



Real-world wintertime CO, N₂O and CO₂ emissions of a Central European village

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Abstract. Although small rural settlements are only minor individual sources of greenhouse gases and air pollution, their high overall quantity can significantly contribute to the total emissions of a region or country. The emissions of the rural lifestyle may be remarkably different from that of the urban and industrialized regions, but nevertheless they have been hardly studied so far. In this study, flux measurements at a tall-tower eddy covariance monitoring site and the footprint model FFP are used to determine the real-world wintertime CO, N₂O, and CO₂ emissions of a small village in western Hungary. The recorded emission densities, dominantly derived from residential heating, are 3.5 μg m⁻² s⁻¹, 0.043 μg m⁻² s⁻¹, and 72 μg m⁻² s⁻¹ for CO, N₂O, and CO₂, respectively. While the measured CO and CO₂ emissions are comparable with those calculated using the assumed energy consumption and applying the according emission factors, the nitrous oxide emission exceeds the expected value by a magnitude. This may indicate that the nitrous oxide emissions are significantly underestimated in the emission inventories, and modifications in the methodology of emission calculations are necessary. Using a 3-dimensional forward transport model, we further show that, in contrast to the flux measurements, the concentration measurements at the regional background monitoring site are only insignificantly influenced by the emissions of the nearby village.

1 Introduction

Climate change, primarily caused by the accumulation of greenhouse gases (GHG) in the atmosphere, is one of the biggest challenges humanity faces. In addition to the direct meteorological consequences, it also manifests in different economical and societal problems including food and water insecurity, migration, political crises, loss of biodiversity, etc. (IPCC, 2014). For the development of an effective climate change mitigation strategy, we need to know the amount of greenhouse gases emitted by each source. At country level, the emission is calculated based on statistical activity data and emission factors suggested by international guidelines (IPCC, 2019). However, emission inventory guidelines cannot specify emission factors for each activity and specific



40 conditions/circumstances, resulting in distortion and uncertainty of the officially reported inventory values, the
essential input of the European Union's emission control policy. The climatic consequences of the anthropogenic
GHG emission depend on the actual amount of GHG emissions rather than the emissions calculated based on
potentially uncertain emission factors, therefore it is highly desirable to validate and improve the accuracy of
emission factors. Industrial emissions can be estimated with relatively low uncertainty. Household emissions,
45 however, may vary largely depending on the available infrastructure, socio-economic conditions, and cultural
traditions, especially for small settlements. Although (mega)cities dominate the anthropogenic greenhouse gas
emission (Moran et al., 2018), the large number of small settlements with poorly constrained emissions and the
scarce direct measurements at village-environment-scale (Fachinger et al., 2021) contribute to the uncertainty of
the estimates of the total anthropogenic emission.

Since 1993, the tall-tower greenhouse gas monitoring station Hegyhátsál has been operated for the Global
50 Atmosphere Watch program of the World Meteorological Organization and the global cooperative air sampling
network of the National Oceanic and Atmospheric Administration, U.S.A., in a rural environment in Hungary. In
addition to the continuous monitoring of the concentration of direct (CO_2 , CH_4 , N_2O) and indirect (CO)
greenhouse gases, the tower is also equipped with eddy covariance (EC) systems to measure the surface-
atmosphere exchange of certain gases. The site is located as far from any major anthropogenic sources as
55 possible in the densely populated Central Europe. However, from time to time, the footprint area of the EC
systems partly covers the area of a nearby small village (Barcza et al., 2009). There is no industrial or notable
commercial activity in the village; therefore, it is an ideal place to estimate one of the most uncertain terms of the
emission inventories of small settlements: residential heating. EC measurement based emission mapping is
common in urban environments (see e.g. Rana et al. (2021), and references herein). In this study, we show that
60 long-term data series at a tall tower can be used for the determination of the emissions of a small settlement
occupying only a minor area of the footprint area of the EC system. For the derivation of the residential
emissions, the hourly nitrous oxide (N_2O), carbon monoxide (CO), and carbon dioxide (CO_2) fluxes measured at
the tower during the winter seasons (Dec-Feb) between December 2015 and February 2021 were used. To the
best of our knowledge, we are the first to attempt to apply this technique in a rural, natural environment for the
65 determination of the emission of a village covering only a small part of the footprint area of the measurements.

As footprints of the eddy covariance and **concentration** measurements differ by magnitudes (Gloor et al., 2001;
Kljun et al., 2002; Vesala et al., 2008; Barcza et al., 2009), the question arises of whether and to what extent the
emissions from the nearby village impact the concentration measurements at the monitoring station. A forward
transport model was deployed to answer this question.

70 **2 Methods and measurements**

2.1 Basic concept

The eddy covariance (EC) technique is widely used for the determination of surface-atmosphere flux of
atmospheric components within the footprint of the measurements (Franz et al., 2018; Papale, 2020). Although
the majority of EC systems are used for the monitoring of gas exchange of different ecological systems, there is
75 a growing number of EC sites used for the estimation of urban anthropogenic emissions (see e.g. Grimmond et
al. 2002; Vogt et al., 2006; Stagakis et al., 2019; Rana et al., 2021, and references herein). Usually, the emission



density in a city is not spatially homogeneous. An appropriate footprint model can help to attribute the measured flux to the emission in the different source areas.

80 The flux footprint area depends on the measurement height of the EC measuring system and the actual meteorological and surface conditions. Flux footprints of EC measurements performed on a tower of almost 100 m tall may cover an area of up to a hundred square kilometers (Barcza et al., 2009; Desai et al., 2015; Satar et al., 2016; Chi et al., 2019). Within this area, the village in focus may cover only a few percentages. Our concept assumes that both the “natural” landscape (vegetated area, i.e. agricultural fields, forests) and the built-up area (i.e. the village) are homogeneous from the point of view of emission density. The measured flux
85 (F_{measured}) is then the combination of the fluxes originating from the built-up areas (F_{village}) including the houses, farm buildings, backyards, roads and parks within the village, and those of the non-residential areas, which we call here as “natural” landscape (F_{natural}):

$$F_{\text{measured}} = \alpha * F_{\text{village}} + (1 - \alpha) * F_{\text{natural}}, \quad (1)$$

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where α is the weighted ratio of the village within the footprint area. The contribution of the surface flux to the measured one is not uniform over the footprint area (Schmid, 1994; Kljun et al., 2002; Vesala et al., 2008). The relative contribution of a surface source at a specific unit area to the measurement at the tower at each time step can be estimated by the value of the source weight function (footprint function) at the given point. The integral
95 of the footprint function over the infinite x-y plane equals one. The weighted ratio of α in Equation (1) is the integral of the footprint function over the village extent. With a suitable footprint model, if F_{natural} is known, then F_{village} , the emission density of the village within the footprint area can be calculated as follows:

$$F_{\text{village}} = (F_{\text{measured}} - (1 - \alpha) * F_{\text{natural}}) / \alpha \quad (2)$$

100

After the estimation of the emission density in the territory of the village based on the available EC flux measurements, the influence of the emission of the village on the concentration measurements at the monitoring site can be calculated using an appropriate transport model.

105 For the realization of this concept surface-atmosphere flux measurements, land cover information, as well as footprint and transport models are required.

2.2 Monitoring site and instrumentation

The eddy covariance measurements used in this study are carried out on a 117 m tall, free-standing TV/radio transmitter tower owned by Antenna Hungária Corp. It is located in a fairly flat region of western Hungary, close to the western edge of the Pannonian Basin (46°57'N, 16°39'E, 248 m asl), in the vicinity of the small village
110 called Hegyhátsál being in the focus of the present study (Fig. 1).

The eddy covariance system is mounted on the tower at 82 m above the ground. It has been monitoring the vertical flux of CO₂ since 1997, and that of N₂O and CO since 2015. The EC system is based on a GILL R3-50 research ultrasonic anemometer (GILL Instruments Ltd, Lymington, U.K.), a Model 913-0014 Enhanced Performance fast-response N₂O/CO/H₂O analyzer (Los Gatos Research Ltd., San Jose, CA, U.S.A.) with fast-

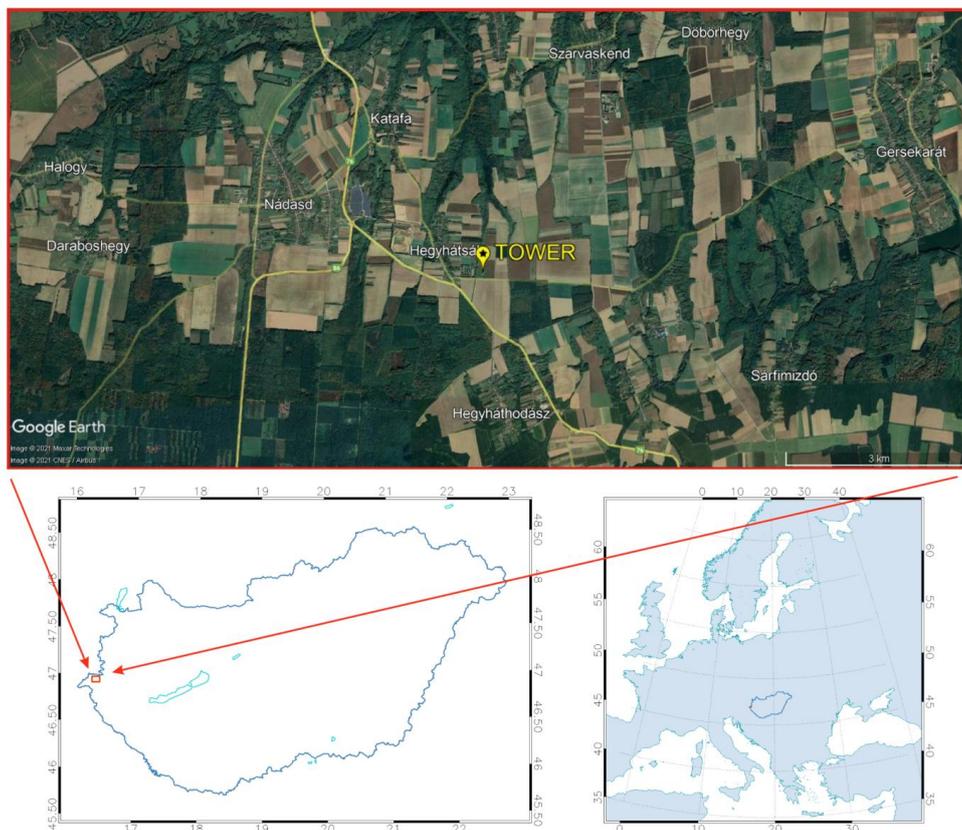


Figure 1: Location and the surrounding region of Hegyhátsál tall-tower GHG monitoring site on a Google Earth satellite image.

115 flow optional accessories, and a Model Li-6262 fast response infrared CO₂/H₂O analyzer (Li-Cor Inc., Lincoln, Nebraska, USA).

In addition to the eddy covariance measurements, the concentration of carbon dioxide and the basic meteorological parameters (wind speed and direction, air temperature, relative humidity) are continuously measured at four elevations along the tower (10 m, 48 m, 82 m, 115 m). In this paper, we use the term
120 “concentration” as a synonym of the actually measured “dry mole fraction”. For a detailed description of the site and instrumentation see Haszpra et al. (2001; 2018) and Barcza et al. (2020).

2.3 Surface-atmosphere flux calculation

The measurements performed at the tower allow the calculation of the turbulence parameters (vertical wind speed and concentration fluctuations) necessary for the calculation of the vertical fluxes of the substances studied. A detailed description of the methodology for the calculation of the surface-atmosphere flux can be
125 found in Haszpra et al. (2001), and Barcza et al. (2020). The EC system and the data evaluation software provide hourly flux values.



A disadvantage of the tall-tower EC systems is that they may be decoupled from the surface by low-level inversions from time to time. At our monitoring station, such conditions are not rare, especially during winter. In these situations, the EC systems cannot provide the actual surface-atmosphere flux data. To avoid decoupled measurements in cases of low-level inversion, information on the height of the boundary layer was used from the ERA5 reanalysis dataset of the European Centre for Medium-Range Weather Forecasts (Copernicus Climate Change Service, 2017) for the grid-point nearest to the monitoring site with hourly resolution. Taking into account the elevation of the EC system of 82 m above the ground, we removed all flux values for the periods when the top of the boundary layer was below 100 m. Removing these data from the data series, 6371 hourly data (49.0 % of the total 13008 winter hours) remained for the study for the six winter seasons (2015/2016-2020/2021) evaluated in this study.

2.4 Environmental and land cover information

The project aims at the determination of the wintertime, residential heating dominated emission of Hegyhátsál village. The village is located in the west-to-northwest sector, 400-1200 m away from the tower. It has 151 inhabitants in 89 households (Hungarian Central Statistical Office, 2019). There is no industrial or notable commercial activity in the village. Approximately half of the single-family houses of the village are connected to the natural gas distribution network and use this fuel also for heating purposes. The other half of the households use solid fuels for heating. Taking into account the socioeconomic conditions in the region, it is reasonable to assume that heating appliances for biomass or other solid fuels are available and occasionally used even in the households connected to the natural gas network.

The land cover of the region of the monitoring tower consists of a mixture of agricultural fields and small forest patches. In addition to Hegyhátsál, the other neighboring villages are about 3 km away from the tower to the north (Katafa), northwest (Nádasd), and south (Hegyháthodász). The nearest settlement worth mentioning in the eastern sector is Gersekarát, located more than 7 km from the tower (Fig. 1). There is hardly any commercial or industrial activity in this dominantly agricultural region.

The local roads connecting the small settlements carry only low traffic (300-600 vehicle units per day). The only major road in the region is the 2x1 lane trans-European E65 running northwest-southeast with 4700 vehicle units per day (Magyar Közút, 2019). Its closest point to the monitoring site is about 500 m to the southwest (Fig. 1).

The prevailing wind directions in winter are northeasterly and southwesterly (Fig. 2), although the monitoring station is located in the zone of westerly wind patterns. However, the Alps rising approximately 100 km to the west of the station significantly modify the regional wind pattern.

For the identification of the land cover type of the potential source areas of the surface-atmosphere fluxes, the National Ecosystem Base Map of Hungary (NÖSZTÉP) (Tanács et al., 2019) was used. This dataset has 56 categories at Level-3 with a spatial resolution of 20 m x 20 m. Within the area of our interest three Level 1 NÖSZTÉP categories occur: urban, cropland, and forest. In our study, cropland and forested land cover types are considered as “natural” landscape areas, while the areas labeled as urban (buildings, roads and other artificial surfaces, vegetated areas in an artificial environment [e.g. backyards, parks, etc.]) represent the villages. There are 1645 grid cells covering the area of Hegyhátsál village, which corresponds to its area of 65.8 hectares (Fig. 3).

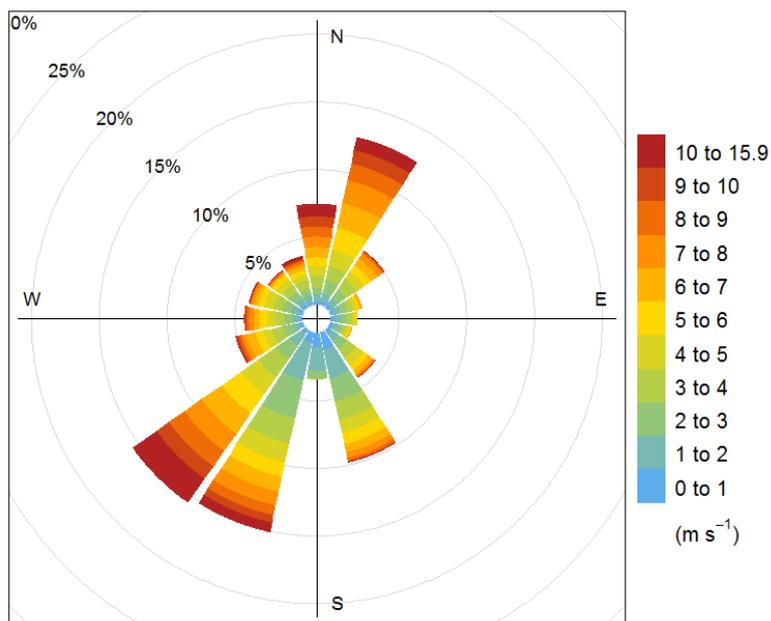


Figure 2: Wintertime (Dec-Feb) frequency distribution of wind directions at 82 m height at the Hegyhátsál tall-tower monitoring site between December 2015 and February 2021. The village is located to the northwest of the tower.



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Figure 3: Definition of the territory of the village based on land cover information. The shaded area covers the buildings, roads in the village, and vegetated areas in artificial environment (parks, backyards, etc.).



175 **2.5 Footprint model**

The flux footprint (or source weight) function is a probability function, describing the relative contribution from each element of the (mainly) upwind surface area source to the measured flux. For the calculation of the source area (footprint) of the flux measurements, the 2-dimensional Flux Footprint Prediction (FFP) model of Kljun et al. (2015) has been applied. The model is based on the LPDM-B backward Lagrangian stochastic particle dispersion model valid for a wide range of atmospheric conditions (Kljun et al., 2002; 2004a; 2004b). While FFP is much less resource-intensive than LPDM-B, it is still applicable for stable, neutral, and convective conditions. The model performance was one of the best in a test of several footprint models against data from a tracer release experiment (Heidbach et al., 2017).

The input parameters of the model are the measurement height above displacement height (z_m), the roughness length (z_0), the Obukhov length (L), the standard deviation of the lateral wind speed (σ_v), the friction velocity (u_*), and the height of the boundary layer (h). The wind direction is an optional input parameter but it is needed for the geographical localization of the source areas. Displacement height was considered to be negligible due to the lack of vegetation in wintertime (see van der Kwast et al., 2009), thus the observation height was used to approximate z_m . The Obukhov length, the standard deviation of the lateral wind speed, the friction velocity, and the wind direction are directly measured or can be calculated from the measurements. The boundary layer height is available from the ERA5 reanalysis data set (see above) for the region of the tower. For the roughness length, 0.15 m is assumed based on an earlier study (Barcza et al., 2009).

FFP assumes the stationarity and horizontal homogeneity of the flow over the time periods of the flux calculations (one hour in our case) and does not include roughness sublayer dispersion near the ground (negligible in this case) nor dispersion within the entrainment layer at the top of the convective boundary layer. The scaling parametrization also sets some limitations. In this study the model was used with the following restrictions:

$$20 \ z_0 < z_m < h_e \quad (3)$$

200 $-15.5 \leq z_m/L \quad (4)$

$$u_* \geq 0.2 \text{ m s}^{-1}, \quad (5)$$

where h_e is the height of the entrainment layer. Accepting that typically $h_e \approx 0.8h$ (Holtslag and Nieuwstadt, 1986; Kljun et al., 2015), (3) does not reduce the available flux data as the fluxes measured during $h < 100$ m have already been excluded from the data set (see above). However, (4) and (5) disqualify approximately a third of our measurement data. So, for footprint calculation, 4277 hourly flux data, 32.9 % of the total winter hours, were available. Theoretically, it is possible to set a lower u_* threshold for the FFP model. However, at low u_* the EC systems mounted high above the ground cannot provide the actual surface-atmosphere flux data alone. In such cases, the storage term has to be considered (Haszpra et al., 2005). As the storage term adds considerable uncertainty to the calculated flux due to the noisy signal, it is preferable to avoid low u_* conditions in the tall-tower flux derivation.

The discretized footprint function, i.e. the output of the model at the 90 % footprint contribution level was integrated for each grid cell of the land cover map giving the contribution of that specific grid cell to the total



flux measured at the monitoring site. The footprint function was also integrated over the area of the village to
215 indicate the total contribution of the emission from the village to the measured flux at the monitoring site.

2.6 Transport model

For the estimation of the influence of the emission from the village on the concentration measurements at the
monitoring tower, the Graz Lagrangian Model (GRAL v14.8 – Oetl et al. 2002; 2015a; 2015b; Romanov et al.,
2020) has been used. This 3D particle dispersion model was originally developed for the dispersion of pollutants
220 from a road tunnel portal but is suitable to describe the 3-dimensional concentration distribution of area sources.
Its input data are wind speed, wind direction, Pasquill-Gifford stability class (determined from the local
meteorological measurements), location of the source area relative to the receptor point, and the yield of the
source homogeneously distributed over the area. The simulation was run at 10 m horizontal and 3 m vertical
resolution.

225 3 Results and discussion

3.1 Emission of the “natural” landscape

The natural sources of carbon monoxide compose of biomass burning, atmospheric oxidation of hydrocarbons,
and direct biogenic emission (Zheng et al., 2019). Open biomass burning (e.g. stubble burning) is prohibited in
the study region. Atmospheric oxidation of hydrocarbons requires hydroxyl radicals. Hydroxyl radicals form in
230 photochemical processes, therefore their concentration is low in the darkest season of the year. The vegetation is
dormant in winter, and biogenic emissions also depend on sufficient light (Bruhn et al., 2013).

The major natural sources of nitrous oxide are denitrification and nitrification processes in soil and water (Tian
et al., 2020). These biochemical processes slow down with decreasing temperature (Benoit et al., 2015;
Butterbach-Bahl et al., 2013). Nitrogen addition to agricultural soil enhances nitrous oxide emission, which is
235 relevant in our case as the surrounding region of the tower is dominated by croplands (Barcza et al., 2009).
Although the average temperature of +1.6 °C during the study period is rather low for biochemical activities, the
croplands forming a significant part of the “natural” landscape around the monitoring site may emit a detectable
amount of N₂O.

The net ecosystem exchange of carbon dioxide in the winter season is positive in our region (Haszpra et al.,
240 2005; Barcza et al., 2020), i.e. the landscape is a net source. The dominantly dormant vegetation assimilates only
a low amount of carbon dioxide. This process might be temperature-dependent during the winter season.
Respiration is decreasing with temperature. The result of the two opposing processes, photosynthetic
assimilation and respiration, is a net emission in winter on average.

For the determination of the emission density of the non-residential landscape (agricultural fields, forests), we
245 selected those footprints where the integrals of the footprint function values over the area of the village were
negligible, that is, the emission in the village did not influence the fluxes measured at the monitoring site (at the
90 % footprint contribution level). Obviously, all these 1147 footprints cover areas in the easterly to southerly
sector, opposite the village Hegyhátsál. To avoid any contamination from remote settlements footprints with
maximum source weight location farther than 5000 m were excluded from the evaluation, which left us with
250 1120 footprints and flux data points. Nevertheless, our selection cannot completely exclude any anthropogenic



emissions. The local roads with little traffic and small settlements of a few households still contribute to the emission of the “natural” landscape”.

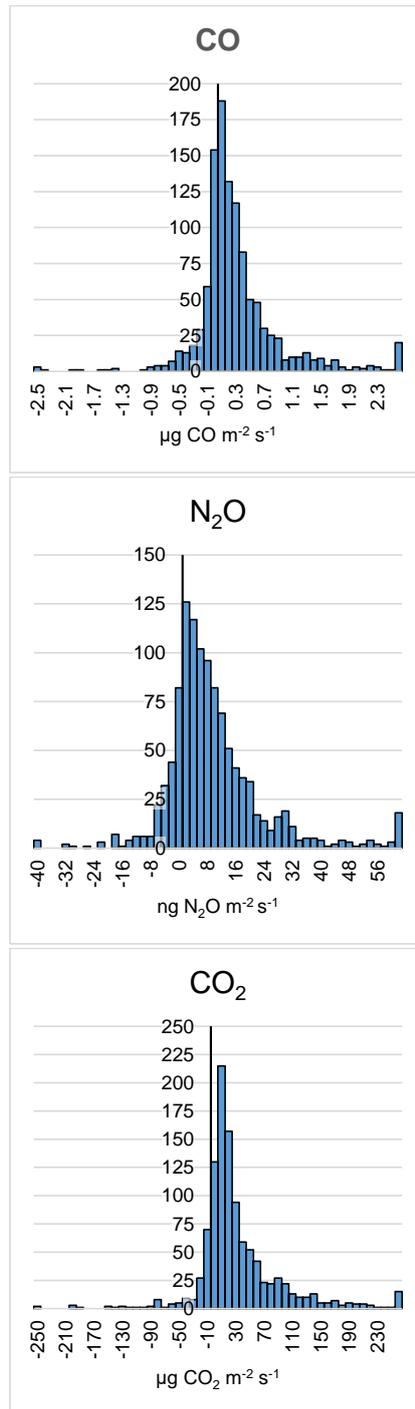
255 The frequency functions of the measured fluxes are skewed towards positive fluxes with a few extremely high values (Fig. 4). Therefore, instead of the arithmetic average, the emission density of the “natural landscape” is characterized by the median of the data sets. The medians are not sensitive to extreme outliers, hence we did not apply any arbitrary outlier filtering algorithm. The emission densities of CO, N₂O, and CO₂ obtained for the natural landscape are 139 ng m⁻² s⁻¹, 5.9 ng m⁻² s⁻¹, and 12 μg m⁻² s⁻¹, respectively. In addition to these median values, to give an impression of the uncertainty of the calculated emission densities, Table 1 also lists the lower and upper quartiles.

260 3.2 Emission from the village

In the ideal case, for the determination of the emission density of the village using the top-down approach (i.e. estimation of the emission through atmospheric measurements), we should select those cases when the footprint of the flux measurements exactly covers the area of the village, not missing any part of it and not including anything but the village itself. Due to the location, size, and shape of the village, this was, however, not possible, 265 as all footprints also included non-village contributions. Increasing the required minimum contribution of the village (the integral of the footprint function over the area of the village, α in Eq. (1) and Eq. (2)) results in a decrease in the number of the available hourly flux data points (Fig. 5). The emission density of the village has been calculated for 25 %, and 30 % footprint weighted coverage ($\alpha=0.25$ and $\alpha=0.30$) (Fig. 6). In addition to the median flux, Table 1 gives the estimated lower and upper quartiles, and the number of footprints available for the calculations. The low number of cases is also due to the prevailing wind directions that avoid the village (see 270 Fig. 2). Nevertheless, the small difference between the emission densities derived for 25 % and 30 % footprint weighted coverages indicates that these coverages can satisfactorily be used to derive the emission densities of the village. At 30 % footprint weighted coverage, the interquartile ranges are a bit narrower also suggesting a bit lower uncertainty.

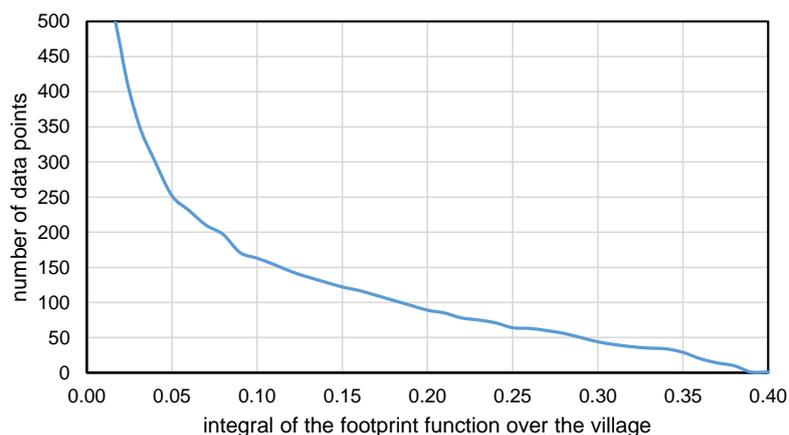
275 The correlation between the air temperature and the top-down emission density of CO, N₂O, and CO₂ is negative and statistically significant at $p<0.05$ significance level (-0.369, -0.343, and -0.480, respectively). The negative correlation between the air temperature and the emission supports that the measured flux originates dominantly from residential heating, which is more intensive at low temperatures. This assumption is further strengthened by the high positive linear correlation (+0.503) between CO and CO₂ emissions.

280 Taking into account the area of the village (65.8 ha) and the median emission densities presented for 30 % coverage in Table 1, the total wintertime (3 months) CO, N₂O, and CO₂ emissions of the village are 17.9 Mg (metric ton), 0.216 Mg, and 364 Mg, respectively (Table 2).



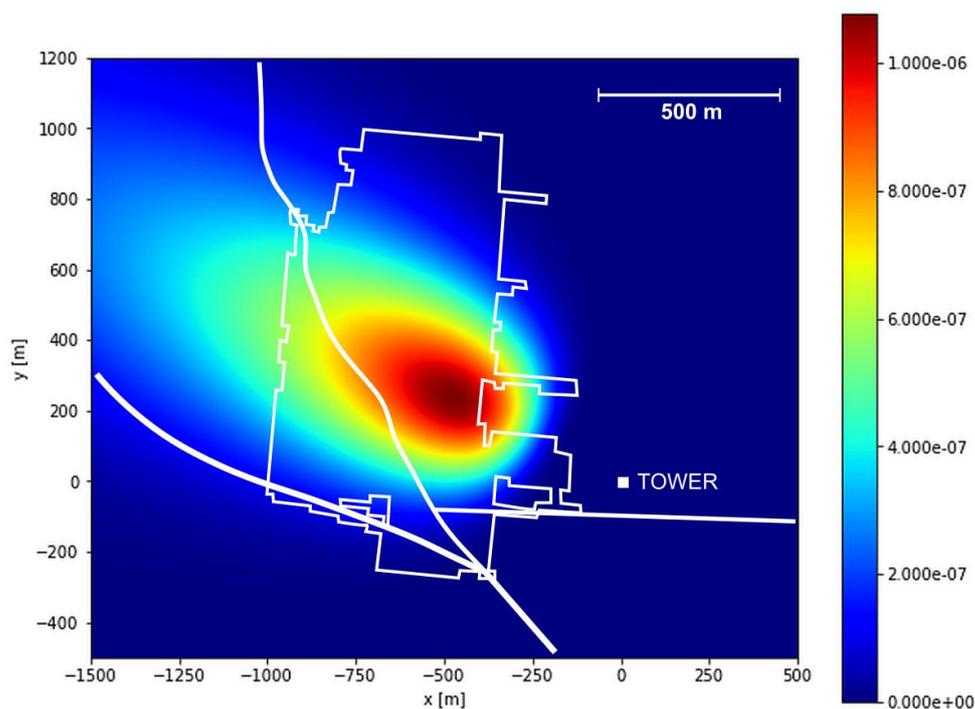
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Figure 4: Frequency distributions of wintertime “natural” landscape (agricultural fields, forests) emissions.



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Figure 5: The number of hourly data points when the integral of the footprint function over the village area was greater than or equal to the value indicated in the x-axis. At $x=0$, the total number of data points is included (4277 – outside axis range).



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Figure 6: Example of a single flux footprint (10:00-11:00 h LST, 27 December 2016). The integral of the footprint function over the village is 0.282.



300 Alternative emission estimation can be obtained using the bottom-up method, i.e. using published or expert-
 based emission factors. Most of the houses in the village are several decades old single-family houses, traditional
 brick constructions without insulation. According to the expert estimates, a house of average size may need
 approximately 57 GJ energy for winter heating (personal communication, Unit of National Emission Inventories,
 Hungarian Meteorological Service). For the 89 households of the village, this results in approximately 5 TJ. As
 305 only half of the houses are connected to the natural gas network it means that a maximum of 2.5 TJ energy may
 come from natural gas and a minimum of 2.5 TJ originates from solid fuels, respectively. (Liquid fuel is not used
 for residential heating in Hungary.) The default emission factors for natural gas are 26 kg CO TJ⁻¹,
 0.1 kg N₂O TJ⁻¹, and 56.1 Mg CO₂ TJ⁻¹, respectively, while for solid fuels they are around 4 Mg CO TJ⁻¹,
 4 kg N₂O TJ⁻¹, and 100 Mg CO₂ TJ⁻¹, respectively, depending on the actual fuel type (wood, lignite, etc.) (IPCC,
 310 2006; European Environmental Agency, 2019). Assuming these values, the overall heating emissions are
 estimated as 10 Mg, 10 kg, and 390 Mg for CO, N₂O, and CO₂, respectively. As it can be assumed that even the
 households with access to natural gas use some solid fuels for heating, the real emission values may be
 somewhat higher. In the extreme case, if no natural gas would be used at all, the corresponding values were
 20 Mg, 20 kg, and 500 Mg for CO, N₂O, and CO₂.

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**Table 1. Emission density of the natural landscape (0 % village coverage) and that of the village calculated at 25 %
 and 30 % footprint weighted coverage of the village, respectively. n gives the number of footprints available for the
 calculations. Q25, Q50, and Q75 indicate the lower quartile, the median, and the upper quartile of the emission
 320 densities calculated on the basis of the hourly flux values.**

village contribution (%) number of footprints	0 % (n=1120)	25 % (n=64)	30 % (n=44)
CO [ng m ⁻² s ⁻¹]	Q25	-14	1459
	Q50	139	3403
	Q75	405	6029
N ₂ O [ng m ⁻² s ⁻¹]	Q25	0.9	8.6
	Q50	5.9	42.7
	Q75	13.2	82.5
CO ₂ [μg m ⁻² s ⁻¹]	Q25	0	7
	Q50	12	86
	Q75	42	282

**Table 2. Winter season emissions of Hegyhátsál village calculated applying the top-down (present study) and bottom-
 up approach, and their ratios**

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	top-down (TD) estimation	bottom-up (BU) estimation	TD/BU
CO	17.9 Mg	8 Mg	2.2
N ₂ O	216 kg	8 kg	27
CO ₂	364 Mg	310 Mg	1.2



Under the given climatic conditions, the heating season starts around mid-October and lasts until mid-April. Based on the heating day distribution, 65-70 % of the heating energy is used during December-February, i.e. the study period. A conservative estimation of 10-15 % share of solid fuels in the households accessing natural gas would give the approximate emission values of 8 Mg, 8 kg, and 310 Mg for CO, N₂O, and CO₂, respectively, for the December-February period (Table 2).

These statistics-based, bottom-up numbers underestimate the CO and CO₂ emissions calculated by the top-down approach. However, taking into account the rough estimation of the bottom-up emission and the uncertainties of the top-down approach, the results are similar enough to support the applicability of our method. Our measurements show a higher CO:CO₂ emission ratio (0.049 vs. 0.026), which indicates the contribution of incomplete combustion that points to biomass rather than natural gas burning. The measurement-based (top-down) N₂O emission is 27-times higher than the statistical-based (bottom-up) one. Even if we assume the upper limit of 15 kg TJ⁻¹ for the N₂O emission factor from solid fuel, the resulting value is still a magnitude lower than the measured one. This emission factor may also indicate biomass (or even waste) burning. Our previous study based on concentration measurements (Haszpra et al., 2019) also showed a higher N₂O:CO₂ ratio relative to the official emission estimates. It suggests that the N₂O emission factor for residential heating is significantly underestimated for the actual conditions. The illegal burning of solid household, municipal or agricultural waste, which is not rare in villages of poor socioeconomic conditions (Hoffer et al., 2020), may also modify the CO:N₂O:CO₂ emission ratios. To prove or disprove the presence of waste burning would only be possible through local measurements of characteristic organics in the atmosphere.

3.3 Influence of local emissions on the regional background concentration measurements

Hegyhátsál tall tower monitoring station is registered in the WMO GAW program (<https://gawsis.meteoswiss.ch/GAWSIS/#/search/station/stationReportDetails/0-20008-0-HUN>) and the ICOS network (<https://meta.icos-cp.eu/labeling/>) as a regional background monitoring site, receiving as little direct anthropogenic pollution as possible in the densely populated, highly industrialized Europe. As the footprints of the eddy covariance measurements (see e.g. Fig. 6) and those of the concentration measurements (Fig. 7) differ significantly, it is appropriate to check by how much the concentration measurements are influenced by the nearby village. The GRAL model was applied for the 2017/2018 winter season to estimate the influence of the village's emissions on the concentration measurements. For 1798 hours of 2160 hours of the study period (83.2 %), emissions from the village did not reach the measurement sensor at all, mainly due to the prevailing wind directions. (The prevailing wind directions were northeasterly and southwesterly, while the village is located in the west-northwest sector relative to the measurement tower. See Fig. 2 and 3.) In a few cases with winds from west-northwest, i.e. from the village, the pollution could not reach the measurement elevation at 82 m above the ground due to the shallow boundary layer.

Fig. 8 shows the frequency distribution of the excess concentrations derived from the emission in the village. The excess burden given in mass per volume unit is converted into concentration given in dry mol fraction assuming standard pressure (972 hPa at 248+82 m above sea level) and air temperature of +0.6 °C (average over the period of December 2017 - February 2018). Due to the high emission relative to the background concentration, the carbon monoxide concentration is the most sensitive to local pollution. The excess concentration exceeds 2 nmol mol⁻¹ in only 0.74 % of the hours. Without considering a specific hour of the study



period, the maximum excess was $2.9 \text{ nmol mol}^{-1}$. For comparison, the recommended network compatibility of CO concentration measurements within the scope of WMO/GAW network is 2 nmol mol^{-1} (WMO, 2020).

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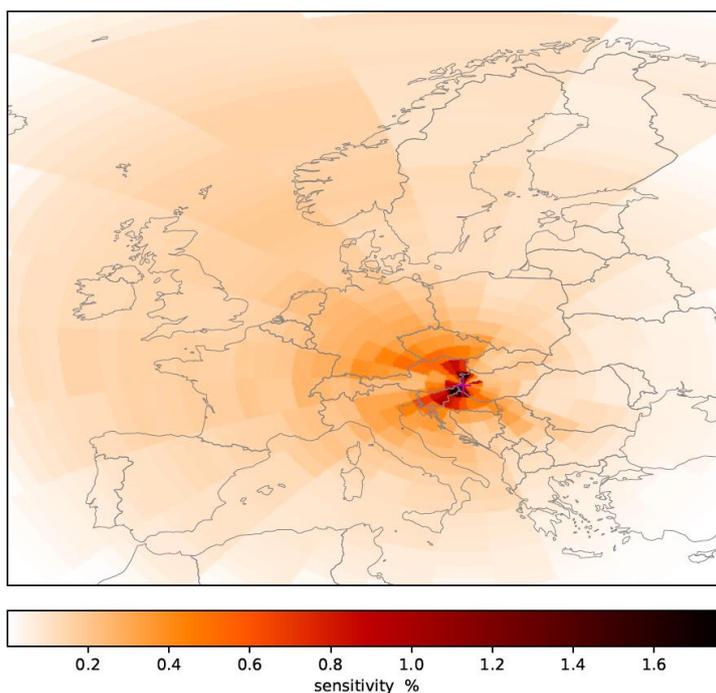


Figure 7: Footprint climatology of the concentration measurements performed at 115 m elevation above the ground in 2019 calculated by the STILT model (Source: ICOS Carbon Portal, <https://stilt.icos-cp.eu/worker/>).

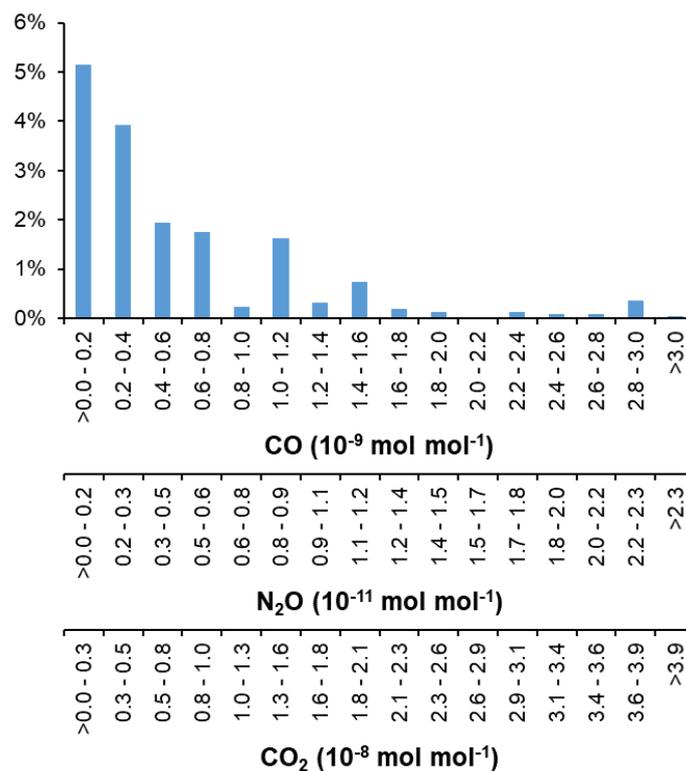
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The 'hot spot' event mentioned above occurred between 11 and 12 o'clock local time on 4 February when the light wind (0.8 m s^{-1}) from the center of the village (wind direction: 300°) directly carried pollution to the measurement sensor in an extremely unstable atmosphere (Pasquill-Gifford stability class A). This process increased the background concentration by $13.6 \text{ nmol mol}^{-1}$. This single-hour measurement has to be highlighted in the quality control process and flagged as a regionally non-representative, locally influenced event.

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Carbon dioxide and nitrous oxide behave similarly to carbon monoxide during the short transport time from the village to the tower. Consequently, their excess concentrations caused by the emission in the village are proportional to their emission densities relative to those of carbon monoxide, and the shape of the frequency distributions is the same (Fig. 7). In the case of carbon dioxide, there were only two events (hourly data points) when the excess concentration exceeded $0.04 \text{ } \mu\text{mol mol}^{-1}$ ($0.06 \text{ } \mu\text{mol mol}^{-1}$, and $0.18 \text{ } \mu\text{mol mol}^{-1}$ in the extreme case discussed above). The values are within the uncertainty of the measurements. Emissions of nitrous oxide were relatively low causing only a maximum of $0.03 \text{ nmol mol}^{-1}$ excess (the extreme value discussed above is $0.10 \text{ nmol mol}^{-1}$), which was practically undetectable.

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Figure 8: Frequency distribution of the excess concentrations at the measuring point allocated to emission sources in the village. Cases of zero excess (83.2 %) are not presented. The distribution is the same for all components but the scales are different. Note: scales for N₂O and CO₂ are rounded and based on the CO scale.

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The wintertime (Dec-Feb) average excess of CO, N₂O, and CO₂ concentrations are 0.10, <0.01, and 1.34 nmol mol⁻¹, respectively. In the lack of industrial and commercial activities, residential heating is the dominant emission source in the village. This means that the excess concentrations may be even lower in the non-winter seasons. The low values confirm the regional representativeness of the measured concentration data, i.e. that the Hegyhátsál tower qualifies as a regional background monitoring site.

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4 Summary and conclusion

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In this study, we have shown that tall-tower eddy covariance measurements can be used for the determination of the emission of a region even if it occupies only a small portion of the footprint area of the measurements. The study is presumably the first one aiming at the direct measurement of GHG emission of a small rural settlement, while similar measurements have already been performed in urban environments that exhibit different emission characteristics. The results revealed that while the statistical-based calculations of carbon monoxide and carbon dioxide emissions do not differ significantly from the real-world top-down measurements, the statistical-based nitrous oxide emission is significantly underestimated. Further in-depth studies are needed, which will possibly



410 result in a correction of the emission factors. The relatively high CO to CO₂ ratio and the high N₂O emission density suggest the higher than “officially” assumed ratio of biomass burning, and a possibility of illegal waste burning.

Using a 3D transport model, we clarified that the village, as a local pollution source, hardly influences the concentration measurements at the nearby greenhouse gas monitoring station at 82 m height. Hence the site can be qualified as a regional background monitoring site.

415 *Data and code availability.* The raw data of the calculations are available from the corresponding author. The FFP footprint model code is publicly available at <https://footprint.kljun.net/>, while the GRAL model code can be downloaded from the Technische Universität Graz (<https://gal.tugraz.at/index.php/download>).

420 *Author contributions.* Conceptualization, part of the calculations, writing the paper (L.H.); flux calculations (Z.B.); forward transport modeling (Z.F.); landscape information and processing (A.K.); footprint modeling support and editing the paper (N.K.). All authors have read and agreed to the published version of the manuscript.

425 *Competing interests.* The authors declare that they have no conflict of interest.

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References

- 440 Barcza, Z., Kern, A., Davis, K. J., and Haszpra, L.: Analysis of the 21-years long carbon dioxide flux dataset from a Central European tall tower site, *Agricultural and Forest Meteorology*, 290, 108027, <https://doi.org/10.1016/j.agrformet.2020.108027>, 2020.
- Barcza, Z., Kern, A., Haszpra, L., and Kljun, N.: Spatial representativeness of tall tower eddy covariance measurements using remote sensing and footprint analysis, *Agricultural and Forest Meteorology*, 149, 795-807, <https://doi.org/10.1016/j.agrformet.2008.10.021>, 2009.
- 445 Benoit, M., Garnier, J., and Billen, G.: Temperature dependence of nitrous oxide production of a luvisolic soil in batch experiments, *Process Biochemistry*, 50, 79-85, <https://doi.org/10.1016/j.procbio.2014.10.013>, 2015.



- Bruhn, D., Albert, K. R., Mikkelsen, T. N., and Ambus, P.: UV-induced carbon monoxide emission from living vegetation, *Biogeosciences*, 10, 7877-7882, <https://doi.org/10.5194/bg-10-7877-2013>, 2013.
- 450 Butterbach-Bahl, K., Baggs, E. M., Dannenmann, M., Kiese, R., and Zechmeister-Boltenstern, S.: Nitrous oxide emissions from soils: how well do we understand the processes and their controls?, *Philosophical Transactions of the Royal Society B: Biological Sciences*, 368, 20130122, <https://doi.org/doi:10.1098/rstb.2013.0122>, 2013.
- Chi, J., Nilsson, M. B., Kljun, N., Wallerman, J., Fransson, J. E. S., Laudon, H., Lundmark, T., and Peichl, M.: The carbon balance of a managed boreal landscape measured from a tall tower in northern Sweden, *Agricultural and Forest Meteorology*, 274, 29-41, <https://doi.org/10.1016/j.agrformet.2019.04.010>, 2019.
- 455 Copernicus Climate Change Service: ERA5: Fifth generation of ECMWF atmospheric reanalyses of the global climate, Copernicus Climate Change Service Climate Data Store (CDS), <https://cds.climate.copernicus.eu/cdsapp#!/home>, Accessed on 29 January 2020, 2017.
- Desai, A. R., Xu, K., Tian, H., Weishampel, P., Thom, J., Baumann, D., Andrews, A. E., Cook, B. D., King, J. Y., and Kolka, R.: Landscape-level terrestrial methane flux observed from a very tall tower, *Agricultural and Forest Meteorology*, 201, 61-75, <https://doi.org/10.1016/j.agrformet.2014.10.017>, 2015.
- 460 European Environmental Agency: EMEP/EEA air pollutant emission inventory guidebook 2019, EEA Report No. 13/2019, 2019.
- Fachinger, F., Drewnick, F., and Borrmann, S.: How villages contribute to their local air quality – The influence of traffic- and biomass combustion-related emissions assessed by mobile mappings of PM and its components, *Atmospheric Environment*, 263, 118648, <https://doi.org/10.1016/j.atmosenv.2021.118648>, 2021.
- Franz, D., Acosta, M., Altimir, N., Arriga, N., Arrouays, D., Aubinet, M., Aurela, M., Ayres, E., López-Ballesteros, A., Barbaste, M., Berveiller, D., Biraud, S., Boukir, H., Brown, T., Brümmner, C., Buchmann, N., Burba, G., Carrara, A., Cescatti, A., Ceschia, E., Clement, R., Cremonese, E., Crill, P., Darenova, E., Dengel, S., 470 D’Odorico, P., Filippa, G., Fleck, S., Fratini, G., Fuß, R., Gielen, B., Gogo, S., Grace, J., Graf, A., Grelle, A., Gross, P., Grünwald, T., Haapanala, S., Hehn, M., Heinesch, B., Heiskanen, J., Herbst, M., Herschlein, C., Hörtnagl, L., Hufkens, K., Ibrom, A., Jolivet, C., Joly, L., Jones, M., Kiese, R., Klemedtsson, L., Kljun, N., Klumpp, K., Kolari, P., Kolle, O., Kowalski, A., Kutsch, W., Laurila, T., de Ligne, A., Linder, S., Lindroth, A., Lohila, A., Longdoz, B., Mammarella, I., Manise, T., Jiménez, S. M., Matteucci, G., Mauder, M., Meier, P., 475 Merbold, L., Mereu, S., Metzger, S., Migliavacca, M., Mölder, M., Montagnani, L., Moureaux, C., Nelson, D., Nemitz, E., Nicolini, G., Nilsson, M. B., de Beeck, M. O., Osborne, B., Löfvenius, M. O., Pavelka, M., Peichl, M., Peltola, O., Pihlatie, M., Pitacco, A., Pokorný, R., Pumpanen, J., Ratić, C., Rebmann, C., Roland, M., Sabbatini, S., Saby, N. P. A., Saunders, M., Schmid, H. P., Schrumpf, M., Sedláč, P., Ortiz, P. S., Siebicke, L., Šigut, L., Silvennoinen, H., Simioni, G., Skiba, U., Sonntag, O., Soudani, K., Soulé, P., Steinbrecher, R., 480 Tallec, T., Thimonier, A., Tuittila, E.-S., Tuovinen, J.-P., Vestin, P., Vincent, G., Vincke, C., Vitale, D., Waldner, P., Weslien, P., Wingate, L., Wohlfahrt, G., Zahniser, M., and Vesala, T.: Towards long-term standardised carbon and greenhouse gas observations for monitoring Europe’s terrestrial ecosystems: a review, *International Agrophysics*, 32, 439-455, <https://doi.org/10.1515/intag-2017-0039>, 2018.
- Gloor, M., Bakwin, P., Hurst, D., Lock, L., Draxler, R., and Tans, P.: What is the concentration footprint of a tall 485 tower?, *J. of Geophysical Research*, 106D, 17831-17840, 2001.



- Grimmond, C. S. B., King, T. S., Cropley, F. D., Nowak, D. J., and Souch, C.: Local-scale fluxes of carbon dioxide in urban environments: methodological challenges and results from Chicago, *Environmental Pollution*, 116, S243-S254., 2002.
- 490 Haszpra, L., Ferenczi, Z., and Barcza, Z.: Estimation of greenhouse gas emission factors based on observed covariance of CO₂, CH₄, N₂O and CO mole fractions, *Environmental Sciences Europe*, 31, 95, <https://doi.org/10.1186/s12302-019-0277-y>, 2019.
- Haszpra, L., Barcza, Z., Davis, K. J., and Tarczay, K.: Long term tall tower carbon dioxide flux monitoring over an area of mixed vegetation, *Agricultural and Forest Meteorology*, 132, 58-77, 2005.
- 495 Haszpra, L., Barcza, Z., Bakwin, P. S., Berger, B. W., Davis, K. J., and Weidinger, T.: Measuring system for the long-term monitoring of biosphere/atmosphere exchange of carbon dioxide, *J. of Geophysical Research*, 106D, 3057-3070, <https://doi.org/10.1029/2000JD900600>, 2001.
- Heidbach, K., Schmid, H. P., and Mauder, M.: Experimental evaluation of flux footprint models, *Agricultural and Forest Meteorology*, 246, 142-153, <https://doi.org/10.1016/j.agrformet.2017.06.008>, 2017.
- 500 Hoffer, A., Jancsek-Turóczy, B., Tóth, Á., Kiss, G., Naghiu, A., Levei, E. A., Marmureanu, L., Machon, A., and Gelencsér, A.: Emission factors for PM₁₀ and polycyclic aromatic hydrocarbons (PAHs) from illegal burning of different types of municipal waste in households, *Atmos. Chem. Phys.*, 20, 16135-16144, <https://doi.org/10.5194/acp-20-16135-2020>, 2020.
- Holtslag, A. A. M. and Nieuwstadt, F. T. M.: Scaling the atmospheric boundary layer, *Boundary-Layer Meteorology*, 36, 201-209, <https://doi.org/10.1007/bf00117468>, 1986.
- 505 Hungarian Central Statistical Office: Gazetteer of Hungary, 1st January 2019, https://www.ksh.hu/docs/hun/hnk/hnk_2019.pdf, 2019.
- IPCC: 2006 IPCC Guidelines for National Greenhouse Gas Inventories - Prepared by the National Greenhouse Gas Inventories Programme (Eds.: Eggleston H.S., Buendia L., Miwa K., Ngara T. and Tanabe K.) <https://www.ipcc-nggip.iges.or.jp/public/2006gl/>, 2006.
- 510 IPCC: Climate Change 2014: Impacts, Adaptation, and Vulnerability. Part A: Global and Sectoral Aspects. Contribution of Working Group II to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change [C. B. Field, V. R. Barros, D. J. Dokken, K. J. Mach, M. D. Mastrandrea, T. E. Bilir, M. Chatterjee, K. L. Ebi, Y. O. Estrada, R. C. Genova, B. Girma, E. S. Kissel, A. N. Levy, S. MacCracken, P. R. Mastrandrea, L. L. White (eds.)], Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2014.
- 515 IPCC: 2019 Refinement to the 2006 IPCC Guidelines for National Greenhouse Gas Inventories (B. E. Calvo, K. Tanabe, A. Kranjc, J. Baasansuren, M. Fukuda, S. Ngarize, A. Osako, Y. Pyrozhenko, P. Shermanau, S. Federici [eds.]), IPCC, Switzerland, <https://www.ipcc-nggip.iges.or.jp/public/2019rf/>, 2019.
- Kljun, N., Rotach, M. W., and Schmid, H. P.: A Three-Dimensional Backward Lagrangian Footprint Model For A Wide Range Of Boundary-Layer Stratifications, *Boundary-Layer Meteorology*, 103, 205-226, 2002.
- 520 Kljun, N., Calanca, P., Rotach, M. W., and Schmid, H. P.: A Simple Parameterisation for Flux Footprint Predictions., *Boundary-Layer Meteorology*, 112, 503-523., 2004a.
- Kljun, N., Calanca, P., Rotach, M. W., and Schmid, H. P.: A simple two-dimensional parameterisation for Flux Footprint Prediction (FFP), *Geosci. Model Dev.*, 8, 3695-3713, <https://doi.org/10.5194/gmd-8-3695-2015>, 2015.



- 525 Kljun, N., Kastner-Klein, P., Fedorovich, E., and Rotach, M. W.: Evaluation of Lagrangian footprint model using data from wind tunnel convective boundary layer., *Agricultural and Forest Meteorology*, 127, 189-201., 2004b.
- Magyar Közút: Az országos közutak 2018. évre vonatkozó keresztmetszeti forgalma <https://internet.kozut.hu/kozerdeku-adatok/orszagos-kozuti-adatbank/forgalomszamlalas/>, 2019.
- 530 Moran, D., Kanemoto, K., Jiborn, M., Wood, R., Többen, J., and Seto, K. C.: Carbon footprints of 13 000 cities, *Environmental Research Letters*, 13, 064041, <https://doi.org/10.1088/1748-9326/aac72a>, 2018.
- Oettl, D.: Evaluation of the Revised Lagrangian Particle Model GRAL against wind-tunnel and field observations in the presence of obstacles, *Boundary-Layer Meteorology*, 155, 271-287, <https://doi.org/10.1007/s10546-014-9993-4>, 2015a.
- Oettl, D.: Quality assurance of the prognostic, microscale wind-field model GRAL 14.8 using wind-tunnel data provided by the German VDI guideline 3783-9, *Journal of Wind Engineering and Industrial Aerodynamics*, 142, 104-110, <https://doi.org/10.1016/j.jweia.2015.03.014>, 2015b.
- Oettl, D., Sturm, P. J., Bacher, M., Pretterhofer, G., and Almbauer, R. A.: A simple model for the dispersion of pollutants from a road tunnel portal., *Atmospheric Environment*, 36, 2943-2953., 2002.
- Papale, D.: Ideas and perspectives: enhancing the impact of the FLUXNET network of eddy covariance sites, *Biogeosciences*, 17, 5587-5598, <https://doi.org/10.5194/bg-17-5587-2020>, 2020.
- 540 Rana, G., Martinelli, N., Famulari, D., Pezzati, F., Muschitiello, C., and Ferrara, R. M.: Representativeness of carbon dioxide fluxes measured by eddy covariance over a Mediterranean urban district with equipment setup restrictions, *Atmosphere*, 12, 197, 2021.
- Romanov, A. A., Gusev, B. A., Leonenko, E. V., Tamarovskaya, A. N., Vasiliev, A. S., Zaytcev, N. E., and Philippov, I. K.: Graz Lagrangian Model (GRAL) for pollutants tracking and estimating sources partial contributions to atmospheric pollution in highly urbanized areas, *Atmosphere*, 11, 1375, 2020.
- 545 Satar, E., Berhanu, T. A., Brunner, D., Henne, S., and Leuenberger, M.: Continuous CO₂/CH₄/CO measurements (2012–2014) at Beromünster tall tower station in Switzerland, *Biogeosciences*, 13, 2623-2635, <https://doi.org/10.5194/bg-13-2623-2016>, 2016.
- 550 Schmid, H. P.: Source areas for scalars and scalar fluxes., *Boundary-Layer Meteorology*, 67, 293-318., 1994.
- Stagakis, S., Chrysoulakis, N., Spyridakis, N., Feigenwinter, C., and Vogt, R.: Eddy Covariance measurements and source partitioning of CO₂ emissions in an urban environment: Application for Heraklion, Greece, *Atmospheric Environment*, 201, 278-292, <https://doi.org/10.1016/j.atmosenv.2019.01.009>, 2019.
- 555 Tanács, E., Belényesi, M., Lehoczki, R., Pataki, R., Petrik, O., Standovár, T., Pásztor, L., Laborczi, A., Szatmári, G., Molnár, Z., Bede-Fazekas, Á., Kisé Fodor, L., Varga, I., Zsembery, Z., and Maucha, G.: Országos, nagyfelbontású ökoszisztéma-alaptérkép: módszertan, validáció és felhasználási lehetőségek (A national, high-resolution ecosystem basemap: methodology, validation, and possible uses), *Természetvédelmi Közlemények*, 25, 34-58, <https://doi.org/10.17779/tvk-jnatconserv.2019.25.34>, 2019.
- 560 Tian, H., Xu, R., Canadell, J. G., Thompson, R. L., Winiwarter, W., Suntharalingam, P., Davidson, E. A., Ciais, P., Jackson, R. B., Janssens-Maenhout, G., Prather, M. J., Regnier, P., Pan, N., Pan, S., Peters, G. P., Shi, H., Tubiello, F. N., Zaehle, S., Zhou, F., Arneth, A., Battaglia, G., Berthet, S., Bopp, L., Bouwman, A. F., Buitenhuis, E. T., Chang, J., Chipperfield, M. P., Dangal, S. R. S., Dlugokencky, E., Elkins, J. W., Eyre, B. D., Fu, B., Hall, B., Ito, A., Joos, F., Krummel, P. B., Landolfi, A., Laruelle, G. G., Lauerwald, R., Li, W., Lienert,



- S., Maavara, T., MacLeod, M., Millet, D. B., Olin, S., Patra, P. K., Prinn, R. G., Raymond, P. A., Ruiz, D. J., van der Werf, G. R., Vuichard, N., Wang, J., Weiss, R. F., Wells, K. C., Wilson, C., Yang, J., and Yao, Y.: A comprehensive quantification of global nitrous oxide sources and sinks, *Nature*, 586, 248-256, <https://doi.org/10.1038/s41586-020-2780-0>, 2020.
- van der Kwast, J., Timmermans, W., Gieske, A., Su, Z., Oliosio, A., Jia, L., Elbers, J., Karssenber, D., and de Jong, S.: Evaluation of the Surface Energy Balance System (SEBS) applied to ASTER imagery with flux-measurements at the SPARC 2004 site (Barrax, Spain), *Hydrol. Earth Syst. Sci.*, 13, 1337-1347, <https://doi.org/10.5194/hess-13-1337-2009>, 2009.
- Vesala, T., Kljun, N., Rannik, Ü., Rinne, J., Sogachev, A., Markkanen, T., Sabelfeld, K., Foken, T., and Leclerc, M. Y.: Flux and concentration footprint modelling: State of the art, *Environmental Pollution*, 152, 653-666., <https://doi.org/10.1016/j.envpol.2007.06.070>, 2008.
- 575 Vogt, R., Christen, A., Rotach, M. W., Roth, M., and Satyanarayana, A. N. V.: Temporal dynamics of CO₂ fluxes and profiles over a Central European city., *Theoretical and Applied Climatology*, 84, 117-126., <https://doi.org/10.1007/s00704-005-0149-9>, 2006.
- WMO: 20th WMO/IAEA Meeting on Carbon Dioxide, Other Greenhouse Gases and Related Measurement Techniques (GGMT-2019), Jeju Island, South Korea, 2-5 September 2019 (Eds.: A. Crowell, H. Lee and M. Steinbacher), GAW Report No., 255, 2020.
- 580 Zheng, B., Chevallier, F., Yin, Y., Ciais, P., Fortems-Cheiney, A., Deeter, M. N., Parker, R. J., Wang, Y., Worden, H. M., and Zhao, Y.: Global atmospheric carbon monoxide budget 2000–2017 inferred from multi-species atmospheric inversions, *Earth Syst. Sci. Data*, 11, 1411-1436, <https://doi.org/10.5194/essd-11-1411-2019>, 2019.