



1 **Quantification of major particulate matter species from a single filter type using**
2 **infrared spectroscopy – Application to a large-scale monitoring network**

3 Bruno Debus¹, Andrew T. Weakley¹, Satoshi Takahama², Kathryn M. George^{1,3}, Bret Schichtel⁴,
4 Scott Copeland⁵, Anthony S. Wexler^{1,6}, Ann M. Dillner^{1*}

5 ¹ Air Quality Research Center, University of California, Davis, California, 95616, USA

6 ² ENAC/IIE, Swiss Federal Institute of Technology Lausanne (EPFL), Lausanne, Switzerland

7 ³ Monitoring and Laboratory Division, California Air Resources Board, Sacramento, CA 95811,
8 USA

9 ⁴ National Park Service, Cooperative Institute for Research in the Atmosphere, Colorado State
10 University, Fort Collins, CO 80523, USA

11 ⁵ Cooperative Institute for Research in the Atmosphere, Colorado State University, Fort Collins,
12 CO, 80523, USA

13 ⁶ Departments of Mechanical and Aerospace Engineering, Civil and Environmental Engineering,
14 and Land, Air and Water Resources, University of California, Davis, California, 85616, USA

15 **Correspondence to:* Ann M. Dillner (amdillner@ucdavis.edu)

16 **Abstract**

17 To enable chemical speciation, monitoring networks collect particulate matter (PM) on different
18 filter media, each subjected to one or more analytical techniques to quantify PM composition
19 present in the atmosphere. In this work, we propose an alternate approach that uses one filter
20 type (teflon or polytetrafluoroethylene, PTFE, commonly used for aerosol sampling) and one
21 analytical method, Fourier Transform Infrared (FT-IR) spectroscopy to measure almost all of the
22 major constituents in the aerosol. In the proposed method, measurements using the typical
23 multi-filter, multi-analytical techniques are retained at a limited number of sites and used as
24 calibration standards while sampling on PTFE and analysis by FT-IR is solely performed at the
25 remaining locations. This method takes advantage of the sensitivity on the mid-IR domain to
26 various organic and inorganic functional groups and offers a fast and inexpensive way of
27 exploring sample composition. As a proof of concept, multiple years of samples collected within
28 the Interagency Monitoring of PROtected Visual Environment network (IMPROVE) are explored
29 with the aim of retaining high quality predictions for a broad range of atmospheric compounds
30 including total mass, organic (OC), elemental (EC) and total (TC) carbon, sulfate, nitrate and



31 crustal elements. Findings suggest that models based on only 21 sites, covering spatial and
32 seasonal trends in atmospheric composition, are stable over a three year period within the
33 IMPROVE network with prediction accuracy ($R^2 > 0.9$, median bias less than 3% for most
34 constituents. Incorporating additional sites at low cost or partially replacing existing, more time
35 and cost intensive techniques are among the potential benefits of one-filter, one-method
36 approach.



37 1 Introduction

38 In the United States, filter-based chemical speciation of ambient aerosols has been in operation
39 for decades to quantify trends, assess transport and atmospheric transformation, identify
40 sources of air pollution, evaluate impacts of pollution regulations, assess impacts on visibility,
41 radiative forcing, human and ecosystem health and evaluate atmospheric and climatological
42 models. The two federally funded speciation networks, the Interagency Monitoring of PROtected
43 Visual Environments (IMPROVE) and the Chemical Speciation Network (CSN) collect 24-hour filter
44 samples using three filter media: polytetrafluoroethylene for analysis by gravimetry, hybrid
45 integrating plate and sphere (HIPS), and x-ray fluorescence (XRF), quartz for thermal optical
46 reflectance (TOR) and nylon for ion chromatography. Over the decades of operation, the
47 analytical methods have evolved with efforts to maintain consistency in trends while also
48 adopting improved methodology and retiring obsolete equipment. Impacts of many of these
49 changes have been addressed in the literature (Hyslop et al., 2015, 2012; White et al., 2016;
50 Zhang et al., 2021; Chow et al., 2007a, 2015) and in data advisories posted on the IMPROVE
51 website (<http://vista.cira.colostate.edu/Improve/data-advisories/>).

52 In this paper, we explore the use of Fourier transform-infrared spectroscopy (FT-IR) to reproduce
53 most of the existing speciation data because the most components exhibit optical activity in the
54 mid-IR. The number and bands of organic compounds are numerous, but generally group
55 frequencies can be found above 1500 cm^{-1} and compound-specific spectral patterns (“fingerprint
56 region”) below this frequency; down to approximately 700 cm^{-1} (for example, (Weakley et al.,
57 2016; Mayo et al., 2004). Graphitic carbon displays peaks near 1600 cm^{-1} due to lattice defects
58 (Tuinstra and Koenig, 1970; Friedel and Carlson, 1971), displacement vibrations near 868 cm^{-1}
59 (Nemanich et al., 1977), and a broad, sloping absorbance between 4000 and 1500 cm^{-1} due to
60 the tail of the electronic transition more strongly observed in the UV (Parks et al., 2021). Inorganic
61 substances containing polyatomic ions such as sulfate, nitrate, and ammonium have strong
62 vibrational modes above 600 cm^{-1} (Mayo, 2004). Crystalline and amorphous geological minerals
63 in the form of oxides (which include hydroxides and oxyhydroxides) have distinct internal
64 vibrational modes influenced by the electronegativity of the metal to which the oxygen is bonded
65 (Busca and Resini, 2006; Chukanov and Chervonnyi, 2016; Margenot et al., 2017).

66 FT-IR spectra with partial least squares (PLS) calibrations have been shown to reproduce OC and
67 EC concentrations using organic and graphitic carbon absorption bands, respectively, at a limited
68 number of sites in the IMPROVE network (Dillner and Takahama, 2015a, b; Reggente et al., 2016),
69 CSN (Weakley et al., 2016, 2018a) and FRM (Weakley et al., 2018b). Takahama et al. (2019)
70 reviews these findings and the overall framework to be used for the two phases of such statistical
71 calibrations: model building (sample selection, spectral preparation, model generation, model
72 selection, model evaluation, and model understanding) and operation (error anticipation and
73 model updating). Inorganic ions and geological mineral absorption bands have been used for



74 chemical speciation of these substances using FT-IR in prior studies (e.g., Cunningham et al.,
75 1974; McClenny et al., 1985; Pollard et al., 1990; Bogard et al., 1982; Foster and Walker, 1984).

76 Organic absorption bands are useful for measuring OC but also provide spectral information
77 needed to add detailed knowledge of composition not currently measured in air quality
78 monitoring networks – such as organic matter (OM) and organic functional group composition –
79 which is the subject of other work (Reggente et al., 2019; Boris et al., 2019, 2021; Burki et al.,
80 2020). Such calibrations, also combined with factor analytic approaches, can provide source
81 characterization on par with more costly mass spectrometric techniques (Boris et al., 2021;
82 Yazdani et al., 2021a; Hawkins et al., 2010; Takahama et al., 2011; Liu et al., 2012; Corrigan et al.,
83 2013).

84 Although FT-IR shows promise for measuring many constituents in aerosol, it is not without its
85 challenges. One limitation is that not all PM constituents can be measured, or measured with
86 high sensitivity, from the FT-IR spectrum. For instance, NaCl and MgCl₂ do not have IR-active
87 substituents. While a multitude of spectral signatures associated with mineral dust arise from
88 their constituent bonds – e.g., the metal-oxygen bonds in oxides (the oxide form is explicitly
89 assumed in estimating dust mass concentrations from elemental composition for the IMPROVE
90 network), some must be predicted from correlation with other constituents (e.g., some forms of
91 iron) if IR-activity is lacking. Other substances are IR-active but have weak responses, such as
92 graphitic carbon (Niyogi et al., 2006; Parks et al., 2021). The absorption and scattering by the
93 PTFE filter also pose challenges for quantitative analysis. The PTFE-based material changes over
94 time due to change in manufacturer or manufacturing process, and is difficult to fully characterize
95 a priori or treat with simple blank subtraction techniques. PTFE absorption limits full access to
96 the range of spectroscopic information in the mid-IR, for instance in the region of carbon-oxygen
97 bonds that can lead to less than full recovery of OM mass. Additionally, scattering leads to
98 broadly-varying slope in the group frequency region. This scattering artifact is minimized by
99 baselining (Kuzmiakova et al., 2016) and using many standards that have a range of scattering
100 and absorption observed in the network (Debus et al., 2019), yet these techniques can still lead
101 to errors in quantification. Weakley et al., (2018b) demonstrated that calibrations built using one
102 brand of filter can be accurately extended to another brand of PTFE filter with numerically
103 marginal but statistically significant increase in method error (e.g., +2% error for $\alpha=0.05$).
104 However, these studies are insufficient to generalize findings to all types of sampling filters.

105 The goal of this work is to assess the capability of using FT-IR to measure the aerosol chemical
106 composition at all IMPROVE sites. FT-IR quickly and non-destructively collects information-rich
107 spectra from routinely collected PTFE filter samples. Ambient samples from strategically-
108 selected IMPROVE sites are used for calibration and reasonably mimic the composition, matrix
109 effects and substrates of the unknowns, all of which theoretically lead to accurate estimations of
110 concentrations. Using all samples from selected sites reduces maintenance, shipping, processing



111 and coordinating required to maintain intermittent quartz and nylon filter sampling at every site.
112 Sites are selected using data from 2015 and are used for calibrating samples from 2015-2017.
113 Samples from all other (non-calibration) IMPROVE sites are predicted and compared to routine
114 IMPROVE data to assess the quality of prediction. Aerosol components to be measured include
115 TC, OC, EC, inorganic ions, soil elements, particulate mass, and light absorption.

116 2 Methods

117 2.1 Network data

118 IMPROVE samples collected every third day at all North American sites (Section S1) from January
119 2015 through December 2017 are included in this study. Fine particulate matter (aerodynamic
120 diameter less than 2.5 micrometers) is deposited on 25 mm diameter filters
121 polytetrafluoroethylene (PTFE) and quartz filters by sampling air at a nominal flowrate of 22.8
122 liters per minute from midnight to midnight local time. Parallel 37 mm nylon filters are collected
123 at the same flow flowrate. PTFE filters are analyzed by multiple instruments and archived for
124 future analysis. Nylon filters and a portion of each quartz filter undergoes destructive analysis
125 and a remaining part of the quartz filters are retained for archive.

126 Over the period covered in this study, two different TOR instruments were employed to measure
127 OC, EC and TC. Quartz filters sampled prior to 2016 were analyzed on a DRI Model 2001 thermal
128 optical carbon analyzers (Chow et al., 1993) while filters collected beginning in January of 2016
129 were analyzed on a DRI Model 2015 multi-wavelength thermal optical carbon instrument (Magee
130 Scientific – Berkley, USA)(Chow et al., 2015). Both instruments use the IMPROVE_A protocol
131 (Chow et al., 2007b), which outlines the temperature step, gaseous environment in the
132 instrument and that reflectance is used to define the split point between OC and EC. To correct
133 for gas phase adsorption onto the quartz filter, the monthly median field blank OC concentration
134 is subtracted from each OC measurement during that sample month. Carbon concentrations are
135 reported in $\mu\text{g}/\text{m}^3$.

136 An in house Hybrid Integrating Plate and Sphere (HIPS) system evaluates light absorption from
137 the PTFE filters in the IMPROVE network (White et al., 2016). In this work, the measured
138 absorption coefficient (*Fabs*) is converted into a TOR EC equivalent concentration assuming a
139 *Fabs* / EC ratio of $10 \text{ m}^2\text{g}^{-1}$ (Malm et al., 1994). The resulting value, referred to as HIPS Black
140 Carbon (HIPS BC), is used as part of a quality control procedure to evaluate potential outliers in
141 TOR EC measurements.

142 Data from gravimetry and X-ray fluorescence (XRF) analysis obtained from PTFE filters and ion
143 chromatography from the nylon filters are also used in this study. Additional information on
144 routine IMPROVE methods can be found on the IMPROVE website
145 (<http://vista.cira.colostate.edu/Improve/>). IMPROVE data are available online at
146 (<http://views.cira.colostate.edu/fed/>).



147 2.2 Outlier removal

148 Data were screened for outliers to eliminate their influence on the results. Out of the ~61,500
149 total number of samples in the three-year period, fewer than 800 were excluded from the
150 analysis due sampling issues or missing TOR, XRF or IC data. In addition, 65 samples collected at
151 the Wheeler Peak Wilderness (New Mexico) site between November 2015 and April 2016 were
152 excluded due to an EC contamination caused by a diesel-powered ski lift.

153 Potential outliers in TOR measurements were investigated by regressing TOR EC against HIPS BC
154 concentrations. Samples with differences exceeding a predefined threshold value ($0.68 \mu\text{g}/\text{m}^3$)
155 were flagged as potential outliers (section S2). The status of these samples was confirmed by
156 building separate TOR EC and HIPS BC calibrations. The poor agreement between TOR EC and FT-
157 IR EC concentrations contrasts with the nearly 1:1 relationship HIPS BC and FT-IR BC predicted
158 values indicating that TOR EC concentrations were likely compromised (Section S2). For the
159 period considered in this study, 112 samples with faulty TOR EC values were identified and
160 excluded from further analysis. The number of valid sample spectra retained in this study is
161 61,462.

162 2.3 Fourier-transform infrared (FT-IR) analyses

163 Since 2015, all PTFE in the IMPROVE network have been analyzed by infrared spectroscopy for
164 research and evaluation purposes. FT-IR measurement occurs after gravimetric analysis and prior
165 to XRF and HIPS to prevent possible loss of volatile species under the mild vacuum in XRF. Three
166 FT-IR spectrometers including one Tensor 27 and two Tensor 2 instruments (Bruker Optics,
167 Billerica, MA) equipped with a pre-aligned mid-IR source and a liquid nitrogen-cooled wide-band
168 mercury cadmium telluride (MCT) detector were used for spectra acquisition in the range $4000 -$
169 420 cm^{-1} by averaging 512 scans at a nominal resolution of 4 cm^{-1} . The single beam signal
170 associated with each PTFE filter was converted to an absorbance spectrum using the most recent
171 zero reference signal, updated hourly.

172 Previously, it was determined that calibration transfer between multiple FT-IR instruments is not
173 required as long as their spectral response is carefully matched by controlling a set of key
174 environmental and instrumental parameters (Debus et al., 2019). Briefly, each mercury cadmium
175 telluride (MCT) detector is connected to an automatic liquid nitrogen micro dosing system
176 (NORHOF, Ede, Netherlands) designed to improve signal stability and maintain a high signal to
177 noise ratio. The repeatability and reproducibility of the filter position relative to the IR beam is
178 controlled via a house-built sample chamber ($4.0 \times 5.1 \times 4.5 \text{ cm}$) mounted inside the instrument
179 sample compartment. Details regarding the chamber design have been published elsewhere
180 (Debus et al., 2019). Finally, the contribution of water vapor and carbon dioxide to the signal was
181 minimized by continuously purging both the sample chamber and the optical bench with a VCD
182 Series CO_2 adsorber / dryer system (PureGas LLC, Broomfield, CO). For each sample, the
183 acquisition procedure involves a 4 minutes purge period followed by a spectrum collection lasting



184 about 1 minute. An in-house macro interfaced to the OPUS software (Bruker Optics, Billerica,
185 MA) controls each step. PTFE filters were measured in transmission mode without sample
186 preparation. No interpolated data (from zero-filling) are included in the final raw spectra.
187 Collected spectra are subjected to weekly quality control procedures detailed in (Debus et al.,
188 2019). Duplicate and replicate measurements were also performed to evaluate instrument
189 stability and found to be within +/- 10%.

190 2.4 Multivariate Calibration Model - Partial Least Squares (PLS) Regression

191 While the presence of certain category of atmospheric compounds can be identified qualitatively
192 from an FT-IR spectrum, an accurate quantification of their concentration requires calibration.
193 PLS is a commonly used algorithm to relate a multi-wavenumber measurement to any particular
194 sample properties such as concentration (Wold et al., 2001). In brief, PLS maximizes the co-
195 variance between a set of response variables (species measurements) and a reference
196 measurement (FT-IR spectra) from which equivalent predicted values are desired. In so doing,
197 the optimal combination of response variables best describing the reference measurement is
198 identified and the selected features are used to build a multivariate calibration. With all least-
199 squares calibration methodologies, concentration-dependent biases in residuals that are
200 determined by the quality of fit (R^2) and dynamic range of the data are expected due to the nature
201 of least-squares estimation (Besalú et al., 2006; Draper and Smith, 1998, pp. 63-64,173,638). For
202 further discussion of these biases, see Section S1.

203 The applicability of PLS to quantify carbonaceous aerosol species (Reggente et al., 2016; Weakley
204 et al., 2016, 2018a) or functional groups (Boris et al., 2019; Ruthenburg et al., 2014) collected on
205 PTFE filters in various monitoring networks and field campaigns has been successfully
206 demonstrated. A complete review of the implementation of PLSR calibration in the framework
207 of atmospheric particulate matter characterization has been recently published (Takahama et al.,
208 2019).

209 To evaluate model performance, FT-IR predicted concentrations were regressed against their
210 reference measurement to quantify residuals and a series of metrics. Reported figures of merit
211 include the coefficient of determination (R^2), bias, error and the method detection limit (MDL).
212 Residuals are defined as the difference between predicted and reference concentrations, bias
213 corresponds to the median residual while error is the median absolute residual. To facilitate inter-
214 model comparison, relative performance metrics were calculated by normalizing the values by
215 their reference value. FT-IR PLSR calibration MDL was estimated from field blank predicted
216 concentrations as the 95th percentile minus the median residuals, as is done for other species in
217 the IMPROVE network [http://vista.cira.colostate.edu/improve/wp-](http://vista.cira.colostate.edu/improve/wp-content/uploads/2021/07/IMPROVE-SOP-351_Data-Processing-and-Validation_2021_final.pdf)
218 [content/uploads/2021/07/IMPROVE-SOP-351_Data-Processing-and-Validation_2021_final.pdf](http://vista.cira.colostate.edu/improve/wp-content/uploads/2021/07/IMPROVE-SOP-351_Data-Processing-and-Validation_2021_final.pdf).
219 Performance is reported for all samples together regardless if the samples were included in the
220 calibration. This enables comparison between models with different samples used for calibration.



221 For further insight into model prediction accuracy, the distribution in FT-IR residuals is
222 qualitatively compared with residuals from collocated measurements. Collocated quartz filters
223 are collected at the Everglades (FL), Hercules-Glades (MO), Medicine Lake (MT) and Phoenix (AZ)
224 sites. Similarly, collocated Teflon filters are sampled at Mesa Verde (CO), Proctor Maple Research
225 Facility (VT), Saint Marks National Wildlife Refuge (FL), Yosemite (CA) and Phoenix (AZ) sites while
226 collocated nylon filters are featured at the Phoenix (AZ), Frostburg Reservoir (MD), Mammoth
227 Cave (KY) and San Gabriel (CA) sites.

228 Data handling and analysis was performed in Matlab R2018a (The MathWorks, Inc, Natick, MA,
229 United States) using the statistics and signal processing toolboxes. PLS was computed via the
230 libPLS Matlab package (v1.9) (Li et al., 2018).

231 2.5 FT-IR Calibrations for Predicting PM Composition

232 This section presents the design of calibrations for quantifying the concentration of major
233 atmospheric species by taking advantage of the composition-based information embedded
234 within an FT-IR spectrum. In practice, spectra are calibrated against reference measurements
235 from TOR, XRF, IC, HIPS and gravimetric analysis with the aim of predicting concentrations of
236 atmospheric constituents using only spectra of PTFE filters as input.

237 A multilevel model (Snijders and Bosker, 2011; Takahama et al., 2019) is proposed in which
238 dedicated calibration models for subgroups of samples are constructed, and applied according
239 to a predetermined selection criterion for each sample. This model considers two subgroups: i)
240 samples determined to be dominated by biomass burning, which are calibrated with similar
241 samples, and ii) the remaining samples, which are calibrated with samples from a limited number
242 of sites. To establish baseline performance metrics for comparison, a “Global model” in which a
243 single calibration (for each species) is constructed from all samples considered together and
244 described in Section S1 (Supplement).

245 The first step in the development of the Multilevel model consists of screening for biomass
246 burning samples. These samples are removed from consideration during the site selection
247 process. A simple detection method combining estimates of key functional group spectral peak
248 areas and spectral dissimilarity metrics were used to segregate biomass burning samples from all
249 other samples. Next, a Gaussian Mixture Model (GMM) was applied to the spectra of all non-
250 biomass burning samples. The GMM exploits the specificity of the infrared signal for organic and
251 inorganic species. The GMM was implemented with the aim of clustering the non-biomass
252 burning FT-IR spectra into groups sharing similar spectral features (Section 2.5.2). Those groups
253 were later used as part of the methodology for selecting sites with representative atmospheric
254 composition. Spectra from the year 2015 were used as a benchmark to validate the biomass
255 burning detection strategy, build the GMM and establish the list of representative sites where



256 multi-filter collection/multi-analyses should be retained (section 2.5.2). The identified biomass
257 burning samples are used to build a calibration for biomass burning samples (Section 2.5.1).

258 2.5.1 Biomass burning model

259 FT-IR spectra were used to estimate functional group areas and calculate spectral dissimilarities
260 metrics to segregate biomass burning samples from all other samples. Although this paper
261 focuses on using FT-IR to measure the major aerosol components in routine speciated aerosol
262 monitoring networks, FT-IR is more frequently used to measure organic functional groups (e.g.
263 (Russell et al., 2011; Ruthenburg et al., 2014; Boris et al., 2019). Specific regions in the IR spectra
264 correspond to specific functional groups. Peak areas, calculated from baseline corrected spectra
265 (see Section S3 for baseline procedure), for carbonyl, OH and CH were used rather than functional
266 group calibrations for simplicity. Because the relative functional group peak area tends to
267 increase significantly as the cumulative peak area decreases, typically for low mass deposition
268 samples, an estimate of spectral dissimilarities, the squared Mahalanobis distance (D_i^2), for each
269 site is also considered to minimize false detection. The Mahalanobis distance (Mahalanobis,
270 1936; Cios et al., 1998) is a measure of the spectral dissimilarity between a given spectrum at a
271 site and the mean spectrum at the site. Taking advantage of D_i^2 and relative functional group
272 areas, a set of criteria were established from observations at known wildfire sites during wildfire
273 season (O'Dell et al., 2019). First, samples collected under heavy smoke conditions whose spectra
274 fulfill $C-H \geq 2\%$, $C=O \geq 15\%$ and $D_i^2 \geq 3 \overline{D}^2$ were detected (Section S3). This group of spectra
275 tend to have large D_i^2 values and, consequently, the $3 \overline{D}^2$ threshold often excludes samples with
276 low to moderate biomass burning contributions. For a more inclusive detection, spectra from the
277 first group were removed from consideration, the D_i^2 values are updated for each sample and
278 the plots were regenerated. The cut-off value for the relative carbonyl functional group area was
279 lowered to 8% while other parameters were not changed. Spectra identified by the first and
280 second rounds are considered biomass burning samples. This procedure is performed for each
281 site and for each year of sample collection (Section S.3)

282 Recent work has shown that smoke samples may be identified using techniques such as cluster
283 analysis and labeling (Burki et al., 2020) similar to the GMM used here and through detection of
284 molecular markers – levoglucosan and lignin – or peak profiles in FT-IR spectra (Yazdani et al.,
285 2021a, b). For the large data set in this work (~20,000 samples in 2015), cluster analysis
286 resulted in multiple clusters that could be associated with smoke-impacted samples likely due
287 to the variations in fuel, oxidation conditions, and contributions from other sources. Therefore,
288 for this work we selected a single group of smoke-impacted samples based on specific organic
289 features known to be present in FT-IR spectra. While the criteria for smoke-impact labeling can
290 be defined differently according to each intended purpose, the method presented here is
291 demonstrated to sufficiently partition the samples for building accurate submodels to predict
292 concentrations of PM constituents.



293 While ions and crustal species are not necessary correlated with wildfire emissions, the Biomass
294 Burning sub-model for accounts for interferences that are necessary to track in order to maintain
295 high prediction accuracy for samples collected on smoky days.

296 2.5.2 Limited Sites Model

297 To assess major PM_{2.5} composition regimes in the network and to identify representative sites to
298 use as calibration standards in the Limited Sites model, screening of all FT-IR spectra (except
299 samples identified as biomass burning samples) across all locations and seasons was performed
300 by building a Gaussian Mixture Model (GMM) (Bilmes, J. A., 1998; Hastie, T et al., 2009). The basic
301 idea behind GMM is to group FT-IR spectra into clusters of similar spectral shape using a
302 probabilistic approach describing the likelihood that any given spectrum belongs to a particular
303 class. To minimize the concentration dependence and emphasize composition, raw spectra were
304 transformed to second derivative spectra using a 2nd order, 21 point, Savitzky-Golay filter
305 (Savitzky and Golay, 1964), differenced with filter blank spectra and divided by their respective
306 Euclidean norm (Bro and Smilde, 2003). Additional details about the GMM pre-processing and
307 implementation as well as cluster interpretation are provided in Section S4.

308 After classification, a single site per cluster was selected to represent the atmospheric
309 composition captured in that cluster. For any given cluster, the retained location was defined as
310 the site with the largest number of classified spectra with the highest probabilities of belonging
311 to that cluster. To prevent misleading site selection and enhance spatial coverage, the following
312 set of decision rules were used: *i)* if the same site is representative of two clusters, it is ascribed
313 to the cluster with the largest number of classified spectra from that site, *ii)* if none of the
314 retained sites accounts for a given spatial region or known source type in the network, an
315 additional site with the highest number of classified spectra is selected from a nearby cluster,
316 and *iii)* only sites under continuous operation between 2015 and 2017 are eligible for selection.
317 Criteria *ii)* was invoked once to add a site in the Midwest to improve spatial coverage in that
318 region and to capture prescribed fire emissions in Kansas. All non-biomass burning samples from
319 selected sites were used as FT-IR calibration standards for all species and all non-biomass burning
320 samples are predicted with these models. Once established, the selected sites are not re-
321 evaluated but instead were used in all subsequent years as would occur in practice.

322 2.5.3 Application of Multilevel Model

323 Multilevel model is the combined FT-IR predicted concentrations from the Limited sites and
324 Biomass burning models. Multilevel modeling will be discussed in the context of carbonaceous
325 aerosols before extending the modeling to other atmospheric constituents with detectable
326 infrared signatures. In addition to OC and EC, species evaluated for FT-IR prediction include PM_{2.5}
327 mass, soil elements (silica, aluminum, calcium, titanium, iron), anions (sulfate, nitrate) and HIPS
328 BC. Next, the years 2016 and 2017 will be examined to assess the long-term stability of the



329 proposed Multilevel strategy by screening for smoke samples and re-calibrating each year using
330 the sites selected using 2015 data.

331 3 Results and discussion

332 In the following sections, the quality of FT-IR based calibrations for quantifying aerosol
333 composition across continental US and their long-term applicability to large speciation
334 monitoring networks will be assessed. Section 3.1 describes the selected calibration samples for
335 the Biomass Burning and Limited Sites models. In Section 3.2, Biomass Burning and Limited
336 model performance will be briefly reviewed before exploring the Multilevel FT-IR predictions for
337 all samples. Initially focused on carbonaceous species on PTFE samples collected in 2015, FT-IR
338 predictions will be extended to other atmospheric constituents and years.

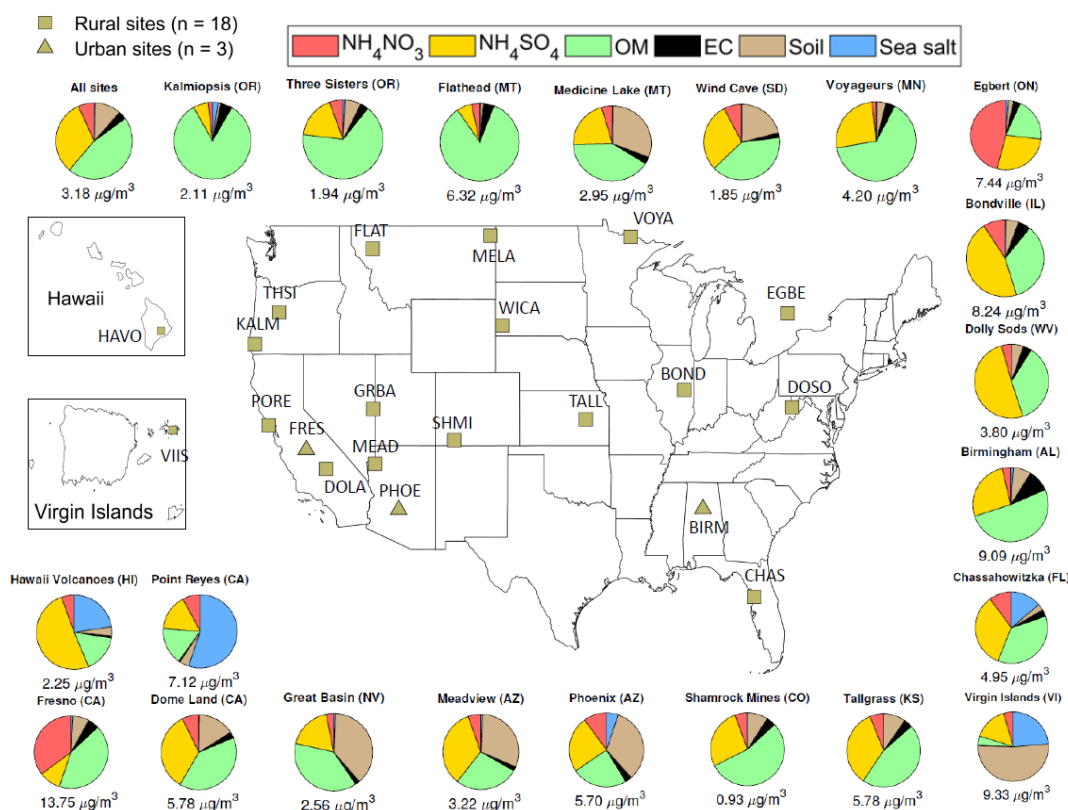
339 3.1 Multilevel modeling – Calibration sample selection

340 3.1.1 Biomass burning sample selection

341 Using the methods described above, 492 samples impacted by biomass burning emissions were
342 identified in 2015 (2.5 % of the network), 288 samples in 2016 (1.5 %), and 817 samples in 2017
343 (3.7 %). The mean OC concentration of the biomass burning samples range was 5.6 – 8.3 $\mu\text{g}/\text{m}^3$
344 with maximum concentrations extending from 44.5 to 97 $\mu\text{g}/\text{m}^3$ over the three year period.
345 Similarly, per year, the mean EC concentration varies between 0.61 – 0.9 $\mu\text{g}/\text{m}^3$ with maximums
346 up to 2.9 – 3.9 $\mu\text{g}/\text{m}^3$. Mean OC/EC ratios are larger than 7, in agreement with past literature
347 (Schichtel et al., 2008; Sorooshian et al., 2011). Analysis of the detected samples shows reliable
348 spatial and seasonal distributions, consistent with biomass burning emissions predominantly in
349 summer and fall across the Pacific North West and Northwestern US (Section S3). Two-thirds of
350 the identified samples were selected (Section S5) as calibration standards for the calibration and
351 resulting model was applied to the remaining third of the smoke impacted samples.

352 3.1.2 Limited Sites model – clusters and retained sites

353 Figure 1 shows the spatial distribution of the 21 sites selected for Limited sites model. From a
354 spatial standpoint, retained sites appear reasonably scattered across the network including
355 Hawaii and the Virgin Islands. Clusters are represented by a distribution of urban and rural sites.
356 One urban cluster is represented by Fresno and contains mostly urban samples from Fresno and
357 Phoenix. All other clusters contain mostly rural and pristine sites. However, two other urban sites
358 were retained, Phoenix and Birmingham. The Phoenix cluster contains samples from the
359 southwest in the spring. The Birmingham site along with the Tallgrass site represent a non-
360 western cluster in the spring and summer.



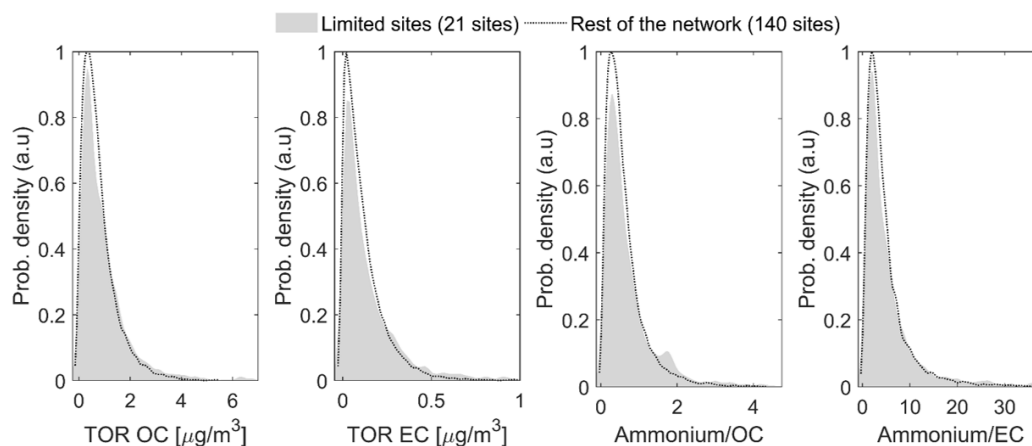
363 **Figure 1:** Spatial distribution, median $\text{PM}_{2.5}$ concentration and composition of the 21
 364 representative sites. Sites are identified by the four letter site code which is the first four letters
 365 of a single word site name (Fresno = FRES) or the first two letters of the first and second word for
 366 two word site names (Dome Land = DOLA). The composition is obtained from IMPROVE
 367 measurements and the IMPROVE reconstructed fine mass equation. The top left pie chart
 368 representing the median $\text{PM}_{2.5}$ composition across all sites and samples is given for comparison.

369 The clusters are also seasonally distributed (Section S6): five clusters are dominated by fall -
 370 winter samples, ten clusters containing summer samples (along with varying number of spring
 371 and fall samples), two clusters are predominately spring and one is spring - fall. Three clusters
 372 have little seasonality.

373 Because FT-IR spectra are clustered based on composition, the first step in assessing the
 374 representativeness of the 21 sites is to compare the concentration ranges. For this purpose,



375 distributions in TOR OC and EC concentrations excluding biomass burning samples are compared
376 for the 21 sites used for calibration and the 140 remaining sites. In Fig. 2, the two probability
377 density functions are very similar for both OC and EC despite large differences in sample
378 populations (2572 and 16,543, respectively). In addition to matching the range of carbonaceous
379 concentrations observed in the rest of the network, the presence of species interfering with
380 organic functional groups should also be accounted for by the calibration. Because ammonium
381 absorptions overlap with carbonaceous absorptions, situations where ammonium to OC and
382 ammonium to EC ratios are different between calibration samples and samples to be predicted
383 were associated with additional sources of bias and error (Dillner and Takahama, 2015a, b).
384 Although not measured in IMPROVE, ammonium concentration is approximated from nitrate and
385 sulfate assuming both species are fully neutralized. The corresponding probability distribution in
386 Fig. 2 confirms the equivalence between the ranges of ammonium/OC and ammonium/EC
387 concentrations spanned by the Limited sites samples and the overall network. In section S6, Fig.
388 S6-3 shows reasonable agreement between the selected sites and the rest of the network for
389 PM_{2.5} mass, ions, elements and HIPS BC. Together, these results suggest the list of 21 sites is a
390 suitable representation of network variations in OC and EC and their relative proportion to
391 ammonium, and for all other predicted constituents.



393 **Figure 2:** Comparison of probability density function for OC, EC and ammonium concentrations
394 in 2015 between the 21 sites retained for Limited calibration and the rest of the network.

395 The spatial and seasonal as well as the urban and rural diversity supports the compositional
396 diversity of the selected sites as shown in Fig. 1. The three urban sites have distinct
397 characteristics. At the Fresno, CA site, the composition is dominated by nitrate (35 %) and organic
398 matter (42 %) with an autumn – winter pattern consistent with agriculture and residential wood
399 burning activities (Ngo et al., 2010) as well as with the formation of secondary aerosols during



400 stagnation events and a low inversion layer (Watson and Chow, 2002). Phoenix, AZ site features
401 a strong soil component (33%) associated with spring dust storms and windblown dust and equal
402 proportions of ammonium sulfate (25 %) and OM (24 %) occurring mostly in spring and summer.
403 The ammonium sulfate and organic matter has been attributed to regional power production and
404 traffic (Brown et al. 2007). In contrast, Birmingham samples show little seasonal trend with
405 elevated OM (52 %) and EC (10 %) fractions originating from various combustion processes
406 including vehicle exhaust, biomass burning and biogenic secondary organic aerosols (Blanchard
407 et al. 2016). The other dominant species at this site is ammonium sulfate (26 %), characteristic of
408 coal burning and industrial activities in the East (Watson et al. 2015).

409 Among rural sites, four noticeable patterns in $PM_{2.5}$ composition are distinguishable. The first
410 corresponds to OM fractions accounting for more than two-thirds of the filter mass. High OM
411 samples are encountered at four locations in Northwestern US, namely the Kalmiopsis (OR),
412 Three Sisters (OR), Flathead (MT) and Voyageurs (MT) sites. Samples from Voyageurs (MN) and
413 Flathead (MT) sites are from Summer-Fall and present elevated median $PM_{2.5}$ concentrations
414 ($4.20 \mu\text{g}/\text{m}^3$ and $6.32 \mu\text{g}/\text{m}^3$, respectively) and very large percentage of OM consistent with
415 biomass burning emissions. Kalmiopsis (OR) and Three Sisters (OR) samples have a lower and
416 nearly identical median $PM_{2.5}$ concentration ($\approx 2 \mu\text{g}/\text{m}^3$) but differ in their monthly distribution
417 with the former displaying more winter samples than any other season whereas the later shows
418 little seasonality.

419 The second type of sites have high OM and sulfate concentrations. Both Shamrock Mines (CO)
420 and Tallgrass Prairie (KS) sites have larger OM than sulfate content. However, the Colorado site
421 has more autumn – winter samples, represents samples in the Rockies and Alaska and an overall
422 small median $PM_{2.5}$ concentration ($< 1 \mu\text{g}/\text{m}^3$). The Kansas site has a majority of spring samples,
423 representing non-western samples and has a significantly larger $PM_{2.5}$ concentration ($\approx 6 \mu\text{g}/\text{m}^3$)
424 that is attributed to prescribed burning near the Tallgrass site (Whitehill et al. 2019). Other sites
425 have higher median sulfate concentrations ($\sim 50\%$) than OM concentrations ($\sim 40\%$) such as Dolly
426 Sods (WV) and Bondville (IL). The monthly sample distribution indicates seasonal influences:
427 Bondville (IL) samples are mostly from the summer and the concentrations are relatively high
428 while the Dolly Sods (WV) site samples are mostly not in the summer with lower concentrations.
429 Because the spectra were normalized to minimize influence of concentration, these two clusters
430 likely have different organic composition even though the relative amount of OM is the same.
431 Finally, situations where sulfate and OM are present in equal proportions ($\approx 36\%$) are reported
432 at the Dome Land (CA) and Chassahowitzka (FL) pristine sites mainly featuring spring – summer
433 and winter samples, respectively.

434 A third group of noteworthy $PM_{2.5}$ compositions at rural sites contain a large fraction of ($> 20\%$
435 of the total mass). The Virgin Islands (VI) site presents the highest soil fraction across the network
436 52% of the total $PM_{2.5}$ mass, mostly originating from long-range Sahara soil dust transport



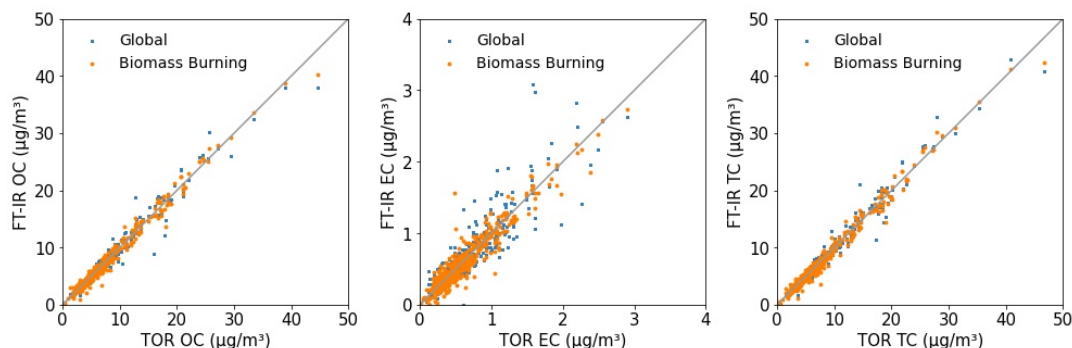
437 (Holmes and Miller 2004). In addition to sulfate and OM, elevated soil contributions are observed
438 for the Wind Cave (SD), Meadview (AZ), Medicine Lake (MT), and Great Basin (NV) sites with soil
439 content between 20 and 40%. Although the seasonality is somewhat different between these
440 sites, they all have many samples from the spring suggesting the dust is due at least in part to
441 spring dust storms and may also contain resuspended road dust and more localized dust sources.

442 A fourth and final distinct category of PM_{2.5} composition includes a collection of sites with unique
443 local atmospheric pollution sources, specific to those locations. The Hawaii Volcanoes (HI) site
444 where sulfur emitted as part as the volcanic activity, contains 51% sulfate along with sea salt (23
445 %). Another location with unique composition is the Point Reyes (CA) site where the sea salt
446 contribution reaches 55% of the median filter mass for the clustered samples, larger than any
447 other marine site in the network. Finally, the Egbert (ON) Canadian site, representing the upper
448 Midwest in winter is dominated by nitrate (46 %), sulfate (27 %) and OM (20 %).

449 As described above, the 21 sites retained for the Limited sites sub-calibration present seasonal,
450 regional and compositional features consistent with known or expected trends in PM_{2.5} across
451 the network. The median PM_{2.5} mass at those locations covers a broad range of concentrations
452 ranging from 0.93 µg/m³ to 13.75 µg/m³ and includes both urban and rural sites. Capturing the
453 large variability in PM_{2.5} composition and concentration is essential to ensure the proposed site
454 list is a representative subset of the parent network. However, it should be mentioned that the
455 proposed site list is not unique but is one of the many feasible solutions since sites whose samples
456 clustered together in the GMM are likely inter-exchangeable.

457 3.2 Evaluation of Biomass Burning Model

458 Prior to describing the overall results from the Multilevel model, the Biomass Burning model is
459 evaluated to determine if the biomass burning model improves predictions for those samples.
460 To evaluate the quality of the biomass burning model, the predictions are compared to a global
461 model (section S1) which contains a few samples from all 160 sites which are mostly non-smoke
462 samples but does contain a few smoke samples. Visual inspection of Fig. 3 suggests the
463 equivalence of the biomass burning models to the global model at the lower end of the
464 concentration range. However, improvement in prediction accuracy can be claimed at high
465 concentrations for the Biomass Burning model. The gain in model performance is subtle for OC
466 and TC; however, for EC, predictions benefit from having a dedicated calibration for samples
467 impacted by wildfire emissions, with an increase in R² from 0.747 to 0.902 (Section S7).



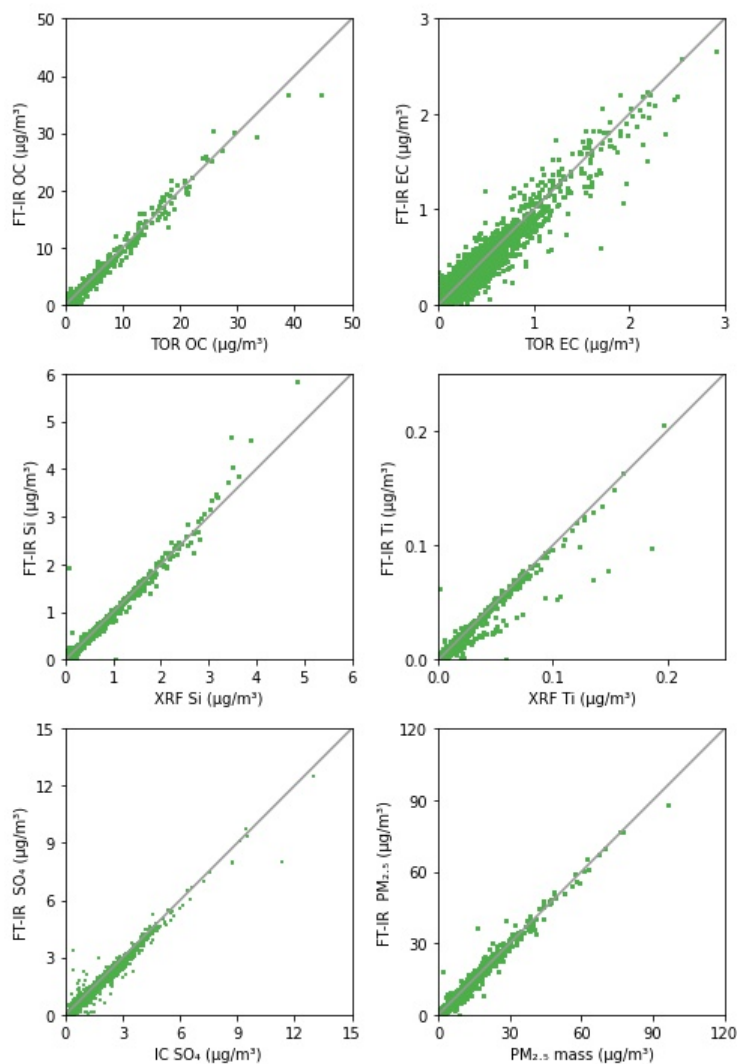
469 **Figure 3:** Inter-model OC (left), EC (middle), TC (right) comparison between global (section S1)
470 and Biomass Burning predicted concentrations for the 492 samples classified as biomass burning
471 in 2015. EC prediction, in particular, benefit from having a dedicated Biomass Burning calibration
472 model.

473 Therefore, we retain the biomass burning model as part of the multilevel model and present
474 the results for the Multilevel model below.

475 3.3 Multilevel modeling – Performance evaluation

476 3.3.1 Carbonaceous aerosol predictions

477 Figure 4 shows the correspondence between FT-IR Multilevel concentrations for OC and EC and
478 TOR measurements for 2015 (plot for TC can be found in Section S9) and Table 1 lists the
479 prediction metrics for all 3 carbonaceous components. The visual agreement between FT-IR
480 and the reference measurements of OC and EC is high but EC shows higher scatter than the
481 other measurements. Table 1 indicates that FT-IR OC and TC has higher prediction quality than
482 EC but both perform satisfactorily. FT-IR OC and TC error is on par with TOR precisions (Table 1)
483 indicating that FT-IR does not add significant additional error to the measurement. FT-IR EC
484 predictions, however, have higher error than TOR precision. With respect to reference (TOR)
485 measurements, concentration-dependent biases in residuals that are determined by the quality
486 of fit (R^2) and dynamic range of the data are expected due to the nature of least-squares
487 estimation (Besalú et al., 2006; Draper and Smith, 1998). For bias defined as FT-IR predictions
488 minus the reference (TOR) measurement, least-squares estimator causes an apparent linear
489 bias which is positive at the low end of the concentration range and negative at the high end of
490 the concentration range. (see Section S8 for further discussion). The satisfactory agreement
491 between FT-IR and TOR concentrations as well as the equivalence agreements using the global
492 model (Section S1) support the validity of the proposed Multilevel modeling in the context of
493 carbonaceous aerosols prediction from PTFE filters in large speciation networks.



495 **Figure 4:** Comparison of predicted FT-IR OC, EC, Si, Ti, SO_4 and mass concentrations using the
496 Multilevel model against their reference measurements. Each subplot contains all 19,608
497 samples collected in the year 2015.

498 **Table 1:** Summary of Multilevel model performance for IR-active atmospheric constituents for
499 19,608 spectra analyzed by FT-IR in the year 2015.



Species	R ²	Bias [µg/m ³]	Bias (%)	Error [µg/m ³]	Error (%)	Reference Data Error ¹ (%)	MDL [µg/m ³]	< MDL (%)
<i>OC</i>	0.983	0.01	1.6	0.08	12	10	0.06	0.9
<i>EC</i>	0.912	0	1.7	0.02	30	15	0.04	20.7
<i>TC</i>	0.984	0.01	1.2	0.08	12	11	0.07	1.3
<i>BC</i>	0.92	0	-0.3	0.03	23	---	0.04	19.3
<i>Si</i>	0.983	0	2.2	0.01	11	13	0.01	9.7
<i>Al</i>	0.985	0	2.2	0	12	10	0	4.7
<i>Ca</i>	0.979	0	1.1	0	13	9	0	6.9
<i>Ti</i>	0.941	0	2.7	0	21	16	0	14.9
<i>Fe</i>	0.95	0	1.1	0	25	8	0.01	19
<i>SO₄</i>	0.983	0	0.1	0.03	6	2	0.03	0.9
<i>NO₃</i>	0.927	0.02	15.3	0.07	54	8	0.07	21.8
<i>PM_{2.5} Mass</i>	0.985	0.03	1	0.18	6	6	0.25	1.1

500 ¹Median relative error for TOR, XRF, IC and gravimetric analysis. OC, EC and TC median relative error estimated
 501 from collocated sampling as measurement error/uncertainty is not reported by IMPROVE for this components.
 502 For all other components, the normalized error was calculated as the uncertainty divided by the concentration
 503 prior to selecting the median. BC is not reported by IMPROVE so measurement error is not estimated.

504 In addition to OC, EC and TC, light absorption which is predominantly due to black carbon, is also
 505 a measure of one fraction of the carbonaceous aerosol. FT-IR calibrations are found to be
 506 adequate for replicating HIPS BC measurements (Section S9). As expected, the corresponding
 507 model is similar in performance to its EC with R² and relative error of 0.920 and 23.3 %,
 508 respectively (Table 1). FT-IR BC residuals have a broader interquartile range than in the HIPS BC
 509 collocated data (Section S9). We attribute this effect to a difference in signal to noise ratio and
 510 sensitivity to chemical interferents between the two analytical techniques. While HIPS exploits
 511 the strong absorption properties of refractory carbon in the visible domain, the weak absorptivity
 512 of EC in the mid-infrared domain (Niyogi et al., 2006) and the presence of overlapping species
 513 makes the quantification less accurate.

514 3.3.2 Elemental oxide predictions

515 Taking advantage of known mineral absorbance bands in the mid-infrared range (Hahn et al.,
 516 2018; Madejová and Komadel, 2001; Senthil Kumar and Rajkumar, 2013) (Section S9), Multilevel
 517 calibrations for soil elements were evaluated for the five crustal elements commonly used to
 518 estimate soil: silicon, aluminum, calcium, titanium, and iron (Table 1 and Fig. 4 for Si and Ti). All
 519 models present a satisfactory agreement between XRF and FT-IR predicted concentrations (R² >
 520 0.94). The quality of prediction of the elemental oxides falls into two groups. The first group



521 includes silicon, aluminum and calcium and is characterized by moderate relative errors (11 – 13
522 %), similar in magnitude to the FT-IR OC model (12 %) and have similar errors to XRF
523 measurements indicating similar to OC and TC that FT-IR does not add additional uncertainty.
524 The second group includes titanium and iron which have larger relative errors (20.9 – 24.8 %),
525 analogous to HIPS BC and EC models (23.3 – 30 %). Comparing residuals to collocated XRF
526 measurements (Section S9) shows that the FT-IR based models have a larger interquartile range.
527 For Fe, XRF uncertainty is quite low and FT-IR adds additional uncertainty to the measurement.
528 XRF Ti measurements have higher error than the other elements but there is an incremental
529 increase in error due to FT-IR. In addition, cross plots of titanium concentrations show a
530 bifurcation (Fig. 4). While most samples fall near their expected titanium concentration, samples
531 collected at the Sycamore Canyon (AZ) site present a systematic negative bias, consistent across
532 years, tentatively attributed to a site-specific soil composition not accounted for by the Limited
533 calibration. Takahama et al. (2019) demonstrated several methods to identify the possible
534 occurrence of anomalous predictions in OC and EC based on comparison of new sample spectra
535 to calibration spectra based on projected PLS scores and regression residual vectors. These
536 samples with systematic negative bias in titanium predictions can presumably be identified using
537 such an approach, provided that compositional differences are detected in the IR spectrum.
538 Although distinct IR fingerprints exist, FT-IR calibrations for quantifying mineral contents should
539 be interpreted with care as specific elements may be indirectly quantified through their
540 correlation with another element even in the absence of clear IR signature (Hahn et al., 2018).
541 For instance, the variable importance in projection (VIP) scores for the Si, Al, and Ti calibrations
542 suggests use of similar spectral variables, with small differences, for prediction of these species
543 (Section S9). However, the 21 GMM sites coverage still meet the necessary requirements for
544 providing a reliable insight into soil composition in the IMPROVE network.

545 3.3.3 Inorganic ions

546 The two most abundant inorganic anions quantified in the network: nitrate and sulfate can also
547 be measured by FT-IR (absorption bands used for prediction are discussed in Section S9). FT-IR
548 sulfate concentrations display a satisfactory agreement with the reference IR measurements (Fig.
549 4). Model performance metrics include R^2 above 0.98 and relative error lower than 10 % as in the
550 FT-IR PM_{2.5} model (Table 1). IC measurements of sulfate are very good have even lower error
551 than FT-IR sulfate. However, FT-IR nitrate concentrations (Section S9) are characterized by a
552 moderate drop in the overall model performance ($R^2 = 0.927$) while relative bias and error exceed
553 15 % and 50 %, respectively and the error far exceeds reference IC nitrate measurement error. A
554 direct comparison against differential nitrate concentrations at collocated sites highlights the
555 broad uncertainty in determining nitrate content from PTFE filters (Section S9). Unlike nylon
556 filters for which nitrate is trapped on the surface, nitrate is known to evaporate from PTFE filters.
557 This causes a discrepancy between the mass of nitrate deposited onto the nylon filter and the
558 mass of nitrate on the PTFE filter (Eldred and Ashbaugh, 2004), making FT-IR calibrations with the



559 nitrate measurements by IC from nylon filters as the reference method error prone. FT-IR based
560 nitrate concentrations, measured in this way, should be considered with caution. A possible
561 alternative is to develop a set of laboratory calibration standards of ammonium nitrate for FT-IR
562 calibration. The nitrate mass on the PTFE would be useful for mass closure exercises on the PTFE
563 filter but would not adequately assess particulate nitrate in the atmosphere.

564 3.3.4 PM_{2.5} mass predictions

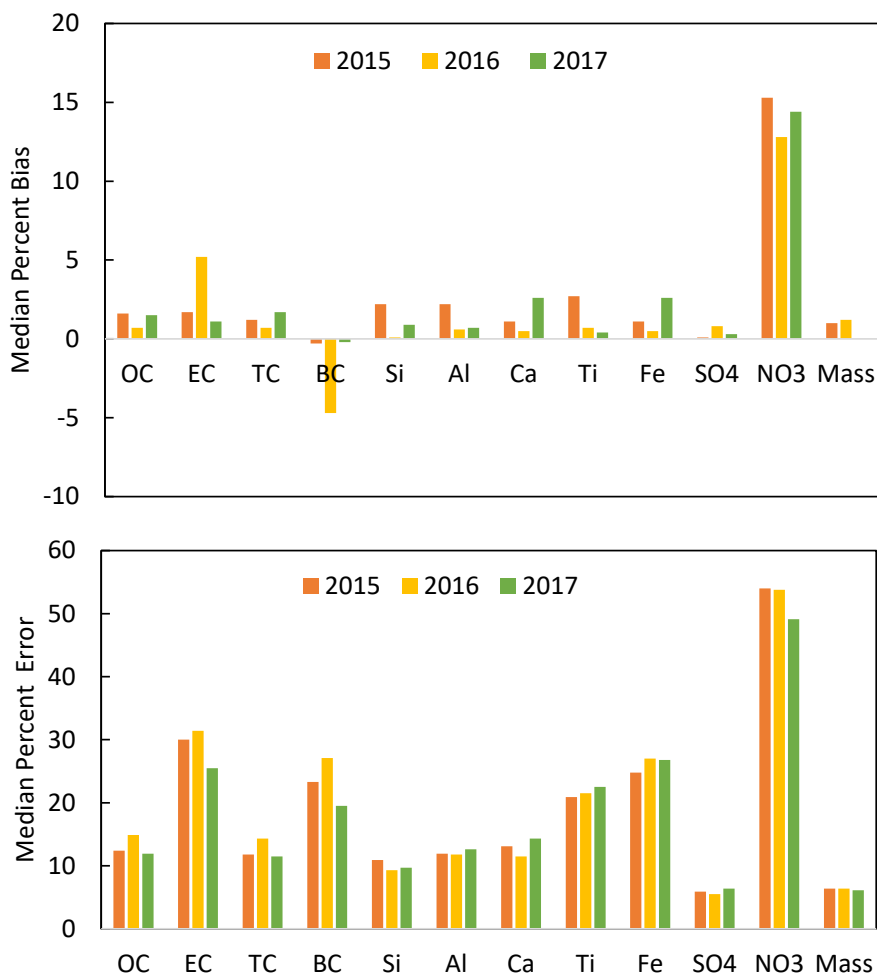
565 Since the major aerosol species are shown to be reasonably well measured by FT-IR, it was
566 anticipated that PM_{2.5} mass calibration would perform well. The PM_{2.5} model presents reliable
567 filter mass predictions ($R^2 = 0.985$) characterized by relative bias and error that are 1/3 to 1/2 of
568 those for OC and on par with gravimetric error (Table 1). The cross plot of gravimetric mass and
569 FT-IR predictions (Fig. 4) shows that PM_{2.5} mass can be accurately predicted across a broad
570 concentration range indicating that FT-IR spectra of PTFE filters do not contain interferences or
571 other limitations that make PM mass predictions error prone.

572 3.4 Long term stability

573 Finally, Multilevel calibrations are extended to 2016 and 2017 to evaluate the inter-year
574 consistency and determine if the assumptions behind Limited Sites and Biomass Burning models
575 remain valid over time. For each sampling year, new calibrations were developed following the
576 framework established for 2015. Models were recalibrated with data from the 21 sites and
577 biomass burning samples were detected by the functional group screening procedure. Fig. 5
578 shows the median relative bias (top) and error (bottom) for the three years of data (cross plots
579 and prediction metrics shown for all predicted species for 2016 and 2017 in Section S10). These
580 results indicate that the modeling methodology provided reasonably consistent results across all
581 three years.

582 Normalized bias for most species is below 3% and normalized error is consistent for all species
583 across all three years. The relative bias for EC and BC are similar to other species in 2015 and
584 2017 but in 2016 they are larger in magnitude than the other two years and different in sign.
585 2016 is the first year of TOR data from the multiwavelength TOR instruments (Chow et al., 2015)
586 so higher bias could be potentially be related the new instruments. However, the HIPS
587 instrument was overhauled beginning in 2017 which provides no explanation for high bias in 2016
588 (http://vista.cira.colostate.edu/improve/Data/QA_QC/Advisory/da0041/da0041_HIPSmodifications.pdf).
589 Further, the EC and BC calibrations are independent of each other except for using
590 the same filters for calibrations (as all species do) so the fact that the median bias is roughly equal
591 but opposite in sign is not due to codependence of the models.

592 In future work, calibrations models will be updated more frequently than annually with the most
593 recent year of ambient samples which may smooth biases and errors due to changes to
594 atmospheric condition and instrument drift.



595

596

597 Figure 5. (top) Median Percent Bias and (bottom) Median Percent Error for each constituent
598 measured for each year.

599

600 4 Conclusion

601 In this paper, we investigate the feasibility of an FT-IR method that uses ambient samples as
602 calibration standards and is adapted for use by a large monitoring network. In this method, all



603 sites in the network collect PTFE filters for FT-IR analysis. A few select sites, used for calibration,
604 would retain all sampling and analyses of current IMPROVE sites to enable re-calibration of the
605 FTIR method on a routine basis. Re-calibration is especially important as atmospheric changes
606 occur and as conditions in the network evolve over time, for example new reference instruments,
607 new or significantly modified FT-IR instruments, changes to sampling protocol or possibly change
608 in filter material. The validity of such a design was evaluated with all PTFE filters collecting PM_{2.5}
609 aerosols at 161 IMPROVE sites in 2015 and then tested for all filters in 2016 and 2017.

610 A multi-level modeling algorithm was used whereby smoke impacted samples are identified and
611 predicted by one model and the rest of the samples are predicted by another model developed
612 from 21 selected IMPROVE sites. The data from the two models are combined to evaluate
613 performance of the FT-IR method. The selection of sites was performed such that if one of the
614 21 sites ceases to operate, another site, selected from the same compositional cluster can be
615 used for calibration.

616 The cross-plots and prediction metrics indicate that the Multilevel model is equivalent to
617 conventional calibrations built from samples from every available site. Reliable performance in
618 predicted concentrations were reported for a broad range of atmospheric constituents with
619 detectable infrared signatures such as OC, EC, TC, sulfate, soil elements (Si, Al, Ca, Ti, Fe), light
620 absorption, and PM_{2.5} mass. Due to volatilization off the PTFE filter, nitrate measurements were
621 found to be unsatisfactory. The calibration method was developed using data from 2015, and the
622 same methodology was applied to 2016 and 2017. The model performance metrics in all three
623 years were similar. Results across ~61,500 FT-IR spectra highlight the suitability of the Multilevel
624 calibration design to quantify multiple atmospheric PM_{2.5} species in large monitoring networks.

625 This work presents an alternative, lower cost, filter analysis method to measure speciated aerosol
626 in an operational routine monitoring network. This could be a valuable addition to routine
627 speciated aerosol monitoring networks, such as IMPROVE, by incorporating monitoring sites that
628 collect samples on only a PTFE filter for subsequent analysis. This would provide the opportunity
629 to have a subset of less expensive monitoring site, which could be used for scoping studies to
630 understand the aerosol composition in unmonitored locations. It could also serve as a network
631 cost savings method by having a subset of network sites collect aerosol samples on only a Teflon
632 filter. However, the inability to measure particulate nitrate is a significant deficiency for using
633 this method to replace existing monitoring sites. The FTIR derived aerosol concentrations are
634 also a semi-independent measurement from the routine speciated aerosol measurements.
635 Therefore, routine FTIR measurements would provide valuable QA/QC information for any
636 speciated monitoring network. In addition, FTIR derived concentrations could be used to
637 substitute for missing concentrations in the case where the Teflon sample is valid, but filter
638 samples or analyses on the nylon or quartz fiber filters are not.



639 For IMPROVE's urban counterpart, the CSN network, after evaluation of the quality of predictions
640 in CSN, this framework could be used to accomplish goals similar to those of IMPROVE.
641 Additionally, this method could be used to predict samples collected in the Federal Reference
642 Method (FRM) network which is a PM mass only network. Finally, this method, with appropriate
643 ambient standards, could be applied at other regional or international monitoring networks or
644 sites to provide low-cost comprehensive composition data.

645

646



647 5 Data availability

648 Data is available at <https://doi.org/10.25338/B8TP8V>.

649 6 Author contribution

650 BD developed the software, performed the formal analysis and visualization for much of the
651 manuscript and wrote the original draft of the manuscript, ATW developed software, performed
652 formal analysis and visualization of the GMM work, ST participated in conceptualization,
653 methodology software, visualization and reviewing and editing the manuscript. KMG developed
654 parts of the biomass burning identification methodology, BS, SC and ASW, provided input
655 throughout the project and reviewed and edited the manuscript, AMD conceptualized of and
656 acquired funding for this project, developed methodology, performed project administration and
657 supervision and reviewed, edited and finalized the manuscript.

658 7 Acknowledgments

659 The authors acknowledge funding from the National Park Service in cooperation with the
660 Environmental Protection Agency (P18AC01222). Thanks to Anahita Amiri Farahani for assisting
661 with figures. We are particularly grateful to Kelsey Seibert for overseeing daily FT-IR operations
662 at the University of California Davis and to the numerous undergraduate students who
663 performed spectra collection from 2015 to 2017.

664 8 References

- 665
666 Besalú, E., Julián-Ortiz, J. V. de., Iglesias, M., and Pogliani, L.: An overlooked property of plot methods,
667 *Journal of Mathematical Chemistry*, 39, 475–484, <https://doi.org/10.1007/s10910-005-9035-z>, 2006.
- 668 Bilmes, J. A.: A gentle tutorial of the EM algorithm and its application to parameter estimation for
669 Gaussian mixture and hidden Markov models, 4:126, 1998.
- 670 Boris, A. J., Takahama, S., Weakley, A. T., Debus, B. M., Fredrickson, C. D., Esparza-Sanchez, M., Burki, C.,
671 Reggente, M., Shaw, S. L., Edgerton, E. S., and Dillner, A. M.: Quantifying organic matter and functional
672 groups in particulate matter filter samples from the southeastern United States – Part 1: Methods,
673 *Atmos. Meas. Tech.*, 12, 5391–5415, <https://doi.org/10.5194/amt-12-5391-2019>, 2019.
- 674 Boris, A. J., Takahama, S., Weakley, A. T., Debus, B. M., Shaw, S. L., Edgerton, E. S., Joo, T., Ng, N. L., and
675 Dillner, A. M.: Quantifying organic matter and functional groups in particulate matter filter samples from
676 the southeastern United States - Part 2: Spatiotemporal trends, 14, 4355–4374,
677 <https://doi.org/10.5194/amt-14-4355-2021>, 2021.



- 678 Bro, R. and Smilde, A. K.: Centering and scaling in component analysis, *Journal of Chemometrics*, 17, 16–
679 33, <https://doi.org/10.1002/cem.773>, 2003.
- 680 Burki, C., Reggente, M., Dillner, A. M., Hand, J. L., Shaw, S. L., and Takahama, S.: Analysis of functional
681 groups in atmospheric aerosols by infrared spectroscopy: method development for probabilistic
682 modeling of organic carbon and organic matter concentrations, 13, 1517–1538,
683 <https://doi.org/10.5194/amt-13-1517-2020>, 2020.
- 684 Busca, G. and Resini, C.: *Vibrational Spectroscopy for the Analysis of Geological and Inorganic Materials*,
685 in: *Encyclopedia of Analytical Chemistry*, John Wiley & Sons, Ltd, 2006.
- 686 Chow, J. C., Watson, J. G., Pritchett, L. C., Pierson, W. R., Frazier, C. A., and Purcell, R. G.: The dri
687 thermal/optical reflectance carbon analysis system: description, evaluation and applications in U.S. Air
688 quality studies, 27, 1185–1201, [https://doi.org/10.1016/0960-1686\(93\)90245-T](https://doi.org/10.1016/0960-1686(93)90245-T), 1993.
- 689 Chow, J. C., Watson, J. G., Chen, L. W. A., Chang, M. C. O., Robinson, N. F., Trimble, D., and Kohl, S.: The
690 IMPROVE_A Temperature Protocol for Thermal/Optical Carbon Analysis: Maintaining Consistency with a
691 Long-Term Database, 57, 1014–1023, <https://doi.org/10.3155/1047-3289.57.9.1014>, 2007a.
- 692 Chow, J. C., Watson, J. G., Chen, L. W. A., Chang, M. C. O., Robinson, N. F., Trimble, D., and Kohl, S.: The
693 IMPROVE_A Temperature Protocol for Thermal/Optical Carbon Analysis: Maintaining Consistency with a
694 Long-Term Database, 57, 1014–1023, <https://doi.org/10.3155/1047-3289.57.9.1014>, 2007b.
- 695 Chow, J. C., Wang, X., Sunlin, B. J., Gronstal, S. B., Chen, L.-W. A., Trimble, D. L., Watson, J. G., Kohl, S. D.,
696 Mayorga, S. R., Riggio, G., Hurbain, P. R., Johnson, M., and Zimmermann, R.: Optical Calibration and
697 Equivalence of a Multiwavelength Thermal/Optical Carbon Analyzer, 15, 1145–1159,
698 <https://doi.org/10.4209/aaqr.2015.02.0106>, 2015.
- 699 Chukanov, N. V. and Chervonnyi, A. D.: *Infrared Spectroscopy of Minerals and Related Compounds*,
700 Springer International Publishing, 2016.
- 701 Cios, K., Pedrycz, W., and Swiniarski, R. W.: *Data Mining Methods for Knowledge Discovery*, Kluwer
702 Academic Publishers, Norwell, MA, USA, 495 pp., 1998.
- 703 Corrigan, A. L., Russell, L. M., Takahama, S., Äijälä, M., Ehn, M., Junninen, H., Rinne, J., Petäjä, T.,
704 Kulmala, M., Vogel, A. L., Hoffmann, T., Ebben, C. J., Geiger, F. M., Chhabra, P., Seinfeld, J. H., Worsnop,
705 D. R., Song, W., Auld, J., and Williams, J.: Biogenic and biomass burning organic aerosol in a boreal forest
706 at Hyytiälä, Finland, during HUMPPA-COPEC 2010, 13, 12233–12256, <https://doi.org/10.5194/acp-13-12233-2013>, 2013.
- 708 Debus, B., Takahama, S., Weakley, A. T., Seibert, K., and Dillner, A. M.: Long-Term Strategy for Assessing
709 Carbonaceous Particulate Matter Concentrations from Multiple Fourier Transform Infrared (FT-IR)
710 Instruments: Influence of Spectral Dissimilarities on Multivariate Calibration Performance, 73, 271–283,
711 <https://doi.org/10.1177/0003702818804574>, 2019.



- 712 Dillner, A. M. and Takahama, S.: Predicting ambient aerosol thermal-optical reflectance measurements
713 from infrared spectra: Elemental carbon, 8, 4013–4023, <https://doi.org/10.5194/amt-8-4013-2015>,
714 2015a.
- 715 Dillner, A. M. and Takahama, S.: Predicting ambient aerosol thermal-optical reflectance (TOR)
716 measurements from infrared spectra: Organic carbon, 8, 1097–1109, <https://doi.org/10.5194/amt-8-1097-2015>, 2015b.
717
- 718 Draper, N. R. and Smith, H.: Applied Regression Analysis, 1998.
- 719 Eldred, Robert. and Ashbaugh, L. L.: Loss of particle nitrate from Teflon sampling filters: Effects on
720 measured gravimetric mass in California and in the IMPROVE network, 54, 93–104,
721 <https://doi.org/10.1080/10473289.2004.10470878>, 2004.
- 722 Friedel, R. A. and Carlson, G. L.: Infrared spectra of ground graphite, 75, 1149–1151,
723 <https://doi.org/10.1021/j100678a021>, 1971.
- 724 Hahn, A., Vogel, H., Andó, S., Garzanti, E., Kuhn, G., Lantzsich, H., Schürman, J., Vogt, C., and Zabel, M.:
725 Using Fourier transform infrared spectroscopy to determine mineral phases in sediments, *Sedimentary
726 Geology*, 375, 27–35, <https://doi.org/10.1016/j.sedgeo.2018.03.010>, 2018.
- 727 Hastie, T, Tibshirani, R, and Friedman, J: The Elements of Statistical Learning: Data Mining, Inference,
728 and Prediction, Second Edition, Springer New York, New York, NY, 2009.
- 729 Hawkins, L. N., Russell, L. M., Covert, D. S., Quinn, P. K., and Bates, T. S.: Carboxylic acids, sulfates, and
730 organosulfates in processed continental organic aerosol over the southeast Pacific Ocean during
731 VOCALS-REx 2008, 115, <https://doi.org/10.1029/2009jd013276>, 2010.
- 732 Hyslop, N. P., Trzepla, K., and White, W. H.: Reanalysis of Archived IMPROVE PM2.5 Samples Previously
733 Analyzed over a 15-Year Period, *Environ. Sci. Technol.*, 46, 10106, 2012.
- 734 Hyslop, N. P., Trzepla, K., and White, W. H.: Assessing the Suitability of Historical PM2.5 Element
735 Measurements for Trend Analysis, *Environ. Sci. Technol.*, 49, 9247–9255,
736 <https://doi.org/10.1021/acs.est.5b01572>, 2015.
- 737 Kuzmiakova, A., Dillner, A. M., and Takahama, S.: An automated baseline correction protocol for infrared
738 spectra of atmospheric aerosols collected on polytetrafluoroethylene (Teflon) filters, 9, 2615–2631,
739 <https://doi.org/10.5194/amt-9-2615-2016>, 2016.
- 740 Li, H.-D., Xu, Q.-S., and Liang, Y.-Z.: libPLS: An integrated library for partial least squares regression and
741 linear discriminant analysis, 176, 34–43, <https://doi.org/10.1016/j.chemolab.2018.03.003>, 2018.
- 742 Liu, S., Ahlm, L., Day, D. A., Russell, L. M., Zhao, Y., Gentner, D. R., Weber, R. J., Goldstein, A. H., Jaoui,
743 M., Offenberg, J. H., Kleindienst, T. E., Rubitschun, C., Surratt, J. D., Sheesley, R. J., and Scheller, S.:



- 744 Secondary organic aerosol formation from fossil fuel sources contribute majority of summertime organic
745 mass at Bakersfield, 117, <https://doi.org/10.1029/2012JD018170>, 2012.
- 746 Madejová, J. and Komadel, P.: Baseline studies of the clay minerals society source clays: Infrared
747 methods, *Clays and Clay Minerals*, 49, 410–432, <https://doi.org/10.1346/CCMN.2001.0490508>, 2001.
- 748 Mahalanobis, P., C.: On the generalized distance in statistics, 2, 49–55, 1936.
- 749 Malm, W. C., Sisler, J. F., Huffman, D., Eldred, R. A., and Cahill, T. A.: Spatial and seasonal trends in
750 particle concentration and optical extinction in the United States, 99, 1347–1370,
751 <https://doi.org/10.1029/93jd02916>, 1994.
- 752 Margenot, A. J., Calderón, F. J., Goynes, K. W., Mukome, F. N. D., and Parikh, S. J.: IR Spectroscopy, Soil
753 Analysis Applications, in: *Encyclopedia of Spectroscopy and Spectrometry (Third Edition)*, edited by:
754 Lindon, J. C., Tranter, G. E., and Koppenaal, D. W., Academic Press, Oxford, 448–454,
755 <https://doi.org/10.1016/B978-0-12-409547-2.12170-5>, 2017.
- 756 Mayo, D. W., Miller, F. A., and Hannah, R. W.: *Course Notes on the Interpretation of Infrared and Raman
757 Spectra*, John Wiley & Sons, Hoboken, NJ, 2004.
- 758 Nemanich, R. J., Lucovsky, G., and Solin, S. A.: Infrared active optical vibrations of graphite, *Solid State
759 Communications*, 23, 117–120, [https://doi.org/10.1016/0038-1098\(77\)90663-9](https://doi.org/10.1016/0038-1098(77)90663-9), 1977.
- 760 Ngo, M. A., Pinkerton, K. E., Freeland, S., Geller, M., Ham, W., Cliff, S., Hopkins, L. E., Kleeman, M. J.,
761 Kodavanti, U. P., Meharg, E., Plummer, L., Recendez, J. J., Schenker, M. B., Sioutas, C., Smiley-Jewell, S.,
762 Haas, C., Gutstein, J., and Wexler, A. S.: Airborne particles in the San Joaquin Valley may affect human
763 health, 64, 12–16, <https://doi.org/10.3733/ca.v064n01p12>, 2010.
- 764 Niyogi, S., Bekyarova, E., Itkis, M. E., McWilliams, J. L., Hamon, M. A., and Haddon, R. C.: Solution
765 Properties of Graphite and Graphene, 128, 7720–7721, <https://doi.org/10.1021/ja060680r>, 2006.
- 766 O’Dell, K., Ford, B., Fischer, E. V., and Pierce, J. R.: Contribution of wildland-fire smoke to US PM_{2.5} and
767 its influence on recent trends, 53, 1797–1804, <https://doi.org/10.1021/acs.est.8b05430>, 2019.
- 768 Parks, D. A., Griffiths, P. R., Weakley, A. T., and Miller, A. L.: Quantifying elemental and organic carbon in
769 diesel particulate matter by mid-infrared spectrometry, *null*, 55, 1014–1027,
770 <https://doi.org/10.1080/02786826.2021.1917764>, 2021.
- 771 Reggente, M., Dillner, A. M., and Takahama, S.: Predicting ambient aerosol thermal-optical reflectance
772 (TOR) measurements from infrared spectra: Extending the predictions to different years and different
773 sites, 9, 441–454, <https://doi.org/10.5194/amt-9-441-2016>, 2016.



- 774 Reggente, M., Dillner, A. M., and Takahama, S.: Analysis of functional groups in atmospheric aerosols by
775 infrared spectroscopy: systematic intercomparison of calibration methods for US measurement network
776 samples, *Atmos. Meas. Tech.*, 12, 2287–2312, <https://doi.org/10.5194/amt-12-2287-2019>, 2019.
- 777 Russell, L. M., Bahadur, R., and Ziemann, P. J.: Identifying organic aerosol sources by comparing
778 functional group composition in chamber and atmospheric particles, 108, 3516–3521,
779 <https://doi.org/10.1073/pnas.1006461108>, 2011.
- 780 Ruthenburg, T. C., Perlin, P. C., Liu, V., McDade, C. E., and Dillner, A. M.: Determination of organic matter
781 and organic matter to organic carbon ratios by infrared spectroscopy with application to selected sites in
782 the IMPROVE network, 86, 47–57, <https://doi.org/10.1016/j.atmosenv.2013.12.034>, 2014.
- 783 Savitzky, A. and Golay, M. J. E.: Smoothing and differentiation of data by simplified least squares
784 procedures, 36, 1627–1639, <https://doi.org/10.1021/ac60214a047>, 1964.
- 785 Schichtel, B. A., Malm, W. C., Bench, G., Fallon, S., McDade, C. E., Chow, J. C., and Watson, J. G.: Fossil
786 and contemporary fine particulate carbon fractions at 12 rural and urban sites in the United States, 113,
787 <https://doi.org/10.1029/2007jd008605>, 2008.
- 788 Senthil Kumar, R. and Rajkumar, P.: Characterization of minerals in air dust particles in the state of
789 Tamilnadu, India through ftir spectroscopy, 2013, 22221–22248, <https://doi.org/10.5194/acpd-13-22221-2013>, 2013.
- 791 Snijders, T. A. B. and Bosker, R. J.: *Multilevel Analysis: An Introduction to Basic and Advanced Multilevel*
792 *Modeling*, SAGE, 370 pp., 2011.
- 793 Sorooshian, A., Wonaschütz, A., Jarjour, E. G., Hashimoto, B. I., Schichtel, B. A., and Betterton, E. A.: An
794 aerosol climatology for a rapidly growing arid region (southern Arizona): Major aerosol species and
795 remotely sensed aerosol properties, 116, <https://doi.org/10.1029/2011jd016197>, 2011.
- 796 Takahama, S., Schwartz, R. E., Russell, L. M., Macdonald, A. M., Sharma, S., and Leaitch, W. R.: Organic
797 functional groups in aerosol particles from burning and non-burning forest emissions at a high-elevation
798 mountain site, *Atmos. Chem. Phys.*, 11, 6367–6386, <https://doi.org/10.5194/acp-11-6367-2011>, 2011.
- 799 Takahama, S., Dillner, A. M., Weakley, A. T., Reggente, M., Bürki, C., Lbadaoui-Darvas, M., Debus, B.,
800 Kuzmiakova, A., and Wexler, A. S.: Atmospheric particulate matter characterization by Fourier transform
801 infrared spectroscopy: a review of statistical calibration strategies for carbonaceous aerosol
802 quantification in US measurement networks, *Atmos. Meas. Tech.*, 12, 525–567,
803 <https://doi.org/10.5194/amt-12-525-2019>, 2019.
- 804 Tuinstra, F. and Koenig, J. L.: Raman Spectrum of Graphite, 53, 1126–1130,
805 <https://doi.org/10.1063/1.1674108>, 1970.



- 806 Watson, J. G. and Chow, J. C.: A wintertime PM_{2.5} episode at the Fresno, CA, supersite, *Atmospheric*
807 *Environment*, 36, 465–475, [https://doi.org/10.1016/S1352-2310\(01\)00309-0](https://doi.org/10.1016/S1352-2310(01)00309-0), 2002.
- 808 Weakley, A. T., Takahama, S., and Dillner, A. M.: Ambient aerosol composition by infrared spectroscopy
809 and partial least-squares in the chemical speciation network: Organic carbon with functional group
810 identification, 50, 1096–1114, <https://doi.org/10.1080/02786826.2016.1217389>, 2016.
- 811 Weakley, A. T., Takahama, S., Wexler, A. S., and Dillner, A. M.: Ambient aerosol composition by infrared
812 spectroscopy and partial least squares in the chemical speciation network: Multilevel modeling for
813 elemental carbon, 52, 642–654, <https://doi.org/10.1080/02786826.2018.1439571>, 2018a.
- 814 Weakley, A. T., Takahama, S., and Dillner, A. M.: Thermal/optical reflectance equivalent organic and
815 elemental carbon determined from federal reference and equivalent method fine particulate matter
816 samples using Fourier transform infrared spectrometry, 52, 1048–1058,
817 <https://doi.org/10.1080/02786826.2018.1504161>, 2018b.
- 818 White, W. H., Trzepla, K., Hyslop, N. P., and Schichtel, B. A.: A critical review of filter transmittance
819 measurements for aerosol light absorption, and de novo calibration for a decade of monitoring on PTFE
820 membranes, 50, 984–1002, <https://doi.org/10.1080/02786826.2016.1211615>, 2016.
- 821 Wold, S., Sjöström, M., and Eriksson, L.: PLS-regression: A basic tool of chemometrics, 58, 109–130,
822 [https://doi.org/10.1016/S0169-7439\(01\)00155-1](https://doi.org/10.1016/S0169-7439(01)00155-1), 2001.
- 823 Yazdani, A., Dudani, N., Takahama, S., Bertrand, A., Prevot, A. S. H., El Haddad, I., and Dillner, A. M.:
824 Characterization of primary and aged wood burning and coal combustion organic aerosols in an
825 environmental chamber and its implications for atmospheric aerosols, 21, 10273–10293,
826 <https://doi.org/10.5194/acp-21-10273-2021>, 2021a.
- 827 Yazdani, A., Dillner, A. M., and Takahama, S.: Estimating mean molecular weight, carbon number, and
828 OM/OC with mid-infrared spectroscopy in organic particulate matter samples from a monitoring
829 network, 14, 4805–4827, <https://doi.org/10.5194/amt-14-4805-2021>, 2021b.
- 830 Zhang, X. L., Trzepla, K., White, W., Raffuse, S., and Hyslop, N. P.: Intercomparison of thermal-optical
831 carbon measurements by Sunset and Desert Research Institute (DRI) analyzers using the IMPROVE_A
832 protocol, 14, 3217–3231, <https://doi.org/10.5194/amt-14-3217-2021>, 2021.
- 833