



35 Notwithstanding, research on the characteristics of cloud vertical structures has always been an important direction
36 of cloud physics research (Zcab et al., 2019). Cloud boundaries are the main information in the study of cloud
37 vertical structure, mainly referring to the cloud bottom and cloud top boundary, of course including the side
38 boundary. The cloud boundary in this paper mainly refers to the cloud bottom and cloud top boundary. In the case
39 of multilayer clouds, it also includes the boundary information of intermediate discontinuous clouds (Zhou et al.,
40 2019; Varikoden et al., 2011; Li et al., 2013; Ward et al., 2004; Zhang et al., 2018; Kuji et al., 2013; Kitova et al.
41 2003; Cao et al. 2021). With the development of remote sensing detection technology, MMCR (Görsdorf et al.,
42 2015; Kollias et al., 2017; Kollias et al., 2007) and lidar (Apituley et al., 2000; Motty et al., 2018; Cordoba et al.,
43 2017) are effective instruments for cloud boundary detection.

44 The common methods of detecting cloud boundaries by lidar include the threshold method and differential
45 zero-crossing method. The threshold method (Kovalev et al., 2005) uses the background signal to measure the
46 amplitude of the echo signal. The first point where the echo signal is higher than the background signal and exceeds
47 the set threshold is the cloud bottom boundary. However, in fact, due to the existence of noise, the point with an
48 obvious increase in amplitude may not be found under the condition of a low signal-to-noise ratio (SNR), so the
49 cloud bottom boundary cannot be judged. The differential zero-crossing method proposed by Pal et al. (Pal et al.,
50 1992) differentiates the echo signal to obtain dP/dr , and the zero crossing point from negative to positive is the
51 cloud bottom boundary. The threshold method, differential zero crossing method and variant detection method are
52 all based on feature points of cloud boundaries (Streicher et al., 1995). It is easily affected by noise, and some
53 indicators must be introduced in the specific implementation process to determine the cloud boundary through
54 complex detail debugging, which brings certain difficulties to accurate cloud boundary detection. Young (Young et
55 al., 1995) designed an independent double-window algorithm to detect cloud bottom and top boundaries by
56 combining the lidar signal and a known atmospheric backscatter signal, but the algorithm needs to manually adjust
57 the window size or the selection of the threshold. Based on the WCT (wavelet covariance transform) method,
58 Morille et al. (Morille et al., 2007) determined the local maxima on both sides of the cloud peak as the cloud
59 bottom and cloud top, but the cloud bottom and cloud top detected by this method will be overestimated and
60 underestimated, respectively. Mao Feiyue (Mao et al., 2011) adopted a multiscale hierarchical detection algorithm,
61 selected the starting and ending points of the feature area as the cloud bottom and cloud peak, and realized the
62 detection of cloud top and cloud bottom through multiple iterative updates.

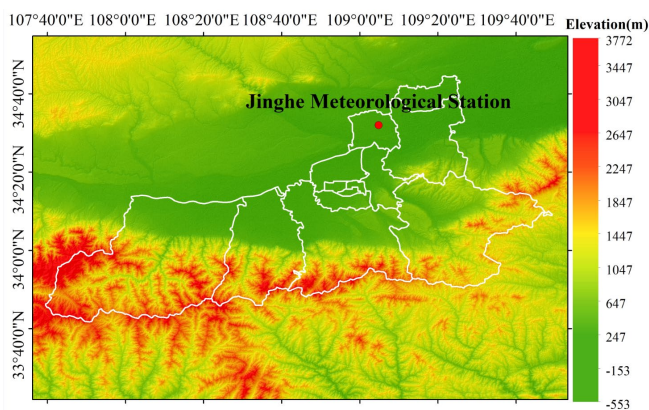
63 The determination of the cloud boundary by MMCR is mainly based on the threshold of the echo reflectivity
64 factor to detect the cloud boundary (Haper et al., 1966; Hobbs et al., 1985; Platt et al., 1994; Brown et al., 1995).
65 Kollias et al. (Kollias et al., 2007) judged the SNR value of a 5×5 grid centered on a distance library. If the SNR of
66 more than 9 consecutive libraries reaches the threshold, the distance library is a cloud signal; otherwise, it is judged
67 as a noncloud signal. Clothiaux et al. (Clothiaux et al., 1999) used 35 GHz millimeter wave cloud measuring radar
68 to analyze different types of clouds and considered that the dynamic range of the cloud reflectivity factor is $-50 \sim 20$
69 dBZ. Due to the existence of certain ground object echoes and biological groups (including insects and other
70 biological particles) in the lower atmosphere, it will interfere with the real cloud echo signal (Luke et al., 2008;
71 Görsdorf et al., 2015; Oh et al., 2016; Melnikov et al., 2013; Melnikov et al., 2015). If the subjective reflectivity
72 factor threshold is directly used to determine the cloud signal, it is not suitable for all cloud types. Therefore, when



73 the cloud signal cannot be accurately identified, it will result in large errors in the detection of cloud boundaries.
74 Research on the macro- and microscopic structures of clouds in a specific area mainly relies on ground-based
75 observations. At present, for better cloud detection, it is necessary to combine lidar and MMCR to observe and
76 study local clouds (Sauvageot et al.,1996; Intrieri et al., 1993; Wang et al., 2000; Sasse et al., 2001). This study
77 will combine the advantages of lidar and MMCR in detecting clouds to achieve high-precision cloud boundary
78 detection and inversion. We effectively identify cloud signals from the power spectrum data of MMCR, and
79 through data quality control, the interference signal caused by floating debris is eliminated to improve the detection
80 accuracy of the cloud boundary. Based on the idea that the MMCR only presents the cloud signal to make cloud
81 boundary detection simple and easy to operate, in this paper, we effectively separate the cloud signal from aerosol
82 and background noise by enhancing and transforming the lidar signal and combining the SNR (Xie et al., 2017) to
83 realize the accurate detection of cloud boundaries. By analyzing the results of cloud boundary detection by two
84 instruments under special weather conditions in Xi'an, the cloud boundary evaluation criteria for the joint
85 observation of the two instruments are established, and the variation characteristics of cloud boundary height over
86 Xi'an in 2021 are statistically analyzed in detail.

87 2 Observation and Instrument

88 Xi'an (107.40 ~ 109.49°E and 33.42 ~ 34.45°N) is located in the Guanzhong Basin in the middle of the Weihe
89 River Basin, bordering the Weihe River and the Loess Plateau to the north and the Qinling Mountains to the south.
90 Xi'an has a semihumid climate. Due to its special geographical location, it is particularly urgent to analyze cloud
91 observations and analyses in Xi'an. The lidar and MMCR are installed at the Jinghe National Meteorological
92 Station in China, placed side by side at a distance of 50 m, and both adopt vertical observation mode to obtain the
93 vertical structure information of sky clouds. Fig. 1 shows the topography of Xi'an and the site location of the Jinghe
94 Meteorological Station.



95
96 Fig. 1. Geographical coverage of Xi'an (107.40-109.49°E, 33.42-34.45°N). The red dot indicates the location of the Jinghe National
97 Meteorological Station in Xi'an.

98 The lidar used in this paper was developed by Xi'an University of Technology. The Ka-band Millimeter-Wave
99 Cloud Radar (MMCR) is the HT101 all-solid-state cloud radar researched by Xi'an Huateng Microwave Co., Ltd.



100 Its main parameters are shown in Table 1 and Table 2.

101 Table 1 Main parameters of the lidar

Indicators	Devices	Main parameter
Launch system	Laser	Nd:YAG; 0.75J@1064nm
Receiving system	Cassegrain telescope	Φ400 mm
	Filter	0.5 nm
Detection system	Detector	APD
	Sampling mode	Analog detection
Spatiotemporal resolution	Time resolution	2 mins
	Range resolution	3.75 m
	Pulse accumulation	2000

102

103 Table 2 Main parameters of MMCR

Indicators	Detailed description	
Radar system	All solid-state; All coherent Doppler; Pulse compression	
Working frequency	35 GHz, and wavelength is 8.6 mm	
Detection altitude range	≥15 km	
Detection blind area	150 m	
Spatiotemporal resolution	Time resolution	5s
	Range resolution	30 m
Scanning mode	Vertical headspace fixed pointing	
Pulse width	1μs, 5μs, 20μs	
Detection accuracy	Z≤0.5 dB, V≤0.5 m/s, W≤0.5 m/s	

104 3 Method

105 Using active instruments to determine cloud boundaries through remote sensing measurements, echo signals in
 106 clear sky areas decay rapidly with increasing detection distance. When the cloud signal is detected, the amplitude of
 107 the echo signal begins to increase sharply. Usually, in the actual observation process, the background noise or
 108 aerosol layer will also increase the amplitude of the echo signal, but the backscattering intensity of the cloud layer
 109 is more continuous and stronger than the aerosol layer and background noise. Therefore, cloud layer and cloud
 110 boundary detection can be realized according to the characteristic changes of echo signals.

111 3.1 Lidar cloud boundary detection

112 When using lidar for detection, the laser beam propagates in a clear atmosphere, and the received echo power
 113 continuously decreases with increasing detection height. However, the beam into the clouds (or aerosols, etc.), the
 114 echo power increases suddenly and becomes stronger at a distance above the cloud bottom. The lidar equation
 115 owing to elastic backscattering can be written as (Motty et al., 2018),
 116

$$P(\lambda, r) = C \cdot \frac{\Delta r \cdot \beta(\lambda, r)}{r^2} \cdot \exp\left[-2 \int_0^r \sigma(\lambda, r') dr'\right] + E(\lambda, r) + N_{back}(\lambda, r) \quad (1)$$



117 where λ is the wavelength of the emitted light, r represents the detection distance, and C is the system constant,
 118 which is determined by the laser energy, the receiving area of the telescope, the quantum efficiency of the detector,
 119 etc. Δr is the detection range resolution of the system, and $\beta(\lambda, r)$ and $\sigma(\lambda, r)$ are the atmospheric backscattering
 120 coefficient and atmospheric extinction coefficient, respectively. $N_{back}(\lambda, r)$ is the background noise received by
 121 the system. $E(\lambda, r)$ represents noise brought to the detection system obtained by calibration.

122 To avoid amplifying the high-level noise signals, we do not perform the distance square correction Eq. (1) and
 123 directly process it as follows:

$$124 \quad P_{new}(\lambda, r) = \frac{P(\lambda, r) - E(\lambda, r) - N_{back}(\lambda, r)}{C \cdot \Delta r} \quad (2)$$

125 For ground-based lidar, the echo signal at a certain height range (>15 km in this study) can be considered
 126 background and electrical noise, $N_{back}(\lambda, r)$ can be estimated with the signal within this range, and the standard
 127 deviation of the noise within the distance range is calculated:

$$128 \quad Sd = \left[\frac{1}{n-1} \sum_{i=1}^n \left(x_i - \frac{1}{n} \sum_{i=1}^n x_i \right)^2 \right]^{\frac{1}{2}} \quad (3)$$

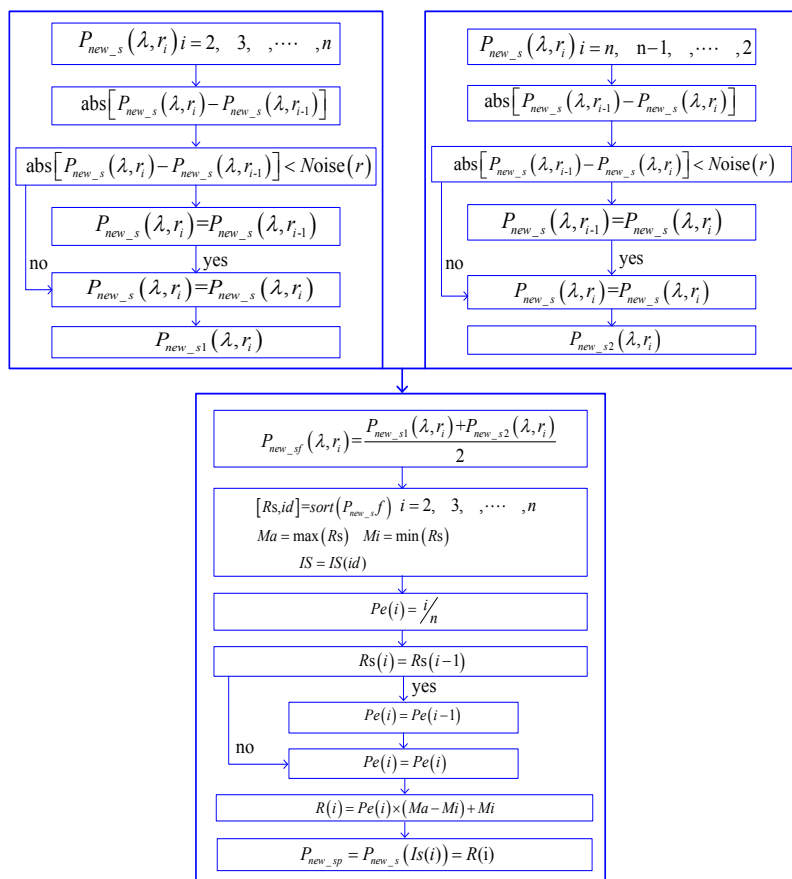
129 where x is the background noise signal. The noise of the lidar signal can be expressed as

$$130 \quad \text{Noise}(r) = k \cdot Sd \quad (4)$$

131 After statistical analysis of the system noise, we set $k=4$ in this paper. Usually, the moving average of $P_{new}(\lambda, r)$
 132 is performed to reduce the influence of random noise. However, the selection of the sliding window directly affects
 133 the quality of the signal. Therefore, in this paper, we use the soft-threshold wavelet denoising method to process
 134 $P_{new}(\lambda, r)$ to obtain $P_{new-s}(\lambda, r)$. To avoid atmospheric turbulence and noise interference, $P_{new-s}(\lambda, r)$ is processed
 135 in one step according to the algorithm flow in Fig. 2, and the enhanced signal $P_{new-sp}(\lambda, r)$ is obtained, as shown in
 136 Fig. 3b) and Fig. 4b). The cloud signal is prominently increased from the background noise and the aerosol signal
 137 compared to Fig. 3a) and Fig. 4a). In this paper, we consider that the echo signal above 15 km is caused by
 138 background and electrical noise. By fitting the echo signal slope in the height range of 15 km~20 km, the slope is
 139 used as the base slope to distinguish the cloud layer and aerosol layer (as shown by the magenta line in Fig. 3b and
 140 Fig. 4b). Without considering the bottom echo signal (0~2 km), the amplitude of the echo signal received by the
 141 lidar will decrease with increasing detection height according to the fitted slope, as shown by the blue line baseline
 142 in Figs. 3b) and 4b). When the beam senses the presence of clouds, the amplitude of the echo signal will exceed the
 143 blue baseline. The SNR of the echo signal is an important parameter to distinguish the cloud layer and aerosol layer
 144 in the echo signal and calculate the SNR of P_{new-sf} with Eq. (5) (Xie et al., 2017),

$$145 \quad SNR(r, \lambda) = \frac{N \cdot P(r, \lambda)}{\sqrt{N \cdot P(r, \lambda) + N \cdot P_{back}}} \quad (5)$$

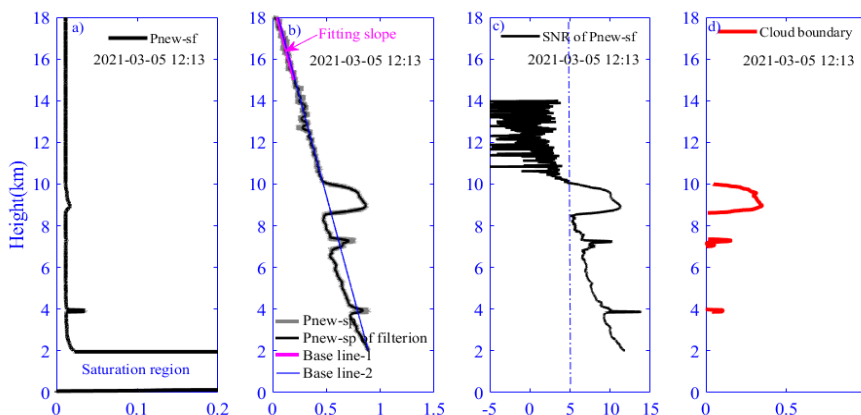
146 where N is the pulse accumulation and P_{back} is the solar background noise power. As shown in Figs. 3c) and 4c), the
 147 SNR of the cloud layer is higher than that of the aerosol layer and background noise, and the SNR in the cloud
 148 layer is approximately equal to 5 (obtained based on multidata statistical analysis in different situations). Combined
 149 with the SNR threshold, the detected cloud information is shown in Figs. 3d) and 4d).



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Fig. 2 The P_{new_sf} of the lidar @1064 nm signal processing flow chart

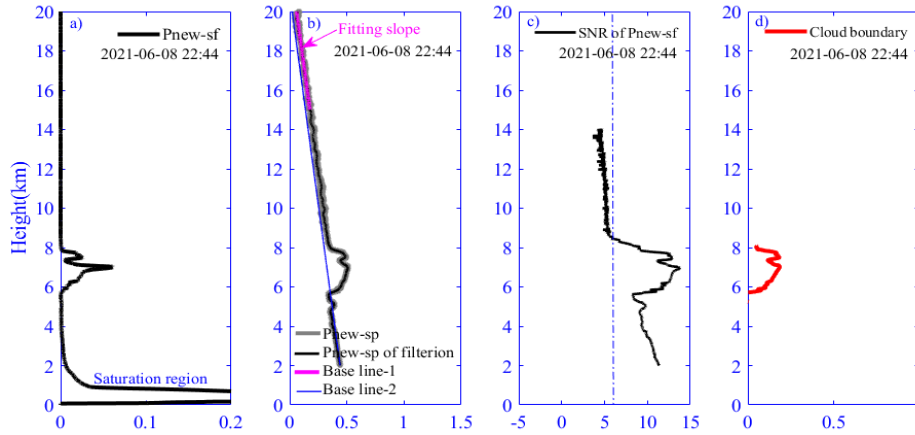


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Fig. 3 Detection results of the lidar at 19:15 on March 4, 2021: a) P_{new_sf} of the 1064 nm signal, b) P_{new_sp} of the 1064 nm signal, c) SNR of P_{new_sf} , and d) cloud information detected



155

156 Fig. 4 Detection results of the Lidar at 22:44 on June 8, 2021 a) P_{new-sf} of the 1064 nm signal, b) P_{new-sp} of the 1064 nm signal, c) SNR
 157 of P_{new-sf} , d) cloud information detected

158 3.2 MMCR cloud boundary detection

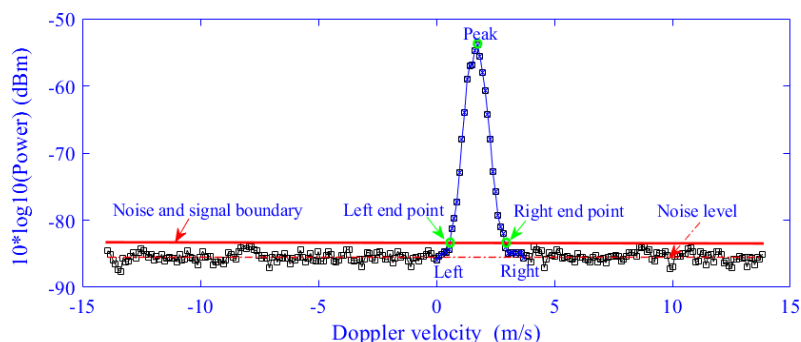
159 Identifying cloud signals from power spectrum of the MMCR is affected by the noise level, especially when the
 160 SNR is low. As shown in Fig. 5, if all the spectral points above the noise level are integrated, it will bring a large
 161 error to the inversion of its characteristic parameters (echo reflectivity, spectral width, radial velocity, etc.).
 162 Therefore, it is necessary to carefully identify the cloud signal in the power spectrum signal. When there is a
 163 meteorological signal in the power spectrum, the general signal has a certain SNR and the number of spectral points,
 164 while the SNR of the noise is very low or the number of continuous spectral points is small, indicating that there is
 165 no meteorological signal (Zheng et al., 2014). Accordingly, by calculating the noise and signal boundary, we count
 166 the number of continuous spectrum signal points greater than the noise and signal boundary. Set the SNR threshold
 167 and the spectral point threshold to evaluate whether each continuous data point is a cloud signal. SNR_{min} refers to
 168 the SNR of the smallest measurable cloud signal in the power spectrum. When the signal is greater than SNR_{min} , it
 169 is considered to have cloud signal; otherwise, there is only noise signal. Fig. 6 shows the algorithm flow chart of
 170 MMCR inversion cloud signal recognition. Referring to the empirical formula proposed by Riddle (Riddle et al.,
 171 1989), the SNR_{min} can be calculated by Eq. (6),

$$172 \quad SNR_{min} = \frac{25 \sqrt{N_F - 2.1325 + \frac{170}{N_P}}}{N_F \cdot N_P} \quad (6)$$

173 where N_F is incoherent accumulation, and N_P is the number of *FFT* (Fast Fourier Transform) sampling points. The
 174 N_F and N_P of the MMCR used in this paper are 32 and 256, respectively, and the SNR_{min} is -17.74 dB by calculating
 175 the SNR_{min} . Adjust the SNR_{min} according to the measured data of the MMCR, and finally determine the SNR_{min} =
 176 -20 dB. Referring to the research results of Shupe et al. (Shupe et al., 2004), N_{ts} is set to 5. When the spectral signal
 177 meets the thresholds of SNR_{min} and N_{ts} , it is considered that there is a cloud signal in the power spectrum, and cloud
 178 feature parameter calculation is performed, flow of cloud signal recognition algorithm is shown in Fig. 6a.



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Fig. 5. Schematic diagram of cloud signal recognition in the power spectrum

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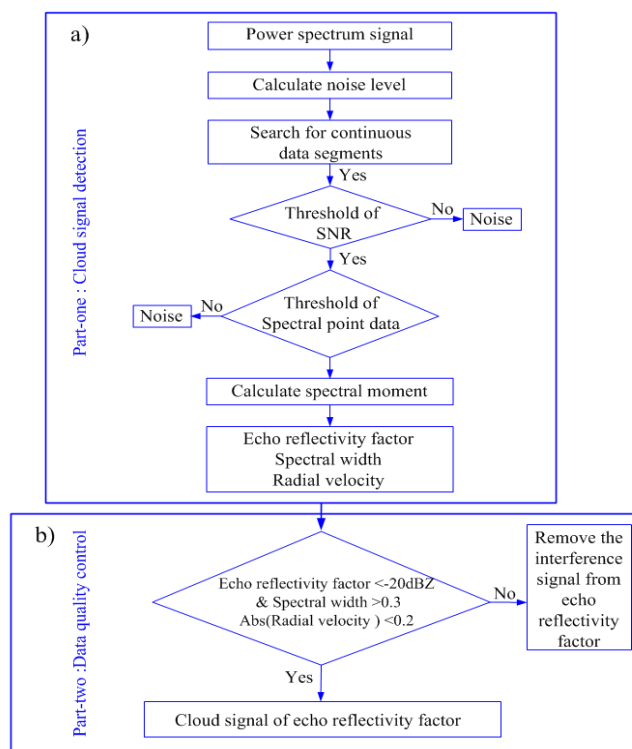
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The echo signals of the floating debris in the bottom atmosphere have the characteristics of a small reflectivity factor, small velocity and large spectral width. To further eliminate interfering wave information, we obtained the data quality control threshold by counting the characteristic changes of planktonic echoes in the boundary layer under cloud-free conditions. As shown in Fig. 6b), when the subjective echo intensity $Z < -20$ dBZ, the absolute value of radial velocity is less than 0.2, and the velocity spectrum width > 0.3 is used as the threshold for removing nonmeteorological information, the expected data quality control requirements can be met. Cloud boundaries are detected using data quality-controlled cloud echo reflectivity factors.

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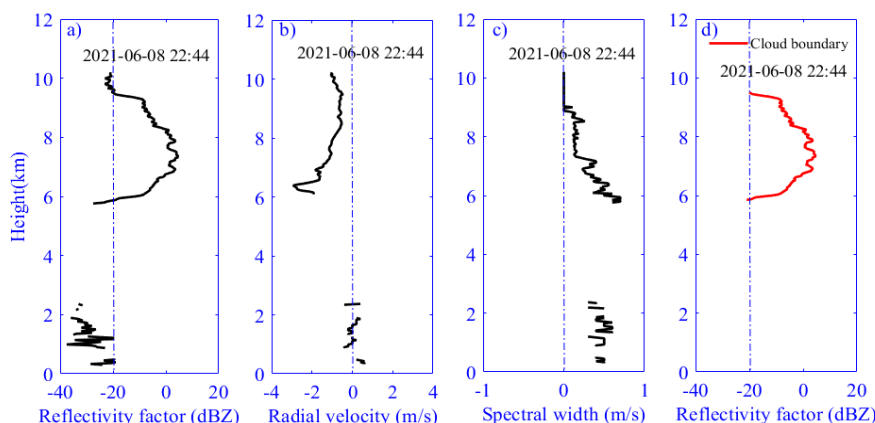


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Fig. 6 Flow chart of MMCR cloud boundary detection



190 According to the algorithm flow of Fig. 6, the power spectrum data at 22:44:00 on June 8, 2021 are analyzed to
191 obtain the meteorological signals of the MMCR reflectivity factor, radial velocity and velocity spectrum width, as
192 shown in Fig. 7a-c). The nonmeteorological signals at the bottom (0~2 km) are effectively eliminated by using the
193 quality control algorithm shown in Fig. 6b). The cloud signal shown in Fig. 7d) realizes the accurate detection of
194 the cloud boundary.



195 Fig. 7 Meteorological signals of MMCR at 22:44 on June 8, 2021. a) echo emissivity factor, b) radial velocity, c) velocity spectrum
196 width, d) reflectivity factor after quality control
197

198 4 Results and discussion

199 4.1 Joint observation and analysis of various types of clouds

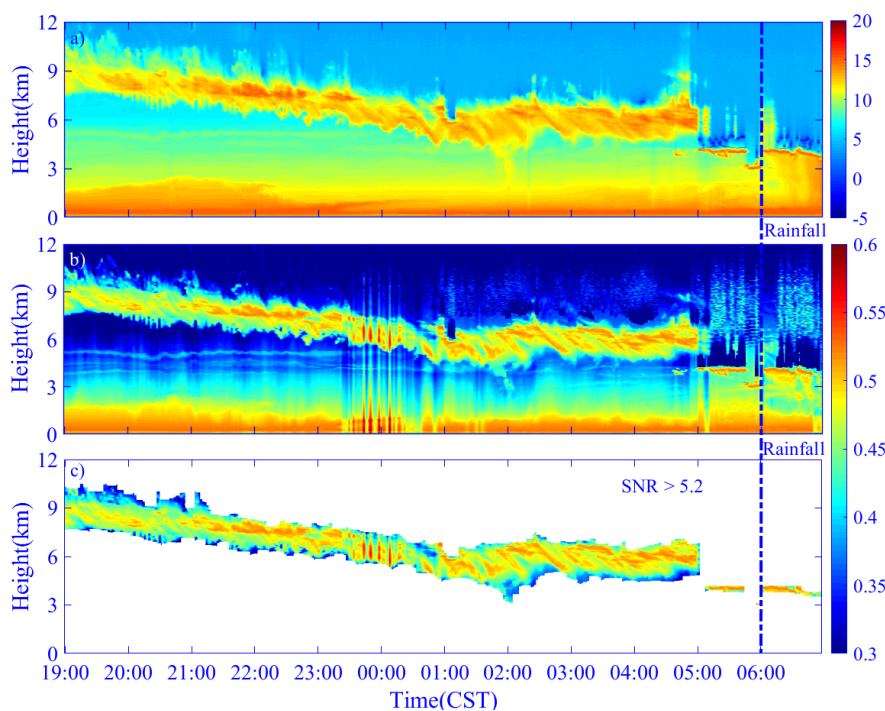
200 Clouds are rapidly changing (Veselovskii et al., 2017). They often appear in the form of single-layer clouds,
201 multilayer clouds and precipitating clouds. Section 4 mainly uses the data inversion method proposed in Section 3
202 to analyze the changing characteristics of clouds under different conditions to obtain reliable cloud macro
203 information. Although the spatial and temporal resolutions of the two detection devices are different, their close
204 proximity allows a good ‘point-to-point’ quantitative comparison between the lidar and MMCR. Before data
205 comparison and analysis, the low spatial resolution of MMCR and the low temporal resolution of the lidar are
206 interpolated to keep the spatial and temporal resolutions of the two consistent (the time resolution is 5 s, and the
207 spatial resolution is 3.75 m).

208 1) Case one studies of a single layer cloud

209 Clouds in the sky often appear as single-layer clouds, and the inversion of macroscopic parameters is simpler than
210 that of multilayer clouds. June 08-09, 2021 (19:00~06:00), the lidar and MMCR jointly monitored the appearance
211 of monolayer clouds in Xi'an. According to the data method described in Section 3.1, the SNR of P_{new_sf} and P_{new_sp}
212 of the echo signal of the lidar @1064 nm are obtained time-height-indicator information (THI) and are shown in
213 Figs. 8a) and 8b). The inversion results show that the thickness of the cloud layer is approximately 2 km, and the
214 height of the cloud bottom decreases from 8 km to 4 km with the passage of observation time. After 05:00, the
215 cloud layer developed deeper, and the laser beam penetrated 0.1 km into the cloud layer and was quickly attenuated.



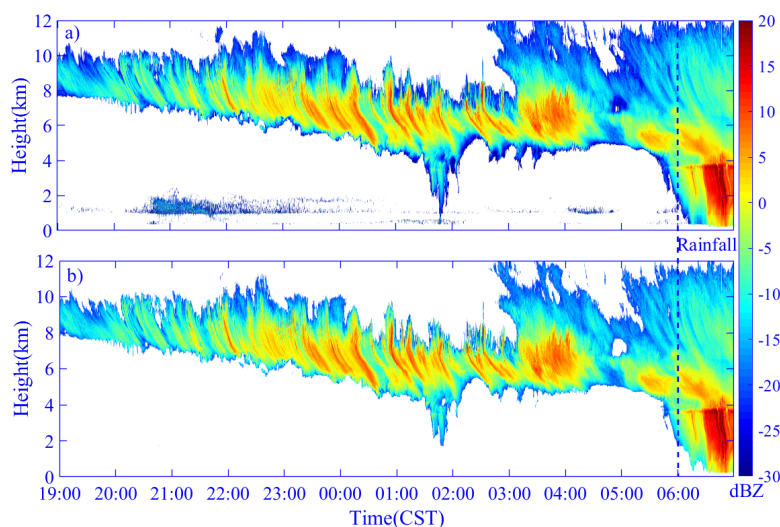
216 Rainfall begins at 06:00, and the lidar cannot continue effective observation and ends the experiment. The SNR in
217 Fig. 8a) causes the SNR of the bottom signal to be large (0~2 km, and the echo signal within the range is not
218 considered in the following cases). The cloud signals have higher SNR than the aerosol and background noise.
219 P_{new_sp} highlights the cloud information from the aerosol signal and background noise and displays the details of the
220 instability of the laser energy from 23:00 to 00:30 in Fig. 8b). Combined with the thresholds of SNR and P_{new_sp} of
221 the cloud signal in Fig. 8a) and Fig. 8b), the cloud layer signal detected from the echo signal is shown in Fig. 8c).



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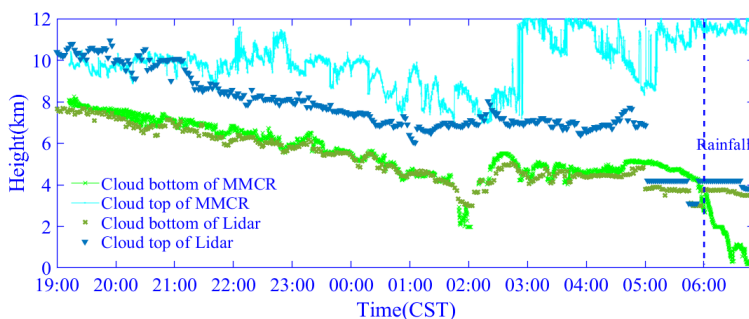
223 Fig. 8 The THI of the echo signal of the lidar @1064 nm from June 08 to 09, 2021. a) SNR of P_{new_sf} , b) P_{new_sp} of the 1064 nm signal,
224 c) cloud information detection results (dotted line indicates rainfall time)

225 Figure 9 shows the cloud echo reflectivity factor of the MMCR at the same observation time period, and the cloud
226 signals observed by the two devices have good macrostructural similarity before 06:00. As shown in Fig. 9a), when
227 the quality control of the echo reflectance factor is not carried out, there are obvious nonmeteorological signals in
228 the range of 0~2 km, and there are also some interference signals around the cloud. If we directly detect the cloud
229 boundary with the echo reflectance factor in Fig. 9a), it will inevitably lead to the underestimation or
230 overestimation of the cloud boundary. We can effectively eliminate the nonmeteorological signals at the bottom
231 atmosphere and the interference signals around the clouds by using data quality control for the echo reflectivity
232 coefficient in Fig. 9b). From the THI of the echo reflectivity of the cloud, the cloud layer starts at 03:00 and
233 gradually develops from 7 km to 12 km (the lidar signal fails to show this detail). When rain appeared at 06:00, the
234 cloud bottom boundary detected by the MMCR became blurred, but lidar could detect effectively (the cloud bottom
235 boundary was ~3.8 km). In this case, we can apply lidar and MMCR to detect cloud bottom and cloud top
236 boundaries, respectively, to achieve high-precision detection of cloud boundaries.



237
238 Fig. 9 The THI of echo reflectivity factor of MMCR from June 08 to 09, 2021. a) echo emissivity factor without quality control, b)
239 echo reflectivity factor with quality control

240 The cloud boundary is retrieved from the cloud signals detected by the lidar and MMCR (Fig. 8c and Fig. 9 b),
241 and the results are shown in Fig. 10. Between 19:00 and 05:00, the cloud bottom boundary height distribution
242 retrieved by the two instruments is agreement. During the period of 21:00~06:00, with the deeper development of
243 clouds, the MMCR can detect more cloud information than the lidar, especially from 03:00 to 06:00. Although lidar
244 cannot penetrate more clouds in this period, it can obtain an effective cloud bottom boundary.



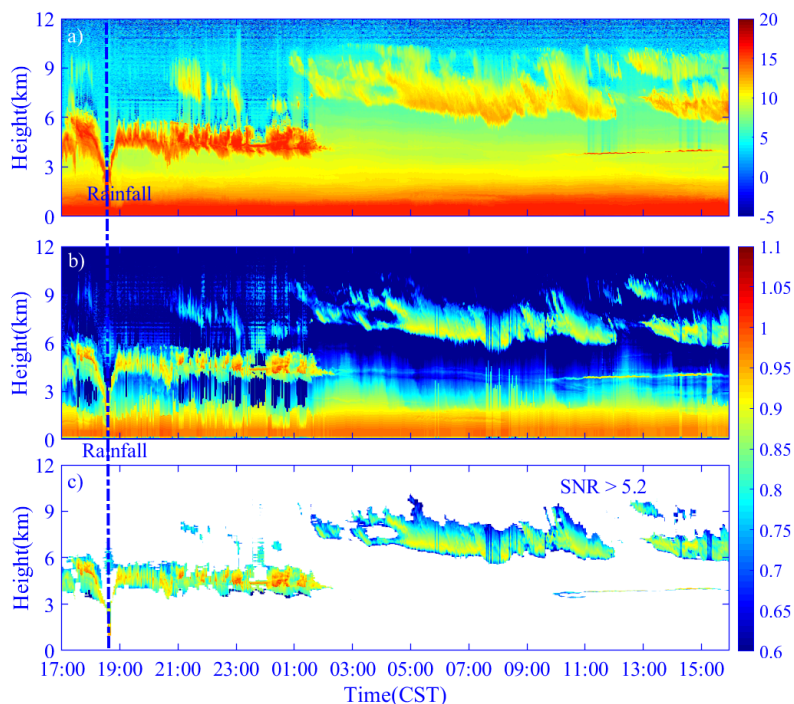
245
246 Fig. 10 Cloud boundary detected by lidar and MMCR from June 08 to 09, 2021

247 2) Case two studies of double-layer clouds

248 From March 4 to 5, 2021, the MMCR and lidar conducted joint observations, with a total observation time of 23
249 hours. By inverting the echo signal of the lidar @1064 nm, we obtained P_{new_sp} of the echo signal and the SNR of
250 P_{new_sf} , and the plotted THIs are shown in Fig. 11a) and Fig. 1b). These THIs display that double layers of clouds
251 appeared in the sky during the observation process. The low-level cloud is located at a height of 4 km, and its
252 thickness is approximately 2 km, the high-level cloud lies at 7 km and its thickness is ~2.7 km. The SNR of the
253 low-level cloud is significantly stronger than that of the high-level cloud, as shown in Fig. 11a. From the
254 characteristic distribution of the P_{new_sp} signal in Fig. 11b), the low-level cloud rained from 18:30 to 18:45, and the



255 cloud bottom height decreased sharply from 4 km to 0.6 km. Subsequently, the cloud layer gradually dissipated
256 from 2 km to 0.05 km, and the dispersal that occurred from 02:00 to 10:00 was too strong for the lidar to detect
257 more detailed information about the low-altitude cloud. We also notice the high-level cloud change characteristics
258 shown in Fig. 11b). During the period from 17:00 to 01:00, there is a relatively weak P_{new_sp} signal in the height
259 range between 7 km and 10 km. This indicates that the high-level cloud may be in the formation stage at this time,
260 and the particle diameter and number concentration of clouds are so small that the lidar can only receive a very
261 weak echo signal. As the observation progresses, the development of high-level clouds is relatively mature, and the
262 structure is relatively stable from 01:00 to 15:00 (except 13:00). Combined with the thresholds of the SNR and
263 intensity information of the cloud signal in Fig. 11a) and Fig. 11b), complete cloud signal detection can be realized,
264 as shown in Fig. 11c).

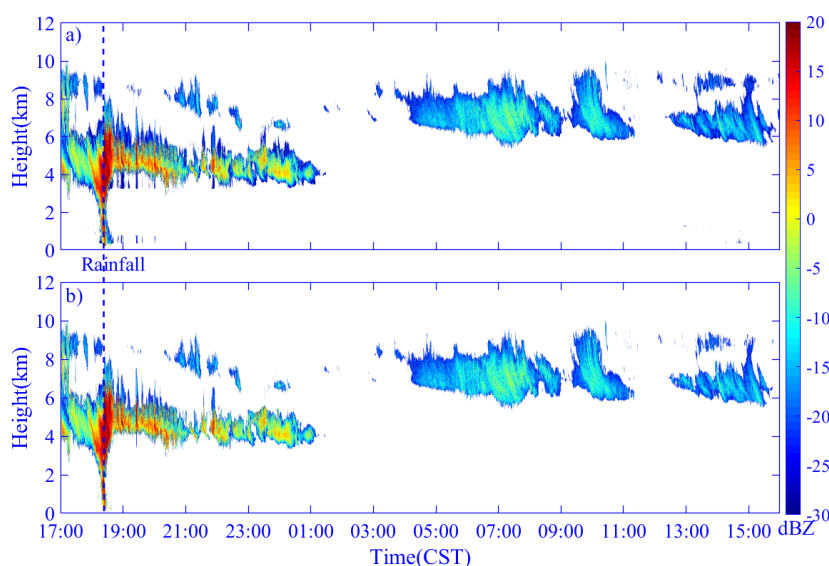


265
266 Fig. 11 The THI of the echo signal of the lidar @1064 nm from March 4 to 5, 2021. a) SNR of P_{new_sf} , b) P_{new_sp} of the 1064 nm signal,
267 c) cloud information detection results (dotted line indicates rainfall time)

268 During the lidar observations, the MMCR also observed double clouds. Figs. 12a) and 12b) show the signal
269 distribution characteristics of the echo reflectivity of MMCR without quality control and after quality control,
270 respectively. It can be seen that in Fig. 12b), after data quality control, the nonmeteorological signals and
271 interference signals at the bottom are effectively eliminated. From the joint observation results of the lidar and
272 MMCR, it can be seen that the appearance and shape of clouds observed by the two are similar, and the occurrence
273 of rainfall is monitored from 18:30 to 18:45. During the period from 15:00 to 01:00, the penetration ability of the
274 MMCR is obviously better than that of the lidar, and more high-level cloud information is obtained. However,
275 between 01:00 and 04:00 at high-level clouds (approximately 8 km), the MMCR detected only part of the debris
276 cloud echo signal, while the lidar detected more cloud information. We can speculate that the main reason is that

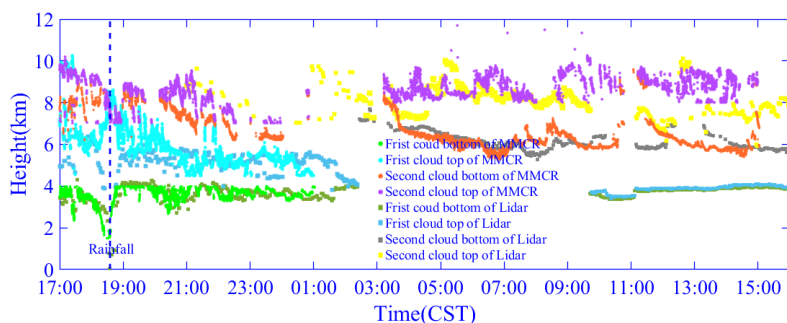


277 clouds were in the growth stage during this time period, their particle diameters were small or their concentrations
278 were low. The echo signal of the MMCR is proportional to the 6th power of the particle diameter, while the echo
279 signal of the lidar is proportional to the 2nd power of the particle diameter, so the lidar can detect clouds that the
280 MMCR cannot detect. From 10:00 to 15:00, the MMCR also failed to detect the thin cloud signal at the lower layer
281 (a height of approximately 4 km). Another reason for MMCR failing to detect thin clouds may also be that the
282 spatial resolution is lower than that of lidar, which makes it unable to detect thin clouds.



283
284 Fig. 12 The THI of echo reflectivity factor of MMCR from March 4 to 5, 2021, a) Echo emissivity factor without quality control, b)
285 Echo reflectivity factor with quality control

286 Based on the cloud signals (Fig. 11c and Fig. 12b) jointly observed by the lidar and MMCR, the height
287 distribution of the double-layer cloud boundaries is detected, as shown in Fig. 13. From the cloud boundary height
288 distribution, it can be seen that the cloud boundary height distribution detected by the lidar and MMCR is relatively
289 consistent for low-level clouds. For high-level clouds, the height of the cloud bottom boundary detected by the two
290 instruments is similar, and the cloud top boundary detected by MMCR is higher than that detected by lidar.
291 However, compared with MMCR, lidar has total supremacy in detecting the information of thin clouds.

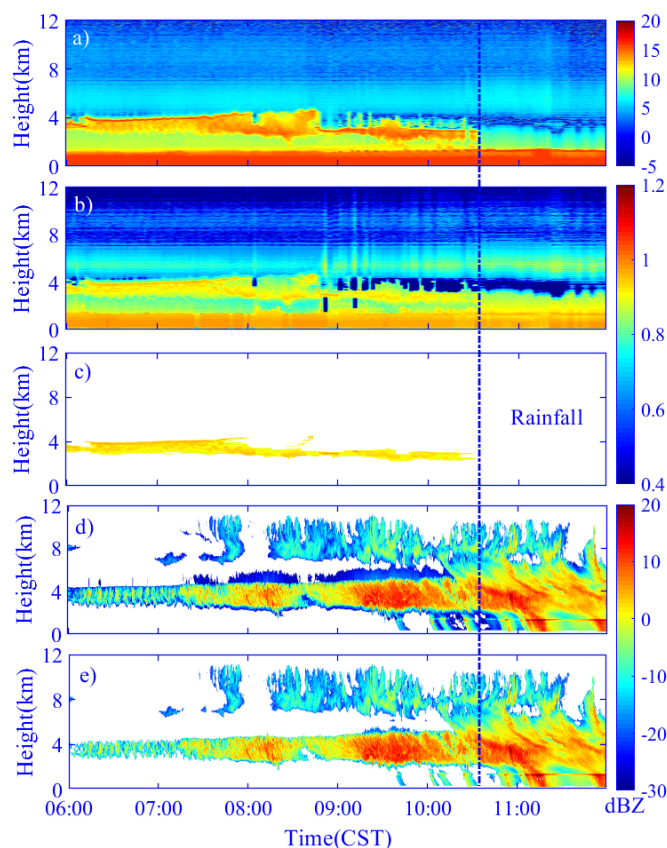


292
293 Fig. 13 Cloud boundary detected by the lidar and MMCR from March 4 to 5, 2021



294 3) Case three studies of precipitating cloud

295 On March 10, 2021, the lidar and MMCR jointly observed the clouds before rainfall for 6 hours (06:00~11:00,
296 and began to rain at 10:45). Fig. 14a) is the distribution of the SNR of P_{new_sf} with time and space, Fig. 14b) lays
297 out the THI of P_{new_sp} of the @1064 nm echo signal, and Fig. 14c) shows the cloud signal detected by the
298 thresholds of the SNR and P_{new_sp} . We inverted the echo reflectivity factor of MMCR and performed data quality
299 control operations on them. The results are shown in Fig. 14d) and Fig. 14e), which are the echo reflections of
300 MMCR without quality control and quality control, respectively. From the comparison, it is obvious that the data
301 quality control can eliminate the interference signal very well, which makes the process of merging the high-level
302 convective cloud and the low-level stratiform cloud clearer.

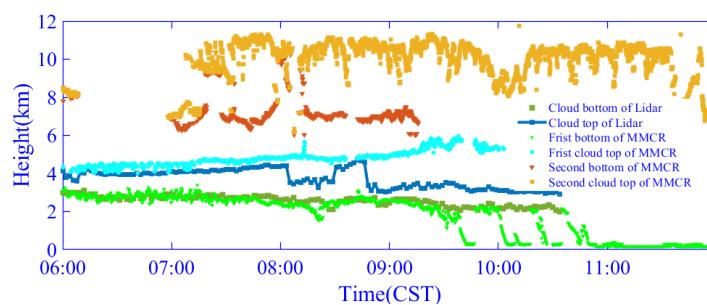


303
304 Fig. 14 The THI of echo signal of the lidar and MMCR on March 10, 2021. a) SNR of the 1064 nm signal, b) P_{new_sp} of the 1064 nm
305 signal, c) cloud information detection results of the lidar, d) echo emissivity factor of the MMCR without quality control, e) echo
306 reflectivity factor of the MMCR with quality control (dotted line indicates rainfall time)

307 By comparing the cloud information detected by the lidar and MMCR (in Fig. 14c) and Fig. 14e)), we can see
308 that during the period from 06:00 to 10:00, the energy of the lidar beam is seriously attenuated at a height of
309 approximately 4 km, resulting in a very weak echo signal and SNR above 4 km. As the observation time progresses,
310 the phenomenon of rain storage (>15 dBZ) occurs in the cloud (Ellis et al., 2011; Williams et al., 2014). The severe
311 attenuation of the lidar in the cloud leads to a sharp decrease in its detection ability, while the millimeter wave still



312 has a strong penetrating ability. When rainfall occurs (at 10:45), neither the lidar nor MMCR can effectively
313 identify the cloud bottom boundary, but MMCR can still detect the cloud top boundary information. The height
314 distributions of cloud boundaries detected by lidar and MMCR are shown in Fig. 15. The height distribution of the
315 cloud bottom and cloud top boundary detected by the two instruments is almost the same from 06:00 to 09:00 (the
316 cloud bottom boundary is approximately 3 km, and the cloud top boundary is approximately 4.1 km). There was a
317 drizzle falling from 09:00 to 10:45, and the lidar obtained an effective cloud bottom boundary. The boundary of the
318 high-level convective cloud at ~8 km and the deep cloud layer from 10:45 to the end of observation can only be
319 detected by MMCR.



320
321

Fig. 15 Cloud boundary detected by the lidar and MMCR on March 10, 2021

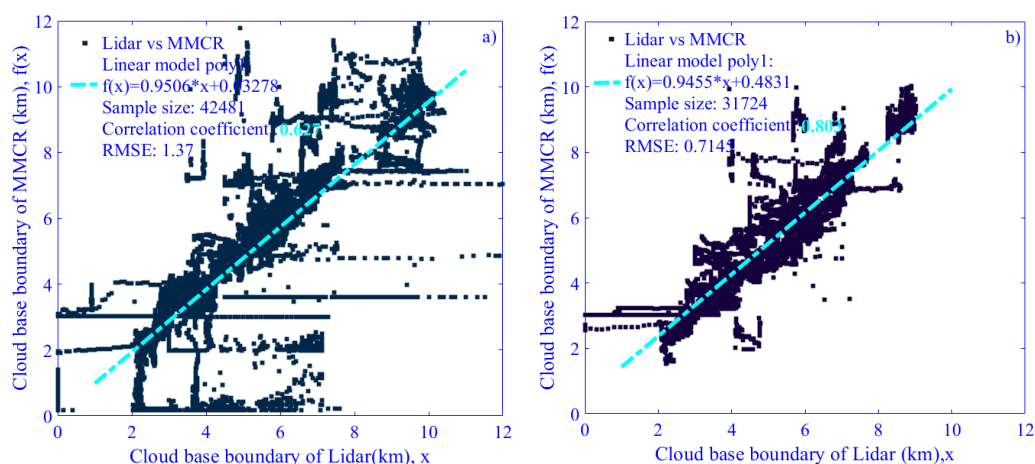
322 From the differences in the height distribution of the cloud boundaries reached by the two devices in the above
323 three different situations, it can be seen that when a single layer of stratiform clouds appears in the sky, the heights
324 of the cloud bottom boundary detected by MMCR and lidar are approximately the same. When there are multilayer
325 clouds, MMCR and lidar have good consistency in the detection results of the cloud bottom boundary height of the
326 low-level cloud, but the energy of the lidar beam of attenuates seriously in the low-level cloud, resulting in the
327 inability to fully obtain the effective bottom boundary of low-level clouds and the height boundary of high-level
328 clouds. In this case, MMCR can obtain more complete height information of the multilayer cloud boundary.
329 Usually, the closer to rainfall, the deeper the cloud layer develops, the beam of the lidar will be seriously attenuated,
330 and more cloud information cannot be obtained. In other words, MMCR still has the ability to penetrate the cloud
331 layer and detect the complete cloud information at this time. Therefore, the joint observation of the lidar and
332 MMCR can comprehensively identify and detect cloud boundary conditions in detail. The difference between the
333 cloud boundaries detected by the two may also be due to the different scattering mechanisms of cloud particles to
334 millimeter-wave electromagnetic waves and laser beams or the difference in the methods used by the two devices
335 to determine the cloud boundary, so that there are some differences in the cloud boundary height results.

336 4.2 Statistics and analysis of cloud boundary distribution characteristics in Xi'an

337 In 2021, the Lidar and MMCR radar conducted cloud observation experiments at the Jinghe meteorological
338 station, in which the MMCR accumulated 302 days of data (7248 hours in total) and the lidar observed 126 days
339 (872.5 hours in total). Due to some unavoidable external reasons, the lidar failed to carry out the observation
340 experiment at the same time as the MMCR. To further analyze the changes in the height distribution of cloud
341 boundaries in Xi'an in 2021, we plan to employ MMCR data to replace the data of periods when the lidar is not
342 running. Accordingly, it is necessary to analyze the correlation of the cloud bottom boundary height detected by the



343 two devices. We randomly select 80 hours of data in the joint observation period (avoid the rainfall period) and
 344 calculate the cloud boundary detection results of the lidar and MMCR according to the data processing methods in
 345 sections 3.1 and 3.2. As shown in Fig. 16, when the quality control of the MMCR is performed, the correlation
 346 between the detected cloud boundary and the lidar detection result is increased from 0.627 (in Fig. 16a)) to 0.803
 347 (in Fig. 16b)). Moreover, under the premise that the difference in cloud boundaries caused by the different detection
 348 principles and detection algorithms of the two devices cannot be avoided, we can use the cloud boundary data
 349 detected by MMCR to replace the missing lidar data.



350
 351 Fig. 16 Correlation between lidar and MMCR cloud bottom boundary. a) without quality control; b) with quality control)

352 From the above three cloud observation cases, it can be seen that MMCR has more advantages than lidar in
 353 detecting cloud top boundaries. Therefore, when calculating the cloud boundary height distribution characteristics
 354 over Xi'an in 2021, we only count the cloud top boundary height detected by MMCR and take it as the actual cloud
 355 top boundary. The statistical rules shown in Table 3 are established for the statistics of cloud bottom boundary
 356 information. The experimental data of 302 days (65 days in spring (January-March), 84 days in summer
 357 (April-June), 65 days in autumn (July-September) and 88 days in winter (October-December) observed in 2021 are
 358 classified and sorted out to ease the statistics and analysis of the variation characteristics of cloud boundary height.

359 Table 3 Statistical rules of cloud bottom boundary information

Detection equipment	Observation	Detection conditions	Record cloud bottom boundary
Both the lidar and MMCR	Case 1	Thin cloud: the lidar detects bottom; MMCR did not detect the cloud bottom	Results of the lidar
	Case 2	Drizzle: the lidar detects bottom; bottom of MMCR is blurred	Results of the lidar
	Case 3	Both the lidar and MMCR detect cloud bottom	Record the lower value of the cloud base boundary
MMCR	Case 4	MMCR detected cloud bottom	Results of MMCR
	Case 5	Drizzle: bottom of MMCR is blurred	No results are recorded

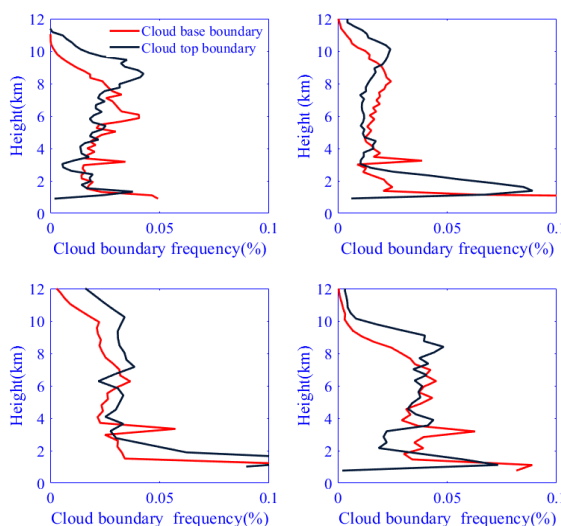
360



361 This study defines "cloud occurrence frequency" as the ratio of cloud occurrence times to total detection times
362 during the analyzed period. It is observed that the total sample size is N , and the sample size of cloud boundaries
363 appearing at different height levels (altitude range from 1.5 km to 12 km is divided into 50 levels) is n_i . The
364 seasonal distribution characteristics of the cloud boundary height are calculated according to Eq. (7),

$$365 \quad y_{cloud} = \frac{n_i}{N} (n_i \in N, i = 1 \dots 50) \quad (7)$$

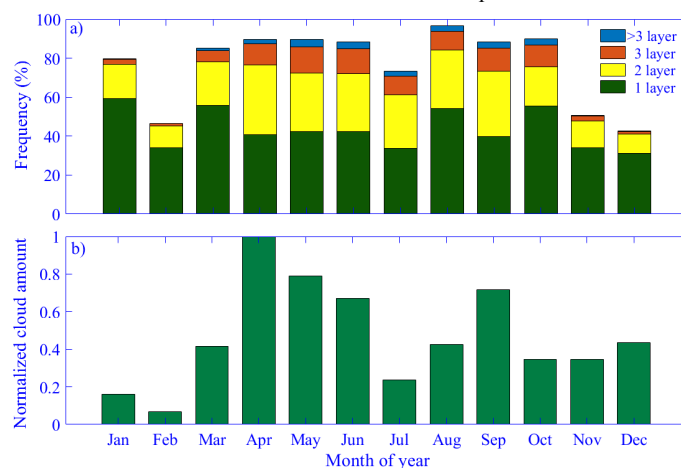
366 Fig. 17 shows that the cloud top boundary occurrence frequency in spring and summer presents a bimodal
367 distribution. In spring, the height of the first peak lies approximately 1.5 ~ 1.9 km, and the second peak is 7.8 ~ 8
368 km. The heights of the first and second peaks are approximately 1.5 ~ 3 km and 8 ~ 12 km, respectively, in summer.
369 In autumn and winter, the frequency of cloud top boundary heights above 2 km is almost in the range of 0.3 to 0.4.
370 For the vertical distribution characteristics of the cloud bottom boundary, there is a triple-mode pattern in four
371 seasons. The frequency distribution characteristics of the cloud bottom boundary height in spring and summer are
372 relatively similar. The first most obvious narrow peak < 1.5 km is the frequency change caused by boundary layer
373 clouds, the second narrow peak is located at 3 ~ 4 km, and the third peaks in spring and summer are located at 6 ~ 8
374 km and 7 ~ 9 km, respectively. From the distribution characteristics of the cloud bottom boundary in summer and
375 spring, it can be guessed that convective and cirrus clouds may be dominant in these two seasons. The frequency
376 distribution of clouds above 8 km in autumn is the largest in the four seasons, and we can speculate that stratus
377 clouds and cumulus clouds are mainly in this season. In winter, the height range of clouds is narrow, and the
378 numerical range is wide, which may be mainly stratiform clouds. This is consistent with the analysis results of
379 Zhao et al. (Zhao et al., 2014) at the SGP site and Hailing Xie (Xie et al., 2017) at the SACOL site. Although there
380 are some differences in the cloud boundary frequency distribution at some heights, the overall change trend is
381 roughly the same.



382
383 Fig. 17 Frequency distribution of cloud boundaries during (a) spring, (b) summer, (c) autumn and (d) winter from January to December
384 2021 at Xi'an Jinghe National Meteorological Station

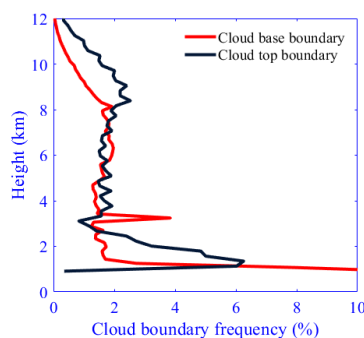


385 Fig. 18a) shows the monthly variation frequency distribution of clouds. The months with the largest (96%) and
 386 smallest (42%) cloud occurrence frequencies are August and December, respectively. Almost more than 34% of the
 387 clouds appear in the form of single layer clouds every month. Compared with January, February, November and
 388 December, the frequencies of double-layer clouds, triple-layer clouds and more clouds in other months are higher.
 389 It is also possible that there are some thin clouds and broken clouds in the upper layer, which are summarized as
 390 multilayer clouds by the algorithm. As shown in Fig. 18b), the normalized monthly distribution of cloud amount
 391 shows that the minimum cloud amount is 0.65 in spring and the maximum is 2.46 in summer, indicating that warm
 392 atmospheric conditions are more conducive to the formation and development of clouds.



393 Fig. 18 The monthly variation in cloud frequency distribution and cloud amount in 2021 a) monthly variation in the frequency of the
 394 number of cloud layers. b) monthly variation in cloud amount
 395

396 Fig. 19 shows the frequency change characteristics of the cloud boundary vertical height distribution in 2021, in
 397 which the frequency of the cloud bottom boundary below the vertical height of 1.5 km is greater than 10%, the
 398 frequency within the height range of 3.06 km and 3.6 km is approximately 3.24%, and the frequency above 8 km is
 399 less than 2%. The frequency of the cloud top boundary at vertical heights has a bimodal distribution; the first
 400 narrow peak is located at 1.5~3.1 km, and the second peak lies at approximately 7.5~10.5 km. Combined with the
 401 changing characteristics of cloud layers, it can be seen that during the observation process in Xi'an in 2021, the
 402 frequency of stratiform clouds below 3.5 km is the largest, and the frequency of high-level ice clouds or cirrus
 403 clouds above 8 km is small.



404 Fig. 19 Frequency distribution of cloud boundaries at vertical heights at Xi'an Jinghe National Meteorological Station in 2021
 405



406 **5 Conclusions**

407 Based on the detection principle of lidar and MMCR, this study realizes the accurate recognition of cloud signals
408 from aerosols and background noise signals by enhancing the lidar echo signal and combining its SNR change. The
409 SNR_{\min} and spectral point continuous threshold are used to identify the cloud signal in the power spectrum of the
410 MMCR, and data quality control technology is implemented for the echo reflectivity factor, which eliminates the
411 interference of nonmeteorological signals on the cloud signal and improves the accuracy of cloud boundary
412 detection.

413 The case analysis results of the joint lidar and MMCR observations show that the two devices have their own
414 advantages in detecting cloud boundaries. 1) For the development of deep clouds, the lidar beam will be seriously
415 attenuated and cannot penetrate the clouds, while the MMCR can penetrate more clouds and obtain the real cloud
416 top boundary. 2) In detecting low-level clouds, the echo reflectivity of MMCR is easy for ground-based clutter
417 interference, and the echo signals observed by lidar can help eliminate clutter to obtain accurate cloud bottom
418 boundaries. 3) When precipitation occurs (except for heavy precipitation), it is difficult to distinguish the cloud
419 bottom height from the echo reflectivity factor of the MMCR, while lidar can reverse the effective cloud bottom
420 boundary to a certain extent. 4) For thin clouds, lidar can obtain more complete information than MMCR.
421 Therefore, when employing the lidar and MMCR to jointly observe the cloud boundary, their respective strengths
422 can be exerted, and their shortcomings can be compensated for each other, making the detection of cloud boundary
423 height more detailed and accurate.

424 Based on the statistical analysis of the changes and distribution of cloud boundaries in Xi'an in 2021, it can be
425 seen that more than 34% of the clouds appear in the form of a single layer of every month. The cloud amount is the
426 lowest in spring and the highest in summer. The seasonal variation in cloud boundary height shows that the
427 distribution characteristics of cloud boundaries in spring and summer are similar, and the frequency of high-level
428 clouds in the range of 8 ~ 10 km is greater than that in the other two seasons. The stratiform clouds appearing
429 below 3.5 km in autumn have the highest frequency, and the high-level ice clouds or cirrus clouds above 8 km in
430 winter are less likely to appear. In this paper, by retrieving the cloud data observed by the lidar and MMCR in 2021,
431 the results of the cloud boundary detected by the two instruments are analyzed and compared to determine the
432 advantages and limitations of the lidar and MMCR in cloud boundary detection, which can provide more
433 information for understanding and studying aerosol-cloud interactions, climate change and forecasting numerical
434 models in Xi'an.

435 **Data availability**

436 Data and code related to this article are available upon request to the corresponding author.

437 **Author contributions**

438 In this paper, Yun Yuan proposed new methods and theories, processed and analyzed the data in the paper, and
439 wrote the manuscript. Huige Di made many revisions and guidance to the manuscript, and put forward many
440 feasible suggestions. Tao Yang collected the experimental data and maintained the equipment. Yuanyuan Liu
441 sorted and classified the original experimental data. Qimeng Li calibrated the experimental system. Qing Yan
442 coordinated the placement and calibration of experimental equipment. Dengxin Hua led the project, instrument
443 development, experimental design and data analysis.



444 **Competing interests**

445 The authors declare that they have no conflicts of interest related to this work.

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