



Comparison of Two Photolytic Calibration Methods for Nitrous Acid

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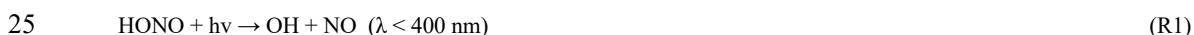
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Abstract. Nitrous acid (HONO) plays an important role in tropospheric oxidation chemistry as it is a precursor to the hydroxyl radical. Measurements of HONO have been historically difficult due to instrument interferences and difficulties in sampling and calibration. The traditional calibration method involves generation of HONO by reacting hydrogen chloride vapor with sodium nitrite followed by quantification by various methods (e.g., conversion of HONO to nitric oxide (NO) followed by chemiluminescence detection). Alternatively, HONO can be generated photolytically in the gas-phase by reacting NO with OH radicals generated by H₂O photolysis. In this work, we describe and compare two photolytic HONO calibration methods that were used to calibrate an iodide adduct chemical ionization mass spectrometer (CIMS). Both methods are based on the water vapor photolysis method commonly used for OH and HO₂ calibrations. The first method is an adaptation of the common chemical actinometry HO_x calibration method, in which HONO is calculated based on quantified values for [O₃], [H₂O], [O₂], and the absorption cross sections for H₂O and O₂ at 184.9 nm. In the second, novel method the HONO concentration is simply determined based on the simultaneous measurements of NO₂ formed by the reaction of NO with HO₂ from the H₂O photolysis. This second, novel approach generally has an improved (lower) calibration uncertainty and is simpler to apply. Calibration uncertainties are typically 30 to 36% (2σ) for the actinometric method and as low as 9% (2σ) for the NO₂ proxy method, limited by the uncertainty of the NO₂ measurements.

20 1 Introduction

Nitrous acid (HONO) is a source of the most important atmospheric oxidant – the hydroxyl radical (OH) - and can therefore play an important role in tropospheric oxidation chemistry. The hydroxyl radical causes the removal of most trace gases from the atmosphere while initiating the formation of secondary pollutants such as ozone (O₃) and secondary organic aerosol (SOA). Photolysis of HONO yields OH and nitric oxide (NO):



This reaction is the primary sink of HONO during the daytime leading to a typical chemical lifetime at mid-day of between 10-20 minutes at mid-latitudes. Sources of HONO include homogeneous formation (R2), direct emissions from combustion (vehicles, biomass burning, etc.) and soils, and numerous heterogeneous processes including heterogeneous reaction of NO₂ with moist terrestrial surfaces, photolysis of particulate nitrate (Ye et al., 2016; Ye et al., 2017), and photolysis of nitric acid (Ye et al., 2016).



The relative importance of these sources varies with environment (Jiang et al., 2022).

HONO photolysis has been reported as a major source of HO_x (HO_x = OH + HO₂) throughout the day in a variety of environments, including urban and highly polluted areas (Whalley et al., 2018; Slater et al., 2020; Ren et al., 2013; Lu et al., 2019)



as well as more pristine environments (Villena et al., 2011; Jiang et al., 2020; Bloss et al., 2007). Vertical distributions of HONO, however, indicate that its significance as a HOx precursor may be limited to near ground level (Li et al., 2014; Young et al., 2012; Villena et al., 2011; Wong et al., 2012; Tuite et al., 2021; Jaeglé et al., 2018).

HONO is notoriously difficult to measure. It can be formed via heterogeneous chemistry within sampling lines or an instrument's inlet. The resulting interferences may pose additional challenges in applying an instrument's zero or in calibration processes that alter the interfering species. Some intercomparison studies have shown substantial differences between HONO measurement techniques. For a Beijing, China based study, a comparison of several HONO measurements showed an overall mixed agreement with major differences observed for a few techniques (Crilley et al., 2019). Measurements in Houston, Texas showed several instruments to mostly agree in capturing variations in HONO, though there were differences in the magnitude of presented [HONO] values (Pinto et al., 2014). Bourgeois et al. (in review, 2022) recently reported an 80% difference between HONO measurements made by cavity-enhanced spectroscopy and iodide-adduct chemical ionization mass spectrometry (CIMS). Closer agreement for two instrument HONO comparisons has been reported by Stutz et al. (2010) (comparing differential optical absorption spectroscopy (DOAS) and mist-chamber ion chromatography (IC)), Cheng et al. (2013) (comparing Long-path absorption photometry (LOPAP) and stripping coil IC), and Dixneuf et al. (2022) (comparing LOPAP and cavity-enhanced absorption spectroscopy), though many of these studies report considerable deviations when HONO mixing ratios were less than ~100 pptv.

Calibrations for HONO are challenging as it is not commercially available and rather must be prepared in situ. Most commonly, HONO is prepared by reacting hydrogen chloride vapor with sodium nitrite (Febo et al., 1995):



This method presents several challenges. A stable source of HCl is required, usually from a heated aqueous solution, a gas cylinder, or a permeation tube. Consistent mixing between the HCl and the NaNO₂ powder is required. This method typically produces high HONO concentrations (above 1 ppmv), requiring dilution, though the temporary unrealistic HONO concentrations can lead to significant HONO loss by its self-reaction and inaccurate HONO quantification. The generated HONO can be quantified by various methods including conversion to NO followed by chemiluminescence detection (Lee et al., 2012; Lao et al., 2020; Villena and Kleffmann, 2022), thermal conversion to NO₂ followed by NO₂ quantification (Gingerysty and Osthoff, 2020), and conversion to aqueous nitrite followed by derivatization and detection by UV-vis (Peng et al., 2020).

More recently, photolytic HONO sources have been utilized. Humidified air is exposed to ultraviolet (UV) light to photolyze H₂O to produce an equal mixture of OH and HO₂, which in the presence of excess NO then converts to HONO. This HONO output is tunable by altering humidity, UV flux, or UV exposure time. The HONO formed has been quantified by inferring the UV flux by measuring the O₃ produced from O₂ photolysis (Dyson et al., 2021; Bottorff et al., 2021) and by thermal dissociation of the HONO followed by measurement of the NO₂ produced (Veres et al., 2015). These methods have an uncertainty of 30 to 36% (2σ), similar to the uncertainty for HOx calibrations based on water vapor photolysis (Dusanter et al., 2008). In this manuscript, we present an alternative photolytic HONO calibration that we refer to as the “NO₂ proxy” method. This method requires a direct NO₂ measurement that is used as a ‘proxy’ to quantify HONO concentrations. We compare this new proxy calibration to the more standard photolytic calibration method as performed by Bottorff et al. (2021) and Dyson et al. (2021). This method has a lower uncertainty (typically ~10%, 2σ) and unlike the actinometric method does not require characterization of the mercury lamp emission spectrum.

2 Methods



2.1 Instrumentation

A Cavity Attenuated Phase Shift (CAPS) spectrometer was used to detect NO₂ (Kebabian et al., 2008). The CAPS also indirectly measured O₃ as it was converted to NO₂ by reaction with excess NO. The CAPS instrument was calibrated using a 2B Technologies Model 306 O₃ Calibration Source, which agreed to within 2.5% with a Thermo Environmental Instruments 49C O₃ Calibrator. We assign an uncertainty of 3% (2σ) to the NO₂ measurements. Temperature and relative humidity (RH) measurements were made using two Vaisala HMP60 probes.

A High-Resolution Chemical Ionization Time of Flight Mass Spectrometer (HR-ToF-CIMS, ToFwerks/Aerodyne Research, Inc.) was used to detect HONO (Bertram et al., 2011; Lee et al., 2014). HONO and concomitant gases are ionized within a laboratory-built ion-molecule reactor (IMR) using reagent iodide (I⁻) ions. The IMR is internally coated with PTFE and sampled the calibration gas at a flow rate of 2.10 SLPM through a stainless-steel critical orifice (nominal diameter of 0.48 mm). Iodide (I⁻) reagent ions in N₂ were sampled at 2.20 SLPM through a similar critical orifice perpendicular to the main sample flow. The I⁻ was prepared by exposing dilute methyl iodide (CH₃I) from a permeation tube to a ²¹⁰Po radioactive source. Humidified N₂ was also added to the IMR perpendicular to the main sample flow at a flow rate of 0.365 SLPM. The pressure in the IMR was held at 80 mbar, controlled by adjusting a valve to the scroll pump (Agilent Technologies IDP-7).

Ions are separated by mass-to-charge ratio (m/z) at a mass resolving power of near 5000 m/Δm. We monitor the HONO iodide adduct I(HONO)⁻ at 173.90575 m/z. The peak-fitting software (ToFWare) accounts for the overlapping contribution from the ¹³C formic acid I(¹³CH₂O₂)⁻ peak at 173.91342 m/z. We account for the humidity dependence of the instrumental response by determining the mole fraction of H₂O(g) (χ_{H2O}) in the IMR by measuring the RH and temperature of the IMR in the exhaust of the scroll pump.

2.2 Calibration Methods

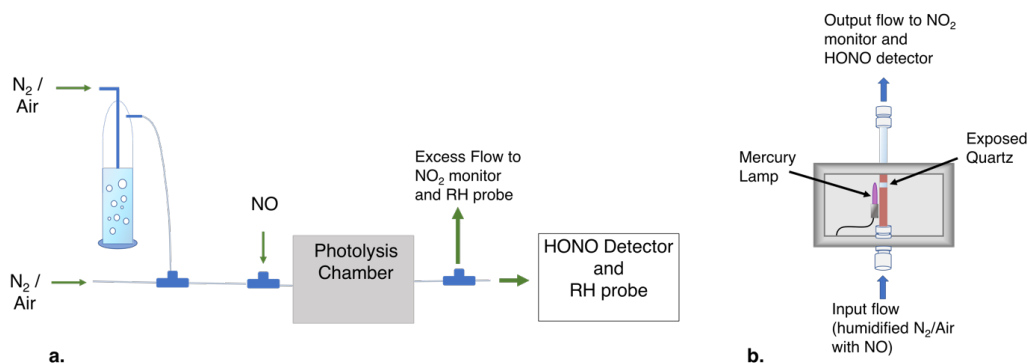


Figure 1: Schematic of the experimental setup (a.) and the photolysis chamber (b.). Not shown: small flow of air when using N₂ as the carrier gas (a.), and the purge flow of N₂ (b.).

We calibrate HONO using two variations of the water vapor HO_x calibration method: one is a modification of the standard actinometric HO_x photolytic calibration and the other we refer to as the “NO₂ proxy” calibration. These calibration methods mainly differ in how HONO is quantified. In both methods, HONO was produced nearly identically. Humidified air (for the actinometric method) or N₂ (for the NO₂ proxy calibration) is mixed with NO and then exposed to 184.9 nm ultraviolet radiation from a low-pressure mercury lamp (Jelight 78-2046-1). The resulting OH and HO₂ from water photolysis form HONO by reaction with excess NO (R4-6a and R2).



105 The HO₂ to HONO pathway is limited by the following reaction:



A schematic of the setup used for both calibrations is shown in Figure 1. The mercury lamp is housed within a 10.8 cm × 26.7 cm × 10.2 cm photolysis chamber (Fig. 1b), and the volume surrounding the lamp is purged with dry N₂ (purge not shown). The humidified air-NO mixture is transported past the mercury lamp within a partially exposed 26.7 cm quartz tube (I.D. = 1.04 cm). HONO sample concentrations are controlled by adjusting the lamp flux with a Variac Variable transformer, adjusting the relative flow rates of the dry and humidified ZA/N₂, or adjusting the absolute flow rates to alter the lamp exposure time. For our example calibrations discussed in this manuscript, we typically used a main N₂ or air flow rate of 5 SLPM with an addition of 200 sccm of 41.02 ppmv NO in N₂ (Airgas) for a total flow rate of 5.20 SLPM and a diluted NO mixing ratio of 1.58 ppmv. Pseudo-first order rate constants calculated using this [NO] for R6a and R2 (the reactions that form HONO) are 322 s⁻¹ and 295 s⁻¹, respectively. Under these conditions, HO_x is converted to HONO within 0.02 s inside the remaining 11.4 cm of the quartz tube. The [NO] chosen must be high enough to minimize OH and HO₂ wall losses. We have ensured that this NO mixing ratio is sufficient in separate experiments by confirming that no additional HONO signal results at increased [NO] values.

115 The resulting calibration gas was sampled by the CIMS (2.1 SLPM) and the NO₂ monitor (1.1 SLPM). The excess flow was vented past an RH/T probe to determine the water mixing ratio in the photolysis cell. A second RH/T probe quantified the water mixing ratio in the CIMS IMR as previously mentioned in Sect. 2.1. Details of the two calibration methods are described in the following sections.

2.2.1 Actinometric Calibration

The water vapor photolysis calibration method has been used for several decades to calibrate OH and HO₂ measurements (Stevens et al., 1994; Lanzendorf et al., 1997; Dusanter et al., 2008). The concentration of HO_x, and therefore HONO, is calculated from the time-integrated photolysis of water vapor:

$$[\text{HONO}] \approx [\text{HO}_x] = (F \cdot t)[\text{H}_2\text{O}]\sigma_{\text{H}_2\text{O}}\Phi_{\text{HO}_x} \quad (1)$$

Where F is the photon flux at 184.9 nm, t is the UV irradiation time, $\sigma_{\text{H}_2\text{O}}$ is the absorption cross section of water at 184.9 nm, and Φ_{HO_x} is the quantum yield of HO_x from water photolysis and equal to 2. F can be quantified using direct actinometric measurements (e.g., using a calibrated phototube), and t can be quantified via characterization of the flow rates and photolysis cell geometry (Faloona et al., 2004). Alternatively, and more commonly among HO_x measurement groups, the product $F \cdot t$ can be determined via “chemical actinometry” (Schultz et al., 1995). In the O₂-O₃ chemical actinometry method, the concentration of O₃ produced by photodissociation of O₂ at 184.9 is used to determine $F \cdot t$:





135 The product of the lamp flux and the exposure time, i.e., the ($F \cdot t$) term, is given by Eq. (2), in which σ_{O_2} is the absorption cross section of O_2 at 184.9 nm and Φ_{O_3} is quantum yield of O_3 from O_2 photolysis ($\Phi_{O_2}=2$):

$$(F \cdot t) = \frac{[O_3]}{[O_2]\sigma_{O_2}\Phi_{O_3}} \quad (2)$$

Substituting this expression for $F \cdot t$ into Eq. (1) gives Eq. (3):

$$[HONO] \approx [HOx] = \frac{[O_3]}{[O_2]\sigma_{O_2}} [H_2O]\sigma_{H_2O} \quad (3)$$

140 The effective value for σ_{O_2} must be experimentally determined for the individual mercury lamp at the experimental O_2 optical depth to account for the non-ideal overlap between the lamp emission spectrum and the O_2 absorption spectrum (Lanzendorf et al., 1997). We use a value of $1.4 \times 10^{-20} \text{ cm}^2 \text{ molec}^{-1}$ for σ_{O_2} for the mercury lamp used for these experiments. We use the JPL-recommended value of $7.1 \times 10^{-20} \text{ cm}^2 \text{ molec}^{-1}$ for σ_{H_2O} (Burkholder et al., 2020).

We apply a small correction to the $[HOx]$ calculated in Eq. (3) in order to obtain $[HONO]$. This correction accounts for the
145 incomplete conversion of HO_2 to HONO due to R6b.

$$[HONO] = (0.5 + 0.5 \cdot (1 - \frac{\beta}{1+\beta})) \cdot [HOx] \quad (4)$$

Equation 4 includes the variable β , which is the relative rate or product ratio of R6b to R6a (i.e., $\beta = k_{R6b}/k_{R6a} = [HNO_3]/[NO_2]$) and depends on temperature, pressure, and humidity (Butkovskaya et al., 2007; Butkovskaya et al., 2009). The Eq. (4) fraction that includes β represents a traditional branching ratio (i.e., $k_{R6b}/(k_{R6a} + k_{R6b}) = [HNO_3]/([NO_2] + [HNO_3])$). Values of β are at most
150 0.04, leading to a 2% correction to equation 1. See Sect. S1 of the supplement for information regarding the formulation of Eq. (4) and the calculation of β .

The O_3 concentration within Eq. (3) is determined with the CAPS NO_2 monitor after its reaction with NO, forming NO_2 . This is measured with dry air flowing in the photolysis chamber so that the NO_2 measured is solely from the reaction of NO with O_3 and not HO_2 . $[H_2O]$ is determined using the measured RH, temperature, and pressure. The CIMS response to HONO is determined
155 by acquiring a background by briefly toggling off the mercury lamp. This background CIMS signal is humidity dependent, so a background is taken at each humidity setting. The NO flow contributes to the CIMS background HONO signal but is constant over time.

2.2.2 Proxy Calibration

For the NO_2 proxy calibration method, we determine $[HONO]$ from the measured value of $[NO_2]$ formed from R6a during
160 HONO production. For each H_2O molecule photolyzed (R4), nearly one NO_2 and two HONO molecules are produced. Therefore, $[HONO]$ is simply given by the measured $[NO_2]$ (Eq. (5)):

$$[HONO] = (2 + \beta) \cdot [NO_2] \quad (5)$$

Where β is added to account for the minor HNO_3 product of the $HO_2 + NO$ reaction (R6b).

For the proxy calibration we use humidified N_2 rather than air and include a small addition of 10 sccm of zero air prior to lamp
165 exposure (not shown in the schematic). The resulting low O_2 concentration is sufficient for the full conversion of H to HO_2 (R5) but results in a negligible amount of O_3 formed by O_2 photolysis (R7-R8), confirmed by toggling the UV source on and off with



dry N₂ flowing. HONO concentrations are quantified using background subtracted [NO₂] values in Eq. (5), which are typically acquired by toggling the mercury lamp off and on. The CIMS signal response is determined simultaneously.

3 Results and Discussion

170 Time series data for the proxy calibration method is shown in Fig. 2. The CAPS NO₂ measurement and the CIMS HONO
signal normalized per one million counts of reagent ion per second ('ncps') are shown at a constant RH setting (40% within the
H₂O photolysis cell). During the first 95 s, HONO is formed by H₂O photolysis via the mercury lamp 184.9 nm emission. This
leads to the stable I(HONO)⁻ signal in the CIMS along with enhanced [NO₂] produced by R6a and measured by the CAPS monitor.
Background I(HONO)⁻ and NO₂ signals are determined by toggling off the mercury lamp (shown at 96 s). The CIMS sensitivity
175 (ncps ppt⁻¹) is equal to the quotient of the normalized background subtracted CIMS signal and the quantified [HONO] which is
calculated by Eq. (5).

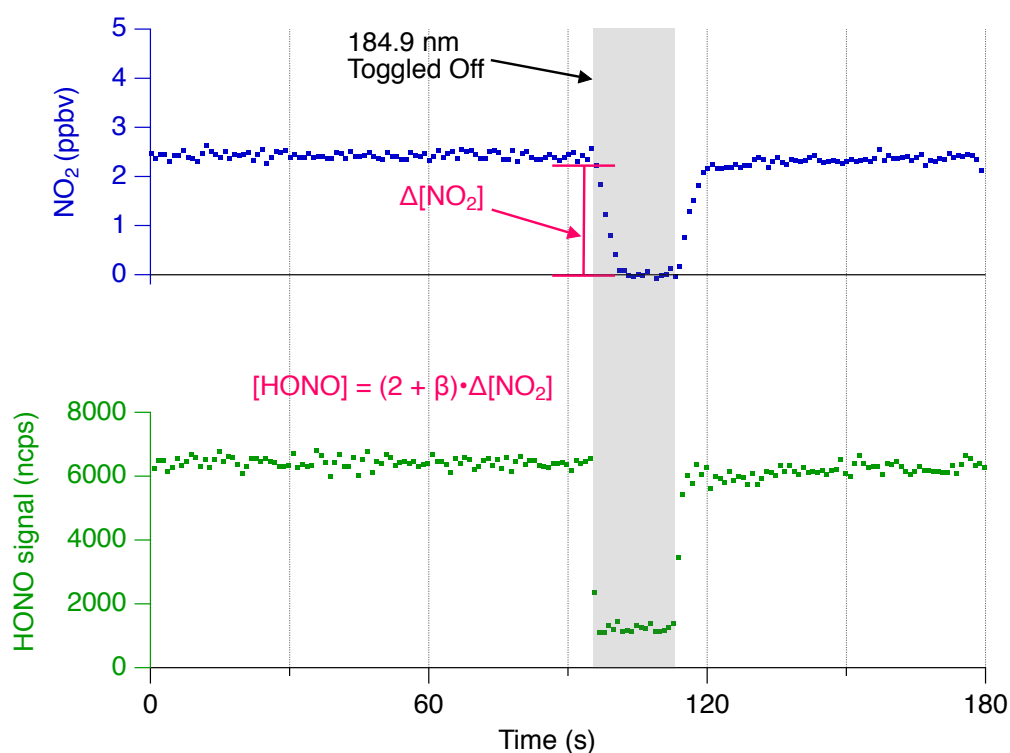
A multipoint NO₂ proxy calibration curve (Fig. 3) shows the linear CIMS signal response to [HONO]. This calibration was
conducted at a constant relative humidity (28% within the photolysis cell and 18% within the CIMS IMR), and [HONO] was
adjusted by altering the mercury lamp flux with a Variac variable transformer. The slope of this curve, 2.89 ± 0.34 ncps ppt⁻¹ (2σ),
180 is the CIMS sensitivity to HONO for this particular mole fraction of water within the CIMS IMR ($\chi_{\text{H}_2\text{O}} = 3.88 \times 10^{-3}$). HONO
mixing ratios ranged from approximately 400 pptv to 3500 pptv, thus demonstrating that a wide range in HONO concentrations
can easily be prepared. The uncertainty for the quantified [HONO] from Eq. (5) (i.e., the x-error bars) is obtained by adding in
quadrature three terms: 1.) the relative uncertainty of the NO₂ background subtraction (based on the 5 s average precision of 27
pptv), 2.) the NO₂ calibration uncertainty (3%, 2σ), and finally 3.) the relative error associated with the $(2 + \beta)$ expression (typically
185 0.14%, 2σ). The uncertainty in the normalized CIMS signal is obtained by adding in quadrature the precision of the I(HONO)⁻
signal with that of the reagent ion. The 2σ error bars range from 5.6 to 34.8% for quantified [HONO] and 6.8 to 11.2% in the
CIMS HONO signal. The Fig. 3 slope (i.e., the CIMS sensitivity to HONO) and its uncertainty was determined using the York
bivariate linear regression method (York et al., 2004). The uncertainty calculations are discussed in greater detail in Sect. S2.1 of
the supplement.

190 A comparison between the more standard O₃ actinometry based calibration and the new proxy calibration method is shown in
Fig. 4. CIMS sensitivities as determined by single point calibrations are shown for a variety of $\chi_{\text{H}_2\text{O}}$ values. The two calibration
methods were conducted consecutively and agree within their provided 2σ errors. Sensitivities ranged from 1.5 to 5 ncps ppt⁻¹ with
the greatest values observed at low $\chi_{\text{H}_2\text{O}}$ settings. The sensitivity determined by the Fig. 3 multipoint calibration (2.89 ± 0.34 ncps
ppt⁻¹ at $\chi_{\text{H}_2\text{O}} = 3.88 \times 10^{-3}$) is consistent with those shown in Fig. 4 at similar $\chi_{\text{H}_2\text{O}}$ values. These CIMS sensitivities are also in line
195 with literature values (Peng et al., 2020; Bourgeois et al., in review, 2022). Unique to this figure is the use of the CIMS IMR $\chi_{\text{H}_2\text{O}}$
to track humidity dependence rather than the partial pressure of H₂O (Lee et al., 2014) or the CIMS signal ratio of the iodide water
adduct I(H₂O)⁻ (m/z 145) to reagent ion I⁻ (m/z 127) (Lee et al., 2014; Peng et al., 2020; Veres et al., 2015; Veres et al., 2020). The
use of $\chi_{\text{H}_2\text{O}}$ allows for a more direct comparison to other CIMS users that may use different IMR pressures and quadrupole voltage
settings that govern the I(H₂O)⁻ to I⁻ ratio.

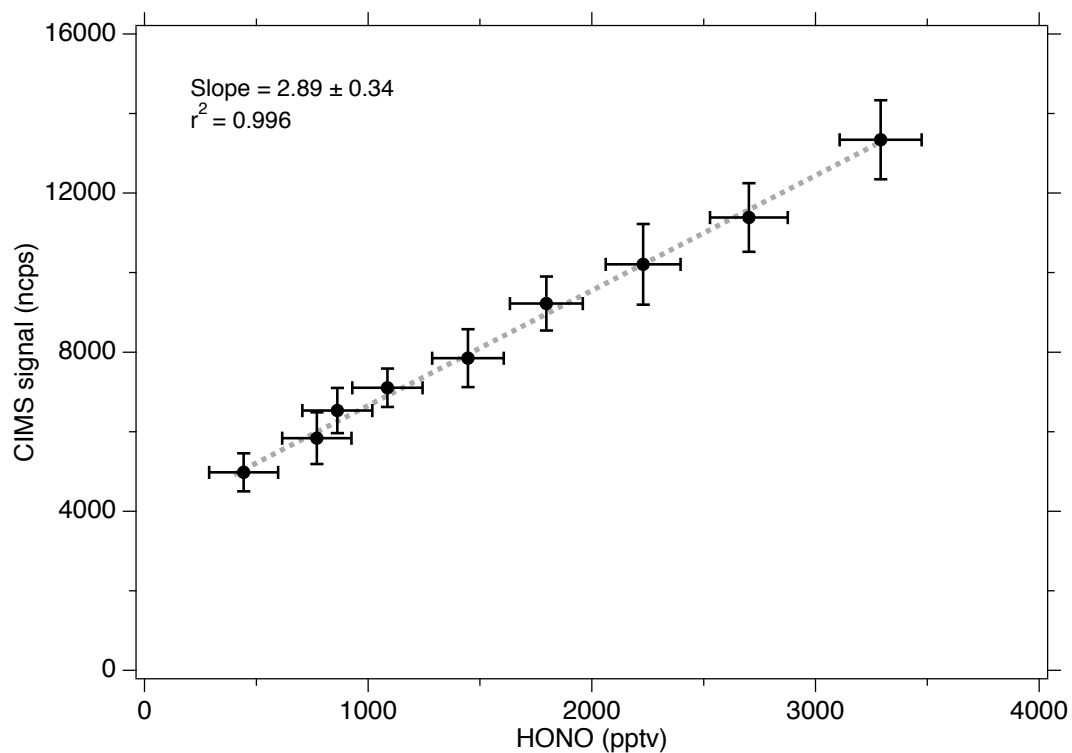
200 The uncertainty in CIMS sensitivity for the proxy method is determined by combining in quadrature the relative error of the
background subtracted CIMS signal with that of the quantified [HONO] value. The calculation for quantified [HONO] uncertainty
and CIMS uncertainty is previously mentioned in the text regarding Fig. 3, though the CIMS uncertainty here slightly differs due
to the background subtraction (see Sect. S2.2 the supplement). As mentioned previously, the total uncertainty is dependent on the



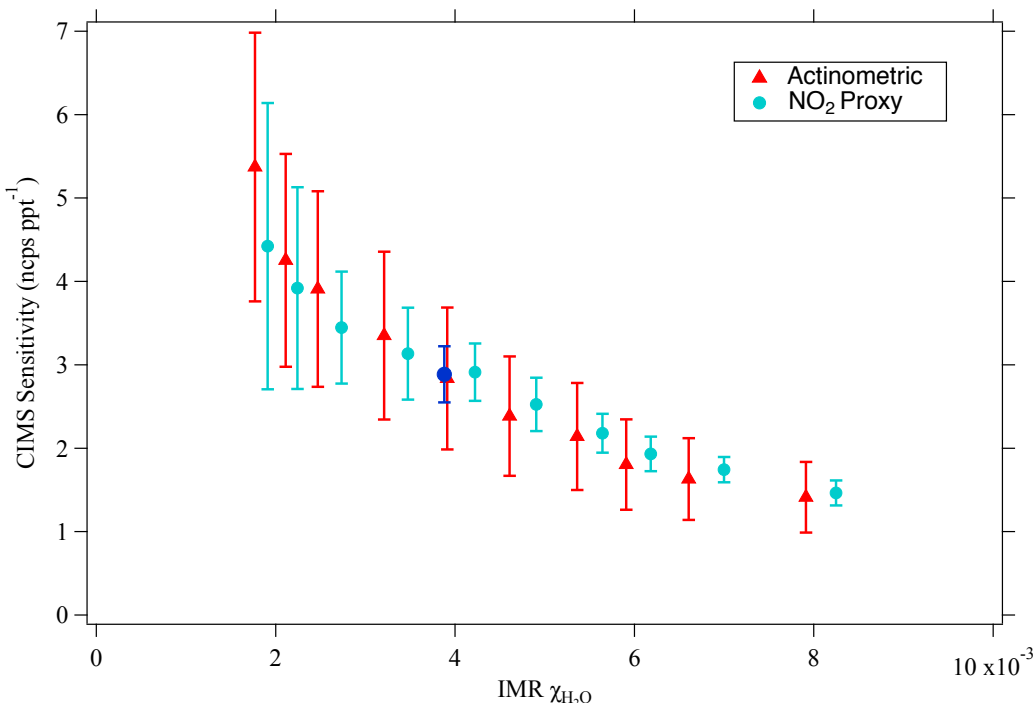
205 quantified [HONO] value, which ranged from 630 (at the lowest $\chi_{\text{H}_2\text{O}}$ value) to 7,800 pptv (at the greatest $\chi_{\text{H}_2\text{O}}$ value) for the shown
Fig. 4 proxy calibration results. The uncertainty at the lowest $\chi_{\text{H}_2\text{O}}$ settings, which were also the lowest HONO concentrations,
were dominated by [HONO] quantification stemming from the background subtraction of NO_2 . Therefore, the uncertainty can be
minimized by using higher HONO concentrations. The presented 2σ uncertainty in using the proxy method ranged from 8.7 to
38.8%. All presented proxy calibration uncertainties besides the two at lowest $\chi_{\text{H}_2\text{O}}$ values (i.e., the smallest [HONO] values) fall
well below the 30% 2σ uncertainty associated with the standard O_3 actinometry calibration. The proxy calibration is therefore a
210 general improvement over the standard actinometry calibration, especially at high [HONO] values.



215 **Figure 2: One-second averaged time series data for a proxy calibration at a constant relative humidity. The iodide HONO adduct signal (cps) is shown normalized per one million reagent ions (ncps). The highlighted section represents the period in which the 184.9 nm mercury lamp is toggled off to obtain background [NO₂] and HONO signal.**



220 **Figure 3:** Calibration curve obtained using the NO₂ proxy calibration method at a constant humidity (RH = 17.7%, [H₂O] = 0.388% as measured in the CIMS scroll pump exhaust). Error bars represent $\pm 2 \sigma$.



225 **Figure 4:** Comparison of the two HONO calibration methods for a range of $\chi_{\text{H}_2\text{O}}$ within the CIMS IMR. The sensitivity determined by multipoint calibration (i.e., the Fig. 3 slope) is plotted as the dark blue circle. Error bars represent $\pm 2\sigma$.

4 Conclusions

230 Two photolytic HONO calibration methods based on reacting NO with the HOx generated by H₂O photolysis at 184.9 nm were presented. This includes a novel approach in which HONO is quantified using the NO₂ formed by the HO₂ + NO reaction as a proxy. The proxy method compares well with the O₃ actinometry based calibration while also having the benefit of a simpler calculation that avoids the need to characterize the emission spectrum of the mercury lamp used. In addition, this proxy method generally has improved uncertainties, typically between 8 and 20% (2 σ) – lower than the 30 to 36% 2 σ uncertainty associated with the actinometric calibration method. These photolytic calibrations require a direct NO₂ measurement (e.g., CAPS or CRDS), a 184.9 nm light source, and a simple quartz tube photolysis chamber. While the proxy calibration method was conducted in N₂ for this manuscript, it is possible to instead perform this method using air in the case that N₂ is incompatible with an instrument (unlike our CIMS). In an air-based proxy calibration, the [O₃] produced by O₂ photolysis would need to be quantified (during dry conditions) and then subtracted from subsequent background subtracted NO₂ signals. In conclusion, these photolytic calibration techniques offer a valuable alternative to the more conventional HONO calibration that is based on reacting hydrogen chloride vapor with sodium nitrite.

240 *Data availability.* The data used in this manuscript is available upon request.



Author contributions. AJL and ECW designed the experiments. AJL carried them out and performed the data analysis. AJL prepared the manuscript with contributions from ECW.

Competing Interests. The authors declare that they have no conflict of interest.

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References

- Bertram, T., Kimmel, J., Crisp, T., Ryder, O., Yatavelli, R., Thornton, J., Cubison, M., Gonin, M., and Worsnop, D.: A field-deployable, chemical ionization time-of-flight mass spectrometer, *Atmospheric Measurement Techniques*, 4, 1471-1479, <https://doi.org/10.5194/amt-4-1471-2011>, 2011.
- 250 Bloss, W., Lee, J., Heard, D., Salmon, R. A., Bauguutte, S.-B., Roscoe, H., and Jones, A.: Observations of OH and HO₂ radicals in coastal Antarctica, *Atmospheric Chemistry and Physics*, 7, 4171-4185, <https://doi.org/10.5194/acp-7-4171-2007>, 2007.
- Bottorff, B., Reidy, E., Mielke, L., Dusanter, S., and Stevens, P. S.: Development of a laser-photofragmentation laser-induced fluorescence instrument for the detection of nitrous acid and hydroxyl radicals in the atmosphere, *Atmospheric Measurement*
- 255 *Techniques*, 14, 6039-6056, <https://doi.org/10.5194/amt-14-6039-2021>, 2021.
- Bourgeois, I., Peischl, J., Neuman, J. A., Brown, S. S., Allen, H. M., Campuzano-Jost, P., Coggon, M. M., DiGangi, J. P., Diskin, G. S., Gilman, J. B., Gkatzelis, G. I., Guo, H., Halliday, H., Hanisco, T. F., Holmes, C. D., Huey, L. G., Jimenez, J. L., Lamplugh, A. D., Lee, Y. R., Lindaas, J., Moore, R. H., Nowak, J. B., Pagonis, D., Rickly, P. S., Robinson, M. A., Rollins, A. W., Selimovic, V., St. Clair, J. M., Tanner, D., Vasquez, K. T., Veres, P. R., Warneke, C., Wennberg, P. O., Washenfelder, R.
- 260 A., Wiggins, E. B., Womack, C. C., Xu, L., Zarzana, K. J., and Ryerson, T. B.: Comparison of airborne measurements of NO, NO₂, HONO, NO_y and CO during FIREX-AQ, <https://doi.org/10.5194/amt-2021-432>, in review, 2022.
- Burkholder, J., Sander, S., Abbatt, J., Barker, J., Cappa, C., Crouse, J., Dibble, T., Huie, R., Kolb, C., and Kurylo, M.: Chemical kinetics and photochemical data for use in atmospheric studies, Evaluation No. 19, JPL Publication 19-10, Jet Propulsion Laboratory, Pasadena, 583-586, 2020.
- 265 Butkovskaya, N., Kukui, A., and Le Bras, G.: HNO₃ Forming Channel of the HO₂+ NO Reaction as a Function of Pressure and Temperature in the Ranges of 72– 600 Torr and 223– 323 K, *The Journal of Physical Chemistry A*, 111, 9047-9053, <https://doi.org/10.1021/jp074117m>, 2007.
- Butkovskaya, N., Rayez, M.-T., Rayez, J.-C., Kukui, A., and Le Bras, G.: Water vapor effect on the HNO₃ yield in the HO₂+ NO reaction: experimental and theoretical evidence, *The Journal of Physical Chemistry A*, 113, 11327-11342, <https://doi.org/10.1021/jp811428p>, 2009.
- 270 Cheng, P., Cheng, Y., Lu, K., Su, H., Yang, Q., Zou, Y., Zhao, Y., Dong, H., Zeng, L., and Zhang, Y.: An online monitoring system for atmospheric nitrous acid (HONO) based on stripping coil and ion chromatography, *Journal of Environmental Sciences*, 25, 895-907, [https://doi.org/10.1016/s1001-0742\(12\)60251-4](https://doi.org/10.1016/s1001-0742(12)60251-4), 2013.
- Crilley, L. R., Kramer, L. J., Ouyang, B., Duan, J., Zhang, W., Tong, S., Ge, M., Tang, K., Qin, M., Xie, P., Shaw, M. D., Lewis, A. C., Mehra, A., Bannan, T. J., Worrall, S. D., Priestley, M., Bacak, A., Coe, H., Allan, J., Percival, C. J., Popoola, O. A. M.,



- Jones, R. L., and Bloss, W. J.: Intercomparison of nitrous acid (HONO) measurement techniques in a megacity (Beijing), *Atmospheric Measurement Techniques*, 12, 6449-6463, <https://doi.org/10.5194/amt-12-6449-2019>, 2019.
- Dixneuf, S., Ruth, A. A., Häsel, R., Brauers, T., Rohrer, F., and Dorn, H.-P.: Detection of nitrous acid in the atmospheric simulation chamber SAPHIR using open-path incoherent broadband cavity-enhanced absorption spectroscopy and extractive
280 long-path absorption photometry, *Atmospheric Measurement Techniques*, 15, 945-964, <https://doi.org/10.5194/amt-15-945-2022>, 2022.
- Dusanter, S., Vimal, D., and Stevens, P.: Measuring tropospheric OH and HO₂ by laser-induced fluorescence at low pressure. A comparison of calibration techniques, *Atmospheric Chemistry and Physics*, 8, 321-340, <https://doi.org/10.5194/acp-8-321-2008>, 2008.
- 285 Dyson, J. E., Boustead, G. A., Fleming, L. T., Blitz, M., Stone, D., Arnold, S. R., Whalley, L. K., and Heard, D. E.: Production of HONO from NO₂ uptake on illuminated TiO₂ aerosol particles and following the illumination of mixed TiO₂/ammonium nitrate particles, *Atmospheric Chemistry and Physics*, 21, 5755-5775, <https://doi.org/10.5194/acp-21-5755-2021>, 2021.
- Faloona, I. C., Tan, D., Leshner, R. L., Hazen, N. L., Frame, C. L., Simpas, J. B., Harder, H., Martinez, M., Di Carlo, P., and Ren, X.: A laser-induced fluorescence instrument for detecting tropospheric OH and HO₂: Characteristics and calibration, *Journal of*
290 *Atmospheric Chemistry*, 47, 139-167, <https://doi.org/10.1023/B:JOCH.0000021036.53185.0e>, 2004.
- Febo, A., Perrino, C., Gherardi, M., and Sparapani, R.: Evaluation of a high-purity and high-stability continuous generation system for nitrous acid, *Environmental science & technology*, 29, 2390-2395, <https://doi.org/10.1021/es00009a035>, 1995.
- Gingerysty, N. J., and Osthoff, H. D.: A compact, high-purity source of HONO validated by Fourier transform infrared and thermal-dissociation cavity ring-down spectroscopy, *Atmospheric Measurement Techniques*, 13, 4159-4167,
295 <https://doi.org/10.5194/amt-13-4159-2020>, 2020.
- Jaeglé, L., Shah, V., Thornton, J., Lopez-Hilfiker, F., Lee, B., McDuffie, E., Fibiger, D., Brown, S., Veres, P., and Sparks, T.: Nitrogen oxides emissions, chemistry, deposition, and export over the Northeast United States during the WINTER aircraft campaign, *Journal of Geophysical Research: Atmospheres*, 123, 12,368-312,393, <https://doi.org/10.1029/2018JD029133>, 2018.
- Jiang, Y., Xue, L., Gu, R., Jia, M., Zhang, Y., Wen, L., Zheng, P., Chen, T., Li, H., and Shan, Y.: Sources of nitrous acid
300 (HONO) in the upper boundary layer and lower free troposphere of the North China Plain: insights from the Mount Tai Observatory, *Atmospheric Chemistry and Physics*, 20, 12115-12131, <https://doi.org/10.5194/acp-20-12115-2020>, 2020.
- Kebabian, P. L., Wood, E. C., Herndon, S. C., and Freedman, A.: A practical alternative to chemiluminescence-based detection of nitrogen dioxide: Cavity attenuated phase shift spectroscopy, *Environmental science & technology*, 42, 6040-6045,
<https://doi.org/10.1021/es703204j>, 2008.
- 305 Lanzendorf, E., Hanisco, T., Donahue, N., and Wennberg, P.: Comment on: "The measurement of tropospheric OH radicals by laser-induced fluorescence spectroscopy during the POPCORN Field Campaign" by Hofzumahaus et al. and "Intercomparison of tropospheric OH radical measurements by multiple folded long-path laser absorption and laser induced fluorescence" by Brauers et al, *Geophysical research letters*, 24, 3037-3038, <https://doi.org/10.1029/97GL02899>, 1997.
- Lao, M., Crilley, L. R., Salehpoor, L., Furlani, T. C., Bourgeois, I., Neuman, J. A., Rollins, A. W., Veres, P. R., Washenfelder, R. A., Womack, C. C., Young, C. J., and VandenBoer, T. C.: A portable, robust, stable, and tunable calibration source for gas-phase nitrous acid (HONO), *Atmospheric Measurement Techniques*, 13, 5873-5890, <https://doi.org/10.5194/amt-13-5873-2020>,
310 2020.



- Lee, B., Wood, E., Wormhoudt, J., Shorter, J., Herndon, S., Zahniser, M., and Munger, J.: Effective line strengths of trans-nitrous acid near 1275 cm⁻¹ and cis-nitrous acid at 1660 cm⁻¹, *Journal of Quantitative Spectroscopy and Radiative Transfer*, 113, 1905-1912, <https://doi.org/10.1016/j.jqsrt.2012.06.004>, 2012.
- 315 Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurtén, T., Worsnop, D. R., and Thornton, J. A.: An iodide-adduct high-resolution time-of-flight chemical-ionization mass spectrometer: Application to atmospheric inorganic and organic compounds, *Environmental science & technology*, 48, 6309-6317, <https://doi.org/10.1021/es500362a>, 2014.
- Li, X., Rohrer, F., Hofzumahaus, A., Brauers, T., Häseler, R., Bohn, B., Broch, S., Fuchs, H., Gomm, S., and Holland, F.:
320 Missing gas-phase source of HONO inferred from Zeppelin measurements in the troposphere, *Science*, 344, 292-296, <https://doi.org/10.1126/science.1248999>, 2014.
- Lu, K., Fuchs, H., Hofzumahaus, A., Tan, Z., Wang, H., Zhang, L., Schmitt, S. H., Rohrer, F., Bohn, B., and Broch, S.: Fast photochemistry in wintertime haze: consequences for pollution mitigation strategies, *Environmental science & technology*, 53, 10676-10684, <https://doi.org/10.1021/acs.est.9b02422>, 2019.
- 325 Peng, Q., Palm, B. B., Melander, K. E., Lee, B. H., Hall, S. R., Ullmann, K., Campos, T., Weinheimer, A. J., Apel, E. C., and Hornbrook, R. S.: HONO emissions from western US wildfires provide dominant radical source in fresh wildfire smoke, *Environmental science & technology*, 54, 5954-5963, <https://doi.org/10.1021/acs.est.0c00126>, 2020.
- Pinto, J. P., Dibb, J., Lee, B. H., Rappenglück, B., Wood, E. C., Levy, M., Zhang, R. Y., Lefer, B., Ren, X. R., Stutz, J., Tsai, C., Ackermann, L., Golovko, J., Herndon, S. C., Oakes, M., Meng, Q. Y., Munger, J. W., Zahniser, M., and Zheng, J.:
330 Intercomparison of field measurements of nitrous acid (HONO) during the SHARP campaign, *Journal of Geophysical Research: Atmospheres*, 119, 5583-5601, <https://doi.org/10.1002/2013jd020287>, 2014.
- Ren, X., Van Duin, D., Cazorla, M., Chen, S., Mao, J., Zhang, L., Brune, W. H., Flynn, J. H., Grossberg, N., and Lefer, B. L.: Atmospheric oxidation chemistry and ozone production: Results from SHARP 2009 in Houston, Texas, *Journal of Geophysical Research: Atmospheres*, 118, 5770-5780, <https://doi.org/10.1002/jgrd.50342>, 2013.
- 335 Schultz, M., Heitlinger, M., Mihelcic, D., and Volz-Thomas, A.: Calibration source for peroxy radicals with built-in actinometry using H₂O and O₂ photolysis at 185 nm, *Journal of Geophysical Research: Atmospheres*, 100, 18811-18816, <https://doi.org/10.1029/95JD01642>, 1995.
- Slater, E. J., Whalley, L. K., Woodward-Massey, R., Ye, C., Lee, J. D., Squires, F., Hopkins, J. R., Dunmore, R. E., Shaw, M., and Hamilton, J. F.: Elevated levels of OH observed in haze events during wintertime in central Beijing, *Atmospheric Chemistry and Physics*, 20, 14847-14871, <https://doi.org/10.5194/acp-20-14847-2020>, 2020.
- 340 Stevens, P., Mather, J., and Brune, W.: Measurement of tropospheric OH and HO₂ by laser-induced fluorescence at low pressure, *Journal of Geophysical Research: Atmospheres*, 99, 3543-3557, <https://doi.org/10.1029/93JD03342>, 1994.
- Stutz, J., Oh, H.-J., Whitlow, S. I., Anderson, C., Dibb, J. E., Flynn, J. H., Rappenglück, B., and Lefer, B.: Simultaneous DOAS and mist-chamber IC measurements of HONO in Houston, TX, *Atmospheric Environment*, 44, 4090-4098, <https://doi.org/10.1016/j.atmosenv.2009.02.003>, 2010.
- 345 Tuite, K., Thomas, J. L., Veres, P. R., Roberts, J. M., Stevens, P. S., Griffith, S. M., Dusanter, S., Flynn, J. H., Ahmed, S., and Emmons, L.: Quantifying nitrous acid formation mechanisms using measured vertical profiles during the CalNex 2010 campaign and 1D column modeling, *Journal of Geophysical Research: Atmospheres*, 126, e2021JD034689, <https://doi.org/10.1029/2021JD034689>, 2021.
- 350 Veres, P. R., Roberts, J. M., Wild, R. J., Edwards, P. M., Brown, S. S., Bates, T. S., Quinn, P. K., Johnson, J. E., Zamora, R. J., and de Gouw, J.: Peroxynitric acid (HO₂NO₂) measurements during the UBWOS 2013 and 2014 studies using iodide ion



- chemical ionization mass spectrometry, *Atmospheric Chemistry and Physics*, 15, 8101-8114, <https://doi.org/10.5194/acp-15-8101-2015>, 2015.
- 355 Veres, P. R., Neuman, J. A., Bertram, T. H., Assaf, E., Wolfe, G. M., Williamson, C. J., Weinzierl, B., Tilmes, S., Thompson, C. R., and Thames, A. B.: Global airborne sampling reveals a previously unobserved dimethyl sulfide oxidation mechanism in the marine atmosphere, *Proceedings of the National Academy of Sciences*, 117, 4505-4510, <https://doi.org/10.1073/pnas.1919344117>, 2020.
- Villena, G., Kleffmann, J., Kurtenbach, R., Wiesen, P., Lissi, E., Rubio, M. A., Croxatto, G., and Rappenglück, B.: Vertical gradients of HONO, NO_x and O₃ in Santiago de Chile, *Atmospheric Environment*, 45, 3867-3873, <https://doi.org/10.1016/j.atmosenv.2011.01.073>, 2011.
- 360 Villena, G., and Kleffmann, J.: A source for the continuous generation of pure and quantifiable HONO mixtures, *Atmospheric Measurement Techniques*, 15, 627-637, <https://doi.org/10.5194/amt-15-627-2022>, 2022.
- Whalley, L. K., Stone, D., Dunmore, R., Hamilton, J., Hopkins, J. R., Lee, J. D., Lewis, A. C., Williams, P., Kleffmann, J., and Laufs, S.: Understanding in situ ozone production in the summertime through radical observations and modelling studies during the Clean air for London project (ClearLo), *Atmospheric Chemistry and Physics*, 18, 2547-2571, <https://doi.org/10.5194/acp-18-2547-2018>, 2018.
- 365 Wong, K., Tsai, C., Lefer, B., Haman, C., Grossberg, N., Brune, W., Ren, X., Luke, W., and Stutz, J.: Daytime HONO vertical gradients during SHARP 2009 in Houston, TX, *Atmospheric Chemistry and Physics*, 12, 635-652, <https://doi.org/10.5194/acp-12-635-2012>, 2012.
- 370 Ye, C., Gao, H., Zhang, N., and Zhou, X.: Photolysis of Nitric Acid and Nitrate on Natural and Artificial Surfaces, *Environ Sci Technol*, 50, 3530-3536, <https://doi.org/10.1021/acs.est.5b05032>, 2016.
- Ye, C., Zhang, N., Gao, H., and Zhou, X.: Photolysis of Particulate Nitrate as a Source of HONO and NO_x, *Environ Sci Technol*, 51, 6849-6856, <https://doi.org/10.1021/acs.est.7b00387>, 2017.
- 375 York, D., Evensen, N. M., Martinez, M. L., and De Basabe Delgado, J.: Unified equations for the slope, intercept, and standard errors of the best straight line, *American Journal of Physics*, 72, 367-375, <https://doi.org/10.1119/1.1632486>, 2004.
- Young, C. J., Washenfelder, R. A., Roberts, J. M., Mielke, L. H., Osthoff, H. D., Tsai, C., Pikelnaya, O., Stutz, J., Veres, P. R., Cochran, A. K., VandenBoer, T. C., Flynn, J., Grossberg, N., Haman, C. L., Lefer, B., Stark, H., Graus, M., de Gouw, J., Gilman, J. B., Kuster, W. C., and Brown, S. S.: Vertically resolved measurements of nighttime radical reservoirs in Los Angeles and their contribution to the urban radical budget, *Environ Sci Technol*, 46, 10965-10973, <https://doi.org/10.1021/es302206a>, 380 2012.