NH_{4\_Bal} \sim 0.75–0.9$) (see Fig. S30). Lower NH_{4_Bal} can be indicative of more acidic aerosol (Nault et al., 2021; Schueneman et al., 2021), making particle-phase NH_4NO_3 less thermodynamically stable. In comparison, the NH_{4_Bal} for the ground-based urban-influenced studies, (SOAR, MILAGRO, DAURE) were consistently near unity (Aiken et al., 2009; Docherty et al., 2011; this work for DAURE, not shown). However, such effects alone would result in higher f_{pRONO2} in the aircraft studies, not lower as observed, due to sulfate not balanced by ammonium and acidity making ammonium nitrate thermodynamically unstable. Therefore, other factors must be at play, such as very different sources being sampled, lower temperatures and higher RH for the aircraft measurements (making NH_4NO_3 more thermodynamically stable; see Sect. 5.1, Fig. S11), dilution shifting the curves, or higher acidity shortening the lifetime of $pRONO_2$ (such as accelerating hydrolysis). At the lower range of pNO_3 (<0.2 µg m^{-3}) the f_{pRONO2} is substantially different following the order KORUS < DC3 < SEAC^{4}RS. Considering again the NH_{4_Bal} (Fig. S30), for SEAC^{4}RS the aerosol inorganics are much less balanced by ammonium ($NH_{4\_Bal} \sim 0.08–0.75$) compared to DC3 ($NH_{4\_Bal} \sim 0.5–0.8$) and KORUS ($NH_{4\_Bal} \sim 0.5–0.9$) at the lower pNO_3 range, suggesting a possible role of acidity and NH_3 availability. On the other hand, it does not appear that acidity plays a dominant role in favoring the high f_{pRONO2} at the rural/remote ground-based studies, as BEACHON tended to be fully balanced ($NH_{4\_Bal} \geq 0.9$) while SOAS was not ($NH_{4\_Bal} \sim 0.5–0.7$) (Fry et al., 2013; Hu et al., 2016).

Many different chemical and physicochemical processes interplay to control the concentrations and relative proportions of NH_4NO_3 and $pRONO_2$ in the atmosphere. Fig. 6 shows a schematic of those key processes. The differentiation can be viewed as effectively beginning with the branching of the radical-radical reaction of NO_x with OH vs RO_2 or VOCs ($NO+RO_2$, $NO_2+RC(O)O_2$, $NO_3+RC=CR'$) to produce gas-phase HNO_3 vs $RONO_2$. The relative amount of these pathways can vary widely, in large part controlled by relative amounts of NO_x concentrations compared to VOC reactivity; the $RONO_2$ formation pathway can become dominant below modest NO_x concentrations, particularly at biogenically-influenced rural sites (e.g., Browne and Cohen, 2012; Romer, 2018). However, the partitioning of HNO_3 and $RONO_2$ into the particle phase can depend on numerous factors such as NH_3 availability, RH, temperature, particle acidity, $RONO_2$ volatility, or OA concentrations. Subsequent chemical, photochemical, evaporation, and deposition losses of gas and particle components will also exert controls on concentrations and lifetimes. In large part, the general trend shown in Fig. 5, over more than three orders of magnitude pNO_3, may be driven by the ability of HNO_3 formation in the presence of sufficient NH_3 at increasing pollutions levels (i.e., NO_x) to overwhelm more modest $pRONO_2$ formation, combined with the high volatility of NH_4NO_3 prone to evaporation upon dilution. In contrast, at rural and remote locations, the formation of $RONO_2$ becomes more favorable, producing $pRONO_2$ of which a substantial portion is not prone to rapid chemical or evaporative loss, thus dominating widespread background nitrate composition. However, this is a very simplified picture of the complex processes at play and more detailed investigations combining corresponding measurements with modeling to better understand the dominant processes controlling the trends shown in Fig. 5 are needed. In a recent study of eleven aircraft campaigns from throughout the globe, Nault et al. (2021) showed overall trends of decreasing pH and NH_{4_Bal} with remoteness (as indicated by decreasing total inorganic PM_1), which was not well-represented in many current models. While there may be some connections between that phenomena and the one shown in Fig. 5 (e.g., via acidity and NH_3 availability), inorganic PM_1 concentration is more closely related to remoteness than pNO_3, as it is often dominated by sulfate, which is less chemically reactive and less volatile than $pRONO_2$ and NH_4NO_3, and its formation is less coupled to VOC conditions. For a ground-based study in a Chinese megacity during fall, a strong trend of increasing inorganic fraction of pNO_3 with increasing calculated aerosol pH (pH=1.5-3.5) was observed, which was attributed to numerous coincident factors during pollution episodes favoring NH_4NO_3 precursor availability and gas-to-particle partitioning (Chen et al., 2021).

We note that the data included in Fig. 5 are generally weighted toward warmer periods or regions. Xu et al. (2015a) reported wintertime (within Nov-Feb) measurements of organic and inorganic nitrate at two urban and one rural site in the southeast US. Campaign averages of pNO_3 ranged 0.8–1.4 $\mu g\ m^{-3}$ (with 1σ variability of ±90–100%) and average f_{pRONO2} was 0–30% across the sites and the apportionment methods considered. pNO_3 and inorganic nitrate showed strong diurnal cycles, peaking mid-morning with minima mid-to-late afternoon. Nitrate apportionment vs pNO_3 was not reported, so it is unclear if similar trends to those in Fig. 5 were present (e.g., if f_{pRONO2} increased during afternoon pNO_3 minima). However, on average all three campaigns fell in the chemical coordinate space of the urban-influenced studies shown in Fig. 5. The fact that the rural site was similar to the urban sites may be due to the cooler winter temperature (and higher RH) as well as reduced biogenic influences, compared to warm rural studies shown in Fig. 5. A few other studies have shown AMS data as supplementary material, that suggest similar relationships to those in Fig. 5 for individual studies. Those include plots of NO^+ vs NO_2^+ ions which appear to have higher ratios of NO^+/NO_2^+ at lower signals (Docherty et al., 2015; Zhou et al., 2016) or decreasing NO_2^+/NO^+ ratios with decreasing pNO_3 (Kiendler-Scharr et al., 2016). Additionally, a recent analysis of three datasets in the North China Plain (urban summer/winter, rural winter), showed a strong decreasing trend in f_{pRONO2} vs PM_1 during the urban summer measurements and weak trends for the wintertime measurements (and lower overall f_{pRONO2}) (Xu et al., 2021). Those observations are generally consistent with the trends with pNO_3 during summer and with seasonality discussed above.

8 Conclusions

We have explored the viability of using the NO_x^+ ion ratios produced in the AMS spectrum from nitrates to separate and quantify NH_4NO_3 and $pRONO_2$ concentrations in ambient aerosols. The use of NH_4NO_3 calibration NO_x^+ ratios and an inferred NO_x^+ ratio for $pRONO_2$ that tracks the NH_4NO_3 ratio ("Ratio-of-Ratios") is investigated and tested. An extensive range of data and approaches are utilized for this investigation including: a diverse collection of ambient field datasets, chamber studies, oxidation flow reactors, pure compounds, comparisons to AMS PMF methods and other $pRONO_2$ or related measurements, and a compilation of a broad literature survey.

It is shown that the method is robust and effective under typical ambient sampling conditions. Methods and practical considerations for calculating concentrations are described. The Ratio-of-Ratios NO_x^+ ratio method produced similar results to conducting PMF on the expanded mass spectra series (including both OA and NO_x^+ ions) to apportion nitrates. While using the PMF method may have advantages of improved signal-to-noise and can provide connections between $pRONO_2$ and OA sources, it is much more labor-intensive and can lead to substantial biases if not explored and applied carefully.

A broad survey of nitrate apportionment shows a pervasive relationship of increasing (decreasing) $pRONO_2$ relative contributions to nitrate with decreasing (increasing) total nitrate concentrations. Those trends generally follow from urban-influenced to rural/remote regions. However, there are some clear differences in those trends between different sampling regions and conditions.

Previous studies reporting nitrate quantification using AMS NO_x^+ ratios (or PMF using NO_x^+ ions) have employed a range different approaches and assumptions, based on generally limited information. In some instances, likely substantial biases were present and rarely has the accuracy of the results been considered. This investigation will help provide a more consistent, accurate and transparent approach to quantification and exploration of bulk particle-phase nitrates in the atmosphere with AMS (and related instrumentation). Comparisons of this method to other instrumentation capable of quantifying bulk or speciated particle-phase organic nitrates, in the laboratory and field, should be an ongoing focus to help better constrain uncertainties, identify biases, and improve this method (and others).

Acknowledgements

This research was supported by NASA grants 80NSSC18K0630 and 80NSSC19K0124, as well as US NSF grant NSF AGS-1822664, NOAA grant NA18OAR4310113. BBP acknowledges support from a US EPA STAR Graduate Fellowship (FP-91761701-0). This work has not been formally reviewed by the US EPA. The views expressed are solely those of the authors, and the US EPA does not endorse any products or commercial services mentioned in this work. We thank John Crounse and Paul Wennberg (Caltech) for use of nitric acid measurements in Figs. S11, S12.

Figures

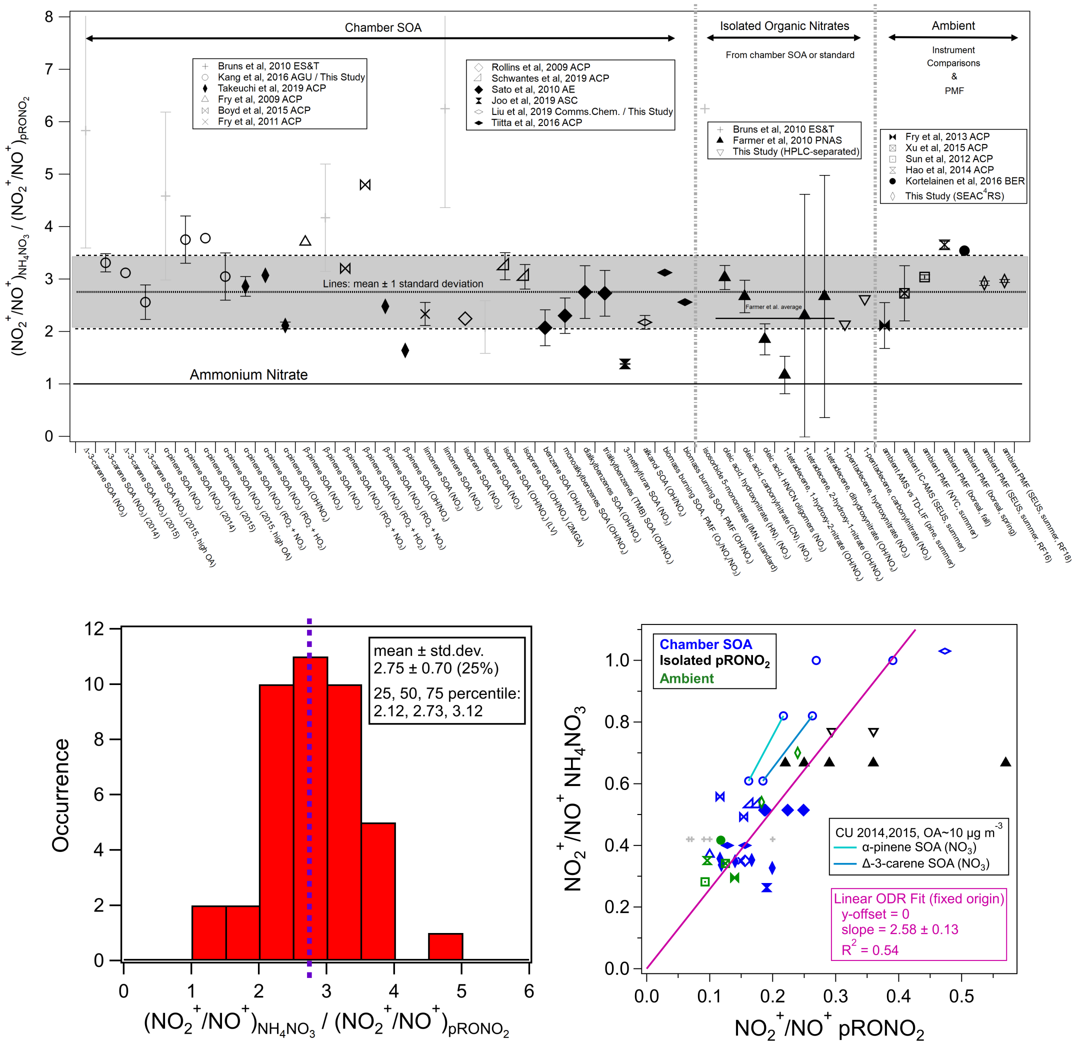

Figure 1. (Top) Survey of "Ratio-of-Ratios" (*RoR*) computed from NO_2^+/NO^+ ratios reported for chamber studies, pure organic nitrates, and field observations (using instrument comparisons or PMF separation). The mean (2.75) and standard deviation (±0.70, ±25%) are also shown (standard error for n=41: ±0.11, ±4.0%). The light grey shading ("+" markers) indicates data that were not used in the average here, nor in the fits below (see Table S1 for rationale). Details of the values used to compute the ratios and uncertainties, data source, and any additional calculations for the information included in Figure 1 are provided in Table S1. (bottom left) Histogram and statistics of *RoR*. (bottom right) scatter plot of R_{NH4NO3} vs. R_{pRONO2}. Linear least-squares lines are shown with orthogonal distance regression ODR fit (with intercept constrained through the origin since offsets from unconstrained fits were not significant and for consistency with apportionment equation). The data connected by cyan and green lines are averages from experiments conducted in our lab with two different AMSs (with substantially different calibration

R_{NH4NO3}) while sampling the same SOA particles produced using the same two precursors mixtures. See Fig. S3 for the equivalent scatter plot, instead using NO^+/NO_2^+ ratios and swapping the axes (R_{pRONO2} vs R_{NH4NO3}).

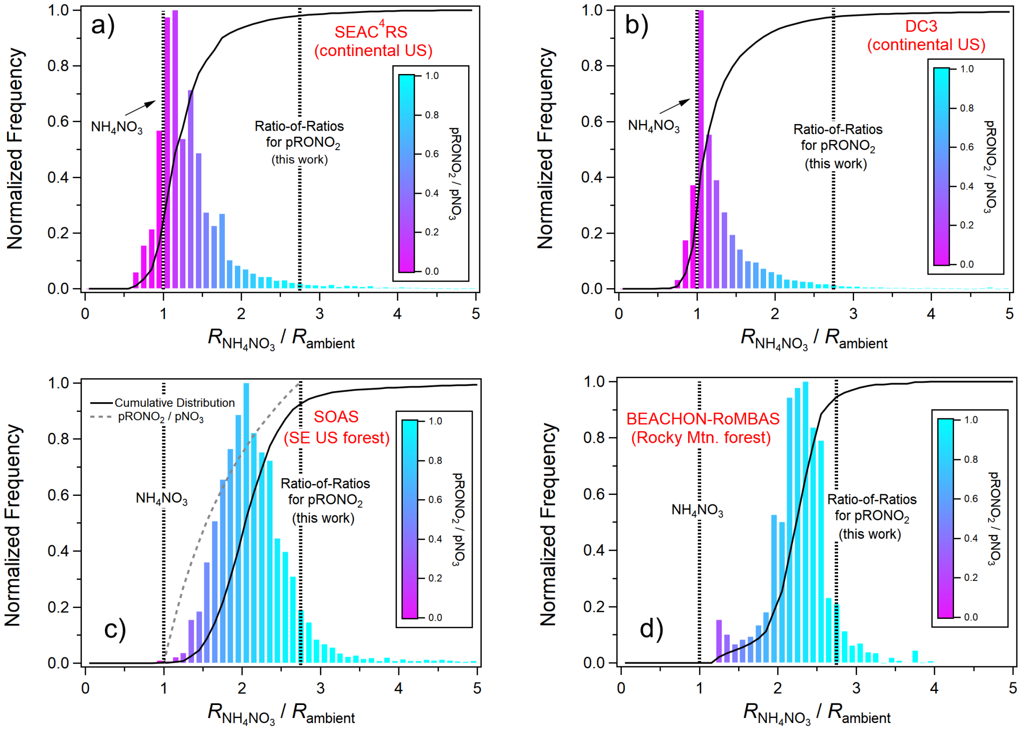

Figure 2. Histograms of ambient NO_x^+ ratios for aircraft and ground-based campaigns. The data is shown as the calibration R_{NH4NO3} divided by $R_{ambient}$, so that all data are on the same reference coordinates. The histograms are weighted by pNO_3 concentration. Cumulative distributions are shown in all plots and an additional curve only on the SOAS panel shows the f_{pRONO2} ($pRONO_2/pNO_3$) for these coordinates (would be identical on all panels). The data used were 1-minute averages and screened for pNO_3 detection limits for the aircraft campaigns (SEAC⁴RS, DC3), and 1-hour averages for the ground-based campaign (SOAS, BEACHON-RoMBAS). Measured R_{NH4NO3} for these studies were as follows: SEAC⁴RS (range 0.40–1.49, mean and stdev. 0.80±0.31); DC3 (0.71±0.04); SOAS (0.44±0.02); BEACHON-RoMBAS: (0.295±0.005). See Fig. S6 for equivalent plots where distributions are *not* weighted by mass concentration).

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Preprint. Discussion started: 6 September 2021

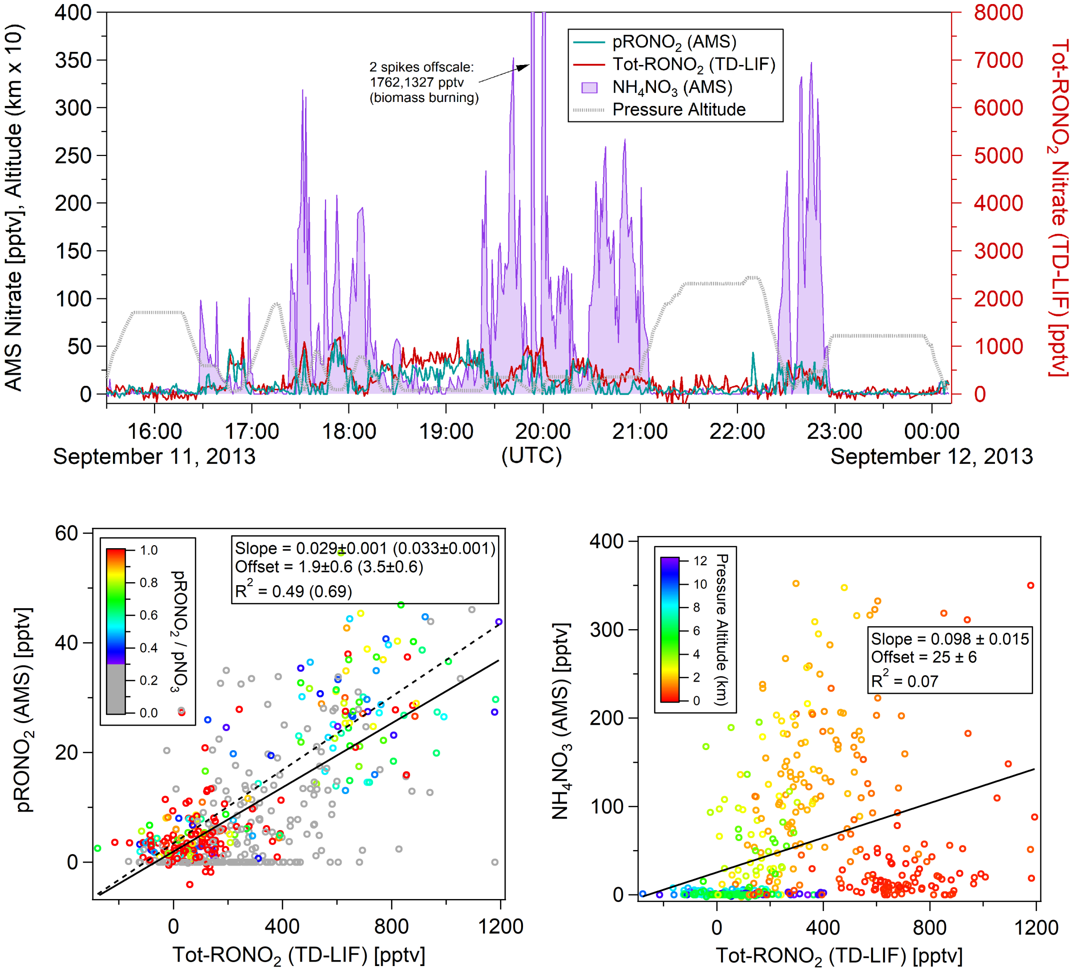

Figure 3. Comparisons of AMS pRONO$_2$ and NH$_4$NO$_3$ with TD-LIF total (gas+particles) organic nitrate (Tot-RONO$_2$) during a SEAC^{4}RS flight (RF16) in the Southeast US (1-min averages). The time series (top) and scatterplots of pRONO$_2$ (bottom left) or NH$_4$NO$_3$ (bottom right) vs Tot-RONO$_2$ are shown. Measured calibration R_{NH4NO3} (consistent with PMF results in Sect. 5.2.2), a *RoR* of 2.75, and Eqs. 2/3 were used to apportion the AMS nitrate. Linear least-squares lines are orthogonal distance regression (ODR). For the pRONO$_2$ vs Tot-RONO$_2$ plot (bottom left), an additional line (dotted) and fits (parentheses) are shown for data including only when f_{pRONO2} (pRONO$_2$/pNO$_3$) is greater than 0.3 (and datapoints with f_{pRONO2}<0.3 are greyed). Figure S10 shows the flight track and timing of different source types sampled.

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Preprint. Discussion started: 6 September 2021

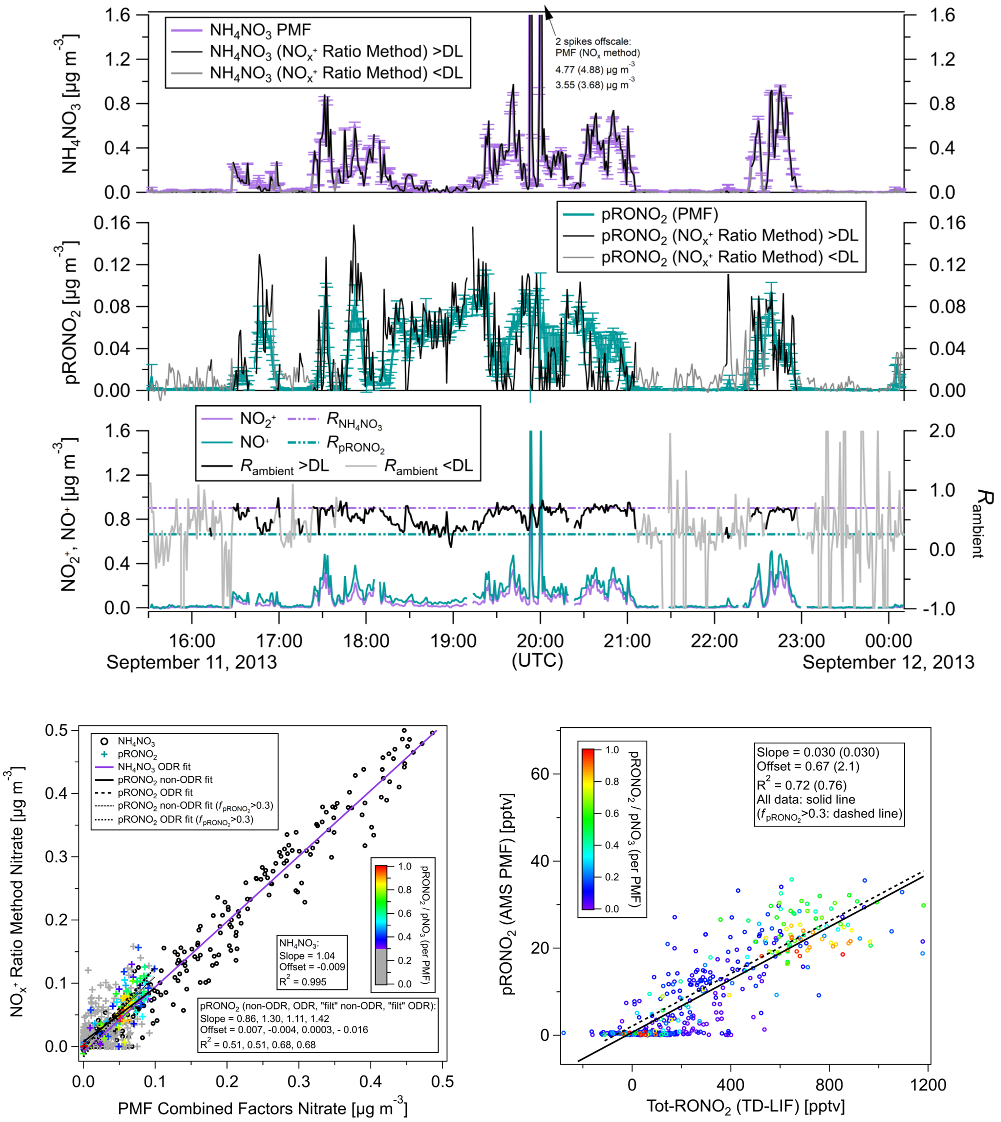

Figure 4. Comparison of NO_x^+ ratio vs PMF methods for calculation of NH_4NO_3 and pRONO$_2$ as time series (top three panels) and as scatter plots (bottom left) for same flight shown in Fig. 3. Concentration time series calculated using the *RoR* method (as well as the measured NO_x^+ signals and ratios) are shown for all data as well as only when above the $R_{ambient}$ detection limit (DL). Bottom right: PMF pRONO$_2$ vs TD-LIF Tot-RONO$_2$ (equivalent to Fig. 3 bottom left, which instead shows pRONO$_2$ from *RoR* method). pRONO$_2$ in scatterplots are colored by the f_{pRONO2} (pRONO$_2$/pNO$_3$) as computed using the PMF method. Regression line fits/slopes/offsets and correlation coefficients are shown using different fitting methods and criterion as indicated in legends (where "filt" indicates fits where data is limited to f_{pRONO2}>0.3). All PMF-derived concentrations are averages (and standard deviations) of 100 bootstrapping runs (similar results using seeding runs shown in Fig. S21).

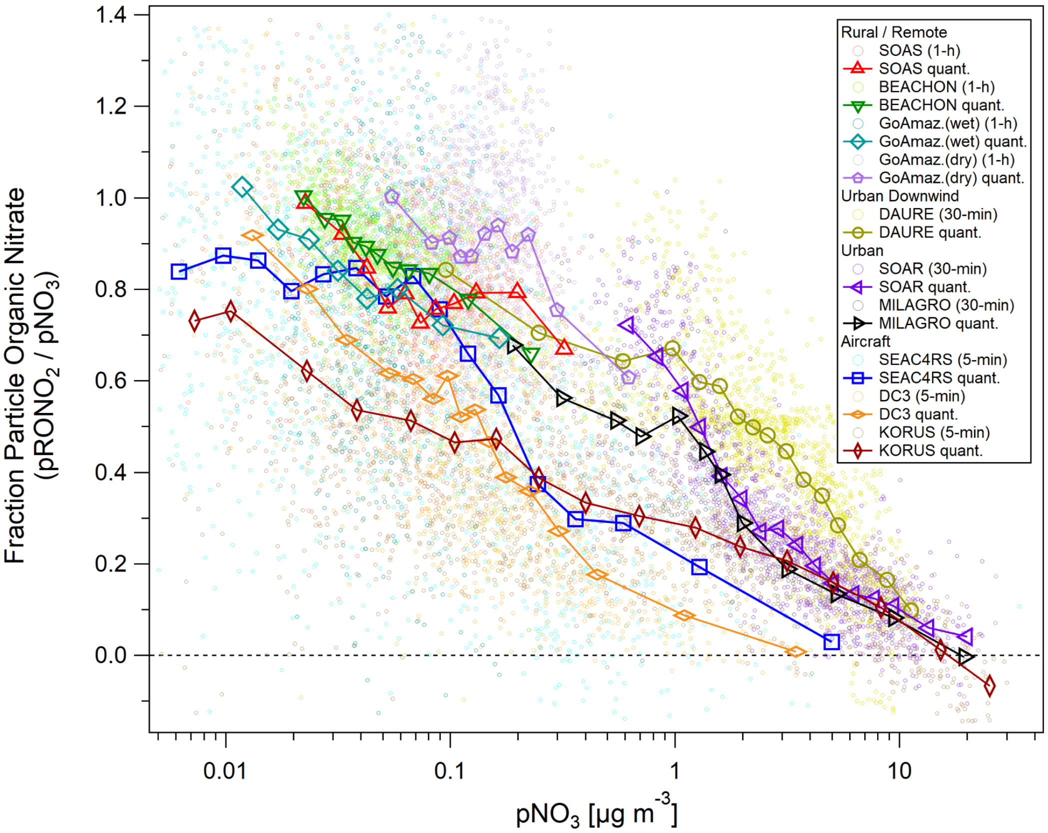

Figure 5. Fraction of total non-refractory submicron nitrate that is organic (f_{pRONO2}) vs. total nitrate concentration (pNO_3) for several ground and aircraft campaigns. Campaigns span: late-winter to summer across the northern hemisphere and wet/dry seasons near the equator; from ground level to the upper troposphere; and urban to remote locations. NO_x^+ ion signals were first averaged and then data was conservatively screened for detection limits (S/N>1-3) using both NO_x^+ ions (small circles). Quantile averages (means, 7–15 bins) are also shown for each campaign. Additionally, for all campaigns, one additional average was calculated and included with the quantile averages for the highest 1% of pNO_3 in order to extend the pNO_3 by a factor of ~1.5–3 (undersampled chemical regime, but with high S/N). The average of the lowest 1% of pNO_3 for the MILAGRO campaign is also included.

https://doi.org/10.5194/amt-2021-263
Preprint. Discussion started: 6 September 2021

855

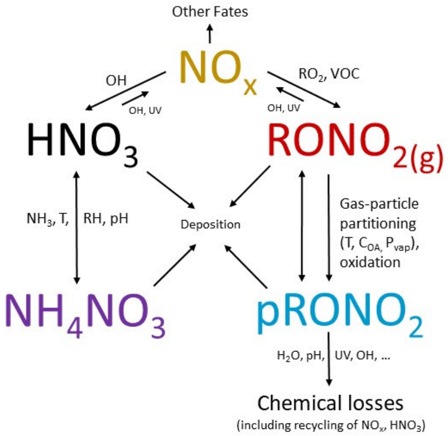

Figure 6. Schematic of key processes controlling particle-phase NH_4NO_3 and $pRONO_2$.

860

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