Aerosol models from the AERONET data base: application to surface reflectance validation

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Abstract. Aerosols play a critical role in radiative transfer within the atmosphere, and they have a significant impact on climate change. In this paper, we propose and implement a framework for developing an aerosol model using their microphysical properties. Such microphysical properties as the size-distribution, the complex refractive index, the percentage of sphericity, are derived from the global AERosol RObotic NETwork (AERONET). These measurements, however, are typically retrieved when almucantar measurement procedures are performed (i.e., early mornings and late afternoons with clear sky), and might not have a temporal correspondence to a satellite overpass time, so a valid validation of satellite-derived products can't be carried out. To address this problem of temporal inconsistency of satellite and ground-based measurements, we developed an approach to retrieve these microphysical properties (and the corresponding aerosol model) using the optical thickness at 440 nm, τ_{440}, and the *Ångström* coefficient, $\alpha_{440\text{-}870}$. Such aerosol models were developed for 851 AERONET sites within the last 28 years. Obtained results suggest that empirically microphysical properties can be retrieved with uncertainties of up to 23%. Exception is the imaginary part of the refractive index *ni*, for which the derived uncertainties reach up to 38%. These specific parametric models of aerosol can be used for the studies when retrieval of microphysical properties is required, as well as validation of satellite-derived products over land. Specifically, we demonstrate the usefulness of the aerosol models to validate surface reflectance records over land derived from optical remote sensing sensors. We then quantify the propagation of uncertainties of the proposed aerosol model on the surface reflectance that is used as a reference from radiative transfer simulations. Results indicate that individual aerosol microphysical property can impact uncertainties of surface reflectance retrievals between 3.5×10^{-5} to $1\text{x}10^{-3}$ in reflectance units. The overall impact of microphysical properties combined yields an overall uncertainty in surface reflectance <0.004 (in reflectance units). That corresponds, for example, to 1 to 3% of the retrieved surface reflectance in the red spectral band (620-670 nm) by the Moderate Resolution Imaging Spectroradiometer (MODIS) instrument. These uncertainties values are well below the specification (0.005+0.05ρ, ρ is the retrieved surface reflectance) used for the MODIS atmospheric correction.

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1 Introduction

Aerosols play a key role in the atmosphere as an important climate forcing in climate assessment (IPCC, 2018; IPCC, 2019) and their better characterization would improve our knowledge of their properties for a better assessment of their impacts (Dubovik and King, 2000; Andreaa et al., 2002; Dubovik et al., 2002b; Roger et al., 2009; Omar et al, 2008; Nousiainen, 2011; Dubovik et al., 2011; Ginoux et al., 2012; Boucher et al. 2013; Calvo et al., 2013; Lenoble et al., 2013; Fuzzi et al. 2015; Derimian et al., 2016; Klimont et al., 2017; Torres et al., 2017; Bond et al., 2018; Contini et al., 2018; De Sá et al. 2019; Li et al., 2019, Mallet et al., 2020…).

In general, the use of the specific aerosol model depends on the temporal and spatial scales. Approximate models are generally adequate for long-term studies, when intra-annual or intra-seasonal variability of aerosols is of less importance; however, studies that require capturing aerosol variability in space and time would require a more specific and precise characterization. The AERosol RObotic NETwork (AERONET) network (Holben et al.,1998) was created in the early nineties and continues operation today. Over the last 30 years, this network has provided information on the aerosol characteristics for approximately 1000 globally distributed sites. AERONET estimates several microphysical properties of aerosols (i.e., the size-distribution, the complex refractive index, and the percentage of sphericity). These parameters are derived during the almucantar measurement procedures which are typically carried out early morning and late afternoon under clear sky conditions. As a result, it is usually not possible to have these aerosol microphysical properties when an Earth observation satellite passes over an AERONET site. To address this problem, we propose a method to retrieve microphysical properties using a parametric model with two variables: the optical thickness at 440 nm, τ_{440}, and the *Ångström* coefficient, $\alpha_{440\text{-}870}$. We selected these two parameters because they are widely accessible (e.g., from the AERONET Network which provides several measurements per clear sky hour, from the satellite itself or from climatology). We used 851 AERONET sites for which the data were in a sufficient quantity and representative. Thus, we can derive a dynamic aerosol model for each of these AERONET sites. These parametric models of aerosol can be used for the studies when retrieval of microphysical properties is required, as well as validation of satellite-derived products over land.

In the context of satellite products validation, the surface reflectance retrieval requires a good characterization of the aerosol properties, particularly for sensors with various and narrow spectral bands (Justice et al, 2013). Therefore, uncertainties in the aerosol models would impact uncertainties in the surface reflectance records derivation. By incorporating aerosol model into a radiative transfer model, one can generate reference surface reflectance, which can be used for validating satellite-derived surface reflectance. It is essential, in this case, that a careful validation be performed on a global and continuous basis, including a wide range of land and, consequently, reflectance conditions. One approach is the direct comparison with 'ground- truth' measurements, but this presents several challenges related to the scale and nature of the ground measurements and their representativeness at coarse and medium satellite pixel resolutions, since the global representativeness of the pixel may differ from the point measurements. Nevertheless, at a finer spatial resolution (pixels less than 30m), ground measurements may occur. Indeed, with a good protocol and good radiometry, direct ground truth measurements can be performed for validation

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(Helder and al., 2012; Czapla-Myers et al., 2015; Czapla-Myers et al., 2016; Badawi et al., 2019; Bouvet et al., 2019). There are also other approaches. For example, we use an indirect approach for the validation of satellite products from MODIS and VIIRS (Vermote et al., 2002; Vermote et al., 2014), for the NASA Harmonized Landsat-8/Sentinel-2 project (Vermote et al., 2016; Claverie et al., 2018), or for the CEOS ACIX Working Group for atmospheric correction intercomparison (Doxani et al., 2018). In the former, we compare a surface reflectance retrieved from satellite data to a surface reflectance reference determined from the Top of Atmosphere (TOA) reflectance corrected using the accurate radiative transfer 6SV code (Vermote et al., 1997; Kotchenova et al., 2006; Kotchenova et al., 2007; Kotchenova et al., 2008) and detailed measurements of the atmosphere. An intermediate step consists of validating the Aerosol Optical Thickness product derived from various sensors such as MODIS, MISR, OMI, POLDER and Landsat, which is further used as an input to the atmospheric correction process. Numerous studies have applied this validation approach (e.g., Martonchik et al., 1998; Remer et al., 2005; Herman et al., 2005; Masek et al, 2006; Keller et al., 2007; Martonchik et al., 2009; Dubovik et al., 2011; Levy et al, 2013; Vermote et al, 2016; Levy et al., 2018; Doxani et al., 2018). In the last part of this paper, we evaluate the uncertainties of our aerosol microphysical properties on the definition of the surface reflectance (to be used as reference) in the MODIS red band.

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2 Description of the aerosol model

2.1 Aerosol microphysical description

There are two ways to describe an optical aerosol model: using optical properties or using the microphysical properties. The optical properties (scattering and absorbing coefficients; phase matrix) are derived from the following microphysical properties: the size-distribution (which gives the diameter distribution of the aerosol population), the complex refractive index (which gives characteristics of the light scattered by the particle for the real part and the absorbing quality of the particle for the imaginary part), and the sphericity (which describes the aerosol shape and non-sphericity aspect) (Hansen and Travis, 1974; Van der Hulst, 1981; Lenoble, 1993; Liou, 2002; Mishchenko et al., 2002; Bohren and Huffman, 2010; Lenoble et al., 2013). Thus, to avoid losing information about the microphysical properties (i.e., the aerosol composition), we prefer to describe the aerosol model using its microphysical properties rather than its optical properties (knowing that it will give us possibility to compute the optical properties from the microphysical ones). The size distribution characterization may be variable in its chemical or optical description, i.e., mass and numbers, respectively. This results in a different shape and description of the size-distribution. For an optical approach, the Gaussian Distribution is widely used as the most appropriate model for the aerosols size-distribution (Whitby, 1978; Shettle and Fenn, 1979; amongst other subsequent studies). In order to design an optical aerosol size-distribution in its vertical description, a combination of a Gaussian's Law for each aerosol mode is suitable (the fine mode and the coarse mode identified hereafter by *f* and *c*), even if it can be much more complex at a small scale (Liou, 2002; Hsu et al., 2004; Roger et al., 2009; Dubovik and King, 2000; Dubovik et al., 2011; Lee et al., 2015). In this way, the particle volume size-distribution can be described by the derivative of the particle volume at a specific radius V(r) by the natural logarithm of the radius:

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As important inputs for atmospheric correction, we need the aerosol properties exactly when the satellite overpasses one of the AERONET validation sites. For the purpose of the surface reflectance product validation, we decided to create a dynamic aerosol model for each AERONET site when all data are sufficiently available and representative. This paper describes how we define and design these aerosol models for the 850 AERONET selected sites (here the optical model is defined by the microphysical properties). These models can then be used for the atmospheric correction validation, as we do here, or for other purposes when an aerosol model is needed (e.g. local studies or creation of an aerosol climatology).¶
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$$\frac{dV(r)}{dlnr} = \frac{C_{vf}}{\sqrt{2\pi}\sigma_f} exp\left[-\frac{(lnr-lnr_{vf})^2}{2\sigma_f^2}\right] + \frac{C_{vc}}{\sqrt{2\pi}\sigma_c} exp\left[-\frac{(lnr-lnr_{vc})^2}{2\sigma_c^2}\right], \quad (1)$$

where the six microphysical parameters that described this model are: C_{vf} (the particle volume concentration of the fine mode), C_{vc} (the particle volume concentration of the coarse mode), ~~r_{vf}~~ et ~~r_{vc}~~ (the particle median volume radius of the fine and coarse mode), ~~σ_f~~ and ~~σ_c~~ (the standard deviation of the Gaussian's law of the fine and coarse mode).

The phase function of aerosols is usually normalized (Lenoble, 1985), thus the size distribution doesn't need to be defined in an absolute manner. We then may define the relative volume concentration $\%C_{vf}$ and $\%C_{vc}$ (scaled between 0 and 1), rather than C_{vf} and C_{vc} (discussed latter in this paper). The complex refractive index of the aerosol, $n = n_r + i\, n_i$, is the second required microphysical parameter. The real part (n_r) describes the scattering properties of the aerosol, while the imaginary one (n_i) describes absorption properties. Both parts have to be known for a given wavelength. Finally, the percentage of sphericity $\%S_{ph}$ can be considered as well, to take into account the non-sphericity of aerosols (Mishchenko et al., 2000; Dubovik et al., 2002b; Herman et al., 2005), in contrast to a "spherical approach" (Mie, 1908). This non-sphericity mostly applies the coarse mode.

2.2 Description of the dataset

Aerosol microphysical properties data were extracted from the AERONET measurements (Holben et al., 1998; Dubovik and King, 2000; Dubovik et al., 2000; Sinyuk et al., 2007; Gilles et al., 2019). We used the Level 2.0 (quality assured) of the "Version 3 Direct Sun" and of the "Version 3.0 Inversions", except for the percentage of sphericity $\%S_{ph}$ for which we used Level 1.5 (in July 2021, Level 2.0 was not yet available for this parameter).

From these datasets, we selected all records corresponding to (1) aerosol optical thicknesses at 4 wavelengths 440, 675, 870, 1020 nm, (2) aerosol *Ångström* coefficients between 440 and 870 nm, which allows us to determine the aerosol optical thickness at 550 nm, and (3) microphysical properties C_{vf}, C_{vc}, $\%C_{vf}$, $\%C_{vc}$, r_{vf}, r_{vc}, σ_r, σ_c, nr_{440}, nr_{650}, nr_{850}, nr_{1020}, ni_{440}, ni_{650}, ni_{850}, ni_{1020}, $\%S_{ph}$.

A minimum threshold of 50 measurements of inversion product was used to exclude all sites without a sufficient number of measurements. We also ensure that all seasons are represented in the dataset. As mentioned above, one possible application of our aerosol microphysical model is the validation of satellite products in an operational context, whereby the atmospheric correction is performed when the aerosol loading is not too high. Thus, we decided to limit the dataset to aerosol optical thicknesses at 550 nm lower than 0.8.

Out of 1,139 available AERONET sites, we selected 851 globally distributed sites (Figure 1), resulting in ~1.3 million retrievals of aerosol microphysical properties. To characterize the representativeness of these sites, we analyzed the type of land cover surface around the selected AERONET sites. As shown in Figure 2, Urban (24%), Cropland (22%), Forest (17%), Grassland/Shrubland (16%) and Coastal areas and Islands (16%) are more or less equally represented.

For the measurements, AERONET instruments consist of two detectors mounted on robots a system developed by *Cimel-France*. One for the measurement of solar (and now lunar) extinction which provides the aerosol optical thicknesses (and then

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the *Ångström* coefficients) and the water vapor content. The other detector measures the luminance of the day sky using two protocols: the almucantar and the principal plane (see Holben et al. 1998, Tables 1 and 2). The almucantar procedure and measurements were used by Dubovik and King (2000) to derive the aerosol microphysical properties. Nevertheless, due to the observations protocol, the atmospheric condition (particularly its turbidity and homogeneity), the processing, and the retrieval purpose, the aerosol microphysical properties retrievals are not provided within a single retrieval. There are 3 different sets of retrievals:

(1) the size distribution C_{vf}, C_{vc}, $\%C_{vf}$, $\%C_{vc}$, r_{vf}, r_{vc}, σ_r, σ_c. This set of parameters is always available when aerosol microphysical properties retrievals are performed by AERONET. For this study, this first block provides a little less than 1.3M sets of retrievals for the whole 851 AERONET sites used.

(2) the complex refractive index for 4 wavelengths nr_{440}, nr_{650}, nr_{850}, nr_{1020}, ni_{440}, ni_{650}, ni_{850}, ni_{1020}. This set has a lower occurrence in terms of retrievals, only 0.17M sets of retrievals from 400 sites,

(3) the percentage of sphericity $\%S_{ph}$. This third set is available for the same 851 sites as (1) and provides a little less than 1.3M sets of retrievals. We decided to limit the non-sphericity at a 30% minimum. Indeed, deriving the non-sphericity integrated over the whole atmospheric column is challenging. Indeed, in almost all cases, particles are randomly oriented and the accumulation of all orientation along the vertical column generates a minimum of sphericity.

The AERONET network has existed since 1993. Figure 3 shows the number of AERONET sites we used for this study since 1993. For the last 9 years, we used more than 350 sites, 250 sites and 350 sites respectively for characterizing the size distribution, the refractive index and the sphericity. The decrease observed in 2020 is because all data have not yet been validated.

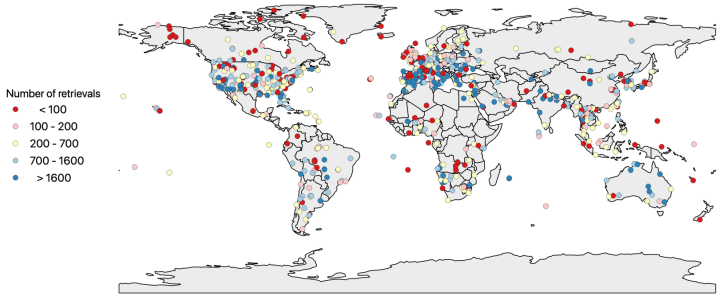
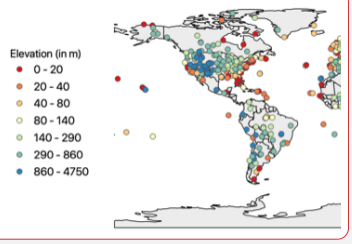


Figure 1. Location of the 851 AERONET sites with their number of retrievals.



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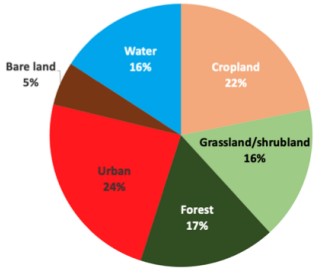

Figure 2. Representativeness of land surface types around the selected AERONET sites for the entire selected dataset.

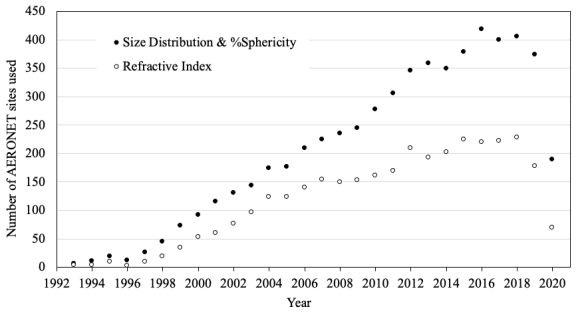

Figure 3. Number of AERONET sites selected for this study over the years. The size distribution and the refractive index are level 2.0 while the sphericity is level 1.5 (see text). The decrease in 2020 is because all data have yet to be validated.

A technical description of values for all aerosol microphysical properties for the 851 AERONET sites is presented Table 1., showing the percentile at 1%, 5%, 95% and 99% and the median value for each of the properties. This gives a global overview of aerosol microphysical properties over land.

Table 1. Description of the database of aerosol microphysical properties, for aerosol optical thickness τ at 440 nm, and the Ångström coefficient α (440,870).

	τ_{440}	$\alpha_{440\text{-}870}$	% C_{vf}	C_{vf} ($\mu m^3/\mu m^2$)	r_{vf} (μm)	σ_f	% C_{vc}	C_{vc} ($\mu m^3/\mu m^2$)	r_{vc} (μm)	σ_c	nr_{440}	ni_{440}	% S_{ph}
Percentile 0.01	0.016	0.11	5.9	0.0020	0.093	0.34	12	0.0010	1.2	0.51	1.33 (†)	0.001	30 (*)

Percentile 0.05	0.031	0.28	9.3	0.0030	0.11	0.37	25	0.0040	1.4	0.55	1.36	0.002	30 (*)
Median	0.14	1.26	33	0.014	0.14	0.47	67	0.026	2.1	0.68	1.47	0.006	63
Percentile 0.95	0.62	1.85	75	0.071	0.20	0.63	91	0.21	3.0	0.79	1.58	0.024	99
Percentile 0.99	0.89	2.03	88	0.11	0.24	0.72	94	0.39	3.4	0.85	1.60 (†)	0.036	99

(*) according to our threshold at 30%

(†) according to the AERONET threshold

AERONET sites don't have the same number of observations (See Figure 4). In the database we developed, one site may contain several thousands of selected retrievals for each aerosol microphysical properties. For example, eight sites provided more than 10,000 sets of retrievals for the Size Distribution i.e., *Sede Boker* (Israel), *Solar Village* (Saudi Arabia – no longer in the network), *GSFC* (USA), *Burjassot* (Spain), *El_Arenosillo* (Spain), *Carpentras* (France – no longer in the network), *Sevilleta* (USA), *Granada* (Spain). On the other hand, one site may contain less than 100 sets (this is the case for 138 sites). It means that 1 site may represent the equivalent of hundreds of other sites. To avoid the impact of those too well-represented sites, we show in Table 2 another way to present the similar information as Table 1. By applying a single median value per AERONET site for each aerosol microphysical parameters retrieval, we have 651 values for each microphysical parameters (400 for the refractive index). Then, we derive a median value reported Table 2. In this case, the median values don't change much (except for $\%S_{ph}$), but the range between both percentiles is reduced, by 20 to 50%. With the assumption of a median value per site, Figure 5 shows the frequency of τ_{440} and $\alpha_{440\text{-}870}$, while Figures 6, 7 and 8 show the frequencies of each aerosol microphysical property from our selected dataset.

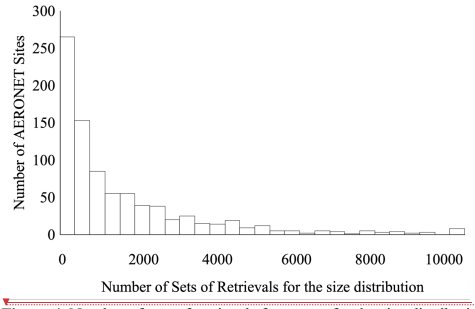

Figure 4. Number of sets of retrievals frequency for the size distribution.

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Table 2. Same than Table 1, but by affecting 1 median value of each microphysical parameter per each AERONET site, and then by deriving the median value of the 851 sites (400 for refractive index) for each microphysical parameter.

	τ_{440}	$\alpha_{440-870}$	% C_{vf}	C_{vf} ($\mu m^3/\mu m^2$)	r_{vf} (μm)	σ_f	% C_{vc}	C_{vc} ($\mu m^3/\mu m^2$)	r_{vc} (μm)	σ_c	nr_{440}	ni_{440}	% S_{ph}
Percentile 0.01	0.031	0.34	11	0.0032	0.12	0.40	24	0.0031	1.7	0.60	1.40	0.0025	30
Percentile 0.05	0.066	0.55	17	0.0057	0.13	0.42	36	0.010	1.8	0.62	1.42	0.0032	34
Median	0.19	1.31	42	0.021	0.15	0.47	58	0.031	2.2	0.67	1.47	0.0065	71
Percentile 0.95	0.55	1.76	64	0.065	0.17	0.55	84	0.18	2.7	0.72	1.52	0.020	93
Percentile 0.99	0.67	1.88	76	0.088	0.19	0.60	89	0.24	2.9	0.74	1.54	0.026	97

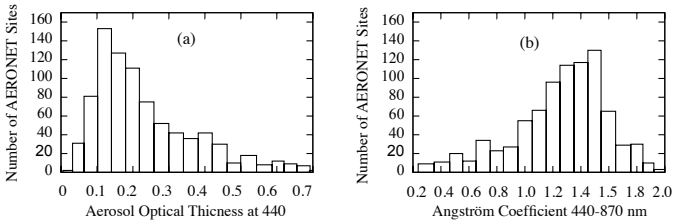

Figure 5. Aerosol optical thickness at 440nm frequency (a) and the Ångström coefficient frequency (b).

2.3 Metrics used

The results of the retrievals are evaluated using three performance metrics: accuracy, precision, and uncertainty (APU):

- The accuracy A represents the average bias of the estimates:

$$A = \frac{1}{N}\sum_{i=1}^{N}(C_i - R_i) \quad (2)$$

- The precision P is the deviation around the mean value:

$$P = \sqrt{\frac{1}{N-1}\sum_{i=1}^{N}(C_i - R_i - \mathrm{A})^2} \quad (3)$$

- The uncertainty U encompasses all errors and is derive from A and P

$$U = \sqrt{\frac{1}{N}\sum_{i=1}^{N}(C_i - R_i)^2} = \sqrt{A^2 + \frac{N-1}{N}P^2} \quad (4)$$

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were C_i is the computed value with our proposed model, R_i is the reference values, and if N is the number of data.

The relative uncertainty is defined here as: U/V where V can be the mean value of a specific site or of the whole set of a specific parameter.

3 Aerosol microphysical properties

3.1 Parameterization of the aerosol microphysical properties

Two measurements protocols are followed to acquire AERONET data. The Aerosol Optical Thicknesses (AOT) is regularly measured every 15 minutes following a direct measurement of the Sun when cloud-free. For the retrieval of the aerosol model microphysical properties, as specified above, the protocol required an almucantar measurement (Holben et al., 1998; Dubovik and King, 2000), which is performed early in the morning or late afternoon. The main issue is that this AERONET measurement might not be coincident with the Earth observation satellites overpass times. Moreover, for various reasons (e.g., inhomogeneous sky, small clouds, calibration procedure…) some measurements might be missing. We can obviously interpolate data between two available measurements, but we miss the variability of the considered aerosols. As an illustration, Figure 9 shows an example of the impact of changing the aerosol model for size distribution from early morning (7:21:30 local time) to late afternoon (16:28:45 local time). In this example, there is an increase in coarse aerosols between the morning and the evening, but we don't exactly know when that occurred.

In 2002, Dubovik et al. suggested to determine each microphysical parameter with a direct regression (Equation 5) using the Aerosol Optical Thickness at 440nm from the AERONET dataset.

$$\text{Aerosol Microphysical Parameter} = a + b.\tau_{440} \quad (5)$$

For each AERONET site, this approach has been used so far for the official validation of the MODIS and VIIRS surface reflectance products (Vermote et al., 2002; Vermote et al., 2014), for the NASA HLS (Harmonization Landsat-Sentinel) project (Claverie et al., 2018; Vermote et al., 2016), and for the CEOS ESA/NASA ACIX exercise (Doxani et al, 2018). Our objective here is to better account for the temporal and spatial variability of the aerosol microphysical parameters, which can't be only related to the aerosol optical thickness itself. In an operational context, another possible and simple variable available for the aerosol description is the Ångström coefficient α (Ångström, 1929). Indeed, it's well accepted that this coefficient is related to the aerosol size (which is important in term of light-matter interaction). If we take the example given in Figure 9, we can see from Figure 10 that the aerosol optical thickness doesn't change between the two almucantar procedures, while the Ångström coefficient does. The value of the latter decreases, indicating a bigger particle represented by a bigger coarse mode, which is consistent with Figure 9. Another reason to select a multiplication of the optical thickness τ and the Ångström

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coefficient α is conceptual. The aerosol optical thickness τ is an extensive parameter, the Ångström coefficient α is an intensive parameter, and it's preferable to have a multiplication of a couple of intensive/extensive variables in physical parametrization, as their multiplication remains an extensive parameter. Indeed, an intensive parameter can be used for identifying a sample while an extensive parameter can be used for describing this sample.

We decided to select the Ångström coefficient for the 440 and 870 nm wavelengths, i.e., $\alpha_{440\text{-}870}$. Even if the Ångström coefficient has a dynamic behavior over the visible range and it is not entirely constant, $\alpha_{440\text{-}870}$ is a good compromise between all values. At the end, we selected τ_{440} and $\alpha_{440\text{-}870}$ as variables of the regression. Within the AERONET network, these variables are available every 15 minutes under clear sky condition for all sites.

We can also use the water vapor content, as it's a very important parameter in terms of the microphysical properties. Some aerosols are hydrophilic, other are hydrophobic. Water vapor also modifies the size of the aerosol and its absorption capacity. We explored this option, but it didn't improve the retrieval in term of uncertainties. The aerosol optical thickness parameter already includes the effect of the water vapor over the aerosol size distribution, and it explains in part why there were no improvements.

One limited aspect of our approach is that these two parameters τ_{440} and $\alpha_{440\text{-}870}$ directly correspond to the aerosol scattering, and we may not properly characterize the aerosol absorption (Fraser and Kaufman, 1985; Vermote et al., 2007; Russell et al., 2010; Giles et al., 2012; Lenoble et al., 2013; Tsikerdekis et al., 2021). Therefore, the complexity of the radiative transfer through the atmosphere partially allows mitigation of this phenomenon. Indeed, coupling between the scattering and the absorption of light allows us to indirectly capture the aerosol absorption information.

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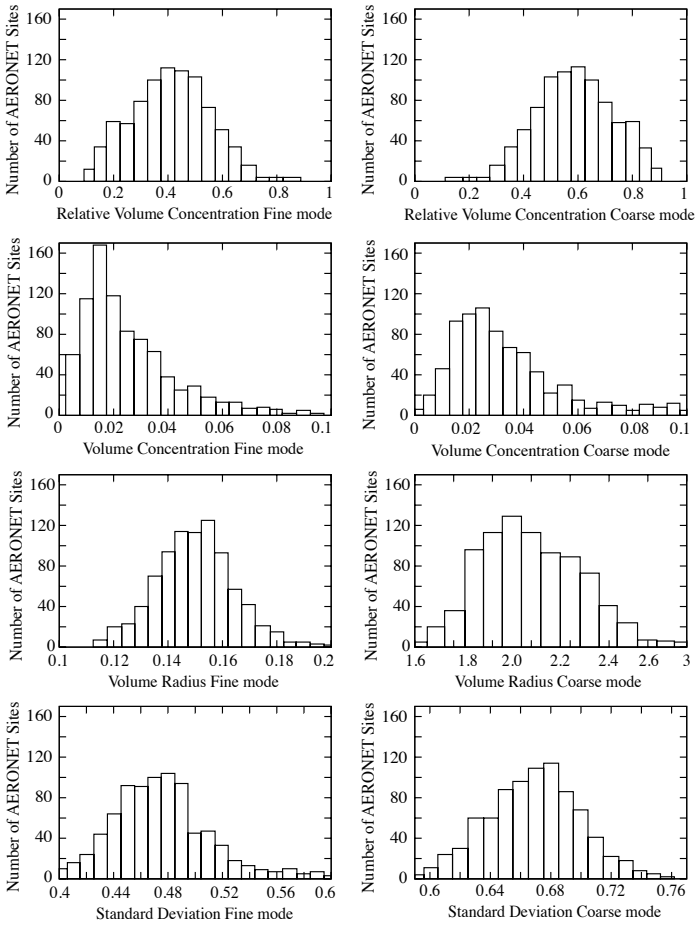
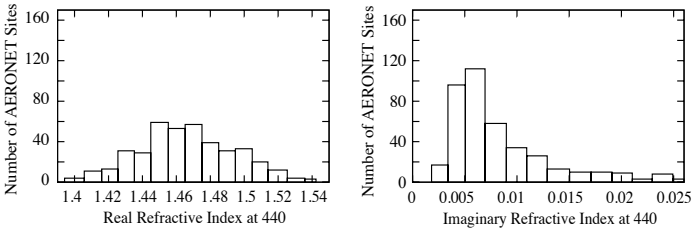


Figures 6. Size distribution parameters frequency for the fine mode (left) and the coarse mode (right).



Figures 7. Real (left) and imaginary (right) refractive frequency.

435

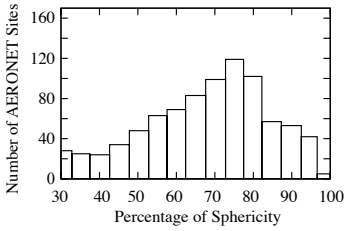
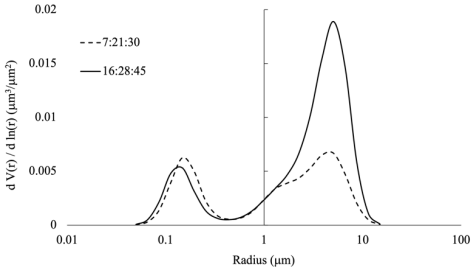
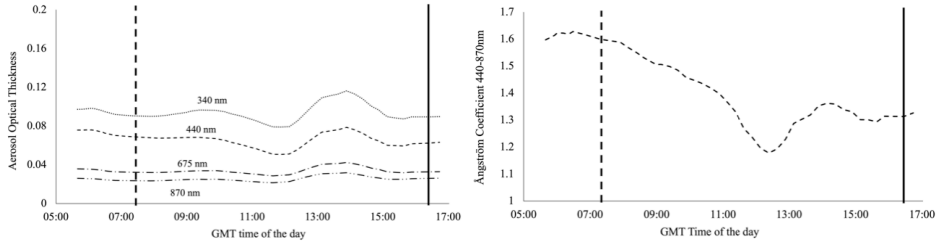


Figure 8. Percentage of sphericity frequency.



440

Figure 9. Example of an aerosol size-distribution from AERONET with a change between 2 almucantar procedures occurring between the early morning and late afternoon observations (data acquired at the Aubiere site in July 2014).



Figures 10. Daily variability of the Aerosol Optical Thickness (left) and of the Ångström coefficient 440-870nm (right) for the example of Figure 9 (data acquired at the Aubiere site in July 2014).

With our AERONET database (over the 400 sites where we have all microphysical properties), we explored several mathematical formulations for a regression between an aerosol microphysical property, called *AMP* in the following equations, and the two variables τ_{440} and $\alpha_{440\text{-}870}$. We used a similar idea after the *Dubovik*'s law (Equation 5). We first tested Equation 5. Then, we tested a linear regression with the Ångström coefficient $\alpha_{440-870}$.

$$AMP_i = a_i + b_i.\alpha_{440\text{-}870} \quad (6)$$

Where i represents one of the microphysical properties (e.g., C_{vf}, C_{vc}, $\%C_{vf}$, $\%C_{vc}$, r_{vf}, r_{vc}, σ_r, σ_c, nr_{440}, nr_{650}, nr_{850}, nr_{1020}, ni_{440}, ni_{650}, ni_{850}, ni_{1020}, $\%S_{ph}$).

Finally, we tested several mathematical formulations using our two predicted variables, and we found that each aerosol microphysical parameter, P, can be optimally described by:

$$AMP_i = \left(a_i + b_i.\tau_{440}{}^{c_i}\right).\left(d_i + e_i.\alpha_{440-870}{}^{f_i}\right) \quad (7)$$

In practice, to better use Equation 7 and for the stability of retrievals, all 6 coefficients a_i, b_i, c_i, d_i, e_i and f_i are not derived with a single interaction. The aerosol microphysical parameters meanly depend on τ_{440} or on $\alpha_{440-870}$ (they barely depend on both at the same level). Thus, to get a stable retrieval of the 6 coefficients, we used a "so-called" residue approach by checking which of the $\left(a_i + b_i.\tau_{440}{}^{c_i}\right)$ or $\left(d_i + e_i.\alpha_{440-870}{}^{f_i}\right)$ is the most representative (i.e., with the best regression coefficient) regarding the behaviour of the microphysical parameters. Following this procedure, we apply the first regression law $\left(a_i + b_i.\tau_{440}{}^{c_i}\right)$ or $\left(d_i + e_i.\alpha_{440-870}{}^{f_i}\right)$ to derive (a_i, b_i, c_i) or (d_i, e_i, f_i) respectively. Then, according which one has the best correlation coefficient and using the remaining residue, we apply the second regression law $\left(d_i + e_i.\alpha_{440-870}{}^{f_i}\right)$ or $\left(a_i + b_i.\tau_{440}{}^{c_i}\right)$ to derive the

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missing triplet of coefficients. Table 3 shows the percentage of occurrence for τ_{440} or $\alpha_{440-870}$ as the most representative variable for all microphysical parameters and for all available AERONET sites (see Figure 2).

Table 3. Percentage of occurrence for the aerosol optical thickness τ_{440} and the Ångström coefficient $\alpha_{440-870}$ as giving the regression coefficient for each microphysical parameter.

	% C_{vf}	C_{vf} ($\mu m^3/\mu m^2$)	r_{vf} (μm)	σ_f	% C_{vc}	C_{vc} ($\mu m^3/\mu m^2$)	r_{vc} (μm)	σ_c	nr_{440}	ni_{440}	% S_{ph}
τ_{440}	6	100	50	22	6	79	29	62	39	22	18
$\alpha_{440-870}$	94	0.1	50	78	94	21	71	38	61	78	82

In Table 3, for 7 of the 11 parameters, $\alpha_{440-870}$ is more correlated with the microphysical parameter than τ_{440}. It confirms that the use of α is pertinent to define these parameters. As expected, C_{vf} and C_{vc} are mostly driven by τ_{440} (Sinyuk et al, 2020), while $\%C_{vf}$ and $\%C_{vc}$ are driven by $\alpha_{440-870}$. Parameters C_{vf} and C_{vc}, which are extensive parameters, are directly related to the volume loading (mass) of the aerosol, and, in fine, to the number of particles (accumulation of particles). Thus, it's not surprising that C_{vf} is more correlated to τ_{440} than C_{vc}. Indeed, we know that the fine mode optically reacts more efficiently in the visible light than the coarse mode in terms of extinction (Van der Hulst, 1981), considering that the number of particles present in the fine mode is usually much higher than the number of particles of the coarse mode. By the same reasoning, $\%C_{vf}$ and $\%C_{vc}$, which are intensive parameters, are not sensitive to accumulation, but rather to the spectral dependency of the aerosol extinction, meaning that $\%C_{vf}$ and $\%C_{vc}$ are more correlated to $\alpha_{440-870}$. In the AERONET processing, the complex refractive index is applied when the AOT is higher than 0.4 at 440nm. This limits the variability in term of AOT and probably increases artificially the occurrence for $\alpha_{440-870}$.

We applied our approach for the 3 mathematical formulations given by Equations 5, 6 and 7 over the whole selected dataset and present the results in Table 4.

Table 4. Mean relative uncertainties (in percent) for each retrieved aerosol microphysical properties modelled using several mathematical formulations over the whole dataset.

	% C_{vf}	C_{vf} ($\mu m^3/\mu m^2$)	r_{vf} (μm)	σ_f	% C_{vc}	C_{vc} ($\mu m^3/\mu m^2$)	r_{vc} (μm)	σ_c	nr_{440}	ni_{440}	% S_{ph}
$a+b.\tau_{440}$	34.1	31.8	11.9	10.1	21.9	51.6	15.2	6.9	3.1	39.5	26.7
$a+b.\alpha_{440-870}$	24.3	66.0	12.0	9.2	16.1	59.4	14.5	7.0	3.1	38.4	23.6
$(a+b.\tau_{440}{}^{c}).(d+e.\alpha_{440-870}{}^{d})$	22.6	30.3	11.4	8.8	15.0	35.0	14.1	6.7	3.0	37.5	22.8

In terms of accuracy A (Equation 2), results show very low values. Except for C_{vf}, C_{vc}, and $\%C_{vf}$ which present an accuracy up to 2%, accuracies of all other microphysical parameters are below 0.1%. For uncertainty U (Equation 4), the third mathematical formulation gives the overall best results (Table 4). As expected, τ_{440} better represents C_{vf} while in contrast $\alpha_{440-870}$ better

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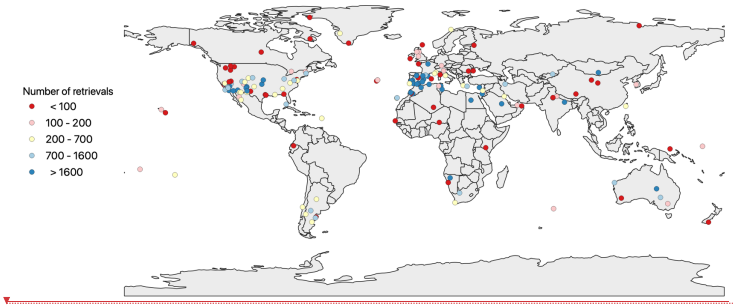
represents the $\%C_{vf}$. Finally, including both variables, we get a non-negligible improvement for both volume concentrations (absolute and relative). For the other microphysical properties, we don't observe much of improvement, but the Equation 7 gives consistently better results. One point to be noted is that all microphysical properties provided by the AERONET network have lower uncertainties than those presented in Table 4 (Dubovik et al., 2000; Sinyuk et al., 2020).

As pointed out, $\%C_{vf}$ and $\%C_{vc}$ globally present a better uncertainty than for C_{vf} and C_{vc}, but for exactly 20% of sites the volume concentration of the fine mode C_{vf} is more accurate than the relative volume concentration $\%C_{vf}$ (Figure 11). We are unable to find a clear reason to explain that. The only tiny explanation is that aerosols over these sites present a tendency described by (1) lower concentrations than the average (both fine and coarse modes), meaning relatively low optical thickness, (2) a relative lower Ångström coefficient, and (3) a relative lower absorption. Nevertheless, according to the radiative transfer theory used to define the optical properties (Phase Matrix, Scattering and Absorption coefficients), the phase matrix is normalized at the end. Thus, either the couple of volume concentrations (C_{vf},C_{vc}) or the couple of relative volume concentrations ($\%C_{vf}$,$\%C_{vc}$) can (it should be a couple) be used depending on the uncertainty for one AERONET site. It should be noted that, in all cases, the uncertainty U of $\%C_{vc}$, $U_{\%Cvc}$, is always lower than that of C_{vc}, U_{Cvc}.



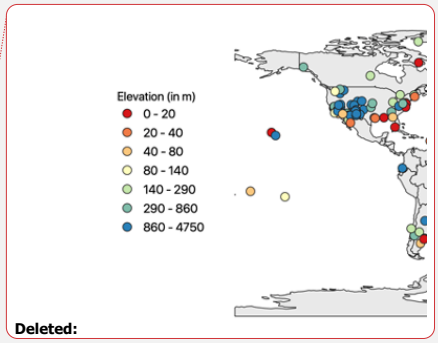


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Figure 11. AERONET sites for which C_{vf} is better represented than $\%C_{vf}$.

Table 5 shows the new uncertainties U of $\%C_{vf}$, $U_{\%Cvf}$, and the new uncertainties U of C_{vf}, U_{Cvf}, when we only select sites for which $U_{\%Cvf} > U_{Cvf}$ (80% of cases) or $U_{Cvf} > U_{\%Cvf}$ (20% of cases) respectively. The improvement is visible if we use both $\%C_{vf}$ and C_{vf} according to the lowest uncertainties.

Table 5. Uncertainties (in percent) for each retrieved aerosol microphysical properties model (as for Table 4), but after selecting sites for %Cv with $U_{\%Cvf} > U_{Cvf}$ (†, 80% of cases) and for Cv with $U_{Cvf} > U_{\%Cvf}$. (††: 20% of cases).

	$\%\ C_{vf}$	C_{vf} $(\mu m^3/\mu m^2)$	r_{vf} (μm)	σ_f	$\%\ C_{vc}$	r_{vc} (μm)	σ_c	nr_{440}	ni_{440}	$\%\ S_{ph}$

$(a + b.\tau_{440}^{c}).(d + e.\alpha_{440-870}^{f})$	22.0 (†)	22.0 (††)	11.4	8.8	15.0	14.1	6.7	3.0	37.5	22.8

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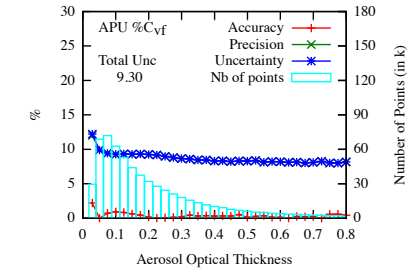
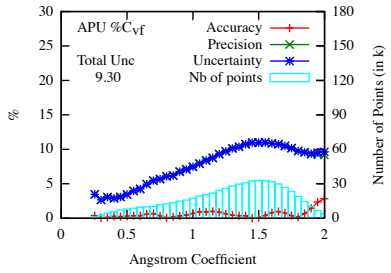
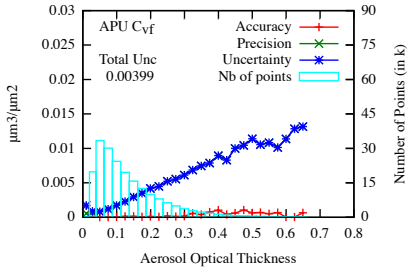
As pointed out previously, we have 50% of sites without any refractive indexes. One solution to improve the number of sites is to define mean parameters (*a*, *b*, *c*, *d*, *e*, *f*) for *nr* and *ni* by kind of environment (urban, urban coastal, forest, non-forest land, desert… for example). In that context, we undertook a preliminary study which included all data independently of the site to retrieve mean parameters. It gave a relative uncertainty *U* of 3.0% for *nr* with no change compared to Tables 4 and 5. In contrast, for *ni*, it showed a relative uncertainty of 52% for *ni* which represents about 40% higher than those shown in Tables 4 and 5, but this study includes all data without distinguishing of environment. If we are able to specifically define the environment of the missing sites, we should get a relative uncertainty closer to 37.5% (Tables 4 and 5) rather closer to 52%, which remains acceptable.

3.2 Retrieved microphysical properties from the whole dataset

To expand on Table 4, Figures 12 give the APU of the retrieved microphysical properties over the whole dataset versus τ_{440} and $\alpha_{440-870}$. The interesting point of these figures is the dependency of uncertainties with τ_{440} and $\alpha_{440-870}$. Indeed, except for C_{vf} and C_{vc}, uncertainties are quite stable with the aerosol optical thickness. On the contrary, most of uncertainties present variation with the Ångström coefficient. This confirms the importance of considering $\alpha_{440-870}$ in the regression. Another point is the correlation between Tables 4 and 5 and Figures 12. When the variability of the uncertainty with the $\alpha_{440-870}$ is important (Figure 12), the variability of the microphysical properties is more important as well (Tables 4 and 5). It should be noted that for $\%C_{vf}$ and for C_{vf}, the APU are for selected sites only (see Table 5).


APU %Cvf
Total Unc
9.30
Accuracy
Precision
Uncertainty
Nb of points
%
Number of Points (in k)
Aerosol Optical Thickness


APU %Cvf
Total Unc
9.30
Accuracy
Precision
Uncertainty
Nb of points
%
Number of Points (in k)
Angstrom Coefficient


APU Cvf
Total Unc
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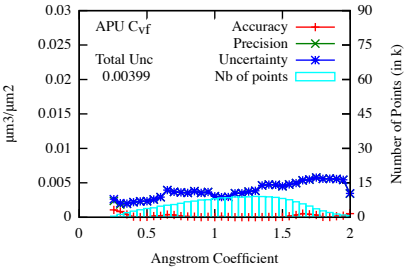
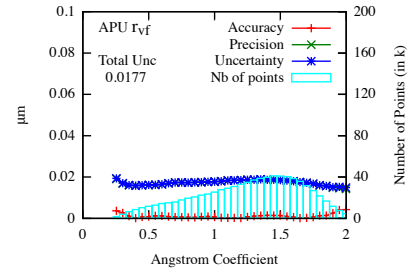
Accuracy
Precision
Uncertainty
Nb of points
µm3/µm2
Number of Points (in k)
Aerosol Optical Thickness


APU Cvf
Total Unc
0.00399
Accuracy
Precision
Uncertainty
Nb of points
µm3/µm2
Number of Points (in k)
Angstrom Coefficient


APU rvf
Total Unc
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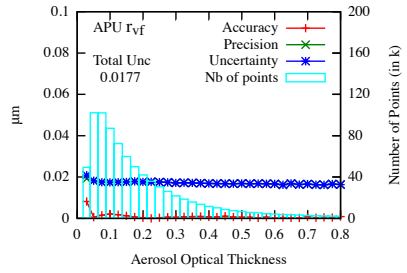
Accuracy
Precision
Uncertainty
Nb of points
µm
Number of Points (in k)
Angstrom Coefficient


APU rvf
Total Unc
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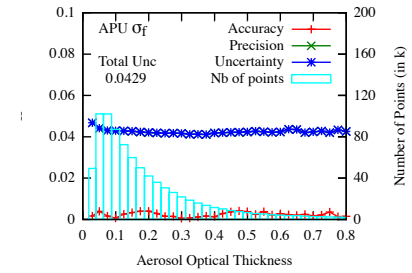
Accuracy
Precision
Uncertainty
Nb of points
µm
Number of Points (in k)
Aerosol Optical Thickness


APU σf
Total Unc
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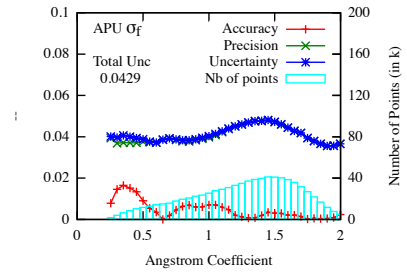
Accuracy
Precision
Uncertainty
Nb of points
--
Number of Points (in k)
Aerosol Optical Thickness


APU σf
Total Unc
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Accuracy
Precision
Uncertainty
Nb of points
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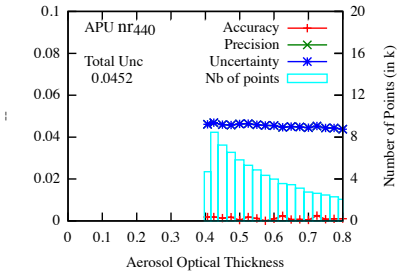
Number of Points (in k)
Angstrom Coefficient


APU nr440
Total Unc
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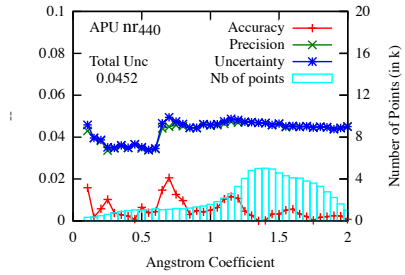
Accuracy
Precision
Uncertainty
Nb of points
Number of Points (in k)
Aerosol Optical Thickness


APU nr440
Total Unc
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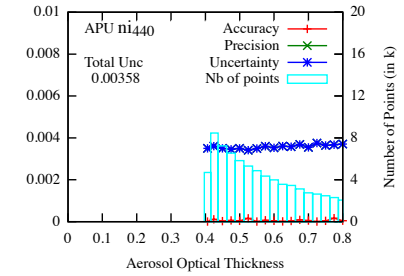
Accuracy
Precision
Uncertainty
Nb of points
Number of Points (in k)
Angstrom Coefficient


APU ni440
Total Unc
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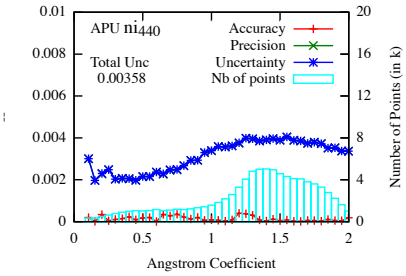
Accuracy
Precision
Uncertainty
Nb of points
Number of Points (in k)
Aerosol Optical Thickness


APU ni440
Total Unc
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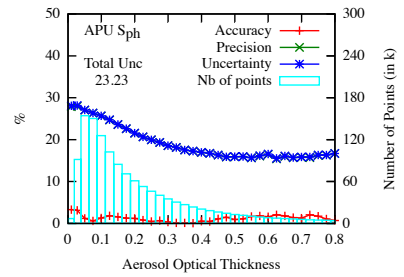
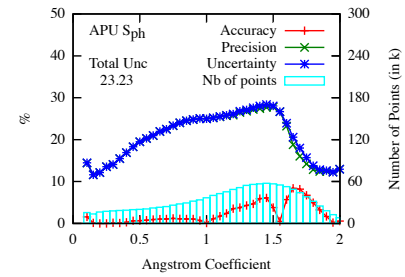
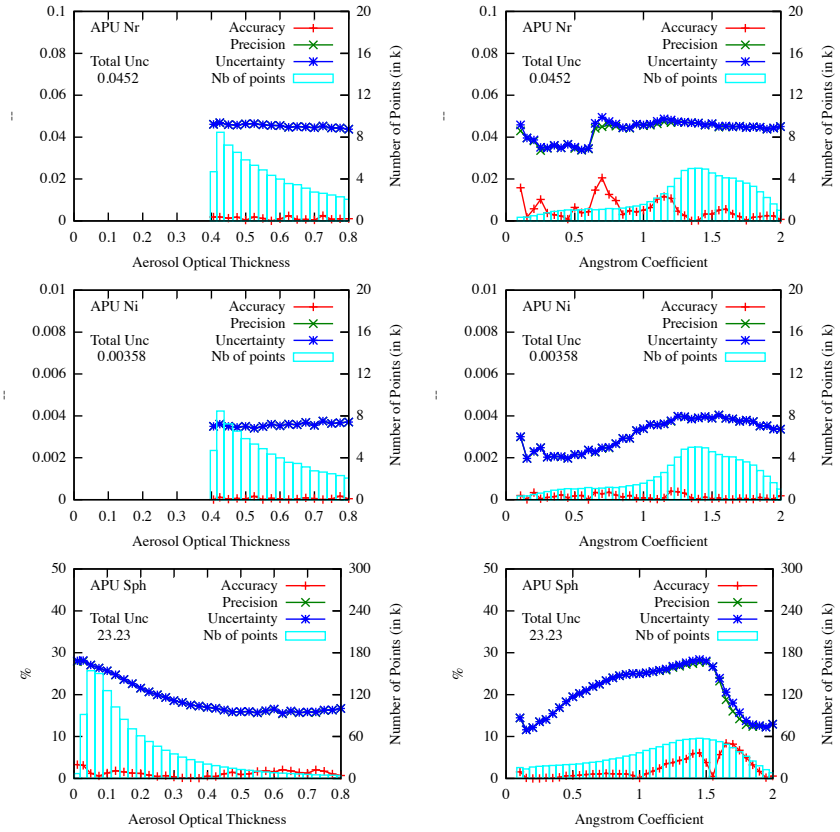
Accuracy
Precision
Uncertainty
Nb of points
Number of Points (in k)
Angstrom Coefficient


APU Sph
Total Unc
23.23
Accuracy
Precision
Uncertainty
Nb of points
%
Number of Points (in k)
Aerosol Optical Thickness


APU Sph
Total Unc
23.23
Accuracy
Precision
Uncertainty
Nb of points
%
Number of Points (in k)
Angstrom Coefficient

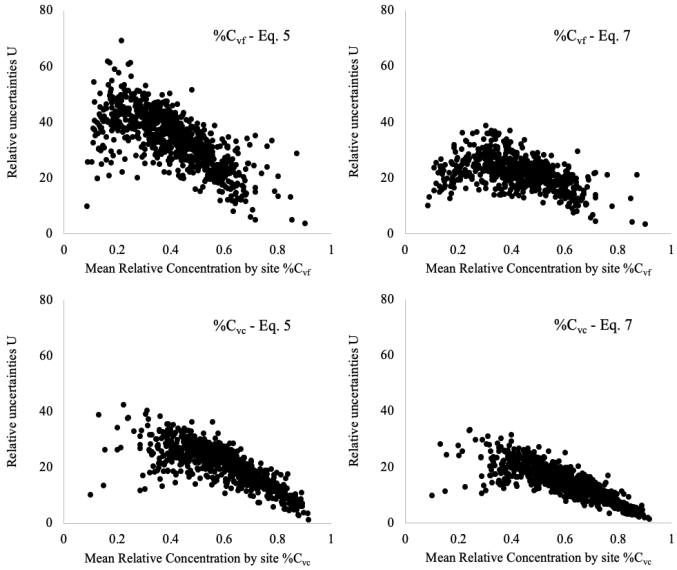


Figures 12. APU for the retrieval of each microphysical parameter (from top to bottom: the 8 parameters describing the size distribution - fine and coarse modes, the 2 parameters for the refractive index at 440 nm, and the parameter for the sphericity), versus the aerosol optical thickness at 440 nm (left) and the Ångström coefficient between 440 and 870 nm (right). "Total Unc" represents the total uncertainty of the microphysical parameter.

3.3 Retrieved microphysical properties considering each AERONET site

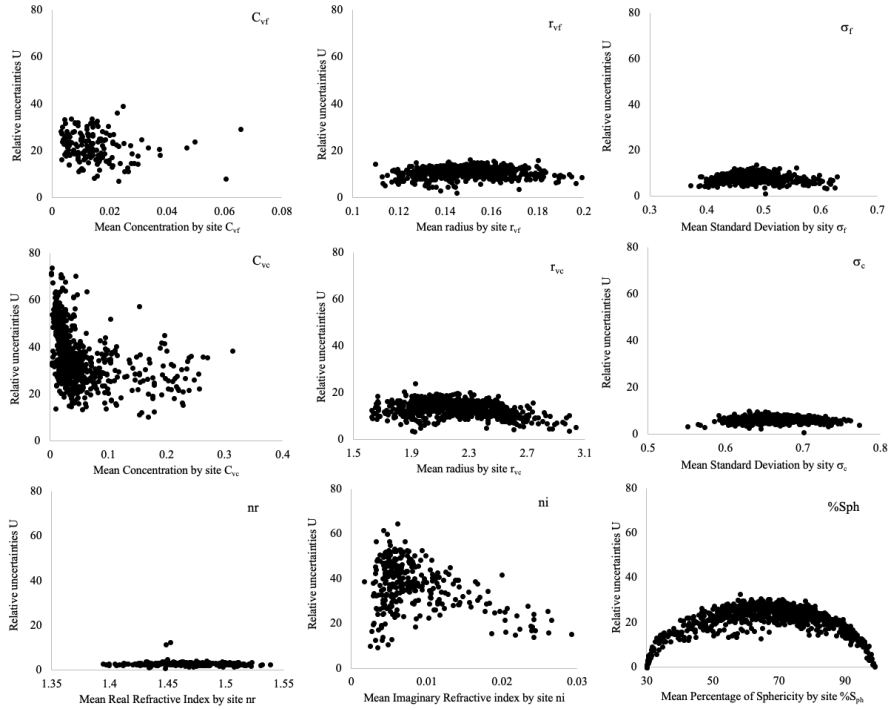
The use of $\alpha_{440\text{-}870}$ mostly improves the retrieval of both $\%C_{vf}$ and $\%C_{vc}$ (Tables 4 and 5). Figure 13 shows the comparison between uncertainties on $\%C_{vf}$ and $\%C_{vc}$ using Equation 5 or Equation 7 versus the mean value of $\%C_{vf}$ and $\%C_{vc}$ for each

AERONET site (one dot represents one AERONET site). For $\%C_{vf}$, we only consider sites where $U\%C_{vf} < UC_{vf}$. These figures highlight the improvement of retrievals (about a 1/3 less). We can also point out that relative uncertainties are lower for high and low values of $\%C_{vf}$.



595 **Figures 13.** Comparison of the relative uncertainties (%) when using Equation 5 (left) and Equation 7 (right) are used to derive $\%C_{vf}$ and $\%C_{ve}$. One point corresponds to one AERONET site.

Figures 14 give the relative uncertainty for the other microphysical properties site by site, but only using Equation 7 (for C_{vf}, we only consider sites where $UC_{vf} < U\%C_{vf}$). Again, except for the volume concentration C_{vf} and C_{ve}, we can notice the "arch"
600 effect generating a lower relative uncertainty for lower values and for higher values of the considered properties. It's not shown here, but this "arch" effect is even more important with absolute uncertainties. At the end, we are able to characterize the uncertainties for each aerosol microphysical property and for each AERONET site.



Figures 14. Relative Uncertainties of the aerosol microphysical properties versus the property itself using Equation 7 (one point corresponds to one AERONET site).

3.4 Impact of the uncertainties on the surface reflectance product over land

As previously mentioned, this work is to support atmospheric correction validation over land. Thus, one question is: how does the uncertainty of the retrieved aerosol microphysical property affect the surface reflectance product validation? To address this issue, we decided to define, for each aerosol microphysical property, the impact of its uncertainty (Table 5) on the atmospheric correction, and the determination of the surface reflectance over land. For that purpose, we defined a synthetic database of TOA reflectances for each AERONET site and for each specific satellite band. To generate this database, we used the 6S code (Vermote et al., 1997; Kotchenova et al., 2006; Kotchenova et al., 2007; Kotchenova et al., 2008) with the following inputs: (1) a set of 80 viewing conditions (solar angle, view angle, azimuth angle) describing all satellite angular

configurations possible, (2) a set of different atmospheres (pressure, temperature, water vapor), (3) a set of surface reflectances (from 0 to 0.6 depending on the wavelength), and (4) a set of 40 aerosol microphysical properties with associated τ_{440} and $\alpha_{440-870}$ picked up in the real AERONET database. Then, we applied the atmospheric scheme developed for the Land Surface Reflectance Code (LaSRC) algorithm for MODIS, VIIRS, Landsat-8, Sentinel-2 (Vermote et al., 2002; Vermote et al., 2014; Vermote et al., 2016; Claverie et al., 2018; Doxani et al., 2018). First, using each set of input, we computed the TOA reflectance. Then, inducing 20 cases of random uncertainties for each aerosol microphysical properties, we applied an atmospheric correction to get the surface reflectance ρ_{surf} to be compared to the one used as input. Table 6 gives the uncertainties we get for the MODIS red channel (band 1, 620-670 nm). For example, % Cvf is generated with an uncertainty of 22.0%. This uncertainty generates, once we proceed an atmospheric correction scheme, an uncertainty of 0.00014 on the surface reflectance (in reflectance unit). The main relative uncertainty appears for the uncertainty Uni_{440} of the imaginary part of the refractive index (relies to the aerosol absorption), 1.0 10^{-3} in terms of surface reflectance, followed by the uncertainty of the radius of the fine mode. In a decreasing order of magnitude, Ur_{vf} and Unr_{440} appear around a third lower. Then, another step below, appears UC_{vf} and $U\%\ C_{vf}$.

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Table 6. Surface reflectance uncertainties (for the MODIS Red channel) due to the initial aerosol model uncertainties (in reflectance unit).

	$\%\ C_{vf}$	C_{vf}	r_{vf}	σ_f	$\%\ C_{vc}$	r_{vc}	σ_c	nr_{440}	ni_{440}	$\%\ S_{ph}$
Initial relative Uncertainty (Table 5)	22.0 %	22.0 %	11.4 %	8.8 %	15.0 %	14.1 %	6.7 %	3.0 %	37.5 %	22.8 %
In-fine uncertainties on surface reflectances	1.4 10^{-04}	1.5 10^{-04}	3.9 10^{-04}	4.0 10^{-05}	6.7 10^{-05}	5.5 10^{-05}	2.6 10^{-05}	3.6 10^{-04}	1.0 10^{-03}	6.0 10^{-05}

Many atmospheric correction schemes use a blue channel to retrieve the aerosol properties, so it's interesting to assess the impact of the aerosol model with the atmospheric reflectance in the blue channel. Figure 15 shows, for an example with the MODIS blue channel (band 3), the dependency between the uncertainties on ρ_{surf} in the red channel and the atmospheric reflectance in the blue channel. This uncertainty is always low, below 0.005, for a range of reasonable atmospheric reflectance values. This figure also shows that this aerosol reflectance in the blue channel is almost linearly correlated to the uncertainties of the surface reflectance in the red channel. This means that a QA flag can be directly defined using the atmospheric reflectance in the blue channel rather than the optical thickness (Vermote et al., 2002; Vermote et al., 2014).

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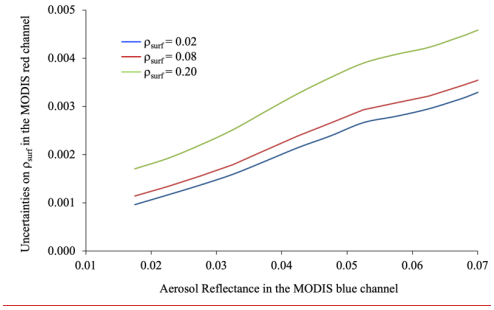

Figure 15. Uncertainties of ρ_{surf} in the MODIS red channel versus the aerosol reflectance in the MODIS blue.

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Finally, Figure 16 represents, in fine, the impact of the aerosol model uncertainties retrieved using Equation 7 on the surface reflectance retrieval ρ_{surf} in the MODIS red spectral band. Uncertainties, shown for two ranges of aerosol optical thicknesses at 550nm τ_{550} - less than 0.25 and less than 0.50, are clearly always below the MODIS specification required for the surface reflectance ($0.005 + 0.05*\rho_{surf}$). For ρ_{surf} ranged between 0.10 and 0.40, the uncertainty on ρ_{surf} is relatively between 1 and 2%. It confirms that our aerosol model description for the AERONET sites can be used with a good confidence for the satellite atmospheric correction.

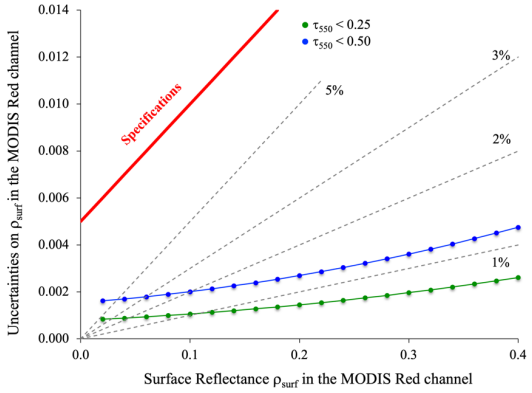

Figure 16. Uncertainties of ρ_{surf} versus ρ_{surf} in the MODIS red channel. Green and blue lines correspond to the uncertainties for 2 ranges of aerosol optical thicknesses at 550 nm (<0.25 and 0.5). Relative uncertainties on ρ_{surf} (1, 2, 3 and 5%) are highlighted in dot lines. The Red line indicates the MODIS specifications for surface reflectance retrieval.

4 Conclusion

This study was aimed at defining and building an aerosol model based on the microphysical parameters obtained for 851 AERONET sites. The AERONET network provides the aerosol microphysical parameters during the almucantar procedures (early morning, late afternoon), which might not be at the time when a satellite passes over an AERONET site. Thus, we upgraded the methodology used by Dubovik et al. (2002) to define the aerosol microphysical parameters and then the aerosol optical properties. Using the optical thickness at 440 nm τ_{440} and the Ångström coefficients $\alpha_{440\text{-}870}$ of aerosols, we characterized each microphysical parameter of the aerosols (C_{vf}, C_{vc}, $\%C_{vf}$, $\%C_{vc}$, r_{vf}, r_{vc}, σ_f, σ_c, nr_{440}, nr_{650}, nr_{850}, nr_{1020}, ni_{440}, ni_{650}, ni_{850}, ni_{1020}, $\%S_{ph}$) for each AERONET site. Compared to initial values, retrievals of the microphysical parameters are done with an acceptable uncertainty (from 6.6% to 20.7%), with the imaginary part of the refractive index being the least well-rendered parameter (c. less than 40%), which is not a surprise since this parameter is the most difficult to retrieve from optical measurements. The study shows different behaviors according to the value of each microphysical parameters, showing an arch effect resulting from lower uncertainty for the highest values and the lowest values of the microphysical parameters.

One use of this characterization is the validation of space-borne remote sensing sensors products, and in particular for the validation of the atmospheric correction over land, but this can be extended to other applications requiring aerosol information. In terms of atmospheric correction over land, this method can be used to define a surface reflectance reference as we do for the validation of surface reflectance products for sensors such as MODIS, VIIRS, Landsat and Sentinel-2. An impact study of the uncertainties of each aerosol microphysical parameters showed that the aerosol models used to define a reference surface reflectance provide a maximum uncertainty always lower than 0.004 in reflectance unit, or of 1 to 3 % (for surface reflectance higher than 0.05 in the MODIS red channel), well below the specifications often used for atmospheric correction. It's worth emphasizing that the imaginary part of the aerosol refractive index generates the more important uncertainties (0.001 in reflectance unit) and corresponds to a major part of the total uncertainty. Nevertheless, it will be important to further test these findings using additional datasets for validation (number of sites and number of comparisons).

Data availability: Data and all a_i, b_i, c_i, d_i, e_i and f_i coefficients are currently available in our web site (https://salsa.umd.edu).

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Competing interests: The authors declare that they have no conflict of interest.

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Annex: Nonparametric model approach

To test the ability of the optical thickness and the Ångström coefficient to be reliable for reproducing the aerosol models, we used a nonparametric approach. A Random Forest (RF) regression model was built with AOT and angstrom coefficient as inputs and all other parameters as outputs (dependent variables). The data were randomly split into training (50%) and test sets. The split was done in order to analyze the robustness of the model. Performance of the model (APU diagram) was assessed on testing data. The RF model had 100 trees and maximum depth of trees was limited to 15 to avoid overfitting.

Figures A1 give examples of results of this nonparametric approach (for parameters describing the fine mode of the size distribution only, but the conclusion can be generalized to all microphysical parameters). Comparing to Figures 12, we have similar results for presented examples of retrieved microphysical properties. This indicates that the use of the optical thickness τ_{440} and the Ångström coefficient $\alpha_{440\text{-}870}$ is consistent.

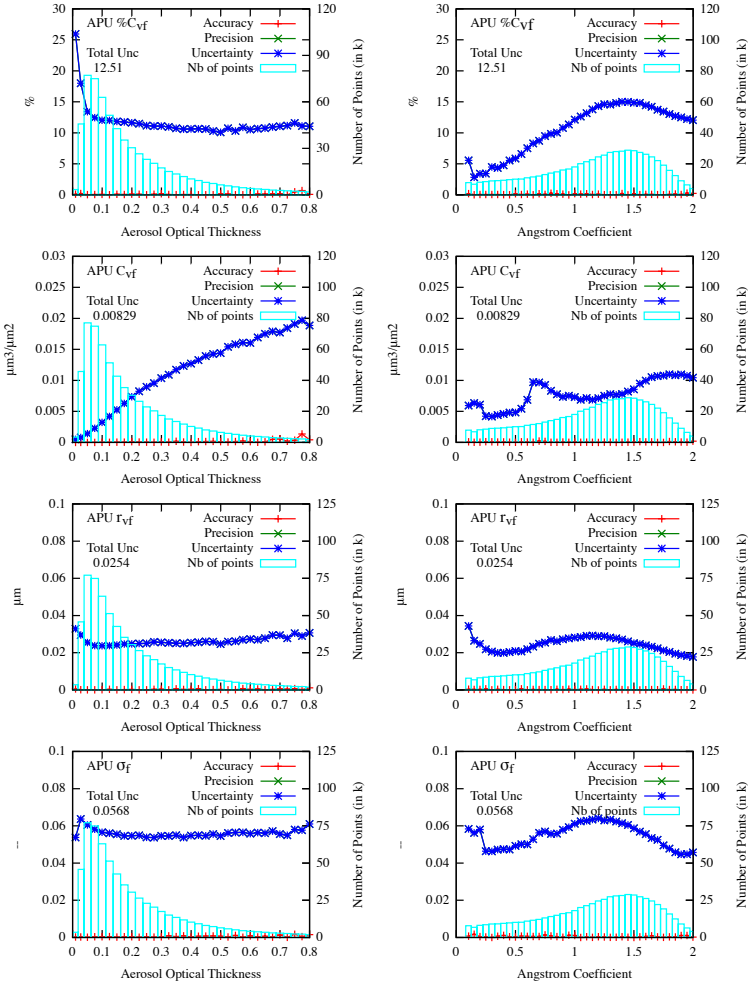

Figure A1. APU for each microphysical parameter (fine mode of the size-distribution only) retrieved from a Random Forrest approach versus the aerosol optical thickness at 440 nm (left) and the Angstrom coefficient between 440 and 870 nm (right).

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