

# 1 Retrieval of Solar-induced Chlorophyll Fluorescence from Satellite

## 2 Measurements: Comparison of SIF between TanSat and OCO-2

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17 **Abstract.** Solar-induced chlorophyll fluorescence (SIF) is emitted during photosynthesis in plant leaves. It constitutes a  
18 small additional offset to reflected radiance and can be observed by sensitive instruments that with high [signal-to-noise ratio](#),  
19 and spectral resolution. The Chinese global carbon dioxide monitoring satellite (TanSat), as its mission, acquires greenhouse  
20 gas column density. The advanced technical characteristics of the hyper-spectrum grating spectrometer (ACGS) onboard  
21 TanSat enable SIF retrieval from space observations in the O<sub>2</sub>-A band. In this study, one-year SIF data [for each](#) sounding  
22 [was retrieved from](#) Orbiting Carbon Observatory-2 (OCO-2) and TanSat [measurements](#) using [the JAPCAS/SIF algorithm](#). A  
23 comparison between the SIF results retrieved from OCO-2 spectra and the official OCO-2 SIF product (OCO2\_Level  
24 2\_Lite\_SIF.8r) shows their strong linear relationship ( $R^2 > 0.85$ ) and suggests the reliability of the SIF retrieval algorithm.  
25 The global distribution showed that the SIF retrieved from the two satellites shared the same spatial pattern for all seasons  
26 with the gridded SIF difference less than  $0.3 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ , and they also agreed well with the official OCO-2 SIF product  
27 with the difference less than  $0.2 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ . The retrieval uncertainty of seasonal-gridded TanSat SIF is less than  $0.03$   
28  $\text{W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$  whereas the uncertainty of each sounding ranges from 0.1 to  $0.6 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ . The relationship between  
29 [annually-averaged](#) SIF products and FLUXCOM gross primary productivity (GPP) was also estimated for six vegetation  
30 types in a  $1^\circ \times 1^\circ$  grid over the globe, indicating that the SIF data from the two satellites have the same potential in  
31 quantitatively characterizing ecosystem productivity. The spatiotemporal consistency between TanSat and OCO-2 and their  
32 comparable data quality make the comprehensive usage of the two mission products possible. Data supplemented by TanSat  
33 observations are expected to contribute to the development of global SIF maps with more spatiotemporal detail, which will  
34 advance global research on vegetation photosynthesis.

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## 44 **1 Introduction**

45 Terrestrial vegetation accounts for a large part of the ecosystem, with its photosynthesis and respiration processes playing  
46 important roles in the global carbon cycle. Incoming radiation is absorbed, reflected, and/or transmitted by plant leaves. A  
47 portion of the absorbed radiation is used by the chlorophyll in plant leaves for carbon fixation, while the rest is either  
48 dissipated as heat or re-emitted as solar-induced chlorophyll fluorescence (SIF) at longer wavelengths (Frankenberg et al.,  
49 2011a, 2014). In contrast to the traditional remotely sensed vegetation indices obtained from some studies (Frankenberg et  
50 al., 2011b; Guanter et al., 2014; Li et al., 2018; Sun et al., 2017a; Yang et al., 2015; Zhang et al., 2014), SIF offers the  
51 potential to measure photosynthesis activity and gross primary production (GPP), due to the strong correlation between these  
52 measures (Frankenberg et al., 2011b; Guanter et al., 2012, 2014). The fluorescence emission ( $F_s$ ) adds a low-intensity  
53 radiance less than  $10 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$  and fills in the solar absorption features of the reflected spectrum (Frankenberg et al.,  
54 2011a). The filling-in effect of the solar lines (Fraunhofer lines) is the basic principle applied to measure SIF from space  
55 using the capabilities of hyperspectral observation (Frankenberg et al., 2011b; Guanter et al., 2012).

56 The first attempt at SIF research based on space-based observations was performed using images acquired by the Medium  
57 Resolution Imaging Spectrometer (MERIS) onboard the ENVIRONMENTAL SATellite (ENVISAT) (Guanter et al., 2007). This  
58 led to a new idea for conducting SIF studies on a global scale. The first global SIF map was retrieved with high-resolution  
59 spectra from the Greenhouse-gases Observing SATellite (GOSAT) (Joiner et al., 2011; Frankenberg et al., 2011b). After that,  
60 SIF retrievals were implemented from a variety of satellite measurements, such as those from the Global Ozone Monitoring  
61 Experiment-2 (GOME-2) instruments onboard meteorological operational satellites, SCIAMACHY on board ENVISAT, and  
62 Orbiting Carbon Observatory-2 (OCO-2) (Joiner et al., 2016; Köhler et al., 2015). The TROPOspheric Monitoring  
63 Instrument (TROPOMI) on board Sentinel 5 Precursor (S-5P) provides more efficient SIF observations in terms of global  
64 coverage and new opportunities for exploring the application potential of SIF data in the terrestrial biosphere as well as in  
65 climate research (Doughty et al., 2019; Köhler et al., 2018b). Furthermore, an upcoming European Space Agency mission  
66 called FLUorescence EXplorer (FLEX), the first satellite dedicated to SIF emission observation, will launch in the middle of  
67 2024 (Drusch et al., 2017). Many studies on SIF applications have been initiated with the accumulation of SIF products in  
68 recent years. The responses of satellite-measured SIF to environmental conditions have been applied to drought dynamics  
69 monitoring and regional vegetation water stress estimation (Lee et al., 2013; Sun et al., 2015; Yoshida et al., 2015). As a  
70 proxy of photosynthesis, SIF acts as a powerful constraint parameter in estimating carbon exchange in an ecosystem between  
71 the atmosphere, ocean, and soil; as such, the analysis of the relationship between SIF and GPP has become an important  
72 research topic (Li et al., 2018; Köhler et al., 2018a; Sun et al., 2017a; Zhang et al., 2018). The strong linear relationship  
73 between them paves the way for improving terrestrial ecosystem model simulation of GPP, along with consequent  
74 improvement of global carbon flux estimation (MacBean et al., 2018; Yin et al., 2020). GPP estimations based on satellite-  
75 measured SIF have proven to be an effective method validated by in-situ flux observations (Joiner et al., 2018; Qiu et al.,  
76 2020). However, uncertainty in the factors that determine the relationship between SIF and GPP still exists and is a key

77 limitation in the application of SIF to flux estimation. Based on multi-satellite SIF products, eddy covariance flux tower  
78 observations, and ecological models, the relationship between SIF and GPP under different environmental conditions has  
79 been discussed in a number of studies to analyze the dominant factors for the growing status of different biomes, such as  
80 temperature, soil moisture, and vegetation types (Chen et al., 2020; Doughty et al., 2019; Li et al., 2020; Qiu et al., 2020; Yin  
81 et al., 2020).

82 The Chinese global carbon dioxide monitoring satellite (TanSat) was launched in December 2016. Aiming at acquiring CO<sub>2</sub>  
83 concentrations as OCO-2, TanSat flies in a sun-synchronous orbit at approximately 700 km in height with a 16-day repeat  
84 cycle and an equator crossing time of ~1:30 p.m. local time (Cai et al., 2014; Liu et al., 2018; Yang et al., 2018). Onboard  
85 TanSat, the hyperspectral Atmospheric Carbon-dioxide Grating Spectrometer (ACGS) is designed to separately record solar  
86 backscatter spectra in three channels centered at 0.76 μm (O<sub>2</sub>-A band), 1.61 μm (weak CO<sub>2</sub> absorption band), and 2.06 μm  
87 (strong CO<sub>2</sub> absorption band). With the recorded spectra, many Optimal Estimation Method (OEM) full physics retrieval  
88 algorithms have been developed and applied for XCO<sub>2</sub> retrievals (Boesche et al., 2009; Butz et al., 2009, 2011; O'Dell et al.,  
89 2012; Reuter et al., 2010; Yang et al., 2015b; Yoshida et al., 2011, 2013). The Institute of Atmospheric Physics Carbon  
90 Dioxide Retrieval Algorithm for Satellite Remote Sensing (IAPCAS) algorithm has been applied for TanSat retrieval (Yang  
91 et al., 2018; Yang et al., 2021) and was also previously tested on GOSAT and OCO-2 missions (Yang et al., 2015b).  
92 However, the fluorescence feature causes substantial biases when retrieving surface pressure and scattering parameters from  
93 the O<sub>2</sub>-A band, and the associated errors propagate into the XCO<sub>2</sub> retrievals. In previous XCO<sub>2</sub> retrieval, the surface  
94 emissions were well modeled as a continuum zero offset of the O<sub>2</sub>-A band to reduce errors (Frankenberg et al., 2011a, 2012;  
95 Butz et al., 2009, 2010; Joiner et al., 2012). The high spectral resolution of ~0.044 nm and a signal-to-noise ratio of ~360  
96 in the O<sub>2</sub>-A band makes it possible to obtain SIF from space measurements, with a spatial resolution of 2 km × 2 km in nadir  
97 mode (Liu et al., 2018).

98 Various approaches have been used to infer SIF from satellite measurements (Frankenberg et al., 2011b, 2014a, 2014b;  
99 Guanter et al., 2007, 2012, 2015; Joiner et al., 2011, 2013, 2016; Köhler et al., 2015, 2018b). The SIF signal induces a  
100 filling-in effect of solar lines, which can be used for SIF retrieval, as the fractional depth of solar Fraunhofer lines does not  
101 change during radiation transmission in the atmosphere. To recognize the filling-in features by SIF, high-resolution spectra  
102 and an instrument spectral response function (ISRF) are required to describe subtle changes in the spectral absorption lines.  
103 With the detailed spectral features, a method was developed based on solar line fitting and the Beer-Lambertian law. This  
104 method is robust and accurate when the spectrum is out of the influence of telluric absorptions, even in the presence of  
105 aerosols (Frankenberg et al., 2011a; Joiner et al., 2011); in the current study, this method was applied to develop the  
106 IAPCAS/SIF algorithm. Another SIF retrieval method is the data-driven algorithm based on the singular value  
107 decomposition (SVD) technique (Joiner et al., 2011; Guanter et al., 2012), which has been broadly applied in GOSAT, OCO-  
108 2, TanSat and TROPOMI SIF retrieval (Joiner et al., 2011; Guanter et al., 2012, 2015; Frankenberg et al., 2014a; Du et al.,  
109 2018; Köhler et al., 2018b). In the data-driven method, the spectrum is represented as a linear combination of the SIF signal  
110 and several singular vectors that are trained from non-fluorescent scenes by SVD; thus, the SIF signal can be obtained with

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111 linear least-squares fitting (Du et al., 2018; Guanter et al., 2012). The first TanSat SIF map was obtained by the SVD  
 112 method (Du et al., 2018). In a previous study, a new TanSat SIF product retrieved by IAPCAS/SIF algorithm was introduced  
 113 and the two kinds of TanSat SIF product by IAPCAS/SIF and SVD methods were compared. The preliminary comparison  
 114 between the two TanSat SIF products shows that the two SIF products share a similar global pattern and signal magnitude  
 115 for all seasons while different biases still exist in four seasons (Yao et al., 2021). The different biases in four seasons may be  
 116 caused by the different training samples of the SVD method, which indicates that the training samples have a significant  
 117 impact on the retrieval results. In order to obtain stable SIF data products from TanSat and other subsequent satellite  
 118 missions, it is particularly important to establish a stable and high-precision SIF inversion algorithm. To detailed validate the  
 119 IAPCAS/SIF algorithm and test the potential of comprehensive usage of multi-satellites SIF data in analysis, in this study,  
 120 we detailed the TanSat SIF retrieval using the IAPCAS/SIF algorithm and made a comparison of SIF products between  
 121 TanSat and OCO-2 in variety of temporal-spatial scales, although a preliminary test was shown in previous work.

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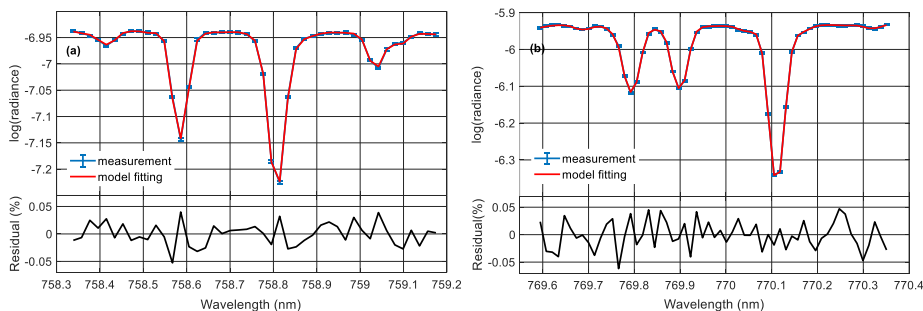
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## 122 2 Data and retrieval algorithm

### 123 2.1 Retrieval Principle and Method

124 We used TanSat version 2 Level 1B (L1B) nadir-mode earth observation data in the retrieval process. The measurements  
 125 covered the period from March 2017 to February 2018. Polarized radiance in the O<sub>2</sub>-A band with a spectral resolution of  
 126 0.044 nm was provided in the L1B data, and two micro-windows near 757 nm (758.3-759.2 nm) and 771 nm (769.6-770.3  
 127 nm) were chosen to retrieve the top-of-atmosphere (TOA) SIF while avoiding the contamination from strong lines of  
 128 atmospheric gas absorption. The retrieval was independent for each micro-window as shown in Figure 1. To avoid  
 129 duplication of information, we use the SIF product at 757 nm as the example in the analysis.



130 **Figure 1: The fitted spectra and residuals for the (a) 757 nm and (b) 771 nm micro-windows of TanSat measurement. The error**  
 131 **bar of the measured spectra depicts the estimated precision of each TanSat sounding.**  
 132

133

142 Filling-in on solar lines by chlorophyll fluorescence in the O<sub>2</sub>-A band can be detected in the hyperspectral measurements  
 143 from TanSat. This effect on spectral radiance is different from the impact of atmospheric and surface processes, e.g.,  
 144 scattering and absorption. For example, scattering by aerosols and clouds does not change the relative depth of clear solar  
 145 lines, unlike the SIF emission signal. We applied the differential optical absorption spectroscopy (DOAS) technique to  
 146 IAPCAS/SIF algorithm for TanSat measurement (Frankenberg, 2014b; Sun et al., 2018).

147 The TOA spectral radiance ( $L_{TOA}^\lambda$ ) at wavelength  $\lambda$  can be represented as follows:

$$148 L_{TOA}^\lambda = I_t^\lambda \cdot \mu_0 \cdot \left( \rho_0^\lambda + \frac{\rho_s^\lambda \cdot T_1^\lambda \cdot T_1^\lambda}{\pi} \right) + F_{TOA}^\lambda \quad (1)$$

149 where  $I_t^\lambda$  is the incident solar irradiance at the TOA,  $\mu_0$  is the cosine of the solar zenith angle (SZA),  $\rho_0^\lambda$  is atmospheric path  
 150 reflectance,  $\rho_s^\lambda$  is surface reflectance, and  $T_1^\lambda$  and  $T_1^\lambda$  are the total atmospheric transmittances along the light-path in the  
 151 downstream and upstream directions, respectively.  $F_{TOA}^\lambda$  is the SIF radiance at TOA.

152 The first term on the right of Eq. (1) represents the transmission process of solar radiance. In the micro-windows used in SIF  
 153 retrieval, gas absorption is very weak and smooth, and hence, the atmosphere term  $\mu_0 \cdot \left( \rho_0^\lambda + \frac{\rho_s^\lambda \cdot T_1^\lambda \cdot T_1^\lambda}{\pi} \right)$  can be simplified to a  
 154 low-order polynomial that varies with  $\lambda$  (Joiner et al., 2013; Sun et al., 2018); this is always valid as long as the spectrum  
 155 fitting range is out of sharp atmospheric absorptions. In the retrieval, the spectral radiance measurement was converted to  
 156 logarithmic space by the instrument and the radiative transfer process  $f(F_s^{rel}, \mathbf{a})$  was represented as follows:

$$157 f(F_s^{rel}, \mathbf{a}) = \log(\langle I_t + F_s^{rel} \rangle) + \sum_{i=0}^n a_i \cdot \lambda^i \quad (2)$$

158 where  $\langle \rangle$  denote the convolution with the ISRF from line-by-line spectra. The polynomial coefficient  $\mathbf{a}$  determines the  
 159 wavelength dependence polynomial for the atmosphere term; in the retrieval, we used a second-order polynomial ( $n = 2$ ).  
 160 The radiance is normalized to the continuum level; hence,  $I_t$  is a normalized disk-integrated solar transmission model, and  
 161  $F_s^{rel}$  is the normalized relative SIF. In the micro-window, SIF was regarded as a constant signal due to its small changes.

162 Although the atmospheric gas absorption was very weak in the micro-window, the weak absorption and the far-wing effects  
 163 (O<sub>2</sub> lines) can still change spectral features, which induces errors in spectrum fitting. In other physical-based retrievals, the  
 164 surface pressure data of the European Centre for Medium-Range Weather Forecasts (ECMWF) is usually used as the true  
 165 surface pressure to simulate the molecular absorption cross-section. However, there is still a difference between the true  
 166 surface pressure and the model surface pressure, so we introduced a factor here to reduce the influent of the inaccurate  
 167 surface pressure. In IAPCAS/SIF algorithm, we used the ECMWF interim surface pressure (0.75° × 0.75°) to estimate O<sub>2</sub>  
 168 absorption firstly and then modified the absorption feature by a scale factor. The scale factor is obtained simultaneously in  
 169 SIF retrieval to reduce the error induced by the uncertainty in surface pressure. As described by Yang (2020), there is also a  
 170 continuum feature in TanSat LIB data that needs to be considered for the high-quality fitting of the O<sub>2</sub>-A band. However, in  
 171 this study, this continuum feature was not corrected, as the impact of such a smooth continuum variation in the micro-  
 172 window is weak and the polynomial continuum model is capable of compensating for most of this effect.

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177 The state vector list in the retrieval includes the relative SIF signal  $F_s^{rel}$ , a wavenumber shift, the scale of  $O_2$  column  
178 absorption for surface pressure correction, and coefficients of the polynomial. The continuum level radiance  $I_{cont}$  within the  
179 fitting window is calculated using the radiance outside the absorption features in the micro-window and is then used for the  
180 actual SIF signal calculation thus:  $F = F_s^{rel} \cdot I_{cont}$ .

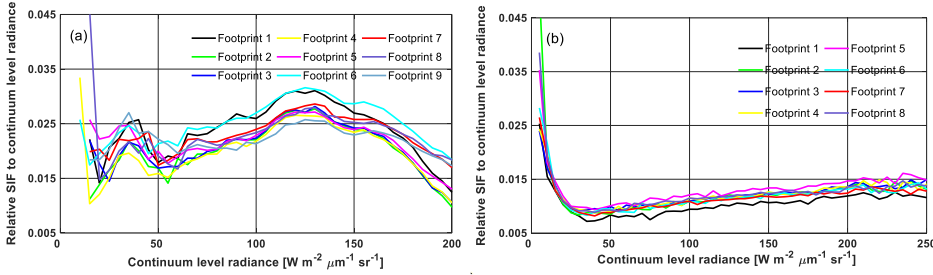
181 In the IAPCAS/SIF algorithm, we used an OEM for state vector optimization in the retrieval process. Unlike  $XCO_2$  retrieval,  
182 SIF retrieval employs a state vector with fewer elements and a much simpler forward model, so there is no need to perform  
183 complex radiation transmission calculations. Considering the low complexity of SIF retrieval, the Gauss-Newton method  
184 was applied in inversion iteratively to find the optimal solution.

## 185 **2.2 Bias Corrections**

186 A systematic error remains in the raw SIF retrieval output if no bias correction is performed; similar results have been  
187 reported in GOSAT and OCO-2 SIF retrieval studies (Frankenberg et al., 2011a, 2011b; Sun et al., 2018). This is because the  
188 SIF signal is weak (e.g., typically ~1-2% of the continuum level radiance), which means that even a small issue in the  
189 measurement, such as a zero-offset caused by radiometric calibration error, could induce significant bias. Unfortunately, the  
190 lack of knowledge on in-flight instrument performance makes it difficult to perform a direct systematic bias correction in the  
191 measured spectrum. In the retrieval, a continuum level radiance bin fit was used to estimate the bias. The bins have a  
192 continuum level radiance interval of  $5 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ . In each bin, the mean bias was estimated using all non-fluorescence  
193 measurements, and a piecewise linear function was built from the mean bias of each continuum level radiance interval.

194 The non-fluorescence soundings that were used in the bias estimation were based on the dataset “sounding\_landCover” in  
195 TanSat L1B data. This dataset depends on the MODIS land cover product and provides a scheme consisting of 17 land cover  
196 classifications defined by the International Geosphere-Biosphere Programme. These retrieved measurements marked as  
197 “snow and ice,” “barren,” and “sparsely vegetated” were chosen to estimate the bias. Calibrations compensated for most of  
198 the instrument degradations, but this alone was not perfect. To reduce the impact from the remaining minor discrepancies,  
199 we built the bias correction function daily to obtain bias for each sounding via interpolation of the continuum level radiance  
200 (Sun et al., 2017b, 2018).

201 The bias curves shown in Figure 2 differ significantly between TanSat and OCO-2. This is mostly due to the differences in  
202 instrument performance and radiometric calibration. In general, the TanSat bias curves exhibited two peaks at radiance levels  
203 of approximately  $40$  and  $125 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ , separately, and most biases were larger than  $0.015$ . For OCO-2, the curves  
204 dropped sharply at low radiance levels, reaching the valley at a radiance level of approximately  $40 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ , and then  
205 increased slowly with the radiance level.



206  
 207 **Figure 2: Variations in the bias correction curves of continuum level radiance from (a) TanSat on July 7, 2017, and (b) Orbiting**  
 208 **Carbon Observatory-2 (OCO-2) on June 16, 2017. The different colors in the legend present different footprints of the satellite**  
 209 **frame.**

### 210 2.3 Data Quality Controls

211 Only data that passed quality control were used in further applications. There were two data quality control processes for the  
 212 SIF products: pre-screening and post-screening. Pre-screening focused mainly on cloud screening; only cloud-free  
 213 measurements were used in SIF retrieval. A surface pressure difference (SPD), defined as:

$$214 \Delta P_0 = |P_{retrieval} - P_{ECMWF}| \quad (3)$$

215 was used to evaluate cloud contamination along with a chi-square test

$$216 \chi^2 = \sum \frac{(y_{sim} - y_{obs})^2}{y_{noise}^2} \quad (4)$$

217 where  $y_{sim}$ ,  $y_{obs}$ , and  $y_{noise}$  represent the model fitting spectrum, observation spectrum, and spectrum noise, respectively.  
 218  $P_{retrieval}$  is the apparent surface pressure obtained from O<sub>2</sub>-A band surface pressure retrieval, assuming a single scattering  
 219 atmosphere.  $P_{ECMWF}$  is the ECMWF interim (0.75° × 0.75°) surface pressure. A “cloud-free” measurement was required to  
 220 simultaneously satisfy an SPD of less than 20 hPa and a  $\chi^2$  value of less than 80. Here, post-screening was applied to filter  
 221 out “bad” retrievals; this screening process involved the following steps: (1) SIF retrievals with reduced  $\chi^2$  ( $\chi_{red}^2$ ) values  
 222 ranging from 0.7 to 1.3 were considered “good” fitting, (2) continuum level radiance outside the range of 15 ~ 200 W m<sup>-2</sup>  
 223 μm<sup>-1</sup> sr<sup>-1</sup> was screened out to avoid scenes too bright or too dark, and (3) soundings with the SZA higher than 60° were also  
 224 filtered out.

### 225 2.4 IAPCAS versus IMAP-DOAS OCO-2 SIF Retrieval

226 Before applied to TanSat retrievals, we tested the IAPCAS/SIF algorithm on the OCO-2 L1B data first  
 227 (OCO2\_L1B\_Science.8r) and then compared the retrieval results with the OCO-2 L2 Lite SIF product (OCO2\_Level  
 228 2\_Lite\_SIF.8r) retrieved by the Iterative Maximum A Posteriori-Differential Optical Absorption Spectroscopy (IMAP-  
 229 DOAS) algorithm (Frankenberg, 2014b). The Lite product provides the SIF value for each sounding on a daily basis and  
 230 hence the SIF comparison could be performed on the sounding scale for each month.

231 Table 1 displays the relationship of OCO-2 SIF values between the IAPCAS/SIF and IMAP-DOAS at 757 nm micro-  
 232 window for each month. Overall, the two SIF products were in good agreement. The linear fitting of the two SIF products  
 233 suggests that they are highly correlated, as indicated by the strong linear relationship with  $R^2$  mostly larger than 0.85 and the  
 234 root mean square error (RMSE) of about  $0.2 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ . Good consistency between the two SIF products implies the  
 235 reliability of the IAPCAS/SIF algorithm; thus, it was further applied to TanSat SIF retrieval. However, there was still a small  
 236 bias in the comparisons, which was due, most likely, to the impact of differences in the bias correction method, retrieval  
 237 algorithm, and fitting window.

238 **Table 1:** Summary of the relationship between the Institute of Atmospheric Physics Carbon Dioxide Retrieval Algorithm for  
 239 Satellite Remote Sensing (IAPCAS) OCO-2 and Iterative Maximum A Posteriori-Differential Optical Absorption Spectroscopy  
 240 (IMAP-DOAS) OCO-2 solar-induced chlorophyll fluorescence (SIF) products at 757nm micro-window.

month	Number of soundings	Slope	Intercept	$R^2$	RMSE/ $\text{W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$
2017/03	1097277	0.85	0.034	0.86	0.18
2017/04	1119464	0.86	0.045	0.87	0.19
2017/05	1054235	0.88	0.041	0.88	0.19
2017/06	1014848	0.91	0.032	0.90	0.19
2017/07	965309	0.92	0.011	0.91	0.19
2017/09	211219	0.88	0.005	0.81	0.23
2017/10	473359	0.88	0.031	0.88	0.17
2017/11	579009	0.87	0.022	0.85	0.19
2017/12	645134	0.87	0.020	0.88	0.16
2018/01	788655	0.87	0.019	0.88	0.17
2018/02*	629995	0.86	0.024	0.87	0.18

241 \* Due to the lack of OCO-2 measurements in August 2017, the comparison is only performed for 11 months.

## 242 3 Results and Discussion

### 243 3.1 Comparison between TanSat and OCO-2 SIF Measurements

244 [The comparison between TanSat and OCO-2 SIF Measurements is a useful and powerful method for further verification of](#)  
 245 [the IAPCAS/SIF algorithm. The reason for adopting OCO-2 data is that OCO-2 and TanSat have similar observation modes,](#)  
 246 [including scanning method, transit time, spatial resolution, spectral resolution, and spectral range. The similarities mean that](#)  
 247 [the SIF product from the two satellite missions can be directly compared. Directly comparing OCO-2 and TanSat SIF](#)  
 248 measurements could provide information on joint data application at the sounding scale for further studies. However, an  
 249 identical sounding overlap barely exists because the two satellites often have different nadir tracks on the ground, which is  
 250 induced by the different temporal and spatial intervals of the two satellite missions. Fortunately, the ground tracks of the two  
 251 satellites were relatively close from April 17 to April 23, 2017. A couple of overlapping orbits were found in the  
 252 measurements obtained from Africa with the orbit number of 1733 from TanSat and 14890a from OCO-2 (Figure 3). In the



253 comparison, the OCO2\_Level 2\_Lite\_SIF.8r product was used to present the SIF emission over the study area. These  
254 overlapping measurements encompassed multiple land cover types, in which the SIF varied within an acceptable time  
255 difference (<5 min).

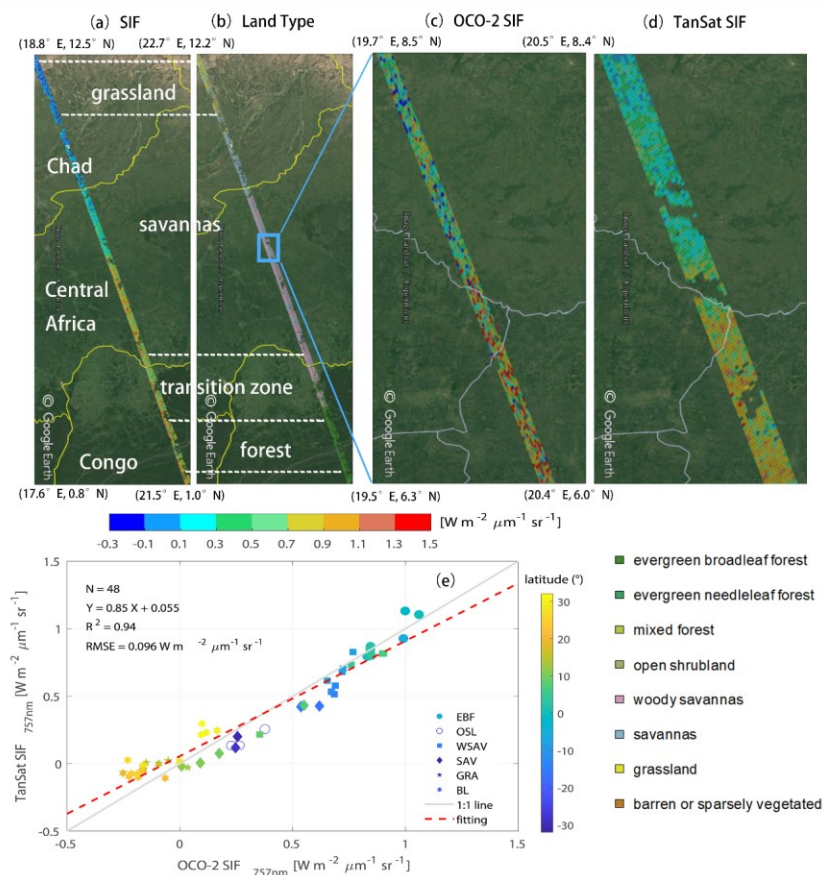
256 Overall, measurements from the two satellites indicated SIF variation with land cover type. The SIF emission over evergreen  
257 broadleaf forests was larger than that over savannas, and grasslands exhibited the lowest SIF emission in April (Figure 3a,b).

258 The mean SIF emission over evergreen broadleaf forests was approximately  $0.9\text{-}1.1 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ , whereas those over  
259 savannas and grasslands were  $0.5\text{-}0.7 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$  and less than  $0.1 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ , respectively (Figure 3c,d).

260 Furthermore, we also found a significant difference in the SIF emission intensity over tropical savannas, which was observed  
261 by both satellites (Figure 3c,d).

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Figure 3: Overlapping orbits of TanSat and OCO-2 on April 19, 2017 over Africa displayed in Google Earth, (a) the SIF measurements from both the two satellites and (b) the footprint land cover type were compared. Compared to OCO-2, TanSat has a wider swath width. A zoom-in view over savannas shows variations in the SIF signal measured by (c) OCO-2 and (d) TanSat. The land surface image shown in Google earth is provided by Landsat/Copernicus team. Following the International Geosphere-Biosphere Programme classification scheme, the vertical legend on the bottom right corner depicts the land cover type that occurs in the study area. The middle horizontal color bar represents the intensity of the SIF radiance. (e) Small-area SIF comparison between OCO-2 and TanSat; each data point represents the mean SIF of a degree in latitude (colors) along the track. The marker legend that is shown on the bottom right of the plot indicates the dominant land cover (defined as the majority land cover type of each sounding) in each small area. There are six land cover types including evergreen broadleaf forest (EBF), open shrubland (OSL), woody savanna (WSAV), savanna (SAV), grassland (GRA), and barren land (BL). The red dashed line represents the linear fit between the two SIF products with statistics shown in the upper left of the plot. The gray line indicates a 1:1 relationship for reference.

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278 Because the footprint sizes of the two satellites are different, it is difficult to make a direct footprint-to-footprint comparison.

279 Therefore, we made the comparison between the two satellite measurements based on a small area average. Each small area

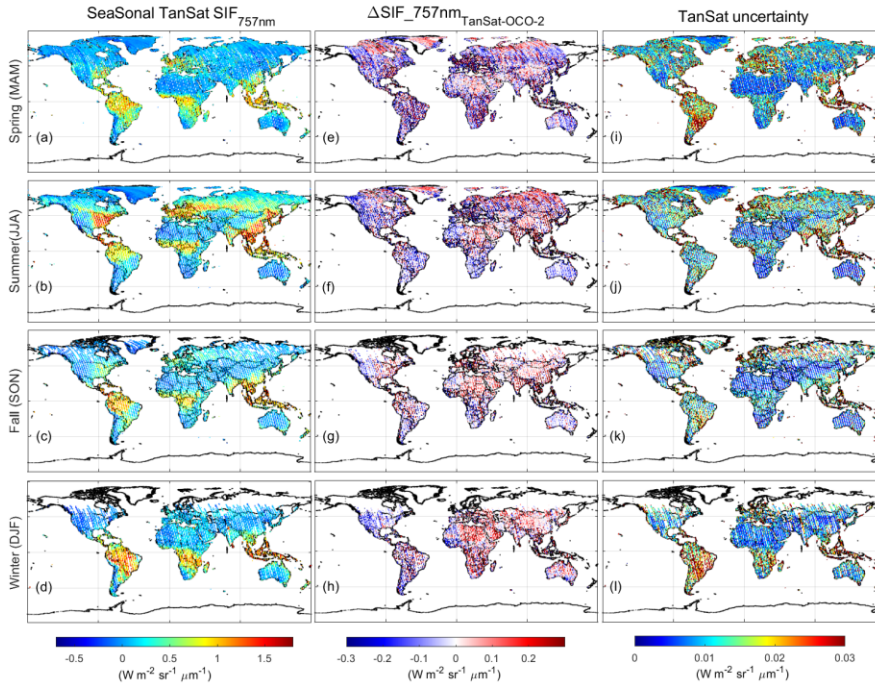
280 spans a degree in latitude and continues along the track. The small area-averaged SIF comparison is shown in Figure 3e. The

281 results indicate good agreement, with an  $R^2$  of 0.94 and an RMSE of  $0.096 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ . Additional ground-based SIF

282 measurement setups (Guanter et al., 2007; Liu et al., 2019; van der Tol et al., 2016; Yang et al., 2015a; Yu et al., 2019)

283 should allow for direct evaluation of satellite retrieval accuracy in the future.

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286 **Figure 4: Global TanSat SIF (left, a-d), differences between TanSat and IAPCAS OCO-2 SIF values (middle, e-h), and the grid-**

287 **cell retrieval uncertainty estimated from TanSat (right, i-l) at  $1^\circ \times 1^\circ$  spatial resolution. The maps in each row represent a**

288 **Northern Hemisphere season, i.e., spring (MAM), summer (JJA), fall (SON), and winter (DJF).**

289

290 Figure 4 shows the global SIF comparison between IAPCAS/SIF retrieved OCO-2 and TanSat; this comparison is only  
 291 performed at  $1^\circ \times 1^\circ$  spatial resolution. In general, the difference in SIF globally is mostly less than  $0.3 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$  for  
 292 all seasons, and on average, the smallest difference appears in fall. There are regional biases observed in North Africa, South  
 293 Africa, South America, and Europe in all seasons except fall. This is mainly caused by the differences in instrument  
 294 performance between TanSat and OCO-2, such as the Instrument Respond Function and the Signal-to-Noise. The instrument  
 295 performance difference is represented by the different structural characteristics of the bias curves. The bias correction  
 296 compensates for most of the bias caused by instrument performance; however, small biases could remain. Furthermore, the  
 297 hundreds of kilometers of distance between the OCO-2 and TanSat footprints, for example, over different vegetation regions,  
 298 will also cause some measurement discrepancies. The global distribution of the two satellites was also compared with the  
 299 official OCO-2 SIF data on the global scale, the results show that the difference between the retrieved SIF maps and the  
 300 official map is less than  $0.2 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ , indicating that the retrieved SIF data from OCO-2 and TanSat both have good  
 301 SIF characterization capabilities on a global scale. The uncertainty  $\sigma$  of each sounding was estimated to validate SIF  
 302 reliability and is provided in the product.  $\sigma$  is derived from the retrieval error covariance matrix,  $S_e = (K^T S_0^{-1} K)^{-1}$ , where  $K$   
 303 is the Jacobian matrix from the forward model fitting and  $S_0$  is the measurement error covariance matrix that is calculated  
 304 from the instrument spectrum noise. In general,  $\sigma$  ranges from 0.1 to  $0.6 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$  for both TanSat and OCO-2  
 305 measurements in the 757 nm fitting window, which is of a similar magnitude and data range as those of previous studies (Du  
 306 et al., 2018; Frankenberg et al., 2014a). Meanwhile, the standard error of the mean SIF in each grid  $\sigma_{meas}$  was estimated to  
 307 represent the gridded retrieval error and natural variability, which is calculated from TanSat SIF values with  $\sigma_{meas} = \frac{\sigma_{std}}{\sqrt{n}}$   
 308 and  $\sigma_{std} = \sqrt{\frac{\sum_{i=1}^n (SIF_i - \overline{SIF})^2}{n}}$ , where  $\sigma_{std}$  represents the standard deviation of the grid cell with  $n$  soundings,  $SIF_i$  is the  
 309 retrieved SIF values of each sounding, and  $\overline{SIF}$  is the mean SIF value for all measurements in the grid. As depicted in the  
 310 right column of Figure 4, the  $\sigma_{meas}$  of each grid cell is much lower than the precision of a single sounding. The  $\sigma_{meas}$  for  
 311 South America is larger than that for any other region on the globe (Figure 4i-l). This is similar to that of OCO-2 SIF  
 312 retrieval and caused by fewer effective measurements due to the South Atlantic Anomaly (Sun et al., 2018). The difference  
 313 in SIF emission values between the two satellites indicates that the collaborative usage of two satellite SIF products still  
 314 requires analysis of the impact of instrument differences, although the two satellite SIF products share the same  
 315 spatiotemporal pattern on a global scale.

### 316 3.2 SIF Global Distribution and Temporal Variation

317 The SIF emission intensity reflects the growth status of vegetation due to its correlation with photosynthetic efficiency;  
 318 hence, the overall global vegetation status can be represented by global SIF maps for each season. TanSat SIF over a whole  
 319 year's cycle, from March 2017 to February 2018, is represented seasonally as a  $1^\circ \times 1^\circ$  grid spatially. The seasonal variation

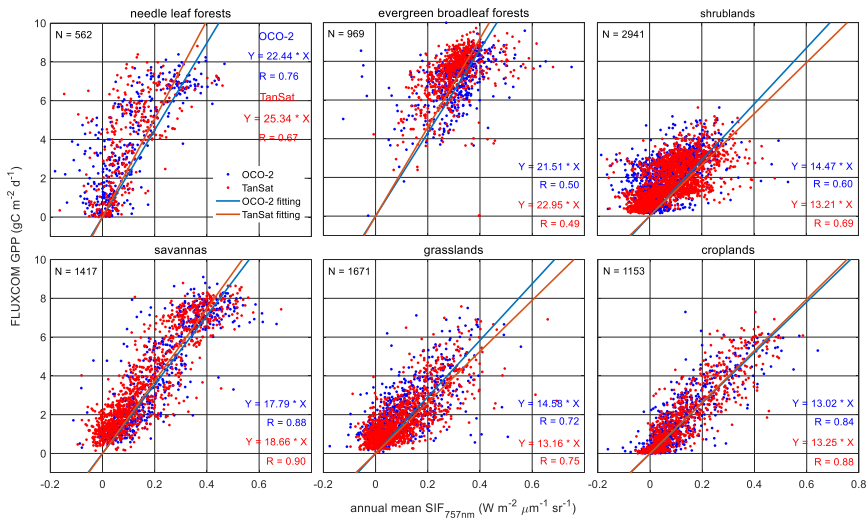
320 in SIF emission is clear in the Northern Hemisphere, i.e., it is enhanced from spring to summer and then decreases (Sun et al.,  
321 2018).

322 In general, the SIF emission varied with latitude and the vegetation-covered areas near the equator maintained a continuous  
323 SIF emission throughout the year. Large SIF emissions in the Northern Hemisphere, above  $1.5 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$ , mostly from  
324 the eastern U.S., southeast of China, and southern Asia in summer, were due to the large areas of cropland. There was also  
325 an obvious SIF emission of  $1\text{-}1.2 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$  observed over Central Europe and northeastern China during the summer.  
326 In these regions, croplands and deciduous forests contribute to SIF emissions. In the Southern Hemisphere, the strongest SIF  
327 emission occurred in the Amazon, with a level of approximately  $1\text{-}2 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$  in DJF (Northern Hemisphere winter),  
328 where there is an evergreen broadleaf rainforest. Africa, which is covered by evergreen broadleaf rainforests and woody  
329 savannas, had an average SIF value of  $0.7\text{-}1.5 \text{ W m}^{-2} \mu\text{m}^{-1} \text{ sr}^{-1}$  during the year.

330 The SIF-GPP relationship over different vegetation types was also investigated by comparing the annual mean satellite SIF  
331 measurements with the FLUXCOM GPP (Jung et al., 2020; Tramontana et al., 2016) dataset in a  $1^\circ \times 1^\circ$  grid over the globe.  
332 The FLUXCOM GPP dataset used in the study comprises monthly global gridded flux products with remote sensing and  
333 meteorological/climate forcing (RS+METEO) setups, which are derived from mean seasonal cycles according to MODIS  
334 data and daily meteorological information (Jung et al., 2020; Tramontana et al., 2016). [In the correlation analysis, the high  
335 spatial resolution \( \$0.5^\circ \times 0.5^\circ\$ \) of the FLUXCOM GPP was first resampled to  \$1^\circ \times 1^\circ\$  to keep the same temporal-spatial scale  
336 of SIF and GPP data.](#) The satellite-measured SIF is an instantaneous emission signal that varies with incident solar radiance  
337 within the day. To reduce the differences caused by the observation time and SZA at different latitudes, we applied a daily  
338 adjustment factor to convert the instantaneous SIF emission into a daily mean SIF (Du et al., 2018; Frankenberg et al., 2011b;  
339 Sun et al., 2018). The daily adjustment factor  $d$  is calculated as follows:

$$340 \quad d = \frac{\int_{t=t_0-12h}^{t=t_0+12h} \cos(SZA(t)) \cdot dt}{\cos(SZA(t_0))} \quad (5)$$

341 where  $t_0$  is the observation time in fractional days and  $SZA(t)$  is a function of latitude, longitude, and time for calculating  
342 the SZA of the measurements. The annual averaged SIF is calculated from the daily mean SIF. To evaluate the relationship  
343 between SIF and GPP on the periodic scale of vegetation growth status, annually-averaged data were used in the regression  
344 fitting analysis.



345  
 346 **Figure 5: Relationship between annual mean SIF and FLUXCOM gross primary production (GPP) from March 2017 to February**  
 347 **2018. Blue and red dots represent OCO-2 and TanSat SIF grids, respectively. Fitted lines and statistics for OCO-2 and TanSat**  
 348 **are shown in each plot.**

349 Figure 5 shows the linear fits for six vegetation types, including needle leaf forest, evergreen broadleaf forest, shrubland,  
 350 savanna, grassland, and cropland. Recent studies have shown a strong linear correlation between SIF and GPP. The TanSat  
 351 SIF and the OCO-2 official SIF data were used to estimate the SIF-GPP correlation. To make a direct comparison of the  
 352 relationship between SIF and GPP among various vegetation types, we used non-offset linear fitting to indicate the  
 353 correlation between satellite SIF and FLUXCOM GPP. For savanna and cropland, there were strong relationships between  
 354 the mean SIF and GPP with an R-value above 0.84. The fitting results show that the SIF products of the two satellites have  
 355 similar capabilities in characterizing GPP, especially for the evergreen broadleaf forest, savanna, and cropland, with slopes  
 356 of approximately 21, 18, and 13, respectively. For shrubland and grassland, the slope of OCO-2 SIF with GPP is higher than  
 357 that of TanSat and has a worse correlation. For forests, OCO-2 SIF presents a better correlation with GPP, especially in the  
 358 needle leaf forest. As a whole, for the same vegetation type, the SIF-GPP correlations for the two satellites are rather similar,  
 359 indicating that the two satellite SIF products have similar capabilities in characterizing GPP. It shows the strong feasibility of  
 360 the comprehensive application of different satellite SIF products. For different vegetation types, the SIF-GPP correlations  
 361 were significantly different, indicating the different ability of SIF to characterize GPP of different vegetation. It represents  
 362 that vegetation type is a key factor in determining the SIF-GPP relationship. The markedly different fitting slopes across  
 363 various biomes suggest that the application of SIF in GPP estimation needs more detailed analysis although the evidence of  
 364 the strong linear relationship between them.

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366 **4 Conclusions**

367 In this paper, we introduced the retrieval algorithm IAPCAS/SIF and its application in TanSat and OCO-2 measurements.  
368 One-year (March 2017-February 2018) TanSat SIF data was introduced and compared with OCO-2 measurements in this  
369 study. The TanSat and OCO-2 SIF products based on the IAPCAS/SIF algorithm are available on the Cooperation on the  
370 Analysis of carbon SATellites data (CASA) website, [www.chinageoss.org/tansat](http://www.chinageoss.org/tansat). Comparisons between TanSat and OCO-2  
371 measurements directly, using a case study, and indirectly, with global  $1^{\circ}\times 1^{\circ}$  grid data, showed consistency between the two  
372 satellite missions, indicating that the coordinated usage of the two data products is possible in future studies. [The correlation](#)  
373 [analysis between SIF and GPP further verified the feasibility of the comprehensive application of SIF products from](#)  
374 [different satellite missions. Meanwhile, it should be noticed that the difference in the ability of satellite SIF products to](#)  
375 [characterize different vegetation types in data applications.](#) With increasing satellites becoming available for SIF  
376 observations, space-based SIF observations have recently expanded in range to provide broad spatiotemporal coverage. The  
377 next-generation Chinese carbon monitoring satellite (TanSat-2) is now in the preliminary design phase, which is designed to  
378 be a constellation of six satellites to measure different kinds of greenhouse gases and trace gases in a more efficient way,  
379 including CO<sub>2</sub>, CH<sub>4</sub>, CO, NO<sub>x</sub>, as well as SIF. SIF measurements from TanSat-2 will provide global data products over  
380 broader coverage areas with less noise. The improvement in the spatiotemporal resolution of SIF data will benefit GPP  
381 predictions based on the numerous studies of the linear relationship between SIF and GPP. In future work, the measurement  
382 accuracy should be validated directly using ground-based measurements to ensure data quality.

383 **Data availability**

384 The SIF products of TanSat and OCO-2 by IAPCAS/SIF algorithm are available on the Cooperation on the Analysis of  
385 carbon SATellites data (CASA) website ([www.chinageoss.org/tansat](http://www.chinageoss.org/tansat)).

386

387 **Author contributions**

388 L.Y. and D.Y. developed the retrieval algorithm, designed the study, and wrote the paper. Y.L. led the SIF data process and  
389 analysis. Y.L., D.Y., Z.C., and J.W. contributed to manuscript organization and revision. C.L. and Y.Z. provided information  
390 on the TanSat instrument performance. L.T. provided TanSat in-flight information. M.W. and S.W. provided information on  
391 the scientific requirement for data further application. N.L. and D.L. led the TanSat data application. Z.Y. led the TanSat in-  
392 flight operation.

393 **Competing interests**

394 The authors declare that they have no conflict of interest.

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