

Retrieval of the Sea Spray Aerosol Mode from Submicron Particle Size Distributions and Supermicron Scattering during LASIC

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Abstract. Improved quantification of sea spray aerosol concentration and size is important for determining aerosol effects on clouds and climate, though accurately capturing the size distribution of the sea spray mode remains limited by the availability of supermicron size distributions. In this [work](#), we introduce a new approach to retrieve lognormal mode fit parameters for a sea spray aerosol mode by combining submicron size distributions with supermicron scattering measurements using a Mie inversion. Submicron size distributions were measured by an Ultra-High Sensitivity Aerosol Spectrometer (UHSAS), and supermicron scattering was taken as the difference between $< 10 \mu\text{m}$ and $< 1 \mu\text{m}$ 3-wavelength integrating nephelometer measurements (NEPH). This UHSAS-NEPH method was applied during background marine periods of the Department of Energy Atmospheric Radiation Measurement Layered Atlantic Smoke Interactions with Clouds (LASIC) campaign on Ascension Island (November 2016 – May 2017) when the contribution of sea spray aerosol was expected to represent a large fraction of the aerosol mass and total scattering. Lognormal sea spray modal parameters were retrieved from comparisons between nephelometer measurements and a look-up table of Mie theory-simulated scattering coefficients for low error solutions that minimized the $0.4 - 1 \mu\text{m}$ residual in the UHSAS size distribution. [We evaluated the UHSAS-NEPH method with a set of clean marine measurements in the North Atlantic that included supermicron size and chemical measurements, showing that measured supermicron size distributions are needed to constrain the sea spray number concentration but mass concentration was reasonably characterized using supermicron scattering. For LASIC](#), the UHSAS-NEPH method retrieved sea spray mode properties for approximately 88% of background [marine times when the scattering variability and total particle concentrations were low \(\$\leq \pm 5 \text{ Mm}^{-1}\$ and \$< 400 \text{ cm}^{-3}\$, respectively\)](#), ranging from [0.6 to 1.9 \$\mu\text{m}\$ in mass mean diameter \(\$1.47 \pm 0.17 \mu\text{m}\$ \), 1.1 to 3.97 in modal width \(\$2.4 \pm 0.3\$ \), and mass concentration of 0.18 to 23.0 \$\mu\text{g m}^{-3}\$ \(\$8.37 \pm 4.1 \mu\text{g m}^{-3}\$ \)](#). The measured nephelometer scattering at 3 wavelengths was found to [constrain the mode width marginally](#) at the largest particle sizes in the absence of additional size and chemical measurements for defining parameters for the Mie solutions. Comparing UHSAS-NEPH retrievals to those of a fitting algorithm applied only to the submicron UHSAS number size distribution showed that correlations between retrieved mass concentration and the available mass-based sea spray tracers (coarse scattering, wind speed, and chloride) are low when supermicron measurements are not considered. This work demonstrates the added value of

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supermicron scattering measurements for retrieving reasonable sea spray mass concentrations, providing the best-available, observationally-constrained estimate of the sea spray mode properties when supermicron size distribution measurements are not available.

1 Introduction

Sea spray aerosol contributes the largest natural source of particles to the global aerosol mass budget (Lewis and Schwartz, 2004). Wind-driven breaking waves and bubble bursting at the ocean surface produce sea spray particles composed of organic components and sea salts that are injected into the atmosphere (Russell et al., 2010; O'dowd et al., 1997). Field measurements have shown that sea spray aerosol makes up 10 – 30% of the particles necessary for cloud formation, known as cloud condensation nuclei (CCN), at low supersaturations in marine regions (Modini et al., 2015; Quinn et al., 2017; Sanchez et al., 2021) and thus have important implications for modeled cloud properties and climate feedbacks (Horowitz et al., 2020). Model predictions of sea spray concentration are determined by a number of different emission parameterizations (e.g., Gong, 2003; De Leeuw et al., 2011; Salter et al., 2015), which leads to uncertainties in the sea spray mass production ($2.2 - 118 \text{ } 10^{12} \text{ kg yr}^{-1}$; De Leeuw et al., 2011), the shortwave scattering direct climate effect ($-2.2 - -0.15 \text{ W m}^{-2}$; Ayash et al., 2008), and the aerosol-cloud indirect climate effect ($-2.9 - +0.3 \text{ W m}^{-2}$; Paulot et al., 2020).

Ambient measurements of sea spray size distributions across submicron and supermicron sizes requires merging measurements from multiple instruments, often with a differential mobility analyzer (DMA) that measures submicron sizes (10 nm to 1 μm diameter) and an aerodynamic particle sizer (APS) for coarse sizes (0.5 to 10 μm diameter) (Modini et al., 2015; Saliba et al., 2019; Quinn et al., 2017). This DMA-APS-based technique has uncertainties controlled by the limited size range, resolution, and timing of each instrument, as well as by the ambient conditions. Merging mobility and aerodynamic measurements requires varying the particle density and shifting the size distribution until there is agreement between both instruments in the overlapping diameter range (Khlystov et al., 2004) due to uncertainties in particle densities of marine aerosol (Tang et al., 1997). The generally low number concentration of sea spray aerosol at supermicron sizes also causes poor counting statistics in the largest size bins of DMAs, which impacts the range of overlap to which the retrieval is sensitive (Russell et al., 1996a; Russell et al., 1996b).

Coarse mode sea spray has been retrieved using automated routines that fit a region of the merged size distribution, typically defined as particles larger than 0.4 μm diameter (0.4 – 10 μm), using a single lognormal mode (Saliba et al., 2019; Quinn et al., 2017). This diameter range often appears as a partial peak or “shoulder” in measured number and mass size distributions and is largely composed of sea salt particles during marine conditions (Quinn et al., 2017; Saliba et al., 2019; Zheng et al., 2018; Sanchez et al., 2021; Murphy et al., 1998). Lognormal fitting routines use parameters defining the shape of the mode

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(number concentration, geometric mean diameter, and geometric standard deviation) with observation-based constraints for the parameters (Modini et al., 2015; Quinn et al., 2017; Hussein et al., 2005) or as an unconstrained approach with parameters that vary freely (Saliba et al., 2019; Sanchez et al., 2021). Correlations of the single lognormal mode to wind speed and sea salt mass concentration provide justification for identifying the coarse mode as the “sea spray mode” during clean marine conditions (Saliba et al., 2019; Modini et al., 2015; Quinn et al., 2017; Lewis and Schwartz, 2004). When DMA and supermicron size distribution measurements were not available, submicron size-resolved measurements from the Ultra-High Sensitivity Aerosol Spectrometer (UHSAS) have been used for UHSAS-only retrieval of the sea spray mode (Sanchez et al., 2021; Zheng et al., 2018; Zheng et al., 2021). These methods have provided some demonstrated skill in predicting the number and cloud forming properties of sea spray particles less than 1 μm , but may not be sufficient to adequately characterize the mass concentration and optical properties of the supermicron fraction of sea spray (Murphy et al., 1998; Chamaillard et al., 2006; Kleefeld et al., 2002; O’ Dowd et al., 2010).

Given the uncertainties associated with merging of multiple instrument size distributions and the limited availability of supermicron size distribution measurements in marine regions, alternative methods should be considered to adequately characterize the modal properties of sea spray aerosol. Supermicron scattering measurements from nephelometers are commonly included as part of long-term atmospheric observations (Uin et al., 2019; Schmale et al., 2022) and provide an attractive alternative to supermicron size distribution measurements for constraining coarse sea spray properties (Demott et al., 2016; Testa et al., 2021). The premise of the approach proposed here is the observation that sea spray particle mass concentration often correlates to the supermicron scattering during clean marine conditions (Kleefeld et al., 2002; Chamaillard et al., 2006; Quinn et al., 1998; O’ Dowd et al., 2010). Scattering measurements can be translated to equivalent size distributions by employing Mie theory (Mie, 1908) using an Inverse Mie Method (IMM) with assumptions of particle size, composition, and concentration for optical properties of the particle population (Bluvshtein et al., 2017). Similar approaches that combine observed or simulated size distributions with scattering measurements have found that optical properties alone, without the addition of a particle sizer, are not sufficient to estimate properties of the aerosol size distribution (Frie and Bahreini, 2021; Shen et al., 2019; Lv et al., 2018).

In this work, we retrieved sea spray aerosol modal properties by fitting a single lognormal mode constrained by supermicron scattering at 3 wavelengths from a nephelometer to measured mass size distributions from an Ultra-High Sensitivity Aerosol Spectrometer, known here as the UHSAS-NEPH method. Since sea spray aerosol concentrations are most relevant to CCN in “clean” marine environments (Quinn et al., 2017) and the addition of non-marine sources (e.g. dust) tend to mask supermicron sea spray contributions, we directed the method at measurements that are largely reflective of clean marine conditions on Ascension Island during the Layered Atlantic Smoke Interactions with Clouds (LASIC) campaign. LASIC measurements provide an example where the sea spray size distribution needs to be retrieved, but no supermicron size distributions were measured. For comparison and support of UHSAS-NEPH with previously developed techniques, the method was also applied

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to an additional dataset that included nephelometer and submicron size measurements as well as size-resolved sea salt mass concentration and supermicron size distributions for the validation of reasonable sea spray mass retrieval.

2. Measurements

2.1 LASIC

Measurements from the Department of Energy Atmospheric Radiation Measurement (DOE ARM) site on Ascension Island, Saint Helena (7.96696° S, 14.34981° W) during the Layered Atlantic Smoke Interactions with Clouds (LASIC) campaign are used to develop the sea spray mode retrieval. LASIC captured the annual and seasonal cycles of aerosol and cloud properties during an 18-month (April 2016 – October 2017) deployment of the ARM Mobile Facility 1 (AMF1) (Miller et al., 2016; Zuidema et al., 2016). AMF1 measurements were collected at an isolated site on the windward flank of Green Mountain, away from the island’s airport and other inhabited areas (Zhang and Zuidema, 2019). The prevailing wind direction measured by meteorological instrumentation during the campaign was $115 \pm 10^\circ$ (east-southeasterly), indicating persistent sampling of offshore maritime air. The ARM Mobile Aerosol Observing System, a component of AMF1, housed the instruments, described in the following subsections, from an inlet situated 10 m above ground level at an altitude of 365 m above sea level (Uin et al., 2019).

Episodic intrusions of airborne biomass burning particles are carried into the Ascension Island marine boundary layer from South African wildfires annually during June – October (Zuidema et al., 2016). These events contrast sharply with the clean boundary layer that persists for the remainder of the year (November – May; Pennypacker et al., 2020). Non-marine aerosol particles are limited during this “background” season, though a few transport events of African dust that entrain into the boundary layer have been documented as occurring during austral summer and fall months (January – April) in the Southeast Atlantic (Kishcha et al., 2015; Virkkula et al., 2006). For background (non-biomass burning) times without dust events, the aerosol population is expected to be largely from marine sources, of which sea spray represent a large fraction of the aerosol mass concentration. This work will focus on LASIC background season observations (November 2016 – May 2017) (Table 1).

2.1.1 Submicron Particle Size Distributions

Two particle sizing instruments were operated during the LASIC campaign: a TSI Scanning Mobility Particle Sizer (SMPS; TSI Inc., Shoreview, MN, USA) and an Ultra-High Sensitivity Aerosol Spectrometer (UHSAS; Droplet Measurement Techniques (DMT) Inc., Longmont, CO, USA). The SMPS measured the aerosol size distribution ranging from 10 nm to 460 nm dry mobility diameter, which did not have sufficient overlap with the $> 0.4 \mu\text{m}$ diameter range relevant for retrieving the sea spray mode (Fig. 1), meaning that it did not provide constraints on the sea spray mode retrieval and was not used here.

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The UHSAS operated with 99 size channels at logarithmic spacing to cover diameters from 60 nm to 1 μm optical diameter at a temporal resolution of 1 Hz that were averaged to 1 min. The UHSAS was calibrated using polystyrene latex spheres with refractive index of approximately 1.59 and has a particle counting efficiency of approximately 100% for particle concentrations below 3000 cm^{-3} and sizes larger than 0.1 μm (Cai et al., 2008). The counts of particles per bin were converted to number size distributions using the sample flow rate (typically 50 $\text{cm}^3 \text{min}^{-1}$) and the sample accumulation time (10 s). [UHSAS sample line relative humidity \(RH\) was not recorded during LASIC, so the RH of UHSAS size distribution measurements was estimated using the UHSAS internal temperature along with ambient temperature and relative humidity from the Mobile Aerosol Observing System meteorological instrumentation. The UHSAS RH was found to be \$55 \pm 8\%\$ during LASIC.](#)

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UHSAS artifacts at large size bins have been reported for measurements in marine air masses (Pennypacker and Wood, 2017; Sanchez et al., 2021). These artifacts appear as two consistent and narrow modes at optical particle diameters of 0.6 and 0.85 μm , which likely represent the splitting of the sea spray mode by partial drying of salt that has been sampled from high ambient relative humidity (Fig. 2). These two modes constitute low contributions to the particle number concentration (Fig. 2a), but an appreciable amount to the mass concentration (Fig. 2b) of the measured size distributions. We expect that the 0.6 μm mode is the dried part of the salt mode, similar to distributions reported by [Sanchez et al. \(2021\)](#), while the narrow 0.85 μm mode is the remainder of the salt mode that is only partially dried. A treatment for these artifacts to fit the sea spray mode is described in Section 3.4.

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2.1.2 Supermicron Scattering

1 min averaged scattering coefficients (b_{scat}) were measured by a TSI 3563 3-wavelength integrating nephelometer at red (700 nm), green (550 nm), and blue (450 nm) light wavelengths over an angular integration range of 7° to 170° (Anderson et al., 1996) and impactor size cuts of 1 and 10 μm that alternated at intervals of about 1 h. [2-h averages of the scattering measurements for each impactor size cut were used to derive supermicron scattering coefficients at each wavelength \(\$b_{\text{scat},10\mu\text{m}}\(\lambda\)\$ \) by differencing the \$b_{\text{scat}}\$ at 1 \$\mu\text{m}\$ from \$b_{\text{scat}}\$ at 10 \$\mu\text{m}\$.](#) Scattering coefficients were corrected to account for known truncation errors due to significant coarse sea salt particle forward scattering at angles less than 7° (Anderson and Ogren, 1998). The detection limit of the nephelometer for typical operating conditions is between 0.1 and 0.3 Mm^{-1} depending on the wavelength (Anderson et al., 1996).

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[Particle scattering measurements during LASIC were not available at standard dry conditions \(\$< 40\%\$ \) as operating conditions only allowed for limited heating, typically producing \$60 \pm 4\%\$ RH at the nephelometer inlet for ambient conditions of \$88 \pm 8\%\$ RH. This average RH of 60% means particles were not dried to the efflorescence point for sea salt mixtures \(\$< 40\%\$ \) \(Ming and Russell, 2001\). The supermicron scattering \(at 450 nm\) did not show a significant correlation to the instrument RH \(\$R = 0.22\$, \$p = 0.19\$; Fig. S1\), but the correlation increases to \$R = 0.36\$ \(\$p < 0.05\$ \) for \$\text{RH} > 60\%\$, indicating that at higher RH the](#)

scattering may need to be corrected for humidity. Measurements of particle scattering at a series of pre-set RH were also collected during LASIC to provide hygroscopic growth factors ($f(RH)$) to correct scattering from 65% RH to the heated conditions (Zieger et al., 2010; Gasso et al., 2000). However, because the uncertainty for $f(65\% \text{ RH})$ was estimated to be $\geq 30\%$, which was approximately fourfold greater than the 8% for $< 1 \mu\text{m}$ and 7% for $< 10 \mu\text{m}$ scattering uncertainties for the heated measurements, we did not apply this correction. This uncertainty was driven by the limited and non-overlapping times for which $f(65\% \text{ RH})$ was available for $< 1 \mu\text{m}$ and $< 10 \mu\text{m}$ scattering, each typically spanning only $30 \pm 5\%$ of the 2-h averaging period. Without sufficient and simultaneous $f(RH)$ measurements of the humidity dependence of scattering, and given the additional uncertainties associated with correcting optical size distribution measurements with humidity and composition dependent refractive indices (Kassianov et al., 2015), correcting the measurements to a standard RH was also not possible. Instead, we restricted the measurements to include only those for which the average nephelometer humidity matched the average UHSAS RH. Restricting nephelometer measurements to those that had RH below 60% gives an average RH of 55% with 10% measurement uncertainty ($55 \pm 10\%$), while still retaining 78% of the measurements for this analysis.

2-h supermicron scattering during the LASIC background season had an average value of $12.0 \pm 6.3 \text{ Mm}^{-1}$ ($0.3 - 41.1 \text{ Mm}^{-1}$) as measured by the nephelometer at 550 nm and made up $70 \pm 7\%$ of the total scattering for particles less than $10 \mu\text{m}$ diameter. The combined use of the submicron particle size distribution from UHSAS and the supermicron scattering coefficients from the nephelometer provide the basis for naming this method UHSAS-NEPH.

2.1.3 Uncertainty and Variability of the Size Distributions and Scattering Measurements

Measurement variability and instrument error are incorporated into the sea spray mode retrieval to account for uncertainties in the Mie theory-based inversion of scattering and size distribution measurements (Table 2; Section 3).

UHSAS sizing uncertainty is within 2.5% of the particle size (Uin, 2016) with variations of -10% to +4% based on calibrated particles with known refractive index between 1.44 and 1.58 (Moore et al., 2021). The reported systematic uncertainty of number size concentration for accumulation mode ($0.1 - 1 \mu\text{m}$) particles measured by UHSAS has been shown to be 3.9% due to calibration, flow, and pressure biases (Kupc et al., 2018). This instrument error propagates to -27.5% to +12.4% for higher moments of the size distribution such as surface area and volume (Kupc et al., 2018; Brock et al., 2019). We therefore adopt a size uncertainty value (σ_D) of 2.5% as defined by the UHSAS instrument manufacturer (Uin, 2016) and 10% for the concentration uncertainty (σ_{PNSD}) that has been used in previous inversion procedures (Bluvshstein et al., 2017; Frie and Bahreini, 2021). Measured size distribution variability was calculated for the UHSAS size distribution at each diameter bin ($\sigma_{\text{PNSD, meas}}(D_p)$) as the standard deviation of the 2-h averages.

Systematic uncertainties of the particle scattering are mainly due to non-idealities at each measurement wavelength and angular sensitivities of the nephelometer (Anderson and Ogren, 1998). These features promote the use of a scattering uncertainty
255 ($\sigma_{\text{sca,inst}}(\lambda)$) value of 5% that has been used in previous inversion procedures (Frie and Bahreini, 2021; Bluvstein et al., 2017). Measured scattering variability was calculated for the supermicron scattering at each wavelength ($\sigma_{\text{sca},1-10\mu\text{m}}(\lambda)$) as the standard deviation during the 2-h average. A list of the measured size distribution and scattering variables and their associated uncertainties and variabilities are provided in Table 2. Particle losses due to aspiration and transmission in the [ARM Mobile Aerosol Observing System](#), were assessed using the particle loss calculator (PLC; Von Der Weiden et al., 2009), sample line
260 configurations and geometry from Bullard et al. (2017), and a particle density of 1 g cm⁻³. Losses were found to be less than 10% for particles smaller than 1 μm diameter and greater than 50% for particles larger than approximately 6 μm . [We did not correct UHSAS size distributions for these losses given the uncertainty of that correction](#), but note that nephelometer supermicron scattering measurements may underestimate retrieved sea spray number and mass concentrations at the largest diameters. Using the mean sea spray mode statistics of UHSAS-NEPH, [the particle losses in the instruments](#) can equate to an
265 underestimation of roughly $13 \pm 8\%$ in the sea spray mass and $0.8 \pm 4\%$ in number using nephelometer scattering.

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2.1.4 Ancillary Variables

As an alternative to size-resolved filter measurements of sodium, which were not collected during LASIC, measurements from an aerosol chemical speciation monitor (ACSM; Aerodyne Research, Billerica, MA USA) were used to evaluate sea salt mass retrievals. ACSM provided mass and chemical composition (organics, sulfate, nitrate, ammonium, and chloride) of non-
270 refractory submicron aerosols. Since sea salt does not volatilize efficiently at 600°C, the ACSM measurement of non-refractory chloride provides a trace signal from NaCl [that is detectable](#) in the absence of large sources of non-refractory chloride (Frossard et al., 2014; Ovadnevaite et al., 2012) and has been used as a tracer to identify sea salt aerosol contributions to CCN (Humphries et al., 2021). We use chloride measurements only for January 2017 [through](#) May 2017, as these data were quality assured. 2-h averaged ACSM chloride concentration showed statistically significant ($p < 0.05$) positive correlation with two common sea
275 spray tracers: wind speed ($R = 0.2$) and nephelometer supermicron scattering at 550 nm ($R = 0.33$) (Fig. S2). This observation supports the potential of LASIC ACSM chloride measurements to serve as a chemical tracer for sea spray mass in the evaluation of [UHSAS-NEPH](#).

We additionally incorporated measurements of 1 min averaged condensation nuclei concentrations above 3 nm (CN_3) from a
280 TSI Ultrafine Condensation Particle Counter (CPC) 3776, wind speed and rain intensity from a Vaisala WXT-520, refractory black carbon (rBC) concentration from a DMT Single Particle Soot Photometer, 470 nm particle absorption from a Radiance Research Particle Soot Absorption Photometer (PSAP), and carbon monoxide (CO) mixing ratio from a Los Gatos Research trace gas analyzer (Miller et al., 2016). These measurements were used to identify clean marine periods (Section 2.2), [assess](#)

environmental influence on retrieval performance, and evaluate the retrieved sea spray masses. Ancillary measurements were averaged to 2-h resolution to match the timing of the supermicron scattering coefficients and size distribution averages.

2.2 Measurement Screening

Non-marine aerosol particles, specifically those from combustion sources, have been shown to influence the performance of sea spray mode retrievals by contributing number and mass concentrations that overlap with the region used to fit a mode in ambient size distributions (Modini et al., 2015; Saliba et al., 2019). Here we focus on periods when boundary layer air masses are assumed to have a multi-day marine history in order to reduce non-marine sources and ensure retrieval results that are consistent with sea spray. In the case of LASIC observations, boundary layer intrusions of biomass burning aerosol can affect particle optical properties by increasing absorption and reducing scattering for sub- and supermicron particles (Delene and Ogren, 2002; Denjean et al., 2020). In particular, UHSAS instruments show under-sizing of particles when highly absorbing biomass burning aerosol are introduced into the particle population due to heating from the instrument beam and subsequent particle vaporization and shrinkage (Howell et al., 2021). Although this effect is more likely to impact particles smaller than the assumed sea spray size ($D_p \leq 0.4 \mu\text{m}$), any impact on potential submicron sea spray aerosol contributions would have an influence on the sea spray mode retrievals. To limit non-marine influences on UHSAS-NEPH retrieval, we first isolated measurements during “clean marine” periods of the LASIC background season from November 2016 through May 2017 by applying the following criteria (Fig. 3):

- (1) CN_3 concentration less than 600 cm^{-3} , which was the approximate 90th percentile particle concentration during the LASIC campaign,
- (2) CO mixing ratio (a proxy for continentally-sourced air) below the limit of ambient marine boundary layer background levels observed during LASIC (70 ppbv; Pennypacker et al., 2020),
- (3) rBC concentration below the combustion source threshold of 50 ng m^{-3} used by Saliba et al. (2020) in the remote marine North Atlantic, and
- (4) $< 10 \mu\text{m}$ Scattering Angstrom Exponent (SAE_{10}) values less than 1 for 450 and 700 nm nephelometer scattering. SAE_{10} characterizes the wavelength dependence of particles and takes on small (< 1) values during periods in which coarse aerosol, such as sea salt, have a large mass contribution (Shen et al., 2019; Mulcahy et al., 2009).

Saharan dust and continental aerosol transport from southern Central Africa into the remote tropical Atlantic boundary layer has been a commonly observed contributor to the surface-level aerosol population at Ascension Island (Swap et al., 1996). The mass concentration of transported dust particles are largely in the supermicron size range (Miller et al., 2021; Denjean et al., 2016) and overlap the fitting region used in UHSAS-NEPH. To exclude interference in the retrieval from dust particles, we used measurements of the sub- $10 \mu\text{m}$ single-scattering albedo at 470 nm ($\text{SSA}_{470\text{nm}}$) from NEPH scattering and PSAP absorption to identify times with possible dust influence (SI Text 1). An $\text{SSA}_{470\text{nm}}$ threshold of 0.95 was used to distinguish

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335 [between sea salt and dust aerosol contributions to coarse scattering based on the relationship of SAE₁₀ and average SSA reported for Saharan dust \(Di Biagio et al., 2019; Von Hoyningen-Huene et al., 2009; Haywood et al., 2003\). This restriction removed 68 2-h periods.](#)

We additionally removed periods when the rain intensity exceeded 1 mm h⁻¹ at any hour during the 2-h average to ensure
 340 minimal influence of precipitation on the retrieval, namely wet scavenging of sea spray aerosol by rain droplets. 14 periods exceeded this rain intensity restriction. The combination of all criteria identified 909 2-h (non-raining) clean marine periods, which accounted for approximately 40% of all available background season observations. The clean marine criteria reduced the average criteria values by 12% for CN₃, 7% for CO, 70% for rBC, and 15% for SAE₁₀ after applying these restrictions. The 909 available periods provided persistent marine conditions that were low in aerosol concentration ($300 \pm 90 \text{ cm}^{-3}$) and
 345 in combustion influence ($15.7 \pm 12.7 \text{ ng m}^{-3}$), as well as having a large scattering contribution from coarse particles (SAE₁₀ = 0.66 ± 0.15).

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3 Sea Spray Mode Retrieval (UHSAS-NEPH)

This section outlines the procedure used to retrieve sea spray modes from scattering measurements and submicron mass size distributions (Fig. 4). We describe the relationship of particle scattering to particle size (Section 3.1) and how this theoretical
 350 relationship is used to identify a group of probable sea spray mode solutions that are consistent with the measured supermicron scattering, as well as with literature reported ranges of modal properties (Section 3.2). Mode solutions are then constrained with measured submicron mass size distributions ($D_p > 0.4 \mu\text{m}$) to isolate and retrieve the most consistent sea spray modal properties (Sections 3.3 – 3.4).

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3.1 Simulating Sea Spray Mode Scattering using Mie Theory

355 A look-up table of scattering coefficients was developed by employing a modified Mie theory code based on the algorithm for red (700 nm), green (550 nm), and blue (450 nm) wavelengths (Bohren and Huffman, 1998; Mätzler, 2002). These wavelengths were chosen to match those used by the 3-wavelength integrating nephelometer operated during LASIC (Section 2.1.2). Each of the three red, green, blue (RGB) scattering coefficients is attributed to a combination of lognormal mode fitting parameters (N_i , D_g , and σ_g) that describe the shape of the sea spray mode. We use the canonical lognormal mode form to represent the
 360 number size distribution of sea spray aerosol with the following equation (Seinfeld and Pandis, 2006):

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$$\frac{dN}{d\log D_p} = \frac{N_t}{\sqrt{2\pi \log_{10} \sigma_g}} e^{-\frac{(\log_{10} D_p - \log_{10} D_g)^2}{2(\log_{10} \sigma_g)^2}} \quad (1),$$

370 where N_r is the number concentration of particles (cm^{-3}), D_p is the particle diameter (μm), D_g is the geometric mean diameter (μm), and σ_g is the geometric standard deviation (or mode width; unitless). These values are $1 - 99 \text{ cm}^{-3}$ for number concentration, $0.05 - 1.19 \mu\text{m}$ for mean diameter, and $1 - 4$ for mode width, which provides over 157,000 possible sea spray mode solutions (Table 3). The range of fitting parameters were chosen to reflect those reported in laboratory experiments and field measurements (Table 6; Prather et al., 2013; Modini et al., 2010; Modini et al., 2015; Quinn et al., 2017; Saliba et al., 375 2019). [Similar look-up tables have previously been used for optical measurement inversions](#) (e.g., Lv et al., 2018; Veselovskii et al., 2002).

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The scattering coefficients ($b_{sca,MIE}(\lambda)$) are then related to the size distribution by integrating Eq. (2) over all particle diameters ($0.01 - 10 \mu\text{m}$),

$$380 \quad b_{sca,MIE}(\lambda) = \int_0^\infty \frac{\pi D_p^2}{4} Q_{sca}(\lambda, m, D_p) \frac{dN}{d \log D_p} d \log D_p \quad (2).$$

Q_{sca} is the scattering efficiency, λ is the light wavelength, and m is the particle core refractive index ($m = n + ik$). [To match the NEPH and UHSAS RH of 55% during LASIC \(Section 2.1.2\), a constant \$m\$ value of \$1.45 + i0\$ was selected to simulate sea spray particle scattering. This value is lower than the average refractive index reported for dry sea salt \(real component = \$1.5 - 1.6\$, imaginary component \$< 10^{-6}\$ \) \(Wang and Rood, 2008; Randles et al., 2004; Bi et al., 2018\) and was calculated as a mass-weighted mixture of salt with water, where water has a refractive index of \$1.33\$ \(Wang and Rood, 2008\) \(SI Text 2\). We found no substantial variation in the retrieved fit parameters for the range of \$1.4 + i0\$ to \$1.6 + i0\$, but correlations to sea spray tracers \(chloride and wind speed\) were lower for \$1.40 + i0\$ \(high RH\) and \$1.6 + i0\$ \(low RH\) \(SI Text 2\). The mid-range value of \$m = 1.45 + i0\$ was used to approximate the scattering of sea salt for the measured RH of 55%.](#)

Deleted: Within the Mie simulation, the refractive index, m , is assumed to be constant at $1.56 + i0$, which is typical for sea salt particles (Bi et al., 2018; Kent et al., 1983). This value provides a reasonable estimate of the scattering at marine sites where sea salts contribute a majority of supermicron scattering mass, with organic components contributing minor if not offsetting differences to the refractive index.

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3.2 Selecting the Most Probable Mie Solutions

Since many solutions are within the constraints of three wavelengths of measurements, we use the error between the nephelometer scattering and the Mie theory solutions to remove mode solutions that are not within the calculated error. The probability of solutions that meet the error threshold are then evaluated and only the top 5% most probable [solutions](#) are selected.

The measured supermicron scattering coefficient ($b_{sca,1-10\mu\text{m}}(\lambda)$) is compared to scattering coefficients computed for each simulated sea spray size distribution and Mie theory [value of \$b_{sca,MIE}\(\lambda\)\$](#) in the look-up table by calculating the absolute error at each wavelength (λ), using Eq. (3):

$$\Delta b_{sca}(\lambda) = |b_{sca,1-10\mu m}(\lambda) - b_{sca,MIE}(\lambda)| \quad (3).$$

The scattering error ($\Delta b_{sca,RGB}$) is then computed by propagating the absolute error at each wavelength following Eq. (4):

$$\Delta b_{sca,RGB} = \sqrt{\sum_{\lambda=R,G,B} [\Delta b_{sca}(\lambda)]^2} \quad (4).$$

The total sample space is then reduced by selecting solutions from the look-up table that fall below the error threshold ($\sigma_{sca,RGB}$) calculated for that measurement time,

$$\Delta b_{sca,RGB} < \Delta \sigma_{sca,RGB} \quad (5a),$$

where,

$$\Delta \sigma_{sca,RGB} = \sqrt{\sum_{\lambda=R,G,B} \left[\left(\sigma_{sca,1-10\mu m}(\lambda) \right)^2 + \left(\sigma_{sca,inst} \right)^2 \right]} \quad (5b).$$

This error threshold incorporates the measured scattering variability at each wavelength for the averaging period ($\sigma_{sca,1-10\mu m}(\lambda)$) and accounts for the instrument error ($\sigma_{sca,inst}$, which is constant; Table 2). Figure 5 illustrates a time series of these error thresholds and the percent reduction of the Mie solution sample space ($N = 157,850$) for each retrieval during the background season of the LASIC. Using these error thresholds, the solution space is reduced by 98% on average with a range of 83 – 99% reduction, resulting in approximately 1000 – 2000 possible solutions each time.

The majority of solutions that are below the error threshold typically have a sea spray mode shape similar to those previously reported in literature (Quinn et al., 2017; Saliba et al., 2019; Bates et al., 2012), namely mass mean diameters within or near the coarse mode size range (1 – 10 μm) and mode width of 2 – 3, decreasing to below 0.1 cm^{-3} number concentration before the 10 μm cutoff. However, some of the solutions that meet the error threshold criterion are either too wide or large relative to the reported range of sea spray modes (Table 6, Fig. S4), which is a limitation of having only three scattering wavelengths to constrain the mode. To remove the outlier solutions and to reduce the sample space to a more consistent group of solutions, we apply a restriction on Mie solutions to consider only the most probable fitting parameters (N_T , D_g , σ_g) based on their frequency of occurrence. To demonstrate this restriction, we considered the normalized probabilities of the fitting parameters from Mie solutions that fall below the scattering error threshold, $\Delta \sigma_{sca,RGB}$, for one retrieval during LASIC (Fig. 6a-c). This retrieval was selected as its $\Delta \sigma_{sca,RGB}$ was within the average value for all LASIC observations during the clean marine background season ($3.1 \pm 2.1 Mm^{-1}$) and is representative of most cases assessed.

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450 Mie solutions that meet the error threshold constrained the mode number concentration (N_T) and mean diameter (D_g) of the sea spray mode as shown by likely (probability $> 75\%$) N_T and D_g falling within narrow ranges of the low error solutions (4% and 10% of the sample space, respectively) compared to a less constrained range for σ_g (45% of the sample space) (Fig. 6a-c). Since σ_g has a wider range of probable values, this parameter is effectively constrained by considering the joint probabilities of $N_T|\sigma_g$ and $D_g|\sigma_g$. These probabilities are computed as

$$P(N_T, \sigma_g) = P(N_T|\sigma_g) \cdot P(\sigma_g) \quad (6)$$

$$P(D_g, \sigma_g) = P(D_g|\sigma_g) \cdot P(\sigma_g) \quad (7)$$

460 The highest joint probability values restrict solutions to localized regions of the solution space (Fig. 6d-e). Selecting only the upper 5th normalized probability percentile of joint probabilities further restricts the full Mie solution sample space by an additional 16% and 13% for $N_T|\sigma_g$ and $D_g|\sigma_g$, respectively (Fig. 6d,e).

465 To assess which of these joint probability restrictions provided the most realistic sea spray mode results, retrieved mass concentrations and fitting parameters during the clean marine background season of LASIC were compared (Fig. S5). Restricting low error solutions using $D_g|\sigma_g$ joint probability led to retrieved sea spray mass concentrations that were 20 – 30% higher than the $N_T|\sigma_g$ combination, had mean mass diameters smaller than $N_T|\sigma_g$ ($0.68 \pm 0.08 \mu\text{m}$ vs. $1.47 \pm 0.17 \mu\text{m}$), and much broader mode widths (3.8 ± 0.2 vs. 2.4 ± 0.3). Although the retrieved mass mean diameters using the $D_g|\sigma_g$ joint probability were within the range of reported values from literature (Table 6; Quinn et al., 2017; Saliba et al., 2019; Bates et al., 2012), the exceptionally broad widths outside of the reported range suggests that the restriction using $D_g|\sigma_g$ does not effectively constrain the mode width as well as the $N_T|\sigma_g$ combination. For these reasons, only sea spray modes with fitting parameters in the top 5th percentile of $N_T|\sigma_g$ joint probability are used for this method, resulting in typically 300 – 500 for each measurement time.

475 3.3 Perturbing the Size Distribution

Uncertainty and variability have been shown to impact the assumed size distributions and optical properties of aerosol using inversion techniques (Viskari et al., 2012; Frie and Bahreini, 2021). Here, instrument uncertainty and measurement variability are incorporated into the fitting method by introducing random noise into the size distribution based on the size (σ_D) and number concentration ($\sigma_{\text{PNSD, meas}}(D_p)$, $\sigma_{\text{PNSD, inst}}$) uncertainties (Section 2.1.3; Table 2). Perturbations are simultaneously made to the measured number size distribution and particle diameters. Each bin of the size distribution is perturbed by introducing Gaussian noise that samples a random number from a normal distribution. The sampled normal distribution is defined by mean

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(μ) and standard deviation (σ) parameters where μ is the measured PNSD averaged over the time interval ($\sigma_{PNSD, meas}(D_p)$) and σ is the combined errors of the temporal variability and instrument concentration uncertainty (which is constant),

$$\sigma = \sqrt{(\sigma_{PNSD}(D_p))^2 + (\sigma_{PNSD\ inst.})^2} \quad (6).$$

The diameters in the size distributions are perturbed by shifting the size bins by the same value with a μ of 0 and σ as the instrument size uncertainty (σ_D). Size distributions are perturbed 100 times to provide a sample space of $N_{perturb} + 1$ (the measured size distribution and 100 perturbations). Each probable Mie solution retrieved in Section 3.2 is then tested for the $N_{perturb} + 1$ cases within the fitting region of the measured submicron size distribution described in Section 3.4.

3.4 Fitting Modes to the Measured Size Distribution

The last step in retrieving sea spray modal properties is to select Mie solutions that most closely match the shape of the measured size distribution in a specified fitting region. [Given the correlation of supermicron scattering and sea spray mass concentration](#) during clean marine conditions (Chamaillard et al., 2006; Kleefeld et al., 2002; Quinn et al., 1998), all PNSDs are converted to particle mass size distributions (PMSDs) using

$$\frac{dM}{d \log D_p} = \frac{\pi}{6} \rho D_p^3 \frac{dN}{d \log D_p} \quad (7),$$

with the assumptions of spherical particle homogeneity [and constant sea spray density, \$\rho\$. UHSAS and NEPH measurements were collected at 55% RH, so \$1.3 \text{ g cm}^{-3}\$ was used as the estimated sea spray particle density by calculating a mass-weighted mixture of salt with water \(SI Text 3\).](#)

To account for the consistent UHSAS artifacts at 0.6 and 0.85 μm (Section 2.1.1), we [restricted the measured size distributions](#) used to fit the Mie theory-simulated size distributions [for diameters larger than 0.4 \$\mu\text{m}\$ \(0.4 to 1 \$\mu\text{m}\$ UHSAS range\)](#) to 0.38 – 0.83 μm (the closest UHSAS diameter size bins within the specified range), which weights the comparison toward smaller sizes and effectively reduces the influence of the largest artifact while maintaining the shape of the accumulation mode “shoulder” (Fig. 1). For each Mie solution ($PMSD_{MIE}$), a residual sum of squares is computed between the measured and perturbed PMSDs as,

$$Fit\ RSS(j, k) = \sum [PMSD(D, j) - PMSD_{MIE}(D, k)]^2 \quad (8),$$

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525 where index j represents each measured or perturbed size distribution and k is the index of the low error Mie solution. The fit
RSS provides a quantifiable metric for comparing the scattering-retrieved mode to the expected sea spray fitting region. Chi-
square “goodness of fit” calculations were also calculated and provided similar retrieved sea spray modal properties to those
retrieved using the fit RSS minimization. The [high probability](#) Mie solution with minimum fit RSS for the measured and noise-
perturbed size distributions is chosen to establish a range of mode fits. This Monte Carlo approach allowed for the selection
530 of Mie solutions that capture the shape of the accumulation mode “shoulder” given uncertainties associated with measurement
variability and instrument error and provides a statistically robust sample space for retrieving a unique sea spray mode solution.
From the range of fit RSS-minimized sea spray modes, 95% confidence intervals of each fitting parameter are calculated to
further constrain the most probable solution that fits the measured size distribution in the expected sea spray size range. On
average, 30 – 40 solutions remain from this perturbation analysis for each measurement time with variabilities of 4% in number
535 concentration, 3% in geometric mean diameter, and 1% in geometric standard deviation based on the sample means and upper
and lower limits of the 95% confidence intervals. The low variabilities of these fitting parameters demonstrate consistent mode
retrievals within the perturbations and stability of the retrieval procedure. Lastly, the fitting parameter solutions that are both
highly probable (Section 3.2) and within the 95% confidence interval of fit RSS-minimized modes are averaged to produce a
single lognormal sea spray [mode](#) with uncertainty of the solution defined as the upper and lower bounds of the 95% confidence
540 interval.

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4 Evaluation of NEPH-constrained Sea Spray Retrieval using Supermicron Size Measurements

[This UHSAS-NEPH method was developed for LASIC measurements because that study lacked supermicron size distribution
measurements. To evaluate the performance of this nephelometer-constrained sea spray mode retrieval, we first compared the
545 scattering-constrained retrieval method during clean marine periods of the first cruise of the North Atlantic Aerosols and
Marine Ecosystems Study \(NAAMES 1\) deployed 6 November – 30 November 2015 \(Behrenfeld et al., 2019; Saliba et al.,
2019\), for which sea spray mode retrievals using supermicron size distributions and salt mass concentrations were also
available. NAAMES 1 was selected as a case study of the nephelometer constraint on supermicron size because it had the most
persistent clean conditions of the four ship deployments \(Saliba et al., 2020\). The low particle concentrations during NAAMES
550 1 \(median \$CN_{total} = 94 \text{ cm}^{-3}\$ \) also included the lowest sea spray aerosol concentrations \(mean \$N_{sea \text{ spray}} = 5 \text{ cm}^{-3}\$ \) of the four
cruises \(Saliba et al., 2019\).](#)

[During NAAMES 1, submicron \(0.1 – 0.8 \$\mu\text{m}\$ \) and supermicron \(0.5 – 10 \$\mu\text{m}\$ \) size distributions were measured by a Scanning
Electrical Mobility Spectrometer \(SEMS, model 2002 BMI\) and an Aerodynamic Particle Sizer \(APS; model 3321 TSI\),
555 respectively. Scattering coefficients at \$< 1 \text{ }\mu\text{m}\$ and \$< 10 \text{ }\mu\text{m}\$ diameter size cuts were measured by a 3- \$\lambda\$ \(450, 550, 700 nm\)
integrating nephelometer \(model 3563, TSI\). In contrast to the LASIC measurements, nephelometer and size distribution
measurements were operated at dry conditions \(\$RH < 40\%\$ \), meaning that dry particle properties could be assumed. Offline
analysis of filters using ion chromatography provided size-resolved \$< 10 \text{ }\mu\text{m}\$ salt \(\$\text{Na}^+\$ \) mass concentrations at a 24-h resolution.](#)

Saliba et al. (2019) retrieved the sea spray aerosol mode by fitting a lognormal mode to the shoulder of the merged ambient size distributions (SEMS-APS) during NAAMES 1. This method expanded upon the techniques described in previous work for fitting lognormal modes to measured marine size distributions (Modini et al., 2015; Quinn et al., 2017; Hussein et al., 2005) by allowing for the diameter and width of the mode to vary without laboratory constraints, which, when constrained, frequently retrieved parameters close to the interval limits. 15-min averaged $< 10 \mu\text{m}$ SEMS-APS sea spray mode mass concentrations during NAAMES 1 were previously compared to the filter measurements of $< 10 \mu\text{m}$ Na^+ mass and wind speed and found correlations of 0.7 and 0.6, respectively (Saliba et al., 2019). These correlations support the interpretation of the SEMS-APS retrieved mode as sea spray aerosol. For comparison of the UHSAS-NEPH retrieval method during 2-h NAAMES 1 measurements, SEMS was used as a replacement for UHSAS (SEMS-NEPH) with refractive index ($m = 1.56 + 0i$) and sea spray particle density (2.0 g cm^{-3}) for dry conditions.

Figure 7 and Table 4 show comparisons of the SEMS-APS and SEMS-NEPH sea spray retrievals for 2-h averaged SEMS, APS, and nephelometer measurements. SEMS-NEPH estimates the sea spray mode with double the number concentration, but only 10% more mass concentration on average than SEMS-APS (Fig. 7a,d). The main differences between these two variables can be explained by the retrieval constraints of each method: SEMS-APS constrains the number size distribution while the Mie-based scattering constraint of SEMS-NEPH is dependent upon the mass concentration. The difference in retrieved number concentrations between the two methods may have implications for attributing particle contributions to CCN. A factor of 2 difference between SEMS-NEPH and SEMS-APS estimates of particle number could modify the commonly observed 10 – 30% contribution of sea spray to CCN at supersaturations of 0.1 – 0.4% (Quinn et al., 2017; Sanchez et al., 2021; Modini et al., 2015) as more particles contribute to an already low CCN concentration during clean marine conditions.

The majority of integrated mass comparisons fall within a reasonable range of the 1:1 line except for some periods of low SEMS-APS mass concentrations ($< 3 \mu\text{g m}^{-3}$) (Figure 7d). The larger mass concentration estimates of SEMS-NEPH could be attributable to observed supermicron sea spray mass from the nephelometer that is not fully resolved by the APS number concentration estimate. This is supported by a stronger correlation of retrieved mass with supermicron scattering for SEMS-NEPH ($R = 0.72$) than what is observed using SEM-APS ($R = 0.56$) (Fig 7f), although it should be noted that the supermicron scattering has been used to constrain SEMS-NEPH solutions. Since the width of the sea spray mode from the scattering-based retrieval has been shown to be a poorly constrained parameter and is often narrower for SEMS-NEPH (2.0 ± 0.3) than SEMS-APS (2.4 ± 0.3), it may be the source of discrepancies between the other modal properties (Fig. 7c). These discrepancies are particularly apparent for the mode diameters, which are 20% larger in mean number size and 30% smaller in mean mass size for SEMS-NEPH in comparison to SEMS-APS. The scattering-based approach uses the probability occurrence of modal width to retrieve an optimal value, but the lack of number size distribution measurements at larger sizes ($> 0.8 \mu\text{m}$) from the SEMS provides insufficient information to constrain this parameter.

595 For the 2-h averaged concentrations, sea spray mass concentrations show moderate correlations of 0.50 for SEMS-APS and
0.53 for SEMS-NEPH with wind speed (Fig. 7g). The SEMS-APS correlation of mass to wind speed is slightly weaker than
the correlation previously reported by Saliba et al. (2019) for 15-min averages ($R = 0.6$), but maintains a moderate correlation
suggesting that the estimated mass concentration from both retrievals can be attributed to sea spray production by wind.
600 Ambient $< 10 \mu\text{m}$ sodium (Na^+) mass analyzed with ion chromatography moderately correlates with the SEMS-NEPH retrieval
of sea spray mass (Fig. 7h). Although the SEMS-NEPH correlation is weaker than the mass concentration correlation to sodium
using SEMS-APS ($R = 0.7$), the SEMS-NEPH correlation is consistent with other studies merging measured sub- and super-
micron size distributions to fit sea spray modes and estimate salt mass in clean marine conditions (Quinn et al., 2017; Modini
et al., 2015). Together these results show the efficacy of nephelometer-constrained estimates of supermicron sea spray
605 particles. The moderate correlations and relative agreement between parameters of the retrieval methodologies with sea spray
tracers indicate reasonable retrievals at dry relative humidity, even though supermicron size distributions provide more
accurate retrieval of sea spray number properties.

5 Performance of UHSAS-NEPH Retrievals

610 For LASIC, sea spray modal properties were retrieved for 906 of the 909 2-h clean marine background periods using UHSAS-
NEPH. The 3 periods for which the sea spray mode could not be retrieved were due to missing supermicron scattering
measurements for at least one wavelength. To ensure that the algorithm retrievals were sufficiently consistent with both
UHSAS and NEPH and representative of marine aerosol, we restricted the results to those with (1) low residual errors between
the retrieval and measurements in the fitting region ($0.38 - 0.83 \mu\text{m}$), (2) low measured scattering variability during each 2-h
time period, and (3) limited influence from potential non-marine aerosol sources, namely CN_3 concentrations $< 400 \text{ cm}^{-3}$.

615 The fit RSS (Section 3.4) is the difference between the measured size distribution fitting region and the Mie solution
determined from 3-wavelength supermicron scattering. Low residuals (fit RSS < 2) indicate good agreement between the
measured region of the accumulation mode “shoulder” ($0.38 - 0.83 \mu\text{m}$) and the modal fit retrieved from scattering, hence low
RSS shows the retrieved mode is well constrained by the measured UHSAS size distribution and NEPH scattering. The average
620 fit RSS for the LASIC dataset that met the marine criteria was 1.26 ± 0.89 (Fig. 8a), which indicates that the retrieved modes
are generally within the uncertainty of the measured size distribution in the $0.38 - 0.83 \mu\text{m}$ fitting region, but solutions that
are not constrained by the UHSAS size distribution are also retrieved as evident by large values of fit RSS for various retrieval
periods (Fig. 8a). A visual inspection of the retrieved mode fits suggested that values exceeding a residual threshold of 5 should
be rejected as there is not sufficient agreement in the overlap region to consider the solution acceptable. Retrieved modes with
625 fits above this threshold tended to have solutions that were either larger than the measured accumulation mode “shoulder” or
had peaks in regions that the size distribution was low. These high fit RSS likely indicate that the supermicron scattering
measurements were influenced by particles other than sea spray, which were not effectively constrained by the supermicron

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scattering. We also examined RSS thresholds of 1 through 10. There were no significant changes in correlations to wind speed or chemical signatures (Section 5.2) for lowering the threshold to values of 1 – 4, and there were decreases in those correlations for thresholds above a value of 6 (SI Table 7).

To assess the limitations of the algorithm due to the observed variability in measured scattering, we examined the relationship between the scattering error threshold and the fit RSS. The scattering error threshold, $\Delta\sigma_{\text{sca,RGB}}$, is defined as the combined effect of temporal variability and instrument uncertainty and had an average value of $3.1 \pm 2.1 \text{ Mm}^{-1}$ during LASIC. Low fit RSS (< 5) appear to coincide with low values of $\Delta\sigma_{\text{sca,RGB}}$ ($R = 0.39$), which suggests well-constrained sea spray mode solutions when the nephelometer scattering variability was low during the 2-h average (Fig. 8b). The relationship of low $\Delta\sigma_{\text{sca,RGB}}$ with low fit RSS generally persists to a threshold value of about $\Delta\sigma_{\text{sca,RGB}} = 5 \text{ Mm}^{-1}$ before the fit RSS values increase in magnitude. Using this observed relationship between $\Delta\sigma_{\text{sca,RGB}}$ and fit RSS as a measure of the algorithm being sufficiently constrained to provide reasonable fits, we applied a restriction of acceptable scattering uncertainty tolerance of 5 Mm^{-1} .

Figure 8c illustrates a moderate correlation between fit RSS and CN_3 concentration ($R = 0.58$). Fit RSS is generally below the threshold of 5 for CN_3 concentrations less than 400 cm^{-3} . Above this concentration the fit RSS increases to higher values more consistently. The 600 cm^{-3} CN_3 criterion used to screen for clean periods is above the average and variability of the clean marine background season after the restriction was applied ($300 \pm 90 \text{ cm}^{-3}$). Because supermicron sea spray particles contribute low number concentrations to the aerosol budget, increases in particle number concentration likely indicate non-marine aerosol sources that were not excluded by the $< 600 \text{ cm}^{-3}$ clean marine criteria applied here. Therefore we have excluded periods when the total aerosol concentrations exceeds 400 cm^{-3} from the retrieval evaluation.

This additional screening of the LASIC dataset after applying the restriction of fit RSS (< 5), scattering error tolerance ($< 5 \text{ Mm}^{-1}$), and using $\text{CN}_3 < 400 \text{ cm}^{-3}$ provided sea spray mode retrievals for 88% (794) of the available clean marine background periods during LASIC. These 794 sea spray retrievals were used to evaluate UHSAS-NEPH during LASIC.

6 Evaluation of UHSAS-NEPH Sea Spray Retrieval during LASIC

Since sea salt composition measurements were not collected during the LASIC campaign, four methods were used to evaluate sea salt identification: (1) comparison of UHSAS-NEPH to a modified version of a sea spray size distribution fitting algorithm that has been validated previously with salt composition (Saliba et al., 2019; Modini et al., 2015; Quinn et al., 2017), (2) correlation to the non-refractory chloride signal measured by the aerosol chemical speciation monitor (ACSM), (3) correlation of supermicron scattering and sea spray mass, where coarse scattering is taken as a tracer for sea spray during clean marine conditions (Kleefeld et al., 2002; Chamailard et al., 2006; Quinn et al., 1998), and (4) correlation of retrieved mass to wind

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speed, since it is widely used as a proxy for sea spray mass production (Lewis and Schwartz, 2004) and model flux parameterization (Gong, 2003; De Leeuw et al., 2011; Ma et al., 2008).

6.1 UHSAS-only Comparison

We applied the [SEMS-APS](#) fitting algorithm described [in Section 4](#) to [measured](#) UHSAS number size distributions (hereafter identified as UHSAS-only) and compared sea spray mode results with those retrieved using UHSAS-NEPH. [Sanchez et al. \(2021\)](#) have recently applied the algorithm to submicron UHSAS size distributions obtained from aircraft measurements in the marine boundary layer of the Southern Ocean and found it to be a good approximation of sea spray contribution to CCN number concentration by comparison to quantified sub- and supermicron sea salt particles using electron microscopy. The mode fitting parameters N_T and D_g of the UHSAS-only method were converted to mass-derived values for comparison with UHSAS-NEPH using Eq. (7) for the same particle diameter range ($D_p = 0.01 - 10 \mu\text{m}$). UHSAS-only sea spray modes were fit for 90% of the 2-h average size distributions during the clean marine background season. Times that could not be fit generally included noise in the measured size distribution or other common singularities (see supplement of [Saliba et al. \(2019\)](#)). Summary statistics comparing parameters retrieved using UHSAS-NEPH and UHSAS-only are provided in Table [5](#).

Integrated sea spray mode mass concentrations ranged from 0.008 to $23.5 \mu\text{g m}^{-3}$ employing UHSAS-only and 0.18 to $23.0 \mu\text{g m}^{-3}$ with UHSAS-NEPH. Average sea spray mass concentrations of $1.3 \pm 2.2 \mu\text{g m}^{-3}$ and $8.37 \pm 4.1 \mu\text{g m}^{-3}$ were observed for UHSAS-only and UHSAS-NEPH, respectively, which indicates that using UHSAS-only provides lower sea spray mass retrievals because UHSAS-only fits are constrained solely by the accumulation mode shoulder. Sea spray mass concentrations from both methods exhibit generally consistent concentrations during the clean marine periods of LASIC with no apparent seasonality (Fig. [9a](#)), but there are distinct differences in the retrieved mode diameters and widths (Fig. [9b,c](#)).

Sea spray mode retrievals using UHSAS-only were much smaller in mean mass diameter than UHSAS-NEPH with averages of $0.68 \pm 0.01 \mu\text{m}$ and $1.47 \pm 0.17 \mu\text{m}$ respectively (Fig. [9b](#)). The range of mass mean diameters were similar for both retrieval methods, 0.4 to $1.6 \mu\text{m}$ for UHSAS-only and 0.6 to $1.9 \mu\text{m}$ for UHSAS-NEPH, though the majority of UHSAS-only mean diameters were submicron (Fig. [9b](#)). Just over 2% (14) of UHSAS-only retrievals had mass mean diameters in the coarse mode ($> 1 \mu\text{m}$) compared to 91% (723) of UHSAS-NEPH. The peak in mass mean diameters at sizes within the coarse mode using UHSAS-NEPH can be attributed to the additional contributions of supermicron mass identified by the nephelometer supermicron scattering that are not constrained by the UHSAS submicron size distribution. These predominately supermicron mean diameters are consistent with the assumption of sea salt particles contributing a large amount of mass and scattering within the coarse mode, which is simply not captured by the submicron UHSAS distributions alone.

On average, UHSAS-only retrievals were narrower at 1.8 ± 0.4 than those retrieved with UHSAS-NEPH at 2.4 ± 0.3 (Fig. [9c](#)). The range of mode widths using UHSAS-only varied from particularly narrow (1.3) to a fairly broad and unconstrained

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width of 5.3, compared to 1.4 to 3.97 for UHSAS-NEPH. UHSAS-only widths are determined solely by the shape of the large accumulation mode shoulder in the UHSAS number size distribution and include a variety of widths based on how flat or sharp the slope of this shoulder may be. Mode widths retrieved from both methods were predominately narrower than a value of 3 with only 2% of UHSAS-only modes and 1% of UHSAS-NEPH modes greater than this value. Average UHSAS-only mode narrowness again reflects the absence of supermicron size distribution measurements. Sanchez et al. (2021) reported a similarly narrow average mode width (1.44 ± 0.25) for marine boundary layer sea spray aerosol retrieved with UHSAS-only. Conversely, Yu et al. (2019) reported a broad mode width (geometric standard deviation of 2.7) and volume peak at supermicron diameter (approximately $2 \mu\text{m}$) for 24-h averaged sea salt particle size distributions at Ascension Island using Aerosol Robotic Network retrievals, which is more consistent with UHSAS-NEPH observations. These results indicate that UHSAS-only may provide a good estimate of sea spray number concentration, which predominantly consists of submicron-sized particles (Sanchez et al., 2021), but the lack of supermicron measurements makes it unable to adequately identify mass contributions from larger particles.

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Comparing the sea spray mode fitting parameters to those found in literature shows that retrieved modal properties for both methods are within the range of reported values of mass mean dry diameter ($0.25 - 1.6 \mu\text{m}$) and mode width ($1.4 - 3$) (Fig. 9b,c and Table 6), while noting that these LASIC retrievals are for 55% RH. The median of reported mass diameter ($0.88 \mu\text{m}$) falls in between the statistical mode values (peaks in histograms) of the retrieval methods at the upper end of UHSAS-only and lower end of UHSAS-NEPH (Fig. 9b), showing consistency with other ambient sea spray mode measurements (Quinn et al., 2017; Sanchez et al., 2021; Modini et al., 2015). Retrievals using only the UHSAS number size distributions generally show better agreement in terms of mode size with laboratory-based bubble bursting and breaking wave flume studies, which had mass mean diameters that were less than $1 \mu\text{m}$ (Prather et al., 2013; Bates et al., 2012; Modini et al., 2010). UHSAS-NEPH retrievals are more consistent with field observations of the sea spray size distribution across several ocean basins, including open ocean studies in the Pacific and Atlantic, in which measurements of supermicron size distributions were incorporated (Quinn et al., 2017; Saliba et al., 2019; Modini et al., 2015). The UHSAS-NEPH retrieval of sea spray mode width was on the lower end of the reported laboratory and field measurement values, while the narrow UHSAS-only modes were generally outside the spread of the majority of reported values. These differences of the ranges in retrieved values show that the mode width is the least constrained parameter derived by UHSAS-only and UHSAS-NEPH, although the scattering-based constraint provides some apparent improvement compared to UHSAS-only when supermicron mass contributions are considered.

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Figure 10 displays six UHSAS-NEPH sea spray mode fits that are characteristic of the retrieval procedure. Cases of different modal properties (diameter, width, mass) and the comparison of this retrieval with the UHSAS-only algorithm are presented. For both methods, the accumulation mode ($0.4 - 1 \mu\text{m}$) is generally well characterized by the retrieved modes following the shape of this broad shoulder closely. Differences between the UHSAS-only and UHSAS-NEPH retrieved mass size

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distributions become more apparent just before the 1 μm size limit of the UHSAS. For narrow modes ($\sigma_g \sim 2$), using the UHSAS-only method appears to generally be sufficient for quantifying sea spray mode mass concentration in the absence of supermicron scattering measurements (Fig. 10a). In these cases, the lower contribution of coarse particles measured by the nephelometer supermicron scattering adds little information at the tail of the size distribution. The limitations of fitting a sea spray mode based solely on the shape of the accumulation mode “shoulder” in the number size distribution are illustrated in Fig. 10b,c,d. The broadness of the shoulder at submicron sizes (0.38 – 0.83 μm) forces the UHSAS-only retrieval to include more particles from submicron sizes and fewer from the supermicron regime. This leads to lower mass concentrations in UHSAS-only compared to UHSAS-NEPH. Mode retrievals using only the UHSAS size distribution likely underestimate much of the mass at supermicron sizes as seen in the UHSAS-only modes tailing off more sharply in the coarse regime (Fig. 10b,c,e,f), with up to 5 $\mu\text{g m}^{-3}$ of sea spray mass difference in the cases considered here.

6.2 Sea Spray Tracers

We next compared sea spray mode mass concentrations from UHSAS-only and UHSAS-NEPH with available sea spray tracers to evaluate the extent to which the retrieved modal properties represent realistic sea spray size distributions.

The submicron mass concentration measured by the ACSM provided a trace chloride signal that could be used to examine sea salt mass concentrations from the UHSAS-only and UHSAS-NEPH retrievals. Comparing retrieved submicron sea spray mass with the ACSM chloride mass, we found correlations of $R = 0.39$ ($p < 0.05$) for UHSAS-only and $R = 0.35$ ($p < 0.05$) for UHSAS-NEPH (Fig. 11a,b). While the ACSM signal is a very indirect chemical measurement of refractory chloride, these positive correlations provide support for the ability of both methods to identify retrieved modes as sea spray. The higher correlation of UHSAS-only over UHSAS-NEPH may be a result of the submicron sampling range of the ACSM, the diameter range in which UHSAS-only is solely constrained. Additionally, much of the sea spray mass retrieved from UHSAS-NEPH is concentrated in the supermicron regime, which is not observed in ACSM submicron chloride measurements.

Correlations between sea spray mass and the supermicron scattering coefficient were assessed for UHSAS-only and UHSAS-NEPH using retrievals that followed the criteria discussed in previous sections. Sea spray mass concentrations from UHSAS-only were consistently less than 2 $\mu\text{g m}^{-3}$ and were not able to quantitatively explain the variability in the measured supermicron scattering at 550 nm, indicated by a very weak correlation ($R = 0.01$, $p < 0.05$; Fig. 11c). Although the stronger correlation between sea spray mass concentration and scattering with UHSAS-NEPH ($R = 0.84$, $p < 0.05$, Fig. 11d) is an expected result given that the retrieved modes are constrained by scattering measurements, the lack of correlation from UHSAS-only emphasizes the importance of including supermicron particle size measurements to adequately characterize how optical properties are influenced by the sea spray size distribution.

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830 The functional relationship between sea spray mass concentration and wind speed was evaluated using linear regression on UHSAS-only and UHSAS-NEPH retrievals and resulted in correlation coefficients of 0.1 and 0.2, respectively ($p < 0.05$) (Fig. 11e,f). These correlations were lower than values in the range of 0.4 to 0.9 reported for numerous basins of the global ocean (Liu et al., 2021; Russell et al., 2010; Feng et al., 2017; Saliba et al., 2019). The calm and generally invariant wind speed observed during the LASIC background season (7.1 ± 1.4 m/s) could explain the poor correlations for both methods as the lack of dynamic range in wind speed at Ascension Island reduces the degree to which it explains the variability in concentration. 835 The variability may instead be impacted by marine air mass transport of sea spray aerosol from source regions away from Ascension that reach the island rather than local wind conditions, which would not be resolved in the correlation (Grythe et al., 2014). The improved correlation of UHSAS-NEPH in comparison to UHSAS-only shows some added value of applying the supermicron scattering constraint to characterize the sea spray mode. [Evaluation of the method using North Atlantic measurements \(Section 4\)](#), where more dynamic wind conditions [were](#), observed, shows that use of the supermicron scattering constraint could be comparable to the constraint of measured supermicron size distributions when estimating sea spray production from wind speed [as has been done in previous work](#) (Saliba et al., 2019; Modini et al., 2015; Quinn et al., 2017). These results indicate that the incorporation of nephelometer scattering as a constraint for supermicron particle size [provides a reasonable replacement for measured supermicron mass size distributions](#).

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7. Concluding Remarks

845 In this work we have presented a new method that combines measured submicron size distributions and 3-wavelength supermicron scattering to estimate observationally-constrained sea spray modal properties at a remote marine site using a Mie inversion (UHSAS-NEPH). When the retrieval was limited to [marine periods with low aerosol concentration](#) ($CN_3 < 400 \text{ cm}^{-3}$), reasonable sea spray size distributions were obtained [88%](#) of the time. UHSAS-NEPH had larger fit residuals for higher ambient scattering variability ($\Delta\sigma_{\text{sea,RGB}} > 5 \text{ Mm}^{-1}$) and high relative humidity within the nephelometer ($> 60\%$), which affected the consistency with which the Mie solutions could be constrained by scattering measurements and UHSAS submicron mass size distributions. 850

Retrieved sea spray modes ranged in mean mass diameter from 0.6 to [1.9 \$\mu\text{m}\$ \(\$1.47 \pm 0.17 \mu\text{m}\$ \)](#), modal width from [1.1 to 3.9](#) ([2.4 \$\pm\$ 0.3](#)), and mass concentration from [0.18 to 23.0 \$\mu\text{g m}^{-3}\$ \(\$8.37 \pm 4.1 \mu\text{g m}^{-3}\$ \)](#), which are consistent with other field-based measurements of the sea spray aerosol mode. By comparing retrieved modes to available tracers of sea salt, we have shown that estimates of supermicron size, such as from Mie Inversion techniques, are necessary to resolve expected sea spray mass correlations with [scattering and wind speed at Ascension Island](#). The observed positive correlation of UHSAS-NEPH submicron sea spray mode mass with measured submicron chloride signal ($R = 0.35$) provided indirect chemical support of UHSAS-NEPH modes as sea spray. UHSAS-NEPH showed stronger correlations ($R = 0.84$) to the supermicron scattering in comparison to fitting only based on the submicron size distribution accumulation mode ($R = 0.01$), as expected from using the 860

890 scattering measurement to constrain UHSAS-NEPH solutions. Incorporation of scattering measurements as an estimate of
supermicron [mass concentration](#) improved the weak wind speed correlation ($R = 0.2$) relative to using only the submicron size
distribution ($R = 0.1$) for sea spray retrieval. This result was consistent with sea spray production even though the relationship
was not as strong as that found in prior studies with a greater dynamic range of wind speed and available supermicron size
distributions. Other environmental parameters such as sea surface temperature and its impacts on surface tension and kinematic
895 viscosity could also be considered when assessing the relationship between sea spray production and modal properties relative
to these retrieval methods at Ascension Island (Saliba et al., 2019; Liu et al., 2021; Salter et al., 2014), although [salt](#)
[measurements provided a more direct evaluation of the method when applied to North Atlantic observations \(Section 4\)](#).

We have demonstrated that 3-wavelength scattering measurements constrained with submicron size distributions yield sea
900 spray mode estimates that are consistent with sea salt during clean marine periods of LASIC. Inclusion of additional scattering
wavelengths and chemical measurements of particles would provide additional constraints on refractive index and scattering
efficiency by allowing for temporally resolved adjustments in parameters used for Mie simulations. The retrieval procedure
outlined in this work is a self-contained code with look-up table and is available to the broader scientific community. Future
use of this combined size distribution and scattering-based approach to other measured marine datasets that lack supermicron
905 size distributions can expand the array of sea spray observations and improve upon size, mass, and emission characterization
of marine aerosol in climate models, further constraining the uncertainty in natural aerosol impacts on radiative forcing.

Code Availability

The sea spray mode retrieval algorithm and Mie theory scattering look-up table are available as a MATLAB function and
MATLAB matrix file at the UCSD digital archives (<https://doi.org/10.6075/J0GT5NCR>). The Mie codes used to simulate sea
910 spray scattering are available at the same location.

Data Availability

All LASIC data are publicly available from the ARM data discovery (<https://adc.arm.gov/discovery/>) (last access: 5 January
2022). Specific direction to each measurement dataset are provided as DOI references in Table 1. Retrieved sea spray modal
parameters using LASIC and NAAMES measurements can be found at the UCSD digital archives
915 (<https://doi.org/10.6075/J0GT5NCR>). [Scattering and chemical measurements from NAAMES 1 are available at](#)
<https://saga.pmel.noaa.gov/data/download.php?cruise=NAAMES1> (last accessed: 18 April 2022); [SEMS size distributions](#)
[from NAAMES 1 are available at https://library.ucsd.edu/dc/object/bb0856963d](https://library.ucsd.edu/dc/object/bb0856963d) (last accessed: 24 April 2022).

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Author Contribution

Conceptualization and Methodology: **JLD, GS, LMR**. Code development, Formal Analysis: **JLD, GS, ASW**. Supervision and Funding acquisition: **LMR and DL**. Writing – original draft: **JLD**. **All authors** contributed to the review and editing of the manuscript.

Competing interests

The authors declare that they have no conflict of interest.

Financial Support

This research was supported by the Director, Office of Science, Office of Biological and Environmental Research, Climate and Environmental Sciences Division of the U.S. Department of Energy under Contract DE-SC0021045.

Acknowledgements

The authors graciously thank Paquita Zuidema, the LASIC science team, and ARM instrument mentors for their assistance in the interpretation of measured size distributions from SMPS and UHSAS, and nephelometer scattering measurements [collected during the LASIC campaign](#). We thank Timothy Bates and Patricia Quinn for NAAMES I scattering and composition measurements. We also extend our thanks to Roya Bahreini for productive discussions on ways to incorporate measurement uncertainties and variabilities into the Mie Inversion methodology. The authors express thanks to Ian Eisenman for sharing computing resources that made the execution of Mie scattering simulations and sea spray retrieval possible. [Two anonymous reviewers are thanked for their constructive comments that have helped improve the further development of the retrieval method and presentation of results in this study.](#)

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Table 1. LASIC measurements analyzed in this study.

Variable	instrument	measurement	temporal resolution	measured relative humidity (RH) (%)	availability ^a	data access
Particle Size Distributions	Ultra-High Sensitivity Aerosol Spectrometer (UHSAS)	dry particle size distributions	1 Hz	55 ± 8^b	2016 Nov. – 2017 May	https://doi.org/10.5439/1333828
		D _p : 0.06 – 1 μm (optical diameter)				
	Scanning Mobility Particle Sizer (SMPS)	dry particle size distributions	5 min	55 ± 8^b	2016 Nov. – 2017 May	https://doi.org/10.5439/1225453
		D _p : 0.01 – 0.46 μm (mobility diameter)				
Particle Scattering	TSI 3563 3-wavelength Integrating Nephelometer (NEPH)	< 1 μm and < 10 μm dry total scattering coefficients at 450, 550, and 700 nm wavelengths	1 min (alternating impactor size cuts approximately every 55 min)	60 ± 4^c	2016 Nov. – 2017 May	https://doi.org/10.5439/1259232
Ancillary	Aerodyne Research Aerosol Chemical Speciation	mass concentration of non-refractory submicron	15 min		2017 Jan. – 2017 May	https://doi.org/10.5439/1762267

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Monitor (ACSM)	chloride aerosol				
TSI Ultrafine Condensation Particle Counter 3776	condensation nuclei concentration of particles > 3 nm	1 min		2016 Nov. – 2017 May	– https://doi.org/10.5439/1046186
Vaisala WXT-520	wind speed and rain intensity	1 min		2016 Nov. – 2017 May	– https://doi.org/10.5439/1025153
DMT Single Particle Soot Photometer	refractory black carbon concentration	15 min		2016 Nov. – 2017 May	– https://iop.archive.arm.gov/arm-iop/2016/asi/lasic/se-dlacek-sp2/
Los Gatos Research trace gas analyzer	carbon monoxide mixing ratio	1 min		2016 Nov. – 2017 May	– https://doi.org/10.5439/1046183
Radiance Research Particle Soot Absorption Photometer (PSAP)	< 1 µm and < 10 µm dry total absorption coefficients at 470, 552, and 660 nm wavelengths	1 min	≤ 25%	2016 Nov. – 2017 May	– https://doi.org/10.5439/1339528

^afull availability defined as typical Ascension Island background season (November – May).

^bNot reported in dataset. Estimated from internal temperature, ambient temperature, and ambient RH.

210 ^cNephelometer measurements used for retrieval were restricted to periods when the RH was below 60% (Section 2.1.2).

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1215 **Table 2.** Scattering and size distribution measurements and the associated uncertainties and variabilities used in the sea spray
[mode](#) retrieval method.

Measurement			Variability or Uncertainty		
Scattering	$b_{\text{sca},1-10\mu\text{m}}(\lambda)$	supermicron scattering at red (700 nm), green (550 nm), and blue (450 nm) light wavelengths	$\sigma_{\text{sca},1-10\mu\text{m}}(\lambda)$	standard deviation of supermicron scattering for averaging time at each wavelength	Range /Selected Value 0.4 – 40.0 Mm⁻¹
			$\sigma_{\text{sca,inst}}$	instrument-defined scattering uncertainty ▼	5% Deleted: (5%)
	PNSD	particle number size distribution	$\sigma_{\text{PNSD,meas}}(D_p)$	standard deviation of PNSD resolved at each size bin	0.01 – 300 cm⁻³
Size Distribution	D_p	PNSD diameters	$\sigma_{\text{PNSD,inst}}$	instrument-defined concentration uncertainty ▼	10% Deleted: (10%)
			σ_D	instrument-defined sizing uncertainty ▼	2.5% Deleted: (2.5%)

Table 3. Lognormal mode fitting parameters and resolution (step) used to derive Mie scattering sea spray mode solutions.

Parameter	unit	minimum value	maximum value	step
N_t	cm ⁻³	1.0	99	2.0
D_g	μm	0.05	1.19	0.015
σ_g		1.0	4.0	0.075

225 **Table 4.** Comparison of retrieved sea spray modal parameters using measured supermicron size distributions (SEMS-APS; Saliba et al., 2019) and Mie inversion of nephelometer supermicron scattering (SEMS-NEPH; this study) during clean marine periods of the NAAMES 1 cruise (6 November – 30 November 2015). Values of the 2-h integrated number (N_T) concentration, mass concentration (M_T), number mean diameter ($D_{g,number}$), mass mean diameter ($D_{g,mass}$) and mode width (σ_g) are provided as the mean \pm 1 standard deviation.

	N_T (cm^{-3})	M_T ($\mu\text{g m}^{-3}$)	$D_{g,number}$ (μm)	$D_{g,mass}$ (μm)	σ_g
SEMS-NEPH	<u>7.1 \pm 2.1</u>	<u>6.1 \pm 2.7</u>	<u>0.6 \pm 0.1</u>	<u>1.1 \pm 0.3</u>	<u>2.0 \pm 0.3</u>
SEMS-APS	<u>4.2 \pm 3.4</u>	<u>5.5 \pm 3.4</u>	<u>0.5 \pm 0.2</u>	<u>1.5 \pm 0.3</u>	<u>2.4 \pm 0.3</u>

Table 5. UHSAS-only and UHSAS-NEPH number concentrations, mass concentrations, and size distribution fitting parameters. Values are mean \pm 1 standard deviation. Bracketed values are the minimum and maximum ranges.

	UHSAS-only	UHSAS-NEPH
N_T (cm^{-3})	8 ± 7 [0, 151]	6 ± 3 [0, 34]
M_T ($\mu\text{g m}^{-3}$)	1.3 ± 2.2 [0.008, 23.5]	8.37 ± 4.1 [0.18, 23.0]
$D_{g,\text{number}}$ (μm)	0.42 ± 0.10 [0.050, 0.54]	0.51 ± 0.10 [0.16, 1.0]
$D_{g,\text{mass}}$ (μm)	0.68 ± 0.01 [0.4, 1.6]	1.47 ± 0.17 [0.6, 1.9]
σ_g	1.8 ± 0.4 [1.3, 5.3]	2.4 ± 0.3 [1.1, 3.97]

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Table 6. Literature reported values of sea spray modal parameters. Number mean diameters ($D_{g,number}$) were converted to mass mean diameters ($D_{g,mass}$) using Eq. (7), integrating over particle sizes 0.01 – 10 μm , and averaging over a total particle concentration range of 1 – 100 cm^{-3} . Values are averages unless noted as an upper or lower bound.

Reference	Experiment type	Ocean basin	Parameter		
			$D_{g,number}$ (μm)	$D_{g,mass}$ (μm)	σ_g
Lewis and Schwartz (2004) ₂	field measurements (RH: 80%)		0.3	1.3	3
Sellegri et al. (2006) ₂					
Keene et al. (2007) ₂	laboratory-based bubble bursting (RH: variable)		0.05 (lower bound)	0.25 (lower bound)	2.8
Fuentes et al. (2010) ₂			0.1 (upper bound)	0.48 (upper bound)	
Modini et al. (2010) ₂					
Bates et al. (2012) ₂					
Zabori et al. (2012)					
Prather et al. (2013)	laboratory-based breaking wave flume (RH: 10 ± 15%)	N.E. Pacific	0.16	0.88	3
Modini et al. (2015)	field measurements (RH: < 40%)	N.E. Pacific	0.14 (lower bound) 0.26 (upper bound)	0.5 (lower bound) 1.3 (upper bound)	2.5 (lower bound) 3 (upper bound)
Quinn et al. (2017)	field measurements (RH: variable, mostly < 50%)	Pacific, Southern, Arctic, and Atlantic	0.3	1.08	2.5
Saliba et al. (2019)	field measurements (RH: < 40%)	N. Atlantic	0.5	1.6	2.3
Sanchez et al.	field measurements	Southern Ocean	0.6	0.71	1.4

<div> <div></div> <div>(2021)</div> <div>This study</div> </div>	<div> <div>(RH: ambient)</div> <div>field measurements</div> <div>(RH = 55 ± 10%)</div> </div>	S. Atlantic	0.4 (UHSAS-	0.68 (UHSAS-	1.8 (UHSAS-
			only)	only)	only)
			0.5 (UHSAS-	1.47 (UHSAS-	2.4 (UHSAS-
			NEPH)	NEPH)	NEPH)

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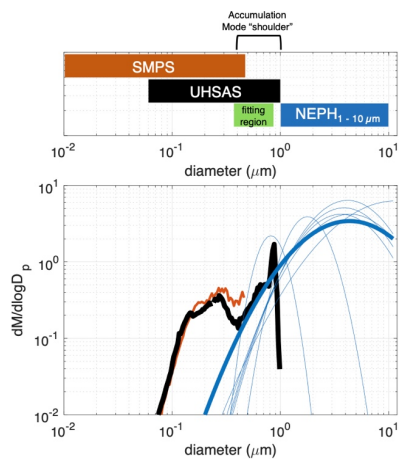
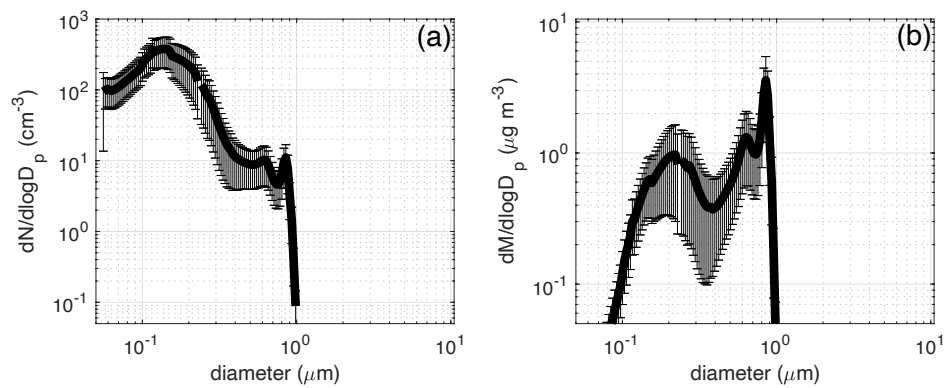


Figure 1. Schematic demonstrating the sea spray mode retrieval method using Mie theory-simulated size distributions, 3-wavelength integrating nephelometer supermicron scattering measurements, and UHSAS submicron mass size distributions (UHSAS-NEPH). The retrieval shown is for a 2-h averaging period beginning 29 November 2016 14:00 UTC. (top panel) Instrument size ranges and mode fitting region for size distributions. (bottom panel) Mass size distributions ($\mu\text{g m}^{-3}$) measured by the SMPS (orange) and UHSAS (black), probable Mie theory-simulated lognormal sea spray mode solutions (thin blue), and best constrained Mie solution (thick blue). Note the UHSAS instrument artifact at $D_p = 0.85 \mu\text{m}$ (see text for description).

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1275 **Figure 2.** Average (solid black line) and variability (1 standard deviation; error bars) of the UHSAS (a) number (cm^{-3}) and (b) mass ($\mu\text{g m}^{-3}$) size distributions during the clean marine background season of LASIC (November 2016 – May 2017).

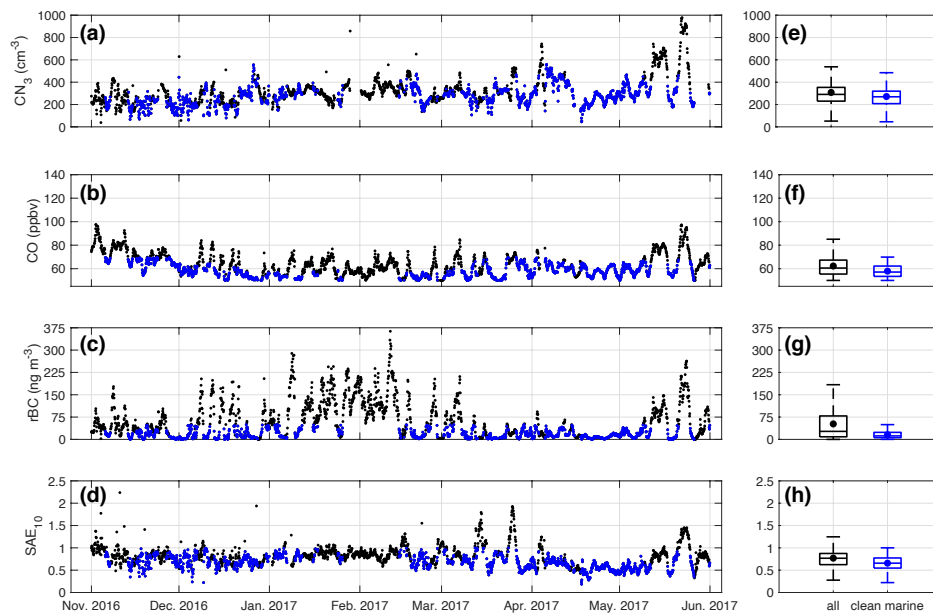
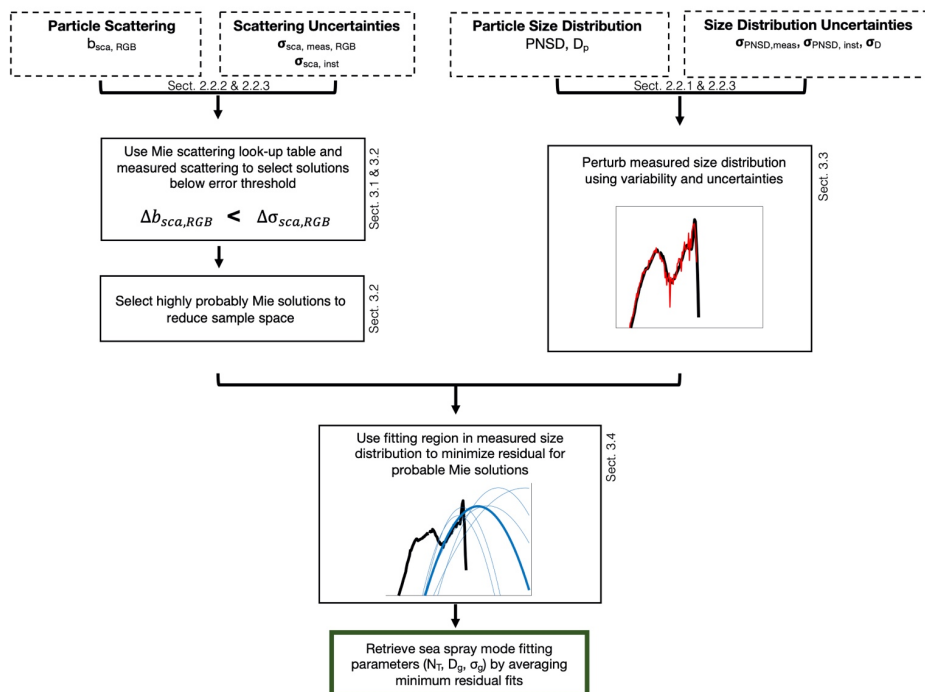


Figure 3. Time series (a-d) and box-and-whisker plots (e-h) of 2-h average variables used to determine clean marine periods during the LASIC background season (November 2016 – May 2017). (a-d) Periods that meet the criteria thresholds described in Section 2.2 are symbolized by blue dots. (e-h) Circles within the box-and-whisker plots are the means and horizontal lines are the median and interquartile ranges (25% and 75%) for the background season (black) and clean marine periods (blue).

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1290 **Figure 4.** Flow chart describing the UHSAS-NEPH retrieval algorithm. Direction to descriptive procedures of each step are identified by main text section at right (below) the boxes.

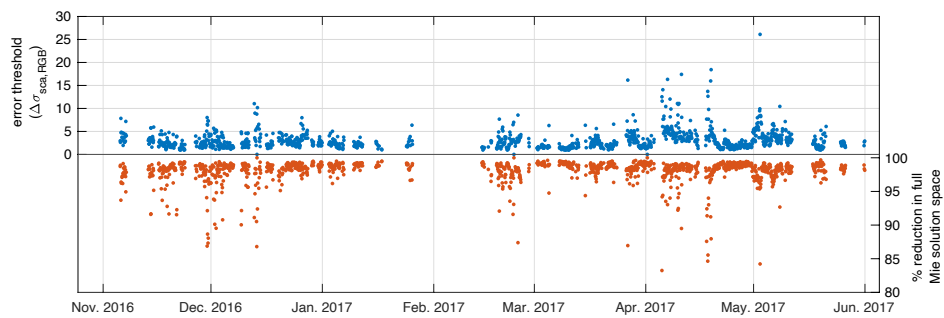


Figure 5. (top, blue) Time series of the scattering error threshold ($\Delta\sigma_{sca,RGB}$, Mm^{-1}) and (bottom, orange) percent reduction of the Mie look-up table solution space ($N = 157,850$) for UHSAS-NEPH retrievals during the background season of LASIC.

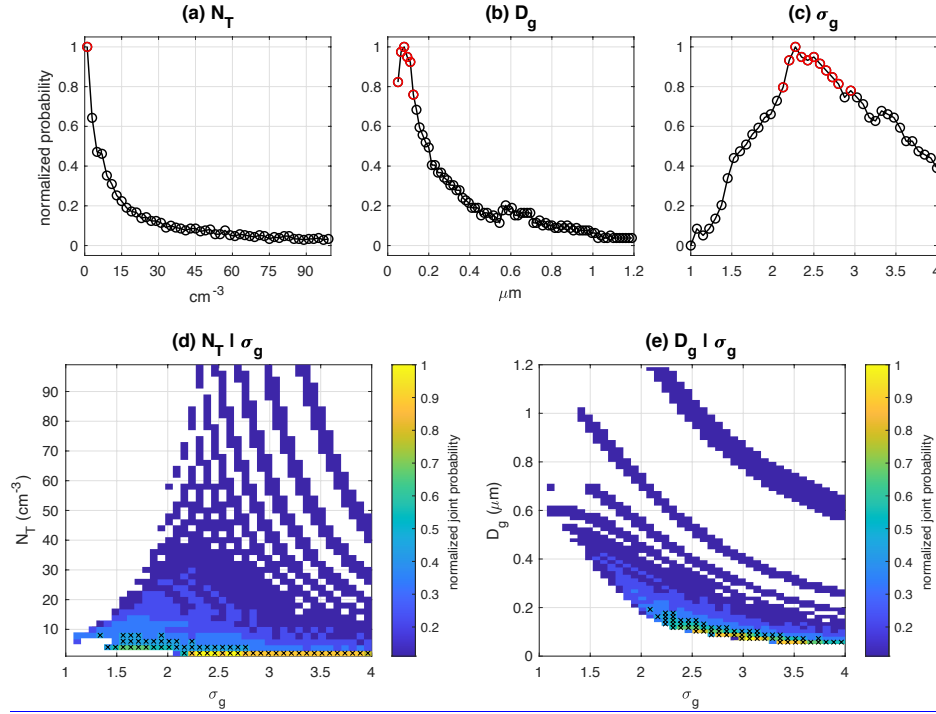


Figure 6. (a-c) Normalized probability distributions of sea spray mode fitting parameters obtained for one retrieval (3 December 2016 22:00 UTC). (a-c) Fitting parameter values with an occurrence probability of greater than 75% are symbolized by red circles. (d-e) Normalized joint probabilities (color bar) for fitting parameter combinations of $N_T | \sigma_g$ and $D_g | \sigma_g$ from the same retrieval. Mie solutions that are within the top 5th joint probability percentile for each combination are symbolized by black crosses.

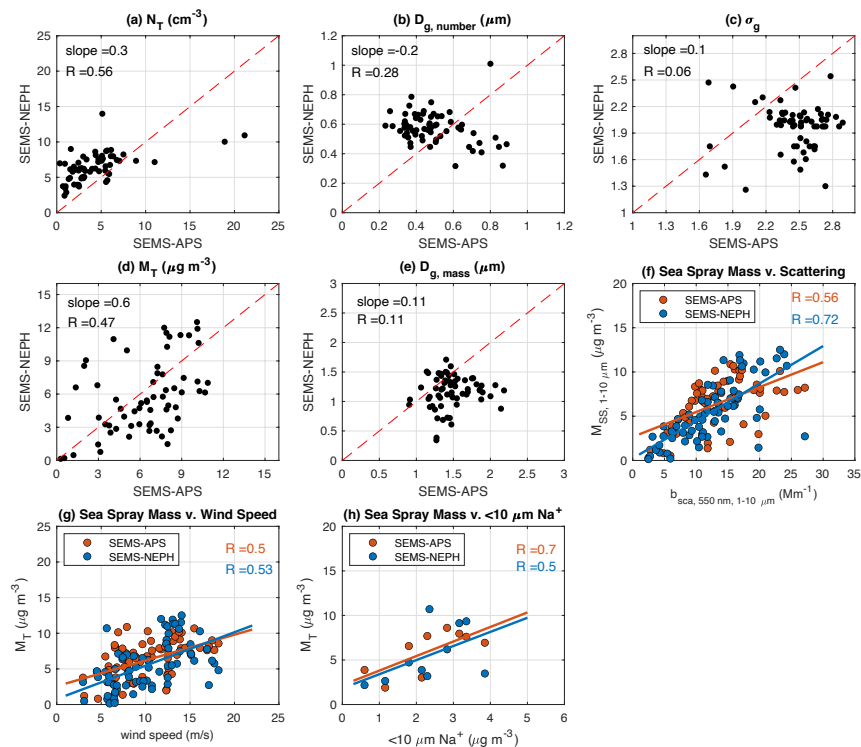


Figure 7. Comparison of retrieved sea spray modal parameters using measured supermicron size distributions (SEMS-APS; Saliba et al., 2019) and Mie inversion of nephelometer supermicron scattering (SEMS-NEPH; this study) during clean marine periods of the NAAMES 1 cruise (6 November – 30 November 2015). 2-h integrated (a) number (cm^{-3}) and (d) mass ($\mu\text{g m}^{-3}$) concentrations, (b) number and (e) mass mean diameters (μm), and (c) mode width. (a-e) Slope of linear best fit and Pearson correlation coefficients at top left. Dashed red lines represent a 1:1 line. Sea spray mass correlation with supermicron scattering (f), wind speed (g), and $<10 \mu\text{m Na}^+$ mass concentration (h) for SEMS-APS (orange) and SEMS-NEPH (blue) methods. Na^+ mass measurements are from offline filter analysis using ion chromatography. Sea spray mass estimates in (h) are averaged over the filter collection times (24-h) which provided 9 samples. Pearson correlation coefficients indicated at top right and linear best fit colored by method in (f,g,h).

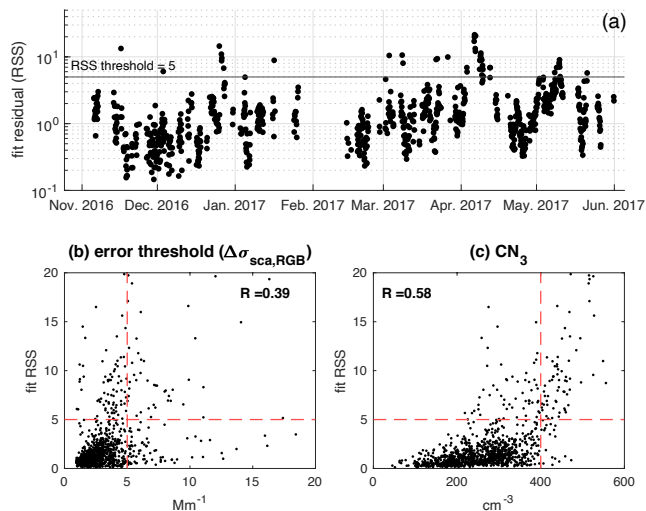


Figure 8. (a) Times series of the fit residual (residual sum of squares, RSS) between the retrieved sea spray mode and UHSAS mass size distribution within the 0.38 – 0.83 μm fitting region. The horizontal line in (a) delineates an RSS restriction threshold of 5 used in this procedure. Note logarithmic y-axis. (b,c) Selected variables used to restrict sea spray mode retrievals defined by the fit residual (fit RSS, y-axes). (b) error threshold ($\Delta\sigma_{\text{sca,RGB}}$, Mm^{-1}) and (c) condensation nuclei concentration of particles $> 3 \text{ nm}$ optical diameter (cm^{-3}). Retrieval restriction thresholds are symbolized by dashed red lines in (b,c). Pearson correlation coefficients ($p < 0.05$) are provided with panels (b,c).

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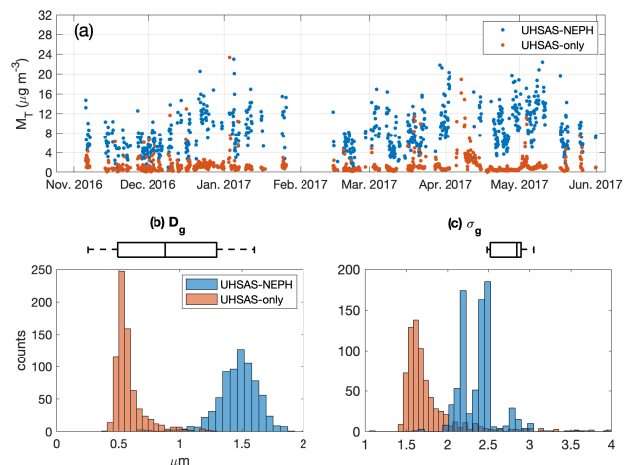


Figure 9. (a) Time series of sea spray mass concentration (M_T , $\mu\text{g m}^{-3}$) retrieved from UHSAS-NEPH (blue) and UHSAS-only (orange) during clean marine periods of the LASIC background season (see Table S3 for summary statistics). Histograms of the (b) geometric mean mass diameter (D_g) and (c) geometric standard deviation (σ_g mode width) for the retrieved sea spray modes using UHSAS-NEPH and UHSAS-only. Box-and-whisker plots above the histograms represent the quartile ranges (25th and 75th) and median (vertical line) of sea spray mode fitting parameters reported in laboratory and field measurements (see Table 6).

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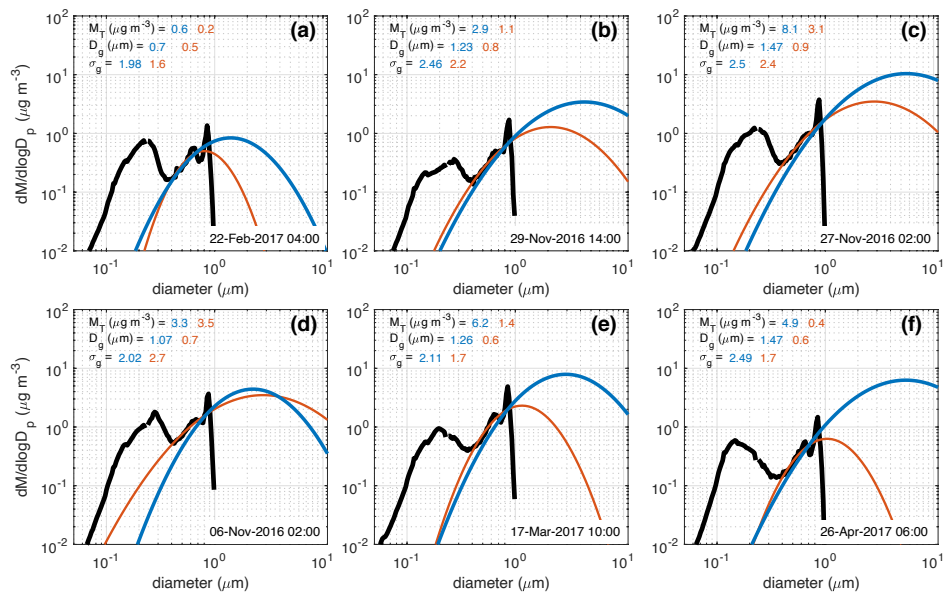


Figure 10. Selected characteristic fits to the sea spray mode. UHSAS mass size distribution (black), UHSAS-NEPH (blue), and UHSAS-only (orange). Panels are oriented to reflect increasing mass, size, and width of the sea spray mode retrieved by UHSAS-NEPH from left to right (a-c, d-f). Mode fitting parameters for the UHSAS-NEPH and UHSAS-only methods are identified in text and specified by method color within each panel: total mode mass (M_T), geometric mean mass diameter (D_g), and geometric standard deviation (σ_g). 2-h average time stamps are provided at the lower right.

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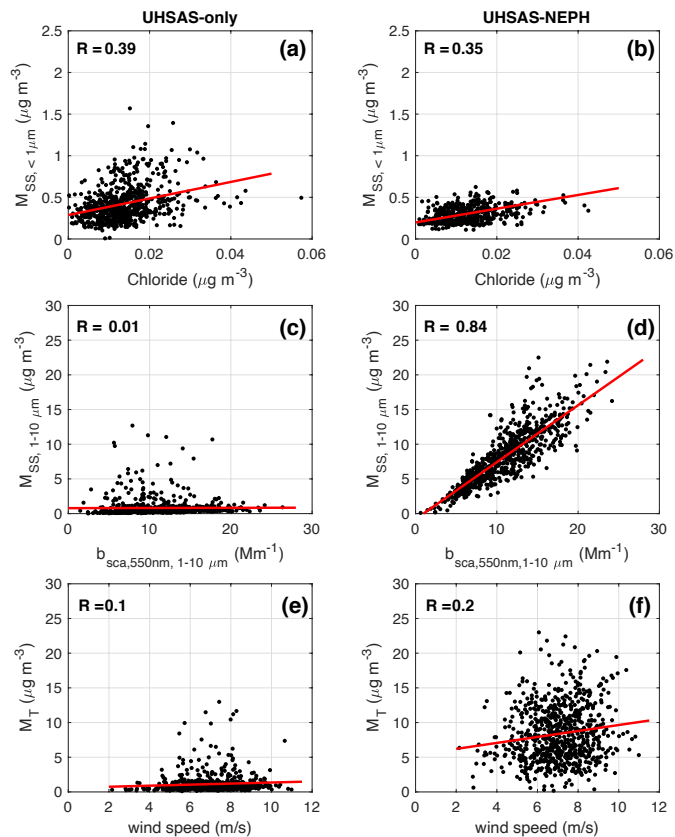


Figure 11. Correlations of submicron sea spray mass ($M_{SS, < 1 \mu m}$; $\mu g m^{-3}$) with ACSM chloride (a,b), supermicron sea spray mass ($M_{SS, 1-10 \mu m}$; $\mu g m^{-3}$) and measured supermicron scattering at 550 nm (c,d), total sea spray mode mass (M_T , $\mu g m^{-3}$) with wind speed (e,f) for UHSAS-only (left) and UHSAS-NEPH (right) methods. Linear regressions are symbolized by red lines and correlation coefficients are provided inside of the panels.

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