Cloud phase and macrophysical properties over the Southern Ocean during the MARCUS field campaign

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Abstract.

To investigate the cloud phase and macrophysical properties over the Southern Ocean (SO), the Department of Energy (DOE) Atmospheric Radiation Measurement (ARM) Mobile Facility (AMF2) was installed on the Australian icebreaker *Aurora Australis* during the MARCUS field campaign [41 to 69 ºS; 60 to 160 ºE] from October 2017 to March 2018. To examine cloud properties over the mid-latitude and Polar regions, the study domain is separated into northern (NSO) and southern (SSO) parts of the SO with a demarcation line of 60 °S. The total cloud fractions (*CF*s) were 77.9 %, 67.6 %, and 90.3 % for the entire domain, NSO and SSO, respectively, indicating that higher *CF*s were observed in the Polar region. Low-level clouds and deep convective clouds are the two most common cloud types over the SO.

A new method was developed to classify liquid, mixed-phase and ice clouds in single-layered low-level clouds (LOW) where mixed-phase clouds dominate with an occurrence frequency (*Freq*) of 54.5 %, while the *Freq* of the liquid and ice clouds were 10.1 % (most drizzling) and 17.4 % (least drizzling). The meridional distributions of low-level cloud boundaries are nearly independent of latitude, whereas the cloud temperatures increased ~8 K and atmospheric precipitable water vapor increased from ~5 mm at 69 °S to ~18 mm at 43 °S. The mean cloud liquid water paths over NSO were much larger than those over SSO. Most liquid clouds occurred over NSO with very few over SSO, whereas more mixed-phase clouds occurred over SSO than over NSO. There were no significant differences for ice cloud *Freq* between NSO and SSO. The ice particle sizes are comparable to cloud droplets and drizzle drops, and well mixed in the cloud layer. These results will be valuable for advancing our understanding of the meridional and vertical distributions of clouds and can be used to improve model simulations over the SO.

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1. Introduction

The Southern Ocean (SO) is one of the cloudiest and stormiest regions on the Earth (Mace et al., 2009; Chubb et al., 2013). Over the SO, most of the aerosols are naturally produced via oceanic sources given the remote environment. The uncertainties of aerosol forcing caused by natural emissions have larger variances than anthropogenic emissions, especially the dimethyl sulfide (DMS) flux contributes significantly to the bias (Carslaw et al., 2013). The SO is a unique natural laboratory to address the natural aerosol emissions and their contributions to the biases because it has rich ecosystems and is remote to human activities (McCoy et al., 2015). However, we have limited knowledge about cloud formation processes within such clean environments and their associated aerosol and cloud properties. The unique nature of the SO region features low-level supercooled liquid and mixed-phase clouds, which is significantly different from the subtropical marine boundary layer (MBL) clouds where warm liquid clouds are dominant (Dong et al., 2014; Wu et al., 2020; Zhao et al., 2020), and also different to the Arctic mixed-phase clouds which are featured with the liquid-topped cloud layer with ice cloud layer beneath (Qiu et al., 2015).

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Large biases in cloud amount and microphysics over the SO in the Coupled Model Intercomparison Project phase 5 (CMIP5) climate models result in a near 30 W m^{-2} shortwave radiation deficit at the top of the atmosphere (TOA) (Marchand et al., 2014; Stanfield et al., 2014, 2015), which further leads to unrealistic cloud feedbacks and equilibrium climate sensitivity (Bony et al., 2015; Stocker et al., 2013). Meanwhile, the efficiency of aerosol-cloud interaction (ACI) over the SO was found to be crucial for the models' sensitivities to the radiation budget. A new aerosol scheme in the Hadley Centre Global Environmental model can dampen the ACI and suppress negative clear-sky shortwave feedback, both of which contribute to a larger climate sensitivity (Bodas-Salcedo et a., 2019).

A climate sensitivity study using CMIP6 general circulation models (GCMs) shows much higher temperature variations across 27 GCMs in response to doubled CO_2 than those in CMIP5, which may have resulted from the decreased extratropical low-level cloud cover and cloud albedo over the SO in CMIP6 (Zelinka et al., 2020). Low-level clouds are a key climate uncertainty and can explain 50 % of the inter-model variations (Klein et al., 2017) because conversion from liquid cloud droplets to ice cloud particles decreases the cloud albedo and reduces the reflected shortwave radiation at TOA. Models, however, have difficulties accurately partitioning the cloud phase (Kalesse et al., 2016). The phase changes in mixed-phase clouds over the Arctic have proved to affect the cloud lifetime and radiative properties significantly, that is, converting from ice cloud particles to liquid cloud droplets may increase the cloud optical depth and the reflected shortwave radiation at TOA (Morrison et al., 2012). In contrast, models that allow mixed-phase clouds to glaciate rapidly can produce 30% more warming from doubling CO_2 (McCoy et al., 2014).

Phase transition processes have been investigated by several groups using both satellite and ground-based measurements. For instance, Mace and Protat (2018) found that there are more mixed-phase clouds over the SO measured from the ship than retrieved from CloudSat and CALIPSO measurements because the satellites cannot accurately measure clouds below ~1 km. Lang et al. (2018) used a model to investigate the clouds under post cold frontal systems and found large biases in model simulations and concluded that the cloud cover and radiative biases over the SO are highly regime dependent. Of all cloud types, low-level clouds are primarily responsible for the biases in the model simulations due to the lack of reliable measurements, which leads to a poor understanding of the conditions where these clouds form and the phase(s) that result. In other words, a physical representation of clouds, especially for low-level clouds, is unclear but truly necessary for improving model simulations. Therefore, reliable observations of the cloud macro-

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and micro-physical properties from ground-based active and passive remote sensors are crucial for the improvement of model simulations.

Previous studies show that cloud phase is primarily dependent on cloud temperature, and the transition from one cloud phase to another will modify the cloud optical properties, which further affects the radiation budgets (Hu et al., 2010; Intrieri et al., 2002; Morrison et al., 2012). Based on satellite observations and retrievals, Hu et al. (2010) found that supercooled liquid water (SLW) clouds are most common in the low-level clouds over the SO, where 80% of low-level clouds contain SLW in a wide range of cloud temperatures from 0°C to −40°C. The formation of SLW clouds is usually related to strong boundary layer convection. However, when ice nuclei exist in the mixed-phase clouds, the ice particles can grow quickly and become bigger through consuming supercooled liquid water drops. The SLW is inherently unstable due to the higher vapor pressure over liquid than over ice and the quicker vapor deposition on ice particles than on liquid droplets (Intrieri et al., 2002). As the supercooled liquid cloud droplets glaciate to ice particles, the cloud layer becomes darker because the ice particles scatter less shortwave radiation and absorb more radiation in the near IR wavelength regime. It is unclear, however, what role these ice particles play in the low-level clouds over the SO, which includes the impact on drizzle development. During HIAPER Pole-to-Pole Observation (HIPPO) campaigns, Chubb et al. (2013) found that there are rarely ice particles in non-drizzling and light drizzling clouds over the SO, which may imply that the ice particles in the mixed-phase clouds may modulate the drizzle formation.

To investigate the aerosol and cloud properties over the SO, a field campaign called the Measurements of Aerosols, Radiation, and Clouds over the Southern Ocean (MARCUS) was conducted using the ship-based measurements between Hobart, Australia, and the Antarctic during the period October 2017-March 2018. The Department of Energy (DOE) Atmospheric Radiation

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Measurement (ARM) Mobile Facility (AMF2) was installed on the Australian icebreaker *Aurora Australis*, which voyaged from Hobart, Tasmania to the Australian Antarctic stations of Casey, Mawson, and Davis, as well as Macquarie Island as illustrated in Fig. 1. Another field campaign, called South Ocean Clouds, Radiation, Aerosol Transport Experimental Study (SOCRATES) field campaign was conducted during austral summer from January 15 to February 26, 2018. In this study, the aircraft in-situ measurements during SOCRATES are used as the reference for the analysis. The SOCRATES domain is shown in the black dotted rectangle box in Fig. 1. The objectives of the MARCUS campaign are to investigate the vertical distribution of boundary layer clouds and reveal the reasons why the mixed-phase clouds are common in the warm season (McFarquhar et al., 2016; McFarquhar et al., 2021). Our study will focus on cloud macrophysical properties and cloud phase along the shiptracks during MARCUS.

MARCUS ship-based instruments include AMF2 cloud radar, lidar, microwave radiometer, micropulse lidar, radiosonde sounding, precision solar pyranometer and precision infrared radiometer, as well as aerosol sensors. Through these comprehensive observations over the SO, we are tentatively answering the following three scientific questions:

(1) What is the total cloud fraction over the SO during MARCUS, as well as vertical and meridional variations in cloud fraction?

(2) What are the dominant cloud types over the SO, their associated cloud phase and macrophysical properties, as well as their vertical and meridional distributions?

(3) What are the vertical and meridional distributions of the low-level clouds over the SO?

This manuscript is organized as follows: the data and method and introduced in section 2. The statistical results for all clouds during MARCUS are summarized in section 3. The low-level cloud

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phase and macrophysical properties are described in section 4, followed by a summary and conclusions in section 5.

2. Data and Method

2.1 Ship-based measurements used in this study

The AMF2 instruments, measurements, and their corresponding uncertainties and references are listed in Table 1. Because AMF2 was designed to support shipboard deployments, the baseline suite of instruments are marine-focused, including the 95-GHz W-band cloud radar (WACR), ceilometer, micropulse lidar (MPL), microwave radiometer (MWR), aerosol observation system (AOS), meteorological measurements (MET, includes the following data: temperature, pressure, specific humidity, wind direction and speed) on the ship, rain gauge and the radiosonde soundings. The combined cloud radar and ceilometer measurements can provide the cloud boundaries as long as there are no optically thin clouds and the cloud-base heights (H_{base}) are not greater than the upper limit (7.7 km) of the ceilometer. The micropulse lidar will be used to identify optically thin clouds and the clouds with H_{base} > 7.7 km. A previous study has shown that these additional clouds detected by the micropulse lidar can be a non-negligible supplement to the total cloud fraction (Mace et al., 2021). A detailed description of the instruments and the cloud parameters during MARCUS can be found in Mace et al. (2021) and McFarquhar et al. (2016 and 2021).

In order to accurately estimate the cloud temperatures, we adopted a linear interpolation method based on the daily balloon soundings (4 to 5 times per day) to achieve a better temporal resolution of the vertical profiles of temperature, pressure, and specific humidity. The method considers MET measurements to ensure vertical continuity and adjacent soundings for temporal continuity. Using these interpolated atmospheric profiles, cloud temperatures can be obtained at a 5-min temporal resolution.

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The cloud liquid water path (*LWP*) and atmospheric precipitable water vapor (*PWV*) are retrieved based on a physical-iterative algorithm using observations of the microwave radiometer brightness temperatures at 23.8 and 31.4 GHz with uncertainties ranging from 15 to 30 g m^{-2} (Marchand et al., 2003). It is important to note that the brightness temperature biases switch signs among different climatological regions because a threshold of 5 ºC in cloud-base temperature was used in their physical retrievals. Since the retrieved LWP and PWV are based on the MWR measured brightness temperatures at two frequencies, any biases on the brightness temperatures will affect these retrievals. Therefore, we propose an extra step to determine the uncertainties during MARCUS. Based on the temperature profiles, we can identify clouds that are not likely to contain liquid (e.g., pure ice-cloud), then we can estimate the LWP uncertainty based on their corresponding retrieved LWP values. From the probability density function (PDF_ analysis, the LWP uncertainty is estimated as 10 g m^{-2} for MARCUS.

To determine the precipitation status, the AOS and rain gauge measurements were used to determine whether rain is reaching the surface qualitatively, but not quantitatively in this study. All the measurements were averaged over 5 minutes, except the radar reflectivity, Doppler velocity, and spectrum width used in Section 4.3.

2.2 Cloud type classification and single-layer low cloud phases

A classification method developed in Xi et al. (2010) was used to categorize different types of clouds using ARM radar-lidar estimated cloud base (H_{base}) and top (H_{top}) heights and cloud thickness (ΔH). A brief description of the classification of cloud types is as follows (Table 1 and Figure 6 in Xi et al., 2010). The single-layered low-level clouds (LOW) is the fraction of time when low clouds with $H_{top} \leq 3$ km occur without clouds above them. Middle clouds (MID) range from 3 to 6 km without any clouds below and above, while high clouds (HGH) have $H_{base} > 6$ km

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with no cloud underneath. Other types of clouds are defined by different combinations of the above three types, middle over low (MOL), high over low (HOL), high over middle (HOM), and the cloud column through the entire troposphere is defined as HML. Three types, MOL, HOM, and HML, include both contiguous and non-contiguous cloud layers, and their thicknesses may be overestimated when clear layer(s) are present between any two cloud layers.

Furthermore, we used the measurements of interpolated sounding, microwave radiometer retrieved *LWP*, radar reflectivity, Doppler velocity and spectrum width to classify the cloud phase in each radar range volume of low-level clouds during MARCUS. The detailed classification method will be introduced in Section 4.1. We also used ERA-Interim reanalysis data to study the environmental conditions during MARCUS. The lower tropospheric stability (LTS) is calculated from the potential temperature difference between the surface and 700 hPa to assess the boundary-layer stabilities when the low-level clouds appeared along the shiptracks. The relative contributions of mixed-phase, liquid and ice clouds to the single-layered low-level clouds as well as their drizzling status are also analyzed in this study. The latitudinal and longitudinal variations of the single-layered low-level clouds as well as their vertical distributions are further explored in this study.

3. Statistical results for all clouds during MARCUS

The occurrence frequencies of total cloud cover and different types of clouds and their associated properties over the entire study domain during MARCUS are presented in Figs. 2 - 4. In order to examine the cloud properties over the mid-latitude and Polar regions, we separate the SO domain into northern (NSO, north of 60°S) and southern (SSO, south of 60°S) parts using a demarcation line of 60ºS. A total of 2,447 hours cloud samples were collected during MARCUS in this study, in which 1,181 hours of samples were located in the NSO and 1,266 hours of samples

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A classification method developed in Xi et al. (2010) was used to calculate the occurrence frequencies of different types of clouds and their corresponding cloud macrophysical properties, e.g., cloud base (H_{base}) and top (H_{top}) heights, cloud thickness (ΔH), and *LWP*.

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were collected from the SSO. It is important to note that adding micropulse lidar measurements increased the total samples of non-liquid-containing clouds by ~20% because micropulse lidar is more sensitive to optically thin clouds than cloud radar. However, micropulse lidar signals are usually attenuated and cannot provide a meaningful signal when the liquid cloud layer is thicker than a couple of hundred meters (Sassen, 1991).

Figure 2 shows the vertical distributions of total cloud cover over the entire domain, as well as over NSO and SSO. For the vertical distributions, the occurrence frequencies of total cloud increase from the first radar gate (~ 226 m) to ~700 m, then monotonically decrease with altitude with a few small increments at different levels, especially over SSO. Comparing the occurrence frequencies of total cloud between NSO and SSO, we can draw the following conclusions. 1) The SSO has more cloudiness than the NSO under 7 km, while the NSO has more cloudiness than the SSO above 7 km. 2) Below 3 km, the occurrence frequencies of clouds over the NSO decrease dramatically from 37 % at an altitude of ~700 m to 16 % at 3 km and from 45 % to 28 % over the SSO, which is similar to the vertical distributions of the low-level clouds over some Northern Hemisphere mid-latitude regions, such as Eastern North Atlantic (ENA, Dong et al., 2014). The occurrence frequencies measured during MARCUS are much lower than these shown in Fig. 8 of Mace et al., (2009) throughout the entire vertical column between the same range of latitudes, especially, the occurrence frequencies during MARCUS are almost half of those measured by CloudSat and CALIPSO from 1 to 3 km. The reason has been explained in Xi et al., (2010), that is, a comparison of occurrence frequencies between measurements of two different platforms can only be performed under an equivalent spatial-to-temporal resolution. In other words, our results were calculated under 5-min temporal resolution, and the results in Mace et al., (2009) were statistically in the 2^{o} gridbox. Therefore, the comparison between these two results is not

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reasonable. To make a fair comparison, one has to know the cloud amount at each area or time step, then the product of amount and frequency is independent of either temporal and spatial measurement.

To compare with other studies, we calculated the cloud fractions (*CF*s) of total and different types of clouds. The total *CF*s were 77.9 %, 67.6 %, and 90.3 % for the entire domain, NSO and SSO, respectively, indicating that 22.7 % more clouds occurred in the Polar region than in the mid-latitude region. The total *CF* over the entire domain is very close to the 76 % calculated by Mace and Protat (2018) using ship-based measurements during the Cloud, Aerosols, Precipitation, Radiation and Atmospheric Composition (CAPRICORN) field experiment. The total *CF* over the SSO is close to that estimated by using the complementarity of CALIOP lidar aboard CALIPSO and CPR aboard CloudSat (DARDAR version 2 data) from Listowski et al. (2019).

Figiure 3 shows the occurrence frequencies of categorized clouds and their cloud boundaries using the maximum H_{top} and the minimum H_{base} if there are two or more layers in each 5-min sample. For example, the mean H_{base} and H_{top} for single-layered low-level (LOW) are 0.92 km and 1.62 km, respectively, listed in Table 2, which are the average values of min H_{base} and max H_{top} in LOW category. As illustrated in Fig. 3a, the single-layered low-level (LOW), deep cumulus or multi-layered (HML), and MOL clouds are the three dominant types of clouds over the SO. Comparing the clouds between NSO and SSO, all types of clouds in SSO have higher frequency of occurrence than those in NSO except HOL. The differences range from less than 1 % (LOW) to more than 10 % (MOL). Comparing the clouds over mid-latitude oceans between the two hemispheres, i.e., between NSO and ARM ENA site (Dong et al., 2014), we find: (1) The total cloud fractions (*CF*s) are close to each other (67.6 % over NSO vs. 70.1 % at ARM ENA); (2) LOW *CF*s are 22.9 % vs. 27.1 %, which is the dominant type of cloud in both regions; and (3)

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Both MOL and HML clouds, including underneath low clouds, are 14.2 % and 16.5 % over NSO, much higher than those (4.2 % and 12.1 %) at ARM ENA site, indicating that there are more MOL and deep convective clouds over NSO than over ENA.

Figure 3b shows the vertical locations of different types of cloud layers, which represent the mean H_{top} and H_{base} listed in Table 2 for any type of cloud. Nearly all H_{top} and cloud thickness (ΔH) values over NSO are higher or deeper than those over SSO, presumably due to stronger solar radiation and stronger convection over NSO. H_{base} values basically followed their cloud-top counterparts with a couple of exceptions. These cloud macrophysical properties are closely associated with large-scale dynamic patterns and environmental conditions. By analyzing the ERA-Interim reanalysis (not shown), the 850 hPa geopotential heights show persistent westerlies with slightly higher geopotential heights over the northwest corner of the domain, which may closely relate to the higher H_{top} over NSO than over SSO. Furthermore, the boundary layer over NSO is relatively more stable than over SSO based on lower troposphere stability (LTS) analysis (12.2-15.32 K over NSO vs. 11.48-13.29 K over SSO).

When we plot the probability density functions (PDFs) of cloud *LWP*s for different types of clouds, we find that the PDFs of *LWP*s for HGH and HOM peak are less than 10 g m^{-2}. These results make physical sense because HGH clouds should not contain any liquid droplets, and most HOM clouds, especially those over SSO, should be ice phase dominant. In addition, the 10 g m^{-2} of *LWP* is close to the uncertainty of the *LWP* retrieval in Marchand et al. (2003). Therefore, this value is used as a threshold for all types of clouds, which leads to less than one percent reduction of the total samples. As shown in Fig. 4a, the *LWP*s (> 10 g m^{-2}) for all types of clouds are much higher over NSO than over SSO because the low-level and MOL clouds in the mid-latitudes contain more liquid water than those in Polar regions. The mean *LWP*s for liquid containing low-

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level and middle-level clouds over NSO, i.e. LOW, MID and HOL, range from ~130 to 150 g m^{-2}, while the mean *LWP*s for MOL and HML are two times higher (~270 g m^{-2}) than the mean *LWP* of LOW, MID and HOL. Note that the mean *LWP*s for most types of clouds over the SSO are much lower than those over the NSO, except for the LOW clouds.

The occurrence frequencies of *LWP*s (> 10 g m^{-2}) over NSO and SSO contradict their cloud *LWP* values as demonstrated in Fig. 4b. To further investigate the amount of available precipitable water vapor (*PWV*), we found that mean *PWV* values in SSO are at least 2 to 3 times less than those in NSO for same types of clouds (figure not shown). Note that the samples of MID, HGH, and HOM clouds are excluded from this study when they have *LWP*s less than 10 g m^{-2}, since these low *LWP*s are within the retrieval uncertainty of cloud *LWP* and hence may not contain any liquid cloud droplets. The higher *LWP*s, larger cloud droplets, drizzle drops and ice particles, and greater drizzling occurrence frequencies over NSO (which is discussed later) will lead to the quick dissipation of clouds over NSO. In contrast to NSO, the SSO cloud *LWP*s and particle sizes are much smaller with less drizzling events, which increases cloud lifetime relative to NSO. The 67.6 % and 90.3 % *CF*s over NSO and SSO provide strong evidence for this argument. We can draw the following conclusions by comparing the cloud macrophysical properties between NSO and SSO in Figs. 3 and 4. The LOW fraction, thickness, and *LWP* over NSO and SSO are comparable to each other. For other types of clouds, cloud thicknesses are similar to each other or slightly deeper over NSO, but the cloud *LWP*s over NSO are much larger than those over SSO, resulting in more precipitation events over NSO. As pointed out in Albrecht (1989), more precipitation events may reduce the cloud lifetime. This argument is consistent with the results shown in Figs. 2 and 3a for all clouds except for HOL. Cloud lifetimes over NSO are shorter than those over SSO, which leads to lower *CF*s over NSO than over SSO.

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Table 2 provides a summary of the mean, standard derivation, minimum and maximum for cloud boundaries, LWP and the percentage of multi-layered cloud for each cloud type over the SO. Non-contiguous (multi-layer) clouds over the SO occur very frequently, especially for HOM and HML. The *LWP* for single-layered clouds is greater than that for multi-layered clouds. The *LWP* for single-layered HML almost doubles that for multi-layered HML.

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4. Single-layered low-level clouds

As discussed in Section 3, single-layered low-level clouds (LOW) are the dominant cloud type in both northern (NSO) and southern (SSO) parts of the SO. Figs. 3 and 4 further reveal that LOW cloud type is the only one having comparable *CF*, cloud, thickness, *LWP* over both NSO and SSO. This warrants further study: Are the cloud phases, properties, and vertical and meridional variations of LOW clouds over these two regions similar to each other or significantly different?

4.1. Cloud phase

In this study, cloud boundaries are determined by combining cloud radar, ceilometer and micropulse lidar measurements at a temporal resolution of 5-min. The cloud phase, liquid water droplets or ice particles, are determined in each radar range volume. A flow chart for classifying the phases of single-layered low-level clouds is drawn in Fig. 5. The determination of warm liquid clouds is straightforward using both cloud-base (T_{base}) and -top (T_{top}) temperatures greater than 0 °C, and cloud *LWP*s greater than the threshold (10 g m^{-2}). The determination of supercooled liquid clouds is slightly complicated. When either T_{base} or T_{top} is below 0° C, and cloud *LWP*s are greater than the threshold, the radar Doppler spectrum width (*WID*) and velocity (V_d) are used for the determination of supercooled liquid water clouds. If the majority (10 seconds of original radar measurements) of *WID* within a 5-min period are less than 0.4 m s^{-1} and V_d are equal to or less than 0.0 m s^{-1} (updrafts) in the volume, then this range volume is defined as supercooled liquid clouds.

Mixed-phase clouds are determined when the median (calculated from 10 seconds of original radar measurements) of *WID* is greater than 0.4 m s^{-1} or V_d is greater than 0.0 m s^{-1} (downdrafts) due to the existence of large ice particles in the clouds. If cloud *LWP* is below the threshold, then it is defined as an ice cloud, otherwise it is defined as a mixed-phase cloud. It is worth mentioning that large ice particles, which grow through vapor deposition or rime processes, dominate the radar reflectivity and are heavier than cloud droplets. Therefore, these large ice particles not only broaden the spectrum width but also have relatively large fall speeds.

To further evaluate our classification method, we compared the classified mixed-phase and ice clouds with the micropulse lidar linear depolarization ratios (*LDR*) as an extra measure. The *LDR* ranges follow the method in Shupe et al., (2005), which are 0.11 < *LDR* < 0.15 for mixed-phase clouds, and *LDR* > 0.15 for ice clouds as listed in Table 1. Table 3a shows the quantitative comparison of the cloud phase identifications between these two classification methods. The numbers represent the counts of each matched 5-min sample, where the diagonal numbers indicate that both methods are identifying the same type of cloud phase. In general, the two methods have 89 % agreement on the phase identification. Secondly, we performed the phase classification directly from microphysical probes onboard G-1 aircraft during SOCRATES and treated them as 'ground-truth' (Mohrmann et al., 2021). Since the in-situ cloud microphysical measurements can tell us the phase of the cloud, it allows us to see the percentage variations of cloud phase, by changing integration time of in-situ sampling to mimic what the radar may observe the cloud for each range volume. Table 3b shows possible cloud phase partitionings that may be detected by cloud radar. As sampling time increases from 1 second to 30 seconds, more mixed-phase clouds and fewer single-phase clouds can be observed.

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Figure 6 shows the determination of mixed-phase and ice clouds through combined measurements of radar reflectivity and spectrum width, lidar LDR and backscatter, and cloud LWP. For the classified ice clouds, cloud LWPs are lower than 10 gm^{-2} (Fig. 6f), most of the Doppler spectrum widths range from 0.08 to 0.16 m s^{-1} (Fig. 6b) and the *LDR* ratios (Fig. 6d) can be greater than ~0.15, representing a narrow range of ice particle size distribution with higher *LDR* ratios. For the classified mixed-phase clouds, cloud LWPs are greater than 10 gm^{-2} and most of the Doppler spectrum widths range from 0.15 to 0.5 m s^{-1}, representing a broad particle size distribution resulting from the mixture of liquid droplets and ice particles. An interesting result occurs where both *LDR* signals (>0.2) and LWPs are much higher during the drizzling periods (Fig. 6a), indicating a mixed-phase cloud with cloud droplets within the cloud layer and large liquid drizzle drops and ice crystals below cloud base.

Based on the Doppler velocity, the mode values for both mixed-phase and ice clouds occur at ~ 0.5 m s^{-1}, where the ice particles are dominant in both types of clouds. The broader particle size distribution with lower *LDR* ratios for mixed-phase clouds and narrower particle size distribution with higher *LDR* ratios for ice clouds further corroborate that the classified results from this study are consistent with the traditional micropulse lidar *LDR* method.

It is important to note that the micropulse lidar signals are usually attenuated and cannot provide a meaningful signal when the liquid cloud layer is thicker than a couple of hundred meters (Sassen, 1991). Arctic mixed-phase clouds are typical with the liquid-dominant layer on the top of the mixed-phase clouds and the ice-dominant layer underneath. The ceilometer-derived cloud-base height represents the base of the liquid-dominant layer near the cloud top, while MPL-derived cloud-base height represents the base of the lower ice-dominant layer (Qiu et al., 2015; Shupe, 2007; Shupe et al., 2005). Over the Arctic, the micropulse lidar signals can penetrate through the

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ice-dominant layer to the liquid-dominant layer. However, the mixed-phase clouds over the Southern Ocean are totally different from those over the Arctic region: they are well mixed (liquid droplets and ice particles) from cloud base to cloud top, which is found in this study. Thus, the micropulse lidar signals can be attenuated in the mixed-phase clouds over the Southern Ocean. Statistical results show that 43 % of micropulse lidar signals were attenuated during MARCUS compared to our classified results.

This classification method is further supported by the onboard cloud radar measurements during the Southern Ocean Clouds Radiation Aerosol Transport Experimental Study (SOCRATES, not shown). In that campaign, the reflectivity measurements were usually greater, and the spectrum widths were much wider when the aircraft observed large ice particles compared to the time periods when liquid cloud droplets were observed. Although the wider spectrum widths might be caused by Doppler broadening of the moving aircraft, further analysis shows that the onboard radar sends the signals (assuming the time of transmitted and received signals is short enough comparing to aircraft speed) in the perpendicular to the movement of the aircraft, that is, there is no relative movement between radar signals and clouds. Thus, the onboard radar spectrum width measurements should be not significantly impacted by Doppler broadening (relative movement in the same direction).

In this study, a total of 6,934 5-min single-layered low-level cloud samples were determined using our classificaiton method, including 697 liquid cloud samples, 3,777 mixed, 1,205 ice, and 1,255 'OTHER' clouds. The category of 'OTHER' clouds represents more than one phase in each column. Note that though the ‘OTHER’ is also mixed-phase cloud, it has different vertical distribution of liquid compared to the ‘mixed’ cloud. It is also worth mentioning that about 5.5 %

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of single-layered low-level cloud phases cannot be determined when the radar measurements were not available during MARCUS, those were not accounted to “OTHER’.

Figure 7 (upper panel) shows the drizzling status for each categorized cloud type, i.e., no rain (yellow-green), virga (brown) and rain (navy blue). The definition of drizzling status follows the method in Wu et al., (2015, 2017) where there are radar reflectivity measurements below the ceilometer/lidar determined cloud base. The major difference for drizzles in the studies of Wu et al. (2015, 2017) and this study is that drizzle is liquid phase at ARM ENA site but could be both liquid and ice phases in this study.

The percentages shown below the x-axis represent the portion of drizzling status in each type of clouds, such as liquid, mixed-phase, ice and 'OTHER' clouds. Figure 7 (bottom panel) also shows the percentages and vertical distributions of classified liquid, mixed-phase, ice, and 'OTHER' clouds for each column in the single-layered low-level clouds, represented by different colors. After classification, the samples in each category are sorted by their H_{top}. In detail, Figure 7 demonstrates that the mixed-phase clouds dominate the single-layered low-level cloud category with an occurrence frequency of 54.5 %. The 'OTHER' and ice clouds have similar occurrence frequencies of 18.1 % and 17.4 %, respectively, while the liquid clouds have the lowest occurrence frequency of 10.1 %. The liquid-topped mixed-phase clouds (included in 'OTHER'), which frequently occur in the Arctic region (Qiu et al., 2015), are rarely found over the SO. The existence of ice particles in mixed-phase clouds should strongly depend on the distribution of ice nuclei (IN), whereas spatially unevenly distributed IN may result in the OTHER type of clouds.

Based on the results in Fig. 7, we draw the following conclusions. Most of the ice clouds are without icy precipitation, and the percentages with virga and precipitation below the cloud base are 12 % and 15 %, respectively. The percentages of non-drizzling, virga and drizzling mixed-

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phase clouds are 50 %, 21 %, and 29 %. The liquid and 'OTHER' clouds have similar percentages, they are 36 %, 25 % and 39 % for liquid clouds, and 35 %, 22 % and 44 % for 'OTHER' clouds. For liquid and 'OTHER' clouds, the drizzling frequencies are independent of H_{top}. In contrast, for mixed-phase and ice clouds, the drizzling frequencies strongly depend on H_{top}, i.e., higher drizzling frequencies occur mostly at higher H_{top}.

The properties of single-layered low-level clouds are summarized in Table 4. The liquid clouds have the lowest H_{base} and H_{top} but more available water vapor than other types of clouds. Since the 'OTHER' clouds are a transitional stage among mixed-phase, liquid and ice clouds, they have the highest H_{top}, deepest cloud layer and largest *LWP*. The ice clouds occur in relatively dry environments and have the highest H_{base} at 1.218 km and thinnest cloud layer. The cloud variables for mixed-phase clouds fall between Liquid and "OTHER". Since *LWP*s in mixed-phase clouds have larger standard deviation, which implies that SLW is more common at higher *LWP*s and ice is more common at lower *LWP*s.

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4.2. Meridional variations of cloud properties

Figure 8 shows the meridional variation in single-layered low-level cloud properties during MARCUS. As illustrated in Fig. 8a, the meridional distributions of H_{base}, H_{top} and ΔH are nearly independent of latitude, however, their corresponding temperatures (T_{base} and T_{top}) increased about 8 K from 69 °S to 43 °S, though there were slight fluctuations. These results suggest that the cloud and sea surface temperatures have minimal impact on the cloud boundaries over the SO, which is consistent with the findings in McFarquhar et al. (2016). The meridional variation of *LWP*s follows those of T_{base} and T_{top}, with an increasing trend from south to north. It is important to point out that a big drop in *LWP* at ~50 °S results from fewer occurrences of low-level clouds there, indicating that the cloud samples at some latitudes are not statistically significant. The atmospheric *PWV*

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increased dramatically from ~ 5 mm at 69 $^{\circ}$S to ~18 mm at 43 $^{\circ}$S, presumably due to increased sea surface and atmospheric temperatures.

Figure 9 shows the latitudinal and meridional distributions of categorized liquid, mixed-phase, ice and 'OTHER' in single-layered low-level clouds over the SO during MARCUS. Each circle represents the exact location and time along the ship track. Mixed-phase clouds occurred everywhere over the SO during the MARCUS field campaign and became dominant in November, December and February. Liquid clouds dominated in March, while ice clouds dominated in January. The 'OTHER' clouds are a kind of transitional phase falling in between the mixed-phase and ice/liquid clouds because there are no stand-alone occurrences in any month during MARCUS.

4.3 Vertical distribution of cloud properties

The vertical distributions of classified liquid, mixed-phase, and ice clouds in LOW category are presented in Figs. 10-12. The focus of this section will be comparisons of cloud macrophysical properties between the north (NSO) and south (SSO) regions of the domain. Figure 10a shows the vertical distributions of liquid clouds, which were capped at ~ 1.6 km, mostly in the marine boundary layer. The vertical occurrence frequencies are up to 27 % over NSO, while they were less than 4 % over SSO, i.e., liquid clouds occurred fairly often over the mid-latitude region, but very few occurred over the Polar region. On the contrary, the occurrence frequencies of mixed-phase clouds between NSO and SSO are opposite to liquid clouds as illustrated in Fig. 10b, though the differences are not so obvious. Mixed-phase clouds increased with altitude until ~1.6 km, then decreased monotonically towards 3 km. The highest frequencies were ~37 % at 0.6 km over SSO and ~27 % at 1.5 km over NSO. The vertical distributions of ice clouds are similar to those of mixed-phase clouds (Fig. 10c), however, there were no significant differences between NSO and SSO. It is worth mentioning that the vertical distributions of mixed-phased clouds over SO are

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quite different to those from DOE ARM Northern Slope Alaska (NSA) site, where the low-level mixed-phase clouds are commonly featured with a liquid-topped layer. (e.g., Qiu et al., 2015).

To further investigate the vertical distributions of classified liquid, mixed-phase, and ice clouds over NSO and SSO, we plot the normalized vertical distributions (cloud base as 0, cloud top as 1) of radar reflectivity, Doppler velocity and spectrum width in Figs. 11 and 12, respectively. In this study, the threshold of -50 dBZ was used to determine the cloud boundary over the SO instead of the threshold of -40 dBZ radar reflectivity used at the ARM ENA site (Dong et al., 2014). If we used the threshold of -40 dBZ over the SO, then there would be only 73% cloud samples available. If we used the threshold of -50 dBZ, then we would have 90.4% cloud samples, which gained additional 17.4% on top of the -40 dBZ threshold. About 9.6% of radar reflectivities during MARCUS are lower than -50 dBZ for all LOW cloud samples, but without ceilometer and MPL lidar signials. Thus these 9.6% cloud samples were eliminated in Figs 11-12.

Figures 11a-11c represent the normalized vertical distributions of radar reflectivity, Doppler velocity and spectrum width of liquid clouds. Liquid clouds had the lowest reflectivity near the cloud top because of cloud-top entrainment., The reflectivity had a nearly constant median value of ~ -22 dBZ from cloud top height (~ 0.8 for normalized height) of the cloud layer to the cloud base. Most of the reflectivities were less than -15 dBZ, which is a threshold to distinguish cloud droplets and drizzle drops in each radar range volume (Wu et al., 2020). Most of the Doppler velocities were greater than 0.0 m s^{-1}, indicating that downwelling motion is dominant in liquid clouds. The profiles of Doppler velocity and spectrum width increased smoothly from the cloud top to base, suggesting that larger cloud droplets and broader size distributions exist near the cloud base, which is attributable to more drizzle drops near the cloud base as illustrated in Fig. 7.

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The vertical distributions of mixed-phase clouds in Figs. 11d-11f are similar to those of liquid clouds. The more occurrences of larger reflectivity measurements and larger median values of spectrum width near the cloud base are most likely due to the presence of moderate ice particles and/or drizzle drops. The nearly same median values of reflectivity, Doppler velocity and spectrum width (but slightly larger standard deviations in each level in mixed-phase clouds) in both liquid and mixed-phase clouds suggest that the ice particle sizes in mixed-phase clouds are comparable to cloud droplets and drizzle drops. The nearly uniform vertical distributions of Doppler velocity and spectrum width indicate well-mixed liquid cloud droplets and ice particles throughout the cloud layer in the mixed-phase clouds over NSO.

Compared to liquid and mixed-phase clouds, ice clouds had much lower reflectivities and narrower spectrum width as shown in Figs. 11g-11i. Almost all reflectivity measurements were less than -25 dBz with a median value of -35 dBz at the cloud base, resulting from small or moderate ice particles but much lower concentration. A nearly constant Doppler velocity within the cloud layer further supports the discussion of mixed-phase clouds above, i.e., the ice particle sizes are independent of cloud height and comparable to liquid cloud droplets in the low-level clouds over the SO. Because there are no mechanisms for growing large ice particles in such shallow ice clouds, the accretion process cannot take place. From the statistical results in Fig. 7, these ice particles have relatively little chance to become virga or raindrops and usually dissipate or transition to other types of clouds.

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Since there are not enough liquid cloud samples over the Polar region, only the mixed-phase and ice clouds results are shown in Fig. 12. Compared to the vertical distributions of ice clouds over NSO, the median values of reflectivity and Doppler spectrum width over SSO were lower and narrower, indicating a lack of large ice particles in the Polar region. The small ice particles in

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the Polar region were also reflected in their mixed-phase clouds. Compared to the vertical distributions of the mixed-phased clouds over NSO, the median values of reflectivity and Doppler spectrum width over SSO were dramatically lower (-35 dBz at SSO vs. -22 dBz at NSO; 0.25 m s^{-1} at SSO vs. 0.32 m s^{-1} at NSO). Figure 12 illustrates that the ice particle sizes over SSO are smaller, their size distributions are narrower than those over NSO, indicative of lack of large ice particles over SSO.

5. Summary and Conclusions

In this study, we presented the statistical results of clouds over the Southern Ocean (SO), and its northern (NSO) and southern (SSO) parts during MARCUS Intensive observational period (IOP). We used the method developed in Xi et al., (2010) to calculate the occurrence frequencies of different types of clouds and their corresponding cloud macrophysical properties. We developed a new method to classify liquid, mixed-phased, and ice clouds in the single-layered low-level clouds as well as their corresponding drizzling status. Lastly, we explored the meridional and vertical distributions of these classified cloud properties. Analysis of the MARCUS cloud phase and macrophysical properties has yielded the following conclusions.

1) The total cloud fractions (*CF*s) were 77.9 %, 67.6 %, and 90.3 % for the entire domain, NSO and SSO, respectively, indicating that 22.7 % more clouds occurred in the Polar region than in the mid-latitude region. The SSO had more clouds under 7 km, while the NSO had more clouds above 7 km. Below 3 km, the occurrence frequencies of clouds over NSO decrease dramatically from 37 % at an altitude of ~0.7 km to 16 % at 3 km, which is similar to the vertical distributions of the low-level clouds over some Northern Hemisphere mid-latitude regions, such as Eastern North Atlantic.

2) The single-layered low-level (LOW), deep connective or multi-layered (HML), and MOL clouds are the three dominant types of clouds over the SO. Comparing the clouds between NSO and SSO, all types of clouds in SSO are higher than those in NSO except HOL. The LOW fraction, thickness, LWP over both NSO and SSO are comparable to each other. The mean *LWP*s for LOW, MID and HOL clouds over NSO, range from ~130 to 150 g m^{-2}, while the mean NSO *LWP*s (~270 g m^{-2}) for MOL and deep convective clouds (HML) are two times higher than the same types of clouds over SSO. The mean *LWP*s of clouds over SSO are much lower than the *LWP*s over NSO. Over the Southern Ocean, the single-layered or contiguous clouds usually have higher LWP than their counterparts of multi-layered or non-contiguous clouds. There are more non-contiguous HML and HOM than contiguous ones.

3) A new method was developed to classify liquid, mixed-phase and ice clouds in the single-layered low-level clouds (LOW) based on comprehensive ground-based observations. The mixed-phase clouds are dominant in the LOW cloud category with an occurrence frequency of 54.5 %. The 'OTHER' and ice clouds had similar occurrence frequencies of 18.1 % and 17.4 %, respectively, while the liquid clouds had the least occurrence frequency of 10.1 %. The percentages of non-drizzling, virga and drizzling for mixed-phase clouds were 50 %, 21 %, and 29 %, and the drizzling frequencies of mixed-phase clouds strongly depend on H_{top}, that is, higher drizzling frequencies occurred mostly at higher H_{top}.

4) The meridional distributions of H_{base}, H_{top} and ΔH are nearly independent on latitude, however, their corresponding temperatures increased about 8 K from 69 °S to 43 °S. The meridional variation of *LWP*s mimics that of cloud temperatures, having an increasing trend from south to north. The mean *PWV* increased dramatically from ~ 5 mm at 69 °S to ~18 mm at 43 °S due to increased sea surface and atmospheric temperatures. More liquid clouds occurred over NSO

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but very few occurred over SSO, whereas more mixed-phase clouds occurred over SSO than over NSO. There were no significant differences in ice clouds occurrences between NSO and SSO.

5) The nearly same median values of reflectivity, Doppler velocity and spectrum width in both liquid and mixed-phase clouds over NSO suggest that the ice particle sizes in mixed-phase clouds are comparable to cloud droplets and drizzle drops. The uniform vertical distributions of Doppler velocity and spectrum width suggest well-mixed liquid cloud droplets and ice particles throughout the cloud layer in the mixed-phase clouds over NSO, which are quite different from those over the DOE ARM NSA site where the liquid-topped mixed-phase low-level clouds are common. The median values of reflectivity and Doppler spectrum width over SSO were lower and narrower than those over NSO, indicating lack of large ice particles in the polar region.

These results provide comprehensive statistical properties of all clouds over the SO during MARCUS, including the occurrence frequencies of different types of clouds and their corresponding cloud macrophysical properties. We also examined the meridional and vertical distributions of the classified cloud properties. These statistics can be used as a ground truth to evaluate satellite retrieved cloud properties and model simulations over the SO. The results of this study will help to advance our understanding of the clouds over the SO, which may lead to improved model simulations, as well as better representation of global climate.

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Data availability. Data used in this study can be accessed from the DOE ARM's Data Discovery at https://adc.arm.gov/discovery/

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Author contributions. The idea of this study is discussed by BX, XD, and XZ. BX and XZ performed the analyses and BX wrote the manuscript. BX, XD, XZ and PW participated in scientific discussions and provided substantial comments and edits on the paper.

Competing interests. The authors declare that they have no conflict of interest.

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References.

Albrecht, B. A.: Aerosols, cloud microphysics, and fractional cloudiness, Science, doi:10.1126/science.245.4923.1227, 1989.

Bodas-Salcedo, A., Mulcahy, J. P., Andrews, T., Williams, K. D., Ringer, M. A., Field, P. R. and Elsaesser, G. S.: Strong Dependence of Atmospheric Feedbacks on Mixed-Phase Microphysics and Aerosol-Cloud Interactions in HadGEM3, J. Adv. Model. Earth Syst., doi:10.1029/2019MS001688, 2019.

Bony, S., Stevens, B., Frierson, D. M. W., Jakob, C., Kageyama, M., Pincus, R., Shepherd, T. G., Sherwood, S. C., Siebesma, A. P., Sobel, A. H., Watanabe, M. and Webb, M. J.: Clouds, circulation and climate sensitivity, Nat. Geosci., doi:10.1038/ngeo2398, 2015.

Chubb, T. H., Jensen, J. B., Siems, S. T. and Manton, M. J.: In situ observations of supercooled liquid clouds over the Southern Ocean during the HIAPER Pole-to-Pole Observation campaigns, Geophys. Res. Lett., doi:10.1002/grl.50986, 2013.

Dong, X., Xi, B., Kennedy, A., Minnis, P. and Wood, R.: A 19-month record of marine aerosol-cloud-radiation properties derived from DOE ARM mobile facility deployment at the Azores. Part I: Cloud fraction and single-layered MBL cloud properties, J. Clim., doi:10.1175/JCLI-D-13-00553.1, 2014.

Hu, Y., Rodier, S., Xu, K. M., Sun, W., Huang, J., Lin, B., Zhai, P. and Josset, D.: Occurrence, liquid water content, and fraction of supercooled water clouds from combined CALIOP/IIR/MODIS measurements, J. Geophys. Res. Atmos., doi:10.1029/2009JD012384, 2010.

Intrieri, J. M., Fairall, C. W., Shupe, M. D., Persson, P. O. G., Andreas, E. L., Guest, P. S. and Moritz, R. E.: An annual cycle of Arctic surface cloud forcing at SHEBA, J. Geophys. Res. Ocean., doi:10.1029/2000jc000439, 2002.

Klein, S. A., Hall, A., Norris, J. R. and Pincus, R.: Low-Cloud Feedbacks from Cloud-Controlling Factors: A Review, Surv. Geophys., doi:10.1007/s10712-017-9433-3, 2017.

Kalesse, H., de Boer, G., Solomon, A., Oue, M., Ahlgrimm, M., Zhang, D., Shupe, M. D., Luke, E. and Protat, A.: Understanding rapid changes in phase partitioning between cloud liquid and ice in stratiform mixed-phase clouds: An arctic case study, Mon. Weather Rev., doi:10.1175/MWR-D-16-0155.1, 2016.

Lang, F., Huang, Y., Siems, S. T. and Manton, M. J.: Characteristics of the Marine Atmospheric Boundary Layer Over the Southern Ocean in Response to the Synoptic Forcing, J. Geophys. Res. Atmos., doi:10.1029/2018JD028700, 2018.

Listowski, C., Delanoë, J., Kirchgaessner, A., Lachlan-Cope, T. and King, J.: Antarctic clouds, supercooled liquid water and mixed phase, investigated with DARDAR: Geographical and seasonal variations, Atmos. Chem. Phys., doi:10.5194/acp-19-6771-2019, 2019.

Mace, G. G., Zhang, Q., Vaughan, M., Marchand, R., Stephens, G., Trepte, C. and Winker, D.: A description of hydrometeor layer occurrence statistics derived from the first year of merged Cloudsat and CALIPSO data, J. Geophys. Res. Atmos., doi:10.1029/2007JD009755, 2009.

Mace, G. G. J. and Protat, A.: Clouds over the Southern Ocean as observed from the R/V investigator during CAPRICORN. Part I: Cloud occurrence and phase partitioning, J. Appl. Meteorol. Climatol., doi:10.1175/JAMC-D-17-0194.1, 2018.

Mace, G. G., Protat, A., Humphries, R. S., Alexander, S. P., McRobert, I. M., Ward, J., Selleck, P., Keywood, M. and McFarquhar, G. M.: Southern Ocean Cloud Properties Derived From

CAPRICORN and MARCUS Data, J. Geophys. Res. Atmos., doi:10.1029/2020JD033368, 2021.

Marchand, R., Ackerman, T., Westwater, E. R., Clough, S. A., Cady-Pereira, K. and Liljegren, J. C.: An assessment of microwave absorption models and retrievals of cloud liquid water using clear-sky data, J. Geophys. Res. Atmos., doi:10.1029/2003jd003843, 2003.

Marchand, R., Wood, R., Bretherton, C., McFarquhar, G., Protat, A., Quinn, P., Siems, S., Jakob, C., Alexander, S., Weller, B.: The Southern Ocean Clouds, Radiation Aerosol Transport Experimental Study (SOCRATES), whitepaper available from http://www.atmos.washington.edu/socrates/SOCRATES_white_paper_Final_Sep29_2014.pdf, 2014.

Mccoy, D. T., Hartmann, D. L. and Grosvenor, D. P.: Observed Southern Ocean cloud properties and shortwave reflection. Part II: Phase changes and low cloud feedback, J. Clim., doi:10.1175/JCLI-D-14-00288.1, 2014.

McFarquhar, G., Bretherton, C., Alexander, S., DeMott, P., Marchand, R., Protat, A., Quinn, P., Siems, S., Weller, R., Wood, R.: Measurements of Aerosols, Radiation, and Clouds over Sothern Ocean (MARCUS) Science Plan, DOE ARM Climate Research Facility., DOE/SC-ARM-16-011, available at: http://arm.gov/publications/programdocs/doe-sc-arm-16-011.pdf, 2016.

McFarquhar, G. M., Bretherton, C. S., Marchand, R., Protat, A., DeMott, P. J., Alexander, S. P., Roberts, G. C., Twohy, C. H., Toohey, D., Siems, S., Huang, Y., Wood, R., Rauber, R. M., Lasher-Trapp, S., Jensen, J., Stith, J. L., Mace, J., Um, J., Järvinen, E., Schnaiter, M., Gettelman, A., Sanchez, K. J., McCluskey, C. S., Russell, L. M., McCoy, I. L., Atlas, R. L., Bardeen, C. G., Moore, K. A., Hill, T. C. J., Humphries, R. S., Keywood, M. D., Ristovski, Z., Cravigan, L., Schofield, R., Fairall, C., Mallet, M. D., Kreidenweis, S. M., Rainwater, B., D'Alessandro, J., Wang, Y., Wu, W., Saliba, G., Levin, E. J. T., Ding, S., Lang, F., Truong, S. C. H., Wolff, C., Haggerty, J., Harvey, M. J., Klekociuk, A. R., and McDonald, A.: Observations of Clouds, Aerosols, Precipitation, and Surface Radiation over the Southern Ocean: An Overview of CAPRICORN, MARCUS, MICRE, and SOCRATES, B. Am. Meteorol. Soc., 102, E894-E928, 10.1175/BAMS-D-20-0132.1, 2021.

Mohrmann, J., Finlon, J., Atlas, R., Lu, J., Hsiao, I., Wood, R.: University of Washington Ice-Liquid Discriminator single particle phase classifications and 1 Hz particle size distributions/heterogeneity estimate, Version 1.0. UCAR/NCAR - Earth Observing Laboratory., doi:10.26023/PA5W-4DRX-W50A, Last Access: Nov 01, 2021

Morrison, H., De Boer, G., Feingold, G., Harrington, J., Shupe, M. D. and Sulia, K.: Resilience of persistent Arctic mixed-phase clouds, Nat. Geosci., doi:10.1038/ngeo1332, 2012.

Muradyan, P. and Coulter, R.: Micropulse Lidar (MPL) Instrument Handbook. DOE ARM Climate Research Facility, DOE/SC-ARM-TR-019, 2020. Available at: https://www.arm.gov/publications/tech_reports/handbooks/mpl_handbook.pdf, last access: 25 March 2022.

Qiu, S., Dong, X., Xi, B. and Li, J. L. F.: Characterizing Arctic mixed-phase cloud structure and its relationship with humidity and temperature inversion using ARM NSA observations, J. Geophys. Res., doi:10.1002/2014JD023022, 2015.

Rémillard, J., Kollias, P., Luke, E., and Wood, R.: Marine Boundary Layer Cloud Observations in the Azores, J. Climate, 25, 7381-7398, 10.1175/JCLI-D-11-00610.1, 2012.

Sassen, K.: The polarization lidar technique for cloud research: a review and current assessment, Bull. - Am. Meteorol. Soc., doi:10.1175/1520-0477(1991)072<1848:TPLTFC>2.0.CO;2, 1991.

Shupe, M. D., Uttal, T. and Matrosov, S. Y.: Arctic cloud microphysics retrievals from surface-based remote sensors at SHEBA, J. Appl. Meteorol., doi:10.1175/JAM2297.1, 2005.

Shupe, M.: A ground-based multisensory cloud phase classifier, Geophys. Res. Lett., doi:10.1029/2007GL031008, 2007.

Stanfield, R. E., Dong, X., Xi, B., Kennedy, A., Del Genio, A. D., Minnis, P. and Jiang, J. H.: Assessment of NASA GISS CMIP5 and Post-CMIP5 Simulated Clouds and TOA Radiation Budgets Using Satellite Observations. Part I: Cloud Fraction and Properties, J. Clim., doi:10.1175/jcli-d-13-00558.1, 2014.

Stanfield, R. E., Dong, X., Xi, B., Del Genio, A. D., Minnis, P., Doelling, D. and Loeb, N.: Assessment of NASA GISS CMIP5 and post-CMIP5 simulated clouds and TOA radiation budgets using satellite observations. Part II: TOA radiation budget and CREs, J. Clim., doi:10.1175/JCLI-D-14-00249.1, 2015.

Stocker, T. F., Qin, D., Plattner, G. K., Tignor, M. M. B., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V. and Midgley, P. M.: Climate change 2013 the physical science basis: Working Group I contribution to the fifth assessment report of the intergovernmental panel on climate change., 2013.

Toto, T. and Jensen, M.: Interpolated Sounding and Gridded Sounding Value-Added Products. DOE ARM Climate Research Facility, DOE/SC-ARM-TR-183, 2016. Available at: https://www.arm.gov/publications/tech_reports/doe-sc-arm-tr-183.pdf, last access: 25 March 2022.

Uin, J.: Cloud Condensation Nuclei Particle Counter Instrument Handbook. DOE ARM Climate Research Facility, DOE/SC-ARM-TR-168, 2016. Available at: https://www.arm.gov/publications/tech_reports/handbooks/ccn_handbook.pdf, last access: 25 March 2022.

Wu, P., Dong, X. and Xi, B.: Marine boundary layer drizzle properties and their impact on cloud property retrieval, Atmos. Meas. Tech., doi:10.5194/amt-8-3555-2015, 2015.

Wu, P., Dong, X., Xi, B., Liu, Y., Thieman, M. and Minnis, P.: Effects of environment forcing on marine boundary layer cloud-drizzle processes, J. Geophys. Res., doi:10.1002/2016JD026326, 2017.

Wu, P., Dong, X. and Xi, B.: A climatology of marine boundary layer cloud and drizzle properties derived from ground-based observations over the azores, J. Clim., doi:10.1175/JCLI-D-20-0272.1, 2020.

Xi, B., Dong, X., Minnis, P. and Khaiyer, M. M.: A 10 year climatology of cloud fraction and vertical distribution derived from both surface and GOES observations over the DOE ARM SPG site, J. Geophys. Res. Atmos., doi:10.1029/2009JD012800, 2010.

Zelinka, M. D., Myers, T. A., McCoy, D. T., Po-Chedley, S., Caldwell, P. M., Ceppi, P., Klein, S. A. and Taylor, K. E.: Causes of Higher Climate Sensitivity in CMIP6 Models, Geophys. Res. Lett., doi:10.1029/2019GL085782, 2020.

Zhao, L., Zhao, C., Wang, Y., Wang, Y. and Yang, Y.: Evaluation of Cloud Microphysical Properties Derived From MODIS and Himawari-8 Using In Situ Aircraft Measurements Over the Southern Ocean, Earth Sp. Sci., doi:10.1029/2020EA001137, 2020.

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Table 1. ARM AMF2 instruments and their corresponding measurements and uncertainties used in this study

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Parameter	Instruments/ Methods	Uncertainty	References
Cloud-base height	Ceilometer/MPL	15 m	Rémillard et al., 2012
Cloud-top height	95 GHz cloud radar	43 m	Rémillard et al., 2012
Cloud-base and -top temps	Radiosonde sounding	$0.2\ ^{\circ}C$	Toto and Jensen, 2016
Profiles of reflectivity, Doppler velocity and Spectra	W-band ARM Cloud Radar (WACR), 95 GHz	Sensitivity: -50 dBZ at 2 km	Rémillard et al., 2012
linear depolarization ratios (*LDR*) and backscatter	Micropulse lidar, MPL Liquid: LDR<0.11 Mix: 0.11<LDR<0.15 Ice: LDR> 0.15		Shupe et al. 2005 Muradyan and Coulter, 2020
Cloud LWP	Microwave radiometer	~15-30 $g\ m^{-2}$ Physical retrieval	Marchand et al., 2003
CCN and aerosol properties	Aerosol Observing System	1-min resolution; Uncertainties < 10%	Uin, 2016

Table 2. Mean, standard deviation, minimum and maximum cloud-base heights (H_{base}), -top heights (H_{top}), and *LWP*s (all samples, single-layered, multilayered) of all seven types of clouds over the SO. All cloud heights have a unit of kilometer, and *LWP* has a unit of g m^{-2}

	LOW	**MID**	**MOL**	**HGH**	**HOM**	**HML**	**HOL**
H_{base} ± std min, max	0.92 ± 0.57 0.06, 2.86	4.14 ± 0.61 3.00, 5.84	1.37 ± 0.96 0.06, 5.27	8.51 ± 2.23 6.00, 18.67	4.70 ± 0.80 3.01, 7.72	1.22 ± 0.98 0.06, 7.81	1.14 ± [illegible] 0.07, 10.37
H_{top} ± std min, max	1.62 ± 0.63 0.29, 3.0	4.88 ± 0.68 3.17, 6.0	4.29 ± 0.89 1.39, 5.99	9.75 ± 2.13 6.20,18.79	7.93 ± 1.27 5.47, 17.98	7.81 ± 1.35 3.62,17.38	8.93 ± 1.66 1.79, 17.56
***LWP* ± std** Max *LWP*	122.4±134.2 1470.8	86.7±124.5 501.1	168.7±236.7 1937.1	/ /	40.9±40.8 345.7	169.2±238.4 1819.3	129.8±202. 1785.2
***LWP* ± std (single layer)** max	126.6±138.1 1470.8	88.7±128.9 501.1	193.1±271.9 1937.1	/ /	48.7±51.7 345.7	270.8±349.5 1819.3	/ /
LWP ± std (multi-layer) max	96.2±103.4 842.3	77.2±109.2 305.6	139.0±180.7 1830.2	/ /	32.3±21.3 86.8	148.4±202.4 1690.7	129.8±202. 1785.2
Multi-layer percentage %	18.1	39.6	50.0	44.9	73.1	77.7	100

***** The definition of the cloud types as follow: LOW (H_{base} and $H_{top} \leq 3\ km$); MID ($H_{base} > 3\ km$ and $H_{top} \leq 6\ km$); HGH ($H_{base} > 6\ km$); MOL ($H_{base} < 3\ km$ and $H_{top} \leq 6\ km$); HOM ($3\ km < H_{base} < 6\ km$ and $H_{top} > 6\ km$); HML ($H_{base} < 3\ km$, $H_{top} \geq 6\ km$ with a MID layer); HOL (LOW and HGH appear at the same time).

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Table 3a. Comparison of cloud phase identifications between our classification method and Shupe et al. (2005) method in each 5-min measurements, the unit is number of 5-min samples

Shupe \this study	Liquid (this study)	Mixed-phase (this study	Ice (this study)
Liquid	**468**	**490**	**0**
Mixed-phase	**98**	**3840**	**0**
Ice	**81**	**0**	**1195**

*Numbers denote the cloud sample classifications between two methods. For example, the number 98 denote a total of 98 samples are classified as Mixed-phase using Shupe's method, while are classified as Liquid using this study's method.

Table 3b. The cloud phase partitioning from CDP and 2DS during SOCRATES

Phase partitioning	**1 second**	**10 seconds**	**30 seconds**
Samples #	**27,280**	**2,255**	**836**
Liquid, %	**58.8**	**26.2**	**18.8**
Mixed-phase, %	**38.9**	**69.1**	**77.0**
Ice, %	**2.3**	**4.7**	**4.2**

Note that Cloud Droplet Probe (CDP) measures particle size from 2 to 50 um in diameter; Two-Dimensional Stereo Probe (2DS) measures particle size from 50 to 5000 um in diameter.

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Table 4. Liquid, mixed, ice and OTHER phases of cloud properties within the single-layered low-level clouds

Phase	Samples	H_{base}, km	H_{top}, km	ΔH, km	*LWP*, g m^{-2}	*PWV*, mm
Liquid	697	0.424±0.204	1.327±0.242	0.903	113.6±90.1	15.7±3.5
Mixed	3777	0.834±0.465	1.434±0.617	0.587	119.7±136.6	8.9±5.0
Ice	1205	1.218±0.635	1.737±0.651	0.519	0	8.4±4.5
OTHER	1255	0.700±0.454	1.774±0.571	1.074	141.9±137.5	11.4±5.9

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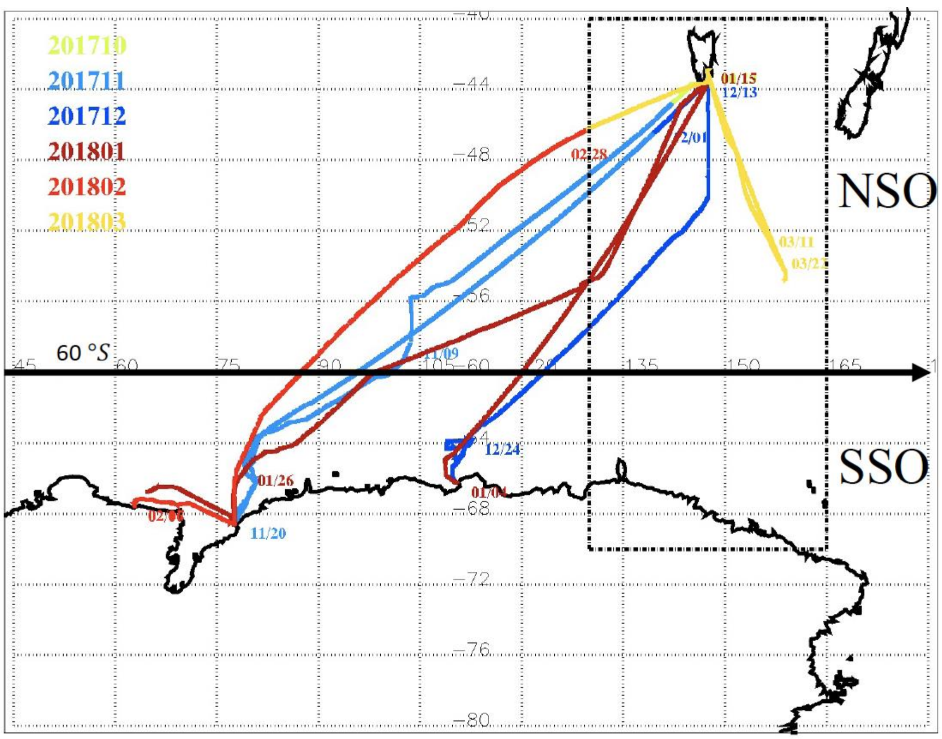

Figure 1. Shiptrack measurements between Hobart, Australia and Antarctica. Different colors represent different month's shiptracks from Oct. 29, 2017 to Mar. 23, 2018 during MARCUS. Along the shiptracks, the study domain is separated into northern (NSO) and southern (SSO) parts of the Southern Ocean with a demarcation line of 60 °S in order to study the clouds over the mid-latitudes (North of 60 °S) and Polar region (South of 60 °S). The black dotted rectangle represents the SOCRATES study domain. Some of the dates are labeled along the shiptracks, indicating the direction of the ship.

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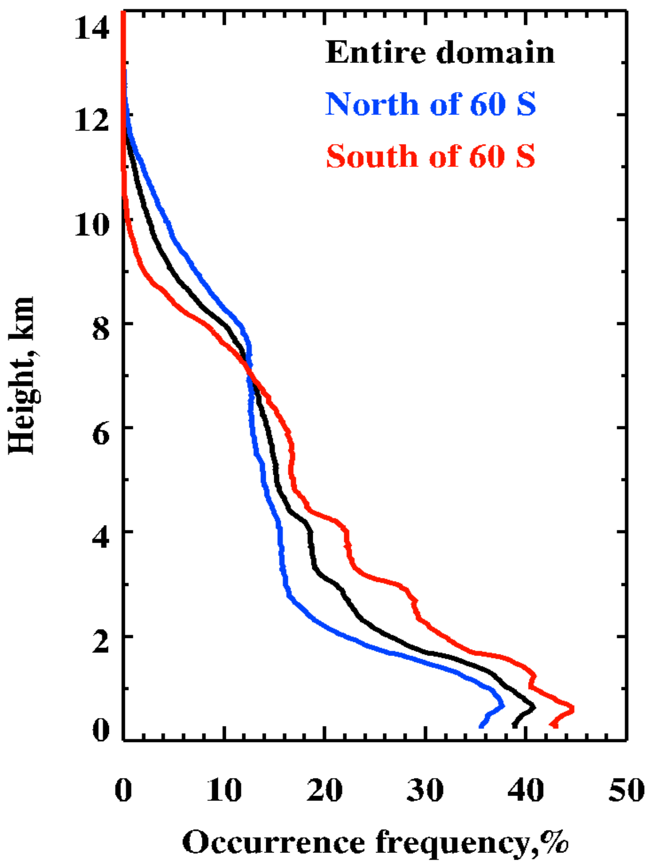

Figure 2. Mean vertical distributions of total clouds derived from ARM radar-lidar observations with a 5-min temporal resolution and a 30-m vertical resolution during MARCUS.

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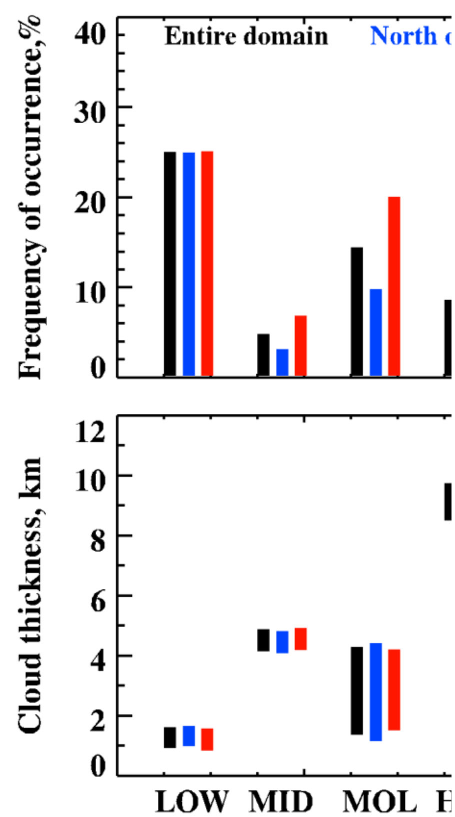

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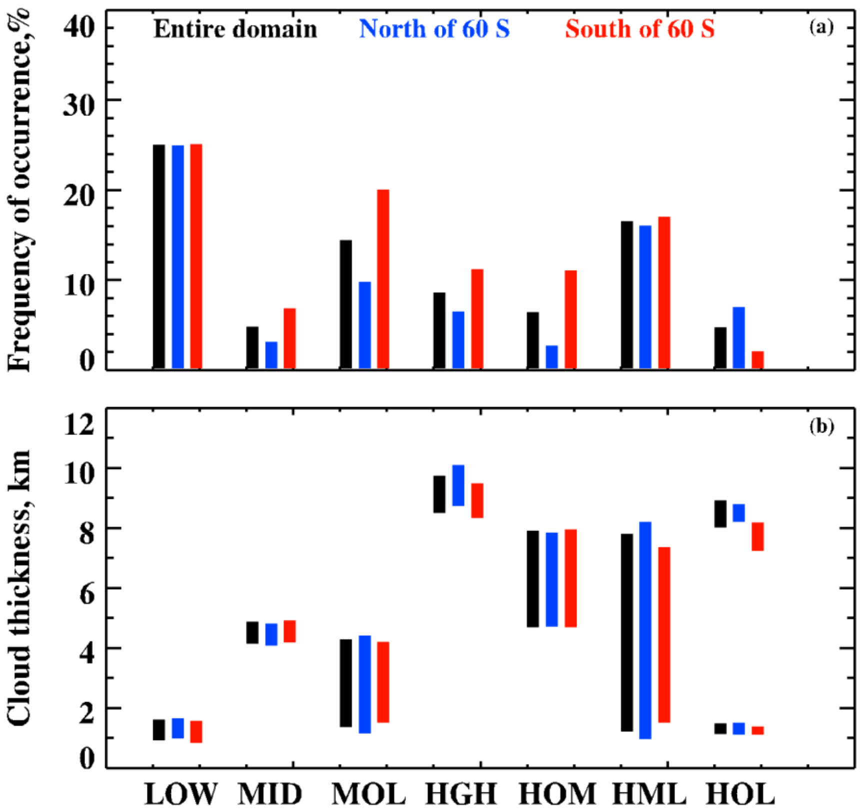

Figure 3. (a) Occurrence frequencies of categorized clouds by their vertical structures. LOW, single - layered low clouds (H_{base} and $H_{\mathrm{top}} \leq 3\ km$); MID, singlelayered middle clouds ($H_{\mathrm{base}} > 3\ km$ and $H_{\mathrm{top}} \leq 6\ km$); MOL, MID over LOW ($H_{\mathrm{base}} < 3\ km$ and $H_{\mathrm{top}} \leq 6\ km$); HGH, singlelayered high clouds ($H_{\mathrm{base}} > 6\ km$); HOM, HGH over MID ($3\ km < H_{\mathrm{base}} < 6\ km$ and $H_{\mathrm{top}} > 6\ km$); HML, HGH over MID and LOW ($H_{\mathrm{base}} < 3\ km$, $H_{\mathrm{top}} \geq 6\ km$ with a MID layer); and HOL, HGH over LOW (LOW and HGH appear at the same time). (b) Cloud thickness for each type of clouds (bar), the top and bottom of the bar represent the maximum cloud-top and minimum cloud-base heights, respectively. Black, blue, and red bars represent the entire domain (Lat:41-69 °S; Long: 60-160° E), north of 60 °S (NSO), and south of 60°S (SSO), respectively, during the MARCUS field campaign (10/2017-3/2018).

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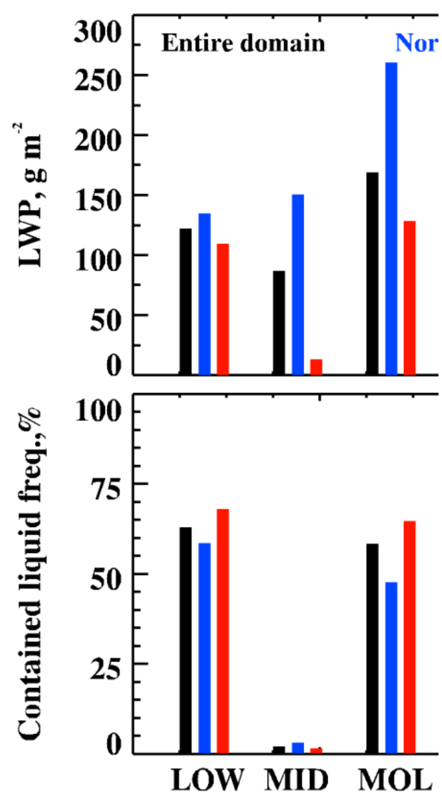

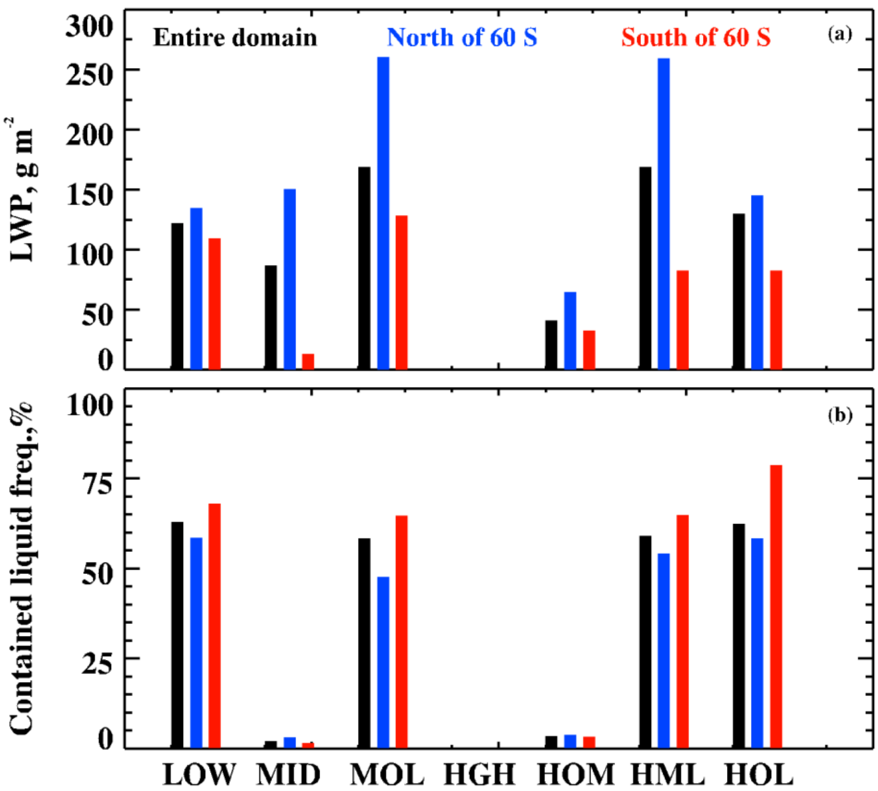

Figure 4. (a) Cloud liquid water paths (*LWPs*) retrieved from microwave radiometer (MWR) measured brightness temperatures using a physical retrieval method for each type of cloud. (b) The occurrence frequencies of LWPs> 10 gm^{-2} for each type of clouds

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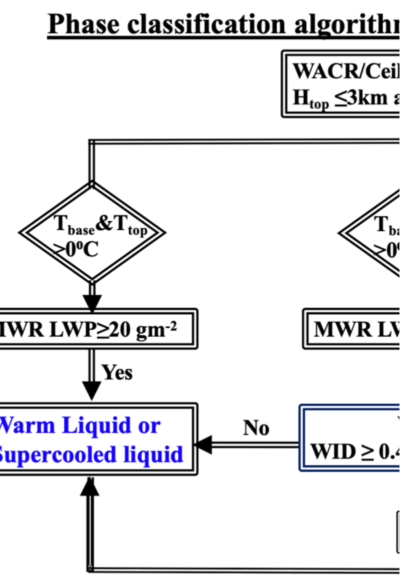

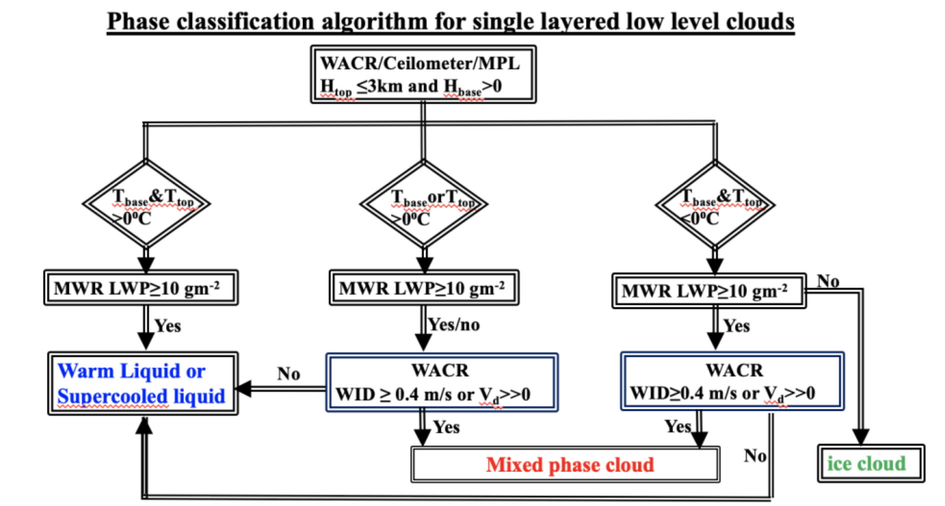

Figure 5. A flow chart for phase classification of single-layered low-level clouds. W-Band (95 GHz) ARM Cloud Radar (WACR) provides radar spectrum width (*WID*) and Doppler velocity (V_d).

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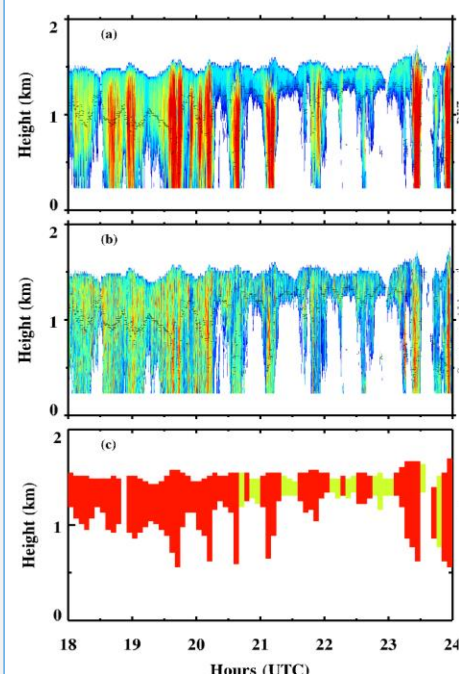

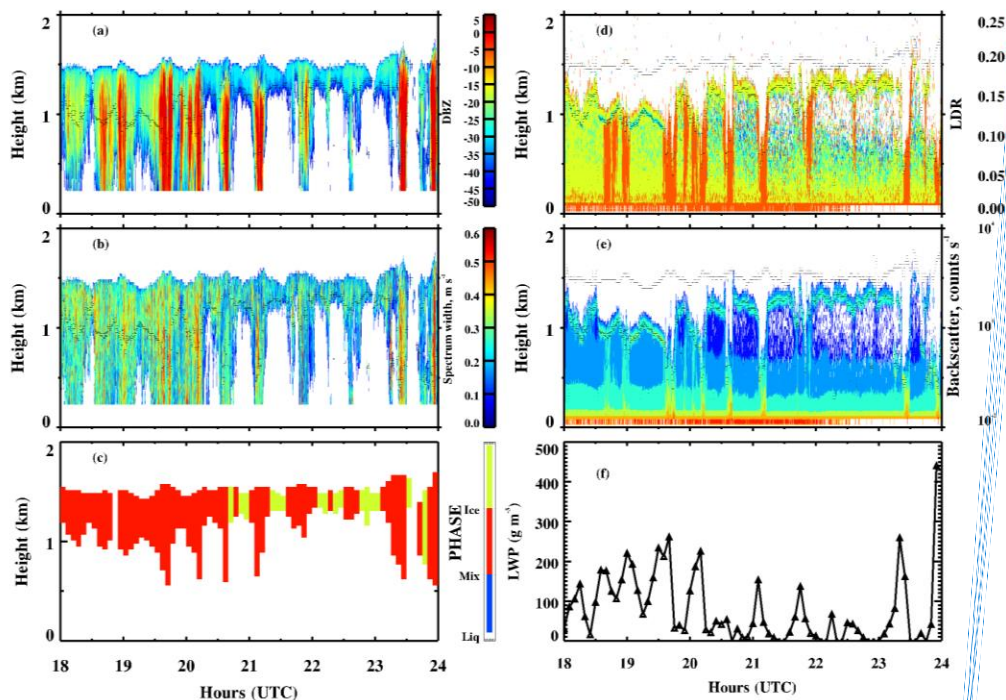

Figure 6. A case study that shows our phase classification (left column) and MicroPulse Lidar (MPL) linear depolarization ratios (*LDR*) and backscatter. W-Band (95 GHz) ARM Cloud Radar (WACR) reflectivity is shown in (a) and spectrum width is shown in (b). Correspondingly, the phase classification in (c); MPL *LDR* in (d) and backscatter in (e); and MWR-derived *LWP* in (f).

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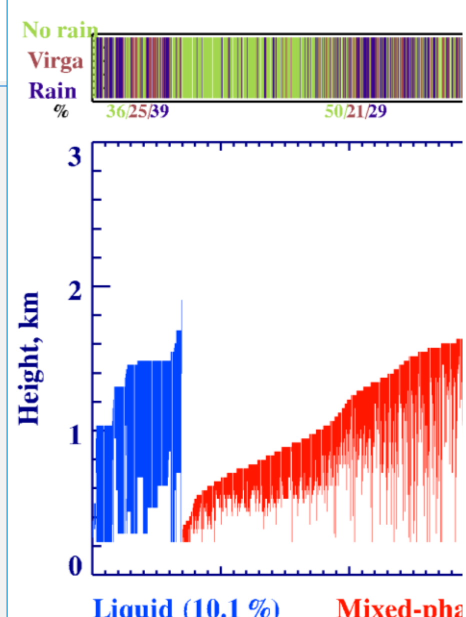

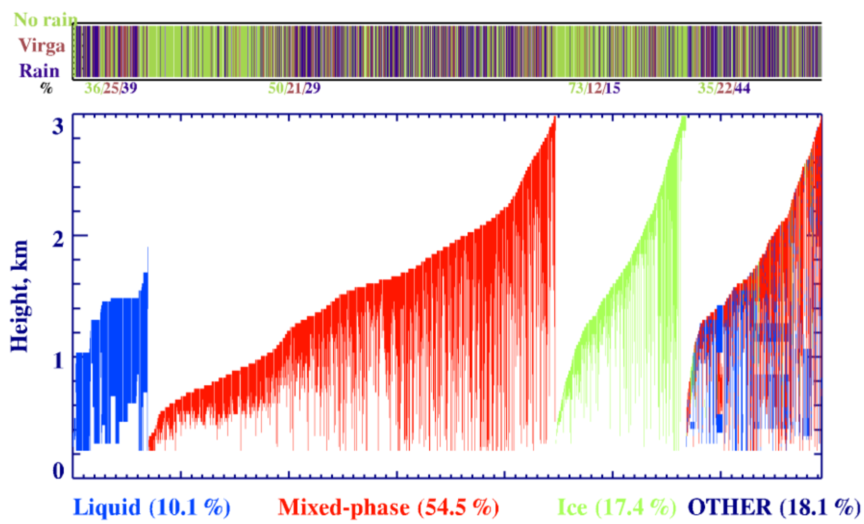

Figure 7. (Upper Panel) The drizzling status for each categorized cloud type, e.g., no rain (yellow-green), virga (brown) and rain (navy blue), the percentages shown below the x-axis represent the portion of drizzling in each type of clouds. (Bottom Panel) The percentages and vertical distributions of classified liquid, mixed-phase, ice, and 'OTHER' clouds for each column in the single-layered low-level clouds, represented by different colors, over the entire domain during MARCUS. Each line represents one 5-min sample. The definition of drizzle here is the radar reflectivity below the ceilometer-derived cloud base, which could be either liquid drizzle drops or ice crystals.

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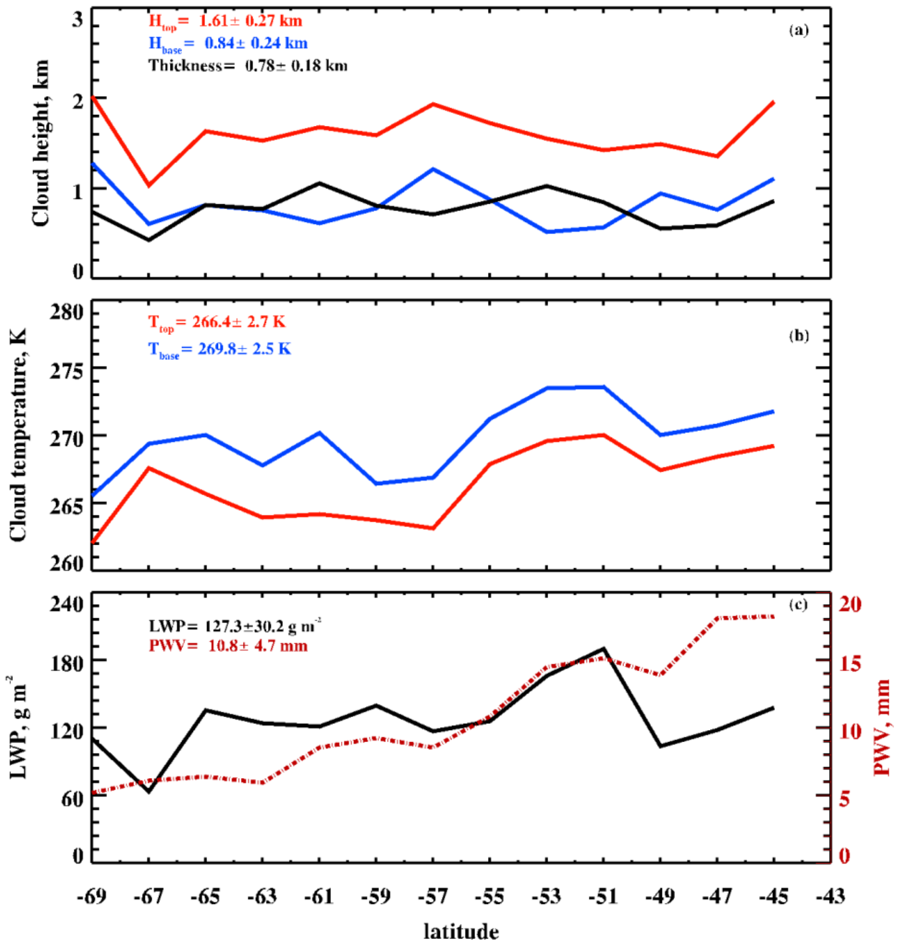

Figure 8. Meridional variations of single-layered low-level cloud properties: (a) cloud-base (H_{base}) and -top (H_{top}) heights, and cloud thickness (ΔH), (b) cloud-base (T_{base}) and -top (T_{top}) temperatures, and (c) cloud liquid water path (LWP) and precipitable water vapor (PWV) over the entire domain during MARCUS.

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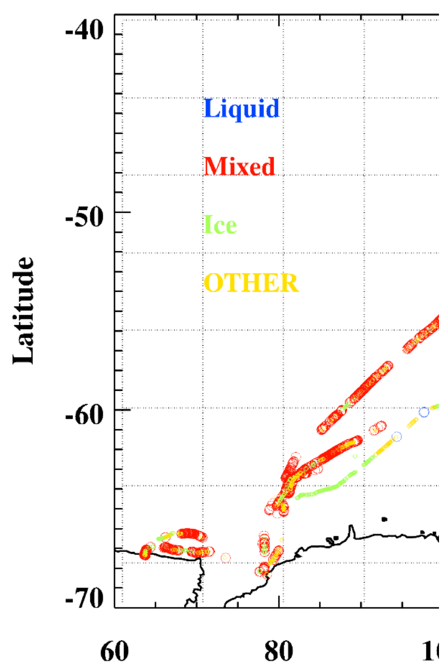

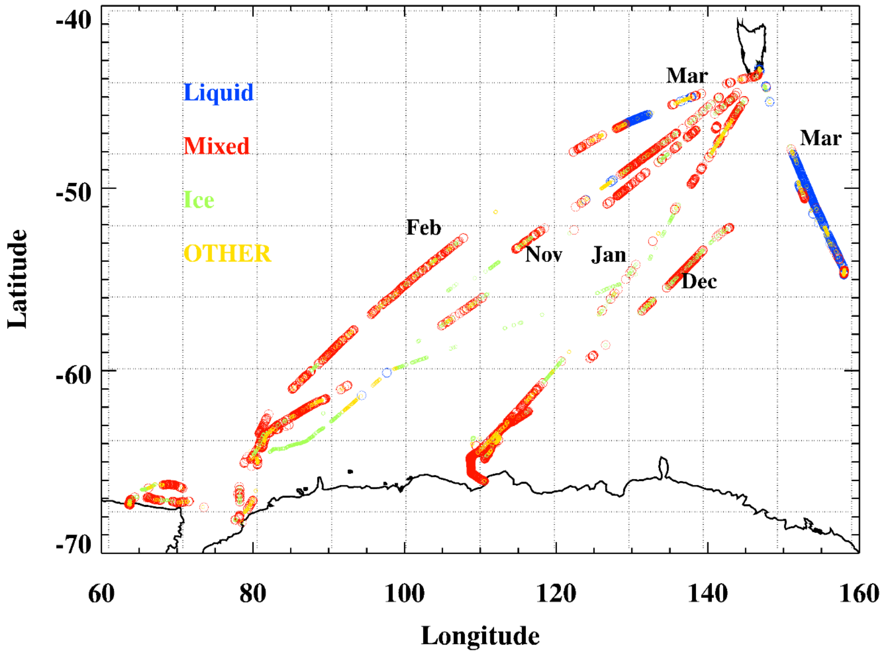

Figure 9. The latitudinal and longitudinal distributions of classified mixed-phase, liquid, and ice clouds in the single-layered low-level clouds. The liquid (blue), mixed (red), ice (light green), and OTHER (yellow) are shown along each shiptrack from October 2017 to March 2018 during MARCUS.

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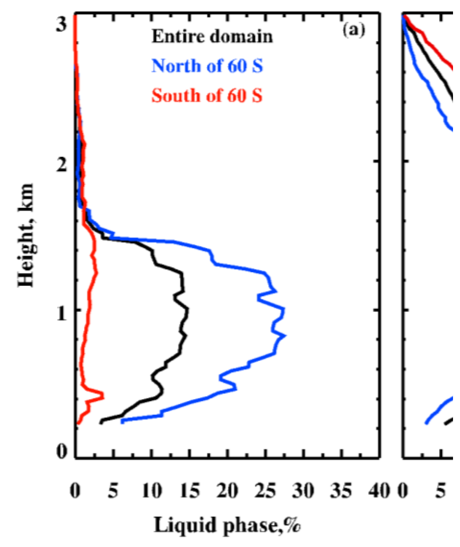

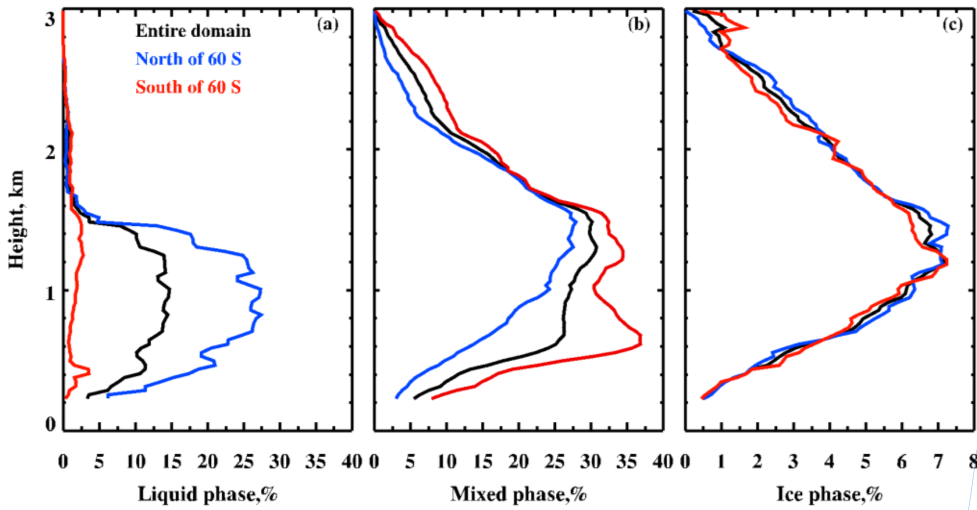

Figure 10. Occurrence frequencies of classified mixed-phase, liquid, and ice clouds over the entire domain (black), North of 60 °S (blue) and South of 60 °S (red) during MARCUS.

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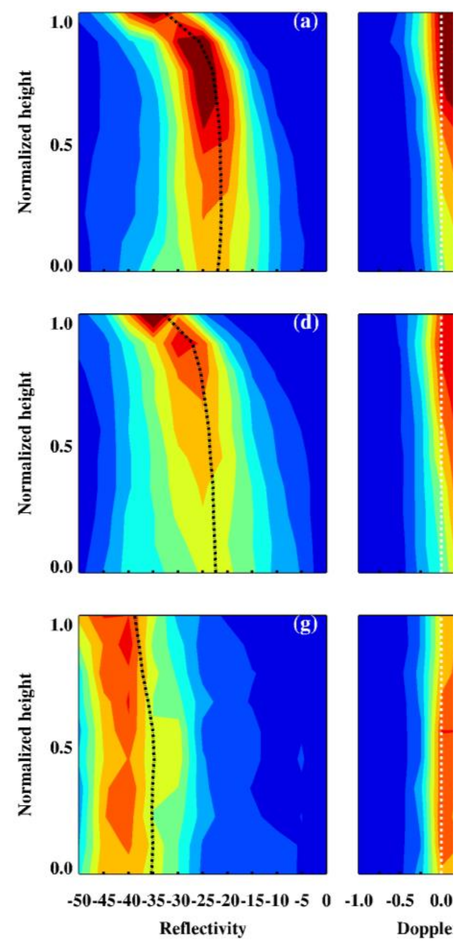

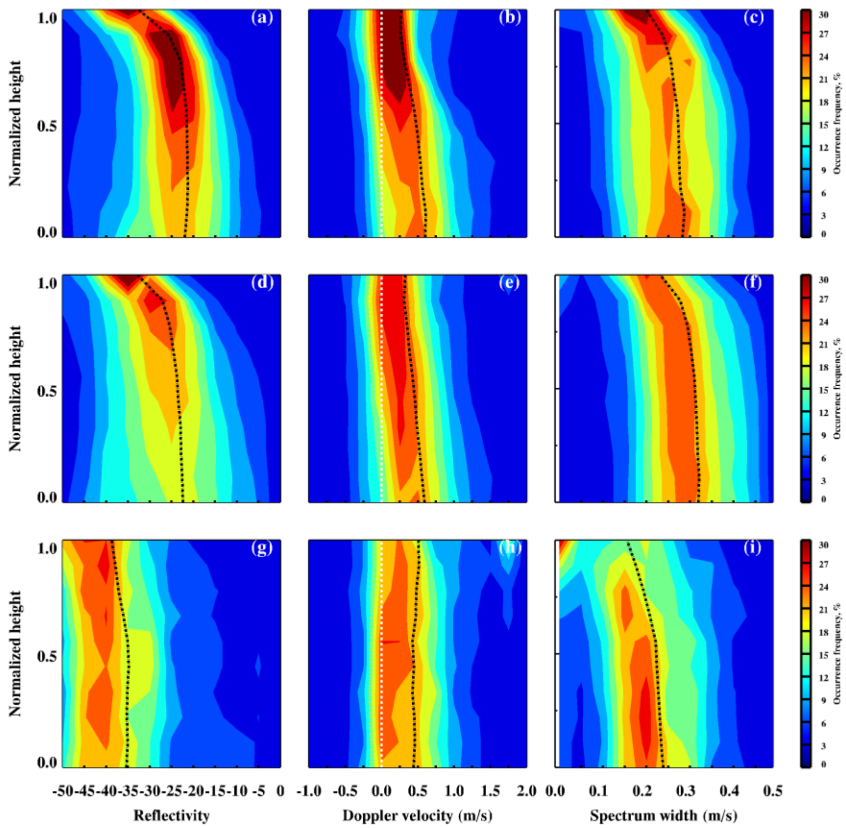

Figure 11. Normalized vertical distributions of radar reflectivity (a), Doppler velocity (b) and spectrum width (c) for the classified liquid (upper panel), mixed-phase (d to f, middle panel) and ice (g to i, bottom panel) clouds over the North of 60 °S during MARCUS Intensive observational period (IOP). Normalized height is defined as $= \frac{H-H_{base}}{H_{top}-H_{base}}$ where cloud base is denoted as 0 and cloud top is 1. The black lines represent the median values and the white lines in Doppler velocity represent the reference of 0.0 m s^{-1}.

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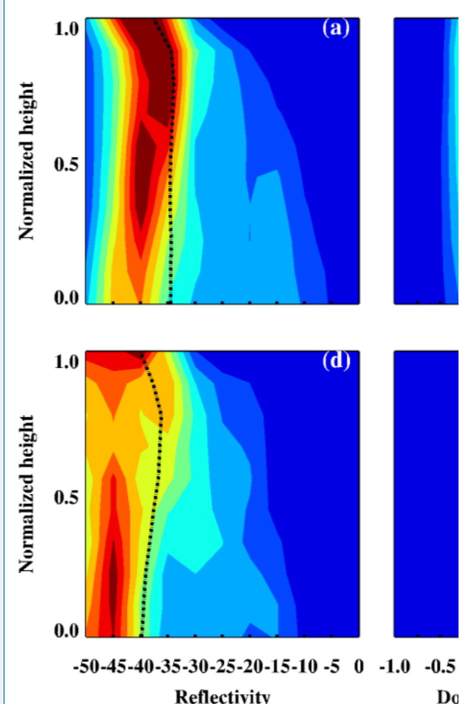

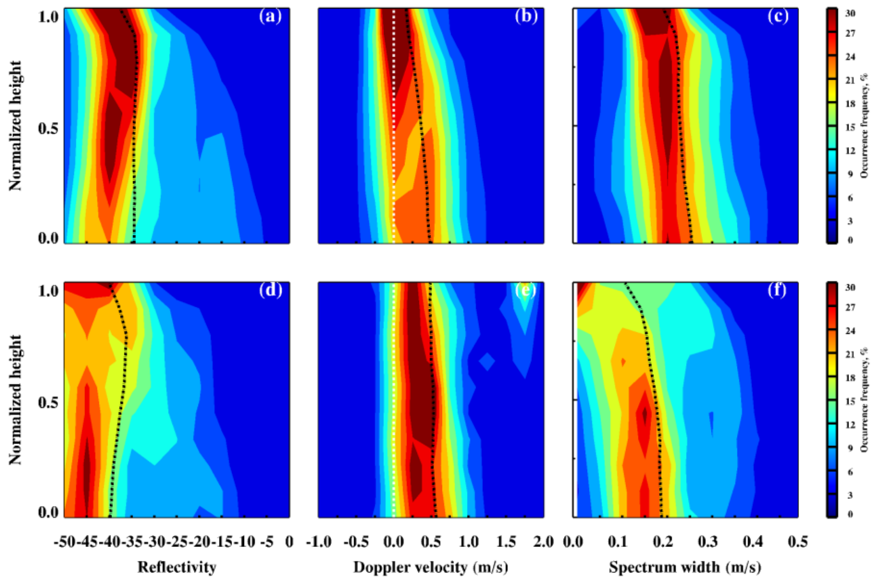

Figure 12. Same as Fig. 11 but only for mixed-phase (a to c, upper panel) and ice (d to f, bottom panel) over the south of 60 °S during MARCUS.

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