

Introducing the MISR Level 2 Near Real-Time Aerosol Product

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Abstract

Atmospheric aerosols are an important element of Earth's climate system, and have significant impacts on the environment and on human health. Global aerosol modeling has been increasingly used for operational forecasting and as support to decision making. For example, aerosol analyses and forecasts are routinely used to provide air quality information and alerts in both civilian and military applications. The growing demand for operational aerosol forecasting calls for additional observational data that can be assimilated into models to improve model accuracy and predictive skill. These factors have motivated the development, testing, and release of a new near real-time (NRT) level 2 (L2) aerosol product from the Multi-angle Imaging SpectroRadiometer (MISR) instrument on NASA's Terra platform. The NRT product capitalizes on the unique attributes of the MISR aerosol retrieval approach and product contents, such as reliable aerosol optical depth as well as aerosol microphysical information. Several modifications are described that allow for rapid product generation within a three-hour window following acquisition of the satellite observations. Implications for the product quality and consistency are discussed as compared to the current operational L2 MISR aerosol product. Several ways of implementing additional use-specific retrieval screenings are also highlighted.

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1. Introduction

Atmospheric aerosols have for long been recognized to influence the climate, environment, and human health (e.g., IPCC, 2013; Lelieveld et al., 2015; Shindell et al., 2013; Turnock et al., 2020). They also affect satellite remote sensing of important geophysical parameters such as ocean color (e.g., Frouin et al., 2019; Gordon, 1997) or greenhouse gas abundance (Butz et al., 2009; Frankenberg et al., 2012; Houweling et al., 2005). Aerosol particles and their properties have been extensively studied in-situ and remotely: from the ground, in the air, and from space. These observational data vary in spatial and temporal coverage, but usually only offer snapshots of local conditions. Since atmospheric aerosols have a life cycle ranging from hours to days, numerical modeling of their emission, transport, and deposition has filled the coverage gaps and extended our understanding of their global impacts. This has given rise to a number of global aerosol reanalyses (Buchard et al., 2017; Gelaro et al., 2017; Inness et al., 2013, 2019; Lynch et al., 2016; Randles et al., 2017; Rienecker et al., 2011) that provide a long-range, gridded, and internally consistent outlook on aerosol burdens around the world. Furthermore, global aerosol modeling has been increasingly used for operational forecasting (e.g., Xian et al., 2019) and as support to decision making, for example in air quality alerts and in non-civilian applications (Liu et al., 2007).

The growing demand for consistent gridded aerosol products has been driving development and steady improvement of numerical predictions. For example, the International Cooperation for Aerosol Prediction initiative was founded in 2010 (Benedetti et al., 2011; Reid et al., 2011), with one of its goals being the development of global multi-model aerosol forecasting ensemble for basic research and operational use (Xian et al., 2019). Still, models suffer from often poorly resolved aerosol emissions and sinks and can be affected by errors in the underlying meteorology. As a result, systematic and sampling-related biases in aerosol fields are often found between model simulations and satellite observations (e.g., Buchard et al., 2015; Colarco et al., 2010; Lamarque et al., 2013; Zhang and Reid, 2009). An effective way to mitigate some of these problems is by assimilating aerosol observations into numerical models (e.g., Bocquet et al., 2015; Fu et al., 2017; Sekiyama et al., 2010; Di Tomaso et al., 2017; Werner et al., 2019; Zhang et al., 2008). Satellite observations of aerosol optical and microphysical properties are inseparable from these data assimilation activities as they offer the necessary data volume, near-global coverage, and frequent repeat cycle. However, an often-considerable latency for generating science-quality “standard” satellite products (8 to 40 hours) renders them unsuitable for operational forecasting. This has led to the development of aerosol

products within the time frame required by modeling centers, usually three hours from satellite overpass. A number of near real-time (NRT) products has emerged.

One example of a platform that provides users with NRT satellite products and imagery is NASA's Land, Atmosphere Near real-time Capability for EOS (LANCE) project (<https://earthdata.nasa.gov/earth-observation-data/near-real-time>). A range of instruments deliver various Level 1 (L1) and Level 2 (L2) data products (<https://earthdata.nasa.gov/collaborate/open-data-services-and-software/data-information-policy/data-levels>), including radiances, land surface properties, and atmospheric thermodynamics and composition within three hours from satellite observation. NRT aerosol products are currently available from the Moderate Resolution Imaging Spectroradiometer (MODIS), Ozone Monitoring Instrument (OMI), and Visible Infrared Imaging Radiometer Suite (VIIRS). NASA's Multi-angle Imaging SpectroRadiometer (MISR) currently provides NRT radiance and cloud motion vector products. The purpose of this paper is to introduce a new MISR NRT L2 aerosol product available within LANCE.

This paper is organized as follows. Section 2 and 3 provide brief descriptions of the MISR instrument and the data processing sequence, respectively. Section 4 first outlines the cloud identification methods employed in the MISR aerosol algorithm and then describes algorithmic modifications introduced in the NRT processing. Adjustments to cloud and retrieval screening parameters and their implications are discussed. The global distributions of the NRT product are analyzed in Section 5. Section 6 provides a summary.

2. MISR instrument and aerosol data product

The MISR instrument flies aboard the NASA Earth Observing System (EOS) Terra satellite, launched in December 1999 to a sun-synchronous descending polar orbit, at an orbital altitude of 705 km, an orbital period of 99 minutes, and an equatorial crossing time of 10:30 a.m. local time. MISR makes 14.56 orbits per day with a repetition cycle (revisit) of 16 days. The orbit tracks are georeferenced to a fixed set of 233 ground paths. With a cross-track swath of about 380 km, total Earth coverage is obtained every 9 days at the equator and every 2 days at high latitudes.

MISR contains nine pushbroom cameras with viewing angles at the Earth's surface ranging from 0° (nadir) to +/- 70.5° oriented along the direction of the flight track. A point on the ground is imaged by all nine cameras in approximately 7 minutes. The cameras make observations of reflected solar radiance in four spectral bands, centered at 446 (blue), 558

(green), 672 (red), and 866 (near-infrared) nm. The spatial resolution depends on the camera and wavelength. The red band has a full 275 m resolution in all cameras. The other three spectral channels are averaged onboard to a 1.1 km resolution in global-mode operation (Diner et al., 1998), with the exception of the nadir camera which preserves the full 275 m resolution in all spectral channels. See <https://misr.jpl.nasa.gov/Mission/> for more details.

MISR employs two processing pathways for aerosol retrievals, one for observations over land (Martonchik et al., 2009), and another for dark water (DW) (Kalashnikova et al., 2013), which applies over deep oceans, seas, and lakes. Previous versions of the MISR aerosol product were extensively validated over the years (e.g., Kahn et al., 2010; Kahn and Gaitley, 2015; Kalashnikova et al., 2013; Shi et al., 2014; Witek et al., 2013) showing high retrieval quality over land and ocean.

The current operational version of the MISR aerosol product, designated as version 23 (V23), was released publicly in June 2018. It introduced multiple algorithmic, data product, and data usability improvements (Garay et al., 2020; Witek et al., 2018a, 2018b). V23 provides aerosol information with a spatial resolution of 4.4 km x 4.4 km packaged in NetCDF-4 format. Initial validation efforts showed that V23 retrievals are more accurate than previous versions, with most pronounced improvements in the DW algorithm (Garay et al., 2020). V23 retrievals over oceans were extensively validated by Witek et al. (2019), indicating excellent agreement with ground-based observations. Other V23 Aerosol Optical Depth (AOD) evaluation efforts show similar results (e.g., Choi et al., 2019; Sayer et al., 2020; Si et al., 2020; Sogacheva et al., 2020). A first regional insight into retrieved particle properties from the MISR V23 aerosol product shows that MISR generally captures the distinct spatial and temporal features of aerosol type in East Asia (Tao et al., 2020). Furthermore, V23 has greatly improved the quality of reported AOD uncertainties, which now realistically represent retrieval errors (Sayer et al., 2020; Witek et al., 2019). This is especially relevant as pixel-level retrieval uncertainties are very important for satellite data assimilation, which is being increasingly used in aerosol modeling studies (Lynch et al., 2016; Shi et al., 2011, 2013; Zhang and Reid, 2010). MISR data and related documentation can be obtained from: <https://asdc.larc.nasa.gov/project/MISR>.

3. NRT latency and data description

MISR currently provides several L1 and L2 near real-time (NRT) radiance and cloud motion vector products (<https://earthdata.nasa.gov/earth-observation-data/near-real-time/download-nrt-data/misr-nrt>). All MISR NRT processing is based on Level 0 data downlinked in observational

131 sessions. These session-based files, representing portions of a single MISR orbit, usually cover
132 between 10 to 50 minutes of observations, as compared to the full orbit period of 98.9 minutes.
133 This session-based processing is necessary to allow for the fast product delivery required for
134 NRT applications.

135 The new NRT L2 aerosol product file content, described in Data Product Specification
136 (https://asdc.larc.nasa.gov/documents/misr/DPS_AEROSOL_NRT_V023.20210430.pdf), is
137 equivalent to the standard aerosol product (Garay et al., 2020). The NRT L2 aerosol product file
138 name convention is:

139 MISR_AM1_AS_AEROSOL_T{yyyymmddHHMMSS}_P{ppp}_O{oooooo}_F13_0023.nc, where
140 'yyyy', 'mm', and 'dd' are the year, month, and day, and 'HH', 'MM' and 'SS' are the hour,
141 minute, and seconds, respectively. Furthermore, {ppp} is the three-digit path identifier (between
142 001 and 233) and {oooooo} is the six-digit orbit number. The NRT L2 aerosol product files are
143 available for download within three hours of acquisition at NASA's Atmospheric Science Data
144 Center (ASDC) (<https://asdc.larc.nasa.gov/project/MISR>).

145 For clarity, it is important to distinguish between the three different MISR L2 aerosol
146 products: NRT, FIRSTLOOK, and standard aerosol (SA) product (see Figure 1). NRT is
147 generated within a three-hour time interval after acquisition and uses the same ancillary inputs
148 as FIRSTLOOK. These include the monthly gridded (1.0 degree) snow/ice mask and surface
149 wind speed from the Terrestrial Atmospheric and Surface Climatology (TASC) database and the
150 seasonal Radiometric Camera-by-camera Threshold Dataset (RCTD) (Diner et al., 1999a). Both
151 NRT and FIRSTLOOK utilize TASC and RCTD datasets from the current month/season in the
152 prior year. The FIRSTLOOK product is generated within two days from acquisition and includes
153 cloud classification parameters obtained from the L1 and L2 cloud products. The SA product is
154 available after final processing is performed on a seasonal basis and within three months past
155 the end of the season, which results in a 3–6-month latency. The final processing utilizes the
156 most recent snow/ice and wind speed data.

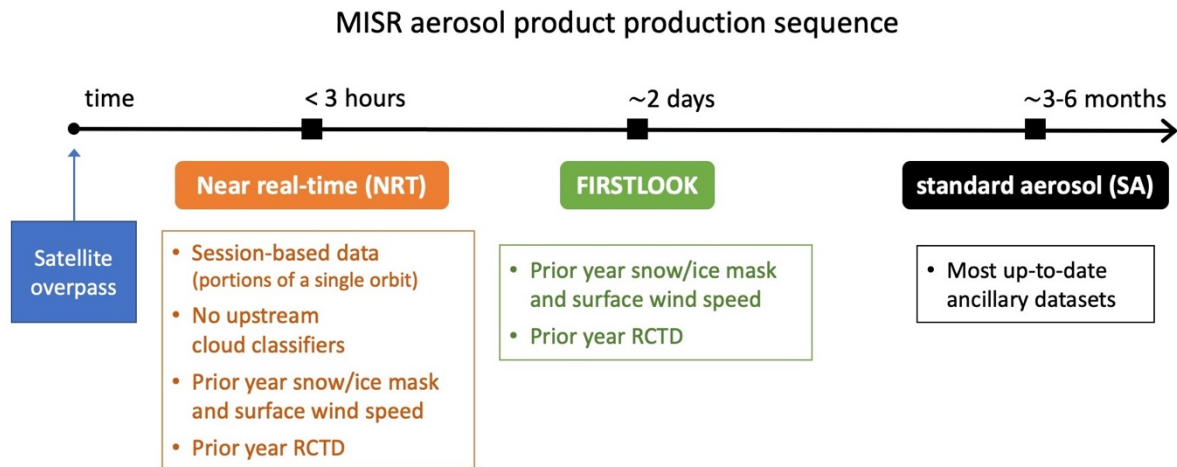


Figure 1 Schematic showing MISR aerosol product delivery timeline. Snow/ice mask and surface wind speed data are monthly averages. RCTD stands for Radiometric Camera-by-camera Threshold Dataset. MISR final production (SA) is processed on a seasonal cycle and is often delayed one to three months past the end of each season, which results in up to 6-month latency.

4. Cloud screening in the NRT MISR aerosol product

4.1. Cloud identification

Identification of cloudy pixels is a critical element of all satellite aerosol remote sensing algorithms. MISR employs several cloud identification strategies which can be loosely split into two groups: the first group relies on cloud classifiers previously generated with MISR Level 2 Cloud Detection and Classification algorithm (Diner et al., 1999b), and the second group includes build-in tests that are internal to the aerosol retrieval algorithm (Diner et al., 2008).

4.1.1. Upstream cloud classifiers

The operational MISR aerosol algorithm relies on a range of external input datasets that are either static—for example, a monthly wind speed climatology—or that need to be generated prior to aerosol retrievals in upstream processing. A notable example of such external inputs to the SA and FIRSTLOOK algorithms are cloud classification parameters obtained from the MISR L2 cloud product. An important implication of this dependency is that aerosol processing needs to wait for the cloud product to be generated, creating a time lag that is prohibitive for NRT applications. Typically, the L2 cloud product is generated within about 18 hours of overpass,

and the MISR L2 FIRSTLOOK aerosol processing is completed within about 2 days. In order to produce an L2 aerosol product within an about three-hour time frame, the algorithm needs to operate without the upstream cloud classifiers.

Two specific L2 cloud classification parameters utilized in FIRSTLOOK and SA aerosol processing are the MISR Stereoscopically-Derived Cloud Mask (SDCM) and the Angular Signature Cloud Mask (ASCM) (Diner et al., 1999b; Girolamo and Davies, 1994). In addition to these L2 products, the Radiometric Camera-by-camera Cloud Mask (RCCM) (Diner et al., 1999a; Girolamo and Davies, 1995) retrieved in L1B processing is also employed. All three parameters are reported at 1.1 km x 1.1 km resolution. It should be noted that RCCM also serves as an input to the algorithm that generates SDCM and ASCM, indicating that these parameters are not independent.

In the FIRSTLOOK and SA algorithm, the RCCM, SDCM, and ASCM cloud masks are used together to determine whether a particular 1.1 km x 1.1 km subregion is clear or cloudy. The implication is that if any of the 9 MISR cameras is designated as cloudy in a subregion, this subregion is excluded from aerosol retrieval. The clear/cloudy decision logic depends on the underlying surface type, assigned into three categories: land, water, and snow/ice. Generally, a “clear” outcome is favored over the two most frequently used surface types, land and water, assigning a subregion as cloudy only if the RCCM and SDCM masks indicate a cloud. The logic is considerably more conservative over snow/ice surfaces due to difficulties in distinguishing clouds from the underlying bright features. Details of the cloud mask decision logic over different surface types can be found in Diner et al. (2008).

Analyzing three months of V23 L2 SA product (March, April, May, 2020) indicates that the cloud masks along with the brightness test (see 4.1.2) lead to screening of about 50% of retrievals. As such, they have the largest impact on identifying and removing pixels where clouds might be present. These masks and decision pathways, however, have their deficiencies and additional checks were put in place to further decrease the frequency of cloud-contaminated aerosol retrievals.

4.1.2. Build-in cloud detection methods

In addition to the cloud masks retrieved in the L1B processing (RCCM) and from the L2 Cloud Detection and Classification algorithm (SDCM, ASCM), the MISR aerosol retrieval algorithm relies on three internal tests to further identify cloudy pixels that might have escaped earlier detection. These are (1) the *brightness test*, (2) the *angle-to-angle smoothness test*, and (3) the

angle-to-angle correlation test. Details of these tests can be found in Diner et al. (2008) or Witek et al. (2013), but a short summary is provided here for completeness.

The brightness test is employed to identify clouds that lacked sufficient texture to be picked up by SDCM. For each surface type a fixed threshold is adopted on measured bidirectional reflectance factors (BRFs), and when exceeded in all spectral bands for at least one camera, it renders a subregion unsuitable for aerosol retrieval. The thresholds are set to 1.0, 0.5, and 0.5 for snow/ice, land, and water surfaces, respectively. The value of 1.0 means that the brightness test is effectively turned off over snow/ice. Furthermore, the brightness test does not override subregions that were identified as clear by RCCM.

The angular smoothness test checks for unusually large variations in the measured equivalent reflectances as a function of camera angle, the premise being that in the absence of artifacts or subpixel clouds, the measured radiance should change smoothly from camera to camera. The test is achieved by fitting a polynomial to equivalent reflectances, separately for aft (+nadir) and forward (+nadir) cameras and each spectral band, and checking if the goodness of fit metric (definition in Diner et al., 2008) exceeds a threshold. If in at least one case the test fails, the subregion is eliminated.

Finally, the angle-to-angle correlation test also investigates radiance smoothness and correlation between camera angles, which makes it conceptually similar to the angular smoothness test, but instead utilizes high-resolution information from the red spectral band. It uses 4 x 4 arrays of the 275m spatial resolution red band equivalent reflectances in each 1.1 km x 1.1 km subregion. The test then evaluates spatial variability within the 4 x 4 array for each camera and compares it to a variability within a camera-average template. Variances, covariances, and normalized cross-correlations are calculated (see Diner et al., (2008) for details). If the variability within a camera deviates considerably from the average, this camera might have sub-pixel clouds or other contaminants, and as a result the subregion is excluded from aerosol retrievals.

In the three months of data analyzed in this study (March, April, May 2020), the relative occurrence of retrieval screening due the above-mentioned internal tests are about 4.0% and 0.1% for the correlation and smoothness tests, respectively. These statistics come from analyzing the output field *Aerosol_Retrieval_Screening_Flags* and as such they do not represent the absolute rates of success of each individual test. That is because the tests are performed in a sequential order and if one of them fails, tests that are next in sequence are not performed. For SA product generation, the order is: upstream cloud mask described in 4.1.1, the brightness test, the correlation test, and the smoothness test. For example, the correlation

test is only performed on pixels that already passed the upstream cloud tests as well as the brightness test. Additionally, the brightness test does not have its own flag in the *Aerosol_Retrieval_Screening_Flags* output but is grouped together with the upstream cloud classifiers.

4.2. Retrieval screening using regional cloud parameters

Methods described in section 4.1 focus on identifying and excluding cloudy 1.1 km x 1.1 km subregions from the aerosol retrieval process. The retrieval region consists of 16 (4 x 4) subregions. These methods are highly effective at removing cloud-contaminated pixels, but since they rely on MISR visible wavelengths they might miss certain cloud signatures more easily detected in the infrared spectrum (e.g., Gao et al., 1993). For example, MODIS routinely uses its reflective and emissive infrared channels to detect optically thin cirrus clouds (Ackerman et al., 2010; Levy et al., 2013). As a result, MISR cloud detection methods occasionally fail, which leads to visible outliers in retrieved AODs (Witek et al., 2018b). For that reason, an additional set of screenings is applied in an effort to eliminate such unusually high AOD retrievals (Garay et al., 2020). Two of these additional methods look at overall cloudiness in the retrieval region (consisting of 4 x 4 subregions) as well as in a larger area consisting of 3 x 3 regions (12 x 12 subregions). The Cloud Screening Parameter (CSP) represents the fraction of clear grid cells within a region, whereas Cloud Screening Parameter Neighbor 3x3 (CSP9) is similar to CSP but for the larger area. If CSP is below 0.7 and CSP9 below 0.5, the retrieval is not reported in the final product intended for most users. However, it is still included in the product's AUXILIARY subcategory and annotated with the term "Raw" to indicate that the product has not undergone recommended quality screenings.

4.3. Adjusting cloud screening thresholds

4.3.1. Performance of the prototype NRT product

This subsection presents results and analysis of prototype NRT aerosol retrievals. These are obtained prior to any threshold and screening adjustments included in the final version of the product. To differentiate between the final and the prototype NRT products, the latter is denoted as NRT_{prot}.

As mentioned in the previous section, the NRT processing cannot rely on the cloud masks generated in the L1 and L2 cloud products, namely the RCCM, SDCM, and ASCM. This implies that potentially less screening of cloudy subregions would be applied, increasing the probability of cloud contamination in aerosol retrievals. However, some of the burden of cloud identification is picked up by the built-in cloud tests described in section 4.1.2. The frequency of these tests identifying cloudy pixels increases in NRT processing in comparison to standard processing, in large part mitigating the negative consequences resulting from the lack of the upstream cloud masks. This is well evidenced by examining the normalized probability density functions (*pdfs*) of AOD from spring 2020 (Figure 2). The SA (red) and NRT_{prot} (blue) lines are very similar, indicating that the built-in cloud tests substitute to a significant extent for the missing upstream cloud masks in generating the NRT_{prot} product. The largest difference occurs in the high-AOD range, suggesting that NRT_{prot} has more retrievals in this regime. The black dotted line shows a *pdf* of the NRT_{prot} AOD retrievals that do not have a matching SA retrieval. This is labeled as “NRT_{prot} gained” as it represents additional retrievals obtained in NRT processing due to the lack of external cloud masks. The “NRT_{prot} gained” *pdf* is clearly shifted towards higher AODs, confirming that the NRT_{prot} processing tends to retrieve higher AODs in places where SA is not available.

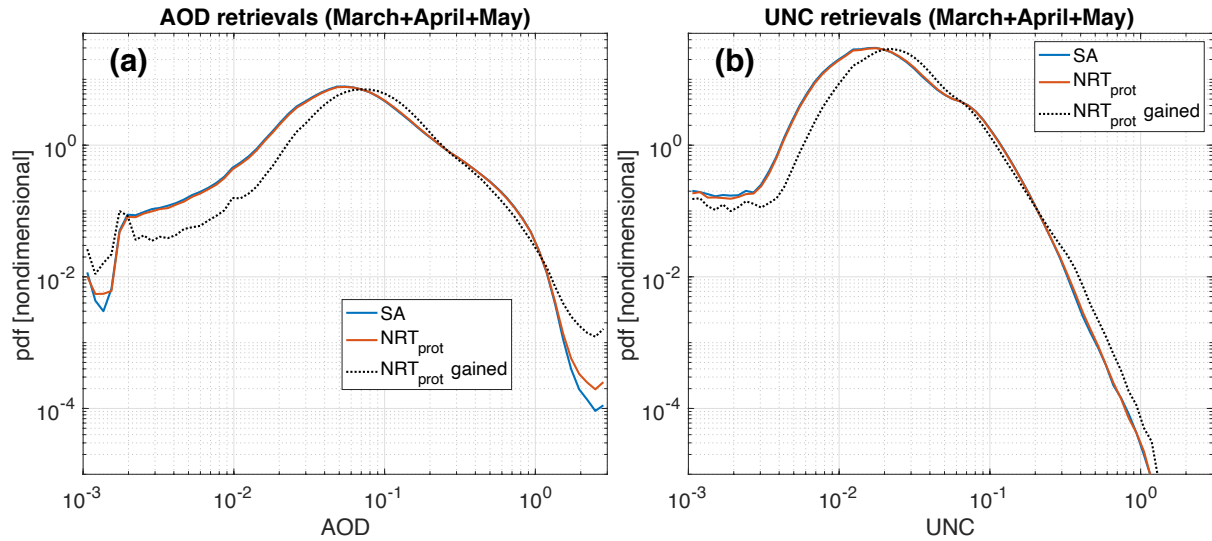


Figure 2 (a) AOD normalized probability density functions from SA, prototype NRT, and prototype NRT retrievals that do not have a matching SA equivalent (labeled as NRT_{prot} gained); (b) same as in (a) but for retrieved AOD uncertainties (UNC). Data statistics for AODs are provided in Table 1.

Figure 3 shows *pdfs* of AOD but with retrievals separated between DW (Fig. 3a) and land (Fig. 3b). These *pdfs* indicate that the retrievals over oceans are the main source of

increased frequency of high-AODs in the NRT_{prot} product. The *pdfs* over land are virtually unchanged, including a slightly flattened but still relatively comparable distribution of the “ NRT_{prot} gained” retrievals (Fig. 3b). The additional statistics of the data presented in Figs. 2 and 3, including the retrieval count, the mean AOD, and the geometric mean AOD, which is better suited for log-normal distributions of AOD (Sayer and Knobelspiesse, 2019), are provided in Table 1. Note that the number of NRT_{prot} gained is not the same as the number of NRT_{prot} minus SA. This is because some SA retrievals do not have their NRT_{prot} equivalent, making the SA count larger than it would have been otherwise.

In the 3-month period analyzed in this study (March, April, May, 2020), the NRT_{prot} processing leads to about 6.4% more retrievals than SA (see Table 1). 5.5 million NRT_{prot} retrievals do not have a matching SA retrieval (NRT gained), and the majority of them (67%) are DW retrievals. The overall geometric means are almost identical in SA and NRT_{prot} , although small variations in this statistic are seen in DW and land categories. The NRT gained have visibly higher mean and geometric mean values, the increase coming mainly from DW retrievals. These basic statistics warrant a further look at the NRT_{prot} performance over DW.

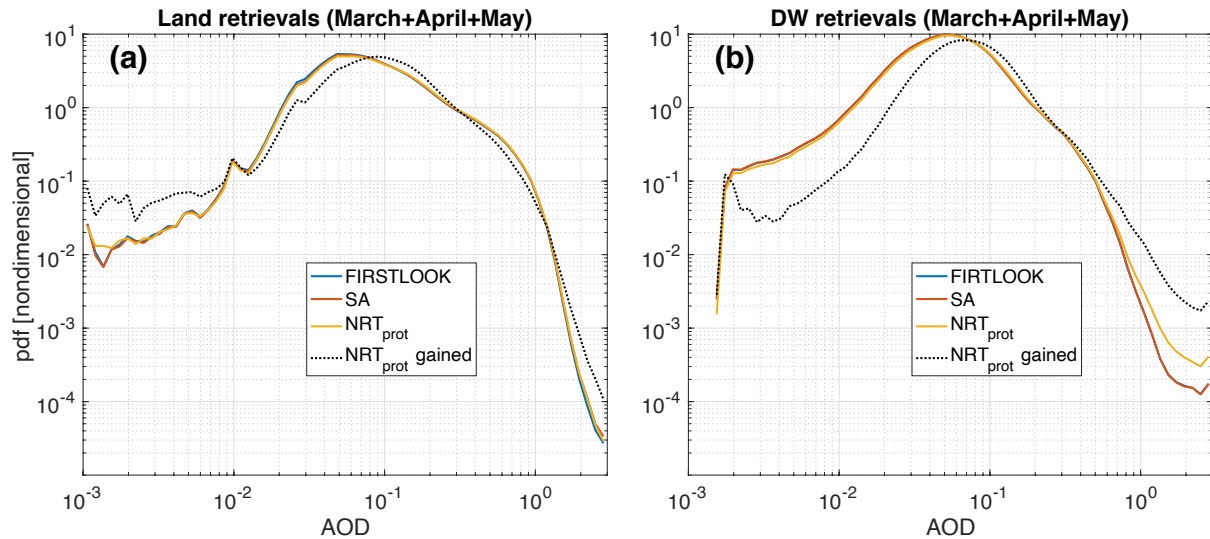


Figure 3 AOD *pdfs* for land (a) and DW (b) retrievals, respectively. Data statistics are provided in Table 1.

| | All retrievals | | | DW | | | Land | | |
|-------------------|----------------|--------------|---------------------|------|--------------|---------------------|------|--------------|---------------------|
| | SA | NRT_{prot} | NRT_{prot} gained | SA | NRT_{prot} | NRT_{prot} gained | SA | NRT_{prot} | NRT_{prot} gained |
| $N (\times 10^6)$ | 49.7 | 52.9 | 5.5 | 27.6 | 30.7 | 3.7 | 22.1 | 22.2 | 1.8 |

| | | | | | | | | | |
|----------------|-------|-------|-------|-------|-------|-------|-------|-------|-------|
| <i>mean</i> | 0.168 | 0.169 | 0.171 | 0.111 | 0.115 | 0.146 | 0.240 | 0.243 | 0.224 |
| <i>geomean</i> | 0.111 | 0.112 | 0.122 | 0.083 | 0.085 | 0.106 | 0.160 | 0.162 | 0.161 |

Table 1 Additional statistics for the data presented in Figs. 2 and 3 (statistic for FIRSTLOOK not shown). NRT gained stands for the prototype NRT retrievals that do not have a matching SA equivalent; geomean stands for the geometric mean AOD.

4.3.2. Sensitivity to CSP and CSP9 thresholds in DW retrievals

One way to screen potentially cloud-contaminated high-AOD retrievals is to adjust thresholds on CSP and CSP9 parameters (Garay et al., 2020). This is furthermore justified by the fact that in the absence of RCCM, SDCM, and ASCM in NRT_{prot} processing, fewer cloudy subregions are identified in a retrieval area and consequently CSP and CSP9 have by default lower values. This argument provides strong justification for investigating sensitivity to increased CSP and CSP9 thresholds in the NRT_{prot} processing.

The SA product uses the thresholds of CSP=0.7 and CSP9=0.5 (Garay et al., 2020); when the values of CSP and CSP9 are below these thresholds in a retrieval region, the aerosol retrieval is removed from the data field recommended for users. Figure 4 and Table 2 show *pdfs* and AOD statistics for different thresholds of CSP and CSP9 parameters in the NRT_{prot} product over dark water surfaces. There are only minor changes in the *pdfs* when the thresholds are increased, including in the high-AOD regime. The mean and geometric mean decrease gradually but slowly; even at the highest considered thresholds (0.85 for CSP and 0.75 for CSP9) these statistics are still above the SA values. At the same time the number of passing NRT_{prot} retrievals decreases considerably faster, with almost 19% of retrievals lost when the highest thresholds are used. These results indicate that adjusting CSP and CSP9 thresholds is not an effective strategy to constraining NRT_{prot} retrievals.

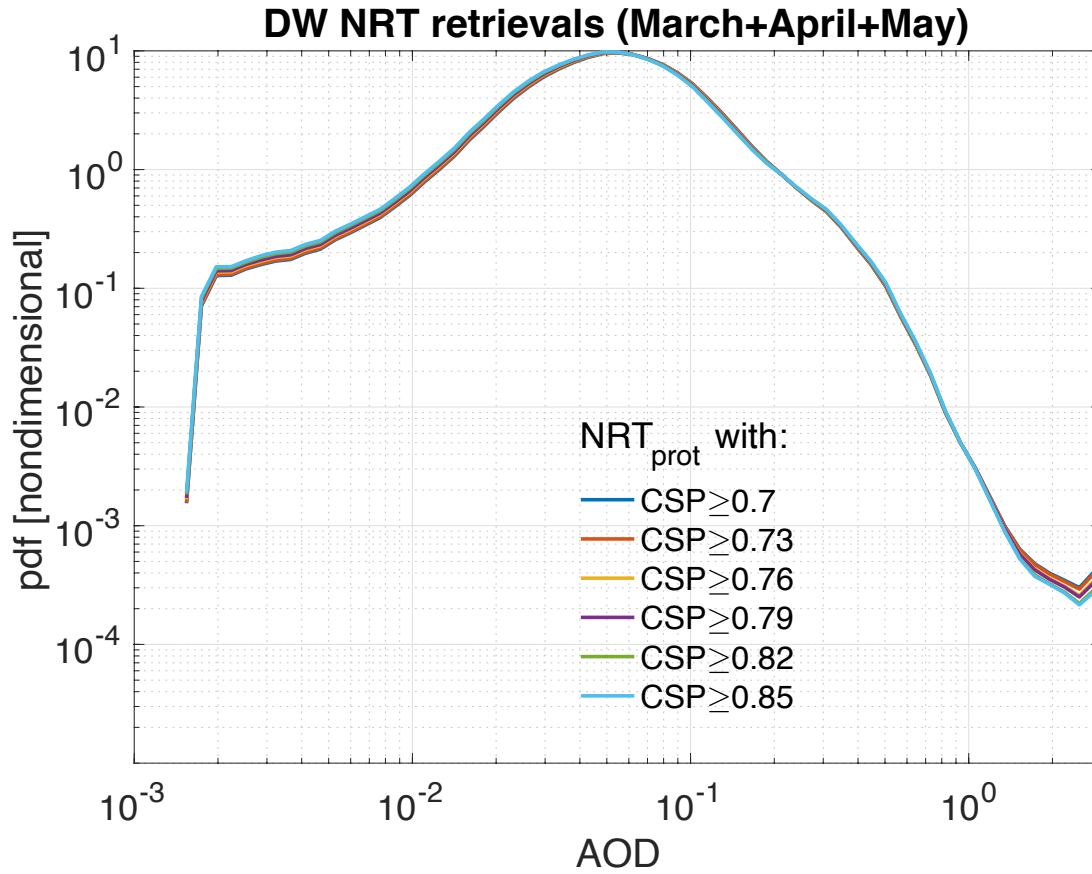


Figure 4 Prototype NRT AOD pdfs over dark water surfaces from spring 2020 obtained with different CSP and CSP9 cloud-screening thresholds. Data statistics are provided in Table 2.

| | | | | | | | |
|-------------------|------------------------|------------------------|------------------------|------------------------|------------------------|------------------------|------------------------|
| $N (\times 10^6)$ | 30.7 | 30.1 (-1.9%) | 28.4 (-7.4%) | 27.7 (-9.8%) | 25.9 (-15.6%) | 24.9 (-18.9%) | SA 27.6 |
| CSP | ≥ 0.7 | ≥ 0.73 | ≥ 0.76 | ≥ 0.79 | ≥ 0.82 | ≥ 0.85 | |
| CSP9 | ≥ 0.5 | ≥ 0.55 | ≥ 0.6 | ≥ 0.65 | ≥ 0.7 | ≥ 0.75 | |
| mean | 0.1151 ± 0.1200 | 0.1149 ± 0.1199 | 0.1145 ± 0.1190 | 0.1144 ± 0.1191 | 0.1142 ± 0.1185 | 0.1143 ± 0.1189 | 0.1110 ± 0.1079 |
| geomean | 0.0850 | 0.0847 | 0.0841 | 0.0839 | 0.0834 | 0.0832 | 0.0826 |

Table 2 Additional statistics for the data presented in Fig. 4. Values for CSP and CSP9 indicate their corresponding thresholds for screening AOD retrievals. The arithmetic mean values are accompanied by their respective \pm one standard deviations.

4.3.3. Sensitivity to ARCI threshold in DW retrievals

V23 of the MISR aerosol product introduced a new parameter, called the aerosol retrieval confidence index (ARCI), that is used to screen high-AOD retrieval outliers caused by cloud contamination and other factors (Witek et al., 2018b). ARCI, defined only for DW retrievals, proved to be an efficient metric at filtering out potentially cloud-contaminated AOD retrievals. In standard processing, retrievals with $ARCI < 0.15$ are removed from the recommended user field, but are retained in the AUXILIARY group. The 0.15 threshold is well supported through statistical analysis (Witek et al., 2018b), although some erroneous AODs still pass this screening method, suggesting that increasing this threshold might be beneficial in NRT processing.

Figure 5 and Table 3 show *pdfs* and AOD statistics for different thresholds of ARCI in the NRT_{prot} product. In this case the differences between ARCI thresholds are quite noticeable, especially in the high-AOD range of retrievals. Increasing the ARCI threshold to 0.2 leads to a loss of about 11% of NRT_{prot} DW retrievals, but the resulting mean and geometric mean are lower than the SA values. At the same time, the absolute number of NRT_{prot} DW retrievals (27.4 million) is still comparable to the number of SA DW retrievals (27.6 million). The *pdfs* and the statistics suggest that increasing the NRT_{prot} ARCI threshold from 0.15 to 0.18 leads to a product that has similar characteristics to SA.

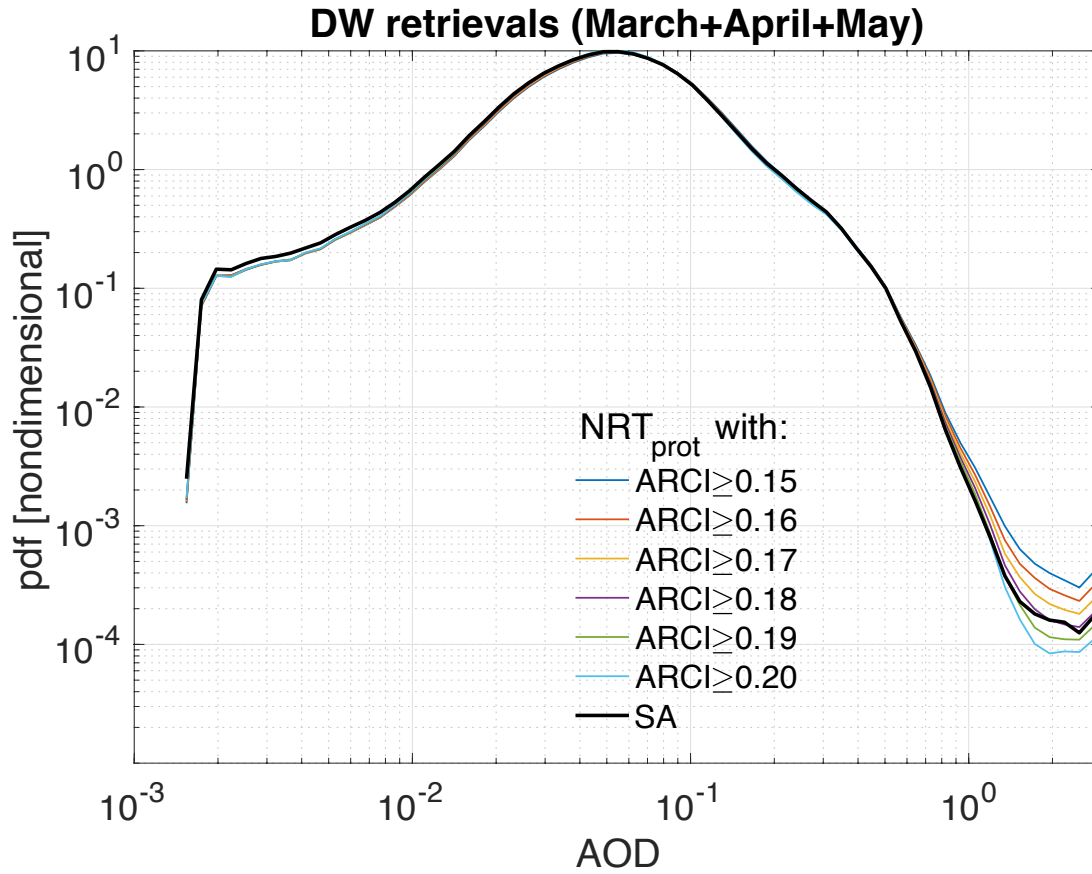


Figure 5 Prototype NRT AOD pdfs from spring 2020 obtained with different ARCI thresholds. Data statistic are provided in Table 3.

| | | | | | | | |
|-------------------|------------------------|------------------------|------------------------|------------------------|------------------------|------------------------|------------------------|
| $N (\times 10^6)$ | 30.7 | 30.0 (-2.2%) | 29.4 (-4.3%) | 28.7 (-6.5%) | 28.0 (-8.6%) | 27.4 (-10.8%) | SA 27.6 |
| ARCI | ≥ 0.15 | ≥ 0.16 | ≥ 0.17 | ≥ 0.18 | ≥ 0.19 | ≥ 0.20 | |
| mean | 0.1151 ± 0.1200 | 0.1137 ± 0.1157 | 0.1124 ± 0.1122 | 0.1112 ± 0.1094 | 0.1100 ± 0.1070 | 0.1090 ± 0.1051 | 0.1110 ± 0.1079 |
| geomean | 0.0850 | 0.0842 | 0.0835 | 0.0828 | 0.0821 | 0.0813 | 0.0826 |

Table 3 Additional statistic for the data presented in Fig. 5.

4.3.4. Recommendation for NRT processing

The statistical analyses presented in the previous sections indicate that the lack of RCCM, SDCM, and ASCM in NRT processing has negative consequences on the product, especially by

allowing more, potentially cloud-contaminated, high-AOD DW retrievals to pass screening criteria. Adjusting build-in cloud screening thresholds on CSP and CSP9 brings only limited benefits at the cost of losing a considerable percentage of retrievals. However, the ARCI threshold adjustments result in much closer statistical correspondence between the NRT_{prot} and standard AOD retrievals. For that reason, a revised ARCI threshold of 0.18 is implemented in NRT processing. Since the unscreened retrievals, as well as the ARCI parameter, are also provided in the AUXILIARY group of the product, users are encouraged to experiment with their own thresholds which might prove more beneficial in specific applications or geographic areas.

4.4. Cloud/clear decision logic over snow/ice

In section 4.1.1 the impact of upstream cloud classifiers in standard processing—namely the RCCM, SDCM, and ASCM—on the subregion’s cloud/clear designation was briefly described. The decision pathway depends on the underlying surface type, which can be either land, water, or snow/ice. Over land and water, the “cloud” outcome is only obtained when both RCCM and SDCM designate the subregion as cloudy. In the absence of RCCM and SDCM the default outcome is “clear”. Over snow/ice, however, the logic is more restrictive and favors the “cloudy” designation (Diner et al., 2008). Specifically, when the upstream cloud classifiers are not available, the subregion designation is set to “cloudy” by default. This has important implications on aerosol retrievals in areas where snow and ice occur seasonally.

The snow/ice surface mask, unlike land and water, is not static and changes every month. Furthermore, the snow/ice mask input to MISR aerosol processing has a 1.0-degree horizontal resolution, which is re-gridded to a 1.1 km resolution corresponding to the resolution of MISR subregion. In FIRSTLOOK processing, the snow/ice mask from the same month but in the previous year is used. The final SA processing is performed when the current year’s monthly snow/ice mask becomes available. The NRT processing, similarly to FIRSTLOOK, relies on the previous year’s snow/ice mask. Additionally, given the lack of upstream cloud classifiers, the snow/ice areas are designated as “cloudy” for aerosol retrieval purposes. This is well visualized in Figure 6 which shows the visible image and the corresponding maps of AOD and Aerosol Retrieval Screening Flag in the NRT processing. The dark blue color (index 5) denotes cloudy regions determined using the snow/ice cloud logic. The box-like nature of the excluded areas is associated with the coarse resolution of the snow/ice mask (1.0 degree). The previous year’s mask might also not be representative of the current conditions on the ground. It is worth noting that the FIRSTLOOK product often suffers from the same exclusion rules as NRT. This is

because of the strict clear/cloud logic over snow/ice surfaces which favors the cloudy outcome; in the case shown in Fig. 6 the AOD gaps in FIRSTLOOK (not shown) look very similar to the NRT product.

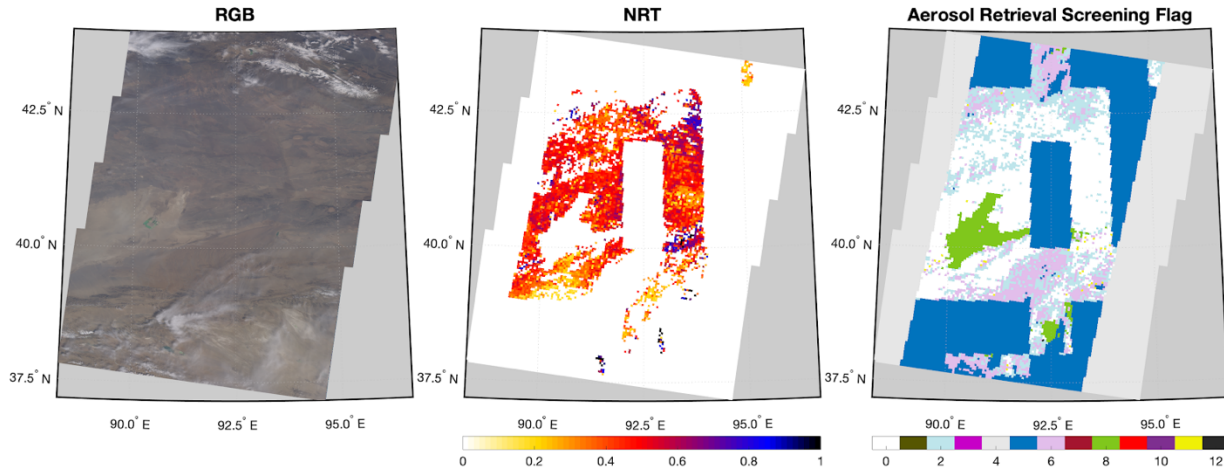


Figure 6 Example of snow/ice masking in NRT AOD retrievals. (Left) Visible image of the retrieval area. (Center) Corresponding NRT AOD retrievals. (Right) NRT Aerosol Retrieval Screening Flag for the same area; the dark blue color denotes regions designated as cloudy.

Several attempts have been made by the MISR science team to improve NRT aerosol retrievals in snow/ice covered areas. However, identifying and isolating snow-covered surfaces in the absence of upstream cloud classifiers proves very challenging. The quality of aerosol retrievals is often negatively affected in such conditions. For that reason, and in an attempt to eliminate as many NRT AOD outliers as possible, the current snow/ice logic is retained in the NRT aerosol processing.

5. NRT and SA differences

In this section, geographic distributions of MISR AOD retrievals from SA and NRT products are analyzed. The datasets encompass three months, March, April, and May of 2020. The NRT AOD retrievals are screened with the revised ARCI threshold of 0.18 as suggested in section 4.3.4. The spatial overlap of the SA and NRT data is achieved using an intersect of the X_Dim and Y_Dim fields in the two data products.

Figure 7 shows the global distributions of geometric mean AOD from the (a) SA and (b) NRT products. The retrievals are gridded at 2-by-2-degree spatial resolution. Fig. 7c shows the AOD difference between the two products (NRT – SA).

The largest AOD differences are seen in areas with climatologically high cloud cover, especially over the Southern Ocean, and over land in areas where potential snow cover could be an issue. Over the Southern Ocean the SA AODs are predominantly higher than the NRT AODs. This is due to the increased ARCI threshold in NRT (0.18 vs. 0.15 in SA) which brings in more aggressive screening of cloud-contaminated retrievals (Witek et al., 2018b). Over land, where the ARCI parameter is not available, the gridded NRT AODs tend to be higher than the SA AODs, which is in part related to the differences in snow/ice mask between the two products. Still, the AOD differences in Fig. 7c are rather small and reflect sampling issues rather than any systematic deficiencies in NRT processing. At the same time the lack of cloud classifiers in NRT does not adversely affect AOD distributions, which is consistent with the statistical analysis presented in section 4.2.3.

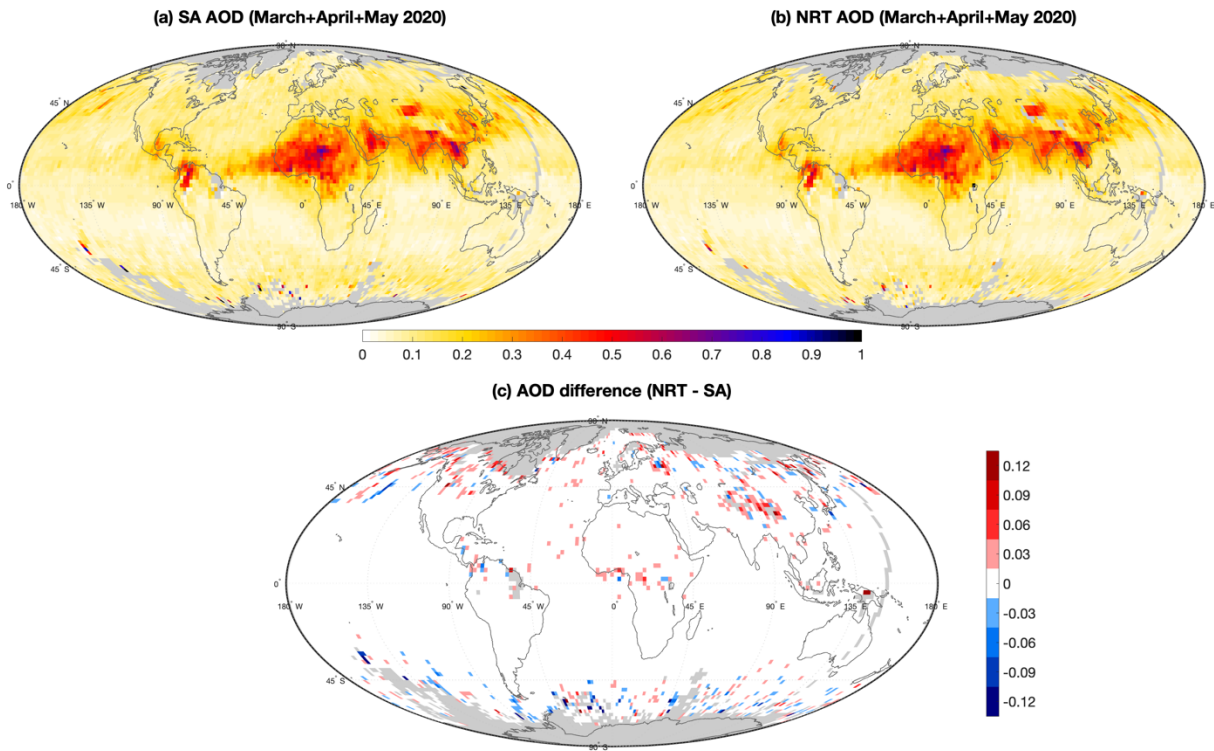


Figure 7 (a) Global distribution of SA AOD geometric mean values across March, April, and May of 2020 on a 2-by-2-degree spatial resolution; (b) same as in (a) but for NRT AOD; and (c) AOD difference between SA and NRT. Grid points with less than 15 retrievals are excluded.

Figure 8 complements Fig. 7 by showing (a) the SA retrieval count distribution as well as (b) the retrieval count difference between the SA and NRT products.

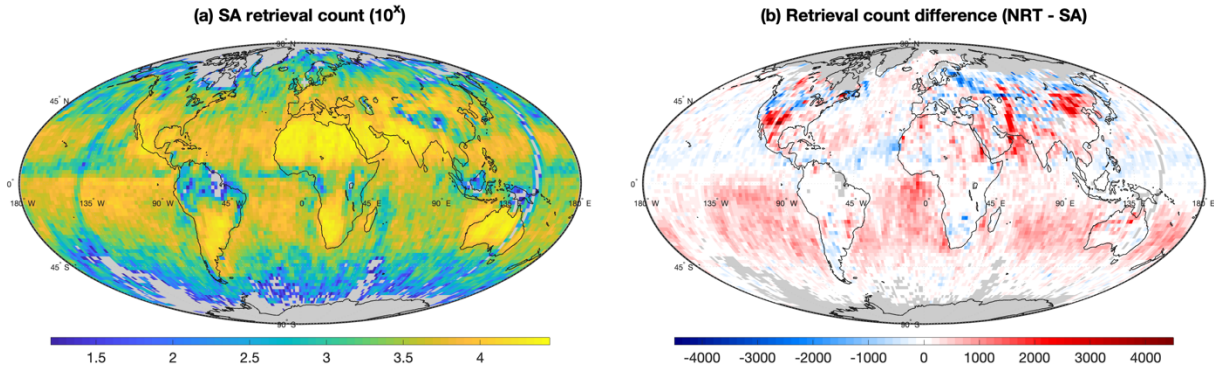


Figure 8 (a) Decimal logarithm of the retrieval count from the SA product in March, April, and May of 2020; (b) retrieval count difference between SA and NRT. Presented values are gridded at 2-by-2-degree spatial resolution and grid points with less than 15 retrievals are excluded.

The highest number of retrievals is found over the subtropical continents where the cloud cover is usually the smallest. Over the subtropical oceans in the Southern Hemisphere the NRT retrieval counts are typically higher than in SA, which results from the absence of upstream cloud classifiers in NRT processing and subsequently fewer subregions being excluded as cloudy. Note that this increase in retrieval count caused by the lack of cloud classifiers is not compensated by the increased ARCI threshold in NRT processing ($\text{ARCI} \geq 0.18$), which always reduces the number of retrievals when compared to the default SA threshold ($\text{ARCI} \geq 0.15$). The lack of hemispheric symmetry in this case is likely due to the seasonal variability (only months in northern spring are analyzed here). Over land the lack of upstream cloud classifiers also results in higher number of NRT retrievals in certain regions, but the surface type exclusion rules reverse this pattern, especially at higher latitudes. The conservative cloud logic over snow/ice surfaces in NRT processing often results in the lower number of NRT retrievals in the high latitudes of the northern hemisphere.

6. Summary

The MISR V23 aerosol product, publicly available since mid-2018, is a high-resolution state-of-the-art data product from NASA's Terra flagship mission. V23 AOD retrievals have remarkable accuracy compared against ground-based observations (Garay et al., 2020; Tao et al., 2020; Witek et al., 2019) and the product is more intuitive and easier to use than previous versions. The product is available within 2 days from satellite overpass as a FIRSTLOOK version, and within 3-to-6 months as a final science-quality SA version that employs the most up-to-date

ancillary datasets. In response to the needs of operational user communities, a new MISR L2 NRT aerosol product has been developed with a 3-hour latency.

The new NRT algorithm does not depend on the upstream cloud classifiers that are generated in L1 and L2 cloud processing. The lack of cloud classifiers is in large part mitigated by the aerosol algorithm's built-in cloud identification methods. Analysis of the prototype NRT product has shown an increased frequency of high-AOD retrievals, especially over oceans and in climatologically cloudy areas, likely due to an increase in cloud contamination. Adjusting the ARCI threshold in DW retrievals proves highly effective at eliminating some of these high-AOD outliers and improves the NRT product's statistical agreement with the SA version. The new NRT aerosol product applies an ARCI threshold of 0.18 to mitigate cloud contamination in the absence of upstream cloud masks in NRT processing. The remaining differences in statistical and geographic distributions between the NRT and SA AODs, which includes information from the L2 cloud product, are small and largely confined to areas with high cloud cover.

The results of this study also serve as an example of the effects of screening threshold adjustments in MISR aerosol retrievals on AOD statistics and distributions. Researchers interested in particular applications and/or specific geographic regions are encouraged to experiment with their own threshold to achieve most optimal results. The NRT aerosol product contains both the recommended product contained within the main science directory "4.4_KM_PRODUCTS" that has the stricter ARCI threshold ($ARCI \geq 0.18$), and the unscreened product without the additional cloud and ARCI filtering designed for more experienced users, located within the AUXILIARY group.

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Data availability

The MISR V23 SA and NRT data is publicly available and can be downloaded from <https://l0dup05.larc.nasa.gov/cgi-bin/MISR/main.cgi>. MISR NRT data is not stored permanently and is only available for three to six months from the time of acquisition; please contact the corresponding author to request the NRT data from the months analyzed in this study.

512 **References:**

- 513 Ackerman, S., Richard, F., Kathleen, S., Yinghui, L., Liam, G., Bryan, B. and Paul, M.:
514 Discriminating clear-sky from cloud with MODIS algorithm theoretical basis document
515 (MOD35), Univ. Wisconsin - Madison, 6th Edn.(October), 129 [online] Available from:
516 <http://citeseerx.ist.psu.edu/viewdoc/summary?doi=10.1.1.385.4885>, 2010.
- 517 Benedetti, A., Reid, J. S. and Colarco, P. R.: International cooperative for aerosol prediction
518 workshop on aerosol forecast verification, in Bulletin of the American Meteorological Society,
519 vol. 92., 2011.
- 520 Bocquet, M., Elbern, H., Eskes, H., Hirtl, M., Aabkar, R., Carmichael, G. R., Flemming, J.,
521 Inness, A., Pagowski, M., Pérez Camaño, J. L., Saide, P. E., San Jose, R., Sofiev, M., Vira, J.,
522 Baklanov, A., Carnevale, C., Grell, G. and Seigneur, C.: Data assimilation in atmospheric
523 chemistry models: Current status and future prospects for coupled chemistry meteorology
524 models, *Atmos. Chem. Phys.*, 15(10), 5325–5358, doi:10.5194/acp-15-5325-2015, 2015.
- 525 Buchard, V., Da Silva, A. M., Colarco, P. R., Darmenov, A., Randles, C. A., Govindaraju, R.,
526 Torres, O., Campbell, J. and Spurr, R.: Using the OMI aerosol index and absorption aerosol
527 optical depth to evaluate the NASA MERRA Aerosol Reanalysis, *Atmos. Chem. Phys.*, 15(10),
528 5743–5760, doi:10.5194/acp-15-5743-2015, 2015.
- 529 Buchard, V., Randles, C. A., da Silva, A. M., Darmenov, A., Colarco, P. R., Govindaraju, R.,
530 Ferrare, R., Hair, J., Beyersdorf, A. J., Ziemba, L. D. and Yu, H.: The MERRA-2 aerosol
531 reanalysis, 1980 onward. Part II: Evaluation and case studies, *J. Clim.*, 30(17),
532 doi:10.1175/JCLI-D-16-0613.1, 2017.
- 533 Butz, A., Hasekamp, O. P., Frankenberg, C. and Aben, U.: Retrievals of atmospheric CO₂ from
534 simulated space-borne measurements of backscattered near-infrared sunlight: Accounting for
535 aerosol effects, *Appl. Opt.*, 48, 3322–3336, doi:10.1364/AO.48.003322, 2009.
- 536 Choi, M., Lim, H., Kim, J., Lee, S., Eck, T. F., Holben, B. N., Garay, M. J., Hyer, E. J., Saide, P.
537 E. and Liu, H.: Validation, comparison, and integration of GOCI, AHI, MODIS, MISR, and
538 VIIRS aerosol optical depth over East Asia during the 2016 KORUS-AQ campaign, *Atmos.*

539 Meas. Tech., 12(8), 4619–4641, doi:10.5194/amt-12-4619-2019, 2019.

540 Colarco, P., Da Silva, A., Chin, M. and Diehl, T.: Online simulations of global aerosol
541 distributions in the NASA GEOS-4 model and comparisons to satellite and ground-based aerosol
542 optical depth, *J. Geophys. Res. Atmos.*, 115(14), doi:10.1029/2009JD012820, 2010.

543 Diner, D. J., Beckert, J. C., Reilly, T. H., Bruegge, C. J., Conel, J. E., Kahn, R. A., Martonchik,
544 J. V., Ackerman, T. P., Davies, R., Gerstl, S. A. W., Gordon, H. R., Muller, J. P., Myneni, R. B.,
545 Sellers, P. J., Pinty, B. and Verstraete, M. M.: Multiangle Image Spectroradiometer (MISR)
546 instrument description and experiment overview, *IEEE Trans. Geosci. Remote Sens.*, 36(4),
547 1072–1087, 1998.

548 Diner, D. J., Di Girolamo, L. and Clothiaux, E. E.: Level 1 Cloud Detection Algorithm
549 Theoretical Basis, Jet Propuls. Lab. Calif. Inst. Technol., D-13397(Rev. B), 1999a.

550 Diner, D. J., Davies, R., Di Girolamo, L., Horvath, A., Moroney, C., Muller, J. P., Paradise, S.
551 R., Wenkert, D. and Zong, J.: Level 2 Cloud Detection and Classification Algorithm Theoretical
552 Basis, Jet Propuls. Lab. Calif. Inst. Technol., D-11399(Rev. D), 1999b.

553 Diner, D. J., Abdou, W. A., Ackerman, T. P., Crean, K., Gordon, H. R., Kahn, R. A.,
554 Martonchik, J. V., McMuldroch, S., Paradise, S. R. and Pinty, B.: Level 2 aerosol retrieval
555 algorithm theoretical basis, Jet Propuls. Lab. Calif. Inst. Technol., D-11400(Rev. G), 2008.

556 Frankenberg, C., Hasekamp, O., O’Dell, C., Sanghavi, S., Butz, A. and Worden, J.: Aerosol
557 information content analysis of multi-angle high spectral resolution measurements and its benefit
558 for high accuracy greenhouse gas retrievals, *Atmos. Meas. Tech.*, 5(7), 1809–1821,
559 doi:10.5194/amt-5-1809-2012, 2012.

560 Frouin, R. J., Franz, B. A., Ibrahim, A., Knobelspiesse, K., Ahmad, Z., Cairns, B., Chowdhary,
561 J., Dierssen, H. M., Tan, J., Dubovik, O., Huang, X., Davis, A. B., Kalashnikova, O., Thompson,
562 D. R., Remer, L. A., Boss, E., Coddington, O., Deschamps, P. Y., Gao, B. C., Gross, L.,
563 Hasekamp, O., Omar, A., Pelletier, B., Ramon, D., Steinmetz, F. and Zhai, P. W.: Atmospheric
564 Correction of Satellite Ocean-Color Imagery During the PACE Era, *Front. Earth Sci.*, 7,
565 doi:10.3389/feart.2019.00145, 2019.

566 Fu, G., Prata, F., Xiang Lin, H., Heemink, A., Segers, A. and Lu, S.: Data assimilation for
 567 volcanic ash plumes using a satellite observational operator: A case study on the 2010
 568 Eyjafjallajökull volcanic eruption, *Atmos. Chem. Phys.*, 17(2), 1187–1205, doi:10.5194/acp-17-
 569 1187-2017, 2017.

570 Gao, B. C., Goetz, A. F. H. and Wiscombe, W. J.: Cirrus cloud detection from Airborne Imaging
 571 Spectrometer data using the 1.38 μm water vapor band, *Geophys. Res. Lett.*, 20(3), 301–304,
 572 doi:10.1029/93GL00106, 1993.

573 Garay, M. J., Witek, M. L., Kahn, R. A., Seidel, F. C., Limbacher, J. A., Bull, M. A., Diner, D.
 574 J., Hansen, E. G. E. G., Kalashnikova, O. V., Lee, H., Nاستan, A. M. and Yu, Y.: Introducing the
 575 4.4km spatial resolution Multi-Angle Imaging SpectroRadiometer (MISR) aerosol product,
 576 *Atmos. Meas. Tech.*, 13(2), 593–628, doi:10.5194/amt-13-593-2020, 2020.

577 Gelaro, R., McCarty, W., Suárez, M. J., Todling, R., Molod, A., Takacs, L., Randles, C. A.,
 578 Darmenov, A., Bosilovich, M. G., Reichle, R., Wargan, K., Coy, L., Cullather, R., Draper, C.,
 579 Akella, S., Buchard, V., Conaty, A., da Silva, A. M., Gu, W., Kim, G. K., Koster, R., Lucchesi,
 580 R., Merkova, D., Nielsen, J. E., Partyka, G., Pawson, S., Putman, W., Rienecker, M., Schubert,
 581 S. D., Sienkiewicz, M. and Zhao, B.: The modern-era retrospective analysis for research and
 582 applications, version 2 (MERRA-2), *J. Clim.*, 30(14), 5419–5454, doi:10.1175/JCLI-D-16-
 583 0758.1, 2017.

584 Girolamo, L. Di and Davies, R.: A Band-Differenced Angular Signature Technique for Cirrus
 585 Cloud Detection, *IEEE Trans. Geosci. Remote Sens.*, 32(4), 890–896, doi:10.1109/36.298017,
 586 1994.

587 Girolamo, L. Di and Davies, R.: The Image Navigation Cloud Mask for the Multiangle Imaging
 588 Spectroradiometer (MISR), *J. Atmos. Ocean. Technol.*, 12(6), doi:10.1175/1520-
 589 0426(1995)012<1215:tincmf>2.0.co;2, 1995.

590 Gordon, R.: Atmospheric correction of ocean color imagery in the Earth Observing System era,
 591 *J. Geophys. Res. - Atmos.*, 102(D14), 17081–17106, doi:10.1029/96JD02443, 1997.

592 Houweling, S., Hartmann, W., Aben, I., Schrijver, H., Skidmore, J., Roelofs, G. J. and Breon, F.

593 M.: Evidence of systematic errors in SCIAMACHY-observed CO₂ due to aerosols, *Atmos.*
594 *Chem. Phys.*, 5(11), 3003–3013, doi:10.5194/acp-5-3003-2005, 2005.

595 Inness, A., Baier, F., Benedetti, A., Bouarar, I., Chabrillat, S., Clark, H., Clerbaux, C., Coheur,
596 P., Engelen, R. J., Errera, Q., Flemming, J., George, M., Granier, C., Hadji-Lazaro, J., Huijnen,
597 V., Hurtmans, D., Jones, L., Kaiser, J. W., Kapsomenakis, J., Lefever, K., Leitão, J., Razinger,
598 M., Richter, A., Schultz, M. G., Simmons, A. J., Suttie, M., Stein, O., Thépaut, J. N., Thouret,
599 V., Vrekoussis, M. and Zerefos, C.: The MACC reanalysis: An 8 yr data set of atmospheric
600 composition, *Atmos. Chem. Phys.*, 13(8), 4073–4109, doi:10.5194/acp-13-4073-2013, 2013.

601 Inness, A., Ades, M., Agustí-Panareda, A., Barré, J., Benedictow, A., Blechschmidt, A.-M.,
602 Dominguez, J. J., Engelen, R., Eskes, H., Flemming, J., Huijnen, V., Jones, L., Kipling, Z.,
603 Massart, S., Parrington, M., Peuch, V.-H., Razinger, M., Remy, S., Schulz, M. and Suttie, M.:
604 The CAMS reanalysis of atmospheric composition, *Atmos. Chem. Phys.*, 19(6), 3515–3556,
605 doi:10.5194/acp-19-3515-2019, 2019.

606 IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to
607 the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, edited by T. F.
608 Stocker, D. Qin, G. K. Plattner, M. M. B. Tignor, S. K. Allen, J. Boschung, A. Nauels, Y. Xia,
609 V. Bex, and P. M. Midgley, Cambridge University Press, Cambridge, United Kingdom and New
610 York, NY, USA., 2013.

611 Kahn, R. A. and Gaitley, B. J.: An analysis of global aerosol type as retrieved by MISR, *J.*
612 *Geophys. Res. Atmos.*, 120(9), 4248–4281, doi:10.1002/2015JD023322, 2015.

613 Kahn, R. A., Gaitley, B. J., Garay, M. J., Diner, D. J., Eck, T. F., Smirnov, A. and Holben, B. N.:
614 Multiangle Imaging SpectroRadiometer global aerosol product assessment by comparison with
615 the Aerosol Robotic Network, *J. Geophys. Res. Atmos.*, 115(23), doi:10.1029/2010JD014601,
616 2010.

617 Kalashnikova, O. V., Garay, M. J., Martonchik, J. V. and Diner, D. J.: MISR Dark Water aerosol
618 retrievals: Operational algorithm sensitivity to particle non-sphericity, *Atmos. Meas. Tech.*, 6(8),
619 2131–2154, doi:10.5194/amt-6-2131-2013, 2013.

620 Lamarque, J. F., Shindell, D. T., Josse, B., Young, P. J., Cionni, I., Eyring, V., Bergmann, D.,
 621 Cameron-Smith, P., Collins, W. J., Doherty, R., Dalsoren, S., Faluvegi, G., Folberth, G., Ghan,
 622 S. J., Horowitz, L. W., Lee, Y. H., MacKenzie, I. A., Nagashima, T., Naik, V., Plummer, D.,
 623 Righi, M., Rumbold, S. T., Schulz, M., Skeie, R. B., Stevenson, D. S., Strode, S., Sudo, K.,
 624 Szopa, S., Voulgarakis, A. and Zeng, G.: The atmospheric chemistry and climate model
 625 intercomparison Project (ACCMIP): Overview and description of models, simulations and
 626 climate diagnostics, *Geosci. Model Dev.*, 6(1), 179–206, doi:10.5194/gmd-6-179-2013, 2013.

627 Lelieveld, J., Evans, J. S., Fnais, M., Giannadaki, D. and Pozzer, A.: The contribution of outdoor
 628 air pollution sources to premature mortality on a global scale, *Nature*, 525(7569), 367–371,
 629 doi:10.1038/nature15371, 2015.

630 Levy, R. C., Mattoo, S., Munchak, L. A., Remer, L. A., Sayer, A. M., Patadia, F. and Hsu, N. C.:
 631 The Collection 6 MODIS aerosol products over land and ocean, *Atmos. Meas. Tech.*, 6(11),
 632 2989–3034, doi:10.5194/amt-6-2989-2013, 2013.

633 Liu, M., Westphal, D. L., Walker, A. L., Holt, T. R., Richardson, K. A. and Miller, S. D.:
 634 COAMPS real-time dust storm forecasting during operation Iraqi freedom, *Weather Forecast.*,
 635 22(1), 192–206, doi:10.1175/WAF971.1, 2007.

636 Lynch, P., Reid, J. S., Westphal, D. L., Zhang, J., Hogan, T. F., Hyer, E. J., Curtis, C. A., Hegg,
 637 D. A., Shi, Y., Campbell, J. R., Rubin, J. I., Sessions, W. R., Turk, F. J. and Walker, A. L.: An
 638 11-year global gridded aerosol optical thickness reanalysis (v1.0) for atmospheric and climate
 639 sciences, *Geosci. Model Dev.*, 9(4), 1489–1522, doi:10.5194/gmd-9-1489-2016, 2016.

640 Martonchik, J. V., Kahn, R. A. and Diner, D. J.: Retrieval of aerosol properties over land using
 641 MISR observations, in *Satellite Aerosol Remote Sensing over Land*, pp. 267–293, Springer
 642 Berlin Heidelberg., 2009.

643 Randles, C. A., da Silva, A. M., Buchard, V., Colarco, P. R., Darmenov, A., Govindaraju, R.,
 644 Smirnov, A., Holben, B., Ferrare, R., Hair, J., Shinozuka, Y. and Flynn, C. J.: The MERRA-2
 645 aerosol reanalysis, 1980 onward. Part I: System description and data assimilation evaluation, *J.*
 646 *Clim.*, 30(17), 6823–6850, doi:10.1175/JCLI-D-16-0609.1, 2017.

647 Reid, J. S., Benedetti, A., Colarco, P. R. and Hansen, J. A.: International operational aerosol
 648 observability workshop, in Bulletin of the American Meteorological Society, vol. 92., 2011.

649 Rienecker, M. M., Suarez, M. J., Gelaro, R., Todling, R., Bacmeister, J., Liu, E., Bosilovich, M.
 650 G., Schubert, S. D., Takacs, L., Kim, G. K., Bloom, S., Chen, J., Collins, D., Conaty, A., Da
 651 Silva, A., Gu, W., Joiner, J., Koster, R. D., Lucchesi, R., Molod, A., Owens, T., Pawson, S.,
 652 Pegion, P., Redder, C. R., Reichle, R., Robertson, F. R., Ruddick, A. G., Sienkiewicz, M. and
 653 Woollen, J.: MERRA: NASA's modern-era retrospective analysis for research and applications,
 654 J. Clim., 24, 3624–3648, doi:10.1175/JCLI-D-11-00015.1, 2011.

655 Sayer, A. M. and Knobelspiesse, K. D.: How should we aggregate data? Methods accounting for
 656 the numerical distributions, with an assessment of aerosol optical depth, Atmos. Chem. Phys.,
 657 19(23), 15023–15048, doi:10.5194/acp-19-15023-2019, 2019.

658 Sayer, A. M., Govaerts, Y., Kolmonen, P., Lipponen, A., Luffarelli, M., Mielonen, T., Patadia,
 659 F., Popp, T., Povey, A. C., Stebel, K. and Witek, M. L.: A review and framework for the
 660 evaluation of pixel-level uncertainty estimates in satellite aerosol remote sensing, Atmos. Meas.
 661 Tech., 13(2), 373–404, doi:10.5194/amt-13-373-2020, 2020.

662 Sekiyama, T. T., Tanaka, T. Y., Shimizu, A. and Miyoshi, T.: Data assimilation of CALIPSO
 663 aerosol observations, Atmos. Chem. Phys., 10(1), 39–49, doi:10.5194/acp-10-39-2010, 2010.

664 Shi, Y., Zhang, J., Reid, J. S., Holben, B., Hyer, E. J. and Curtis, C.: An analysis of the collection
 665 5 MODIS over-ocean aerosol optical depth product for its implication in aerosol assimilation,
 666 Atmos. Chem. Phys., 11(2), 557–565, doi:10.5194/acp-11-557-2011, 2011.

667 Shi, Y., Zhang, J., Reid, J. S., Hyer, E. J. and Hsu, N. C.: Critical evaluation of the MODIS Deep
 668 Blue aerosol optical depth product for data assimilation over North Africa, Atmos. Meas. Tech.,
 669 6(4), 949–969, doi:10.5194/amt-6-949-2013, 2013.

670 Shi, Y., Zhang, J., Reid, J. S., Liu, B. and Hyer, E. J.: Critical evaluation of cloud contamination
 671 in the MISR aerosol products using MODIS cloud mask products, Atmos. Meas. Tech., 7(6),
 672 1791–1801, doi:10.5194/amt-7-1791-2014, 2014.

673 Shindell, D. T., Lamarque, J. F., Schulz, M., Flanner, M., Jiao, C., Chin, M., Young, P. J., Lee,

674 Y. H., Rotstayn, L., Mahowald, N., Milly, G., Faluvegi, G., Balkanski, Y., Collins, W. J.,
 675 Conley, A. J., Dalsoren, S., Easter, R., Ghan, S., Horowitz, L., Liu, X., Myhre, G., Nagashima,
 676 T., Naik, V., Rumbold, S. T., Skeie, R., Sudo, K., Szopa, S., Takemura, T., Voulgarakis, A.,
 677 Yoon, J. H. and Lo, F.: Radiative forcing in the ACCMIP historical and future climate
 678 simulations, *Atmos. Chem. Phys.*, 13(6), 2939–2974, doi:10.5194/acp-13-2939-2013, 2013.

679 Si, Y., Chen, L., Xiong, X., Shi, S., Husi, L. and Cai, K.: Evaluation of the MISR fine resolution
 680 aerosol product using MODIS, MISR, and ground observations over China, *Atmos. Environ.*,
 681 223(December 2019), 117229, doi:10.1016/j.atmosenv.2019.117229, 2020.

682 Sogacheva, L., Popp, T., Sayer, A. M., Dubovik, O., Garay, M. J., Heckel, A., Hsu, N. C.,
 683 Jethva, H., Kahn, R. A., Kolmonen, P., Kosmale, M., de Leeuw, G., Levy, R. C., Litvinov, P.,
 684 Lyapustin, A., North, P., Torres, O. and Arola, A.: Merging regional and global aerosol optical
 685 depth records from major available satellite products, *Atmos. Chem. Phys.*, 20(4), 2031–2056,
 686 doi:10.5194/acp-20-2031-2020, 2020.

687 Tao, M., Wang, J., Li, R., Chen, L., Xu, X., Wang, L., Tao, J., Wang, Z. and Xiang, J.:
 688 Characterization of Aerosol Type Over East Asia by 4.4 km MISR Product: First Insight and
 689 General Performance, *J. Geophys. Res. Atmos.*, 125(13), 1–16, doi:10.1029/2019JD031909,
 690 2020.

691 Di Tomaso, E., Schutgens, N. A. J., Jorba, O. and García-Pando, C. P.: Assimilation of MODIS
 692 Dark Target and Deep Blue observations in the dust aerosol component of NMMB-MONARCH
 693 version 1.0, *Geosci. Model Dev.*, 10(3), 1107–1129, doi:10.5194/gmd-10-1107-2017, 2017.

694 Turnock, S. T., Allen, R. J., Andrews, M., Bauer, S. E., Deushi, M., Emmons, L., Good, P.,
 695 Horowitz, L., John, J. G., Michou, M., Nabat, P., Naik, V., Neubauer, D., O'Connor, F. M.,
 696 Olivié, D., Oshima, N., Schulz, M., Sellar, A., Shim, S., Takemura, T., Tilmes, S., Tsigaridis, K.,
 697 Wu, T. and Zhang, J.: Historical and future changes in air pollutants from CMIP6 models,
 698 *Atmos. Chem. Phys.*, 20(23), 14547–14579, doi:10.5194/acp-20-14547-2020, 2020.

699 Werner, M., Kryza, M. and Guzikowski, J.: Can data assimilation of surface PM_{2.5} and Satellite
 700 AOD improve WRF-Chem Forecasting? A case study for two scenarios of particulate air
 701 pollution episodes in Poland, *Remote Sens.*, 11(20), doi:10.3390/rs11202364, 2019.

702 Witek, M. L., Garay, M. J., Diner, D. J. and Smirnov, A.: Aerosol optical depths over oceans: A
 703 view from MISR retrievals and collocated MAN and AERONET in situ observations, *J.*
 704 *Geophys. Res. Atmos.*, 118(22), 12620–12633, doi:10.1002/2013JD020393, 2013.

705 Witek, M. L., Diner, D. J., Garay, M. J., Xu, F., Bull, M. A. and Seidel, F. C.: Improving MISR
 706 AOD Retrievals with Low-Light-Level Corrections for Veiling Light, *IEEE Trans. Geosci.*
 707 *Remote Sens.*, 56(3), 1251–1268, doi:10.1109/TGRS.2017.2727342, 2018a.

708 Witek, M. L., Garay, M. J., Diner, D. J., Bull, M. A. and Seidel, F. C.: New approach to the
 709 retrieval of AOD and its uncertainty from MISR observations over dark water, *Atmos. Meas.*
 710 *Tech.*, 11(1), 429–439, doi:10.5194/amt-11-429-2018, 2018b.

711 Witek, M. L., Garay, M. J., Diner, D. J. and Smirnov, A.: Oceanic Aerosol Loading Derived
 712 From MISR’s 4.4 km (V23) Aerosol Product, *J. Geophys. Res. Atmos.*, 124(17–18), 10154–
 713 10174, doi:10.1029/2019JD031065, 2019.

714 Xian, P., Reid, J. S., Hyer, E. J., Sampson, C. R., Rubin, J. I., Ades, M., Asencio, N., Basart, S.,
 715 Benedetti, A., Bhattacharjee, P. S., Brooks, M. E., Colarco, P. R., da Silva, A. M., Eck, T. F.,
 716 Guth, J., Jorba, O., Kouznetsov, R., Kipling, Z., Sofiev, M., Perez Garcia-Pando, C., Pradhan,
 717 Y., Tanaka, T., Wang, J., Westphal, D. L., Yumimoto, K. and Zhang, J.: Current state of the
 718 global operational aerosol multi-model ensemble: An update from the International Cooperative
 719 for Aerosol Prediction (ICAP), *Q. J. R. Meteorol. Soc.*, 145, 176–209, doi:10.1002/qj.3497,
 720 2019.

721 Zhang, J. and Reid, J. S.: An analysis of clear sky and contextual biases using an operational
 722 over ocean MODIS aerosol product, *Geophys. Res. Lett.*, 36(15), doi:10.1029/2009GL038723,
 723 2009.

724 Zhang, J. and Reid, J. S.: A decadal regional and global trend analysis of the aerosol optical
 725 depth using a data-assimilation grade over-water MODIS and Level 2 MISR aerosol products,
 726 *Atmos. Chem. Phys.*, 10(22), 10949–10963, doi:10.5194/acp-10-10949-2010, 2010.

727 Zhang, J., Reid, J. S., Westphal, D. L., Baker, N. L. and Hyer, E. J.: A system for operational
 728 aerosol optical depth data assimilation over global oceans, *J. Geophys. Res. Atmos.*, 113(10), 1–

729 13, doi:10.1029/2007JD009065, 2008.

730