Evaluation of the quality of a UAV-based eddy covariance system for measurements of wind field and turbulent flux

- Yibo Sun^{1,2,3}, Shaomin Liu⁴, Xinwen Lin⁵, Ziwei Xu⁴, Bing Geng⁶, Bo Liu^{1,2,3}, Shengnan Ji^{1,2,3}, Junping Jing⁷, Zhiping Zhu^{8,9}, Bilige Sude^{1,2,3}, Zhanjun Quan^{1,2,3}
- 5 ¹State Key Laboratory of Environmental Criteria and Risk Assessment, Chinese Research Academy of Environmental 6 Sciences, Beijing 100012, China.
- 7 ²Institute of Ecology, Chinese Research Academy of Environmental Sciences, Beijing 100012, China.
- 8 ³State Environmental Protection Key Laboratory of Ecological Regional Processes and Functions Assessment, Beijing 9 100012, China.
- 4State Key Laboratory of Earth Surface Processes and Resource Ecology, Faculty of Geographical Science, Beijing Normal
 University, Beijing 100875, China.
- 12 ⁵Collage of Geography and Environment Science, Zhejiang Normal University, Zhejiang 321004, China.
- 13 ⁶Beijing Academy of Social Sciences, <u>Beijing</u> 100101, <u>BeijingChina</u>.
- ⁷National Ocean Technology Center, Tianjin 300112, China.

18 19

20

2122

23

24

25

2627

28

29

30

31

32

33 34

- 15 ⁸Kunming General Survey of Natural Resources Center, China Geological Survey, Kunming 650111, China.
- ⁹Technology Innovation Center for Natural Ecosystem Carbon Sink, Ministry of Natural Resources, Kunming 650100, China.
- 17 Correspondence to: Yibo Sun (sun.yibo@ craes.org.cn) and Shaoming Liu (smliu@bnu.edu.cn)

Abstract. Instrumentation packages of eddy covariance (EC) have been developed for a small unmanned aerial vehicle (UAV) to measure the turbulent fluxes of latent heat (LE), sensible heat (H), and CO_2 (Fe)(Fc) in the atmospheric boundary layer. This study evaluates the measurement performance of this UAV-based EC system. First, the precision (1σ) of georeferenced wind measurement was estimated at 0.07 m s⁻¹. Then, the effect of calibration parameter and aerodynamic characteristics of the UAV on the quality of the measured wind was examined by conducting a set of calibration flights. The results shown that the calibration improved the quality of measured wind field, and the influence of upwash and leverage effect can be ignored in the wind measurement. Third, for the measurement of turbulent fluxes, the measurement error caused by instrumental noise was estimated at 0.03 µmol m⁻² s⁻¹ for Fc, 0.02 W m⁻² for H, and 0.08 W m⁻² for LE. Fourth, the precision (1 σ) of the measurements was estimated at 0.04 m s⁻¹ for wind velocity, 0.08 umol m⁻² s for Fe, 1.61 W m⁻² for H₂ 0.15 W m⁻² for LE, and 0.02 m s⁻¹ for friction velocity (u*). Second, the effect of calibration parameter and aerodynamic characteristics of the UAV on the quality of the measured wind was examined by conducting a set of calibration flights. The results shown that the calibration improved the quality of measured wind field, and the influence of upwash and leverage effect can be ignored in the wind measurement. Third, data from the standard operational flights are used to assess the influence of resonance on the measurements and to test the sensitivity of the measurement under the variation ($\pm 30 \%$) of the calibration parameters around their optimum values, of the system by adding an error of ±30 % to their calibrated value. Results shown that the effect of resonance mainly affect the measurement of CO₂ (\sim 5 %). The pitch offset angle (ε_{θ}) significantly affected the measurement of evertical wind (~30 %) and turbulent fluxesH (~1525 %). The heading offset

angle (ε_{ψ}) only mainly affected the measurement of horizontal wind (~15%), and other calibration parameters had no significant effect on the measurements. The results lend confidence to use the UAV-based EC system, and suggest future directions for optimization and development of the next generation system.

1 Introduction

38

56

57

58

59

60

61

62

63

64

65

66

- 39 In environmental, hydrological and climate change sciences, the measurement of surface fluxes at the regional scale (level of 40 several to tens of kilometerskilometers level) has attracted great interest despite often being considered a gordian knot 41 (Mayer et al., 2022; Chandra et al., 2022). Process-based or remote sensing (RS)-based models are often used to estimate 42 land surface fluxes of matter and energy at continental to global scales with typical spatial resolution from 1-10 km (Hu and 43 Jia, 2015; Mohan et al., 2020; Liu et al., 1999). However, observational data, especially at similar scales to models' 44 estimates, are-is often lacking, which presents a significant challenge for the validation and evaluation of the surface flux 45 products from these models' estimates (Li et al., 2018; Li et al., 2017). On the ground, in the past decades, extensive ground 46 eddy-covariance (EC) flux sites with their composed networks and optical-microwave scintillometer (OMS) sites have been 47 built to provide temporally continuous monitoring of surface fluxes at local (hundreds of meters around the measurement site 48 of ground EC) and path (a distance of a few hundred meters to near 10 kilometers between transmitter and receiver terminal 49 of OMS) scales (Yang et al., 2017; Liu et al., 2018; Zhang et al., 2021; Zheng et al., 2023). Generally speaking, Ffluxes from 50 ground measurements need to be scaled up to kilometers-scale to provide comparable spatial surface "relative-truth" flux 51 data for the process- or RS-based models at larger spatial scales (Liu et al., 2016). However, the spatial density of these flux 52 measurements sites is still low compared to the complex variability heterogeneity of surface fluxes, which means that major 53 scaling bias may exist in the upscaled flux data (Wang et al., 2016; Li et al., 2021). Therefore, regional-scaled flux 54 measurement techniques need to are needed to be developed to complement the ground- and models-based approaches 55 (Vellinga et al., 2010).
 - The aA ircraft-based EC flux measurement method, which has been developed for turbulence measurements for more than 40 years (Lenschow et al., 1980; Desjardins et al., 1982), is considered as the optimum method to measure turbulent flux at regional scale (several hundred square kilometers), thus bridging the scale gap between ground and model-derived methods (Gioli et al., 2004; Garman et al., 2006). To date, several types of aircrafts, including manned or unmanned fixed-wing aircrafts, delta-wing aircrafts, and helicopters, have been used for measurements of turbulent flux by equipping them with the EC sensors to measure three-dimensional (3D) wind, air temperature, and gas concentrations at high-a frequency of 50 Hz (Gioli et al., 2006; Metzger et al., 2012; Thomas et al., 2012; Bange and Roth, 1999). Among them, fixed-wing aircrafts and delta-wing aircrafts are better airborne platforms for EC measurements compared to helicopters due to their tightly coupled structure with the wind sensor and because their flow distortion around the fuselage can be more easily avoided or modeled (Prudden et al., 2018; Garman et al., 2008). A wide range of manned aircrafts has been developed to measure turbulent flux, including single-engine light aircrafts (e.g., Sky Arrow 650, Long-EC, WSMA) (Gioli et al., 2006; Crawford

and Dobosy, 1992; Metzger et al., 2012), twin-engine aircrafts (e.g., Twin Otter, NASA CARAFE) (Desjardins et al., 2016; Wolfe et al., 2018) and larger quad-engine utility aircrafts (e.g., NOAA WP-3D) (Khelif et al., 1999). These airborne flux measurements, in combination with ground measurements, provide an excellent opportunity to produce regional-scaled, spatio-temporal continuous surface flux datasets that can improve our understanding of the interactive processes between the land surface and the atmosphereof land-atmosphere interactions in regional and global change (Chen et al., 1999; Liu et al., 1999; Prueger et al., 2005). However, manned aircrafts are expensive to operate and maintain. Aviation safety and operational regulations require that manned aircrafts must fly above a minimum altitude (400 m above the highest elevation within 25 km on each side of the center line of the air route) and must avoid hazardous conditions such as icing or severe turbulence (Elston et al., 2015). The flow distortion induced by the aircraft itself (from the wings, fuselage, and the propellers) complicates the wind vector measurement from aircraft platform, which means that sophisticated correction procedures should be applied to compensate for the flow distortion effects (Elston et al., 2015; Williams and Marcotte, 2000; Drüe and Heinemann, 2013).

In recent years, interesting in unmanned aerial vehicle (UAV) fixed-wing platforms for atmospheric studies hashave been fast growing, especially because of their lower construction, operation, and maintenance costs compared with manned platforms. High-performance fixed-wing UAVs offer a high payload capacity (5-10 kg) and similar endurance (2-3 h) and operating altitude (up to 3500 m above the sea level) to manned aircrafts, but with much less turbulence disturbance due to their small fuselage size (Reineman et al., 2013). More importantly, the advancements in small, fast, and powerful sensors and microprocessors make it possible to use of UAVs for comprehensive atmospheric measurements (Sun et al., 2021a). Several types of UAVs with different turbulence measurement objectives have been developed and deployed, ranging from small payload capacity (e.g., 140 g SUMO) to medium (e.g., 1.5 kg M²AV, 1.0 kg MASC) and large (e.g., 6.8 kg Manta, 5.6 kg ScanEagle) (Reuder et al., 2016; Båserud et al., 2016; Van Den Kroonenberg et al., 2012; Reineman et al., 2013). A comprehensive overview of the use of these UAVs for turbulence sampling can be found in Elston et al. (2015) and Sun et al. (2021a). For turbulence measurements, the UAVs were equipped with a commercial or custom multi-hole (5- or 9-hole) probe paired with an integrated navigation system (INS) to obtain the wind vector. Small and medium UAVs typically could only measure fast 3D wind vector and air temperature fluctuations for measurements of momentum and sensible heat flux. whereas, large UAVs were equipped with more types (e.g., radiation, imageoptics, or gas concentration) and more accurate sensors for measurement of a larger range of meteorological properties including sensible and latent heat fluxes, CO₂ flux, radiation fluxes as well as surface properties (Reineman et al., 2013; Sun et al., 2021a). UAVs equipped with scientific instruments can be deployed in a variety of application environments and conditions. UAVs offer distinct advantages over manned aircraft in their ability to safely perform measurements and greatly reduce operational costs especially in lowaltitude conditions (below 100 m above the ground level), which are optimal for measuring turbulent flux (Witte et al., 2017). Anderson and Gaston (2013) predict that UAVs will revolutionize the spatial data collection in ecology and meteorology.

The EC method is a well-developed technology for directly measuring vertical turbulent flux (flux of sensible heat, matter latent heat and momentum CO₂) within the atmospheric boundary layers (ABL) (Peltola et al., 2021). It requires accurate time

102 turbulent wind. Each should be measured at sufficient frequency to resolve the flux contribution from small eddies (Vellinga 103 et al., 2013). The measurement of the geo-referenced 3D wind vector, which is the prerequisite for EC measurements, is 104 challenging for airborne platform. Airborne measurement of geo-referenced 3D wind is the vector sum between the aircraft 105 velocity relative to the earth (inertial velocity) and the velocity relative to the air (relative wind vector, or true airspeed). 106 Therefore, accurate measurements of the relative wind as well as the motion and attitude of the platform are essential to 107 accurately measure the geo-referenced wind vector and turbulent flux (Metzger et al., 2011), Garman et al. (2006) estimated the measurement -1σ -precision (1σ) of the vertical wind measurements of a commercial 9-hole turbulence probe (known as 108 109 "Best Air Turbulence Probe", often abbreviated as the "BAT Probe") to be 0.03 m s⁻¹ by combining the precision of the BAT Probe and the integrated navigation device. The BAT Probe is widely used on manned fixed-wing aircrafts, such as Sky 110 111 Arrow 650 ERA (Environmental research aircraft), Beechcraft Duchess, and Diamond DA42, for turbulent flux 112 measurement (Gioli et al., 2006; Garman et al., 2008; Savres et al., 2017). A light delta-wing EC flux measurement aircraft developed by Metzger et al. (2011) reported a 1σ precision of wind measurement of 0.09 m s⁻¹ for horizontal wind and 0.04 113 m s⁻¹ for vertical wind using a specially customized five-hole probe (5HP). On this basis, in combination with a commercial 114 infrared gas analyzer. The 1 σ precision of flux measurement was 0.003 m s⁻¹ for friction velocity, 0.9 W m⁻² for sensible 115 heat flux, and 0.5 W m⁻² for latent heat flux. The smallest resolvable magnitudes for the wind velocity and turbulent flux 116 were estimated from these values by assuming a signal-to-noise ratio of 5:1 (Metzger et al., 2012). The EC flux measurement 117 118 from a UAV platform can now be achieved with a similar reliability to a manned platform. The Manta and ScanEagle UAV-119 based EC measurements developed by Reineman et al. (2013) achieved precise wind measurements (0.05 m s⁻¹ for horizontal and 0.02 m s⁻¹ for vertical wind) using a custom nine-hole probe and a commercial high precision integrated navigation 120 121 system (INS) integrated navigation system (INS), at a lower price and lighter weight than the commercial BAT probe. 122 However, the onboard instrument packages for Manta and ScanEagle UAV are independent of each other in their 123 measurements of turbulent and radiation flux, and the CO₂ flux measurement is lacking. 124 Inspired by these studies. Sun et al. (2021a) used a high-performance fuel-powered vertical take-off and landing (VTOL). 125 fixed-wing platform to integrate the scientific payloads for EC and radiation measurements to obtain a comprehensive 126 measurement of turbulent and radiation flux using an UAV. This UAV-based EC system measured turbulent fluxes 127 including of sensible heat, latent heat, and CO₂, as well as radiation including net radiation and upward- and downward-128

(for ground tower) or spatial (for mobile platform) series of both the transported scalar quantity and the transporting

101

129

130

131

132

133

134

measurement of turbulent and radiation flux using an UAV. This UAV-based EC system measured turbulent fluxes including of sensible heat, latent heat, and CO₂, as well as radiation including net radiation and upward- and downward-looking photosynthetically active radiation (PAR). This system was successfully tested in the Inner Mongolia of China and applied to measure the regional sensible and latent heat fluxes in the Yancheng coastal wetland in Jiangsu, China (Sun et al., 2021a; 2021b). During these field studies, the UAV-based EC measurements achieved a near consistent observational result compared with ground EC measurements (Sun et al., 2021b). However, some shortcomings in the developed UAV-based EC system were also identified. In particular, the noise effects from the engine and propeller were not fully isolated, resulting in high frequency noise in the measured scalars (air temperature, H₂O, and CO₂ concentration). This UAV-based EC system is being continuously improved (in Section 2.1) based on previous field measurements. However, there is no quantitative

evaluation of the measurement eapabilityprecision of the wind field and turbulent flux or as well as of the influence of the resonance noise from the UAV-based EC measurement system operation yet. Previous work using ground EC measurements as a benchmark to assess the measurement performance of the UAV-based EC system has been disputed, due to difference in EC sensors, platforms, measurement height, and source areas (i.e., footprint), as well as the influence of surface heterogeneity, flux divergence, inversion layer and the stochastic nature of turbulence (Sun et al., 2021b; Wolfe et al., 2018; Hannun et al., 2020).

This study attempts to quantitatively evaluate the performance of a-the UAV-based EC system developed by Sun et al. (2021a) in the measurement of wind field and turbulent fluxes. For these purposes, data from two field measurement campaigns, including a set of calibration flights and some standard operation flights, were used in this study. First, the current study investigateds the quality of the measurement of geo-referenced wind vector including measurement error (1 σ) and the improvements for wind measurement after system calibration. Second, using the measured data from standard operation flights, flux measurement error related to instrumental noise was estimated with a method proposed by Billesbach (2011). Errors propagated through the correction terms [i.e., Webb-Pearman-Leuning (WPL) correction for latent heat and CO₂ flux] were also included in our analysis (Webb et al., 1980; Kowalski et al., 2021). Then, the impacts of resonance noise on the measured scalar variance and the flux covariance were also estimated by comparing the real (co)spectra curve with the theoretical reference curve from Massman and Clement (2005). Lastly, the sensitivity of the measured geo-referenced wind vector and turbulent flux to the errors in the calibration parameters (determined by the calibration flight) were assessed by adding an error of ±30 % to their optimum calibration value.

the 1σ measurement error of the geo-referenced wind vector and turbulent flux by propagating the error of each EC sensor along the data process procedure. Then, a set of calibration flights were conducted to assess how the calibration parameters and aerodynamic characteristics of the UAV affect the quality of the wind measurement. The effects of resonance noise on the measured scalar variance and the fluxes were also estimated by comparing the real (co)spectra curve with the theoretical reference curve from Massman and Clement (2005). Lastly, the sensitivity of the measured geo-referenced wind vector and turbulent flux to the calibration parameters (determined by the calibration flight) were assessed by adding an error of ± 30 % to their optimum calibration values.

2 Materials and Methods

2.1 The UAV-based EC system

The UAV platform used for EC measurement is a high-performance, fuel-powered VTOL, fixed-wing UAV, which has minimal requirements for the takeoff location and offers a high payload capacity of up to 10 kg. The UAV has a wing-span of 3.7 m, a fuselage length of 2.85 m, and a maximum take-off weight of 60 kg. The UAV engine is mounted in a pusher configuration, allowing for the turbulence probe to be installed directly on the nose of the UAV, minimizing or eliminating

airflow contamination due to upwash and sidewash generated by the wings (Crawford et al., 1996). Control of the UAV is totally autonomous, and the pilots have the option to enable manual and semi-manual control in emergency conditions. The UAV has a cruise flight speed of 28 to 31 m s⁻¹ with an endurance of almost 3 h, and it has a flight ceiling of up to 3800 m above the sea level. Detailed information on this UAV could be found in Sun et al. (2021a).

166

167

168169

170

171

172

173

174

175

176

177

178

179

180

181

182

183

184

185

186

187

188

189

190

191

192

The fluxflux payloads of the UAV-based EC system include a precision-engineered 5-hole pressure probe (5HP) for measurement of the true airspeed and the attack (α) and sideslip (β) angles of incoming flow relative to the UAV, a dualantenna integrated navigation system (INS) for high accuracy measurement of UAV ground speed and attitude, an open path infrared gas analyzer (IRGA) for recording the atmospheric densities of CO₂ and water vapor, a fast temperature sensor for measurement of the fast temperature fluctuations, and a slow-response temperature probe for providing a mean air temperature reference. The auxiliary payloads include a net radiometer and two photosynthetically active radiation (PAR) radiometers that look upward and downward. The sample rate of the flux payloads is 50 Hz except for the slow-response temperature probe (1 Hz), yielding a turbulence horizontal resolution of approximately 1.2 m at a cruising speed of 30 m s⁻¹. The system was improved according to deficiencies identified after several field measurements with the following adjustments: 1) a laser distance measurement unit was mounted for measuring the distance between the UAV and the ground level, 2) the platinum resistance thermometer was replaced by a thermocouple (Omega T-type COCO-003; \(\phi 0.075 \) mm) for improving the resistance of the high-frequency temperature measurements to vibration noise from the engine, 3) the vibration isolator structure of the IRGA was improved, and 4) the original datalogger (CR1000X, Campbell, USA) was replaced with a lighter one (CR6, Campbell, USA). All the digital and analog signals from the measurement sensors modules are stored and synchronized by the on-board datalogger, and the on-board scientific payloads are designed to be isolated from the electronic components of the UAV to ensure that any problems occurring would not jeopardize the safety of the UAV (Sun et al., 2021a).

In the present study, to estimate the <u>measurement1 σ </u> precision for of the <u>measured</u> geo-referenced wind and turbulent flux, the sensor modules and the <u>ir</u> 1 σ precision of the measurement variables related to EC measurement were used, as presented in Table 1. For the 5HP, the 1 σ measurement precision was acquired from the wind tunnel test after the wind tunnel calibration (Sun et al., 2021a).

Table 1: Summary of tThe sensor modules, measured variables, and their measurement p 1σ precision used to determine the geo-referenced wind velocity and turbulent flux.

Sensor (Module, company, country)	Variables	Precision (1σ)	
GNSS/INS	Roll, Pitch, Heading	0.1°	
(Trimble-BD992-INS, Trimble, USA)	Horizontal velocity	0.007 m s^{-1}	
	Vertical velocity	0.02 m s^{-1}	
5HP	Attack angle	$0.02^{o\#}$	
(Simtec AG-ADP-55, Simtec AG,	Sideslip angle	$0.04^{o\#}$	
Switzerland)	True airspeed	$0.05~\mathrm{m~s^{-1}}^{\#}$	

	Static pressure	1.1 111 α	
	Dynamic pressure	0.003 hPa	
IRGA	CO ₂ density	0.2 mg m^{-3}	
(Campbell-EC150, Campbell, USA)) H ₂ O density	$0.004~{\rm g}~{\rm m}^{-3}$	
Thermistor	Temperature (slow)	0.2 °C	
(BetaTherm 100K6A1IA, Campbell	<u>l, USA</u>)	0.2 °C	
Thermocouple	Temperature (fast)	0.5 °C	
(Omega T-type COCO-003, Omega	<u>, USA</u>)	0.5 C	

Static pressure

1 1 hPa

2.2 Field campaign

2.2.1 In-flight calibration campaign

In order to calibrate the wind measurement component of the UAV-based EC system, Aan in-flight calibration campaign was carried out on 4 September 2022 at the Caofeidian Shoal Harbor (39°2′55″-N, 118°38′48″ E) in the Bohai Sea of northern China. The average water depth of this area is approximately 0-5 m, with a maximum water depth of 22 m. At low tide, a large area of the tidal flat is exposed; while at high tide, only the barrier islands are visible (Xu et al., 2021). The assumptions for calibration flight include 1) low turbulent transport condition (i.e., no disturbance), 2) a constant mean horizontal wind, and 3) mean vertical wind near zero (Drüe and Heinemann, 2013; Vellinga et al., 2013; Van Den Kroonenberg et al., 2008). This allows identical wind components for several consecutive straights in opposite or vertical flight directions. These assumptions are usually well satisfied above the ABL or under stable atmospheric conditions (Drüe and Heinemann, 2013). Over the sea surface, due to its uniform and cool surface property, the turbulence fluctuations are weaker than that over the land surface (Mathez and Smerdon, 2018), making where a more ideal environment to conduct calibration flight.

The optimum atmospheric conditions for the in-flight calibration include 1) no large turbulent transport, 2) a constant mean horizontal wind, and 3) mean vertical wind near zero (Van Den Kroonenberg et al., 2008). These conditions are easier to meet over the sea surface than land, due to the uniform nature of the sea surface. Additionally, the sea-atmosphere interaction is relatively weaker than the land-atmosphere interaction (Mathez and Smerdon, 2018).

The in-flight calibration campaign included three flight maneuvers, including a 'box' maneuver, 'racetrack' maneuver, and 'acceleration-deceleration' maneuver. The trajectory of the calibration flight is shown in Figure 1, with different color corresponding to different flight maneuvers. The calibration flight was executed between 7:28-7:48 a.m. (China Standard Time, CST) to coincide with the ebb tide stage. During this time, the average water depth was approximately 1.1 m, and the average flight altitude was 400 m ($\sigma = \pm 0.78$ m) above the sea level. Considering the conditions of the regional uniform and cool underlying surface and and the stable atmospheric conditions—in the early morning, we assume no disturbance from underlying surface was present during the calibration flight.

[#] Results from the wind tunnel test.

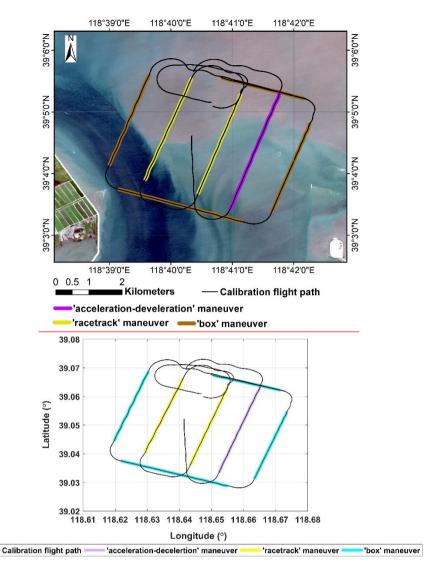


Figure 1. Flight trajectory of the calibration campaign on 4 September 2022 at the Caofeidian Shoal Harbor in the Bohai Sea of northern China. The land surface image is from Sentinel-2A satellite image with true color combination acquired on 1 September 2022.

In this study, the 'box' maneuver is used to determine the mounting misalignment angle in the heading (ϵ_{ψ}) and pitch (ϵ_{θ}) between the 5HP and the center of gravity (CG) of the UAV. The flight path is a box in which the four straight legs are flown at constant cruising speed, flight altitude, and heading for 2 minutes. The 'racetrack' maneuver is used to evaluate the quality of the calibration parameters acquired from the previous 'box' maneuver. The flight path consists of two parallel straight flight tracks connected by two 180° turns. Each straight flight section lasts 2 minutes at constant speed and flight altitude. Lastly, the 'acceleration-deceleration' maneuver is used to check the influence of lift-induced upwash from the wing to the measured attack angle by the 5HP. During this maneuver, the aircraft is kept straight and level at constant pressure

altitude. When beginning this maneuver, the aircraft accelerates to its maximum airspeed (35 m s⁻¹). Then, the airspeed reduces gradually to near its minimum airspeed (25 m s⁻¹) and back up to its maximum airspeed. The pressure-altitude of the aircraft is maintained throughout this maneuver, and the entire maneuver lasts one minute. This maneuver creates a series continuous changed pitch (θ) and attack (α) angle. A relationship between the measured incident flow attack angles (α) by the 5HP and the measured pitch angle by the INS of close to 1:1, indicates that the effect from the fuselage-induced flow distortion on the wind measurements is negligible (Garman et al., 2006).

2.2.2 Standard operation flight campaign

The reliability of the EC measurement from UAV is susceptible to several factors, mainly including instrumental noise, resonance noise, and the quality of the calibration parameter. In order to evaluate the flux measurement error related to instrumental noise, the effects of resonance on the measured scalar and to investigate the sensitivity of the measured georeferenced wind and turbulent flux to uncertainty in the calibration parameter, Previous studies have shown that the measured scalars were affected by the vibration noise from the engine and propeller of the UAV (2021b; Sun et al., 2021a). In order to evaluate the effects of resonance on the measured scalar and to investigate the sensitivity of the measured georeferenced wind and turbulent flux to uncertainty in the calibration parameters, we used data from 7 flights in the Dagang district (38°54′27″ N, 117°48′16″ E) in Tianjin, China between 8 and 16 August 2022. This area is located on the west coast of the Bohai Sea and is a coastal alluvial plain with altitudes between 1-3 m (Chen et al., 2017). The flight path, shown in Figure 2, includes three parallel transect lines of approximately 4 km in length each and at 1-2 km intervals. All flights occurred during the daytime, and were performed in the same trajectory at low altitude about 90 m above sea level. The flight area covered three different underlying surfaces: land, coastal zone, and water surfaces, that can represent typical flux intensity characteristics for different surface conditions.

117°38'0"E

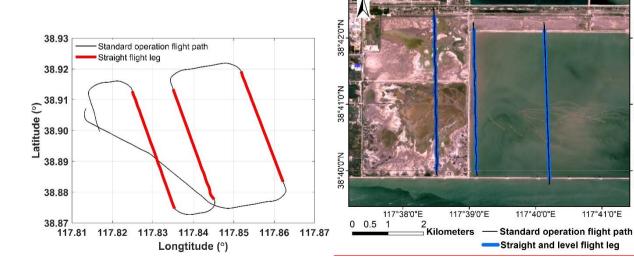
117°39'0"E

117°40'0"E

117°41'0"E

38°42'0"N

38°41'0"N



- Figure 2. Flight trajectory of the standard operation flight campaign, 8-16 August 2022, at Dagang district, Tianjin, China. The
- 252 land surface image is from Sentinel-2A satellite image with true color combination acquired on 27 August 2022.
- During the operation flight campaign, the atmospheric stability conditions changed from stable (Monin-Obukhov stability
- parameter, z/L = 1.93) to very unstable (z/L = -10.28) as measured by the UAV, where <u>z</u> is the flight height above the
- 255 ground level, L is the Obukhov length. The stable condition mostly occurred on flight path located over the sea surface,
- 256 while the unstable condition mostly occurred on flight path located over the land surface. The average net radiation of each
- 257 transect flight line changed from -40 W m⁻² over the sea surface to 626 W m⁻² over the land surface. These flight data
- 258 provide various measurement conditions for us to evaluate the performance of the UAV-based EC system.

2.3 Data processing

259

260

271

- 261 The raw data collected with the on-board datalogger (CR6, Campbell, USA) is subsequently saved in Network Common
- 262 Data Form (netCDF) format. It includes dynamic and static pressure, attack, and sideslip angle of incoming flow; slow (1 Hz)
- 263 and fast (50 Hz) air temperature; mass concentration of H₂O and CO₂; as well as the full navigation data (including 3D
- 264 location, ground speed, angular velocity, and attitude, etc.) of the UAV. The subsequent data processing includes three basic
- processing stages in order to calculate flux data from raw measured data.
- In the first stage, Before processing the raw data into geo-referenced wind vector and turbulent flux, a moving average
- 267 filter was used to detect outliers in each variable. Detected outliers were removed and replaced by values obtained by linear
- 268 interpolation. Outliers tend to be rare. However, if outliers constitute more than 20 % of the data points, the corresponding
- 269 flight data should be discarded. The cleaned raw data was then used to calculate the geo-referenced wind vector, (co)spectra,
- and turbulent fluxes.

In the second stage, geo-referenced 3D wind vector is calculated. 2.3.1 Wind measurements

- The full form of the equations of motion for calculating the geo-referenced wind vector by theour UAV-based EC system
- 273 is described in detailed in Supplement Part A. From the aircraft platform, geo-referenced turbulent wind vector is measured
- 274 in two independent reference coordinate systems: the relative true airspeed (\hat{U}_a) measurement in the aircraft coordinate
- 275 system and the ground speed of the aircraft (U_n) in the geo-referenced coordinate system. The geo-referenced wind (U) is the
- 276 vector sum of the relative true airspeed (\widehat{U}_a) , the UAV's motion (U_n) and the tangential velocity due to the rotational motion
- 277 of the aircraft ("lever arm" effect), which is described in Eq. (S2). In this stage, the acquired calibration parameters (ϵ_{th} and
- 278 ϵ_{θ}) from the calibration flight are substituted into the Eq. (S8) to correct the mounting angle offset errors between the 5HP
- and the CG of the UAV. The final equations for geo-referenced wind vector calculation (Eqs. S15 to S17) revealed that the
- 280 lever arm effects due to the spatial separation between the tip of the wind probe and the CG of UAV can influence the wind
- 281 measurements.

The relative wind vector (\widehat{U}_a) measured by the aircraft is susceptible to flow distortion because the airplane must distort the flow to generate lift and thrust, and the aircraft's propellers, fuselage, and wings are the main sources of flow distortion as flow barriers (Metzger et al., 2011). For fixed wing aircrafts, the wind probe mounted on the nose of the UAV and extended as far forward of the fuselage as possible could avoid significant influence from flow distortion from the fuselage and propellers. Nevertheless, the wind measurement is still subject to lift induced upwash due to the wings distorting the flow to generate lift and thrust. The influence of upwash decreases with linear distance from the wing and appears as an additional vertical component of airflow ahead of the wind probe (Crawford et al., 1996; Garman et al., 2008). The magnitude of upwash influence generally increases with airplane size and airspeed, typically ranging from 0.5 to 2.5 m s⁻¹ as reported by the manned fixed wing aircraft (Garman et al., 2008). Therefore, for EC measurements by manned fixed wing aircrafts, the upwash effects must be corrected for wind measurements (Garman et al., 2008; Kalogiros and Wang, 2002). However, wind measurements using a multi-hole probe on the UAV seldom need this correction due to the fuselage size and because the airspeed is very low compared to a manned aircraft. This is considered in the equations for relative wind calculation (in Supplement) used in this study as well. The 'acceleration deceleration' flight maneuver (Section 2.2.1) was used to assess whether the lift induced upwash could be safety ignored by the UAV-based EC system.

The lever arm effects due to the spatial separation between the tip of the wind probe and the CG of UAV can influence the wind measurements (Eqs. S15 to S17). Typically, the separation distance (L) is small, and the influence of the lever arm effects can be ignored when the L is less than about 10 m (Lenschow, 1986). In the current UAV-based EC system, the displacements of the 5HP tip with respect to the CG of the UAV along the three axes of UAB body coordinate are: $x^b = 1.459 \text{ m}$, $y^b = 0 \text{ m}$, and $x^b = 0.173 \text{ m}$ (in Supplement Part A). Therefore, in practice, the influence of leverage effects in geo-referenced wind calculation was also ignored in this study. ; thus, the influence of leverage effects in geo-referenced wind calculation was also ignored. This was confirmed by assessing the difference in the geo-referenced wind vectors with and without the leverage effect correction term (in Section 3.1).

In the final stage, 2.3.2 Spectra and turbulent flux calculation

Unlike traditional ground-based EC measurements, those recorded on aircraft platforms are subject to several simultaneous motions, including flow distortion around the fuselage and resonance from the rotation of the engine and propeller. Spectral analysis is an effective way to assess if and to what extent the UAV's motion influences the EC measurements. For this reason, the fast Fourier transform (FFT) method was used to calculate the spectra and co-spectra of the measured turbulent variables. Before calculating the turbulence (co)spectra, condition of the raw turbulence data was performed, including a linear detrend and tapering using the Hamming window to reduce the spectral leakage (sharp edge) according to Kaimal et al. (1989).

Bbased on the EC technology and spatial averaging, the turbulent flux is calculated using the covariances of vertical wind (w) with horizontal wind components (u, v) for vertical flux of momentum (τ) , with virtual potential ir temperature $(\underline{T}_{\underline{a}}\theta_{\underline{v}})$ for sensible heat flux (H), with water vapor density (q) for latent heat flux (LE), and with CO₂ density (c) for CO₂ flux (F_c) .

and with the necessary correction (Webb et al., 1980). The time lag due to the separation between the 5HP tip, the adjacent temperature probe, and the open-path gas analysis did not need to be corrected because the time delay was less than 1 second at the cruise airspeed of 30 m s⁻¹ and sensor separation less than 20 cm. Only the measurement data from the straight-line portion of the flight path can be used in flux calculation. Detailed calculation procedure and formulas of H, LE, and E_c used by the present UAV-based EC system are provided in Supplement Part B, including spatially averaging, coordinate rotation, and necessary correction (i.e., WPL correction for LE and E_c).

Detailed information about the EC method and airborne EC calculation can be found in other studies (Aubinet et al., 2012;

Vellinga et al., 2013; Gioli et al., 2006).

One important aspect for airborne EC measurement is the definition of a proper spatial averaging length to calculate turbulent flux (Sun et al., 2018). Such spatial averaging length depends on the flying altitude, surface characteristics, and atmospheric stability, and could be determined using Ogive analysis (Gioli et al., 2004; Kirby et al., 2008). In this study, the objective is not to quantify the actual flux exchange between the surface and the atmosphere, but rather to assess the sensitivity of the calculated turbulent flux to external parameters. Therefore, the entire measurement measured data of each straight and level flight leg (each with length about 4 km) from the standard operational flight campaign was used to calculate turbulent flux, regardless of the uncertainty of the flux results in fluxes associated with spatial averaging.

2.4 Evaluation strategyscheme

2.4.1 Propagation of sensor errorsWind measurement evaluation

The EC technique relies upon the precise measurement of the geo-referenced wind vector and atmospheric scalars The key to successful aircraft EC measurements lies in the translation of accurately measured, aircraft-orientated, wind vector to geo-referenced orthogonal wind vector (Thomas et al., 2012). Measured from an aircraft, dDetermining the geo-referenced wind vector requires a sequence of thermodynamic and trigonometric equations, and the calculation of turbulent flux also needs a series of data operations and corrections to the turbulent variables (Metzger et al., 2012). The EC data processing procedure and, these equations propagate various sources of error to the measured geo-referenced wind vector and turbulent flux. To estimate the measurement errors in the geo-referenced wind vector, we used the linearized Taylor series expansions of Eqs. (S15) to (S17) derived by Enriquez and Friehe (1995) (in Supplement Part A) to determine the sensitivities of each of the geo-referenced wind vector components with respect to the relevant variables. Then, these sensitivity terms can be combined to compute the overall measurement error (1σ) in the geo-referenced 3D wind vector (Eqs. S21 to S23 in Supplement Part A).

To estimate the errors in the geo-referenced wind vector and turbulent flux, we performed a Monte Carlo simulation of the data processing procedure.

The Monte Carlo simulation method consists of repeated calculation of the target quantity, each time varying the input data randomly within their stated limits of precision. The distribution of the calculated quantity then shows the effects of the

imprecision of the input data (Anderson, 1976). In this study, all input variables (Table 1) used to calculate the georeferenced wind vector and turbulent flux were randomly sampled from the Gaussian distributions with means corresponding to the constant altitude straight line flight and one standard deviation (1σ) widths given by individual component typical precision specifications from the manufacturer. Errors in the input variables are considered uncorrelated. The Monte Carlo process was repeated $N = 10^5$ times, and the measurement precision of the geo-referenced wind vector and turbulent flux was estimated as the standard deviation of the distribution of the simulated results.

The above wind measurement error analysis Monte Carlo error simulation gives the nominal measurement precision of the geo-referenced wind and turbulent flux, but does not consider the influence of environmental changes. Following the methods of Lenschow and Sun (2007), we assess whether the accuracy of wind measurements from the UAV in satisfying the minimum signal level needed for resolving the mesoscale variations of the three wind components in the encountered atmospheric conditions. Firstly, the minimum required signal level for measurement of vertical air speed (ω) under the encountered atmospheric conditions could be estimated as (Lenschow and Sun, 2007):

$$\frac{\partial w}{\partial t} < 0.2\sqrt{2}\sigma_w 2\pi k U_a \tag{1}$$

with the true airspeed (U_a) set to mean cruise speed 30 m s⁻¹, the peak signal magnitude (σ_w) of the power spectra, and the corresponding wavenumber (k) (Thomas et al., 2012). The measurement error of the system in the vertical wind component can be calculated as (Lenschow and Sun. 2007):

$$364 \frac{\partial w}{\partial t} \cong \Theta \frac{\partial U_a}{\partial t} + U_a \frac{\partial \Theta}{\partial t} + \frac{\partial w_{UAV}}{\partial t}$$
 (2)

with $\Theta = \alpha - \theta$, where α is the attack angle, θ is the pitch angle, w_{UAV} is the UAV's vertical velocity. According to Lenschow and Sun (2007), the signal level and mesoscale fluctuation of horizontal wind components (u and v) are considerably larger than that of vertical wind, so the accuracy criteria are not nearly as stringent. The measurement error of the horizontal wind component could be calculated as (Lenschow and Sun, 2007):

$$369 \left| \begin{array}{c} \frac{\partial u}{\partial t} \cong -\frac{\partial U_a}{\partial t} + \frac{\partial u_{UAV}}{\partial t} \end{array} \right. \tag{3}$$

$$370 \left| \begin{array}{c} \frac{\partial v}{\partial t} \cong \Psi \frac{\partial U_a}{\partial t} + v_{tas} \frac{\partial \Psi}{\partial t} + \frac{\partial v_{UAV}}{\partial t} \end{array} \right. \tag{4}$$

371 and,

348

349

350

351

352

353354

355

356

357

358

359

$$372 \mid \Psi \equiv \psi' + \beta \tag{5}$$

where u_{UAV} , v_{UAV} are the UAV's horizontal velocity measured from INS, ψ' is the departure of the measured true heading from the average true heading, and β is the sideslip angle of airflow. If the measurement error of the 3D wind vector from Eqs. (2) to (4) is smaller than the required minimum signal level of the vertical and horizontal wind components, it can be confirmed that the measurement accuracy of the geo-referenced 3D wind vector from UAV is sufficient to resolve the mesoscale variations of the three wind components in the encountered atmospheric conditions.

In addition, 2.4.2 Wind measurement evaluation

The key to successful aircraft EC measurements lies in the translation of accurately measured, aircraft-orientated, wind vectors to geo-referenced orthogonal wind vectors (Thomas et al., 2012). Agacurate measurements of geo-referenced wind vectors typically not only depend on the measurement precision of the sensors (i.e., 5HP and INS), but also related to the quality of the calibration parameters, as well as and the geometry structure of the UAV-based EC system (i.e., flow distortion and leverage effect). For evaluation of the effect of the latter two aspects, a calibration flight campaign (Section 2.2.1) was performed to determine the calibration parameter (ϵ_{ψ} , ϵ_{θ}), check its quality, as well as to ascertain the effects of the lever arm and up-wash by the wings. The methods for acquiring the calibration parameter were given by Vellinga et al. (2013) and Sun et al. (2021a), and the results are reported in Supplement Part C (Figs. S2 and S3). In During the in-flight calibration campaign, a 'racetrack' maneuver was performed to check the quality of the calibration parameters determined from the 'box' flight maneuver. The initial ($\epsilon_{\psi} = 0^{\circ}$, $\epsilon_{\theta} = 0^{\circ}$) and calibrated ($\epsilon_{\theta} = -0.183^{\circ}$, $\epsilon_{\psi} = 2^{\circ}$, in Supplement Part C) set of parameters were used to calculate the geo-referenced wind vector. By comparing the mean and standard deviation of the horizontal and vertical wind vector between the initial and calibrated set, the quality of the geo-referenced wind vector measurement in real environment conditions can be verified.

The relative wind vector (\widehat{U}_q) measured by the aircraft is susceptible to flow distortion because the airplane must distort the flow to generate lift and thrust, and the aircraft's propellers, fuselage, and wings are the main sources of flow distortion as flow barriers (Metzger et al., 2011). For fixed-wing aircrafts, the wind probe mounted on the nose of the UAV and extended as far forward of the fuselage as possible could avoid significant influence from flow distortion from the fuselage and propellers. Nevertheless, Eeffects from the flow distortion around the body of the aircraft, especially the induced upwash by the wings $\underline{\text{can}}_{s}$ can also significantly influence significantly influence the correspondence between measured and freestream flow variables (Garman et al., 2008). The induced upwash by the wings modifies the local angle of attack, causing the measured attack angle (α) to be larger than the free-stream attack angle (α_{∞}). The magnitude of upwash influence generally increases with airplane size and airspeed, typically ranging from 0.5 to 2.5 m s⁻¹ as reported by the manned fixed-wing aircraft (Garman et al., 2008). Therefore, for wind measurements by manned fixed-wing aircrafts, the upwash effects must be corrected (Garman et al., 2008; Kalogiros and Wang, 2002). However, wind measurements using a multi-hole probe on the UAV seldom need this correction due to the fuselage size and because the airspeed is very low compared to a manned aircraft.

In order to access whether the lift-induced upwash could be safety ignored by the current UAV-based EC system, the 'acceleration' flight maneuver was performed. AAccording to Crawford et al. (1996), the pitch angle (θ) by the INS instrument can be utilized as an estimate of the free-stream attack angle (α_{∞}) if the aircraft's vertical velocity is zero,

since it is unaffected by lift-induced upwash and varies directly with α_{∞} when the ambient vertical wind is zero. Under ideal conditions (zero aircraft vertical velocity and zero ambient vertical wind), the approximation relationship of $\theta \cong \alpha_{\infty}$ is valid when $\theta < 6^{\circ}$ (Crawford et al., 1996; Vellinga et al., 2013). Departures from the 1:1 relationship can be caused by airflow distortion around the airplane behind the 5HP. The 'acceleration-deceleration' maneuver produced various pitch and attack angles measured under various airspeeds, which allowed a direct comparison between the pitch angle (θ) and the attack angle (α). If the slope between α and θ is close to unity, it indicates that the influence of lift-induced upwash can be ignored; otherwise, its influence should be corrected using upwash models (Garman et al., 2006). At last Meanwhile, the influence of leverage effects was also evaluated based on the measurement data from the 'acceleration-deceleration' maneuver by considering or ignoring the leverage effect correction term in Eqs. (S15) to (S17).

2.4.2 Flux measurement error caused by instrumental noise

Errors or uncertainties in EC measurements can be systematic or random. Measurement from UAV, they can be attributed to several sources, mainly including instrumental noise, data handing, atmospheric conditions, insufficient flux calculation length, and bumpy flight environment (Mahrt, 1998; Finkelstein and Sims, 2001; Mauder et al., 2013). Knowledge of the measurement precision is of great importance for interpretation of EC measurements especially when detecting small fluxes in terms of turbulent exchange or signal-to-noise (SNR) of the instrumentation. Determination of the flux measurement error from instrument noise is very useful, as it is related not only to the system performance, but also to the minimum resolvable capability for the flux to be measured. In the current study, uncertainty related to instrumental noise (listed in Table 1) was estimated with a directly method proposed by Billesbach (2011). This method can be called as "random shuffle" (denoted as the RS) method and was "designed to only be sensitive to random instrument noise". According to Billesbach (2011), the uncertainty of the flux covariance can be expressed as:

$$\underline{\sigma_{\underline{w'x'}}} \equiv \frac{1}{N} \sum_{i,j=1}^{N} \underline{w'(t_i)x'(t_j)} \tag{6}$$

where x is the target entity of the covariance, N is the number of measurements contained in the block averaging period, $j \in [1...N]$ but the values are in the random order. The idea behind the RS method was that the randomly shuffled will remove the covariance between biophysical (source/sink) and transport mechanisms, leaving only the random "accidental" correlations attributed mostly to instrument noise (Billesbach, 2011). It means that the shuffled component x makes it uncorrelated in time/space and decorrelates x from w, resulting in two independent variables (i.e., $\overline{w'x'} \sim 0$), and any residual value of the covariance is attributed to random instrument noise. In addition to the basic assumptions made in EC flux calculation, one additional assumption in RS method is that the block averaging period should be sufficiently long to accurately represent the lowest significant frequencies contributing to the covariance which could be verified by forming Ogive plots of the covariance (Billesbach, 2011).

In this study, in order to obtain robust estimates of the instrumental noise, $\underline{\sigma}_{\overline{w'x'}}$ was repeatedly calculated 20 times for every straight and level flight leg in operation flight (Fig. 2), and the mean of the absolute values of these 20 replicated

440 estimates for $\sigma_{\overline{w'x'}}$ were used to estimate the random uncertainty related to instrumental noise. According to Rannik et al.

441 (2016), RS method tends to overestimate the covariance uncertainty due to instrumental noise only. Then, the uncertainty in

442 the flux covariance of sensible heat $(\underline{\sigma_{\overline{w'}\overline{D'}}})$, latent heat $(\underline{\sigma_{\overline{w'}\overline{O'}}})$, and CO_2 $(\underline{\sigma_{\overline{w'}\overline{O'}}})$ were estimated using RS method,

443 respectively.

444 <u>It should be noted that the measurement error of EC flux measurement is influenced not only by the uncertainty in the raw</u>

445 covariance but also by the propagated errors form correction terms (i.e., WPL correction) or any lens contamination

446 (Serrano-Ortiz et al., 2008). For EC measurement from our UAV, the signal quality of the IRGA is checked before each

447 flight measurement to ensure that the measurement of gas concentration is not affected by lens contamination. The final

448 uncertainty of flux measurement was evaluated using the partial derivatives of the full flux calculation equation with respect

449 to their flux value derived by Liu et al. (2006) (Eqs. S29 to S31 in Supplement Part B). These equations (Eqs. S29 to S31)

450 ignored the perturbations terms from the errors in the individual scalar (i.e., ρ_v , ρ_c , T) which were proved negligible small

451 (Serrano-Ortiz et al., 2008). At last, after several repetitive calculation of the Eq. (6), their averaged value then be combined

452 to Eqs. (S29) to (S31) for estimating the final flux measurement error due to instrumental noise.

453

454

463

467

469

470

2.4.3-3 Resonance effects

455 Previous work found that the measurement of the atmospheric scalars (e.g., air temperature, H₂O, and CO₂ concentration) is

456 susceptible to resonance effects caused by the operation of the engine and propeller (Sun et al., 2021b). In order to further

457 reduce the resonance effects, the vibration damping structure of the developed UAV-based EC system was further optimized.

The reference (co)spectra curve of Massman and Clement (2005) was used to quantify the influence of the resonance effects

459 remaining after vibration isolation optimization. Massman and Clement (2005) gave the generalization mathematical

460 expression of the models of spectra and co-spectra:

$$461 \quad Co(f) = A_0 \frac{1/f_x}{[1+m(f/f_x)^{2\mu}]^{\frac{1}{2\mu}(\frac{m+1}{m})}}$$
 (6)

462 where f is frequency (Hz), f_x is the frequency at which fCo(f) reaches its maximum value, A_0 is a normalization parameter,

m is the (inertial subrange) slope parameter, and μ is the broadness parameter. To describe co-spectra, m should be 3/4; to

464 describe spectra, m should be 3/2. According to Massman and Clement (2005), $\mu = 7/6$ under stable atmospheric condition

and $\mu = 1/2$ under unstable atmospheric condition. Fast Fourier transform (FFT) method was used to calculate the spectra

466 and co-spectra of the measured turbulent variables. Before calculating the turbulence (co)spectra, condition of the raw

turbulence data was performed, including a linear detrend and tapering using the Hamming window to reduce the spectral

468 leakage (sharp edge) according to Kaimal et al. (1989).

According to Sun et al. (2021b), the noise influence from resonance mainly appears in the high frequency domain,

According to the feature of spectral curve, the frequency range of the noise region was artificially designated to f > 8 Hz for

air temperature, $f = 1 \sim 5$ Hz for water vapor, and $f = 1 \sim 8$ Hz for CO₂. The normalized spectra and co-spectra curve were adopted and the area difference of the designated frequency range beneath the (co)spectra curve between the measured and reference (co)spectra curve was calculated to quantify the influence of resonance noise in the variance and flux <u>covariance</u> of the measurement atmosphere scalars. An example is shown in Figure 3, and also shown is the reference (co)spectra curve of Massman and Clement (2005), with the (co)spectral maximum at $f_x = 0.1$. The red region in Fig. 3 represents the impact extent of the resonance noise in the variance (Figs. 3a to 3c) and flux <u>covariancees</u> (Figs. 3d to 3f) of the measured scalars. The systematic noise deviation in the fluxes of sensible, latent heat and CO₂ could be derived relative to the entire frequency range.

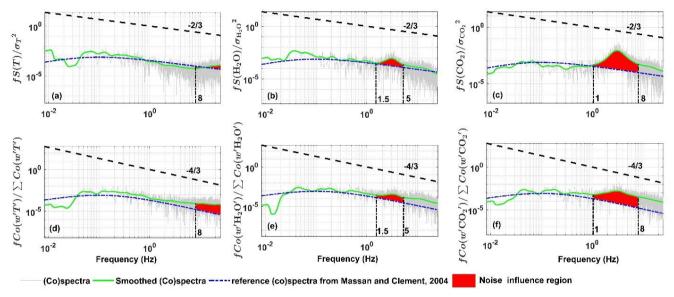


Figure 3. The influence of resonance noise on the spectra (top row) and co-spectra (bottom row) curve of the measured scalars based on the measurement from the standard operation flight campaign on 8 August 2022 at Dagang district, Tianjin, China. The red region is the area difference of the designated frequency range (vertical black dashed-dotted line) beneath the (co)spectral curve between the measured and reference (co)spectral curve.

2.4.4 Sensitivity analysis

To understand the <u>relevance_relative_influence_of_external_the_calibration parameters_onfor_the measurements of georeferenced wind vector and turbulent flux, two sensitivity tests were conducted. The magnitude of the <u>change-perturbation</u> in the wind vector and turbulent flux was investigated as a function of the uncertainties <u>of in the four_calibration</u> parameters, <u>including:</u> three mounting misalignment angles $(\epsilon_{\psi}, \epsilon_{\theta}, \epsilon_{\phi})$ between the 5HP and the CG of the UAV and one temperature recover factor $(\epsilon_r = 0.82)$ used to calculate the ambient temperature (Eq. 3 in Sun et al. 2021a).</u>

First, the sensitivity of the geo-referenced 3D wind vector and turbulent flux to the uncertainties of the individual parameter was investigated. The geo-referenced 3D wind vector and turbulent flux was calculated based on the straight leg (about 4 km) of the standard operational flight with an by addeing and error of ±30 % to the optimum value of each

calibration parameter alternately; except for ϵ_{ϕ} , for which the typical range of $\pm 0.9^{\circ}$ was taken for sensitivity analysis (Vellinga et al., 2013).

Then, in order to test the overall interaction between the parameters, a second sensitivity test was performed to calculate the geo-referenced 3D wind vector and turbulent fluxes by adding $\pm 30\%$ error to all calibration parameters simultaneously, to test the overall interactions between the parameters, a second sensitivity test was run to calculate the geo-referenced 3D wind vector and turbulent flux using an error of $\pm 30\%$ added to all the calibration parameters simultaneously. Lastly, their relative errors (*RE*) with respect to the original value were the relative errors of the calibrated set and the added error set of geo-referenced 3D wind vector and turbulent flux values was calculated to evaluate the perturbation of the wind vector and turbulent flux under the variation of each calibration parameter as well as under simultaneous variation of all calibration parameters. During the sensitivity analysis, the calculated geo-referenced wind and turbulent flux results whose absolute value was less than their least resolvable magnitude (in Table 2) were filtered out to avoid the influence of the errors contained in the measurements themselves on the results.

ing the error contained in the measurement itself from impacting results.

2.4.5 Relative error

In this study, relative error (*RE*) was used to evaluate the influence of different factors on the measurements of georeferenced wind vector and turbulent flux by the UAV-based EC system. It is defined as:

$$509 \quad RE = \frac{|x_0| - |x|}{|x|} \times 100 \% \tag{7}$$

- where '\ ' means the absolute value, x is the 'true' value, x_0 is the influence value. RE > 0 means the exerted influence will
- 511 cause the measurement value to be larger than 'true' value and vice versa.

512 3 Results

513

495

496

497

498 499

500

501

502

503

504

505

506

3.1 Wind measurement evaluation Error analysis

- 514 Evaluation of the wind measurement performance from the UAV-based EC system includes three contents: (1) measurement
- 515 precision and its ability to resolve the mesoscale variations of the wind, (2) checking the quality of the acquired calibration
- 516 parameters, and (3) checking whether the measured wind vector is affected by upwash flow and leverage effects.
- 517 First, according to the equations described in Supplement Part A (Eqs. S18 to S23), the measurement precision for
- 518 horizontal wind components is a function of true airspeed and true heading, while, the measurement precision for vertical
- 519 wind component is largely decided by the true airspeed. The typical values of true airspeed ranging from 25 m s⁻¹ to 35 m s⁻¹
- 520 (interval of 1 m s⁻¹) and the true heading values ranging from 0° to 180° (interval of 30°) were used in the evaluation of
- 521 wind measurement error. Then, the measurement precision (1σ) of the geo-reference 3D wind vector from aircraft was

estimated using the measurement precision of the related parameters from Table 1. The results are shown in Figure 4 for the measurement precision of horizontal wind (σ_u and σ_v in Figs. 4a and 4b) and vertical wind (σ_w in Fig 4c), respectively. The typical values of the measurement precision are ranging from 0.05 m s⁻¹ to 0.07 m s⁻¹ for horizontal wind component u, ranging from 0.02 m s⁻¹ to 0.08 m s⁻¹ for horizontal wind component v, and ranging from 0.05 m s⁻¹ to 0.07 m s⁻¹ for vertical wind component v. When the flight direction is towards due east or due west, the horizontal wind (v and v) has the smallest measurement error.

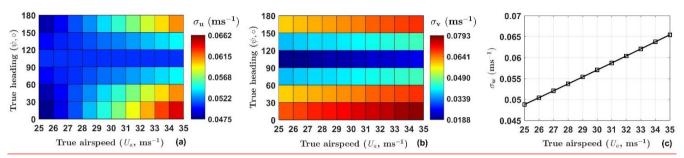


Figure 4. Estimated measurement precision (1σ) of the horizontal wind (a, b) and vertical wind (c) according to the equations described in Supplement Part A (Eqs. S18 to S23).

Generally speaking, an autopiloted UAV can maintain a near-constant true airspeed during the cruise flight phase. For a true airspeed of 30 m s⁻¹ for the current UAV-based EC system during the cruising, the maximum measurement error in the northward, eastward, and vertical velocities of the geo-referenced wind components were calculated as approximately 0.06, 0.07, and 0.06 m s⁻¹, respectively. Then, we assume that a minimum signal-to-noise ratio of 10:1 is required to measure the wind components with sufficient precision for EC measurements (Metzger et al., 2012). Accordingly, in the real environments, horizontal and vertical wind speed greater than 0.7 m s⁻¹ and 0.6 m s⁻¹ can be reliably measured, respectively (Table 2).

After running $N=10^5$ samples through the data processing procedure, the results of Monte Carlo error simulation for geo-referenced wind vector and turbulent flux are summarized in Table 2. The least resolvable magnitude of the measurements of wind and turbulent flux was calculated by assuming the minimum required signal-to-noise ratio of 5:1 (Metzger et al., 2012). The 1σ precision of the geo-referenced wind vector measurement was ± 0.04 m s⁺ for the horizontal and vertical wind components, and the magnitude of the wind velocity greater than 0.2 m s⁺ could be reliably measured. The calculated measurement 1σ precision of the geo-referenced vertical wind component agreed well with the 1σ uncertainty in the vertical wind measurement (0.057 m s⁺) calculated by propagating instrument errors through linear combination, as described by Sun et al. (2021a).

Table 2: The maximum measurement error in the northward (u), eastward (v), and vertical (w) velocities of the georeferenced wind components at the true airspeed of 30 m s⁻¹, Results of the simulated measurement precision (1σ) of the geo-

referenced wind vector and turbulent flux from the Monte Carlo error simulation with N=10⁵ runs and the least resolvable magnitude assuming the minimum required signal-to-noise ratio of 105:1.

Measurements	Measurement precision (1σ)	Least resolvable magnitude	
u -windspeed (m s ⁻¹)	0.0 <u>6</u> 4	0.62	
<i>v</i> -windspeed (m s ⁻¹)	0.0 <u>7</u> 3	0.7	
w-windspeed (m s ⁻¹)	0.0 <u>6</u> 4	0.62	
CO ₂ -flux (µmol m ⁻² -s ⁻¹)	0.08	0.4	
Sensible heat flux (W m ⁻²)	1.61	8.05	
Latent heat flux (W m ⁻²)	0.15	0.75	
Friction velocity (m s ⁻¹)	0.02	0.1	

The simulated 1σ measurement precision of the turbulent flux was $0.08~\mu\text{mol}\ m^2$ -s for the CO₂ flux, $1.61~W\ m^2$ -for the sensible heat flux, $0.15~W\ m^2$ -for the latent heat flux, and $0.02~m\ s^4$ -for the friction velocity, respectively. Using a signal-to-noise ratio of 5:1, the minimum magnitudes for reliably resolving the CO₂ flux, friction velocity, sensible and latent heat fluxes were $0.4~\mu\text{mol}\ m^2$ -s, $0.1~m\ s^4$, $8.05~W\ m^2$, and $0.75~W\ m^2$, respectively.

The Monte Carlo error simulationabove results gave the nominal precision for wind measurements that does not consider the influence of environmental conditions. Changes in the environment will lead to sensor drift, increasingly deteriorating the measurement with flight duration (Metzger et al., 2012; Lenschow and Sun, 2007). Following the methods of Lenschow and Sun (2007), the ability of the limitations of the accuracy of wind field measurements from UAV to resolve the mesoscale variations of the 3D wind components in the encountered atmospheric conditions was assessed. For the vertical wind, the mesoscale variability was defined as the peak signal magnitude of the power spectra curve. The corresponding average wavenumber was determined as 0.09 m⁻¹ based on the straight flight leg (about 4 km, lasting about 120 s) of the standard operational flight. Then, the minimum required signal level for the vertical wind measurement was estimated as $\partial w/\partial t \simeq 0.14 \text{ m s}^{-2}$. The accuracy of the vertical wind measurement using Eq. (2) is estimated as follows. The first term on the right-hand side of Eq. (2) is dominated by the drift in the differential pressure transducer, the value of $\partial U_a = 0.05 \text{ m s}^{-1}$ acquired from the wind tunnel test was applied (Table 1). The histogram of Θ derived from the standard operational flights is shown in Figure 45. The 99 % confidence interval indicates that the value of Θ seldom exceeds $\pm 3^\circ$, i.e., ± 0.053 radians. Thus, the value of the first term was estimated as $2.2 \times 10^{-5} \text{ m s}^{-2}$.

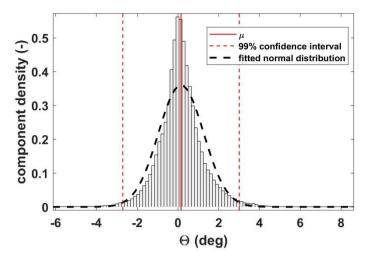


Figure 54. Histogram of Θ derived from the standard operational flight. Component density is scaled so that the histogram has a total area of one. Red vertical lines indicate distribution average (solid) and 99% confidence interval (dashed). The black dashed bell curve displays a reference fitted normal distribution.

The second term in Eq. (2) is a combination of INS pitch accuracy and drift in the measured attack angles. The combined accuracies of these two sensors were applied to derive $\partial\theta=0.0024$ radians. Thus, the second term in Eq. (2) was estimated as 6×10^{-4} m s⁻². Finally, the third term in Eq. (2) was estimated as 1.7×10^{-4} m s⁻², according to the stated accuracy of the vertical velocity from the INS. The overall performance of the vertical wind measurement (7.9×10^{-4} m s⁻²) was accurate enough to resolve the mesoscale variations in vertical air velocity.

The required accuracy of horizontal wind for mesoscale measurement was estimated as 10 times larger than that of vertical wind, i.e., $\partial u/\partial t \simeq \partial v/\partial t \simeq 1.4 \,\mathrm{m\ s^{-2}}$. The measurement accuracy of the horizontal wind component u was estimated as $4.8 \times 10^{-4} \,\mathrm{m\ s^{-2}}$ according to Eq. (3). Like the first term in Eq. (2), with the value of Ψ rarely exceeding ± 0.18 radians, the measurement accuracy of the horizontal wind component v was estimated as $2.7 \times 10^{-2} \,\mathrm{m\ s^{-2}}$ according to Eq. (4). Thus, the measurement accuracy of the horizontal wind components was accurate enough to resolve the mesoscale variations in the horizontal air velocity as well.

3.2 Wind measurement

Second, before checking the quality of the acquired calibration parameters, \mathbf{T}_{t} he calibration results of the offset in pitch (ϵ_{θ}) and heading (ϵ_{ψ}) angles based on the 'box' maneuver are provided in Supplement Part C (Figs. S2 and S3). The final calibration values are $\epsilon_{\theta} = -0.183^{\circ}$ and $\epsilon_{\psi} = 2^{\circ}$. In order to verify the quality of these calibration parameters, a 'racetrack' maneuver was performed. Figure 5–6 shows the verification results by plotting wind vectors and calculating summary statistics for the 'racetrack' maneuver (including turns), using the initial $(\epsilon_{\theta} = \epsilon_{\psi} = 0^{\circ}, \text{Fig. 5a6a})$ and calibrated (Fig. 5b6b) set of parameters. The introduction of the calibration parameter effectively improved the quality of geo-referenced wind vector measurement. The standard deviation for wind direction, $\sigma_{U_{dir}}$, is 4.9° for the calibrated set compared to 8.7° for the

initial set, and the standard deviation of wind speed, σ_U , is 0.52 m s⁻¹ for the calibrated set compared to 1.12 m s⁻¹ for the initial set. The average vertical wind speed is much closer to zero ($\bar{w} = -0.006 \text{ m s}^{-1}$) for the calibrated set than for the initial set ($\bar{w} = 0.1 \text{ m s}^{-1}$). For the horizontal wind, it is evident from Fig. 5-6 that the wind direction and velocity are little affected by sharp turns. On the contrary, the measurement of the vertical wind component is obviously affected by turns in flight, as shown by the large ripple in the vertical wind speed around the scan value of 150 (Fig. 56). It should be noted that the influence of upwash flow and the leverage effect are not considered in the calculated of geo-referenced wind vector.

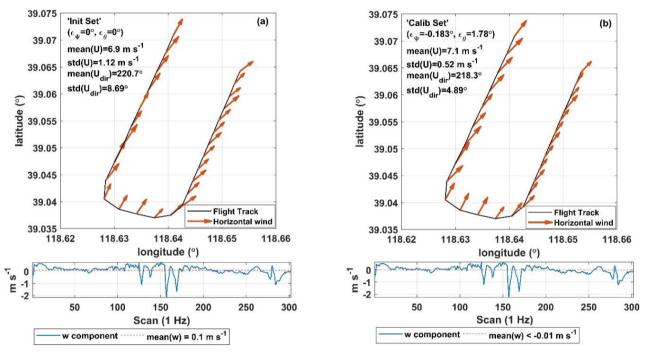


Figure 56. Quality check of the calibration parameter by plotting wind vectors and calculating summary statistics for the 'racetrack' maneuver, using the initial (a) and calibrated (b) set of parameters. The calibration flight was carried out on 4 September 2022 at the Caofeidian Shoal Harbor.

Third, In order to check the influence of the lift-induced upwash on the attack angle measurement from the 5HP, an 'acceleration-deceleration' flight maneuver was performed. During the 'acceleration-deceleration' maneuver, INS data shown a vertical velocity of the UAV at 0.05 ± 0.2 m s⁻¹, the altitude of UAV at 392 ± 0.6 m, the heading of UAV at $199\pm2.4^{\circ}$. The flight conditions met the requirements of the 'acceleration-deceleration' maneuver (Vellinga et al., 2013). The relationship between the pitch angle (θ) measured by INS and the attack angle (α) measured by 5HP is plotted in Figure 67, where the attack angle was not corrected for lift-induced upwash. The slope (0.94) between θ and α is close to its theoretical value of 1, and the intercept (0.16) is close to zero. This result indicates that the lift-induced upwash has only a very small effect on the attack angle, and the influence of upwash could be ignored.

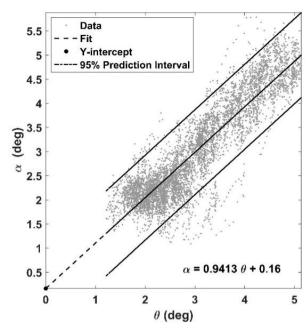


Figure 67. Relationship between the pitch angle (θ) measured by INS and the attack angle (α) measured by 5HP. The fitted linear equation is also shown.

Finally, the geo-referenced wind vector was calculated with and without the correction for the leverage effect based on the measurement data from the 'acceleration-deceleration' flight maneuver. The average relative differences between the corrected and uncorrected horizontal and vertical wind speeds are 0.1 % and 0.2 %, respectively. The standard deviation for horizontal wind speed is 0.307 m s⁻¹ without the level arm term compared to 0.306 m s⁻¹ when the level arm term is introduced. The standard deviation of vertical wind speed is 0.254 m s⁻¹ without the level arm term compared to 0.253 m s⁻¹ with the level arm term. The correction of leverage effect had minimal effect on improving the geo-referenced wind vector measurement; therefore, this correction term can be ignored.

3.2 Flux measurement error caused by instrumental noise

Flux measurement error caused by the instrumental noise gives the lowest limit of the flux that the UAV-based EC system is able to measure. Such measurement error was assessed by combining the covariance uncertainty estimated by RS method and the propagation of errors in flux correction terms. Before estimating the flux covariance uncertainty using RS method, using the measured data from each straight and level flight leg of the standard operational flight (Fig. 2), the normalized integrated cospectra (ogives) curves of sensible heat (Fig. 8a), latent heat (Fig. 8b), and CO_2 (Fig. 8c) flux are formed as a function of wavenumber (k), where $k = 2\pi f/U_a$. As shown in Figure 8, although the heterogeneous turbulence (or mesoscale turbulence) interfered the shape of ogive curves, most curves converged at the high and low frequency ends, which indicated that these segmented data were sufficiently long to represent the lowest significant frequencies contributing to the covariance.

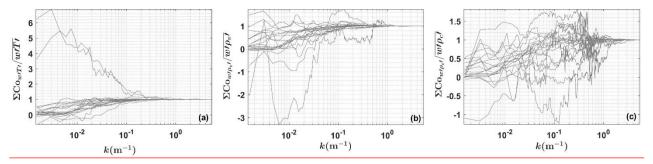


Figure 8. Normalized ogive curves as a function of wavenumber for the flux covariance of sensible heat (a), latent heat (b), and CO₂ (c) from each straight and level flight leg of the standard operational flight in Section 2.2.2.

Then, the results of instrumental noise related relative flux measurement error compared to the magnitude of the flux are shown in Figure 9. It can be seen that the flux measurement error caused by instrumental noise significantly decreases when the flux magnitude increases. It is not surprising since, in theory, instrumental noise is usually close to a constant and the relative flux measurement error caused by instrumental noise will decreases with increasing measurement magnitude. Overall, instrumental noise has the least effect on latent heat flux (ranging from 0.02% to 2.42% in this study) measurements, followed by sensible heat flux (ranging from 0.05% to 8.6% in this study), and has the greatest effect on the measurement of CO₂ flux (ranging from 0.22% to 75.6% in this study). Then, a simple rational function relationship between the relative measurement error and the flux magnitude is fitted according to the measured data, where the constant term in the denominator is set to 0. The fitted coefficient in the numerator can be considered as the flux measurement error caused by instrumental noises, which are 0.03 µmol m⁻² s⁻¹, 0.02 W m⁻², and 0.08 W m⁻² for the measurement of CO₂ flux, sensible and latent heat flux, respectively. At last, using the signal-to-noise ratio of 10:1, the minimum magnitudes for reliably resolving the CO₂ flux, sensible and latent heat fluxes were estimated as 0.3 µmol m⁻² s⁻¹, 0.2 W m⁻², and 0.8 W m⁻², respectively.

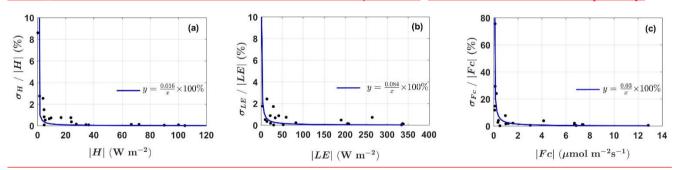


Figure 9. Relative flux measurement error caused by instrumental noise plotted against the magnitude of the flux. Also shown the fitted error curves. Measured data was from the standard operation flights in Section 2.2.2.

3.3 Resonance noise

The resonance noise from the engine and propeller can lead to systematic overestimation of the variance and covariance of the observed atmospheric scalars. Since the noise mainly appears in the high frequency domain of the (co)spectra, the reference (co)spectral curve of Massman and Clement (2005) was used to quantify the systematically bias caused by the resonance noise.

All spectra curves of the variance of measured scalars (including air temperature, H_2O , and CO_2 concentration) approximately followed the reference spectra curve and the reference -2/3 slope in the inertial subrange (Figs. 3a to 3c). The largest scatter occurred in the spectra of CO_2 (Fig. 3c). When comparing the spectra curve with the reference spectra, the resonance noise led to a systematic deviation in the variance of air temperature, H_2O , and CO_2 concentration of 0.1 ± 0.1 %, 1.0 ± 0.79 %, and 4.4 ± 0.66 %, respectively, relative to the entire frequency range. For the flux covariancees of sensible, latent heat and CO_2 , all the co-spectra curves approximately follow the reference co-spectra curve and the reference -4/3 slope in the inertial subrange (Figs. 3d to 3f). Compared with the reference co-spectra, the resonance noise led to a systematic deviation in the flux of sensible, latent heat, and CO_2 of 0.07 ± 0.004 %, 0.3 ± 0.25 %, and 2.9 ± 1.62 %, respectively, relative to the entire frequency range.

The results show that <u>the</u> resonance noise has a very little impact on the measured variance and flux <u>covariancees</u>. Among them, the measurements of CO₂ concentration and flux are most susceptible to the resonance noise, but the impact of this noise is limited to around 5 % of the observed value.

3.4 Sensitivity analysis

In this study, in order to investigate the <u>relevance of the calibration parameters for the measurement of the sensitivity of the</u> geo-referenced wind vector and turbulent flux-measurements to the uncertainties of the external calibration parameter, <u>two</u>a sensitivity tests wereas conducted by adding an error of ± 30 % to the <u>used optimum parameters (ε_{ψ} , ε_{ϱ} , ε_{φ} , ealibrated value of each calibration parameter. Welt is assumed that the maximum uncertainties contained in the calibration parameter is not more than 30 % of its own value.</u>

First, the sensitivity of the geo-referenced 3D wind and turbulent flux to the uncertaintyies in the individual parameters was tested. The *RE* value iswas used to quantify the sensitivity. Tables 3 and 4 show the and the *RE*-results are summarized in Tables 3 and 4. For the measurement of the geo-referenced wind vector, Table 3 shows that the uncertaintyies in the temperature recovery factor (ε_r) and 5HP mounting misalignment error in the roll (ε_ϕ) angle do not contribute significantly to errors in the wind measurements, which were typically smaller than 4% of the observed value in this study. PThe parameter ε_θ had the largest effect on the vertical wind component (up to 30 %), whereas ε_ψ had the largest effect on the horizontal wind component. For the measurement of turbulent fluxes, Table 4 shows that the errors in ε_r and ε_ϕ does not significantly influence significantly the flux measurements ε_θ and ε_ψ had significant effects on the measurement of turbulent flux. Adding

an error of ± 30 % to the calibration parameter ε_{θ} may result in significant perturbation (large *RE* variance) uncertainty in the measured turbulent fluxes including sensible heat, latent heat and CO₂-(large *RE* variance). While, errors in ε_{ψ} to some extent mainly affect the measurement of Similarly, adding an error of ± 30 % to the calibration parameter ε_{ψ} resulted in the largest effect on latent heat fluxes (*RE* may up to 15 %).

Table 3: RE of the sensitivity test for the geo-referenced 3D wind vector (u, v, w). An error factor of ± 30 % was added to each calibrated parameter. The geo-referenced 3D wind vector was calculated based on the straight leg of the standard operational flight.

Parameter Error (%)		RE of g	eo-referenced 3D wind	vector
			mean \pm std	
		u (%)	v (%)	w (%)
	-30	0.04 ± 0.41	-0.004 ± 2	0 ± 0
$\boldsymbol{arepsilon_r}$	30	0.06 ± 0.43	0.27 ± 1.1	-0.07 ± 0.23
c *	-30	0.41 ± 2.51	-0.09 ± 2.05	1.15±2.43
$oldsymbol{arepsilon_{oldsymbol{arphi}}}^*$	30	-0.43 ± 2.61	0.09 ± 1.79	-1.1±2.60
	-30	0.03 ± 0.41	-0.35 ± 2.54	-30.51±6.42
$oldsymbol{arepsilon}_{oldsymbol{ heta}}$	30	0.05 ± 0.45	0.42 ± 1.82	30.37 ± 6.6
c .	-30	2.98 ± 25.06	-2.04 ± 16.3	0±0
$oldsymbol{arepsilon_{\psi}}$	30	-2.97 ± 24.96	2.42 ± 16.63	0±0

^{*} The optimum calibration value is set to 0, ε_{ω} was varied over $\pm 0.9^{\circ}$, which is 30 % of its typical range.

Table 4: RE of the sensitivity test for the turbulent fluxes. An error factor of ±30 % was added to each calibrated parameter.

The turbulent fluxes were calculated based on the straight leg of the standard operational flight.

		RE of turbulent flux			
Parameter	Error (%) mean \pm std				
	• •	<i>fCO2</i> - <i>Fc</i> (%)	H (%)	LE (%)	u* (%)
	-30	1.04±3.04	-0.76±4.82	0.1±0.29	0±0
$oldsymbol{arepsilon}_{r}$	30	-1.0 ± 3.3	0.74 ± 4.8	-0.1 ± 0.29	0.2 ± 1.07
c *	-30	0.07 ± 1.2	0.03 ± 0.7	0.15 ± 1.51	0.54 ± 1.71
ϵ_{φ}^* 30	30	-0.14 ± 0.89	-0.06 ± 0.7	-0.16 ± 1.46	0.12 ± 1.61
	-30	-3.27±11.18	-0.8 ± 9.48	0.19 ± 11.91	-4.08 ± 5.61
$oldsymbol{arepsilon}_{oldsymbol{ heta}}$	30	2.34 ± 10.52	-0.44 ± 8.24	-1.27 ± 9.92	3.73 ± 4.53
6	-30	1.78 ± 5.18	-0.73 ± 4.87	1.89 ± 13.42	0.63 ± 5.75
$oldsymbol{arepsilon_{\psi}}$	30	-0.99 ± 3.96	-0.57 ± 3.26	2.66 ± 11.76	-0.59 ± 4.42

^{*} See Table 3.

The second sensitivity test was performed to evaluate the overall interaction between calibration parameters and the calculation of geo-referenced wind vector and turbulent flux by adding an error of $\pm 30 \%$ to all the calibration members simultaneously. overall sensitivity of the geo-referenced 3D wind vector and turbulent flux to the external calibration parameters was tested by adding an error of $\pm 30 \%$ to all the calibration members simultaneously. Tables 5 and 6 provided a summary of the RE results from the second sensitivity test. For the measurement of geo-referenced wind vector (Table 5),

adding an error of ± 30 % to <u>all</u> the calibration <u>parameters</u>set at the same time resulted in <u>great perturbations to both the horizontal</u> (low *RE* with high variance) and vertical wind components (high *RE* with low variance). For the measurement of turbulent fluxes, 30% error in the calibration parameters can result in errors in measured fluxes more than 10%. In addition, Table 6 reveals that the a high *RE* (near 30 %) for vertical wind and a low *RE* (about 4%) for horizontal wind, which also had high *RE* variance (up to 28 %). For the measurement of turbulent fluxes, the measurement of the latent heat flux (mean RE > 6 %) is more sensitivity to the errors in the calibration parameter than other <u>flux</u> measurements (higher mean and variance of the *RE* compared to other measurements)(mean RE < 3 %) and had a larger *RE* variance (>10 %). In general, CO_2 and sensible heat fluxes as well as friction velocity are not sensitive to errors in the external calibration parameters, but there were also some exceptions where the response to errors was large (e.g., the *RE* variance of CO_2 flux up to 10 %).

Table 5: RE of the sensitivity test for the geo-referenced 3D wind vector (u, v, w) calculated by adding an error of ± 30 % to all the calibrated parameter simultaneously. The geo-referenced 3D wind vector was calculated based on the straight leg of the standard operational flight.

	vector				
Parameter	Error (%)	$mean \pm std$			
		u (%)	v (%)	w (%)	
A 11	-30	4.24±27.89	-3.2±21.1	-29.35±4.63	
All	30	-4.15 ± 27.46	3.55±21.91	29.16 ± 4.86	

Table 6: RE of the sensitivity test for the turbulent flux calculated by adding an error of ± 30 % to all the calibrated parameter simultaneously. The turbulent flux was calculated based on the straight flight leg of the standard operational flight.

		RE of turbulent flux			
Parameter	Error (%)	$mean \pm std$			
		<u>fCO2-Fc (%)</u>	H (%)	LE (%)	u* (%)
A 11	-30	-1.19±10.51	-0.9±8.06	2.71±13.91	-2.92±8.19
All	30	-0.49 ± 10.01	-1.66 ± 5.4	-6.07 ± 13.24	1.74 ± 6.55

4 Discussions

As one in a new generation of airborne flux measurement platforms, the UAV-based EC system can significantly reduce the cost of implementing airborne flux measurement campaigns and greatly promote their wide application at regional scales. The trend of sensor miniaturization further promotes the rapid development of technology in this field. Sun et al. (2021a) developed an UAV-based EC system for measuring the turbulent flux of sensible heat, latent heat, CO₂, as well as radiation fluxes of net radiation and photosynthetically active radiation (PAR). The current study aimed to quantitatively evaluate the basic performance of this systeme UAV-based EC system developed by Sun et al. (2021a) in the measurement of wind vector and turbulent flux.

First, the <u>wind</u> measurement precision (nominal precision) of the UAV-based EC system was <u>evaluated_estimated_by</u> propagating the sensor errors to the geo-referenced wind vector <u>and turbulent flux_using the linearized Taylor series</u>

expansions from Enriquez and Friehe (1995) Monte Carlo error simulation. The 1σ precision for of geo-referenced wind measurement was estimated to be ±0.04-07 m s⁻¹, and the least resolvable magnitude for wind measurement was estimated at 0.7 m s⁻¹ by assuming the minimum signal-to-noise ratio of 10:1. 1σ precision for turbulent flux was 0.08 μmol m⁻² s for the CO₂ flux, 1.61 W m⁻² for the sensible heat flux, 0.15 W m⁻² for the latent heat flux, and 0.02 m s⁻¹ for the friction velocity. As proposed by Lenschow and Sun (2007), a minimum signal to-noise ratio of 5:1 was assumed to be required to measure the wind field and turbulent flux. Using that ration, the least resolvable magnitude for wind and turbulent flux measurement was estimated at 0.2 m s⁻¹ for wind velocity, 0.4 μmol m⁻² s for the CO₂ flux, 8.05 W m⁻² for the sensible heat flux, 0.75 W m⁻² for the latent heat flux, and 0.1 m s⁻¹ for the friction velocity. These derived wind measurement minimum resolvable magnitudes for measurements of wind vector and turbulent flux can be used as a basic reference for the wind measurement capability of the UAV-based EC system, and the measured values of wind vector and sealar fluxes smaller than the minimum resolvable values should be considered unreliable. The accuracy of the sensors was also assessed by examining the collected data in the real environment (Lenschow and Sun, 2007; Thomas et al., 2012). Our results revealed that Tthe overall performance of geo-referenced wind measurement is sufficient accuracy for resolving the mesoscale variations of the 3D wind components under the encountered atmospheric conditions. Therefore, it is possible to capture the mesoscale variability of the atmospheric boundary layer (ABL) over a wide range of spatial scales by performing longer flight paths.

SecondSecond, based on the measurement data from the in-flight calibration campaign, several key factors affecting the accuracy of geo-referenced wind measurement were analysed. First, the UAV-based EC system was calibrated (in Supplement Part C) using measured data from the 'box' flight maneuver to correct the mounting misalignment between the 5HP and the CG of the UAV in the heading ($\epsilon_{\theta} = -0.183^{\circ}$) and pitch ($\epsilon_{\psi} = 2^{\circ}$) angles. The quality of the acquired calibration parameters was verified using the 'racetrack' flight maneuver, and the acquired calibration value effectively improved the observed wind field with smaller variance compared with the wind calculated using their initial value. The measured geo-referenced 3D wind vector was consistent with the assumptions made about the atmospheric condition for calibration campaign (constant horizontal wind and near zero mean vertical wind) especially in the standard operation flight. At the same time, The measurement of the vertical wind component was significantly affected by the in-flight turn (maintaining about 20° roll). Therefore, it is necessary to avoid using the measured data from the turn section for turbulent flux calculation. Compared to-the_other studies (Vellinga et al., 2013; Reineman et al., 2013), the relatively large variance in the ealibrated-horizontal wind and wind direction after calibrated in this study may be caused by the nonstationary condition of the turbulence. This was caused by the reason that the flight altitude of 400 m was not high enough to_totally avoid interaction from the underlying surface.

The current calibration procedure did not include methods to determine the offset angle in roll (ε_{φ}) and the temperature recovery factor (ε_r) because of the small vertical separation (27.3 cm) between the 5HP and the roll axis of the UAV and the small Mach number (<0.1) during operational flight. The default ($\varepsilon_{\varphi} = 0^{\circ}$) and empirical ($\varepsilon_r = 0.82$) value were adopted for

these two ese calibration parameters. The sensitivity analysis shown these two parameters have no large effect on the wind vector and turbulent flux.

It should be noted that Wwind measurements from the airborne platform may be susceptible to flow distortion and rigid-body rotation (leverage effects). HoweverGenerally, the influence of these two aspects—factors were ignored by UAV platform when calculating the geo-referenced wind vector. To confirm that these effects could be safely ignored, data from 'acceleration-deceleration' flight maneuver was used to analyse the effects of lift-induced upwash and the leverage effect on the wind measurements. OurThe results demonstrate that the upwash has almost no effect on the wind measurement, which was indicated by the near 1:1 relationship (0.94 in Fig. 7) between the measured attack angles and pitch angle. The slight departures from the ideal 1:1 relationship may have been caused by the nonstationary condition during the flight. For the influence from the leverage effects, the Odiifferences in 3D wind vector between corrected and uncorrected for the leverage effect. Thus, ignoring the influence of the leverage effect has almost no effect on the measurement of wind. Therefore, we concluded that the geo-referenced 3D wind vector can be measured reliably by the current UAV-based EC system without considering the possible disturbance interference from the lift-induced upwash and leverage effect as indicated by the results.

Third, instrumental noise related relative flux measurement error was estimated by combining the covariance uncertainty estimated by RS method and the propagation of errors in flux correction terms. By assuming that the instrumental noise is close to a constant, we fitted a simple rational function relationship between the relative measurement error and the flux magnitude according to measured data (Fig. 9), and the fitted coefficient in the numerator can be considered as the flux measurement error caused by instrumental noises. The estimated flux measurement error of CO₂, sensible and latent heat flux are 0.03 µmol m⁻² s⁻¹, 0.02 W m⁻², and 0.08 W m⁻², respectively. Since the RS method directly uses the shuffled raw measurement data to calculate the instrumental noise induced flux measurement error, its estimated results inevitably included the effects of resonance noise from the UAV. Using the signal-to-noise ratio of 10:1, the least resolvable magnitude for turbulent flux measurement was estimated to be 0.3 µmol m⁻² s⁻¹ for the CO₂ flux, 0.2 W m⁻² for the sensible heat flux, 0.8 W m⁻² for the latent heat flux, respectively.

ThirdFourth, because the UAV-based EC system has not completely insulated the noise from the operation of the engine and propeller and its effect on the measured scalars, the reference (co)spectra of Massman and Clement (2005) was used to quantify the effect of the resonance noise on the variance and flux of the measured scalars. Previous studies found that the influence of resonance noise mainly appears in the high frequency domain of the power spectra of the measured atmospheric scalars (e.g., air temperature, H₂O, and CO₂ concentration). The frequency range of the noise region was artificially designated for air temperature, water vapor and CO₂. By calculating the area difference of the designated frequency range beneath the (co)spectral curve between the measured and reference (co)spectral curves, the resonance effect could be quantified. The results shown that, overall, resonance noise has little impact on the variance and flux covariancees of the measured scalars. The measurements of CO₂ concentration and its flux covariance were the most susceptible to resonance

noise, but the maximum effect was less than 5 %. It should be noted that this method may overestimate the deviation caused by resonance noise as indicated by the reference (co)spectra curve and the measured (co)spectra not fully overlapping in the inertial subrange (shown in Fig. 3).

785

786

787

788

789

790

791

792

793

794

795

796

797

798

799

800

801

802

803

804

805

806

807

808

809

810

811

812

813

814

815

816817

In general, Ggas detection based on optical absorption methods can achieve fast and high precision gas concentration measurements, but they are extremely sensitive to vibration noise. However, due to the limited space inside the UAV, it is difficult to install all the hardware needed for a complex vibration isolation structure to effectively isolate the impact of vibration on the gas analyser. The weight and the aerodynamic shape of the UAV also present challenges. In the future, a new UAV-based EC system based on a pure electric UAV eould-will be developed. The electro-powered UAV has similar performance to the current fuel-powered UAV but can minimize the impact of vibration noise from the engine and propeller rotation, which makes it possible to completely isolate the resonance effect using a simple vibration isolation structure. Electro-powered UAVs also have other advantages including larger wingspan (lower cruising speed), a constant CG position, and lower operational complexity compared to the current system.

ForthFifth, two sensitivity tests wereas conducted to assess the perturbation of the geo-referenced wind vector-velocity and turbulent flux under variation ($\pm 30\%$) of each calibrationed parameter around its calibrated optimum value ($\epsilon_{th} \equiv$ 2° , $\epsilon_{\theta} = -0.183^{\circ}$, $\epsilon_{\phi} = 0^{\circ}$, $\epsilon_{r} = 0.82$) as well as under simultaneous variation (+30%) of all calibrationed parameters. Their RE was used to evaluate the sensitivity, and values of wind and flux less than their least resolvable magnitude were removed from the calculation. The sensitivity analysis was carried out based on the straight flight leg (about 4 km) by adding an error of ±30 % to the calibrated parameter value and then calculating the RE of the geo-referenced wind vector and turbulent flux between the calibrated and added error sets of results. Values less than the least resolvable magnitude were removed from the dataset. The results revealed that uncertainties in the temperature recovery factor (ε_r) and mounting offset in roll angle (ε_{ω}) do not significantly contribute to an error in the measurement of wind vector and turbulent fluxes. The typical RE for the geo-referenced wind measurements is less than 1.2 % with variance less than 3 %, and the typical RE for turbulent flux is less than 1.1 % with variance less than 5 %. The sign of the added errors of ε_{ω} and ε_{κ} also has no significant effect on the sign of the RE. Calibration parameters that had the largest effect on the measurement of geo-referenced wind vector and turbulent flux are the mounting offset angle in pitch (ε_{θ}) and heading (ε_{ψ}) . Uncertainties in ε_{θ} had a direct effect on the measurement of vertical wind component, and these then these errors propagate to the measured fluxes, resulting in a large fluctuation error contains in the RE (\sim 12 %)measured fluxes (\sim 15 %). A negative error of in the ε_{θ} will resulted inlead to smaller an underestimation of the vertical wind and vice versa. Uncertainties Errors in ε_{ψ} directly affect the measurement of the horizontal wind, and to some extent, the measurement of turbulent flux. The most obvious phenomenon difference is that the added error in ε_{ψ} lead to a great variability (up to 25 %) in the RE of horizontal wind. By checking the relationship between the magnitude of the horizontal wind (u, v) and RE, a near exponential relationships was seen, as shown in Figure 10. As shown in Figure 7, tThe influence of the error in the ε_{ψ} decreased significantly with the increase in the magnitude of the horizontal wind velocity. Additionally, the measurement of latent heat flux may be greatly

affected by the error in ε_{ψ} , which is reflected by the relatively large deviancy (~14 %) of the *RE*. Therefore, calibration parameter ε_{θ} and ε_{ψ} need to be carefully calibrated.

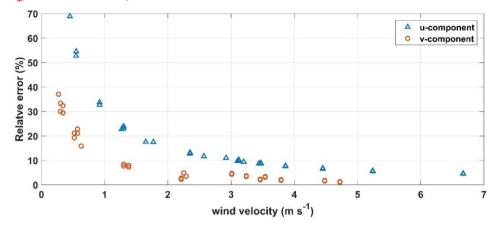


Figure 107. Relationship between the magnitude of the horizontal wind velocity (u, v) and RE from the sensitivity test. Windvelocity of less than 0.2 m s⁺ was excluded.

The second sensitivity test varied all the external calibration parameters simultaneously, which resulted in a slightly larger-but similarly varied REs compared with the first test. The results confirm that the quality of calibration parameter ε_{θ} and ε_{ψ} had the largest effect on the reliability of measurement for geo-referenced wind vector and turbulent flux. The parameter ε_{θ} directly affected the measurement of vertical wind and propagated its errors to the measured turbulent flux. The parameter ε_{ψ} significantly affected the measurement of horizontal wind, and in the measurement of turbulent flux, its effect on latent heat flux was somewhat more pronounced than for other fluxes. Therefore, these two parameters need to be carefully calibrated.

Lastly, it should be noted that the accuracy of the measured geo-referenced wind vector and turbulent flux from the UAV-based EC system is subject to the combination of many factors, mainly including sensor accuracy, UAV powerplant, UAV fluctuation (e.g., variation of the UAV attitude and flight height), and the atmospheric conditions during the measurements, etc. This study mainly focused on assessing the effects of sensor precision and UAV powerplant on the measurement errors of geo-referenced wind vector and turbulent flux. Evaluation results gave the lowest limit of the wind vector and turbulent flux that the UAV-based EC system can measure, this study could not evaluate the accuracy of the measured geo-referenced wind vector and turbulent flux by the UAV-based EC system compared to their actual true value. AnAnother effective way to evaluate the measurement accuracy of this new technique is by comparing measured values with those from the traditionally recognized measurement. However, the direct comparison of flux measurements between aircraft and traditional ground tower is still challenging due to the difference in the measurement height, mechanism (time series for ground EC and space series for aircraft), and instruments (e.g., wind sensor). Previous studies have extensively compared the measurement of fluxes and wind vector between airborne and ground-based EC methods and found consistent results (Gioli et al., 2004; Metzger et al., 2012; Sun et al., 2021b). At the same time, substantial and consistent over- or underestimation of the measured wind and fluxes by UAV compared to ground measurements were observed and reported. These differences may

be due to several factors such as vertical flux divergence (the measurement height of UAV is higher than ground-tower), surface heterogeneity (induced by the larger footprint region of the UAV compared to the ground tower), measurement errors (e.g., window length, resonance noise, etc.) as well as their difference in platform and sensors. Therefore, it is necessary to conduct a comparison test on the same platform and under the same environment to exclude the influence of these factors. Inspired by Reineman et al. (2013), future work can include developing a ground-vehicle-based UAV flux validation platform. This platform could carry both the UAV-based and traditional ground EC system to assess the measurement accuracy of the UAV-based EC system with the measurement of ground EC as the benchmark in a flight-like scenario.

The main objective of this study was to quantitatively evaluate the performance of the developed UAV-based EC system in

5 Conclusions and further works

the measurement of geo-referenced wind vector and turbulent flux. In terms of measuring precision, turbulence measurements from the UAV-based EC system were achieved with sufficient precision to enable reliable measurement of geo-referenced wind and EC flux. Magnitudes larger than 0.2-7 m s⁻¹ for wind velocity, 0.4-3 µmol m⁻² s⁻¹ for CO₂ flux, 80.052 W m⁻² for sensible heat flux, and 0.8 W m⁻² for and latent heat flux, and 0.1 m s⁻¹ for friction velocity could be reliably measured by the UAV-based EC system by assuming the minimum required signal-to-noise ratio of 105:1 for EC application. Based on the data from the calibration flight, the carefully calibrated offset angle in pitch (ϵ_{θ}) and heading (ϵ_{ψ}) were shown to effectively improve the quality of wind field measurements, and the influences of flow distortion and the leverage effect on the wind measurement were minimal and could be ignored. The influence of resonance noise was small on the measurement of air temperature and water vapor (typically < 1 % for their variance and flux covariance), but relatively large on the measurement of CO₂ (around 5 % for variance and flux covariance).

The relevance of the calibration parameters (ϵ_{r} , ϵ_{ϕ} , ϵ_{ψ} , ϵ_{θ}) for the measurement of the geo-referenced wind vector and turbulent flux was also assessed based on two sensitivity tests. uncertainties in the external calibration parameters (ϵ_{r} , ϵ_{ϕ} , ϵ_{ψ} , ϵ_{θ}) to the computation of geo-referenced wind vector and turbulent flux was also assessed based on a sensitivity test. The measurements of the geo-referenced wind vector and turbulent flux were insensitive to the errors in the ϵ_{r} and ϵ_{ϕ} .

turbulent flux. Therefore, these two calibration parameters need to be carefully calibrated. Conducting the UAV-based EC measurement when wind velocity is larger than $\frac{1}{2}$ m s⁻¹ can led to more stable and reliable (RE < 20%) results of the wind speed measurement compared to a relatively windless environmental ($\frac{1}{2}$ m s⁻¹).

The uUncertainties in the calibration parameter ε_{θ} and ε_{ψ} had the strongest effects on the measurements. Because of ε_{θ}

directly effects determining the magnitude of the vertical wind, its error will ereate lead to inaccuracies uncertainties in

vertical wind measurement and then propagate the errors uncertainties to the measurement of turbulent flux. The

uncertainties in ε_{ψ} have a direct effect on the measurement of horizontal wind, and to some extent, the measurement of

Finally, we concluded that Tthe developed UAV-based EC system measured the geo-referenced wind vector and turbulent flux with sufficient precision. The lift-induced upwash and leverage effect had almost no effect on the measurement of georeferenced wind vector. The resonance effect caused by the operation of engine and propeller mainly affected the measurement of CO₂, and its effect on variance and flux covariance was around 5 %. The quality of calibration parameters ε_{ib} and ε_{θ} has a significant effect on the measurement of the geo-referenced wind vector and turbulent flux, underscoring the importance of careful calibration. Although UAV-based EC measurements have many advantages over manned aircraft and tower-based EC measurements, airborne EC measurements themselves have some shortcomings, such as flux results hard to interpret (e.g., influence from surface heterogeneity, flux divergence, etc.), the measurements are restricted to short periods of time, and the interaction between the UAV and turbulence. Future researches may include the development of a new generation UAV-based EC system with the following improvements: 1) a new electro-powered UAV platform with the advantages of being quieter (low noise), having a low cruising speed, and being easy to operate; 2) a ground-vehicle-based validation platform to enable direct comparative evaluation of the UAV-based EC system with traditional ground EC methods under near-identical environmental conditions; 3) a graphics based real-time monitoring system to make it possible to change the flight pattern according to real-time data; and 4) a number of integrated field observation experiments that combining ground-based EC networks, OMS, and multi-source satellite RS to further prompt the development of theory and methodology for scaling transformation. -Ultimately, the versatility of the UAV-based EC system as a low cost and widely applicable environmental research aircraft facilitates further improving our understanding of the energy and matter cycling processes at regional scales.

891 892

893

894

895

874

875

876

877

878

879

880

881

882

883

884

885

886

887

888

889

890

- **Author contributions.** SY, GB and LX planned the field campaign; SY, LB, JJ, ZZ and JS carried out the field measurements. SY-and, LS_and XZ analysed the data and wrote the manuscript draft. XZ, SB, and QZ reviewed and edited the manuscript.
- 896 *Competing interests.* The authors declare that they have no conflict of interest.
- 897 Acknowledgments. This work was supported by the National Natural Science Foundation of China (Grant No. 42101477).
- 898 We would like to thank F-EYE UAV Technology Co. Ltd. for building, maintaining, and operating the UAV in this study.
- Data availability. Data for this research are not publicly available due to its proprietary nature currently. The UAV calibration flight data and the standard operation flight data in this study are available upon request to the corresponding author.

902 References

- 903 Anderson, K. and Gaston, K. J.: Lightweight unmanned aerial vehicles will revolutionize spatial ecology, Frontiers in
- 904 Ecology and the Environment, 11, 138-146, https://doi.org/10.1890/120150, 2013.
- 905 Bange, J. and Roth, R.: Helicopter-Borne Flux Measurements in the Nocturnal Boundary Layer Over Land a Case Study,
- 906 Boundary-Layer Meteorology, 92, 295-325, 10.1023/A:1002078712313, 1999.
- 907 Båserud, L., Reuder, J., Jonassen, M. O., Kral, S. T., Paskyabi, M. B., and Lothon, M.: Proof of concept for turbulence
- 908 measurements with the RPAS SUMO during the BLLAST campaign, Atmos. Meas. Tech., 9, 4901-4913, 10.5194/amt-9-
- 909 4901-2016, 2016.
- 910 Billesbach, D. P.: Estimating uncertainties in individual eddy covariance flux measurements: A comparison of methods and
- 911 a proposed new method, Agricultural and Forest Meteorology, 151, 394-405,
- 912 https://doi.org/10.1016/j.agrformet.2010.12.001, 2011.
- 913 Chandra, N., Patra, P. K., Niwa, Y., Ito, A., Iida, Y., Goto, D., Morimoto, S., Kondo, M., Takigawa, M., Hajima, T., and
- 914 Watanabe, M.: Estimated regional CO2 flux and uncertainty based on an ensemble of atmospheric CO2 inversions, Atmos.
- 915 Chem. Phys., 22, 9215-9243, 10.5194/acp-22-9215-2022, 2022.
- 916 Chen, J. M., Leblanc, S. G., Cihlar, J., Desjardins, R. L., and MacPherson, J. I.: Extending aircraft- and tower-based CO2
- 917 flux measurements to a boreal region using a Landsat thematic mapper land cover map, Journal of Geophysical Research:
- 918 Atmospheres, 104, 16859-16877, https://doi.org/10.1029/1999JD900129, 1999.
- 919 Chen, W., Wang, D., Huang, Y., Chen, L., Zhang, L., Wei, X., Sang, M., Wang, F., Liu, J., and Hu, B.: Monitoring and
- 920 analysis of coastal reclamation from 1995-2015 in Tianjin Binhai New Area, China, Scientific Reports, 7, 3850,
- 921 10.1038/s41598-017-04155-0, 2017.
- 922 Crawford, T. L. and Dobosy, R. J.: A sensitive fast-response probe to measure turbulence and heat flux from any airplane,
- 923 Boundary-Layer Meteorology, 59, 257-278, 10.1007/BF00119816, 1992.
- 924 Crawford, T. L., Dobosy, R. J., and Dumas, E. J.: Aircraft wind measurement considering lift-induced upwash, Boundary-
- 925 Layer Meteorology, 80, 79-94, 10.1007/BF00119012, 1996.
- 926 Desjardins, R. L., Brach, E. J., Alvo, P., and Schuepp, P. H.: Aircraft Monitoring of Surface Carbon Dioxide Exchange,
- 927 Science, 216, 733-735, 10.1126/science.216.4547.733, 1982.
- 928 Desjardins, R. L., Worth, D. E., MacPherson, J. I., Bastian, M., and Srinivasan, R.: Flux measurements by the NRC Twin
- 929 Otter atmospheric research aircraft: 1987–2011, Adv. Sci. Res., 13, 43-49, 10.5194/asr-13-43-2016, 2016.
- 930 Drüe, C. and Heinemann, G.: A Review and Practical Guide to In-Flight Calibration for Aircraft Turbulence Sensors, Journal
- 931 of Atmospheric and Oceanic Technology, 30, 2820-2837, 10.1175/JTECH-D-12-00103.1, 2013.
- 932 Elston, J., Argrow, B., Stachura, M., Weibel, D., Lawrence, D., and Pope, D.: Overview of Small Fixed-Wing Unmanned
- 933 Aircraft for Meteorological Sampling, Journal of Atmospheric and Oceanic Technology, 32, 97-115, 10.1175/JTECH-D-13-
- 934 00236.1, 2015.

- 935 Enriquez, A. G. and Friehe, C. A.: Effects of Wind Stress and Wind Stress Curl Variability on Coastal Upwelling, Journal of
- 936 Physical Oceanography, 25, 1651-1671, https://doi.org/10.1175/1520-0485(1995)025<1651:EOWSAW>2.0.CO;2, 1995.
- 937 Finkelstein, P. L. and Sims, P. F.: Sampling error in eddy correlation flux measurements, Journal of Geophysical Research:
- 938 Atmospheres, 106, 3503-3509, https://doi.org/10.1029/2000JD900731, 2001.
- 939 Garman, K. E., Wyss, P., Carlsen, M., Zimmerman, J. R., Stirm, B. H., Carney, T. Q., Santini, R., and Shepson, P. B.: The
- 940 Contribution of Variability of Lift-induced Upwash to the Uncertainty in Vertical Winds Determined from an Aircraft
- 941 Platform, Boundary-Layer Meteorology, 126, 461-476, 10.1007/s10546-007-9237-y, 2008.
- 942 Garman, K. E., Hill, K. A., Wyss, P., Carlsen, M., Zimmerman, J. R., Stirm, B. H., Carney, T. Q., Santini, R., and Shepson,
- 943 P. B.: An Airborne and Wind Tunnel Evaluation of a Wind Turbulence Measurement System for Aircraft-Based Flux
- 944 Measurements, Journal of Atmospheric and Oceanic Technology, 23, 1696-1708, 10.1175/JTECH1940.1, 2006.
- 945 Gioli, B., Miglietta, F., Vaccari, F. P., and Zaldei, A.: The Sky Arrow ERA, an innovative airborne platform to monitor mass,
- 946 momentum and energy exchange of ecosystems. Annals of Geophysics, 49, 109-116, 10.4401/ag-3159, 2006.
- 947 Gioli, B., Miglietta, F., De Martino, B., Hutjes, R. W. A., Dolman, H. A. J., Lindroth, A., Schumacher, M., Sanz, M. J.,
- 948 Manca, G., Peressotti, A., and Dumas, E. J.: Comparison between tower and aircraft-based eddy covariance fluxes in five
- 949 European regions, Agricultural and Forest Meteorology, 127, 1-16, https://doi.org/10.1016/j.agrformet.2004.08.004, 2004.
- 950 Hannun, R. A., Wolfe, G. M., Kawa, S. R., Hanisco, T. F., Newman, P. A., Alfieri, J. G., Barrick, J., Clark, K. L., DiGangi, J.
- 951 P., Diskin, G. S., King, J., Kustas, W. P., Mitra, B., Noormets, A., Nowak, J. B., Thornhill, K. L., and Vargas, R.: Spatial
- 952 heterogeneity in CO2, CH4, and energy fluxes: insights from airborne eddy covariance measurements over the Mid-Atlantic
- 953 region, Environmental Research Letters, 15, 035008, 10.1088/1748-9326/ab7391, 2020.
- 954 Hu, G. and Jia, L.: Monitoring of Evapotranspiration in a Semi-Arid Inland River Basin by Combining Microwave and
- 955 Optical Remote Sensing Observations, Remote Sensing, 7, 10.3390/rs70303056, 2015.
- 956 Kaimal, J. C., Clifford, S. F., and Lataitis, R. J.: Effect of finite sampling on atmospheric spectra, Boundary-Layer
- 957 Meteorology, 47, 337-347, 10.1007/BF00122338, 1989.
- 958 Kalogiros, J. A. and Wang, Q.: Aerodynamic Effects on Wind Turbulence Measurements with Research Aircraft, Journal of
- 959 Atmospheric and Oceanic Technology, 19, 1567-1576, 10.1175/1520-0426(2002)019<1567:AEOWTM>2.0.CO;2, 2002.
- 960 Khelif, D., Burns, S. P., and Friehe, C. A.: Improved Wind Measurements on Research Aircraft, Journal of Atmospheric and
- 961 Oceanic Technology, 16, 860-875, 10.1175/1520-0426(1999)016<0860:IWMORA>2.0.CO;2, 1999.
- 962 Kirby, S., Dobosy, R., Williamson, D., and Dumas, E.: An aircraft-based data analysis method for discerning individual
- 963 fluxes in a heterogeneous agricultural landscape, Agricultural and Forest Meteorology, 148, 481-489.
- 964 https://doi.org/10.1016/j.agrformet.2007.10.011, 2008.
- 965 Kowalski, A. S., Serrano-Ortiz, P., Miranda-García, G., and Fratini, G.: Disentangling Turbulent Gas Diffusion from Non-
- 966 diffusive Transport in the Boundary Layer, Boundary-Layer Meteorology, 179, 347-367, 10.1007/s10546-021-00605-5,
- 967 2021.

- 968 Lenschow, D. H.: Aircraft Measurements in the Boundary Layer, in: Probing the Atmospheric Boundary Layer, edited by:
- 969 Lenschow, D. H., American Meteorological Society, Boston, MA, Boston, https://doi.org/10.1007/978-1-944970-14-7 5,
- 970 1986.
- 971 Lenschow, D. H. and Sun, J.: The spectral composition of fluxes and variances over land and sea out to the mesoscale,
- 972 Boundary-Layer Meteorology, 125, 63-84, 10.1007/s10546-007-9191-8, 2007.
- 973 Lenschow, D. H., Delany, A. C., Stankov, B. B., and Stedman, D. H.: Airborne measurements of the vertical flux of ozone in
- 974 the boundary layer, Boundary-Layer Meteorology, 19, 249-265, 10.1007/BF00117223, 1980.
- 975 Li, X., Liu, S., Xiao, Q., Ma, M., Jin, R., Che, T., Wang, W., Hu, X., Xu, Z., Wen, J., and Wang, L.: A multiscale dataset for
- 976 understanding complex eco-hydrological processes in a heterogeneous oasis system, Scientific Data, 4, 170083,
- 977 10.1038/sdata.2017.83, 2017.
- 978 Li, X., Liu, S., Yang, X., Ma, Y., He, X., Xu, Z., Xu, T., Song, L., Zhang, Y., Hu, X., Ju, Q., and Zhang, X.: Upscaling
- 979 Evapotranspiration from a Single-Site to Satellite Pixel Scale, 10.3390/rs13204072, 2021.
- 980 Li, X., Liu, S., Li, H., Ma, Y., Wang, J., Zhang, Y., Xu, Z., Xu, T., Song, L., Yang, X., Lu, Z., Wang, Z., and Guo, Z.:
- 981 Intercomparison of Six Upscaling Evapotranspiration Methods: From Site to the Satellite Pixel, Journal of Geophysical
- 982 Research: Atmospheres, 123, 6777-6803, https://doi.org/10.1029/2018JD028422, 2018.
- 983 Liu, H., Randerson, J. T., Lindfors, J., Massman, W. J., and Foken, T.: Consequences of Incomplete Surface Energy Balance
- 984 Closure for CO2 Fluxes from Open-Path CO2/H2O Infrared Gas Analysers, Boundary-Layer Meteorology, 120, 65-85,
- 985 10.1007/s10546-005-9047-z, 2006.
- 986 Liu, J., Chen, J. M., Cihlar, J., and Chen, W.: Net primary productivity distribution in the BOREAS region from a process
- 987 model using satellite and surface data, Journal of Geophysical Research: Atmospheres, 104, 27735-27754,
- 988 https://doi.org/10.1029/1999JD900768, 1999.
- 989 Liu, S., Xu, Z., Song, L., Zhao, Q., Ge, Y., Xu, T., Ma, Y., Zhu, Z., Jia, Z., and Zhang, F.: Upscaling evapotranspiration
- 990 measurements from multi-site to the satellite pixel scale over heterogeneous land surfaces, Agricultural and Forest
- 991 Meteorology, 230-231, 97-113, https://doi.org/10.1016/j.agrformet.2016.04.008, 2016.
- 992 Liu, S., Li, X., Xu, Z., Che, T., Xiao, Q., Ma, M., Liu, Q., Jin, R., Guo, J., Wang, L., Wang, W., Qi, Y., Li, H., Xu, T., Ran,
- 993 Y., Hu, X., Shi, S., Zhu, Z., Tan, J., Zhang, Y., and Ren, Z.: The Heihe Integrated Observatory Network: A Basin-Scale
- 994 Land Surface Processes Observatory in China, Vadose Zone Journal, 17, 180072, https://doi.org/10.2136/vzj2018.04.0072,
- 995 2018.
- 996 Mahrt, L.: Flux Sampling Errors for Aircraft and Towers, Journal of Atmospheric and Oceanic Technology, 15, 416-429,
- 997 https://doi.org/10.1175/1520-0426(1998)015<0416:FSEFAA>2.0.CO;2, 1998.
- 998 Massman, W. and Clement, R.: Uncertainty in Eddy Covariance Flux Estimates Resulting from Spectral Attenuation, in:
- 999 Handbook of Micrometeorology: A Guide for Surface Flux Measurement and Analysis, edited by: Lee, X., Massman, W.,
- 1000 and Law, B., Springer Netherlands, Dordrecht, 67-99, 10.1007/1-4020-2265-4 4, 2005.

- 1001 Mathez, E. and Smerdon, J.: Climate Change3. Ocean- Atmosphere Interactions, in: The Science of Global Warming and
- 1002 Our Energy Future, Columbia University Press, 69-100, doi:10.7312/math17282-005, 2018.
- 1003 Mauder, M., Cuntz, M., Drüe, C., Graf, A., Rebmann, C., Schmid, H. P., Schmidt, M., and Steinbrecher, R.: A strategy for
- 1004 quality and uncertainty assessment of long-term eddy-covariance measurements, Agricultural and Forest Meteorology, 169,
- 1005 122-135, https://doi.org/10.1016/j.agrformet.2012.09.006, 2013.
- 1006 Mayer, J., Mayer, M., Haimberger, L., and Liu, C.: Comparison of Surface Energy Fluxes from Global to Local Scale,
- 1007 Journal of Climate, 35, 4551-4569, 10.1175/JCLI-D-21-0598.1, 2022.
- 1008 Metzger, S., Junkermann, W., Butterbach-Bahl, K., Schmid, H. P., and Foken, T.: Measuring the 3-D wind vector with a
- 1009 weight-shift microlight aircraft, Atmos. Meas. Tech., 4, 1421-1444, 10.5194/amt-4-1421-2011, 2011.
- 1010 Metzger, S., Junkermann, W., Mauder, M., Beyrich, F., Butterbach-Bahl, K., Schmid, H. P., and Foken, T.: Eddy-covariance
- flux measurements with a weight-shift microlight aircraft, Atmos. Meas. Tech., 5, 1699-1717, 10.5194/amt-5-1699-2012,
- 1012 2012.
- 1013 Mohan, M. M. P., Rajitha, K., and Murari, R. R. V.: Review of approaches for the estimation of sensible heat flux in remote
- sensing-based evapotranspiration models, Journal of Applied Remote Sensing, 14, 1-31, 10.1117/1.JRS.14.041501, 2020.
- 1015 Peltola, O., Aslan, T., Ibrom, A., Nemitz, E., Rannik, Ü., and Mammarella, I.: The high-frequency response correction of
- 1016 eddy covariance fluxes Part 1: An experimental approach and its interdependence with the time-lag estimation, Atmos.
- 1017 Meas. Tech., 14, 5071-5088, 10.5194/amt-14-5071-2021, 2021.
- 1018 Prudden, S., Fisher, A., Marino, M., Mohamed, A., Watkins, S., and Wild, G.: Measuring wind with Small Unmanned
- 1019 Aircraft Systems, Journal of Wind Engineering and Industrial Aerodynamics, 176, 197-210,
- 1020 https://doi.org/10.1016/j.jweia.2018.03.029, 2018.
- 1021 Prueger, J. H., Hatfield, J. L., Parkin, T. B., Kustas, W. P., Hipps, L. E., Neale, C. M. U., MacPherson, J. I., Eichinger, W. E.,
- and Cooper, D. I.: Tower and Aircraft Eddy Covariance Measurements of Water Vapor, Energy, and Carbon Dioxide Fluxes
- 1023 during SMACEX, Journal of Hydrometeorology, 6, 954-960, 10.1175/JHM457.1, 2005.
- 1024 Rannik, Ü., Peltola, O., and Mammarella, I.: Random uncertainties of flux measurements by the eddy covariance technique,
- 1025 Atmos. Meas. Tech., 9, 5163-5181, 10.5194/amt-9-5163-2016, 2016.
- 1026 Reineman, B. D., Lenain, L., Statom, N. M., and Melville, W. K.: Development and Testing of Instrumentation for UAV-
- 1027 Based Flux Measurements within Terrestrial and Marine Atmospheric Boundary Layers, Journal of Atmospheric and
- 1028 Oceanic Technology, 30, 1295-1319, 10.1175/JTECH-D-12-00176.1, 2013.
- 1029 Reuder, J., Båserud, L., Jonassen, M. O., Kral, S. T., and Müller, M.: Exploring the potential of the RPA system SUMO for
- multipurpose boundary-layer missions during the BLLAST campaign, Atmos. Meas. Tech., 9, 2675-2688, 10.5194/amt-9-
- 1031 2675-2016, 2016.
- 1032 Sayres, D. S., Dobosy, R., Healy, C., Dumas, E., Kochendorfer, J., Munster, J., Wilkerson, J., Baker, B., and Anderson, J. G.:
- 1033 Arctic regional methane fluxes by ecotope as derived using eddy covariance from a low-flying aircraft, Atmos. Chem. Phys.,
- 1034 17, 8619-8633, 10.5194/acp-17-8619-2017, 2017.

- 1035 Serrano-Ortiz, P., Kowalski, A. S., Domingo, F., Ruiz, B., and Alados-Arboledas, L.: Consequences of Uncertainties in CO2
- 1036 Density for Estimating Net Ecosystem CO2 Exchange by Open-path Eddy Covariance, Boundary-Layer Meteorology, 126,
- 1037 209-218, 10.1007/s10546-007-9234-1, 2008.
- 1038 Sun, Y., Jia, L., Chen, Q., and Zheng, C.: Optimizing Window Length for Turbulent Heat Flux Calculations from Airborne
- 1039 Eddy Covariance Measurements under Near Neutral to Unstable Atmospheric Stability Conditions, Remote Sensing, 10,
- 1040 10.3390/rs10050670, 2018.
- 1041 Sun, Y., Ma, J., Sude, B., Lin, X., Shang, H., Geng, B., Diao, Z., Du, J., and Quan, Z.: A UAV-Based Eddy Covariance
- 1042 System for Measurement of Mass and Energy Exchange of the Ecosystem: Preliminary Results, Sensors, 21,
- 1043 10.3390/s21020403, 2021a.
- 1044 Sun, Y., Sude, B., Geng, B., Ma, J., Lin, X., Hao, Z., Jing, W., Chen, Q., and Quan, Z.: Observation of the winter regional
- 1045 evaporative fraction using a UAV-based eddy covariance system over wetland area, Agricultural and Forest Meteorology,
- 1046 310, 108619, https://doi.org/10.1016/j.agrformet.2021.108619, 2021b.
- 1047 Thomas, R. M., Lehmann, K., Nguyen, H., Jackson, D. L., Wolfe, D., and Ramanathan, V.: Measurement of turbulent water
- vapor fluxes using a lightweight unmanned aerial vehicle system, Atmos. Meas. Tech., 5, 243-257, 10.5194/amt-5-243-2012,
- 1049 2012.
- 1050 van den Kroonenberg, A., Martin, T., Buschmann, M., Bange, J., and Vörsmann, P.: Measuring the Wind Vector Using the
- 1051 Autonomous Mini Aerial Vehicle M2AV, Journal of Atmospheric and Oceanic Technology, 25, 1969-1982,
- 1052 10.1175/2008JTECHA1114.1, 2008.
- 1053 van den Kroonenberg, A. C., Martin, S., Beyrich, F., and Bange, J.: Spatially-Averaged Temperature Structure Parameter
- 1054 Over a Heterogeneous Surface Measured by an Unmanned Aerial Vehicle, Boundary-Layer Meteorology, 142, 55-77,
- 1055 10.1007/s10546-011-9662-9, 2012.
- 1056 Vellinga, O. S., Dobosy, R. J., Dumas, E. J., Gioli, B., Elbers, J. A., and Hutjes, R. W. A.: Calibration and Quality Assurance
- 1057 of Flux Observations from a Small Research Aircraft*, Journal of Atmospheric and Oceanic Technology, 30, 161-181,
- 1058 10.1175/JTECH-D-11-00138.1, 2013.
- 1059 Vellinga, O. S., Gioli, B., Elbers, J. A., Holtslag, A. A. M., Kabat, P., and Hutjes, R. W. A.: Regional carbon dioxide and
- 1060 energy fluxes from airborne observations using flight-path segmentation based on landscape characteristics, Biogeosciences,
- 1061 7, 1307-1321, 10.5194/bg-7-1307-2010, 2010.
- 1062 Wang, H., Jia, G., Zhang, A., and Miao, C.: Assessment of Spatial Representativeness of Eddy Covariance Flux Data from
- 1063 Flux Tower to Regional Grid, Remote Sensing, 8, 742, 2016.
- 1064 Webb, E. K., Pearman, G. I., and Leuning, R.: Correction of flux measurements for density effects due to heat and water
- 1065 vapour transfer, Quarterly Journal of the Royal Meteorological Society, 106, 85-100,
- 1066 https://doi.org/10.1002/gi.49710644707, 1980.

- 1067 Williams, A. and Marcotte, D.: Wind Measurements on a Maneuvering Twin-Engine Turboprop Aircraft Accounting for
- 1068 Flow Distortion, Journal of Atmospheric and Oceanic Technology, 17, 795-810, 10.1175/1520-
- 1069 0426(2000)017<0795:WMOAMT>2.0.CO;2, 2000.
- 1070 Witte, B. M., Singler, R. F., and Bailey, S. C. C.: Development of an Unmanned Aerial Vehicle for the Measurement of
- 1071 Turbulence in the Atmospheric Boundary Layer, Atmosphere, 8, 10.3390/atmos8100195, 2017.
- 1072 Wolfe, G. M., Kawa, S. R., Hanisco, T. F., Hannun, R. A., Newman, P. A., Swanson, A., Bailey, S., Barrick, J., Thornhill, K.
- 1073 L., Diskin, G., DiGangi, J., Nowak, J. B., Sorenson, C., Bland, G., Yungel, J. K., and Swenson, C. A.: The NASA Carbon
- 1074 Airborne Flux Experiment (CARAFE): instrumentation and methodology, Atmos. Meas. Tech., 11, 1757-1776,
- 1075 10.5194/amt-11-1757-2018, 2018.
- 1076 Xu, S., Xu, S., Zhou, Y., Yue, S., Zhang, X., Gu, R., Zhang, Y., Qiao, Y., and Liu, M.: Long-Term Changes in the Unique
- 1077 and Largest Seagrass Meadows in the Bohai Sea (China) Using Satellite (1974-2019) and Sonar Data: Implication for
- 1078 Conservation and Restoration, Remote Sensing, 13, 856, 10.3390/rs13050856, 2021.
- 1079 Yang, X., Yong, B., Ren, L., Zhang, Y., and Long, D.: Multi-scale validation of GLEAM evapotranspiration products over
- 1080 China via ChinaFLUX ET measurements, International Journal of Remote Sensing, 38, 5688-5709,
- 1081 10.1080/01431161.2017.1346400, 2017.

1088

- 1082 Zhang, G., Zhang, J., and Meng, P.: Estimation of kilometer-scale heat fluxes over a hilly area in Northern China using an
- 1083 optical-microwave scintillometer, Agricultural Water Management, 244, 106582,
- 1084 https://doi.org/10.1016/j.agwat.2020.106582, 2021.
- 1085 Zheng, C., Liu, S., Song, L., Xu, Z., Guo, J., Ma, Y., Ju, Q., and Wang, J.: Comparison of sensible and latent heat fluxes
- 1086 from optical-microwave scintillometers and eddy covariance systems with respect to surface energy balance closure,
- Agricultural and Forest Meteorology, 331, 109345, https://doi.org/10.1016/j.agrformet.2023.109345, 2023.