



1 **Optimizing a twin-chamber system for direct ozone production rate** 2 **measurement**

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33 **Abstract.** High Ozone Production Rate (OPR) leads to O₃ pollution episodes and adverse human health outcomes.
34 Discrepancies between OPR observation (Obs-OPR) and OPR modeling (Mod-OPR) as calculated from observed and modeled
35 peroxy radical and nitrogen oxides reminds of a yet-perfect understanding of O₃ photochemistry. Direct measurement of OPR
36 (Mea-OPR) by a twin-chamber system emerges with the optimization required for suppressing the wall effect. Herein, we
37 minimized the chamber surface area to volume ratio (S/V) to 9.8 m⁻¹ and the dark uptake coefficient of O₃ to the order of 10⁻⁹.
38 Condition experiments further revealed a photo-enhanced O₃ uptake and recommended an essential correction. We finally
39 characterized a measurement uncertainty of ±27% and a detection limit of 2.8 ppbv h⁻¹ (3SD), which suggests that Mea-OPR
40 is sensitive enough to measure OPR in urban or suburban environments. Application of this system in urban Beijing during
41 the Winter Olympic Games recorded a noontime OPR of 7.4 (±3.8, 1SD) ppbv h⁻¹, which indicates fairly active O₃
42 photochemistry despite the pollution control policy implemented. Mea-OPR *versus* $j(\text{O}^1\text{D})$ slope of 6.1×10^5 ppbv h⁻¹ s⁻¹
43 confirmed fairly active O₃ photochemistry, which was assisted by a high abundance of VOCs and NO_x, atypically high Mea-
44 OPR even under high-NO_x conditions, but mediated by relatively weak ultraviolet (UV) radiation.

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46 **Short summary.** We reported an optimized system (Mea-OPR) for direct measurement of ozone production rate, which
47 showed a precise, sensitive and reliable measurement of OPR for at least urban and suburban atmosphere, and active O₃
48 photochemical production in winter Beijing. Herein, the Mea-OPR system also shows its potential in exploring the fundamental
49 O₃ photochemistry, i.e., surprisingly high ozone production even under high-NO_x conditions.

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63 1 Introduction

64 Tropospheric ozone (O_3) is a hazardous air pollutant and a key product of photochemical smog (Prinn, 2003; Sillman, 2003).
65 The growing abundance of O_3 due to primarily photochemical production has long been associated with the human health risk
66 (Ho et al., 2007), plant growth issues (Ashmore, 2005), and climate change (Forster et al., 2007). Understanding why the O_3
67 level continues increasing in some regions, despite rigorous emissions control policy (Tarasick et al., 2019), is vital to reverse
68 the O_3 increasing trend.

69 In favorable meteorological scenarios for high ozone production rate such as no or light wind under strong solar radiation,
70 O_3 pollution episodes have been repeatedly recorded. The mixing ratio of O_3 (8-hour maximum average) could reach over 100
71 ppbv, doubling or tripling its background values. Photochemical production reactions of O_3 mainly involve volatile organic
72 compounds (VOCs), nitrogen oxides ($NO_x = NO + NO_2$) and solar radiation (Finlayson-Pitts and Jr. Pittes, 1999; Seinfeld and
73 Pandis, 2006). NO_2 photolysis produces NO and O_3 while O_3 reacts with NO to recycle NO_2 . These rapidly cycling processes
74 (within minutes) are defined as the Leighton cycle. Under the photostationary-state (PSS) assumption for the Leighton cycle,
75 neither O_3 production nor loss is occurring. Production rate of O_3 , referred to as $P(O_3)$, increases as peroxy radicals ($HO_2 + RO_2$)
76 efficiently oxidize NO to produce extra NO_2 (Eq. 1). Additionally, photochemical consumption of O_3 , referred to as $L(O_3)$ (Eq.
77 2), represents a considerable fraction of a comprehensive chemical budget of O_3 to offset O_3 production:

$$78 P(O_3) = k_{HO_2+NO}[HO_2][NO] + \sum_i (\alpha_i k_{RO_2+NO}[RO_{2,i}][NO]) \quad (1)$$

$$79 L(O_3) = f_{O(1D)+H_2O} O_3 + k_{OH+O_3}[OH][O_3] + k_{HO_2+O_3}[HO_2][O_3] + \sum_i (k_{O_3+Alkene_i}[O_3][Alkene_i]) + \\ 80 k_{OH+NO_2}[OH][NO_2] + L(O_3+halogens) \quad (2)$$

$$81 OPR = P(O_3) - L(O_3) \quad (3)$$

82 where k_{RO_2+NO} defines the rate coefficient of the reaction between a specific peroxy radical and NO. Alkyl nitrate formation
83 from molecule isomerization of [R-O-O-NO] competes with O_3 production. The yield of O_3 production is defined as α_i for a
84 specific RO_2 as it reacts with NO, while corresponding yield of alkyl nitrate is defined as $1 - \alpha_i$. $f_{O(1D)+H_2O}$ is defined as the
85 probability of an oxygen atom, a photo-fragment of O_3 photolysis, reacting with water vapor to produce two OH radicals. OPR
86 represents the effective or net production rate of O_3 .

87 NO_x and VOCs regulate the photochemical cycling of radicals and therefore OPR. Under low- NO_x conditions, bimolecular
88 reactions among peroxy radicals buffer the accumulation and chain propagation of peroxy radicals. Increasing abundance of
89 NO_x thus favors chain propagation of peroxy radicals, outcompetes bimolecular reactions among peroxy radicals and favors
90 photochemical production of O_3 . Under high- NO_x conditions, chain termination reaction of NO_2 with OH radical, one major
91 route of O_3 photochemical consumption, competes with the chain propagation of peroxy radicals and photochemical production
92 of O_3 . This scenario analysis reveals the nonlinear relationship between OPR and its precursor of NO_x (Cazorla et al., 2012;
93 Guo et al., 2021). The role of VOCs in O_3 production and the common sources of both VOCs and NO_x in some areas further



94 complicate the OPR-NO_x relationship and its temporal-spatial variability (Schroeder et al., 2017). High time resolution
95 characterization of OPR in the specific atmosphere of concern will be essentially helpful to characterize OPR-NO_x relationship
96 and its environmental variability.

97 In view of aforementioned expressions, Obs-OPR or Mod-OPR can be obtained from the measured or model-simulated HO₂
98 and RO₂, given mixing ratio of NO is usually measured in field campaigns (Ren et al., 2003, 2013; Green et al., 2006; Kanaya,
99 2002; Griffith et al., 2015; Tan et al., 2019b). A NO_x-dependent representation of Mod-OPR was now summarized with
100 consistent and considerable underestimation in high-NO_x areas (Ren et al., 2013; Tan et al., 2019b, 2017; Whalley et al., 2018,
101 2021; Brune et al., 2016) while slightly underestimation or even possible overestimation in low-NO_x areas (Ren et al., 2003;
102 Whalley et al., 2018). For instance, Ren et al. developed an observation-prescribed box model and outputted a underestimated
103 Mod-OPR by up to one order of magnitude compared with Obs-OPR when the median NO mixing ratio reached 17 ppbv (Ren
104 et al., 2013), while overestimated Mod-OPR by 1.4–1.7 times when NO was below 1.0 ppbv was also reported (Ren et al.,
105 2003). Evidently, such observation-model discrepancy on OPR decreased our confidence for understanding both the O₃ budget
106 and policy efficiency of regional O₃ pollution control.

107 The idea of direct measurement of OPR, Mea-OPR, which dated as far back as 1971, was first proposed by Jeffries (1971).
108 The principle is based on differential O₃ signal when ambient air continuously flows into two identical and horizontally
109 oriented chambers, one of which (called reaction chamber) is enough transparent for UV and visible radiation to mimic ambient
110 photochemistry and the other one (named reference chamber) is covered or coated by UV-blocking film to filter UV to suppress
111 the O₃ photochemical production (Cazorla and Brune, 2010; Baier et al., 2015; Sadanaga et al., 2017; Sklaveniti et al., 2018).
112 As simultaneous measurements of NO₂ and O₃ (= O_x) in the reaction chamber and reference chamber can cancel out any ozone
113 difference resulting from the rapid interconversion between NO₂ and O₃ in two chambers, a differential O_x signal rather than
114 O₃ was further adopted. We refer to the twin-chamber method for direct measurement of OPR or the measurement result by
115 such method as Mea-OPR in the context.

116 Accurate and precise Mea-OPR relies on several assumptions implied from the twin-chamber measurement scheme. The
117 most important one is that photochemistry in the reaction chamber should be able to mimic the ambient condition while O₃
118 photochemistry in the reference chamber is totally suppressed. Several versions of Mea-OPR, or named MOPS in literature,
119 have been managed to mimic photochemistry in both chambers. Teflon film and quartz glass are usually adopted for the
120 chamber body to guarantee UV transmittance in the reaction chamber (Cazorla and Brune, 2010; Sadanaga et al., 2017;
121 Sklaveniti et al., 2018). Light transmittance achieves easily above 90% for either Teflon (FEP Teflon film, 0.05 mm thick) or
122 quartz tube, providing ideal UV conditions in the reaction chamber. UV blocking film is also effective in creating a “dark”
123 condition for O₃ photochemistry.

124 Another precondition for Mea-OPR is that measurement of O_x could precisely represent the differential O_x signal, ΔO_x,
125 between chambers. Measurement uncertainties in O_x might be enlarged as transferred to be the measurement uncertainties in
126 ΔO_x, thus further amplifying the measurement uncertainties of Mea-OPR in a unit of ppbv h⁻¹ because the gas residence time
127 in Mea-OPR chambers is typically much shorter than 1 hour. Both insensitivity and the drift in measurement baseline of O₃



128 and NO₂ instruments contribute to the measurement uncertainties of ΔO_x (Cazorla and Brune, 2010; Baier et al., 2015). Such
129 stringent requirement in precise ΔO_x measurements requires advanced and tested measurement techniques and well trained
130 technicians to perform the measurement, which might limit the widespread deployment of Mea-OPR. Reliable and sensitive
131 measurement technique for O₃ and NO₂ might be a savior in this situation.

132 Last but not least, the chamber wall effect must be suppressed and quantified. The chamber wall effect might change the
133 abundance of O_x, which presents false signal of ΔO_x that is difficult to be decoupled from photochemical ΔO_x. The chamber
134 wall effect might also change the abundance of photochemical intermediates, leading to unwanted perturbation on O₃
135 photochemistry in the reaction chamber. Sklaveniti et al. (2018) estimated that O₃ uptake loss in their version of Mea-OPR
136 would lead to false Mea-OPR signal of ~20 ppbv h⁻¹ assuming ambient O₃ to be 50 ppbv. It was also found that HONO
137 production in uptake experiment of NO₂ on quartz chamber and Teflon chamber can reach up to tens of ppbv h⁻¹ under
138 irradiated conditions, which would lead to an overestimation of Mea-OPR by approximately 27% on average (Sklaveniti et
139 al., 2018) or Mea-OPR error of around 10 ppbv h⁻¹ (Baier et al., 2015). Effective design to suppress the wall effect is, therefore,
140 the key to ensure the data quality of Mea-OPR. Previous designs are devoted to keeping the air in plug flow motion and a
141 small-diameter (14.0–17.8 cm) chamber is usually chosen for it is easier to manipulate (Cazorla and Brune, 2010; Baier et al.,
142 2015; Sadanaga et al., 2017; Sklaveniti et al., 2018). However, a small diameter works against the wall effect suppression. The
143 S/V ratio was decreased by 20% by employing a larger chamber volume (26.9 L) in the second version of Mea-OPR (Baier et
144 al., 2015), relative to the first generation (11.3 L) (Cazorla and Brune, 2010). Referring to the uptake coefficient calculation
145 formula (Eq. 4), a 20% decrease of S/V will result in a 20% reduction in false Mea-OPR signal due to uptake loss of O₃ for a
146 given uptake coefficient and O₃ abundance. Previous research has also confirmed less O_x loss on Teflon wall surface (Sadanaga
147 et al., 2017) than on quartz surface (Sklaveniti et al., 2018). Sadanaga et al. (2017) first managed to coat the inner wall surface
148 of the quartz chamber with transparent Teflon and effectively reduced the wall effect.

$$149 \gamma_{O_3} = \frac{4 \times \Delta O_{3, \text{uptake}}}{O_{3, \text{amb}} \times \omega_{O_3} \times \tau \times S/V} \quad (4)$$

150 To date, these effective designs have not yet been integrated and evaluated in a state-of-the-art version of Mea-OPR. In this
151 article, we will present our construction of a state-of-the-art Mea-OPR and condition experiments to characterize this Mea-
152 OPR. Later, the employment of Mea-OPR system in Beijing, a megacity in China, provides further validation of it.

153 **2 Experimental section**

154 **2.1 Construction of state-of-the-art Mea-OPR**

155 The schematic of our Mea-OPR is shown in Fig. S1. The quartz chamber inner wall was coated with the transparent Teflon.
156 It was chosen over the Teflon film chamber for Teflon film chamber showed its weakness in working in high wind velocity
157 conditions. The quartz chamber and the Teflon coating together presented an ideal transmittance (> 88%) for both UV and



158 visible light (Fig. S2). The difference in $j(\text{O}^1\text{D})$ between the ambient and the reaction chamber calculated by measured solar
159 flux (Metcon CCD-Spectrograph) in Beijing winter was within 4% (Table S1). UV filter was adopted for the reference chamber
160 (transmittance for UV was 0, Fig. S3) and enabled small temperature and RH differences in both chambers (Fig. S4).

161 Parallel and accurate measurements of ΔO_x were supported by two identical sets of O_3 analyzer (Thermo Scientific, Model
162 49i, LOD: 1.0 ppbv) and NO_2 analyzer (Los Gatos Research, Inc., Model 911-0009, LOD: 50 pptv). To calculate the wall loss
163 of O_x in both chambers in real time, another set of O_3 analyzer (Thermo Scientific, Model 49i, LOD: 1.0 ppbv) and homemade
164 iBBCES- NO_2 instrument (AIOFM, LOD: 168 pptv; 30 min background shifted: 100 pptv) (Fang et al., 2017) were sampling
165 the ambient air. The system was running with a duty cycle in a 23 h plus 1 h, which included 1 h instrument alignment as all
166 instruments shifted to the ambient sampling line at midnight. The purpose of the instrument alignment was to check the
167 working status of the measurements on a daily base, rather than data correction. ΔO_x up to 1 ppbv in a 1 h alignment reminded
168 of instrument calibration and maintenance, which were otherwise conducted on a weekly basis. Under the working protocol,
169 the system performed stably in a 1-month measurement duty from February 5 to March 12, 2022. Multiple calibrations
170 suggested the instrument response shifted within $\pm 0.4\%$ and $\pm 1.9\%$ for O_3 and NO_2 , respectively, in this period (Table S2).
171 The instrument alignment experiment suggested that nighttime ΔO_x was 0.07 (± 0.26) ppbv, within the instrument detection
172 limits (Fig. S5).

173 Our Mea-OPR development put emphasis on wall effect management. The diameter of our Mea-OPR was 41 cm, which
174 produced a S/V ratio that was 1.8 fold less than the lowest value ever published in the literature (Cazorla and Brune, 2010;
175 Baier et al., 2015). This reduced the O_3 wall effect by 1.8 folds even though the same uptake coefficient of O_3 was taken into
176 consideration (Eq. 4), and so did for wall effect of other species, such as NO_2 and HONO. Larger diameter of our Mea-OPR
177 also gave the flexibility of the working flow rate, which spanned from 5 to 20 L min^{-1} while enabling multiple sampling
178 instruments equipped and a residence time of up to 30 min in a plug flow mode (Fig. S1). Another key design to suppress the
179 wall effect was the transparent Teflon coating, which was essential for accurate Mea-OPR by reducing the uptake coefficient
180 of O_3 from 10^{-8} on quartz wall (Sklaveniti et al., 2018) to 10^{-9} on Teflon coating wall under dark conditions.

181 Mea-OPR can be calculated in Eq. (5), wherein ΔNO_2 and ΔO_3 are the difference values of NO_2 and O_3 between both
182 chambers, respectively. τ is the mean gas residence time in chambers (Fig. S6). γ is the uptake coefficient of O_3 in the
183 chamber. ω represents the mean molecular velocity of O_3 in m s^{-1} . D is the diameter of chambers in m. $\text{O}_{3, \text{amb}}$ represents
184 the ambient O_3 concentration in ppbv. φ_{trans} is the ratio of in-chamber $j(\text{O}^1\text{D})$ to ambient $j(\text{O}^1\text{D})$ as determined by the UV
185 transmittance of Mea-OPR system. $\varphi_{\Delta\text{HONO or } \Delta\text{NO}_x}$ is the ratio of OPR in the reaction chamber to that in the ambient owing to
186 the presence of ΔHONO or ΔNO_x between the reaction chamber and the ambient.

$$187 \text{ Mea-OPR} = \left(\frac{\Delta\text{NO}_2 + \Delta\text{O}_3}{\tau} + \frac{(\gamma_{\text{Rea}} \cdot \omega_{\text{Rea}} - \gamma_{\text{Ref}} \cdot \omega_{\text{Ref}}) \cdot \text{O}_{3, \text{amb}}}{D} \right) \cdot \frac{1}{\varphi_{\text{trans}} \cdot \varphi_{\Delta\text{HONO or } \Delta\text{NO}_x}} \quad (5)$$



188 2.2 Condition experiments to characterize wall effect

189 As Mea-OPR was sampling the ambient air, O₃ wall loss and photochemical production were occurring simultaneously and,
190 thus, cannot be decoupled from each other. A condition experiment of zero-NO_x-and-zero-VOCs (referred to as zero-NO_x-
191 and-high-O₃ experiment) was designed and, thus, zero OPR was assumed. High O₃ control (also referred to as ambient O₃ for
192 simplicity) in zero-NO_x-and-high-O₃ experiment favored measurement of O₃ wall losses directly in both chambers on 6 March,
193 2022. In addition, HONO production from NO₂ uptake (Eq. 6) on the Teflon film and quartz surface have been proposed by
194 Baier et al. (2015) and Sklaveniti et al. (2018) as a potential perturbation on photochemical production of O₃ in the reaction
195 chamber. To obtain differential NO_x and HONO signals in chambers relative to the ambient, additional measurements of
196 HONO in the chambers and in the ambient were equipped in a 1-week HONO production experiment as Mea-OPR was
197 deployed in 11–18 February, 2022. The above condition experiments to characterize wall effect were described in detail in the
198 Supplement (S2.6).



200 2.3 Field application of Mea-OPR

201 Our Mea-OPR system was deployed on the top floor of an academic building on the campus of Peking University (39°59'23"
202 N, 116°18'25" E). The observatory as a typical urban and polluted site in Beijing City, China, was impacted by considerable
203 fresh, anthropogenic emissions in the surroundings, such as the 4th ring-road traffic emission. More details about this site were
204 also described elsewhere (Guo et al., 2010; Tang et al., 2018). Notably, the measurement period coincided with the Beijing
205 2022 Winter Olympic Games from 4 to 20 February, when the Municipal Government of Beijing implemented a package of
206 factory and transportation controls to improve air quality. Owing to the aggressive control measures and favorable
207 meteorological conditions, no haze events and substantial reductions in gas pollutant concentrations were observed (Guo et
208 al., 2023; Liu et al., 2022). Therefore, the air pollution conditions during this period can be regarded as relatively clean
209 compared to those polluted episodes in the same period of other years.

210 3 Results and discussion

211 3.1 O₃ uptake evaluation

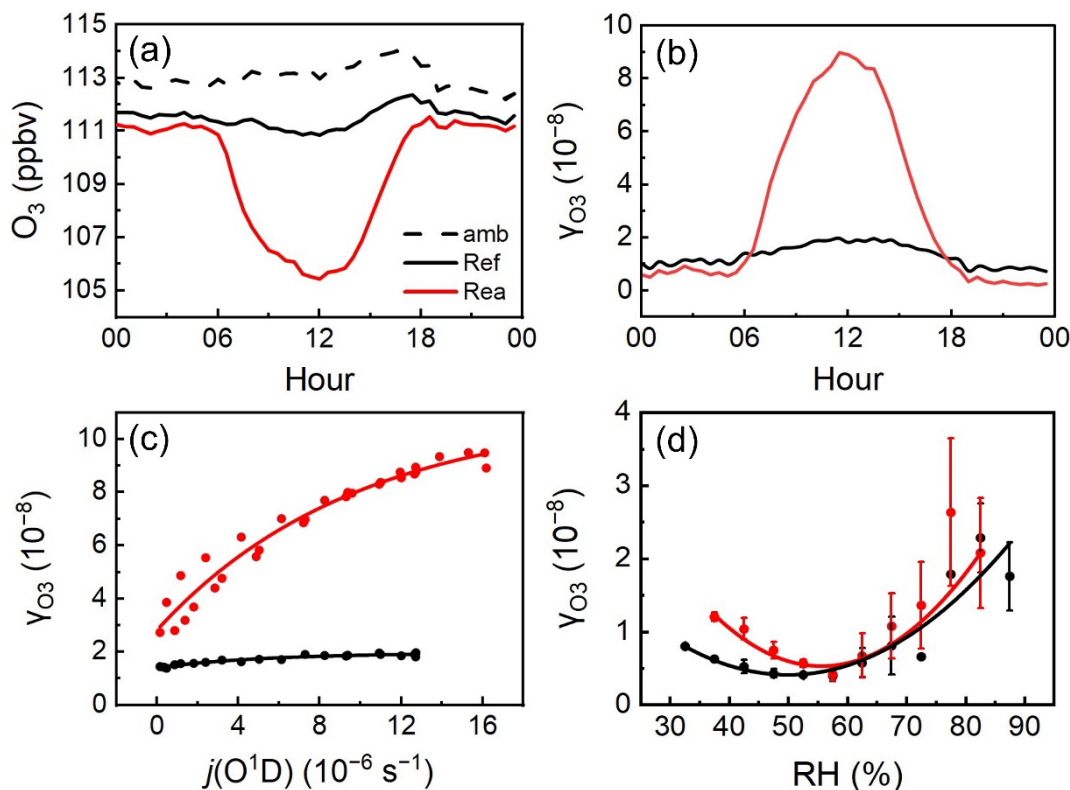
212 A zero-NO_x-and-high-O₃ experiment was conducted for directly measuring the O₃ wall loss in both chambers with little
213 perturbation from O₃ photochemical production in chambers. As shown in Fig. 1a, evident O₃ wall loss ($\Delta\text{O}_{3, \text{raw}} = \text{O}_3$ in the
214 ambient – O₃ in the chamber) were observed and higher O₃ wall loss was observed for the reaction chamber ($\Delta\text{O}_{3, \text{raw}} = 7.7$
215 ppbv) relative to the reference chamber ($\Delta\text{O}_{3, \text{raw}} = 2.3$ ppbv) at noon. The ambient O₃ control was measured at approximately
216 113.0 ppbv just before entering chambers. NO_x concentration was measured around 0.03–1.05 ppbv to double-check the low
217 level of NO_x in the control experiment. A slight increase in NO_x from the morning to the noon was accompanied with increasing



218 $j(\text{O}^1\text{D})$, which might be attributed to the previously-identified unknown source of HONO and NO_x in Teflon chamber (Zhou
219 et al., 2003). The unknown chamber source of HONO was also confirmed later during the 1-week HONO production
220 experiment (see below). The MCM model was conducted to calculate O_3 production in chambers. NO_x and CO were both
221 major prescribed chemical parameters. At noon, the corresponding O_3 production in the reaction chamber, $\Delta\text{O}_3, \text{photochemistry}$,
222 was found to be 2.7 ppbv in the residence time of 20 min (Fig. S7), which comprises 35% of the $\Delta\text{O}_3, \text{raw}$ for the reaction
223 chamber. O_3 production was negligible in the reference chamber at noon because stray light in the reference chamber was too
224 weak to be meaningful for O_3 photochemistry.

225 After eliminating the photochemical contribution, $\Delta\text{O}_3, \text{uptake}$ ($= \Delta\text{O}_3, \text{raw} - \Delta\text{O}_3, \text{photochemistry}$) was obtained and the uptake
226 coefficient of O_3 on chambers could be calculated according to Eq. (4). The uptake coefficient from a typical value of $7.11 \times$
227 10^{-8} on quartz wall (Sklaveniti et al., 2018) (Table S3) was successfully reduced to the present average value of 8.12×10^{-9}
228 on the transparent Teflon coating under dark conditions. Under typical working conditions of Mea-OPR, O_3 uptake loss
229 contributes to a false Mea-OPR signal of 20.3 ppbv h^{-1} at uptake coefficient of 7.11×10^{-8} and S/V ratio of 18 m^{-1} (the least
230 in the literature), relative to a false Mea-OPR signal of 1.29 ppbv h^{-1} at uptake coefficient of 8.12×10^{-9} and S/V ratio of 9.76
231 m^{-1} in our Mea-OPR, assuming ambient O_3 concentration of 50 ppbv. This calculation suggests the essential success in wall
232 effect suppression by our strategy, i.e., to minimize chamber S/V ratio and uptake coefficient of O_3 on the chamber wall, for
233 assuring the data quality of Mea-OPR.

234 During the daytime, UV-dependent γ was observed and inferred a photo-enhanced uptake of O_3 of up to 8.98×10^{-8} at noon,
235 nearly one order of magnitude higher relative to 8.12×10^{-9} at night (Fig. 1b). Low uptake coefficient of 5.17×10^{-9} under
236 dark condition was observed for Teflon surface due to its inert nature (Sadanaga et al., 2017). However, photosensitization
237 reactions of O_3 on organic coating was observed with an uptake coefficient in the range of 10^{-6} – 10^{-5} under near-UV and visible
238 irradiation relative to 10^{-7} – 10^{-6} in the dark (Styler et al., 2009; D'Anna et al., 2009). This suggested that the aerosol particles
239 deposited on the inner wall surface might substantially contribute to the photo-enhanced wall loss of O_3 . A routine water flush
240 cleaning and UV-photochemical-aging cleaning of both chambers were then scheduled after occurrence of heavy pollution
241 episodes ($\text{PM}_{2.5} > 80 \mu\text{g cm}^{-3}$) (Juda-Rezler et al., 2020). RH might modify aerosol phase state and thus was another parameter
242 affecting the uptake coefficient on aerosol particles and apparent wall loss of O_3 . The daytime RH was typically below 50%,
243 approximately the threshold RH of particle phase shift in the city of Beijing (Liu et al., 2016). A relatively stable γ at low RH
244 was also observed. The daytime γ was then expressed as a function of merely $j(\text{O}^1\text{D})$, not considering the complex influence
245 from RH variation below 50% (Fig. 1c). An increasing γ as a function of RH for the nighttime was summarized, with a
246 threshold RH of ca. 60% (Fig. 1d).



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248 **Figure 1: Results of condition experiment designed to measure O₃ wall loss. (a) O₃ concentrations in the ambient (approximately**
 249 **113.0 ppbv) and in both chambers as zero-NO_x-and-high-O₃ air was introduced to both chambers; (b) Calculated half-hour**
 250 **resolution uptake coefficient (γ) of O₃ in both chambers; (c) Fitted exponential relationship between γ and $j(O^1D)$ in daytime, i.e.,**
 251 **$j(O^1D) > 1 \times 10^{-7} \text{ s}^{-1}$ ($R^2 = 0.93$ and 0.96 , respectively); (d) Fitted polynomial relationship between γ and RH during night, i.e., $j(O^1D)$**
 252 **$< 1 \times 10^{-7} \text{ s}^{-1}$ ($R^2 = 0.82$ and 0.82 , respectively). Error bar represents standard deviation of γ in each RH interval of 5%. Fitting**
 253 **method is ordinary least squares estimation.**

254 Sklaveniti et al. (2018) also suspected photo-enhanced O₃ uptake in the reaction chamber as a major measurement
 255 interference for their Mea-OPR system utilizing quartz wall without Teflon coating. However, no quantification and correction
 256 of Mea-OPR interference from the O₃ uptake was suggested. The $j(O^1D)$ -dependent γ for the daytime and RH-dependent γ for
 257 the nighttime as determined in our control experiments provided the first reliable correction for Mea-OPR (Eqs. 7–10). For the
 258 correction, O₃ concentrations measured in the ambient could be used to calculate the $\Delta O_{3, \text{uptake}}$ (Eq. 4). After multiple control
 259 experiments (not shown), we could also assume the uptake coefficient of O₃ being stable between two adjacent control
 260 experiments.

$$261 \quad \gamma_{\text{Ref}} = -5.53 \times 10^{-9} \times \exp\left(-\frac{j(O^1D)}{5.41 \times 10^{-6}}\right) + 1.96 \times 10^{-8} \quad (7)$$

$$262 \quad \gamma_{\text{Rea}} = -8.21 \times 10^{-8} \times \exp\left(-\frac{j(O^1D)}{9.76 \times 10^{-6}}\right) + 1.10 \times 10^{-7} \quad (8)$$

$$263 \quad \gamma_{\text{Ref}} = 1.28 \times 10^{-11} \times \text{RH}^2 - 1.28 \times 10^{-9} \times \text{RH} + 3.60 \times 10^{-8} \quad (9)$$



$$264 \quad \gamma_{\text{Rea}} = 2.37 \times 10^{-11} \times \text{RH}^2 - 2.56 \times 10^{-9} \times \text{RH} + 7.50 \times 10^{-8} \quad (10)$$

265 3.2 In-chamber HONO production evaluation

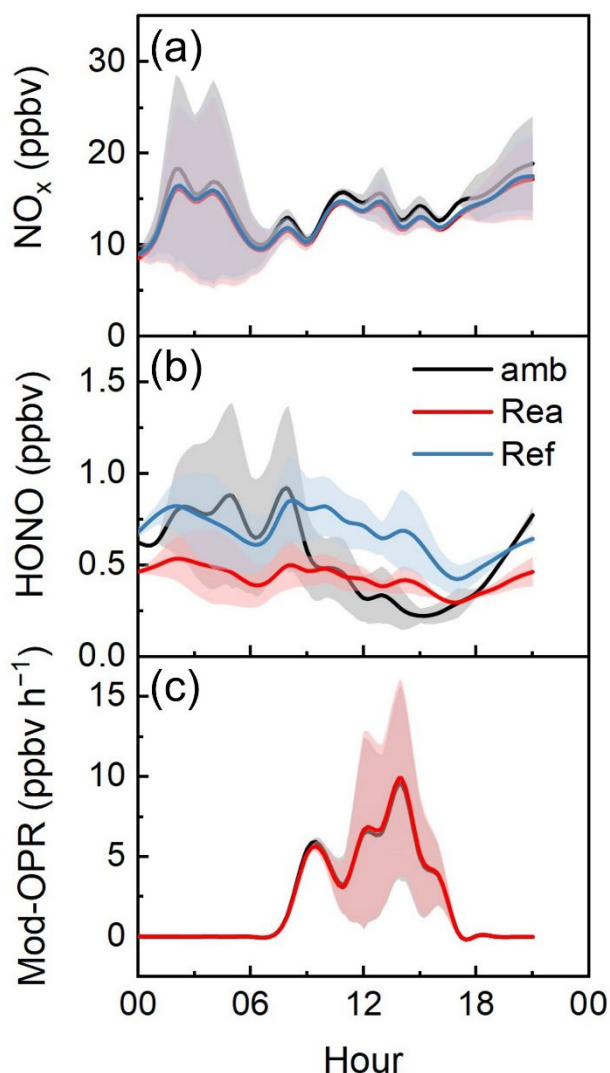
266 In the 1-week HONO production experiment, ΔNO_x (= NO_x in the ambient – NO_x in the chamber) showed no obvious
267 diurnal variation, with an average of 1.01 and 0.77 ppbv in the reaction and reference chamber, respectively (Fig. 2a). Based
268 on Eq. (S3), the average uptake coefficient of NO_2 in the reaction and reference chamber were calculated as 8.66×10^{-8} and
269 6.43×10^{-8} , respectively. Despite of considerable uptake coefficient of NO_2 , compared with that of O_3 , much lower NO_2 level
270 (12.9 ppbv) relative to O_3 inferred negligible wall loss of NO_2 , which was transferred to *ca.* 1.79 ppbv h^{-1} false signal of Mea-
271 OPR at most during the 1-week HONO production experiment. The perturbation on O_3 photochemistry in the reaction chamber
272 by such ΔNO_x was evaluated by MCM model to be negligible due to the insensitive response of O_3 photochemistry to NO_x
273 abundance in the range around 10 ppbv.

274 Uptake loss during the nighttime and production of HONO during the daytime in both chambers were spotted (Fig. 2b).
275 Nighttime loss of HONO on the wall surface, as observed in the lower HONO in both chambers relative to the ambient HONO
276 ($\Delta\text{HONO} = -0.22$ ppbv on average for the reaction chamber), was not suspected even assuming that uptake loss of NO_2 and
277 heterogeneous production of HONO were negligible (Sadanaga et al., 2017). However, Sklaveniti et al. (2018) reported HONO
278 production rate of up to 9 ppbv h^{-1} in controlled ambient concentration of NO_2 under dark condition. HONO uptake on
279 deliquescent aerosol particles at night might account for the HONO loss here (Ren et al., 2020). RH in the reaction chamber
280 scattered at approximately 61% ($\pm 14\%$) and was much higher than the ambient air of 36% ($\pm 14\%$) during the nighttime, which
281 might lead to deliquescence of deposited aerosol particles on the wall surface. As the temperature rose and RH dropped in the
282 early morning, uptake loss of HONO on the wall surface was diminishing. Further decrease in the zenith angle even led to
283 production or release of HONO in both chambers, resulting in a higher HONO concentration in both chambers relative to the
284 ambient air during the daytime ($\Delta\text{HONO} = 0.09$ ppbv on average for the reaction chamber). Either heterogeneous uptake of
285 NO_2 or unknown temperature-related or UV-related chamber source of HONO or releasing of nighttime uptaken HONO might
286 account for the daytime ΔHONO in chambers. Daytime ΔHONO herein appeared to be less than previous reports in laboratory
287 condition (Sklaveniti et al., 2018) and in the ambient of Houston (Baier et al., 2015). Both the inert surface coating and less
288 abundant NO_2 in 1-week HONO production experiment might justify the result.

289 While the small ΔNO_2 and ΔHONO contribute slightly to the measurement uncertainty of ΔO_x , daytime HONO production
290 in the reaction chamber might perturb the O_3 photochemistry therein and challenge the mimic of O_3 photochemistry by the
291 reaction chamber (Baier et al., 2015; Sklaveniti et al., 2018). We then evaluated the O_3 production difference in the reaction
292 chamber relative to the ambient air. Since the radical budget is closely related to O_3 chemistry (Tan et al., 2018a; Ma et al.,
293 2022), the primary RO_x source budget and O_3 production in the reaction chamber and the ambient air were calculated and
294 compared, based on the MCM models prescribed with HONO in the ambient air and the reaction chamber, respectively. When
295 the HONO difference between the reaction chamber and the ambient air was maximum in the daytime, HONO photolysis
296 comprised 22.2% of the total primary RO_x source budget for the ambient air, while constituted 28.2% of the total primary RO_x



297 source budget for the reaction chamber. The overall primary RO_x source budget was at most 8.1% higher in the reaction
298 chamber than in the ambient air, which revealed that ΔHONO considerably perturbed the RO_x source budget owing to the
299 considerable contribution from HONO photolysis. Furthermore, O_3 production enhancement owing to HONO production in
300 the reaction chamber, relative to the ambient air, was calculated as 4.5% at most during the 1-week HONO production
301 experiment (Fig. 2c). It could be seen that although the HONO production in the reaction chamber significantly increased the
302 RO_x source budget and promoted the O_3 production, it might also increase the sink of O_3 accordingly, and in general, it had
303 little effect on the net O_3 production. While in less polluted environments, the perturbation of HONO on O_3 photochemistry
304 was assumed to be less and, therefore, negligible. Currently, NO_2 uptake and HONO production correction are not applied for
305 our Mea-OPR.



306
307 **Figure 2: Diurnal variation of average hourly (a) NO_x and (b) HONO concentrations in ambient air and both chambers. (c)**
308 **Comparison of O_3 production simulation as prescribed with HONO in ambient air and reaction chamber during condition**



309 experiment of HONO production. Shaded areas represent 1SD variation of NO_x, HONO, and OPR for the 1-week HONO production
310 experiment.

311 3.3 Uncertainty evaluation for Mea-OPR

312 The detection limit of 2.8 ppbv h⁻¹ of our Mea-OPR was determined by three times the standard deviation of Mea-OPR
313 during nighttime (21:00–04:00) when OPR was supposed to be near null. Currently, tropospheric OPR in remote and rural
314 areas is believed to be lower than 5 ppbv h⁻¹ (Kanaya, 2002; Xue et al., 2013; Bozem et al., 2017). In polluted regions, OPR
315 can be up to tens of ppbv h⁻¹ (Baier et al., 2015; Kleinman, 2005; Xue et al., 2021; Tan et al., 2021; Whalley et al., 2021; Zhou
316 et al., 2014; Cazorla et al., 2012). Hence, our system is sensitive enough to detect OPR in polluted urban sites, but might not
317 be so in the clean remote sites.

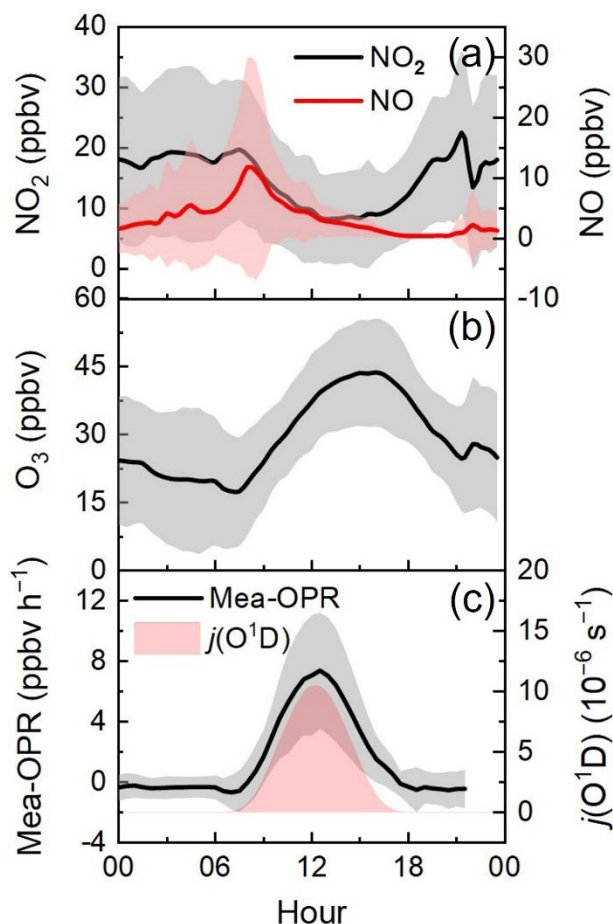
318 The uncertainties of Mea-OPR arise from UV transmittance of the chamber wall, measurements of gas residence time and
319 ΔO_x , and correction from O₃ and NO₂ wall loss, or HONO production in the reaction chamber (Table S4). Assuming a linear
320 relationship among Mea-OPR and $j(O^1D)$, reduction in $j(O^1D)$ in the reaction chamber and stray light in the reference chamber
321 contribute to uncertainties of -4% and -5%, respectively (Table S1). The measurement uncertainty of gas residence time is
322 evaluated from the multiple residence time measurements to be better than $\pm 0.45\%$. Evaluated from 1-hour consistency
323 measurements every night, the measurement uncertainties for ΔNO_2 and ΔO_3 are $\pm 3.8\%$ and $\pm 1.1\%$, respectively. Higher
324 uncertainties were found at low levels of NO_x or O₃, which suggests that continuous improvement in ΔO_x measurement
325 precision will benefit our measurement. The largest uncertainty of Mea-OPR comes from the wall effect of O₃ and in our case
326 the wall effect correction of O₃. The wall effect correction functions of O₃ indicate a higher wall effect of O₃ as $[O_3] \times j(O^1D)$
327 is higher. As a strong O₃ production is also presumed when $[O_3] \times j(O^1D)$ is higher, the wall effect correction of O₃ does not
328 indicate lower measurement quality in strong O₃ production scenario. The uncertainties of the fitted O₃ uptake coefficient in
329 both chambers (Eqs. 7–10) are $\pm 4.4\%$ and $\pm 23\%$, respectively. Wall effect of HONO contributes to an uncertainty of +4.5%,
330 evaluated from the condition experiment. The total uncertainty of our Mea-OPR can be obtained by error propagation as
331 described in detail in the Supplement (S2.7). As shown in Fig. S8, the total uncertainty of the Mea-OPR system decreases with
332 the increase of Mea-OPR. When Mea-OPR is above the detection limit of 2.8 ppbv h⁻¹, the uncertainty of Mea-OPR system is
333 stable and low at an average value of $\pm 27\%$. In addition, when the O₃ uptake coefficient decreases, the uncertainty of Mea-
334 OPR also clearly reduces, which further indicates the importance of reducing O_x uptake in the chambers for the accurate
335 measurement of Mea-OPR.

336 3.4 Field application in urban site of Beijing

337 The mean diurnal profile of ambient $j(O^1D)$, NO, NO₂, O₃ and Mea-OPR during the Beijing 2022 Winter Olympic Games
338 are shown in Fig. 3. NO_x levels ranged from approximately 0.5 to 112.8 ppbv, with an averaged value about 18.2 (± 16.7) ppbv,
339 which was similar to that observed at the suburban site (Huairou Station) of Beijing from January to March, 2016 (Tan et al.,
340 2018b), but much lower than our measurement of 32.0 (± 22.4) ppbv in the winter of 2021 (not shown). Specific pollution



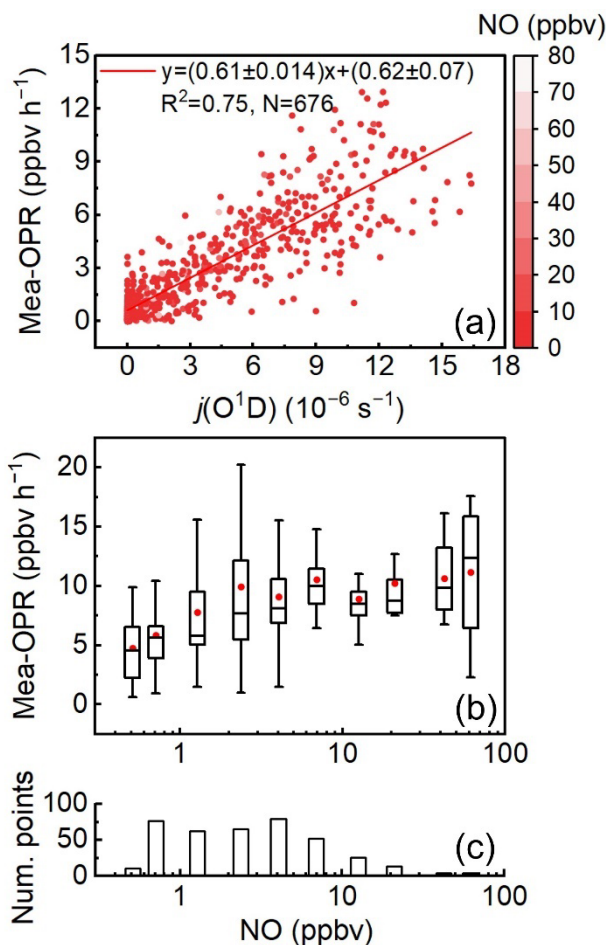
341 control measures for the Beijing 2022 Winter Olympic Games reconcile such a discrepancy. The maximum daily concentration
342 of NO₂ climbed to 26.6 ppbv during the end of our field campaign, which was normal as compared to the winter 2021
343 campaign. O₃ concentration showed afternoon (14:00–16:00) maxima and early morning (6:00–8:00) minima (Fig. 3b), the
344 former of which was a typical feature of photochemical production. The O₃ levels varied from 1.2 to 90.5 ppbv in the daytime,
345 and the mean and median values were approximately 34.4 (±13.9) ppbv and 36.6 ppbv, respectively. In fact, O₃ titration by
346 NO in this urban site was apparent during the morning/afternoon rush hours, but less evident at noon as it was seen from the
347 NO/NO₂ ratio during rush hours (0.20) and noon (0.46), in comparison with the photo-steady-state ratio of 0.14 and 0.44,
348 respectively. Considering the O₃ titration, the urban diurnal O_x concentration was 43.5 (±3.1) ppbv, higher than the regional
349 background of 38.0 ppbv (Xu et al., 2020), suggesting the urban area was still the regional source of O₃ pollution. Noontime
350 $j(\text{O}^1\text{D})$ peaked at approximately $1.0 \times 10^{-5} \text{ s}^{-1}$ on average (Fig. 3c), less than half of the value during the summer (2.6×10^{-5}
351 s^{-1}) (Tan et al., 2019a). Mea-OPR showed similar diurnal variation to $j(\text{O}^1\text{D})$. Mean Mea-OPR peaked at 7.4 ppbv h⁻¹ at around
352 12:30 (Fig. 3c), which was smaller than the O₃ production peak of 20 ppbv h⁻¹ measured using the Obs-OPR method in summer
353 at the PKU urban site (Tan et al., 2019a). The nighttime Mea-OPR approached zero as it was also expected from the low levels
354 of O₃, NO₃, and RO₂ (not shown), accompanied with a nighttime NO of 2.3 ppbv. Therefore, such high Mea-OPR confirmed
355 active O₃ photochemistry and that urban Beijing was still a regional source of O₃ pollution even though the pollution control
356 policy was implemented during the Winter Olympic Games.



357

358 **Figure 3: Mean diel profiles of (a) NO and NO₂, (b) O₃, and (c) Mea-OPR and $j(O^1D)$ at PKU urban site during field campaign.**
359 **Shaded areas represent 1SD variation of measurement parameters.**

360 A linear relationship was established for Mea-OPR *versus* $j(O^1D)$, but an atypical relationship was established for Mea-OPR
361 *versus* NO from the Gaussian-shaped ones (Fig. 4) (Whalley et al., 2021; Cazorla et al., 2012). The linear relationship between
362 Mea-OPR and $j(O^1D)$ justified that our Mea-OPR measurement captured its photochemical pattern well. The fitting slope of
363 the Mea-OPR *versus* $j(O^1D)$ plot was 6.1×10^5 ppbv h⁻¹ s⁻¹. Compared with a study of Xue et al. (2013) in a global background
364 site of Waliguan, the fitting slope in this study was higher, which might suggest the relatively active nature of O₃
365 photochemistry in winter in Beijing. With relatively high NO_x level and assumed VOCs-limited regime for O₃ production in
366 Beijing, O₃ photochemistry was expected to be suppressed under high-NO_x conditions. However, an atypical relationship was
367 observed here, as also being validated in Obs-OPR *versus* NO plot (Whalley et al., 2021; Cazorla et al., 2012). The continuous
368 increase in OPR as NO_x increased even in high-NO_x conditions might have accounted for the active O₃ photochemistry in
369 urban Beijing. Detailed reasons for such atypical relationship are not yet clear, but our Mea-OPR shows its promising capacity
370 to capture both expected and atypical patterns of O₃ photochemistry alongside the change of $j(O^1D)$ and NO.



371

372 **Figure 4: Plot of Mea-OPR against (a) $j(\text{O}^1\text{D})$ and (b) NO for daytime conditions ($j(\text{O}^1\text{D}) > 10^{-7} \text{ s}^{-1}$) during field campaign. The**
373 **interval of NO bin is $\Delta \ln(\text{NO}) = 0.57$ ppbv. Panel (c) shows number of datapoints included in each NO interval.**

374 4 Conclusions

375 Accurate quantification of OPR is an important premise to effective O₃ pollution control strategy and to control the adverse
376 effects of O₃ pollution on human health and climate. Previous studies have shown discrepancies between Obs-OPR and Mod-
377 OPR, which indicates that our understanding of O₃ photochemistry is yet-perfect. Direct measurement of OPR using a twin-
378 chamber system (Mea-OPR) could provide an accurate measurement of OPR, shining light on the emerging conceptual
379 framework of O₃ photochemistry. In this work, we reported an optimized system for direct measurement of ozone production
380 rate, Mea-OPR, and its employment at an urban site of Beijing. Our study optimized the chamber design in several ways, i.e.,
381 considerably increased the chamber volume and, thus, broadened the optional range of flow rate; employment of a large
382 chamber diameter and the Teflon coating on the inner wall of the chamber to suppress the wall effect. We minimized the
383 chamber surface area to volume ratio to 9.8 m⁻¹ and the dark uptake coefficient of O₃ to the order of 10⁻⁹. In addition, photo-



384 enhanced uptake of O₃ on deposited particulate matter was found to be the major error source of Mea-OPR. Condition
385 experiments further revealed a photo-enhanced O₃ uptake and recommended a quantitative correction. We finally characterized
386 a detection limit of 2.8 ppbv h⁻¹ and a measurement uncertainty of ±27%, which suggests that Mea-OPR is sensitive enough
387 to measure OPR in urban or suburban environments. Application of this system in urban Beijing during the Winter Olympic
388 Games recorded a noontime Mea-OPR of 7.4 (±3.8) ppbv h⁻¹ and Mea-OPR *versus* $j(O^1D)$ slope of 6.1×10^5 ppbv h⁻¹ s⁻¹,
389 which indicates fairly active O₃ photochemistry despite the pollution control policy. The fairly active O₃ photochemistry was
390 mainly assisted by a high abundance of VOCs and NO_x, atypically high Mea-OPR even under high-NO_x conditions, but
391 mediated by relatively weak ultraviolet radiation.

392

393 *Data availability.* Contact the corresponding author for data.

394

395 *Supplement.* The following file is available free of charge. Optimizing a twin-chamber system for direct ozone production rate
396 measurement_ SI

397

398 *Author Contributions.* Y.W. and C.Y. built the Mea-OPR system, Y.C. and S.C. conducted the condition experiments,
399 interpreted the data, and wrote the manuscript with revision mainly from Y.W., C.Y., and other authors. Y.W., Y.C., and S.C.
400 contributed equally to this work and should be considered co-first authors. All authors have given approval to the final version
401 of the manuscript.

402

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404

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418 References

- 419 Ashmore, M. R.: Assessing the future global impacts of ozone on vegetation, *Plant. Cell. Environ.*, 28, 949–964,
420 <https://doi.org/10.1111/J.1365-3040.2005.01341.X>, 2005.
- 421 Baier, B. C., Brune, W. H., Lefer, B. L., Miller, D. O., and Martins, D. K.: Direct ozone production rate measurements and
422 their use in assessing ozone source and receptor regions for Houston in 2013, *Atmos. Environ.*, 114, 83–91,
423 <https://doi.org/10/gnxf26>, 2015.
- 424 Bozem, H., Butler, T. M., Lawrence, M. G., Harder, H., Martinez, M., Kubistin, D., Lelieveld, J., and Fischer, H.: Chemical
425 processes related to net ozone tendencies in the free troposphere, *Atmos. Chem. Phys.*, 17, 10565–10582,
426 <https://doi.org/10.5194/acp-17-10565-2017>, 2017.
- 427 Brune, W. H., Baier, B. C., Thomas, J., Ren, X., Cohen, R. C., Pusede, S. E., Browne, E. C., Goldstein, A. H., Gentner, D. R.,
428 Keutsch, F. N., Thornton, J. A., Harrold, S., Lopez-Hilfiker, F. D., and Wennberg, P. O.: Ozone production chemistry in the
429 presence of urban plumes, *Faraday Discuss.*, 189, 169–189, <https://doi.org/10.1039/C5FD00204D>, 2016.
- 430 Cazorla, M. and Brune, W. H.: Measurement of Ozone Production Sensor, *Atmos. Meas. Tech.*, 3, 545–555,
431 <https://doi.org/10.5194/amt-3-545-2010>, 2010.
- 432 Cazorla, M., Brune, W. H., Ren, X., and Lefer, B.: Direct measurement of ozone production rates in Houston in 2009 and
433 comparison with two estimation methods, *Atmos. Chem. Phys.*, 12, 1203–1212, <https://doi.org/10.5194/acp-12-1203-2012>,
434 2012.
- 435 D’Anna, B., Jammoul, A., George, C., Stemmler, K., Fahrni, S., Ammann, M., and Wisthaler, A.: Light-induced ozone
436 depletion by humic acid films and submicron aerosol particles, *J. Geophys. Res.*, 114, D12301,
437 <https://doi.org/10.1029/2008JD011237>, 2009.
- 438 Fang, B., Zhao, W., Xu, X., Zhou, J., Ma, X., Wang, S., Zhang, W., Venables, D. S., and Chen, W.: Portable broadband cavity-
439 enhanced spectrometer utilizing Kalman filtering: application to real-time, in situ monitoring of glyoxal and nitrogen
440 dioxide, *Opt. Express*, 25, 26910, <https://doi.org/10.1364/OE.25.026910>, 2017.
- 441 Finlayson-Pitts, B. J. and Jr. Pittes, J. N.: *Chemistry of the Upper and Lower Atmosphere*, academic press, San Diego,
442 California, 1999.
- 443 Forster, P., Ramaswamy, V., Artaxo, P., Berntsen, T., and Dorland, R. V.: Changes in atmospheric constituents and in radiative
444 forcing, in: *Climate Change 2007: The Physical Science Basis. Contribution of Working Group I to the Fourth Assessment*
445 *Report of the Intergovernmental Panel on Climate Change*, Cambridge University Press, Cambridge, UK and New York,
446 USA, 2007.
- 447 Green, T. J., Reeves, C. E., Fleming, Z. L., Brough, N., Rickard, A. R., Bandy, B. J., Monks, P. S., and Penkett, S. A.: An
448 improved dual channel PERCA instrument for atmospheric measurements of peroxy radicals, *J. Environ. Monit.*, 8, 530,
449 <https://doi.org/10/d95fhs>, 2006.



- 450 Griffith, S. M., Huang, X. H. H., Louie, P. K. K., and Yu, J. Z.: Characterizing the thermodynamic and chemical composition
451 factors controlling PM_{2.5} nitrate: Insights gained from two years of online measurements in Hong Kong, *Atmos. Environ.*,
452 122, 864–875, <https://doi.org/10.1016/j.atmosenv.2015.02.009>, 2015.
- 453 Guo, F., Bui, A. A. T., Schulze, B. C., Yoon, S., Shrestha, S., Wallace, H. W., Sakai, Y., Atkinson, B. W., Erickson, M. H.,
454 Alvarez, S., Sheesley, R., Usenko, S., Flynn, J., and Griffin, R. J.: Urban core-downwind differences and relationships
455 related to ozone production in a major urban area in Texas, *Atmos. Environ.*, 262, 118624,
456 <https://doi.org/10.1016/j.atmosenv.2021.118624>, 2021.
- 457 Guo, S., Hu, M., Wang, Z. B., Slanina, J., and Zhao, Y. L.: Size-resolved aerosol water-soluble ionic compositions in the
458 summer of Beijing: implication of regional secondary formation, *Atmos. Chem. Phys.*, 10, 947–959,
459 <https://doi.org/10.5194/acp-10-947-2010>, 2010.
- 460 Guo, Y., Deng, C., Ovaska, A., Zheng, F., Hua, C., Zhan, J., Li, Y., Wu, J., Wang, Z., Xie, J., Zhang, Y., Liu, T., Zhang, Y.,
461 Song, B., Ma, W., Liu, Y., Yan, C., Jiang, J., Kerminen, V.-M., Xia, M., Nieminen, T., Du, W., Kokkonen, T., and Kulmala,
462 M.: Measurement report: The 4-year variability and influence of the Winter Olympics and other special events on air quality
463 in urban Beijing during wintertime, *Atmos. Chem. Phys.*, 23, 6663–6690, <https://doi.org/10.5194/acp-23-6663-2023>,
464 2023.
- 465 Ho, W. C., Hartley, W. R., Myers, L., Lin, M. H., Lin, Y. S., Lien, C. H., and Lin, R. S.: Air pollution, weather, and associated
466 risk factors related to asthma prevalence and attack rate, *Environ. Res.*, 104, 402–409,
467 <https://doi.org/10.1016/j.envres.2007.01.007>, 2007.
- 468 Jeffries, H. E.: An experimental method for measuring the rate of synthesis, destruction, and transport of ozone in the lower
469 atmosphere, Ph.D. Thesis, Department of Environmental Science and Engineering, University of North Carolina, E.S.E.
470 Publication No.285, 1971.
- 471 Juda-Rezler, K., Reizer, M., Maciejewska, K., Błaszczak, B., and Klejnowski, K.: Characterization of atmospheric PM_{2.5}
472 sources at a Central European urban background site, *Sci. Total Environ.*, 713, 136729,
473 <https://doi.org/10.1016/j.scitotenv.2020.136729>, 2020.
- 474 Kanaya, Y.: Photochemical ozone production at a subtropical island of Okinawa, Japan: Implications from simultaneous
475 observations of HO₂ radical and NO_x, *J. Geophys. Res. Atmos.*, 107, 4368, <https://doi.org/10.1029/2002JD002301>, 2002.
- 476 Kleinman, L. I.: A comparative study of ozone production in five U.S. metropolitan areas, *J. Geophys. Res.*, 110, D02301,
477 <https://doi.org/10.1029/2005JD006401>, 2005.
- 478 Liu, Y., Xu, X., Yang, X., He, J., Ji, D., and Wang, Y.: Significant reduction in fine particulate matter in Beijing during 2022
479 Beijing Winter Olympics, *Environ. Sci. Technol. Lett.*, 9, 822–828, <https://doi.org/10.1021/acs.estlett.2c00532>, 2022.
- 480 Liu, Z., Hu, B., Zhang, J., Yu, Y., and Wang, Y.: Characteristics of aerosol size distributions and chemical compositions during
481 wintertime pollution episodes in Beijing, *Atmos. Res.*, 168, 1–12, <https://doi.org/10.1016/j.atmosres.2015.08.013>, 2016.



- 482 Ma, X., Tan, Z., Lu, K., Yang, X., Chen, X., Wang, H., Chen, S., Fang, X., Li, S., Li, X., Liu, J., Liu, Y., Lou, S., Qiu, W.,
483 Wang, H., Zeng, L., and Zhang, Y.: OH and HO₂ radical chemistry at a suburban site during the EXPLORE-YRD campaign
484 in 2018, *Atmos. Chem. Phys.*, 22, 7005–7028, <https://doi.org/10.5194/acp-22-7005-2022>, 2022.
- 485 Prinn, R. G.: The cleansing capacity of the atmosphere, *Annu. Rev. Environ. Resour.*, 28, 29–57, <https://doi.org/10/b8kg65>,
486 2003.
- 487 Ren, X., Harder, H., Martinez, M., Leshner, R. L., Oligier, A., Simpas, J. B., Brune, W. H., Schwab, J. J., Demerjian, K. L., and
488 He, Y.: OH and HO₂ chemistry in the urban atmosphere of New York City, *Atmos. Environ.*, 3639–3651,
489 [https://doi.org/10.1016/s1352-2310\(03\)00459-x](https://doi.org/10.1016/s1352-2310(03)00459-x), 2003.
- 490 Ren, X., van Duin, D., Cazorla, M., Chen, S., Mao, J., Zhang, L., Brune, W. H., Flynn, J. H., Grossberg, N., Lefer, B. L.,
491 Rappenglück, B., Wong, K. W., Tsai, C., Stutz, J., Dibb, J. E., Thomas Jobson, B., Luke, W. T., and Kelley, P.: Atmospheric
492 oxidation chemistry and ozone production: Results from SHARP 2009 in Houston, Texas, *J. Geophys. Res. Atmos.*, 118,
493 5770–5780, <https://doi.org/10/f5cwzr>, 2013.
- 494 Ren, Y., Stieger, B., Spindler, G., Grosselin, B., Mellouki, A., Tuch, T., Wiedensohler, A., and Herrmann, H.: Role of the dew
495 water on the ground surface in HONO distribution: a case measurement in Melpitz, *Atmos. Chem. Phys.*, 20, 13069–13089,
496 <https://doi.org/10.5194/acp-20-13069-2020>, 2020.
- 497 Sadanaga, Y., Kawasaki, S., Tanaka, Y., Kajii, Y., and Bandow, H.: New system for measuring the photochemical ozone
498 production rate in the atmosphere, *Environ. Sci. Technol.*, 51, 2871–2878, <https://doi.org/10/f9qw59>, 2017.
- 499 Schroeder, J. R., Crawford, J. H., Fried, A., Walega, J., Weinheimer, A., Wisthaler, A., Müller, M., Mikoviny, T., Chen, G.,
500 Shook, M., Blake, D. R., and Tonnesen, G. S.: New insights into the column CH₂O/NO₂ ratio as an indicator of near-surface
501 ozone sensitivity, *J. Geophys. Res. Atmos.*, 122, 8885–8907, <https://doi.org/10.1002/2017JD026781>, 2017.
- 502 Seinfeld, J. H. and Pandis, S. N.: Atmospheric chemistry and physics: from air pollution to climate change, 2nd ed., John
503 Wiley & Sons, Inc., Hoboken, N.J., 1203 pp., 2006.
- 504 Sillman, S.: 9.11 - Tropospheric ozone and photochemical smog, in: *Treatise on Geochemistry*, edited by: Holland, H. D. and
505 Turekian, K. K., Pergamon, Oxford, 407–431, <https://doi.org/10.1016/B0-08-043751-6/09053-8>, 2003.
- 506 Sklaveniti, S., Locoge, N., Stevens, P. S., Wood, E., Kundu, S., and Dusanter, S.: Development of an instrument for direct
507 ozone production rate measurements: measurement reliability and current limitations, *Atmos. Meas. Tech.*, 11, 741–761,
508 <https://doi.org/10/gmttxc>, 2018.
- 509 Styler, S. A., Brigante, M., D’Anna, B., George, C., and Donaldson, D. J.: Photoenhanced ozone loss on solid pyrene films,
510 *Phys. Chem. Chem. Phys.*, 11, 7876, <https://doi.org/10.1039/b904180j>, 2009.
- 511 Tan, Z., Fuchs, H., Lu, K., Hofzumahaus, A., Bohn, B., Broch, S., Dong, H., Gomm, S., Häsel, R., He, L., Holland, F., Li,
512 X., Liu, Y., Lu, S., Rohrer, F., Shao, M., Wang, B., Wang, M., Wu, Y., Zeng, L., Zhang, Y., Wahner, A., and Zhang, Y.:
513 Radical chemistry at a rural site (Wangdu) in the North China Plain: observation and model calculations of OH, HO₂ and
514 RO₂ radicals, *Atmos. Chem. Phys.*, 17, 663–690, <https://doi.org/10.5194/acp-17-663-2017>, 2017.



- 515 Tan, Z., Lu, K., Jiang, M., Su, R., Dong, H., Zeng, L., Xie, S., Tan, Q., and Zhang, Y.: Exploring ozone pollution in Chengdu,
516 southwestern China: A case study from radical chemistry to O₃-VOC-NO_x sensitivity, *Sci. Total Environ.*, 636, 775–786,
517 <https://doi.org/10.1016/j.scitotenv.2018.04.286>, 2018a.
- 518 Tan, Z., Rohrer, F., Lu, K., Ma, X., Bohn, B., Broch, S., Dong, H., Fuchs, H., Gkatzelis, G. I., Hofzumahaus, A., Holland, F.,
519 Li, X., Liu, Y., Liu, Y., Novelli, A., Shao, M., Wang, H., Wu, Y., Zeng, L., Hu, M., Kiendler-Scharr, A., Wahner, A., and
520 Zhang, Y.: Wintertime photochemistry in Beijing: observations of RO_x radical concentrations in the North China Plain
521 during the BEST-ONE campaign, *Atmos. Chem. Phys.*, 18, 12391–12411, <https://doi.org/10.5194/acp-18-12391-2018>,
522 2018b.
- 523 Tan, Z., Lu, K., Jiang, M., Su, R., Wang, H., Lou, S., Fu, Q., Zhai, C., Tan, Q., Yue, D., Chen, D., Wang, Z., Xie, S., Zeng,
524 L., and Zhang, Y.: Daytime atmospheric oxidation capacity in four Chinese megacities during the photochemically polluted
525 season: a case study based on box model simulation, *Atmos. Chem. Phys.*, 19, 3493–3513, <https://doi.org/10.5194/acp-19-3493-2019>, 2019a.
- 527 Tan, Z., Lu, K., Hofzumahaus, A., Fuchs, H., Bohn, B., Holland, F., Liu, Y., Rohrer, F., Shao, M., Sun, K., Wu, Y., Zeng, L.,
528 Zhang, Y., Zou, Q., Kiendler-Scharr, A., Wahner, A., and Zhang, Y.: Experimental budgets of OH, HO₂, and RO₂ radicals
529 and implications for ozone formation in the Pearl River Delta in China 2014, *Atmos. Chem. Phys.*, 19, 7129–7150,
530 <https://doi.org/10.5194/acp-19-7129-2019>, 2019b.
- 531 Tan, Z., Ma, X., Lu, K., Jiang, M., Zou, Q., Wang, H., Zeng, L., and Zhang, Y.: Direct evidence of local photochemical
532 production driven ozone episode in Beijing: A case study, *Sci. Total Environ.*, 800, 148868, <https://doi.org/10/gnzczm>,
533 2021.
- 534 Tang, R., Wu, Z., Li, X., Wang, Y., Shang, D., Xiao, Y., Li, M., Zeng, L., Wu, Z., Hallquist, M., Hu, M., and Guo, S.: Primary
535 and secondary organic aerosols in summer 2016 in Beijing, *Atmos. Chem. Phys.*, 18, 4055–4068,
536 <https://doi.org/10.5194/acp-18-4055-2018>, 2018.
- 537 Tarasick, D., Galbally, I. E., Cooper, O. R., Schultz, M. G., Ancellet, G., Leblanc, T., Wallington, T. J., Ziemke, J., Liu, X.,
538 Steinbacher, M., Staehelin, J., Vigouroux, C., Hannigan, J. W., García, O., Foret, G., Zanis, P., Weatherhead, E.,
539 Petropavlovskikh, I., Worden, H., Osman, M., Liu, J., Chang, K.-L., Gaudel, A., Lin, M., Granados-Muñoz, M., Thompson,
540 A. M., Oltmans, S. J., Cuesta, J., Dufour, G., Thouret, V., Hassler, B., Trickl, T., and Neu, J. L.: Tropospheric Ozone
541 Assessment Report: Tropospheric ozone from 1877 to 2016, observed levels, trends and uncertainties, *Elemen. Sci. Anthro.*,
542 7, 39, <https://doi.org/10.1525/elementa.376>, 2019.
- 543 Whalley, L. K., Stone, D., Dunmore, R., Hamilton, J., Hopkins, J. R., Lee, J. D., Lewis, A. C., Williams, P., Kleffmann, J.,
544 Laufs, S., Woodward-Massey, R., and Heard, D. E.: Understanding in situ ozone production in the summertime through
545 radical observations and modelling studies during the Clean air for London project (ClearfLo), *Atmos. Chem. Phys.*, 18,
546 2547–2571, <https://doi.org/10.5194/acp-18-2547-2018>, 2018.
- 547 Whalley, L. K., Slater, E. J., Woodward-Massey, R., Ye, C., Lee, J. D., Squires, F., Hopkins, J. R., Dunmore, R. E., Shaw, M.,
548 Hamilton, J. F., Lewis, A. C., Mehra, A., Worrall, S. D., Bacak, A., Bannan, T. J., Coe, H., Percival, C. J., Ouyang, B.,



- 549 Jones, R. L., Crilley, L. R., Kramer, L. J., Bloss, W. J., Vu, T., Kotthaus, S., Grimmond, S., Sun, Y., Xu, W., Yue, S., Ren,
550 L., Acton, W. J. F., Hewitt, C. N., Wang, X., Fu, P., and Heard, D. E.: Evaluating the sensitivity of radical chemistry and
551 ozone formation to ambient VOCs and NO_x in Beijing, *Atmos. Chem. Phys.*, 21, 2125–2147, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-21-2125-2021)
552 21-2125-2021, 2021.
- 553 Xu, X., Lin, W., Xu, W., Jin, J., Wang, Y., Zhang, G., Zhang, X., Ma, Z., Dong, Y., Ma, Q., Yu, D., Li, Z., Wang, D., and
554 Zhao, H.: Long-term changes of regional ozone in China: implications for human health and ecosystem impacts, *Elem. Sci.*
555 *Anth.*, 8, 1–27, <https://doi.org/10.1525/elementa.409>, 2020.
- 556 Xue, L. K., Wang, T., Guo, H., Blake, D. R., Tang, J., Zhang, X. C., Saunders, S. M., and Wang, W. X.: Sources and
557 photochemistry of volatile organic compounds in the remote atmosphere of western China: results from the Mt. Waliguan
558 Observatory, *Atmos. Chem. Phys.*, 13, 8551–8567, <https://doi.org/10/f49wrr>, 2013.
- 559 Xue, M., Ma, J., Tang, G., Tong, S., Hu, B., Zhang, X., Li, X., and Wang, Y.: RO_x budgets and O₃ formation during
560 summertime at Xianghe suburban site in the North China Plain, *Adv. Atmos. Sci.*, 38, 1209–1222, <https://doi.org/10/gkf965>,
561 2021.
- 562 Zhou, W., Cohan, D. S., and Henderson, B. H.: Slower ozone production in Houston, Texas following emission reductions:
563 evidence from Texas Air Quality Studies in 2000 and 2006, *Atmos. Chem. Phys.*, 14, 2777–2788,
564 <https://doi.org/10.5194/acp-14-2777-2014>, 2014.
- 565 Zhou, X., Gao, H., He, Y., Huang, G., Bertman, S. B., Civerolo, K., and Schwab, J.: Nitric acid photolysis on surfaces in low-
566 NO_x environments: Significant atmospheric implications, *Geophys. Res. Lett.*, 30, n/a-n/a,
567 <https://doi.org/10.1029/2003GL018620>, 2003.
- 568