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Harmonized retrieval of middle atmospheric ozone from two microwave radiometers in Switzerland

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Abstract. We present new harmonized ozone time series from two ground-based microwave radiometers in Switzerland: GROMOS and SOMORA. Both instruments measure hourly ozone profiles in the middle atmosphere ($20-75$ km) since more than two decades. As inconsistencies in long-term trends derived from these two instruments were detected, a harmonization project was initiated in 2019. The goal was to understand and reduce the discrepancies between the two data records. The harmonization has been completed for the data from 2009 until 2022 and has been successful at reducing the differences observed between both time series. It also explains the remaining differences between both instruments and flags their respective anomalous measurement periods in order to adapt their consideration for future trend computations.

We describe the harmonization and the resulting time series in detail. We also highlight the improvements in the ozone retrievals with respect to the previous data processing. In the stratosphere and lower mesosphere, the seasonal ozone relative differences between both instruments are now within 10 % and show good correlations ($R > 0.7$) except for summertime. We also perform a comparison of these new data series against measurements from the satellite instruments Microwave Limb Sounder (MLS) and Solar Backscatter Ultraviolet Radiometer (SBUV) over Switzerland. Seasonal mean differences with MLS and SBUV are within 10 % in the stratosphere and lower mesosphere up to 60 km and increase rapidly above.

1 Introduction

Ozone is a trace gas of great importance in the earth's atmosphere. It shields the surface of our planet from most of the sun's harmful ultraviolet radiation by absorbing it in the stratosphere (the "ozone layer") and consequently allowing life out of water. In the second half of the twentieth century, it was suggested that anthropogenic emissions of certain chemical compounds, the commonly called ozone-depleting substances (ODSs), were threatening this protective layer (Molina and Rowland, 1974; Crutzen, 1970; Farman et al., 1985; Solomon et al., 1986). As a result, severe depletion of the ozone layer was observed in the spring time over the Antarctic and led to the banning of ODS emissions formalized in the Montreal Protocol in 1987.

Since then, there has been an increased interest in the monitoring of ozone in the middle atmosphere to assess the effect of the Montreal Protocol. The reduction of ODSs emission has led to a decrease in total chlorine concentration since 1997 whereas the increasing greenhouse gases concentration is cooling the upper stratosphere (Anderson et al., 2000; Solomon et al., 2006). From the existing knowledge in middle-atmospheric chemistry, the combination of both factors should lead to

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an observable recovery or even super recovery of ozone concentration at these altitudes (Eyring et al., 2010). In fact, over the polar regions, the stratospheric ozone concentrations have already begun their recovery towards pre-industrial levels (Solomon et al., 2016). Over the mid-latitudes, the situation is less obvious, and ozone recovery seems to differ depending on the altitude and the geographical area of interest (Braesicke et al., 2018; Petropavlovskikh et al., 2019; Tummon et al., 2015). In the upper stratosphere, the latest observations agree on a positive trend of ozone concentration despite a high variability in its significance and magnitude (Fahey et al., 2018; Steinbrecht et al., 2017; Bernet et al., 2019; Godin-Beekmann et al., 2022). In contrast, no clear indication of ozone recovery has been reported yet in the lower stratosphere and some observational evidence of further decline in this region were even reported (Ball et al., 2018). In a context of climate change, there remain also many unknowns regarding the influence of long-term dynamic and composition changes on middle-atmospheric ozone trends depending on the region (von der Gathen et al., 2021). In regards to these uncertainties, there is still a high need for accurate and long-term time series in the research field.

Microwave ground-based radiometers (MWRs) provide continuous, all-weather measurement of ozone in the middle atmosphere and are therefore well suited to estimate long-term trends and cross-validate satellite measurements (Hocke et al., 2007). Compared to other ground-based measurement techniques, they are able to retrieve ozone profiles from the stratosphere well into the mesosphere with a high temporal resolution but at the cost of a quite low vertical resolution.

In Switzerland, two ozone MWRs are operated since more than 20 years close to each other: the GROund-based Millimeter-wave Ozone Spectrometer (GROMOS) in Bern and the Stratospheric Ozone MOnitoring RAdiometer (SOMORA) in Payerne (Fig. 1). They operate in the frame of the Network for the Detection of Atmospheric Composition Change (NDACC) (De Mazière et al., 2018). Such long-term time series of two ozone MWRs combined to a geographic proximity is unique worldwide and therefore offers the opportunity for extensive cross-validations. It also allows to investigate more thoroughly measurement uncertainties, possible instrumental failures, calibration and retrieval errors.

During the first phase of the activity "Long-term Ozone Trends and Uncertainties in the Stratosphere" (LOTUS), inconsistencies were found in ozone trend estimates from these two radiometers (Petropavlovskikh et al., 2019). In addition, Bernet et al. (2019) identified some anomalous period in the Bern time series and highlighted the need to account for these anomalies to compute more accurate trends. However, Bernet et al. (2019) did not investigate the reasons for such anomalies, and the differences between these two time series remained unexplained. Therefore, a harmonization project was initiated jointly by the operators of these two instruments in 2019 with the goal to understand their differences and reconcile their series while keeping them independent.

We present and validate here the new harmonized time series for GROMOS and SOMORA. We focus on the time period when both instruments use the same spectrometer, namely from the month of September 2009 until December 2021. We present the harmonization process applied to the data processing of the two radiometers, including a short description of the new calibration and retrieval routines. We also show the improvements resulting from this harmonization by comparing the new series with their previous versions. As a validation, we performed a cross-comparison between both instruments and compared them against satellite dataset, namely from the Microwave Limb Sounder (MLS) and the Solar Backscatter Ultraviolet Radiometer (SBUV).

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A detailed description of the calibration and retrievals routines have been published in the form of two research reports available on the publication database of the University of Bern (Sauvageat, 2021, 2022) and a full documentation of the time series is available together with the data.

The manuscript is organized as follows. Section 2 presents the instruments, highlighting their similarities and differences. Section 3 presents the harmonization procedure applied to the calibration and retrieval routines. Section 4 presents the new harmonized ozone time series whereas section 5 presents comparisons and cross-validations against satellite measurements. Section 6 summarizes the main conclusions and gives an outlook.

2 Ozone microwave radiometry in Switzerland

Passive microwave radiometry uses the electromagnetic radiation emitted and transmitted in the microwave frequency region to derive geophysical quantities of interest. The large wavelengths associated with this spectral region make this technique relatively insensitive to atmospheric conditions. Therefore, it is suitable for both earth's surface observation from space and sounding of atmospheric trace gases, temperature or winds from satellites or ground-based instruments. In addition, the pressure broadening effect at microwave frequencies enables to retrieve vertical profiles of temperature, winds and abundances (e.g. Connor et al., 1994; Rüfenacht et al., 2012; Krochin et al., 2022).

Ozone possesses many rotational transition lines in the microwave region. Its emission lines at 110.836 and 142.175 GHz are most often used for ground-based observations because of their line intensity and the limited effect of water-vapor absorption at these frequencies.

GROMOS and SOMORA have been designed and built at the Institute of Applied Physics (IAP) at the University of Bern with quite similar components (Calisesi, 2003; Peter, 1997). They observe the ozone emission line around 142 GHz to retrieve hourly ozone profiles in the stratosphere and lower mesosphere (~ 20 to 75 km) using the Optimal Estimation Method. GROMOS is currently operated by IAP in Bern since 1994 and SOMORA is operated by the Federal Office of Meteorology and Climatology MeteoSwiss in Payerne since 2000 (see location in Fig. 1). The main characteristics of both instruments are summarized in Table 1.

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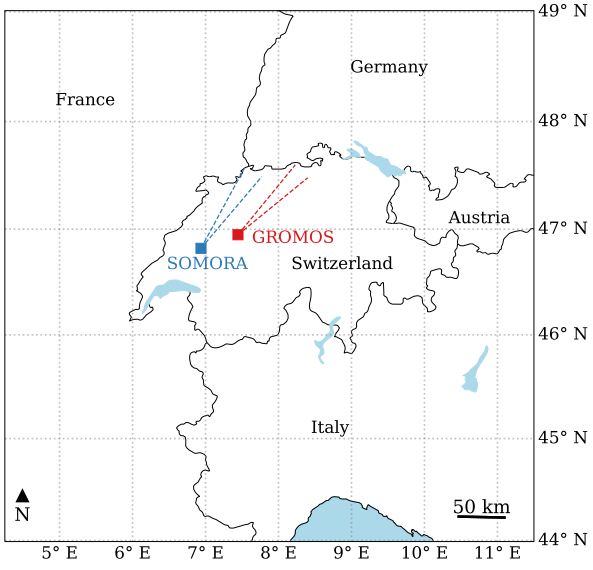

Figure 1. Location of GROMOS and SOMORA, with their approximate viewing directions.

Table 1. GROMOS and SOMORA microwave radiometers

	GROMOS	**SOMORA**
Location	Bern	Payerne
Latitude	46.95° N	46.82° N
Longitude	7.44° E	6.94° E
Altitude	560 m	491 m
Azimuth angle	45°	34°
Elevation angle	40°	39°
Observation frequency	142.175 GHz	142.175 GHz
Spectrometer	Acqiris AC240	Acqiris AC240
Bandwidth	1 GHz (32768 channels)	1 GHz (16384 channels)
Intermediate frequency	3.7 GHz	7.1 GHz
Frequency resolution	30.52 kHz	61.04 kHz
T_{rec}	$\sim$ 2750 K	$\sim$ 2550 K

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2.1 Spectrometers

The spectrometer is a key component of any MWR and can influence significantly its retrieval capabilities. Since 2009, both instruments use the same spectrometer, namely the Acqiris AC240 which is a digital Fast Fourier Transform (FFT) spectrometer (Benz et al., 2005; Muller et al., 2009). On SOMORA, it replaced an acousto-optical spectrometer in September 2009 whereas on GROMOS, it replaced discrete filter banks in July 2009. In both cases, the time series were homogenized using an overlap period of roughly two years and the pre-2009 time series were corrected with respect to the FFT spectrometer time series (e.g. Moreira et al., 2015; Maillard Barras et al., 2020). Whereas both instruments use the same digitizer with the same bandwidth of 1 GHz, it should be noted that the frequency resolution is two times higher for GROMOS than for SOMORA because GROMOS uses an in-phase quadrature (IQ) down-converter and digital sideband separation, which results in twice the number of channels (Murk et al., 2009). As a result, GROMOS should be more sensitive to ozone at higher altitude, but as it also has a higher noise compared to SOMORA, we do not see any significant differences in their sensitivity.

The AC240 is still being used in many MWRs however, it is ageing and has recently been shown to produce a spectral bias compared to more recent spectrometers, most likely impacting ozone retrievals as well (Sauvageat et al., 2021). In this contribution, we only focus on the period where both instruments use the AC240, namely from September 2009 to end of 2021. Therefore, both time series should be similarly impacted by the spectrometric bias and it should not question the results of the comparisons between GROMOS and SOMORA. This might however, influence the comparisons against the satellite observations but there is unfortunately no way to confirm the amplitude of the effect of the bias on the ozone profiles at the moment.

3 Harmonization process

Discrepancies were identified between GROMOS and SOMORA data series and trends (Bernet et al., 2019; Petropavlovskikh et al., 2019; Maillard Barras et al., 2020), for which no explanations could be found. Due to their geographic proximity and similar observation geometry, the differences are too big to be geophysical. The data processing however, was quite different between the instruments and therefore it was decided to reprocess both time series with new and harmonized algorithms. It aims at a complete harmonization of the data processing, from the raw data (level 0) to the ozone profiles (level 2) with the purpose to improve the agreement between both instruments while keeping them fully independent.

The harmonization project can be separated in two distinct parts: the calibration of the radiometric data (level 0 to 1) and the retrievals of ozone profiles (level 1 to 2). Section 3.1 will briefly describe the new calibration and integration routines (see Sauvageat (2021) for details) whereas section 3.2 will describe the retrievals of ozone profiles from the calibrated spectra.

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3.1 Calibration

GROMOS and SOMORA are both total power radiometers with superheterodyne receivers. They measure the atmospheric ozone emission line around 142.175 GHz and use the heterodyne principle to down convert the incoming radiation (RF signal) to an intermediate frequency (IF) by mixing with a local oscillator frequency (LO) which allows for easier signal processing.

Due to their high sensitivity, the operation of microwave radiometers requires continuous calibration (e.g. Ulaby and Long, 2014, chap. 7) and both instruments use a so-called hot-cold calibration scheme. Using a rotating mirror fixed on a path length modulator, they are continuously switching between the atmospheric observation, a hot and a cold calibration target. In both instruments, a heated black-body kept at a constant temperature ($T_{hot} \approx 310$K) is used as hot load whereas liquid nitrogen (LN2) observation is used as cold load. Both instruments use a Martin-Pupplet Interferometer (MPI) to suppress the contribution of the undesired sideband. The pathlength modulator is used to mitigate the standing-waves between the receiver and the calibration targets, which are otherwise causing systematic baseline errors on the calibrated spectra. In parallel to the hot-cold calibration scheme, the instruments also perform tipping curve calibration (Ingold et al., 1998) as cross-validation for the LN2 calibration. Assuming linear transfer characteristics, the atmospheric spectral radiance can then be determined and further converted to brightness temperature using the Planck's law (e.g. Ulaby and Long, 2014, chap. 6).

Despite similar design and raw data content, the previous calibration routines for GROMOS and SOMORA were different. Therefore, a new routine was designed to harmonize the calibration between the two instruments. The calibration essentially converts the raw spectrometer measurements to radiance intensity and integrates them together on a chosen integration time. For this new routine, the calibration results in two different data levels, namely the calibrated spectrum (level 1a) and the integrated spectrum (level 1b).

Harmonized quality control was introduced in order to identify spurious instrumental signals. It flags the most common technical problems at the level 1a (e.g. Noise temperature jumps, LN2 refills, LO frequency shifts...) and combines them into a single instrumental flag value for the level 1b (Sauvageat, 2021).

Considering instrumental issues and technical interruptions for maintenance (e.g. for LN2 refilling or instrument repairs), GROMOS and SOMORA provided respectively 87 % and 89 % of good quality hourly spectra on the period 2009-2021, with over $80'000$ hours of comparable measurements.

3.2 Retrieval setup

In the microwave frequency range, the pressure-broadening effect of atmospheric emission lines is used to retrieve information on the atmospheric constituent profile from the calibrated microwave emission spectra. This so-called retrieval is a well validated technique which has been successfully applied to temperature, wind, and many trace gases like O_3, CO or H_2O (Janssen, 1993, chap. 7). Among the different retrieval techniques, we selected the Optimal Estimation Method (OEM) following the formalism described by Rodgers (2000). This statistical method extracts the best estimate of an atmospheric profile from a set of measurements with noise, a priori information, and a forward model. In addition, the OEM enables to characterize the error budget of the retrievals (Fig. 3 and Fig. 4). In the following, we will briefly present and discuss the new harmonized retrieval

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setup used for GROMOS and SOMORA. More information on this setup is available in Sauvageat (2022). For detailed information on the OEM or its application to ozone profiling instruments, the reader is redirected to Parrish et al. (1992) or Tsou et al. (1995).

3.2.1 Forward model

In the case of ground-based microwave radiometry, the forward model (FM) describes the radiative transfer physics between trace gases emissions and the instrument's receiver. We used the Atmospheric Radiative Transfer Simulator 2.4 (ARTS), an open source software with a special focus on microwave radiative transfer simulations (Eriksson et al., 2011; Buehler et al., 2018). In addition, it offers a fully integrated OEM retrieval environment and includes many tools to help simulate and retrieve the sensor's influence on the radiometric measurements (Eriksson et al., 2006).

ARTS offers many possibilities to define the atmospheric state, a priori data, and simulation grids. We use one-dimensional pressure and temperature profiles from the European Centre for Medium-Range Weather Forecasts (ECMWF) daily operational analysis (6 hour time and $1.125°$ spatial resolution). This dataset is limited to approximately $70\ \mathrm{km}$ altitude and therefore, we extend it using the COSPAR International Reference Atmosphere (CIRA-86) climatology at upper altitudes (Chandra et al., 1990). The frequency grids have been defined to cover the range of GROMOS and SOMORA spectrometers with a refined frequency resolution around the ozone line: it matches the spectrometer resolution at the line center to optimize retrievals at higher altitudes, whereas the spectral resolution is coarser on the line wings to limit computation time.

Table 2. Main parameters used in GROMOS and SOMORA retrievals

Forward model	ARTS
Species	O_3, H_2O, O_2 and N_2
Spectroscopy	Perrin (JPL & HITRAN)
Atmospheric state	1D ECMWF & CIRA 86
O_3 a priori	WACCM
H_2O a priori	ECMWF
FM grid	$\sim 1 - 112\ \mathrm{km}$, $2\ \mathrm{km}$ resolution
Retrieval grid	$\sim 1 - 95\ \mathrm{km}$, $2\ \mathrm{km}$, resolution

As atmospheric species, we use ozone, water vapour, oxygen and nitrogen. For ozone, we use the spectroscopic database from Perrin et al. (2005), which is provided with ARTS 2.4 and is derived from the HITRAN and JPL spectroscopic databases. For water vapour, oxygen and nitrogen, we use the parametrizations provided within ARTS (see Buehler et al., 2005). A

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summary of the main retrievals parameters used for GROMOS and SOMORA can be found in Table 2 and more details are provided in Sauvageat (2022).

3.2.2 Ozone retrieval

The main retrieval quantity is hourly ozone volume mixing ratio (VMR) from the stratosphere to the lower mesosphere, i.e. between ~ 100 and 0.01 hPa. The a priori are monthly ozone profiles extracted from free-running simulations of the Whole Atmosphere Community Climate Model (WACCM) as described in Schanz et al. (2014). Further, depending on the local solar time, we either use a daytime or nighttime a priori ozone profile. The a priori covariance matrix for ozone varies with atmospheric pressure in order to optimize the information from the measurements in the stratosphere and lower mesosphere. It includes exponentially decreasing covariances between pressure levels to reflect the vertical coupling of the atmosphere.

3.2.3 Sensor and noise

The accuracy of the retrievals can be improved by taking the systematic characteristics of the instrument into account. ARTS has dedicated built-in functions that can model the influence of the most relevant components on the atmospheric observations (Eriksson et al., 2006). For GROMOS and SOMORA, we included the effect of the FFT spectrometer channel response ($|\frac{\sin(x)}{x}|^2$) and the effect of the sideband ratio.

The measurement noise is an important quantity for OEM retrievals because it defines, together with the a priori covariance, the information that can be extracted from the measurement at each pressure level. The noise covariance matrix is computed independently for each instrument and each retrieval based on the noise level observed on the integrated spectrum and is considered to be uncorrelated between the different channels similarly as explained in Krochin et al. (2022). It is slightly higher for GROMOS (≈ 0.7 K) than SOMORA (≈ 0.5 K) because GROMOS has a higher receiver noise temperature and a higher frequency resolution.

3.2.4 Additional retrieval quantities

There are other sensors or external influences which are difficult to estimate and correct during the calibration process or to simulate accurately for each spectrum. This is the case for the instrumental baselines and the tropospheric absorption. The instrumental baselines are modulation of the atmospheric spectrum due to the observing system. They can arise during the mixing process, the sideband filtering or can be due to undesired reflections, typically when observing the calibration targets. In ARTS, it is possible to consider them as unknown and add them as additional retrieval quantities.

Around the 142 GHz ozone line, the tropospheric water continuum contributes significantly to the observed spectra and has to be considered during the inversion process. One simple correction method is the so-called tropospheric correction (Ingold et al., 1998) but it is certainly a better solution - also in view of assessment of the error propagation - to include the tropospheric water vapour as a retrieval quantity within ARTS, as has been done previously for such retrievals (e.g. in Palm et al., 2010).

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A frequency shift was also retrieved for each spectrum because the local oscillators of both GROMOS and SOMORA are not perfectly stable and even a slight shift of the reference frequency can bias the ozone profile retrievals.

Despite mitigation of instrumental baselines using different techniques (e.g., mirror wobbling, non-perpendicular aspect of cold load), it is often necessary to retrieve some instrumental baselines as well (Palm et al., 2010). In the case of GROMOS and SOMORA, we include a second-order polynomial and different sinusoidal baselines. In order to avoid the degradation of the retrievals with the addition of too many sinusoidal baselines, we processed first the full time series without any sinusoidal baselines and used the residuals to compute the main sinusoidal baseline periods for each instrument. We observed that the sinusoidal baseline periods remain similar on time-scale of month to years so in practice, only a few period changes were applied during the full extent of the time series for each instrument (see Sauvageat (2022) for details).

3.2.5 Retrieval results

For each retrieval quantity, the OEM returns the statistical best estimates of the results and ARTS returns the corresponding fitted atmospheric spectrum which can be compared against the MWR observation to evaluate the goodness of the fit. Figure 2 shows examples of hourly integrated spectra from GROMOS and SOMORA together with their fitted measurement spectra.

Figures 3 and 4 show the corresponding ozone retrievals and main diagnostic quantities for the spectra shown in Fig. 2. It includes the averaging kernels (AVKs) which are a measure of the sensitivity of the retrieval to the true ozone profile at each pressure level. The sum of the AVKs at each level defines the measurement response (MR). It is an indication of the measurement contribution to the retrieved profile, whereas the remaining information comes from the a priori. In microwave remote sensing, a MR of $80\ \%$ is often used to define the lower and upper boundaries of the retrievals in order to limit the influence of the a priori on the results. Also included as diagnostic quantities are the smoothing and measurement errors computed by the OEM as defined by Rodgers (2000). The smoothing error is a consequence of the limited resolution of the instrument whereas the measurement error arises from the noisy nature of the observations. Finally, we show the Full Width at Half Maximum (FWHM) of the AVKs at each level and the altitude offset (in km) between the AVKs maximum and its corresponding altitude. Both together give an indication on the altitude resolution and the vertical offset between the true and retrieved profiles.

3.3 Uncertainty budget

The retrievals errors presented above do not include systematic errors that can arise during the calibration or the retrievals. It is cumbersome to estimate all possible errors on such complex measurement setup and therefore, we decided to perform a sensitivity analysis on the most important error sources using two reference time periods with low and high atmospheric opacities. The uncertainties considered in our study are listed in Table 3 as well as the perturbations used for the sensitivity analysis. These were determined in different ways for each error source, deriving it either from measurement (e.g. T_{cold}, sideband ratio, window transmittance) or empirical values (e.g. pointing, spectroscopy).

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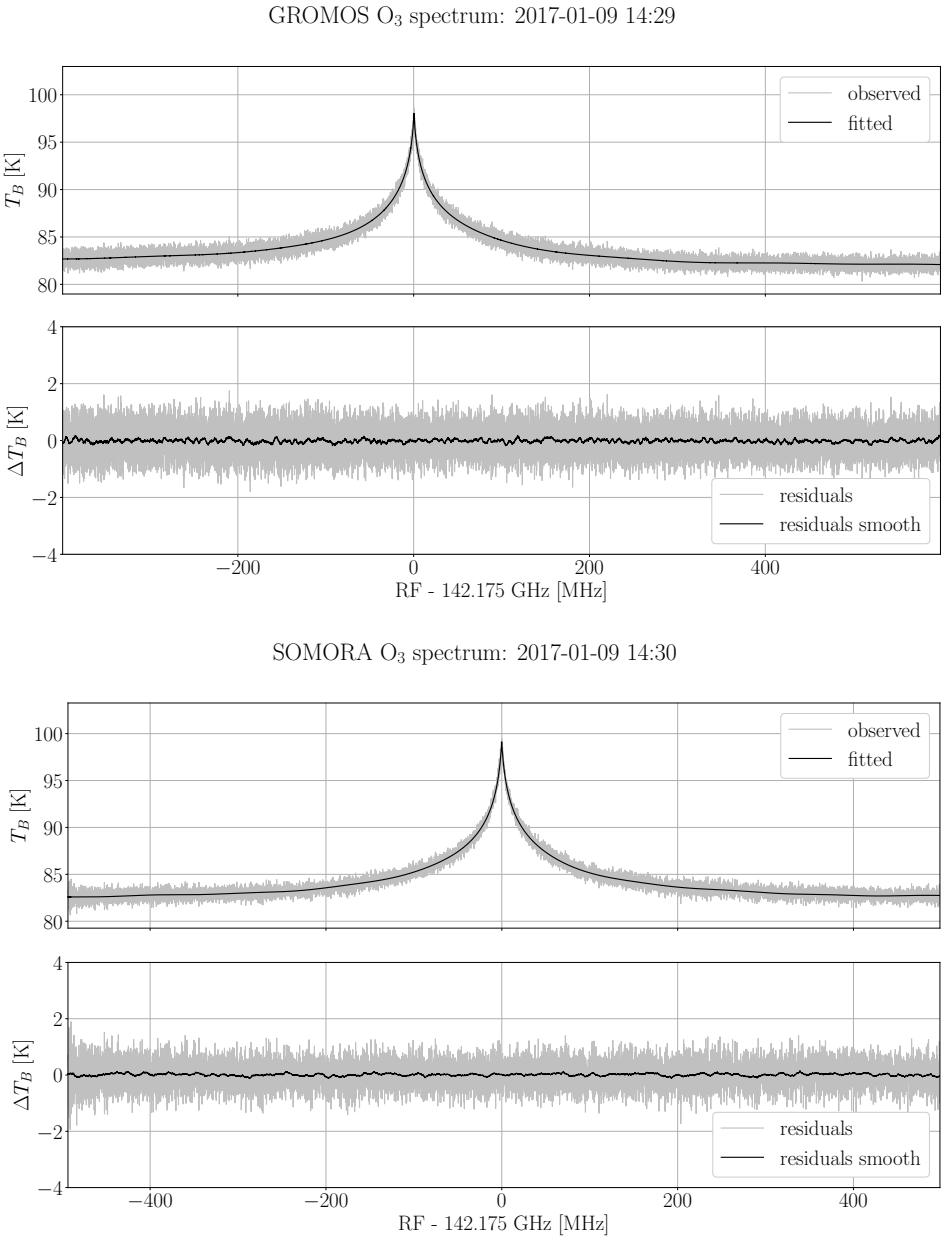

Figure 2. Integrated and fitted spectrum for GROMOS and SOMORA, binned to the same spectral resolution. The lower panels show the residuals, i.e. the differences between the measurement and the fitted spectrum. The smoothed residuals are computed using a running mean over 128 channels.

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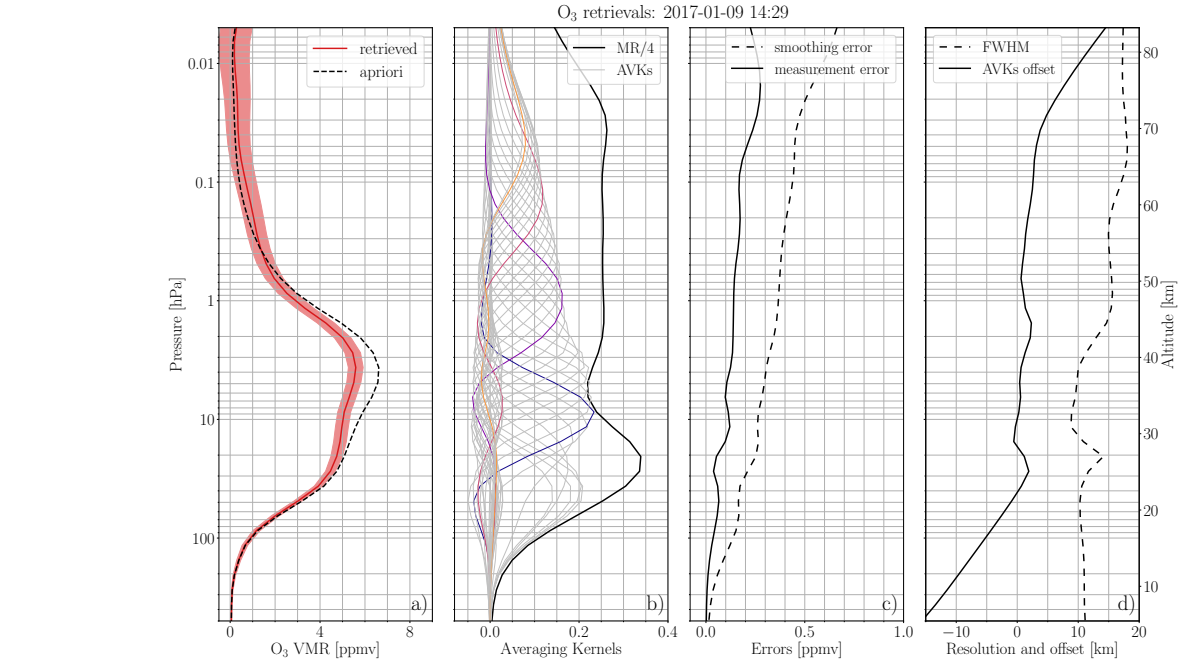

Figure 3. Example of GROMOS retrievals on 09.01.2017 a) shows the a priori and retrieved ozone profiles, b) shows the averaging kernels together with the MR (divided by 4 to fit in the same plot), c) shows the smoothing and measurement error and d) shows the Full Width at Half Maximum (FWHM) and the offset between the AVKs peak and its corresponding altitude. All quantities are retrieved on pressure levels and corresponding altitudes are indicated on the right. See text for more details on each diagnostic quantity.

Table 3. Potential error sources and the perturbations used for the sensitivity analysis

pointing	error on the zenith angle	1°
T_{cold}	cold calibration target temperature	2 K
window transmittance	transmittance of the windows in front of the instrument	3 %
$T_{profile}$	constant offset in atmospheric temperature profile	5 K
spectroscopy	error in spectroscopic line intensity	3 %
sideband ratio	error in MPI path length difference	0.05 mm

The uncertainty budget for GROMOS and SOMORA is presented in Fig. 5 and Fig. 6 in the case of low tropospheric opacities. The high opacity cases for both instrument can be seen in Appendix A (Fig. A1 for GROMOS and Fig. A2 for SOMORA).

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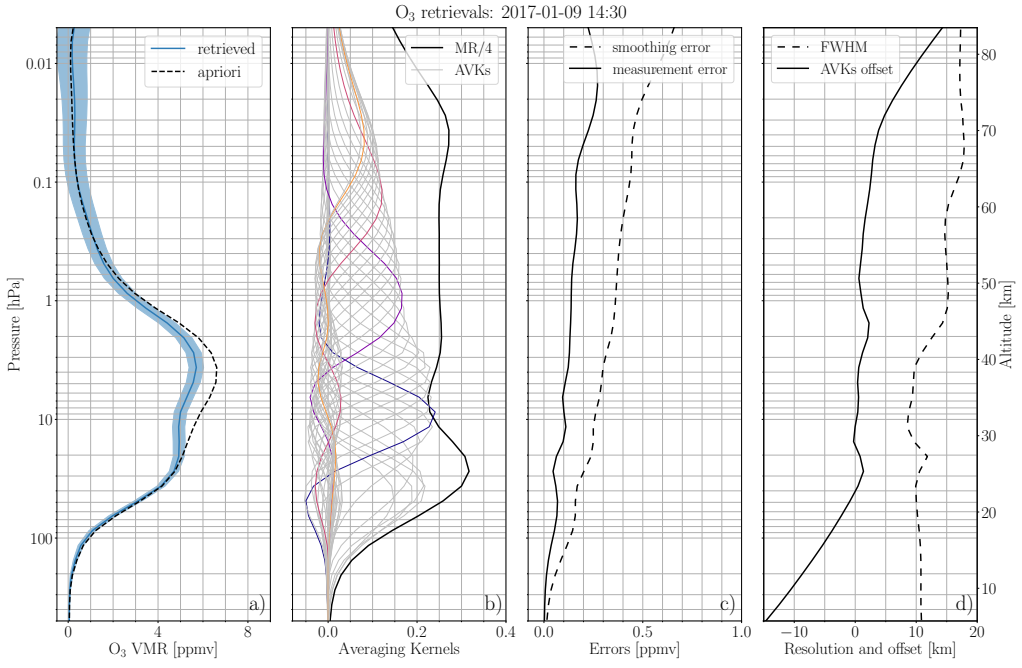

Figure 4. Same as Fig. 3 but for SOMORA.

In general, the sensitivity of GROMOS and SOMORA to the different perturbations is very similar. Notable exception is the GROMOS higher sensitivity to the sideband path length which is a consequence of its lower intermediate frequency. For both instruments, the total uncertainty is dominated by systematic errors below 2 hPa whereas the measurement noise becomes quickly dominant above. In relative terms, the uncertainty is approximately $9-10$ % for GROMOS, respectively $7-8$ % for SOMORA, up to the stratopause and increases then significantly in the mesosphere. In the case of high atmospheric opacity, the uncertainties are higher which is expected since the ozone emission line gets more attenuated by the tropospheric water vapor absorption. In this case, the total relative uncertainty on GROMOS in the stratosphere is $12-15$ %, respectively $10-12$ % for SOMORA. In views of the perturbations and error sources considered in this study, these values compare well with similar ozone radiometers at other locations reported in the literature (e.g., Palm et al., 2010; Kopp et al., 2002).

4 Harmonized ozone time series

Using the new calibration and retrieval routines described previously, we have reprocessed the GROMOS and SOMORA data series for the time where they both use the AC240 spectrometer, i.e. from the end of 2009 until 2021. Fig. 7 shows weekly averaged ozone profiles for GROMOS and SOMORA for the decade 2010-2020. It shows the consistency of the measurements and highlights the very few large interruptions happening on both instruments during this period. Most interruptions are due to instrumental issues (LN2 refilling, LO frequency stability, ...) or to the atmospheric conditions (e.g. high tropospheric opacity

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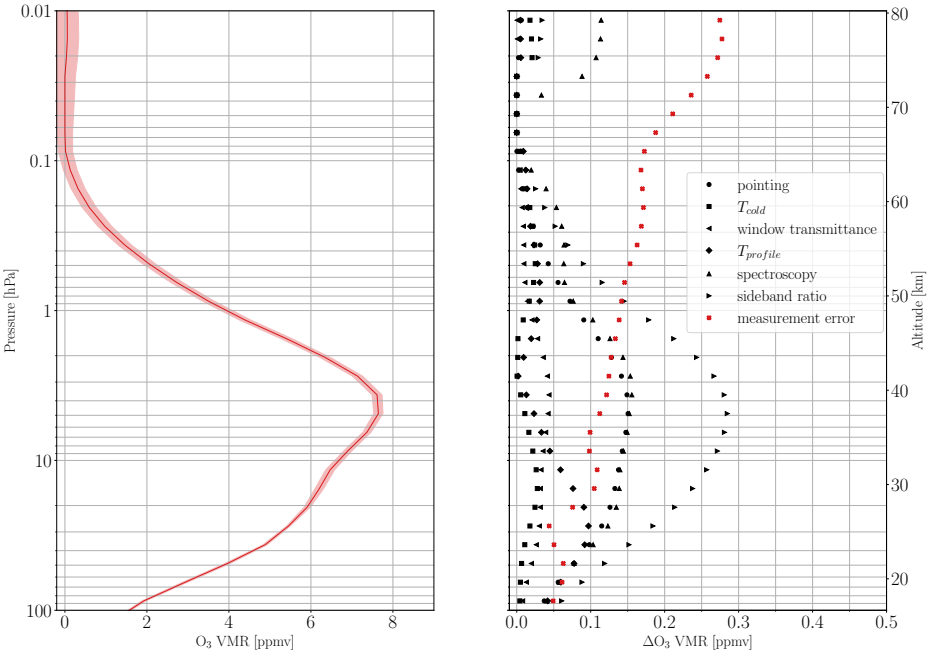

Figure 5. Uncertainty budget for GROMOS at low opacity case. The left panel show the reference ozone profile chosen for the sensitivity analysis. The right panel shows the ozone VMR uncertainties arising from the error sources listed in Table 3.

masking the ozone emission line) and they usually last for a few hours at most. The longer interruptions result from cold load issue or hardware changes which can last for a few days or weeks.

To validate these two data series, we first present a cross-comparison of GROMOS and SOMORA data series and show the improvement resulting from the reprocessing compared to the previous retrieval version. We then compare both instruments against space-based ozone observations from MLS and SBUV above Switzerland.

4.1 Cross-comparison between GROMOS and SOMORA

In views of their close location, GROMOS and SOMORA can be used for cross-validation of their time series. In addition, both instruments have similar viewing direction, altitude range and sensitivity so they can be compared directly. The upper panel in Fig. 10 shows the weekly mean relative differences between GROMOS and SOMORA. In general, GROMOS and SOMORA agree well in most of the middle atmosphere, with relative differences lower than $10\ \%$ in the upper stratosphere and lower mesosphere (from ~ 50 to 0.1 hPa), increasing towards lower and higher altitudes. The higher relative differences at lower and higher altitudes are partly explained by the shape of the ozone VMR profile which intensity is maximum in the

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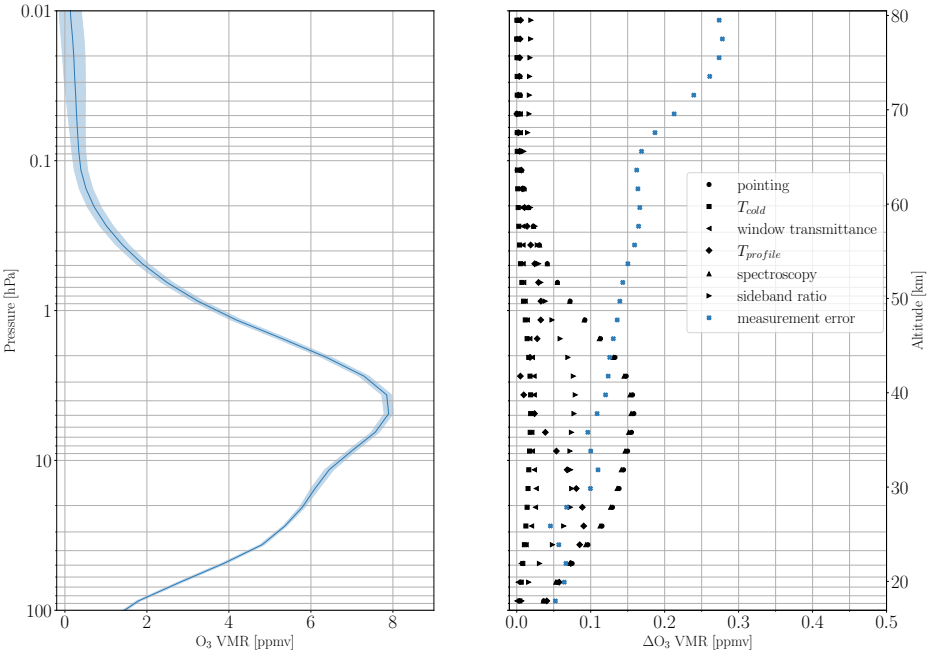

Figure 6. Same as Fig. 5 but for SOMORA.

stratosphere. In general, the lower altitudes are also the most impacted by instrumental baselines which explains the increase of the differences below 50 hPa whereas at higher altitudes, the instrumental noise becomes the dominant factor and the sensitivity of the radiometers decreases quickly. In addition, the diurnal ozone variations typically become much larger in the mesosphere (e.g. around 20 % compared to a few percent in the stratosphere (Haefele et al., 2008)).

We also see some oscillatory patterns in the relative differences, some of which can be identified as clear seasonal patterns (e.g. in the lower stratosphere between 2014 and 2017). The reasons for such seasonal patterns are not fully clear but are most likely the result of seasonal temperature and humidity cycle in the troposphere. Indeed, despite controlled room temperature for both instruments, the higher summer temperature and humidity still influence room and window temperatures and consequently the instruments (e.g. receiver noise temperature). We believe that the hardware components of GROMOS and SOMORA have different sensitivity to such influences, which could explain the seasonal patterns observed in their relative differences. The warmer and wetter troposphere also emits more microwave during the summertime which tends to stronger absorb the ozone spectral lines and decreases the retrieval sensitivity during summer. As discussed in section 3.3, a higher tropospheric opacity also results in larger uncertainties on the retrieved ozone profile. In case of very hot and humid conditions, the troposphere can become optically thick at 142 GHz which can prevent the retrieval of ozone profiles.

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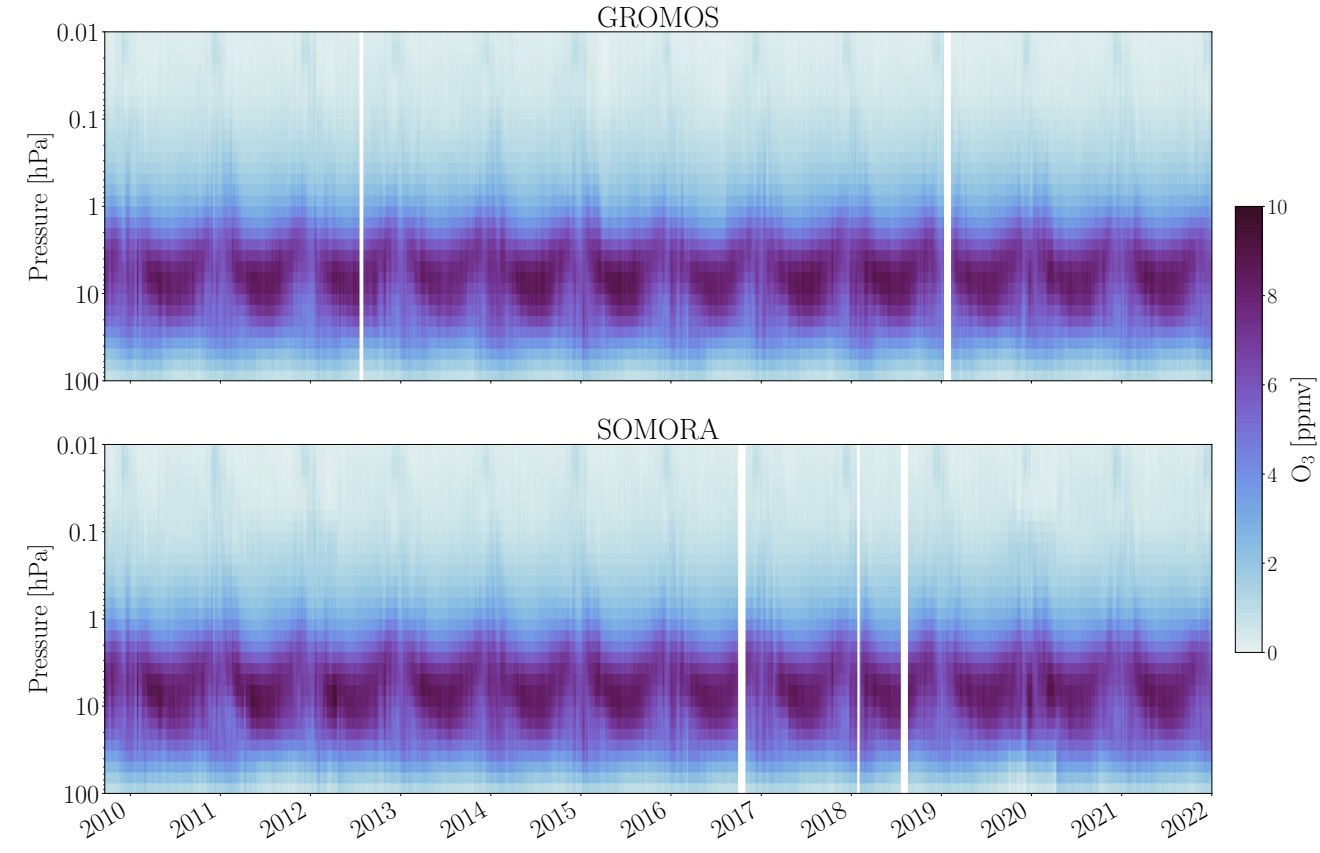

Figure 7. Weekly averaged ozone volume mixing ratios (VMR) profiles for GROMOS and SOMORA.

Table 4. Definition of the three pressure ranges and corresponding altitudes

Region	**Pressure range** [hPa]	**Approximate altitudes** [km]
Lower mesosphere	$0.9 - 0.1$	$\sim 50 - 65$
Upper stratosphere	$5 - 1$	$\sim 38 - 50$
Lower stratosphere	$50 - 10$	$\sim 22 - 32$

These seasonal differences are highlighted in Fig. 8 which show seasonal ozone profile comparisons between GROMOS and SOMORA. The mean seasonal differences between the two instruments are lower than 10 % at all seasons and throughout most of the middle atmosphere and show a negative ozone bias from GROMOS in the upper mesosphere ($p < 0.05$ hPa). The seasonal profiles have strong correlations with Pearson's R coefficients mostly above 0.5 at most pressure levels.

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Figure 9 shows scatter plots of their differences in three pressure level domains corresponding approximately to the lower stratosphere, the upper stratosphere and the lower mesosphere (see Table 4 for the definitions). It shows the net difference in atmospheric opacity between the winter and the summer and highlights the higher ozone variability during the winter time. It results in stronger correlations in the winter time and lower values during the summer season at all pressure ranges. Overall though, there is a good agreement between GROMOS and SOMORA data series in the middle atmosphere.

In addition to these seasonal effects, Fig. 10 highlights some sudden changes in the differences between the two instruments, most of which can be related to specific instrumental issue on either instrument. It can be seen for instance in April 2012, where the cold load observation angle was changed on SOMORA, reducing significantly its baseline. Another example is the strong negative ozone differences during of the summer 2016 which were due to a frequency lock problem of GROMOS. Finally, the large flagged period starting at the end of 2019 marks the beginning of several instrumental issues on SOMORA which were finally solved by the replacement of its spectrometer in September 2020. All these issues have been identified, documented and are flagged accordingly in the new ozone data series. A detailed documentation of the time series can be found together with the data.

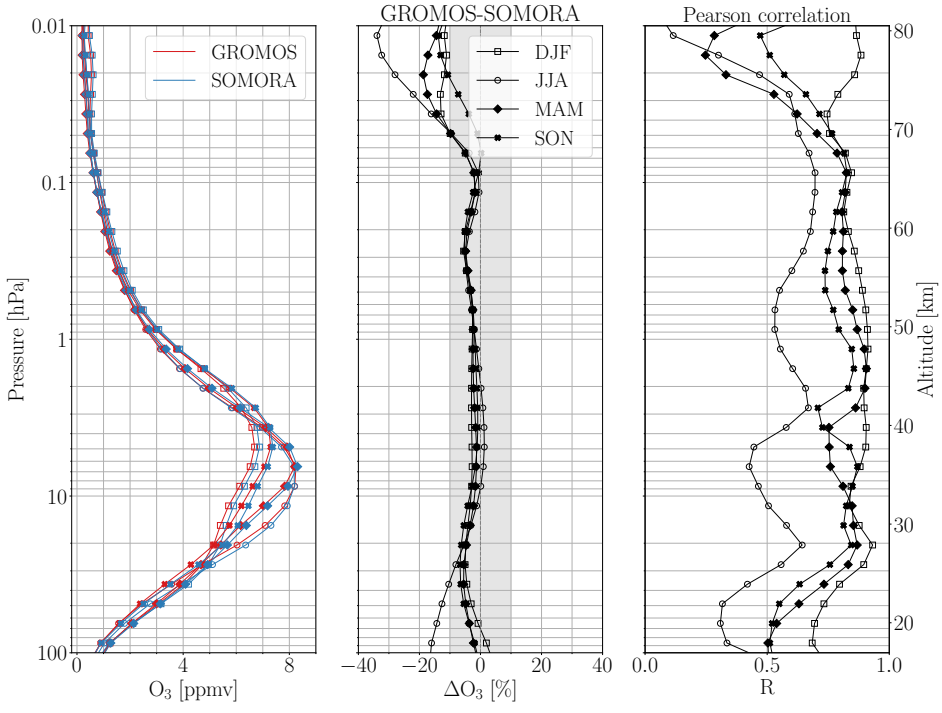

Figure 8. Mean seasonal ozone VMR profiles (left panel), their mean relative differences (middle panel) and correlations (right panel). The shaded area in the middle panels indicates the ± 10 % interval.

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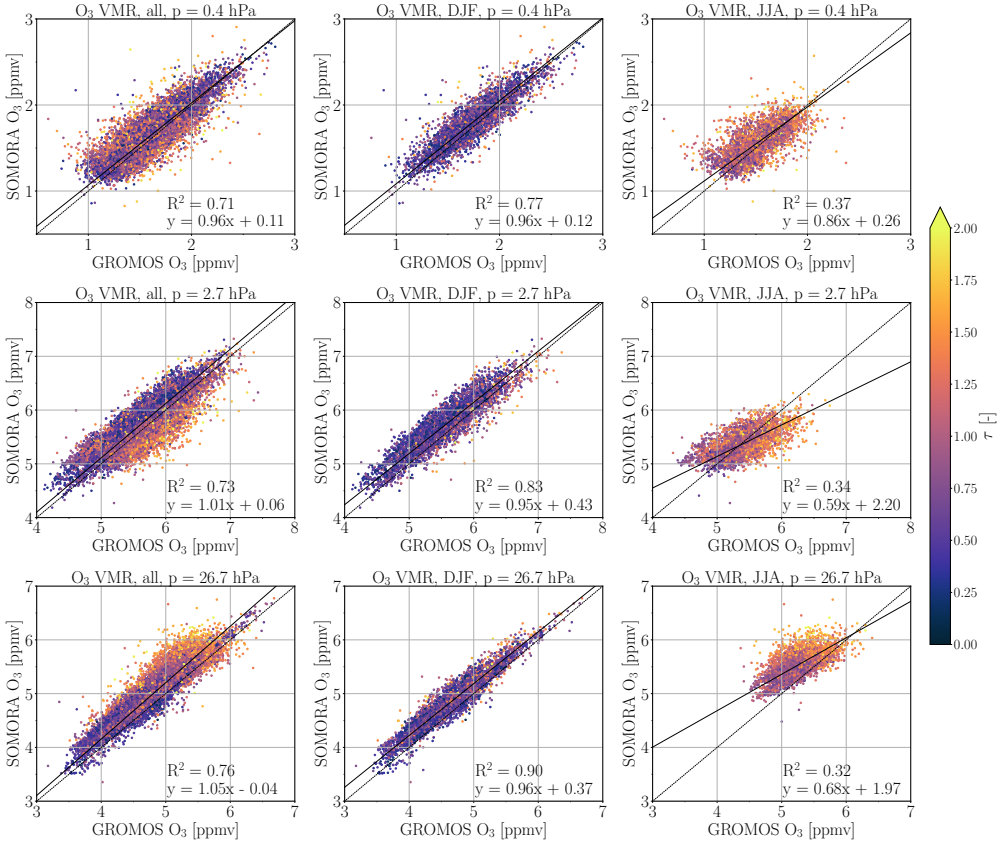

Figure 9. Mean ozone VMR for 3 different levels for the whole series (left panel), the boreal winter season (center) and the boreal summer (right). The 3 pressure levels correspond approximately to the lower ($10 < p < 50$ hPa), upper stratosphere ($1 < p < 5$ hPa) and lower mesosphere ($0.1 < p < 0.9$ hPa). All data points are color-coded based on the atmospheric opacity (τ) computed at SOMORA measurement time and location. The linear regression coefficients and their coefficient of determination R^2 are indicated on each subplot.

4.2 Comparison with previous retrievals

Computing trends for GROMOS and SOMORA is out of scope of this contribution but, still, we would like to provide some first elements of answer to whether this harmonization can help solving the discrepancies found previously between both instruments (Bernet et al., 2019; Petropavlovskikh et al., 2019). Therefore, we compare our new harmonized ozone time series with the previous data version of GROMOS and SOMORA.

Figure 10 shows the daily relative differences between the new harmonized series (upper panel) and the previous retrievals (lower panel) from 2010 to 2021. It highlights the significant improvements introduced by the harmonization process in most of the pressure range covered by the radiometers. Among other changes, it corrects the strong positive ozone bias from GROMOS seen in the mesosphere and reduces the stratospheric ozone difference clearly visible in many years of the previous data series

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at $\sim$ 10 hPa. The previous series were also showing a quite strong seasonal signal which gives confidence that the previously discussed seasonal patterns are not artifacts introduced by the new data processing but are likely linked with the instruments themselves. Whereas the harmonized retrievals improve most of the time period considered, it seems that the problems seen on SOMORA in 2020 are less well treated in the new processing. Indeed, in the previous processing the sine baseline periods were adapted daily during this periods whereas the new processing only considered fixed periods. It indicates that the instrumental baselines on SOMORA varied significantly during this period and highlights the need to treat it carefully for further analysis.

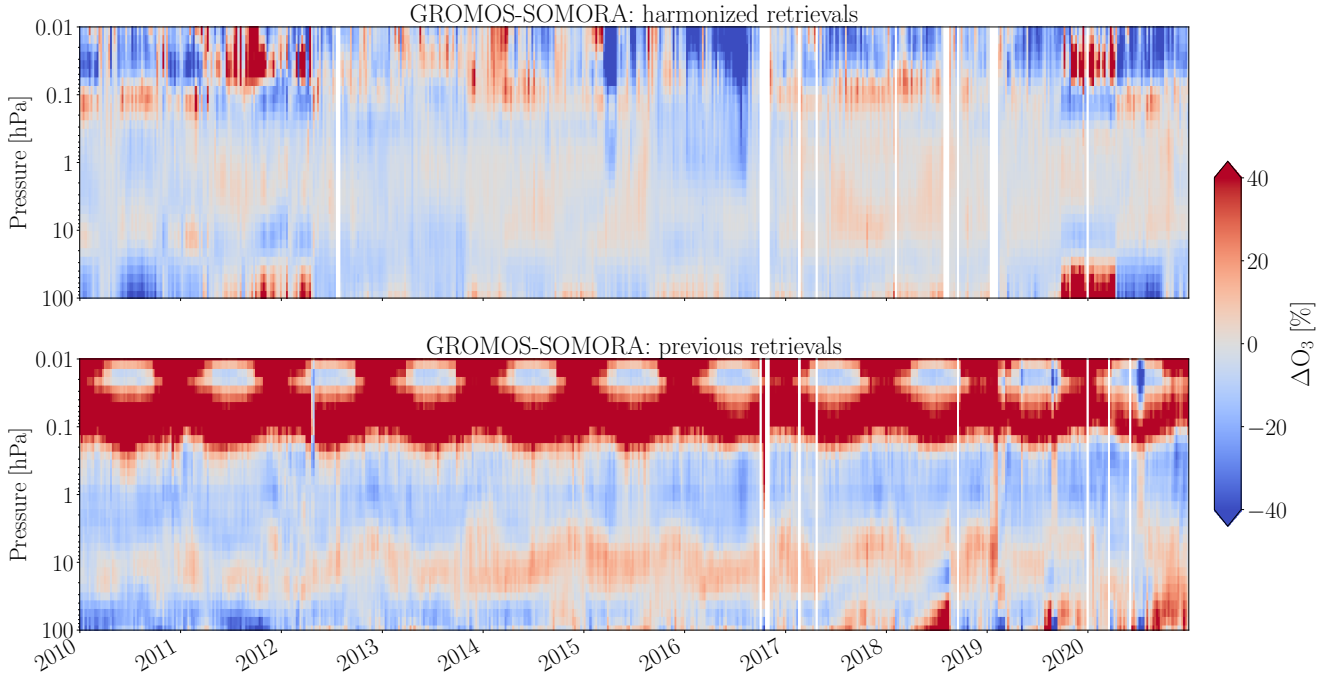

Figure 10. Weekly ozone relative difference between the new (upper panel) and the previous (lower panel) GROMOS and SOMORA series.

From Fig. 10, it is clear that the harmonized processing reduces significantly the differences between GROMOS and SOMORA ozone time series. However, the question remains if it can solve the discrepancies found between their respective trends. Of course, the full reprocessing of the series (including the decade 2000-2010) would be needed to fully answer this question but we present here some preliminary results showing the temporal evolution of the ozone differences between both series in Fig. 11. It shows the weekly mean differences between GROMOS and SOMORA with the previous and new retrieval algorithms in three pressure ranges. Ideally, these differences should be constant to guarantee similar trends from both instruments. Simple linear regressions have been performed on these data and indicates smaller drift intensities at all pressure ranges from the new data processing which are significant above 10 hPa.

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As a consequence, the future trends to be derived for this decade from the new series should be in better agreement than with the previous retrievals. However, even with the new series, we still observe a drift between both instruments in the stratosphere which calls for a careful treatment of spurious data periods for the next trends analysis, as done in Bernet et al. (2021).

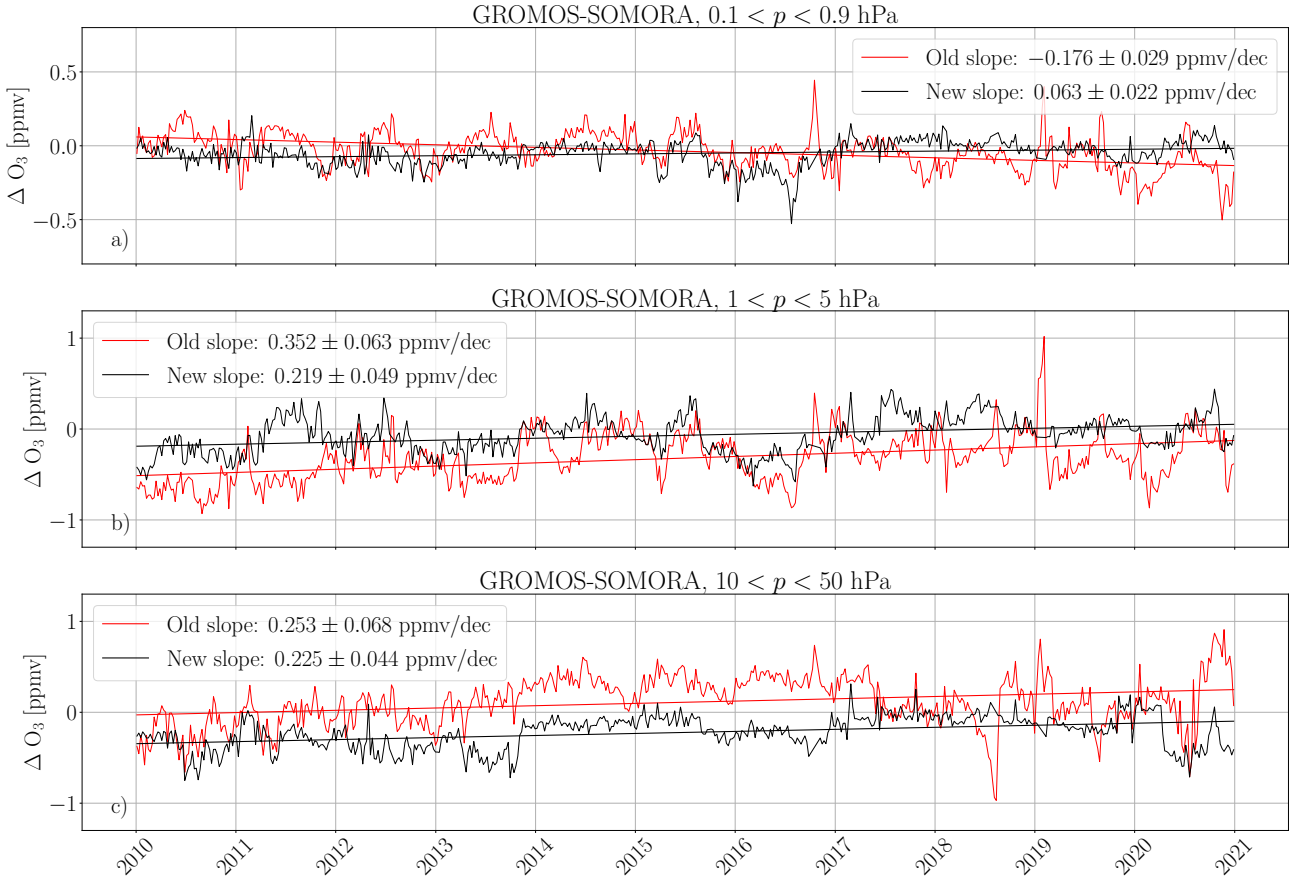

Figure 11. Weekly ozone differences between the previous and the new GROMOS and SOMORA series for the three pressure levels defined in Table 4: a) lower mesosphere, b) upper stratosphere and c) lower stratosphere. A linear fit of the differences is shown as straight line for the previous and the new series. The slope values are indicated with a 95 % confidence interval.

5 Comparison with satellites

Attention was paid to keep GROMOS and SOMORA data processing fully independent. However, they would be both impacted by any bias introduced by the calibration or retrieval algorithms and therefore, we provide further validation by comparing their observations with satellite measurements.

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315 ### 5.1 Aura MLS

As the main validation dataset, we use ozone measurements from the Microwave Limb Sounder (MLS) on the Aura satellite launched in 2004 (Waters et al., 2006). It is operated by the National Aeronautics and Space Administration (NASA) in the frame of the Earth Observing System and has been used extensively for ozone profile validation over many regions and against many other observing systems (e.g., Boyd et al., 2007; Livesey et al., 2008; Hubert et al., 2016).

320 MLS is a passive microwave radiometer observing the ozone emission line around $240\ \mathrm{GHz}$ in a limb sounding geometry. It follows a sun-synchronous orbit which results in two overpasses per day around 1 AM and 1 PM over central Europe. In this work, we have used the latest level 2 ozone retrievals v5 and the recommended data screening described in Livesey et al. (2022). It results in ozone VMR profiles between 261 to $0.001\ \mathrm{hPa}$ with a typical vertical resolution ranging from $\sim 2.5\ \mathrm{km}$ in the lower stratosphere increasing to $\sim 5.5\ \mathrm{km}$ at the mesopause with an accuracy of $5-10\ \%$ in the stratosphere increasing 325 up to $100\ \%$ at $0.01\ \mathrm{hPa}$.

For the following comparisons, we extracted collocated MLS observations to GROMOS and SOMORA. As spatial coincidence criteria, we use $\pm 3.6°$ in latitude and $\pm 10.5°$ in longitude from Bern, an area corresponding approximately to Central Europe. As temporal criteria, we averaged the MWR and the MLS profiles within 3 hours time windows and keep only the time windows where both MLS and the MWR have profiles with sufficient data quality.

330 The MLS vertical resolution of ozone retrievals is much lower than the one from the MWRs. It means that the MWRs will essentially observe a smoothed vertical profile compared to the MLS observations. Therefore, the higher resolved MLS profiles are convolved with the MWR averaging kernels for the comparisons (see Connor et al., 1994; Tsou et al., 1995). This AVKs smoothing also enables to remove the influence of the a priori and follows Eq. 1:

$$x_c = x_a + A(x - x_a) \tag{1}$$

335 where x is the higher resolution profile (MLS), x_a is the a priori profile from the MWR retrievals, A are the averaging kernels and x_c is the resulting convolved profile.

5.2 SBUV/2

In addition to MLS, we also use the latest release of the Solar Backscatter Ultraviolet Radiometer (SBUV/2) Merged Ozone Dataset (MOD) (Frith et al., 2020; Ziemke et al., 2021). This dataset provides daily overpasses over many ground-based ozone 340 measurement stations, including Payerne in Switzerland. It provides stratospheric ozone VMR profiles from 50 to $0.5\ \mathrm{hPa}$ merged according to the new MOD v2 Release 1 derived from SBUV and adjusted for the diurnal cycles to an equivalent local measurement time of 1:30 PM. The vertical resolution from the SBUV retrievals is $\sim 6-7\ \mathrm{km}$ in the middle and upper stratosphere (McPeters et al., 2013; Bhartia et al., 2013) which is closer to the vertical resolution of GROMOS and SOMORA in this region. For this reason, contrary to MLS, we do not apply any AVKs smoothing to the SBUV measurements for the 345 following comparisons.

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5.3 Time series

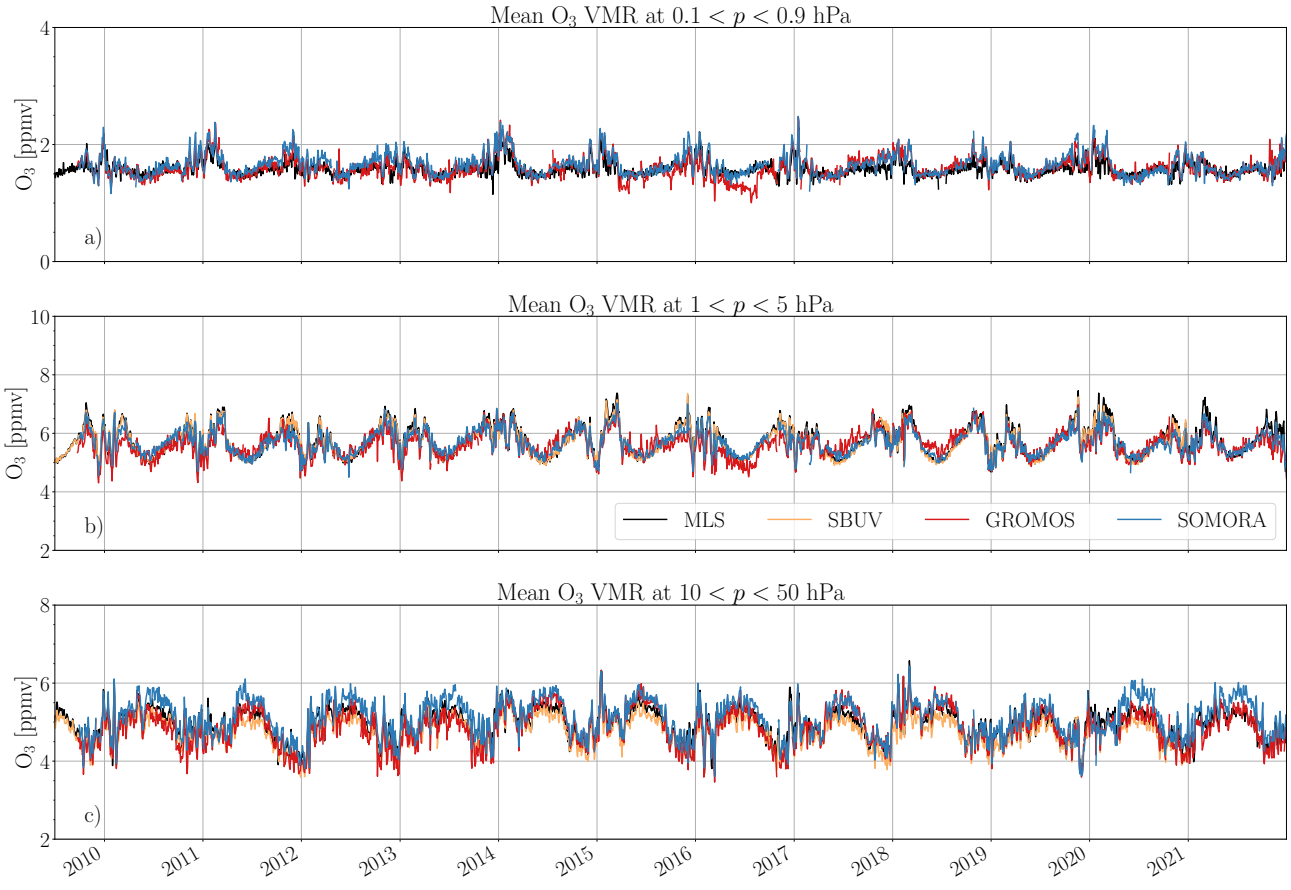

Figure 12. Two days averaged ozone VMR from MLS, SBUV, GROMOS and SOMORA at three pressure intervals: a) lower mesosphere, b) upper stratosphere and c) lower stratosphere. The SBUV dataset extends only up to 0.5 hPa and is therefore not shown in a).

Figure 12 shows daily mean GROMOS and SOMORA time series together with SBUV and MLS measurements on three pressure ranges corresponding to the lower, upper stratosphere and lower mesosphere. It shows the consistencies of GROMOS and SOMORA time series and highlights the good agreement of both MWRs with both satellite datasets during the last decade. As these time series are already averaged on given pressure ranges, we did not apply any AVKs smoothing on the MLS data at this stage. Also it is important to keep in mind that the SBUV daily dataset is adjusted to daytime (1.30 PM) whereas both MLS and the MWRs have both day and night time measurements.

In the stratosphere, clear seasonal pattern are well captured by all dataset and the higher winter ozone variability is clearly visible at all pressure levels. On time scale of a few days, we can see that all four dataset are able to capture well the short

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term ozone variations not only in the stratosphere, but also in the mesosphere where these variations become relatively small compared to the amplitude of the ozone diurnal cycle.

We can see a slight bias of the SOMORA data series in the lower stratosphere. It is especially visible before 2014 and after 2019 as have been mentioned previously. This plot also helps to identify some remaining spurious time periods in the new harmonized series (e.g. GROMOS data in Summer 2016). From a qualitative point of view, we do not observe large drifts from any of the dataset with respect to the others. More work will be needed to confirm the stability from both MWRs, but it gives some confidence that both instruments can be used for trends analysis in the decade 2010-2020.

5.4 Profile comparisons

As quantitative validation, we show seasonal comparisons of MWRs profiles with the satellite datasets. In the following, we mostly focus on the MLS time series because it covers the same altitude range as the MWRs and because SBUV only provides daytime measurements. For the period between 2009 and 2021, we obtain more than 7100 collocated profiles between MLS and each MWR, giving approximately 1700 profiles per meteorological season. Figure 13 and 14 show comparisons between winter (resp. summer) ozone profiles measured by GROMOS, SOMORA, SBUV and MLS. Both figures show the mean seasonal ozone profile from each dataset as well as the relative differences between MLS and the MWRs with and without AVK convolution. The comparisons for spring and autumn are shown in Appendix B (Fig. B1 and Fig. B2).

Both GROMOS and SOMORA show very good agreement with MLS at most seasons and altitudes. Mean seasonal relative differences between both instruments and collocated MLS profiles are within $10\ \%$ in the stratosphere and lower mesosphere (up to $\sim 60\ \mathrm{km}$), corresponding to the expected uncertainties of the MWRs. Above in the mesosphere, the relative differences between the MWRs and MLS grows rapidly and shows some oscillations. For most of the higher mesosphere, the mean seasonal relative differences stay below $50\ \%$ for both instruments but given the errors reported for the MWRs and MLS at these altitudes, we will focus our discussion on the region from ~ 20 to $60\ \mathrm{km}$. The relative differences with SBUV (not shown) are very similar to those with MLS and are below $10\ \%$ in the whole stratosphere for both instruments.

Figure 14 reveals again the summer bias mentioned previously. Taking MLS as a reference, this plot indicates that the summer bias in the lower stratosphere is the result of an overestimation of ozone by SOMORA during this season. The reason for this could be a seasonal change in the instrumental baselines that is not taken into account in the retrieval. For both instruments still, the differences with the convolved MLS profiles are smaller in autumn and winter than in spring and summer when the absorption by the troposphere is stronger.

Moreira et al. (2017) compared the previous GROMOS retrieval dataset to MLS between 2009 and 2016. Similar agreement was found in the middle stratosphere however quickly degrading at lower and higher altitudes. This is in accordance with the results shown in Fig. 10 and confirms the improvement brought by the new data processing. SOMORA showed similar agreement with MLS in the range 25 to $0.1\ \mathrm{hPa}$ between 2004 and 2015 (Maillard Barras et al., 2020). Below $25\ \mathrm{hPa}$, SOMORA showed a positive bias compared to other datasets which gives confidence that this bias is not related to the new data processing.

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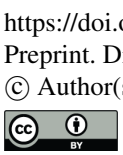

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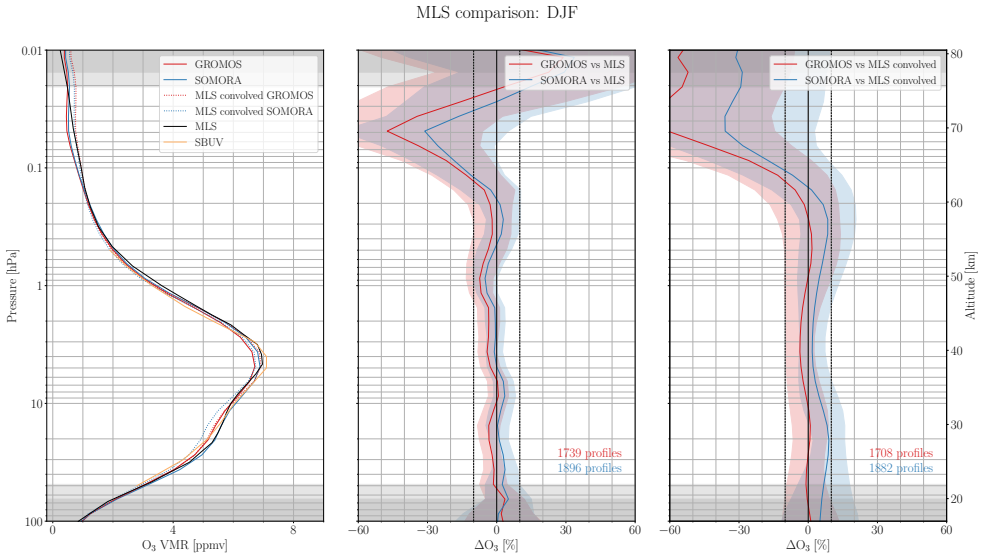

Figure 13. Seasonal comparison with MLS and SBUV during winter months (December, January and February). The middle panel shows the relative differences with MLS whereas the right panel shows the relative differences with the convolved MLS profiles. The color shaded areas show the standard deviation of the differences with MLS and the grey shading indicates the limits where the a priori contribution exceeds 20 %. The dashed vertical lines indicate the ±10 % interval.

Similar comparisons between MWR and MLS has been performed at various locations (e.g. Boyd et al., 2007; Palm et al., 2010; Ryan et al., 2016). Compared to others studies, the Swiss MWRs are in close agreement with the simultaneous MLS profiles. This is confirmed by the mean ozone VMR relative differences between MWR and MLS given in Table 5 for the three pressure ranges. Averaged over these pressure ranges, the differences between MLS and the MWRs are less than 5 % in the as defined stratosphere and lower mesosphere.

Overall, SOMORA and GROMOS profiles are in better accordance with the non-convolved MLS than with the convolved MLS profiles. This can be seen for both instruments and at the three pressure ranges from the seasonal plots and in Table 5.

Table 5. Mean relative VMR differences ((MWR-MLS)/MWR) between MWRs and MLS at three pressure ranges, with and without AVK convolution. In parenthesis, we show the standard deviations of the VMR relative differences in each pressure range.

Pressure range [hPa]	$\Delta O_{3,\textbf{GROMOS}}$ [%]	$\Delta O_{3,\textbf{GROMOS, convolved}}$ [%]	$\Delta O_{3,\textbf{SOMORA}}$ [%]	$\Delta O_{3,\textbf{SOMORA, convolved}}$ [%]
$0.9 - 0.1$	-4.1 (3.2)	-0.9 (4.0)	-0.9 (4.0)	$+5.6$ (4.3)
$5 - 1$	-1.7 (1.1)	$+2.5$ (0.1)	-0.3 (0.8)	$+5$ (0.8)
$50 - 10$	-0.7 (1.0)	$+2.0$ (1.4)	$+4.2$ (1.2)	$+11.6$ (1.4)

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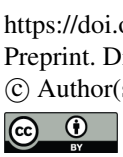

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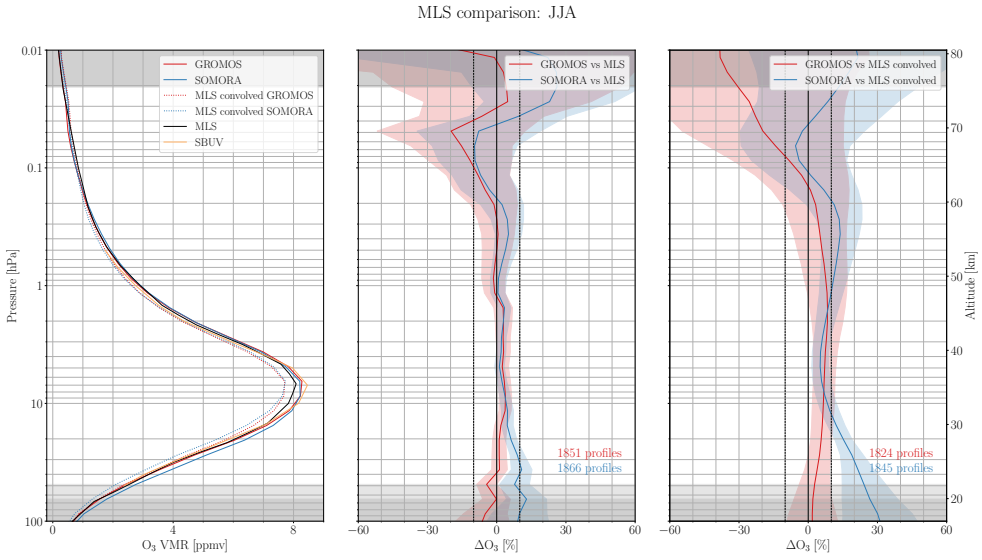

Figure 14. Same as Fig. 13 but for summer (June, July and August)

6 Conclusions

New harmonized data series from two Swiss ozone ground-based microwave radiometers are now available from 2009 until 2021. The reprocessing provides a full harmonization at all levels, from the calibration of the raw data to the retrieval of the ozone profiles. It includes the data inputs and outputs, the flagging systematic, the output temporal resolution and the retrieval grids. The harmonization makes the comparison and the identification of biases easier than in the past. It improves significantly the agreement between both instruments on this time period and reduces the long-term drift of their differences. It should help to resolve the discrepancies previously found in the trend estimates derived from these two time series.

However, despite these significant improvements, systematic differences remain between the two instruments. It includes a seasonal bias, mostly visible in the lower stratosphere in summer, as well as a negative ozone bias of GROMOS in the upper mesosphere. Further work is needed to fully understand these systematic biases but they probably both arise from instrumental sources as they were already seen in the previous retrieval versions. In addition, limited anomalous time periods still remain on both instruments but most of their causes are now identified and documented. The new harmonized data series are also compared against two independent and collocated satellite datasets. Both instruments show a good agreement with SBUV and MLS, with mean relative differences below $10\ \%$ in most of the stratosphere and lower mesosphere (up to $\sim 60\mathrm{km}$).

The new retrieval products of ozone profiles at Bern and Payerne are available and will be submitted to NDACC. We also plan to extend the harmonization process to the older observations from these two instruments in order to provide the full harmonized ozone time series since 1994 (GROMOS) and 2000 (SOMORA). The collocation of two harmonized time series with high temporal resolution also opens the way to unique short-term ozone variations analysis.

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Appendix A: Uncertainty budget at high atmospheric opacities

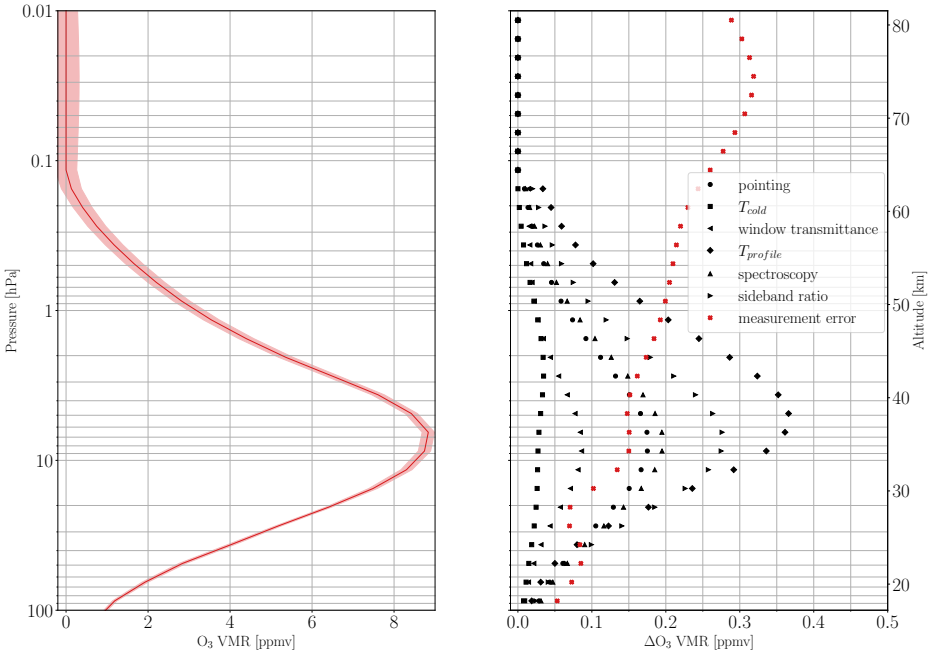

Figure A1. Same as Fig. 5 but for GROMOS in the high opacity case.

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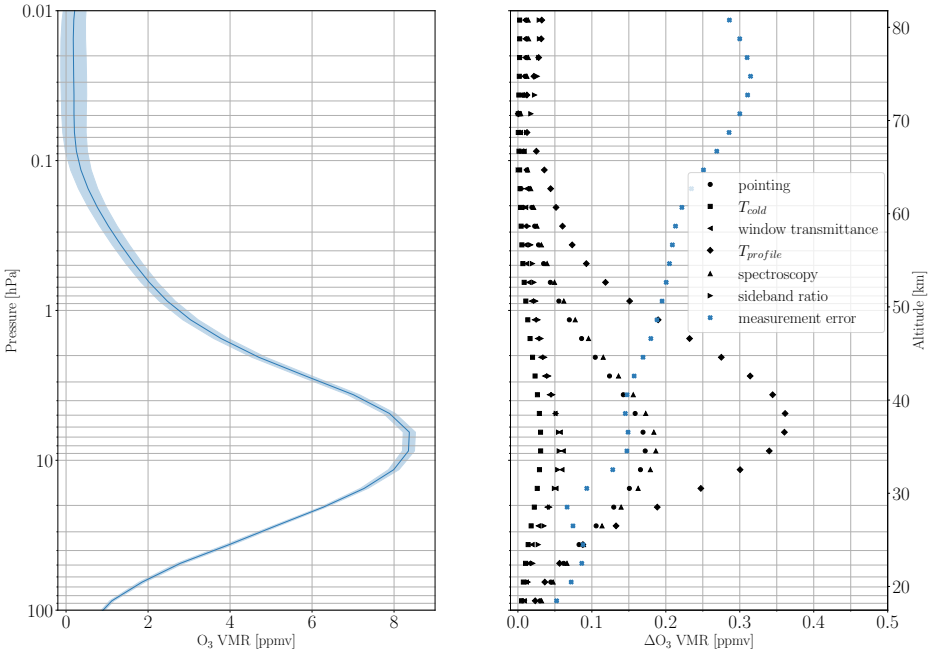

Figure A2. Same as Fig. 5 but for SOMORA in the high opacity case.

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Appendix B: Seasonal comparison with MLS and SBUV

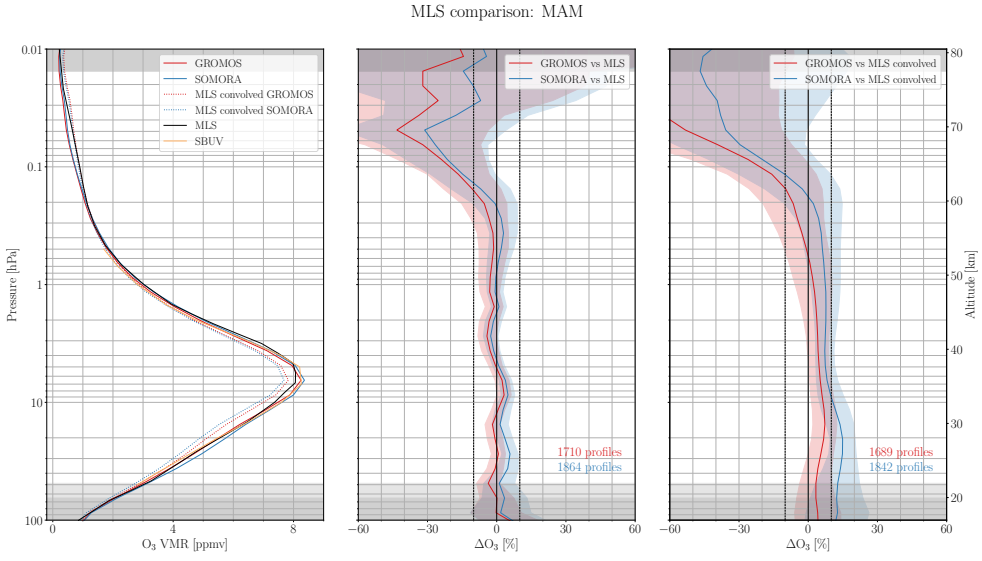

Figure B1. Same as Fig. 13 but for spring (March, April and May)

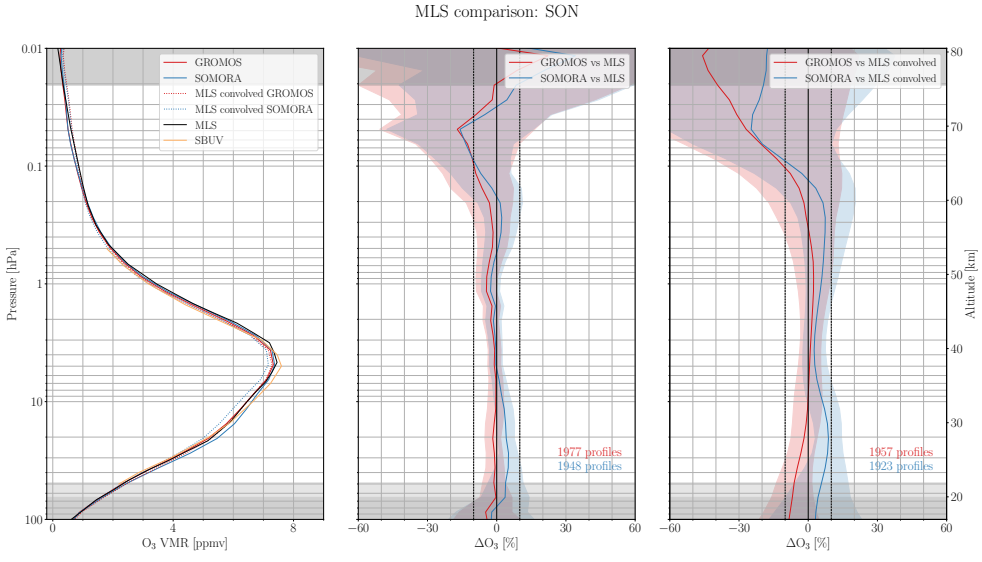

Figure B2. Same as Fig. 13 but for autumn (September, October and November)

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Code and data availability. The GROMOS and SOMORA level 2 data will be available from the Bern Open Repository and Information System (University of Bern) in the form of yearly netCDF files. The new harmonized calibration and retrieval routines are freely available at https://doi.org/10.5281/zenodo.6799357. The analysis code reproducing all the results presented in this manuscript can be found at https://doi.org/10.5281/zenodo.6801529. MLS v5 data are available from the NASA Goddard Space Flight Center Earth Sciences Data and Information Services Center (GES DISC): https://disc.gsfc.nasa.gov/. The SBUV MOD dataset are available at https://acd-ext.gsfc.nasa.gov/Data_services/merged/index.html)

Author contributions. ES performed the harmonization project, carried out the data analysis and prepared the manuscript. EM provided the SOMORA data and helped with the data analysis. KH provided the GROMOS data and helped with the data analysis. AH conceived the project and provided advice on the data analysis. AM conceived the project and helped with the data analysis. All of the authors discussed the scientific findings and provided valuable feedback for the manuscript editing.

Competing interests. The authors declare that they have no conflict of interest.

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