Development and Testing of a Novel Sulfur Dioxide Sonde

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13 Correspondence to: James H. Flynn (jhflynn@central.uh.edu) Abstract. A novel technique has been developed to measure sulfur dioxide (SO2) using a modification of the existing electrochemical concentration cell (ECC) ozonesonde technology. The current sonde-based method to measure SO₂ (i.e., the dual-sonde approach) involves launching two ozonesondes together with one of the sondes having a filter to remove SO₂ at the inlet. The SO₂ profile is determined by taking the difference between the measurements from the two instruments. The dual-sonde method works well in typical tropospheric conditions when $[O_3] > [SO_2]$ but saturates when $[SO_2] > [O_3]$ and has large uncertainties in the upper troposphere/lower stratosphere that would limit its effectiveness in measuring SO2 from an explosive volcanic eruption. Due to these limitations, several modifications were made to create a single-sonde system that would directly measure SO₂ (i.e., the SO₂ sonde). These modifications included (1) a positively biased ECC background current, (2) the addition of an O₃ removal filter, and (3) the addition of a sample dryer. The SO₂ sonde measures SO₂ as a reduction in the cell current. There was a strong correlation ($r^2 > 0.94$) between the SO₂ sonde and a Thermo 43c analyzer during controlled laboratory tests and pre-flight tests. Varying humidity levels affected the SO₂ sonde's sensitivity (avg = 84.6 \pm 31.7 ppbv/ μ A, 1σ RSD = 37%) during initial field tests, which was resolved by adding a sample dryer upstream of the O₃ removal filter and pump inlet. This modification significantly reduced the variability and increased the sensitivity of the SO₂ measurements (avg = 47 ± 5.8 ppby/μA, 1σ RSD = 12%). Field tests included measurements near Kīlauea Volcano (before and during the 2018 eruption of the Lower East Rift Zone), Costa Rica's Turrialba Volcano, and anthropogenic plumes from the Athabasca Oil Sands region of Alberta, Canada. This single SO₂ sonde system is an effective, inexpensive instrument for measuring both ground-based and vertical profiles of SO₂ from anthropogenic and natural sources (i.e., volcanic eruptions) over a wide range of concentrations.

1 Introduction

Sulfur dioxide (SO₂) emissions result from anthropogenic activities, such as power generation and crude oil refining processes, and natural sources, such as volcanoes. In gas form, SO₂ acts as a respiratory irritant leading to complications with asthma and cardiovascular conditions (Chen et al., 2007; Sunyer et al., 2003; Tzortziou et al., 2015, 2018). Gaseous SO₂ can be converted to sulfate aerosols (Zhang et al., 2015), which are highly scattering, reduce visibility, and can have a cooling effect on the surface climate when injected into the stratosphere (Kiehl and Briegleb, 1993; Schmidt et al., 2010). SO₂ acidifies rain, accelerating damage of infrastructure and vegetation, particularly near SO₂ sources such as volcanoes (Delmelle et al., 2002; Krug and Frink, 1983; Tortini et al., 2017). Due to these various climate, environmental, and human health-related impacts, anthropogenic SO₂ has been heavily monitored (Shannon, 1999; Zhang and Schreifels, 2011), and regulations have been enacted to reduce these emissions (EPA, 2000).

The largest natural sources of SO_2 are volcanoes. The eruption of Mt. Pinatubo in the Philippines in June 1991 had global climatic effects and significant impacts on the tropospheric and lower stratospheric composition (Bluth et al., 1992; Parker et al., 1996). Apart from such catastrophic eruptions, SO_2 can be continually emitted from volcanoes. SO_2 plumes from over 90

volcanoes have been reliably detected by satellites, resulting in the injection of an estimated 23 ± 2 Tg yr⁻¹ of SO₂ into the atmosphere (Carn et al., 2017). However, unlike anthropogenic sources of SO₂, most volcanoes lack routine ground monitoring (Galle et al., 2010; Pieri et al., 2013) and few opportunities exist for routine validation of satellite retrievals of SO₂ with *in situ* measurements. Small Unmanned unmanned aerial vehicle (UAV) platforms can measure volcanic plumes at altitudes of 2 km above the take-off altitude (Galle et al., 2010; Diaz et al., 2015): while larger UAVs can measure stratospheric plumes (e.g., Global Hawk). However, the lack and difficulty of monitoring and the possibility of another stratospheric injection of SO₂ motivated the development of an inexpensive but reliable balloon-borne instrument that could be deployed quickly after an eruption to validate satellite observations with *in situ* measurements.

Radiosondes and ozonesondes have been widely used for measurements of various atmospheric parameters (e.g., temperature, air pressure, relative humidity [RH], and wind speed and direction.) and O₃ concentrations), respectively. These measurements Electrochemical concentration cell (ECC) ozonesondes produce vertical O₃ profiles and allow for the validation of satellite based O₃ vertical column density (VCD). A schematic of the ECC is included in Figure S1. The current sonde-based method for measuring SO₂, the dual-sonde method, uses two En-Sci (Environmental Science Inc., Westminster, CO) ECC ozonesondes in tandem (Morris et al., 2010). For the dual-sonde method, an SO₂ removal filter is placed at the pump inlet of one of the ozonesondes, scrubbing SO₂ from the sampled air before it enters the electrochemical concentration cell (ECC). The other sonde samples unfiltered air (i.e., air containing both SO₂ and O₃). Due to the chemical reactions in the cathode cell, the filtered sonde measures O_3 , while the unfiltered sonde measures the difference between O_3 and SO_2 ($[O_3] - [SO_2]$) since SO_2 has an equal (relative to O₃) but negative signal in the ECC (Morris et al., 2010). The SO₂ concentrations are then determined from the difference between the two sonde measurements. This method works well in the troposphere when the SO₂ concentration is less than the O₃ concentration, but not as well in intense plumes, such as those found in eruptive volcanic environments. When the SO₂ concentration exceeds the O₃ concentration, the cell current in the unfiltered sonde becomes zero. The excess SO₂ saturates the dual-sonde and distorts the calculated SO₂ profile. Additionally, in the stratosphere, where the O₃ signal grows much larger than in the troposphere, the combined uncertainty of the measurements of the filtered and unfiltered sondes results in a large lower limit of detection (LLOD), on the order of tens of ppbv. Thus, a field deployment of the dual-sonde method more than a few days after an explosive, tropical volcanic eruption such as Mt. Pinatubo would result in little useful data in the critical upper troposphere/lower stratosphere region.

This study reports on the development of a single instrument capable of *in situ* SO_2 measurements in the presence or absence of O_3 . This sonde can measure SO_2 at much greater concentrations than O_3 without saturating the system and can be configured for a sub-ppbv LLOD (calculated using 3σ) at sea level. Since O_3 is removed from the sample stream, this SO_2 sonde avoids the compounded <u>uncertaintieserrors</u> of the dual-sonde method. Field deployments of the SO_2 sonde include sampling of volcanic emissions from Kīlauea on the Big Island of Hawai'i, U.S., Turrialba Volcano in Costa Rica; and the emissions from petroleum extraction and processing at the Athabasca Oil Sands, Canada. Results from these field tests, covering a wide range

of SO₂ concentrations from both natural and anthropogenic emission sources, are described below. The SO₂ sonde has been used for tethered and free-release balloons but can also be adapted for UAV platforms.

2 Instrumentation

2.1 Ozonesondes

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The standard and modified ECC En-Sci ozonesondes were used for the O₃ and SO₂ sonde measurements in this study. The basic functioning of the ECC ozonesonde is described in Komhyr (1969) and Morris et al. (2010). The ECC sensor is composed of platinum cathode and anode electrodes, each in its own cell, immersed in a diluted and saturated solution of potassium iodide (KI), respectively. The cells are connected by an ion bridge allowing for the transfer of electrical charges while maintaining the separation of the solutions (Eq. 1 and 2). When the cells are charged with the solution, a transient potential difference is generated that isbut dissipated through the redistribution of charge across the ion bridge. The following equilibria are established from these reactions:

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$$3I^- \rightleftharpoons I_3^- + 2e^- \text{ (anode)}$$
 (1)

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$$I_2 \rightleftharpoons 2e^- \rightarrow 2I^- \text{ (cathode)}$$
 (2)

94 Sampled air is diffused pumped into the cathode cell, and the presence of O₃ initiates a reaction (Eq. 3) that causes an imbalance 95 in favor of [I₂] in the cathode solution.

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$$2KI + O_3 + H_2O \rightarrow 2KOH + I_2 + O_2$$
 (3)

To rebalance the <u>electrochemical potential of the</u> cell, the iodine/iodide redox reactions in <u>EpEp</u>. 4 and 5 result in a flow of electrons from the anode to the cathode via the ion bridge. This cell current, measured by an external ammeter, is proportional to the O₃ concentration.

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$$3I^- \rightarrow I_3^- + 2e^- \text{ (anode)}$$
 (4)

$$101 I_2 + 2e^- \rightarrow 2I^- \text{ (cathode)} (5)$$

As is also described in in Komhyr (1969) and Morris et al. (2010), when When SO₂ is present in the sample air, an additional reaction (Eq. 6) occurs in the cathode cell of the ECC, supplying the two electrons needed to rebalance the cathode cell after the O₃ reaction (Eq. 3) (Komhyr, 1969; Morris et al., 2010).

$$105 SO_2 + 2H_2O \rightarrow \frac{2}{5}O_4^{2-} + 4H^+ + 2e^- (6)$$

Thus, each SO_2 molecule in the sampled air has the effect of cancelling the measurement of one O_3 molecule. In effect, the standard ECC ozonesonde reports $[O_3]$ - $[SO_2]$ for its measurement. In most places and at most times, $[SO_2] \ll [O_3]$, so there is not a significant impact on the O_3 measurements, but in places downwind of SO_2 sources (e.g., coal-burning power plants or volcanos), the O_3 measurement will be negatively impacted.

2.2 Instrumentation

Several SO₂ and O₃ instruments were used for validation of the SO₂ sonde during laboratory and field testing. A calibration system was used to produce controlled concentrations of SO₂ and O₃. The calibration system relied on the operation of flow controllers or restrictors, an SO₂ ultra-high purity (UHP) gas cylinder (4.87 ppm; Scott-Marrin, Inc., Riverside, CA) and/or a U.V. Photometric O₃ calibrator (49C PS; Thermo Fisher Scientific, Franklin, MA), and zero air to produce desired pre-set concentrations of SO₂ and/or O₃. The zero-air setup used for the field and laboratory testing was achieved using a dry zero air UHP gas cylinder or else generated by scrubbing ambient air through activated charcoal and Purafil SP (Purafil, Inc., Doraville, GA) canisters. The Thermo 43*i*-TL SO₂ analyzer (LLOD: 60-90 pptv at 5 min averaging) and the 49*i* O₃ analyzer (LLOD: 1.5 ppbv at 5 min averaging) were also used during laboratory testing, while a Thermo 43*c*-TL SO₂ analyzer was used during field testing in Hawai'i. These instruments were set to report 10 s average measurements.

3 Single-sonde SO₂ System and Laboratory Testing

3.1 SO₂ sonde system description

The single-sonde The first version (version 1.0) of the single-sonde SO₂ system (i.e. SO₂ sonde v1.0) included threetwo major modifications to the En-Sci ECC ozonesonde: (1) the application of a positively biased background current to the cathode cell, and (2) the addition of an O₃ removal filter, and (3) a sample dryer (Fig. S1). The first version of the SO₂ system (SO₂ sonde v1.0) included the first two modifications: the bias current and an O₃ removal filter. The bias current sets the upper limit of detection (ULOD) for the SO₂ sonde and is set prior to measurement. The O₃ removal filter is placed in line with the inlet allowing O₃-free air to be sampled in the SO₂ sonde. In a standardthe ECC, O₃ produces a positive response signal while SO₂ produces a negative signal when sufficient O₃ is present (i.e., positive signal) is present. With these two modifications, SO₂ can be measured directly as the reduction of the cell current from the pre-set biased background-current (Flynn and Morris, 2021). Unlike the dual-sonde system, this approach allows for direct SO₂ measurements rather than an inference by subtraction of signals from two separate instruments. A sample dryer was added to the SO₂ sonde in the second version (v1.1) to combat humidity issues discovered after initial field tests. The addition of the dryer corrected the highly varying instrument sensitivity observed in the field. All components of the SO₂ sonde fit within a standard ozonesonde foam box (approximately 8" x 8" x 10") except for the inlet filter. The free-release balloon payload's total mass is approximately 1 kg. The patent publication and Fig. S1 provides a detailed description and schematic of the SO₂ sonde (Flynn and Morris, 2021).

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3.2 Testing of background the bias current

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The background bias current is supplied by inserting into the cathode cell an additional platinum electrode in the cathode cell powered by a 9V battery (Fig. S1) (Flynn and Morris, 2021). To maintain consistent power, the circuit uses a 5V regulator. Varying the resistorresistance allows for a range of bias currents to be introduced. The current version of the SO₂ sonde uses a fixed resistor which requires a priori knowledge of the desired SO₂ concentration range. The desired resistor is installed in series with the battery and the electrode allows for a range of bias currents to be introduced. An earlier Laboratory laboratory tests compared the SO₂ sonde measurements (initially configured without an O₃ removal filter) to those made by a 43i-TL SO₂ analyzer (Fig. 1, Table 1). O₃ and SO₂ gases were introduced using the laboratory calibration setup and a manifold to allow the sonde and the Thermo trace gas instruments to sample the same air. Results in Fig. 1 show 60 s averaged data. The test included (A) input of O₃ without an added background bias current; (B) the same input of O₃ with the addition of a background bias current (equivalent to a signal of approximately 90 ppby of O₃); and the addition of SO₂ to the O₃ with the enhanced background bias signal where the SO₂ concentration was either (C) smaller or (D and - E) larger than the O₃ concentration. During (A), measurements made by O₃ and SO₂ sondes compare well to measurements made by the Thermo instruments (Fig.1, Table 1). The test included (E) the response of the SO₂ sonde towith a stepwise reduction of the O₃ concentration resulting in an equivalent decrease in signal, followed by (G - I) a stepwise reduction in the SO₂ concentration resulting in an equivalent increase in signal. At (F), the SO₂ concentration exceeded the biased background current (90 ppbv), producing a signal equivalent to 2.9 ± 0.1 ppby., resulting in no sonde response. During the full test, tThe sonde successfully measured SO₂ both with and without O₃ with approximately 9797% efficiency.

Examination of the SO₂ sonde data showed that noise was proportional to the measured signal, with 1-σ noise at approximately 0.2 – 0.3% of the measured signal. Because increases in the SO₂ concentrations result in a decreases in the signal (i.e., lower cell currents), the magnitude of the applied background current bias current determines the saturation point (i.e., upper limit of detection [ULOD]) of the SO₂ sonde; saturation occurs when the measured cell current drops to zero. Applying a higher background bias current increases the ULOD but also increases noise and the LLOD. The reported LLODs of bias currents are calculated as 3σ relative to the baseline signal when sampling zero air. During laboratory testing, the LLOD (3σ) was calculated for a range of applied background current bias currents (0.25 to 10.0 μA). The LLOD for the varying bias current of 0.25 to 10.0 μA ranged from approximately 0.002 to 0.084 μA, respectively. Results of calculated LLOD of a 0.25 μA bias current at varying replicated altitudes is included in Table S1. At the surface, the LLOD of 20s averaged measurements is 0.17 ppby. The final version of the SO₂ sonde (v1.1) requires the bias current to be selected prior to measurement. If the bias current is set too low, a measurement of larger than expected SO₂ concentrations can saturate the sensor while a bias current that is set too high will have higher LLOD due to the increase in noise. The applied magnitude of the bias current can be best determined based on known SO₂ sources including volcanic emissions, urban and/or industrial emissions.

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3.3 Testing of O₃ removal filter

Since the ECC responds to both O₃ and SO₂, an O₃ removal filter was developed to remove interference from O₃ in the sample. This proprietary O₃ removal filter is placed upstream of the sonde inlet (Flynn and Morris, 2021). During laboratory testing, the O₃ removal filter was exposed to a continual concentration of 487 ± 3 ppbv of O₃ and a varying concentration of SO₂ ranging from 0 to 111 ± 1 ppbv (Fig. 2). The O₃ was effectively and consistently removed from the sampled air by the O₃ removal filter asduring a stepwise dilution of SO₂ was diluted. The testing included measurements with (white-gray background) and without (gray-white background) the O₃ removal filter. The SO₂ and O₃ concentrations measured by the Thermo 43*i*-TL and 49*i* instruments, respectively, and changes in SO₂ dilution levels are also indicated in Fig. 2. The O₃ removal filter destroyed the O₃ at all SO₂ dilution levels to below the detection limit of the O₃ instrument. By comparing the Thermo 43*i*-TL SO₂ analyzer measurements with and without the O₃ removal filter, SO₂ passed through the filter with 88% efficiency (Fig. S1a3a). The transmission efficiency was calculated by taking the ratio of SO₂ measured by the sonde to that measured by the analyzer. The SO₂ transmission efficiency increased to 97% when testing the O₃ removal filter with the dry zero air UHP gas cylinder (Fig. S1b3b) instead of the zero-airzero-air generator that processes ambient laboratory air (Fig. S1a3a). Additional testing of the O₃ removal filter demonstrated that the filter removed approximately 1 ppm of O₃ at sea level with > 99.9% in O₃ removal efficiency, concentrations below the detection limit of the Thermo 49*i* O₃ monitor.

3.4 Sample Dryer

The SO₂ sonde v1.0 had highly varying sensitivities during the initial field tests. The instrument sensitivity was determined by regression analysis of the sonde's cell current to the SO₂ concentration measured by an SO₂ analyzer. The variability in the sensitivities was hypothesized to be due to differing levels of humidity during each SO₂ sonde launch. SO₂ is soluble in water and through multiphase reactions can be oxidized to sulfuric acid in the atmosphere in the presence of water vapor (e.g., precipitation, clouds, fog, etc.) (Carmichael and Peters, 1979; Zhang et al., 2013; Terraglio and Manganelli, 1967). Factors including liquid water content, aerosol composition, aerosol loading, and pH of the water are important in determining the adsorption and oxidation rates of SO₂ (Liu et al., 2021). When air with elevated humidity is flowing through a filter, SO₂ gas is likely adsorbing on the filter causing lower SO₂ transmission efficiency due to the potential uptake of SO₂ in water on the filter. Several laboratory tests confirmed the need to remove water from the sample upstream of the O₃ removal filter to improve the measurement of SO₂. A desiccant membrane dryer (Perma Pure LLC, Lakewood, NJ) composed of a NafionTM tube in silica gel desiccant was placed in-line upstream of the O₃ removal filter. This sample dryer is lightweight, relatively inexpensive, and does not require power.

Laboratory tests included exposing the SO_2 sonde, with and without a sample dryer, to controlled levels of humidity and SO_2 . Without removing water vapor, the SO_2 transmission efficiency decreases as humidity increases, particularly above 50% RH (Fig. 6). As the O_3 removal filter is humidified, the SO_2 transmission efficiency decreases. With the sample dryer in place,

each of the laboratory SO₂ transmission efficiency (May 17-18 and 21, 2018) tests varied by an average of <1% across a range of 0-85% RH (Fig. 6).

The dryer's useful lifetime was determined by continuously exposing it to high humidity (> 95% RH at approximately 23 °C) sample stream. The downstream RH climbed from 5% to 16% after 2.3 h and to 25% after 6.3 h. At these downstream RH levels, the SO₂ transmission efficiency remained above 95%. A typical SO₂ sonde's measurement time per flight, including pre-flight calibration, is approximately three hours. The dryer's useful lifetime is likely much longer than required for a balloon flight since exposure to 95% RH conditions for several hours is highly unusual outside of hurricanes and tropical systems. SO₂ sonde and Thermo 43c-TL measurements were strongly correlated ($r^2 = 0.99$) during a multipoint calibration conducted using the O₃ removal filter and the dryer under relatively high humidity levels. During that calibration, the SO₂ sonde's sensitivity was 45.43 ± 0.17 ppbv/ μ A. By comparison, the average sensitivity during the initial Hawaii deployment was 84.6 ± 31.7 ppbv/ μ A across 10 sondes. The sample dryer, therefore, improved both the sensitivity and stability of the measurements observed. The addition of the sample dryer is necessary for providing accurate ambient SO₂ measurements.

4. Field Deployments, Part I with SO₂ sonde v1.0

Theis SO₂ sonde v1.0), single-SO₂ sonde without the sample dryer, was deployed and tested in Hawai'i and Costa Rica (Fig. S2). The field sites were close to active volcanoes, which are significant sources of natural SO₂ (Tang et al., 2020; Carn et al., 2017). In Hawai'i, field measurements were made near Kīlauea Volcano on the south-eastern shore of Island of Hawai'i, the largest of Hawai'i's islands. Kīlauea is the youngest volcano on the island and one of Earth's most active volcanoes (Kern et al., 2015; Nadeau et al., 2015). Kīlauea had been in a state of eruption since 1983 (Patrick et al., 2019) with an average SO₂ release rate of approximately 5,500 T/d measured during 2014 – 2017 (Elias et al., 2018). In Costa Rica, field measurements were made near Turrialba Volcano, one of the most active volcanoes in the Central American Volcanic Arc. Studies of emissions from Turrialba prior to 2013 reported SO₂ release rates of up to 4,000 T/d (de Moor et al., 2016; Xi et al., 2016). The Activity activity escalated of Turrialba increased after 2014, raising concerns for air quality and environmental

4.1 Kīlauea, Hawai'i - February 2018

health (de Moor et al., 2016; Tortini et al., 2017).

The first deployment of the SO₂ sonde v1.0 was during NASA's HyspIRI HyTES Hawaii Campaign (H3C) from February 3-10, 2018, near Kīlauea Volcano. The instrument was tested in flights on free-release <u>balloons</u> and a tethered balloon system (TBS), and at ground level with measurements in Hawaii Volcanoes National Park (HVNP) downwind of Kīlauea's summit crater, Halema'uma'u. During the ground-level testing, an SO₂ sonde and a Thermo 43*c*-TL SO₂ analyzer's sample inlet were mounted on the top of a van for co-located sampling.

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Figure $\frac{3a \cdot 4a}{2a}$ depicts the measurements taken during the first encounter with an SO₂ plume while driving through the HVNP on February 3, 2018. The strongly correlated SO₂ sonde and Thermo 43c-TL measurements ($r^2 = 0.99$) reached upward of ~940 ppbv. The SO₂ sonde had a sensitivity of 118.4 ± 0.4 ppbv/ μ A, determined by regression analysis of the sonde's cell current with the Thermo 43c-TL concentrations (Fig. $\frac{3a4a}{2a}$). The SO₂ sonde sensitivity varied significantly during the field deployment. During surface measurements on February 10, 2018, earlier zero-air calibrations measured a sensitivity of 86.5 ± 1.5 ppbv/ μ A₂ while measurements during an SO₂ plume event, with peak concentrations of up to 400 ppbv, found the SO₂ sonde's sensitivity was 73.9 ± 0.6 ppbv/ μ A (Fig. $\frac{3b4b}{2}$). Although the SO₂ sonde sensitivity varied significantly in ten subsequent calibrations (84.6 ± 31.7 ppbv/ μ A), the measurements remained strongly correlated (range: $r^2 = 0.94 - 0.99$). The variability in the sensitivity in the field was likely due to changes in the ambient RH impacting the SO₂ transmission efficiency of the O₃ removal filter. This hypothesis was further confirmed by laboratory RH testing and discussed in Sect. $\frac{53.3}{2}$ and $\frac{3.4}{2}$.

4.2 Turrialba, Costa Rica (Dual-sonde versus SO₂ sonde comparison)

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On March 23, 2018, the University of Houston/St. Edward's University team conducted a traditional SO₂ dual-sonde payload (Morris et al., 2010) as well as the SO₂ sonde v1.0 were launched using a free--release balloon flight from the Universidad de Costa Rica's campus in San Jose (approximately 31 km downwind of Turrialba Volcano) consisting of a traditional SO₂ dualsonde payload (Morris et al., 2010) as well as the SO₂ sonde v1.0. This flight provided the first direct in situ comparison of the two SO₂ sonde methods. Figure 4-5 shows the response of the SO₂ sonde v1.0 and the calculated SO₂ dual-sonde profile. The dual-sonde SO₂ method can only report concentrations of SO₂ up to a maximum of the concentration of O₃ present. Furthermore, because the SO₂ concentration is determined by subtracting the signals from two instruments, its uncertainty is higher than the uncertainty of a measurement from a single instrument. When [SO₂] > [O₃], the dual sonde's unfiltered ozonesonde signal goes to zero, as happened for the Turrialba sonde launch between 3 - 5 km (Fig. 45). The SO₂ saturates the cathode solution in the unfiltered sonde, not recovering until enough ambient O₃ has been processed to rebalance the cell, resulting in a distorted profile (Fig. 45). For this flight, the SO₂ sonde was configured to its maximum range (ULOD of approximately 450 ppbv at standard pressure) and was able to capture both the small plume below 2 km above mean sea level (AMSL) (approximately 18 ppbv) as well as the primary plume between 3 – 4 km AMSL (approximately 230 ppbv). The SO₂ sonde v1.0 was able to capture the full shape of the profile, including the peak values and structure of the plume. The SO₂ sonde v1.0 reports the top of the plume around 4 km AMSL, whereas the dual-sonde remains saturated until closer to 5 km AMSL. Thus, the dual-sonde SO₂ profiles, when saturated by high concentrations of SO₂, erroneously appear to have a greater vertical extent. Further, the SO₂ sonde v1.0 showed no interference from O₃ at altitudes from the surface to altitude at 24.4 km AMSL, with O₃ concentrations in the stratospheric O₃ layer reaching > 4 ppmv (not shown), demonstrating the effectiveness of the O_3 filter. The SO_2 VCD was 8.3 DU (Dobson Units, 1 DU = 2.69×10^{16} molecules cm⁻²) for the SO_2 sonde but was only 3.4 DU for the dual-sonde measurement. Thus, once saturated, the dual-sonde method may is likely to underestimate the SO₂ VCD. Additional laboratory testing is planned to resolve this discrepancy.

5. Post Field Test Improvements and Laboratory Testing

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The variability in the SO_2 sonde v1.0's sensitivity during the initial field tests was hypothesized to be due to varying levels of humidity. SO_2 is soluble in water and through multiphase reactions can be oxidized to sulfuric acid in the atmosphere in the presence of water vapor (e.g. precipitation, clouds, fog. etc.) (Carmichael and Peters, 1979; Zhang et al., 2013; Terraglio and Manganelli, 1967). Factors including liquid water content, aerosol composition, aerosol loading, and pH of the water are important in determining adsorption and oxidation rate of SO_2 (Liu et al., 2021). With increased humidity and presence of a filter, SO_2 gas is likely adsorbing on the filter causing lower SO_2 transmission efficiency due to potential uptake of SO_2 in water on the filter. Several laboratory tests were done to confirm the need to remove water from the sample upstream of the O_3 removal filter. A desiccant membrane dryer (Perma Pure LLC, Lakewood, NJ) composed of a Nafion TM tube in silica gel desiccant was placed in line upstream of the O_3 removal filter. This sample dryer is lightweight, relatively inexpensive, and does not require power.

Laboratory tests included exposing the SO₂ sonde, with and without a sample dryer, to controlled levels of humidity and SO₂. Without removing water vapor, the SO₂ transmission efficiency decreases as humidity increases, particularly above 50% RH (Fig. 5). As the O₂ removal filter is humidified, the SO₂ transmission efficiency decreases due to increased SO₂ loss in the filter. With the sample dryer in place, the SO₂ transmission efficiency varies by an average of <1% across a range of 0-85% RH (Fig. 5).

The dryer's useful lifetime was determined by continuously exposing it to high humidity (> 95% RH at approximately 23 °C) sample stream. The downstream RH climbed from 5% to 16% after 2.3 h and to 25% after 6.3 h. At these downstream RH levels, the SO_2 transmission efficiency remained above 95%. A typical SO_2 sonde's measurement time per flight, including pre flight calibration, is approximately three hours. The dryer's useful lifetime is likely much longer than required for a flight since exposure to 95% RH conditions for several hours is highly unusual outside of hurricanes and tropical systems for balloon measurements. SO_2 sonde and Thermo 43c TL measurements were strongly correlated ($r^2 = 0.99$) during a multipoint calibration conducted using the O_3 removal filter and the dryer under relatively high humidity levels. During that calibration, the SO_2 sonde's sensitivity was 45.43 ± 0.17 ppbv/ μ A. By comparison, the average sensitivity during the H3C campaign was 84.6 ± 31.7 ppbv/ μ A across 10 sondes. The sample dryer, therefore, improved both the sensitivity and stability of the measurements observed.

6-5 Field Deployments, with SO2 Sonde v 1.1 Part II

The updated SO₂ sonde (SO₂ sonde v1.1) with the dryer filter was deployed and tested in near Ft. McMurrayMackay, Canada, and again in Hawai'i in June 2018. Ft. McMurrayMackay is in the Alberta province of Canada and is home to the Athabasca Oil Sands, a large area of bitumen and heavy crude oil surface deposits high in sulfur content. Local processing of these products (e.g., surface mining) and resulting by-products (e.g., tailing ponds) can release significant amounts of SO₂ into the

atmosphere (Bari et al., 2020; McLinden et al., 2016; Simpson et al., 2010). A second field deployment to Hawai'i followed immediately after the deployment to Canada. On May 3, 2018, Kīlauea Volcano on Hawai'i entered a new eruptive phase with an outbreak of a series of fissures in the lower Puna area (Liu et al., 2021; Anderson et al., 2019; Gansecki et al., 2019; Patrick et al., 2020). The active phase volcanic gas emissions resulted in localized evacuations in the Lower East Rift Zone (LERZ), destroying more than 700 homes and displacing thousands of residents, and resulting in poor air quality for much of the southern and western portions of the island (Tang et al., 2020). The eruption event entered a paused phase in early August, and was declared over on December 5, 2018 (Kern et al., 2020).

65.1 Athabasca Oil Sands, Canada

The SO₂ sonde v1.1 was tested in Ft. MacKay-Mackay (near Ft. McMurray; 57.1206° N, 111.4241° W), Alberta, in the Athabasca Oil Sands from June 10 - 16, 2018 (Fig. S2c). This field project, conducted in conjunction with Environment Canada and York University, evaluated SO₂ emissions from industrial activities in and near the oil sands region using a combination of TBS and ground-based measurements. The SO₂ sonde v1.1 was flown on the York TBS payload recording measurements from the ground to 300 m above ground level (AGL; 650 m AMSL). This deployment provided a dilute anthropogenic plume to test the SO₂ sonde in a high-sensitivity, low-range configuration. The average sensitivity of the SO₂ sonde v1.1 during the project was 51 ± 1.2 ppbv/ μ A. The SO₂ sonde was configured to sample in a range from ~0.5-25 ppbv of SO₂. The TBS SO₂ sonde's vertical profiles were averaged into 10 m altitude bins that measured SO₂ concentration ranges that are more representative of anthropogenically-impacted SO₂ rather than large volcanic plumes (Fig. 67). This field deployment also demonstrated the performance of the sonde at sub-ppbv levels of ambient SO₂.

65.2 Kīlauea, Hawai'i - June 2018

In response to the larger eruption that started in May 2018, the SO_2 sonde v1.1 was deployed to Hawai'i for the NASA-funded Big Island SO_2 Survey (BISOS). The SO_2 sonde launches occurred from Kahuku Ranch (19.0549° N, 155.6934° W) and Na'alehu Elementary School (19.0610° N, 155.5788° W) approximately 90 km downwind of Kīlauea's LERZ (Fig. S2d). The site's distance from the source allowed the plume to disperse and dilute as compared with measurements at the vent. An SO_2 plume was detected during seven of the nine free-release balloon launches during the June 2018 BISOS campaign. The ten SO_2 sonde v1.1 calibrations performed during BISOS had an SO_2 sensitivity of 47.0 ± 5.8 ppbv/ μ A and were similar to the laboratory results using dry air (45.43 \pm 0.17 ppbv/ μ A).

With the anticipated levels of SO₂, the sondes were configured to sample in theat the maximum range of 10-450 ppbv of SO₂. Figure 7-8 shows four distinctive SO₂ profiles, and Table 2 includes the VCDs for each flight. No plumes above 5 km AMSL were detected, at which point reductions in air density significantly impacted the LLOD. All but one of the observed SO₂ plumes were below the capping inversion of the planetary boundary layer (PBL). On June 22 (Fig. 7a8a), the ascent profile shows SO₂ below 3 km AMSL peaking at nearly 100 ppbv and additional features between 3-4 km AMSL peaking at

20-35 ppbv (Tang et al., 2020). The latter peaks were correlated with higher RH, perhaps the result of steam from a vent or the ocean entry points having broken through the inversion. The early afternoon June 28 profile (Fig. 7b8b) shows the highest concentration (325 ppbv) for a resolved SO₂ plume during the BISOS campaign. Typical for the trade winds, NOAA HYSPLIT trajectories (Stein et al., 2015) showed the winds were out of the NE, consistent with the plume's transport from vents in the LERZ or the lava ocean entry points. Although the descent profile from a June 29 early afternoon launch lost the signal at 0.58 km AMSL, Fig. 7c-8c shows an SO₂ plume over the ocean with a peak concentration of 188 ppbv at 0.74 km AMSL. HYSPLIT trajectories again showed the winds were out of the NE. Lastly, the SO₂ plume detected during the ascent of the June 30 launch (Fig. 87d) exceeded the ULOD between 1-3 km AMSL for the SO₂ sonde configuration used. The distorted SO₂ enhancement extending above the PBL as determined by the temperature inversion is most likely an artifact of the saturated sonde, similar to what was seen in the dual-sonde profile from Costa Rica (Fig. 45). As the RH remains low above the PBL, it is most likely that the SO₂ is contained entirely within the PBL.

56. Conclusion and Future Work

An innovative new method for measuring vertical profiles of SO_2 from TBS and free-release balloons was successfully tested and demonstrated in controlled laboratory experiments and during four different field deployments covering SO_2 concentrations ranging from 0.5-325 ppbv during flights and up to 940 ppbv during ground measurements. This new method requires three major modifications to the standard ECC ozone sonde: the addition of a positive background bias current in the cathode cell, an O_3 removal filter, and a sample dryer. Relative to the previous dual-sonde method, the new method measures SO_2 using a single-sonde system (i.e., the SO_2 sonde). The SO_2 sonde and Thermo 43c-TL measurements were strongly correlated during laboratory ($r^2 > 0.99$) and field-based ($r^2 > 0.94$) comparisons. Initial field tests and subsequent laboratory testing of SO_2 sonde v1.0 highlighted the need to dry the sample upstream of the O_3 removal filter to achieve consistent results. Follow-up field measurements in the Athabasca Oil Sands and Hawai'i clearly demonstrated the improvement in the SO_2 sonde v1.1's sensitivity and consistency (SO_2 and SO_3 and SO_4 and SO_4 and SO_4 and SO_4 are spectively) as a result of drying the sample.

The SO₂ sonde v1.1 offers several advantages over the dual-sonde method, including the ability to measure [SO₂] independent of [O₃], the capability of sub-ppbv detection limits, faster response and recuperation time when exposed to larger SO₂ plumes, and reduced uncertainty. The lighter weight of the payload requires a smaller balloon and less helium to lift, which may prove advantageous for deployment under some field conditions, particularly where helium supplies are limited. It's compactness and weight can also make it a candidate for small drones and UAV campaigns. Field deployments revealed specific issues and areas for improvement. The present design requires pre-setting the sonde's background-bias current prior to the launch. Thus, some *a priori* estimates of the plume are required to determine the appropriate background-bias current so that the instrument can measure the full range of SO₂ concentrations present. In the current SO₂ sonde v1.1, increasing the ULOD by applying a larger background-bias current also increases the LLOD. Further laboratory experiments are needed to identify the factors that

cause the remaining observed variability in the SO₂ transmission efficiency in the latest instrument version that includes the sample dryer. Much of the testing and calibration completed to date assessed the complete SO₂ sonde system (i.e., sonde, filter, dryer). Building a database of the various individual factors, including pump speeds and filter transmission efficiency, will help us to better characterize the causes of sonde-to-sonde variability and allow future versions of the system to improve performance characteristics so that the system can be made available for operational use. Additionally, future manuscripts topics include intercomparison studies of the SO₂ sonde's vertical profile measurements with other column measurements (i.e., Pandora) and satellite measurements and more in-depth analysis of the SO₂ sonde measurements at the various field deployments.

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Author Contributions

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- 369 Conceptualization by J.H.F. and G.M. Data curation by J.H.F., A.K., S.L.A., M.G.S., E.K., P.W., G.M., E.C., A.A., and J.A.D.
- Formal analysis by A.K., S.L.A., S.Y. and P.W. Funding acquisition by J.H.F. Investigation by A.K., S.L.A., M.G.S., and E.K.
- Methodology by J.H.F. and G.M. Writing original draft preparation by S.Y. Writing review and editing by P.W. G.M.,
- 372 J.A.D. and J.H.F. Supervision by J.H.F.

original draft of this manuscript.

373 **Conflict of Interest:** The authors declare that they have no conflict of interest.

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Table 1: Averaged O_3 and SO_2 concentration measured by the SO_2 sonde version 1.0 and Thermo instruments during different stages of testing indicated in Fig. 1.

O ₃ Thermo (ppbv)	O ₃ Sonde (ppbv)	SO ₂ Thermo (ppbv)	SO ₂ Sonde (ppbv)
103 105 ± 0.4	100 ± 1.3	-0. 3 <u>0</u> ± 0.06	99 - <u>96</u> ± 1. 8 3
104-105, ± 0.5	10 <u>1</u> 2 ± 0.4	-0.4- <u>0</u> ± 0.06	190-188 ± 2.3
103 ± 0.4	100-99 ± 0.4	57 ± 0. <u>37</u> 40	138 135 ± 1.0
103-105 ±	98- <u>97.</u> ± 0.6	116 ± 1.9	81- <u>78</u> ± 1.0
0. <u>əs</u>			_
-0.13 13+			5.32.9 ± 0.1
0.5	0.208	110 = 1.1	0.020
0.44 <u>1.1, ±</u>	0-0.51 ± 0.11	58 ± 0.7	30-29 ± 0.65
	0.40-15 ±	24 ± 0.8	67-64 ± 0.86
0.439	0. 04 03,		
$\frac{0.31-1.3}{0.2931}$	1 <u>0,12-64</u> ± 0.278	-0.25 ± 0.22	91- <u>89</u> ± 0.7 <u>6</u>
	(ppbv) 103 105 ± 0.4 104 105 ± 0.5 103 ± 0.4 103 105 ± 0.55 0.441 1 ± 0.429 0.31 1.3 ±	(ppbv) (ppbv) 103 105 ± 0.4 100 ± 1.3 104 105 ± 0.5 1012 ± 0.4 103 ± 0.4 100 + 9, ± 0.4 103 105 ± 98-97, ± 0.6 0.55	$\begin{array}{c ccccccccccccccccccccccccccccccccccc$

Table 2. The SO_2 vertical column density (VCD) for profiles shown in Fig. $\frac{78}{2}$ from BISOS in June 2018. For profile c, the descent profile VCD is reported for the flight without extrapolation (shown without parentheses) and using linear extrapolation assuming the SO_2 concentration to be 0 ppbv at sea level (shown in parentheses).

Profile	Launch Time (UTC)	SO ₂ VCD
a (ascent)	06/22/2018 00:32	8.6 DU
b (ascent)	06/28/2018 20:45	12.5 DU
c (descent)	06/29/2018 21:36	6.2 (9.8*) DU
d (ascent)	06/30/2018 20:48	79.1 DU**

VCD from extrapolated data

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^{**} Saturation of SO₂ at altitudes of 1 to 3 km AMSL

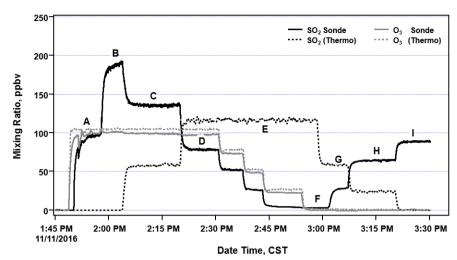


Figure 1: Test of the SO₂ sonde v1.0 (without an O₃ removal filter) with an applied background-bias current responding to O₃ and SO₂. See the text for further details.

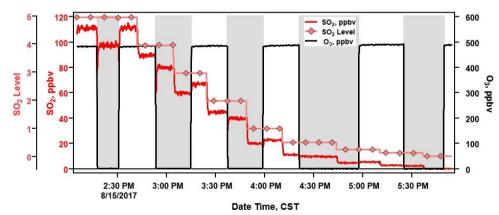


Figure 2: Time of series of a multipoint test of the O_3 filter removal efficiency and impact on SO_2 measurements taken by a Thermo 43*i*-TL SO_2 analyzer. Changes in SO_2 dilution levels are indicated by the blue-pink lines (diamond markers).

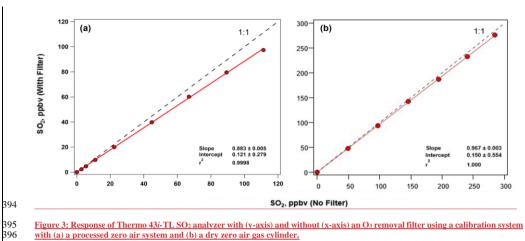


Figure 3: Response of Thermo 43i-TL SO2 analyzer with (y-axis) and without (x-axis) an O3 removal filter using a calibration system with (a) a processed zero air system and (b) a dry zero air gas cylinder.

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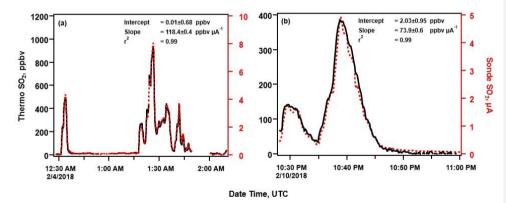


Figure 34: SO2 sonde v1.0 and Thermo Environmental SO2 analyzer measurements at Kīlauca, Hawai'i during H3C for (a) initial SO2 plume encounter on February 3, 2018, and (b) a pre-flight measurement on February 10, 2018, approximately 6 km downwind of Kīlauea's summit crater.

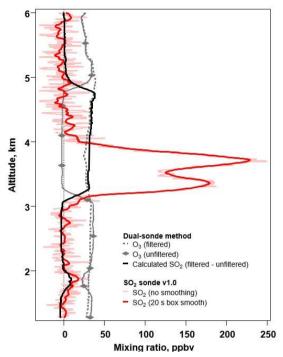


Figure 45: The profiles of a triple-sonde payload, which consisted of a dual-sonde in tandem with an SO₂ sonde v1.0, launched from the Universidad de Costa Rica's campus in San Jose (approximately 31 km downwind of the volcano Turrialba) on March 23, 2018.

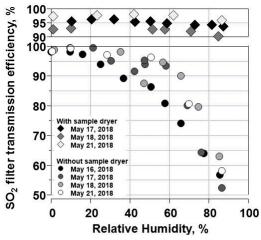


Figure $\frac{56}{2}$: Tests of SO_2 transmission efficiency as a function of relative humidity without (circles) and with $\underline{\text{(diamonds)}}$ an upstream sample dryer- $\underline{\text{(diamonds)}}$.

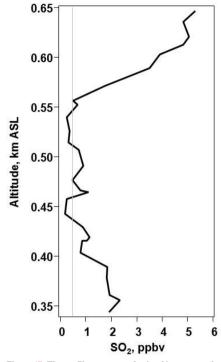


Figure 67: The profile, constructed using 20 s average changes in altitude (ranging from 1 to 15 km), is for a tethered SO₂ sonde v1.1 in the Athabasca Oil Sands region of Alberta, Canada. The SO₂ sonde background bias current was 0.5 μ A, and the LLOD was 0.47 ppbv.

Ambient Tempterature, °C

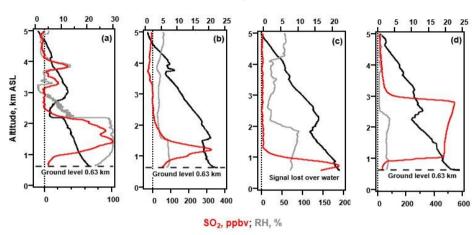


Figure $\frac{78}{2}$: Vertical profiles of SO₂ (20 s box smoothing) from the SO₂ sonde v1.1 during BISOS in June 2018 with free-release balloon launches occurring at the Kahuku Ranch on the Big Island of Hawai'i. Profiles are from (a) $\frac{6}{22}$ 2018 00:32; (b) $\frac{6}{28}$ 2018 20:45; (c) $\frac{6}{29}$ 2018 21:36; and (d) $\frac{6}{30}$ 2018 20:48. All times are UTC.

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