



# A Modified Gaussian Plume Model for Mobile in situ Greenhouse Gas Measurements

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## Abstract.

Atmospheric methane measurements are important for evaluating high resolution methane inventories and monitoring emissions reductions. Despite recent international efforts to harmonize measurement methodologies and techniques, currently there are no standardized or internationally accepted techniques for estimating emissions from mobile in situ concentration measurements. We present measurements from two different mobile in situ methane laboratories, and compare emission rates calculated from four Gaussian plume Bayesian optimal estimation strategies and a statistical algorithm. For mobile transects from the slower flow-rate instrument, we find a significant asymmetric smoothing artifact. The effect of this asymmetry is most significant for short transects of small ( $0\text{-}50\text{ kg CH}_4\text{ day}^{-1}$ ), nearby methane sources, where the plume crossing time is comparable to the mean residence time of the instrument. We develop a model of this effect, demonstrate how this model can be applied to Gaussian plume inversions, and describe its limitations. We use these results to compute emissions rate estimates for two methane sources from Toronto's wastewater management system to demonstrate the use and limitations of Gaussian plume inversions to quantify methane emissions in an urban environment. Overall, we highlight the importance of using observed plume enhancement areas rather than the more commonly used enhancement heights for determining comparable emissions estimates between different mobile laboratories.

## 1 Introduction

Mitigating anthropogenic methane emissions is one of many critical steps necessary to combat climate change. Atmospheric measurements are necessary to monitor and validate the impact of the emissions reductions strategies taken by municipal, regional, and national entities in order for us to track our progression towards new climate targets. Cities have been at the forefront of planning necessary decarbonization pathways, including Toronto's TransformTO aim to achieve net zero  $\text{CO}_{2eq}$  emissions by 2040 (The City of Toronto, 2021). Methane has a global warming potential that is 28 times that of  $\text{CO}_2$  over a 100 year period, and is a valuable energy resource (IPCC, 2013). In the atmospheric carbon cycle, methane eventually reacts to form  $\text{CO}_2$ , so mitigating emissions by flaring, or fuel use is less harmful than allowing direct emissions of methane to the



atmosphere (Kirschbaum, 2014; Tavakkoli et al., 2022). In the urban environment, when compared to the nearly-ubiquitous emissions of CO<sub>2</sub>, there is a relatively small number of highly emitting facilities and point sources of CH<sub>4</sub>, making it a good candidate for mobile survey measurement strategies. There are two significant sources of methane emissions from methane gas fuel use in urban environments. To distinguish emissions from the primarily fossil fuel derived gas used for domestic energy use from the biogas produced at sanitation facilities, we will use the terms fossil gas, and biogas, respectively. In the EDGARv7.0 greenhouse gas emission inventory, cities account for approximately 15% of anthropogenic methane emissions in Canada (Crippa, M. et al., 2022). For the Greater Toronto Area, Canada's largest and most populous metropolitan area, the Facility Level and Area Methane Emissions (FLAME-GTA) inventory is the most detailed and comprehensive assessment of methane emissions (Pak et al., 2021). FLAME-GTA, as with other emissions inventories, relies on bottom-up accounting methods, generally either derived from emissions factors weighted by population density, or production activity. Uncertainties in these emissions factors can be significant; in some cases larger than 50%. Additionally, bottom-up inventories are not exhaustive lists, potentially missing emissions sources, particularly smaller industrial sources which are below reporting thresholds and small point sources from sewers or pneumatic valves in the fossil gas distribution system (Williams et al., 2022). Mobile emissions monitoring using atmospheric in situ measurements is a useful technique for verifying emissions and constraining inventory uncertainties for point sources and facilities, and tracking the seasonal trends of those methane emissions. However, mobile in situ monitoring is not able to easily quantify diffuse emissions sources, such as urban wetlands, post-meter indoor gas leaks, and inefficient appliance combustion, which have recently been shown to be a significant contributor to the urban atmospheric methane burden (Lebel et al., 2022; Herrero Ortega et al., 2019).

There is a large and growing body of literature utilizing vehicle based mobile in situ GHG concentration data to estimate methane emissions from known sources using a variety of techniques. Currently, research groups utilize different gas analyzers, experimental setups, and data analysis techniques for estimating methane emissions from mobile in situ data. In order for this data to be useful for inventory assessment purposes and comparable to other studies, standardized or comparable measurement methodologies should be adopted. Recently, there has been an effort to harmonize methodologies used to quantify emissions from mobile in situ measurements to improve the comparability of estimates from different researchers and organizations (Turnbull et al., 2022; Wielgosz et al., 2023).

Here, we present a brief summary of common methodologies used to analyze mobile in situ CH<sub>4</sub> concentrations. The first such technique is the coincidental controlled release, where a gas is released into the atmosphere with a known emission rate near a GHG emission source with an unknown emission rate. Both gases are measured simultaneously during a mobile survey, allowing for a simple tracer ratio calculation to estimate emissions. This technique has been used to measure emissions from urban infrastructure (Lamb et al., 1995; Shorter et al., 1996), landfills (Mønster et al., 2015), and fossil oil and gas infrastructure (Lamb et al., 2015; Delre et al., 2022).

It is not always practical or possible to conduct coincidental controlled release experiments, as additional instrumentation or site access is often required. When coincidental controlled release experiments are not possible, Gaussian plume inversions of a single measured gas have been used to quantify methane emissions from a variety of different sources, including emissions from the oil and gas sector in Canada (Atherton et al., 2017; Zavala-Araiza et al., 2018; Baillie et al., 2019; O'Connell et al.,



2019; Vogt et al., 2022), the United States (Yacovitch et al., 2015), and Europe (Yacovitch et al., 2018; Kumar et al., 2022). One study in the United States utilized this technique for quantifying emissions from 63 wastewater treatment plants (WWTP) (Moore et al., 2023). Other mobile experiments have utilized multiple GHG analyzers measuring at different inlet heights, which allows for a more thorough analysis of observed downwind plumes, which has been used to quantify CH<sub>4</sub> emissions from active and inactive oil and gas infrastructure (Rella et al., 2015; Lebel et al., 2020). A third prominent technique utilizing mobile platforms for quantifying emission rates from in situ measurements is the Other Test Method (OTM) 33a technique, which estimates emissions from the co-variation in downwind concentration and wind direction for a stationary downwind sensor (Thoma and DeWees, 2014). Mobile measurement platforms are typically parked downwind of sources for 20-30 minute periods to determine how concentrations vary with wind direction. This technique has been used to measure CH<sub>4</sub> emissions from oil and gas infrastructure in the United States (Brantley et al., 2014), Canada (Vogt et al., 2022; Hugenholtz et al., 2021), Europe (Yacovitch et al., 2018; Kumar et al., 2022), and compressed fossil gas facilities in Eastern China (Wang et al., 2022). Collectively, these in situ emissions quantification techniques have enabled researchers to quantify emissions from individual facilities, in addition to updating national inventory estimates on a sector-wide basis (MacKay et al., 2021).

Quantifying emissions from urban sources can be challenging because individual point sources of methane emissions can be harder to correctly identify in an environment with more built infrastructure, wind tunnelling, and gas distribution leakage. However, Gaussian plume inversions of mobile in situ data have been used to quantify urban methane emissions from known sources using vehicle based surveys (Ars et al., 2020; Maazallahi et al., 2020; Moore et al., 2023) and walking surveys (Chen et al., 2020). One study used a computational fluid dynamics model to drive a large eddy simulation to invert observed mobile in situ CH<sub>4</sub> plumes from large sources in Indianapolis (Lamb et al., 2016).

In addition to the aforementioned quantification strategies, which explicitly account for the location and likely source of emissions, statistical algorithms have been developed to estimate emissions rates from leaks in fossil gas distribution based solely on observed atmospheric concentration enhancements measured by mobile in situ laboratories (Von Fischer et al., 2017; Weller et al., 2019). These algorithms are most useful for datasets missing the required meteorological or source location data necessary for specific plume attribution and quantification. They have been widely applied to urban survey data for determining relevant statistics on the size and variability of fugitive methane emissions from urban fossil gas distribution networks (Weller et al., 2020; Ars et al., 2020; Defratyka et al., 2021; Maazallahi et al., 2020; Fernandez et al., 2022; Vogel et al., 2023a).

All of these emissions quantification strategies can be evaluated using controlled release experiments, where the target gas of interest is released into the atmosphere at a known rate. Such experiments have been widely used to evaluate and calibrate different emissions quantification methodologies (Von Fischer et al., 2017; Ars et al., 2017; O'Connell et al., 2019).

In this paper, we analyze a dataset consisting of mobile in situ atmospheric methane concentrations collected in the Greater Toronto Area. We evaluate various Gaussian plume inversion strategies, and other emissions quantification methodologies using data from a controlled release experiment and from surveys downwind of two types of wastewater facilities. In Section 2, we describe the datasets and inversion strategies. We evaluate the impact of different GHG analyzers with different flow rates on our observed mobile transect data, and in Section 2.3, we develop an instrument response function for the instrument with the slower flow-rate. In Section 2.4, we analyze data from a controlled release experiment to evaluate the inversions. In



Section 3, we apply our Gaussian plume inversion methods to quantify emissions from urban wastewater infrastructure. In Section 4, we compare and discuss the emission rate quantification techniques. Finally, in Section 5, we present conclusions and implications for future mobile GHG quantification experiments.

## 2 Methods

In this section, we describe the datasets collected using our mobile laboratories in Section 2.1, and our Gaussian plume inversion techniques in Section 2.2. We describe the need for developing an instrument response function for our slower flow-rate instrument in Section 2.3, and evaluate the inversion techniques with and without the instrument response function among other sensitivity studies in Section 2.4.

### 2.1 Datasets

Since 2017, we have used a Los Gatos Research (San Jose, CA) Ultraportable Greenhouse Gas Analyzer (UGGA) to measure atmospheric concentrations of CO<sub>2</sub>, CH<sub>4</sub>, and H<sub>2</sub>O from a mobile bicycle-trailer-based laboratory around the City of Toronto, Ontario, Canada (Figure 1a). Simultaneous 1 Hz GHG concentrations, meteorological information, and GPS position data are logged on a laptop computer. We have also occasionally used this analyzer in a vehicle based laboratory to survey sites outside of the city, but these data are not evaluated in this work, aside from a side-by-side comparison described in the Supplemental Information (S1.1). Since 2018, we have also used a vehicle based mobile laboratory which is equipped with various Picarro cavity ring-down spectrometers (G1301, G2401) fitted with exterior high flow-rate pumps to record CH<sub>4</sub>, CO<sub>2</sub> and H<sub>2</sub>O concentrations every two seconds (Figure 1b). Both platforms are described further in Ars et al. (2020), and in the Supplemental Information (S2). In this paper, our focus is on the methane and meteorological measurements from both mobile laboratories (Wunch et al., 2018; Ars et al., 2019; Gillespie et al., 2022a, b, 2023).

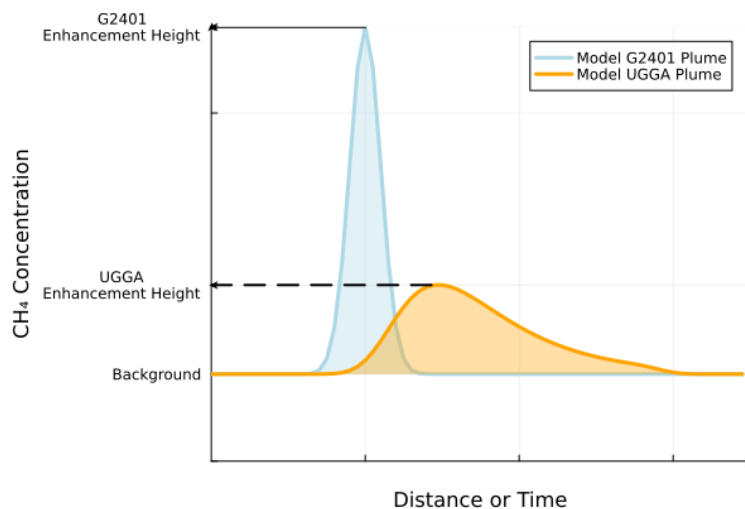
### 2.2 Gaussian Plume Inversions

We estimate instantaneous CH<sub>4</sub> emissions rates by using Bayesian optimal estimation to fit a Gaussian plume dispersion model to our downwind concentration enhancement measurements. In order to estimate an instantaneous emission rate from an emission source, such as a WWTP or power plant, we use the mobile laboratories to transect plumes downwind of methane sources. To isolate the contribution of specific sources, we subtract the fifth percentile concentration to remove the local background concentration, isolating the concentration enhancements from the upwind source. Then, we use these concentration enhancements to model the upwind emission rates.

When optimizing the Gaussian plume model, one must choose which scaling factor to optimize. Two common choices to scale are the enhancement height or the enhancement area. These scaling factors comprise a vector of scaling factors,  $\mathbf{p}$ , with each element of the vector describing an equal distance portion of the transect to be optimized. The enhancement height is the difference between the maximum concentration in the plume and the background concentration. The enhancement area is the integral of the plume enhancement above the background. Both of these quantities are sketched in Figure 2. The plume vector



**Figure 1.** (a) The University of Toronto’s bicycle based laboratory equipped with an LGR UGGA. (b) ECCC’s vehicle based laboratory equipped with a Picarro G1301 or G2401. Both setups use Airmar WX220 weather stations to record meteorological and GPS data.



**Figure 2.** A model representation of the trace gas concentrations in narrow downwind plumes as observed by our mobile laboratories for an identical emission event. The shaded plume areas are the same, while the enhancement height for the UGGA is significantly lower.

is related to the Gaussian dispersion model:

$$125 \quad p = \mathbf{H}f + \epsilon_0 \tag{1}$$

where  $\epsilon_0$  represents the combined uncertainties resulting from measurement uncertainty in the methane concentrations, and uncertainties from the determination in the background concentration.  $\mathbf{H}$  is the operator that transforms the vector of emission rates,  $f$ , which are in units of mass-time<sup>-1</sup>, into the plume vector,  $p$  space, which is in units of ppb, ppb-m, or ppb-s. There



are two options for the plume axis when optimizing the enhancement area: one can integrate under the enhancement curve  
130 using units of time (ppb·s) or distance (ppb·m). When the mobile laboratory maintains a steady velocity during the transect,  
there is little difference between these methods. However, if the mobile laboratory stops during the transect, then a time-based  
inversion would weigh each observation equally in time, while the distance-based model would correctly weigh the data point  
from the stationary location. If a significant portion of the plume transect is spent in stop-and-go traffic, an inversion optimizing  
the concentration-time area might be preferential to take into account more of the observed data.

135 The Bayesian optimal estimate of the emissions,  $f^a$ , can be calculated as

$$f^a = f^b + \mathbf{B}\mathbf{H}^T (\mathbf{R} + \mathbf{H}\mathbf{B}\mathbf{H}^T)^{-1} (p - \mathbf{H}f^b), \quad (2)$$

where  $f^b$  is the vector of prior emissions rates,  $\mathbf{B}$  is a matrix representing the uncertainty in the prior emissions estimate. For  
all of our inversions, we assume an 80% uncertainty in our prior values, so that our inversions are only loosely constrained  
by our priors.  $\mathbf{R}$  is a matrix representing the atmospheric transport model errors, and other measurement errors from the  
140 background determination and instrumentation errors. In our inversion, we assume that  $\mathbf{B}$  and  $\mathbf{R}$  are diagonal matrices. We  
assume constant values for  $\mathbf{R}$ , depending on the transect distance. For our small plumes, we use a value of 1500 ppb·m, and  
for the larger transects, we used a model error of 15 000 ppb·m. The variance-covariance uncertainty matrix,  $\mathbf{A}$ , is defined as

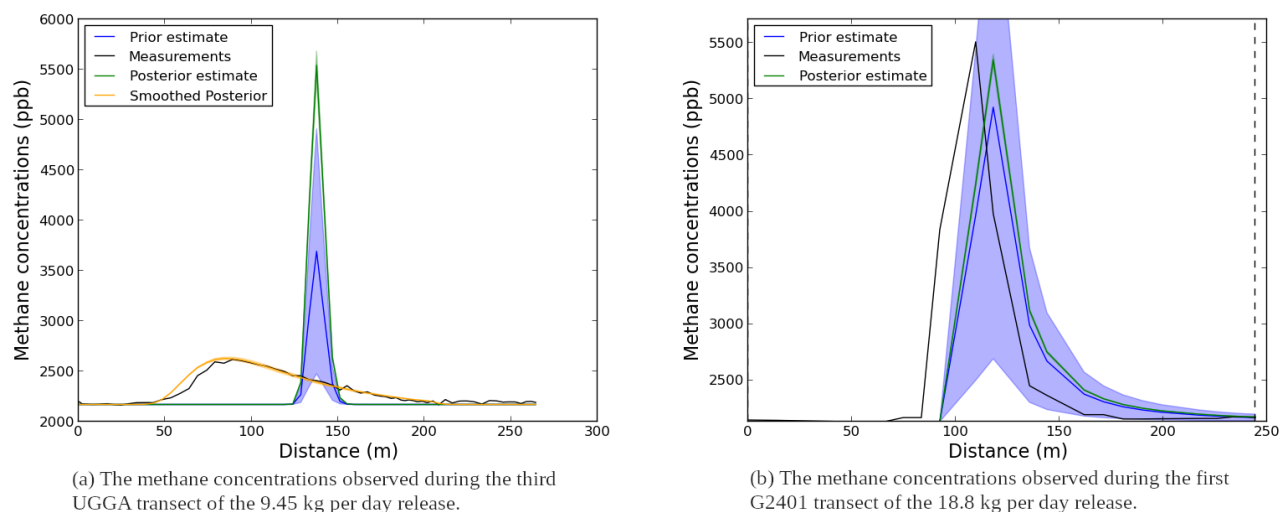
$$\mathbf{A} = (\mathbf{B}^{-1} + \mathbf{H}^T \mathbf{R}^{-1} \mathbf{H})^{-1}. \quad (3)$$

Information in the  $\mathbf{A}$  matrix can also be used to determine how well each source in the optimal estimate vector,  $f^a$ , is con-  
145 strained by the inversion, as well as to provide an estimate of the uncertainty of the inversion.

The Gaussian plume dispersion model we use is the Polyphemus air quality modelling system (Mallet et al., 2007). The  
horizontal and vertical diffusivities in Polyphemus are the Briggs's formulation of the Pasquill-Turner stability classes, which is a  
flexible parameterization of the Gaussian plume model (Briggs, 1973). The stability classes are determined from the insolation,  
surface wind speed, and surface roughness estimated from land cover type. When reliable stationary in situ meteorology data  
150 are available, we use the  $\sigma_a$  method of determining transect-specific stability classes, which uses wind speed and variability in  
azimuthal wind direction to determine stability class (Bailey et al., 2000). We adjust the initial model parameters, principally  
wind direction, to align observed and modelled Gaussian plumes. Then, we use this wind-rotated Gaussian plume dispersion  
model as our operator,  $\mathbf{H}$ .

### 2.3 Instrument Response Function

155 For many of our small plume transects observed with the bicycle laboratory, the standard Gaussian plume model provides a  
poor fit to our observations. Similar to the findings of Takriti et al. (2021), we observe a significant smoothing of the measured  
plumes, which is a consequence of the lower flow-rate of air through the instrument relative to the plume transect time. The  
low flow-rate of the UGGA causes an asymmetric smoothing for short plume transects where the time it takes to cross the  
plume is similar to the mean residence time of the instrument. The UGGA's mean residence time in the detection cavity  
160 is around 8 seconds with our setup. This, combined with the 1Hz sampling frequency resolves the asymmetric instrument



**Figure 3.** Examples of measurements collected by the mobile platforms downwind of the gas outlet during the controlled release experiment. In panel (a), one transect from LGR UGGA measurements is shown from the 9.45 kg·day<sup>-1</sup> release. In panel (b), a Picarro G2401 measurement transect is shown from the 18.8 kg·day<sup>-1</sup> release. The measured CH<sub>4</sub> concentrations are shown in black and plotted as a function of the transect distance in metres. The Gaussian plume prior concentrations are shown in blue, and the Gaussian plume posterior is green. The posterior of the smoothed Gaussian plume area inversion is shown in orange in panel (a).

response function when sampling a narrow concentration perturbation. For the Picarro instrument in vehicle based lab, the mean residence time is roughly half of the sampling frequency, 0.5Hz. This, when combined with the slower sampling frequency, means that observations better represent in situ concentrations in space, which agrees better with the Gaussian plume models. A model representation of the concentrations measured by the two instruments for a transect that lasts on the order of 8 seconds is shown in Figure 2. A similar figure showing real data is in Figure S1.

To more accurately model the instrument response function, we use a normalized heavy-tailed distribution function given by equation 4.

$$f(t, \sigma, \mu, n) = \frac{\frac{1}{\sigma\sqrt{2\pi}} \exp\left(-\frac{(\log t - \mu)^2}{2\mu^2}\right)}{\sum_0^n \frac{1}{\sigma\sqrt{2\pi}} \exp\left(-\frac{(\log t - \mu)^2}{2\mu^2}\right)}, \quad (4)$$

where  $t$  is the measurement time in seconds,  $\mu$  is a variable adjusting the horizontal coordinate of the vertex,  $\sigma$  is a variable adjusting the vertical coordinate of the vertex, and  $n$  is the length of the convolution window. The values of the variables chosen to represent the instrument response function for the UGGA are  $n = 37$ ,  $\mu = 2.2$ , and  $\sigma = 0.65$ . These values were estimated by fitting our observed concentrations to Gaussian plume models, and from a vehicle based campaign where coincident measurements from the higher flow-rate Picarro G2401 and LGR UGGA were conducted. Convolving the observed Picarro measurements with our asymmetric smoothing function (Equation 4) significantly increases the correlation between the mobile



**Table 1.** Summary of main results from the controlled release experiment using the  $\sigma_a$  method for prescribing P-G Stability Classes

Methane Release Rate (kg·day <sup>-1</sup> )	Number of Bicycle Transects	Number of Car Transects	Mean Estimated Emission Rate (kg·day <sup>-1</sup> ) [Bike, Enhancement Area]	Mean Estimated Emission Rate (kg·day <sup>-1</sup> ) [Car, Enhancement Area]
2.36	3	7	1.7 ± 0.5	2.9 ± 1.8
4.72	3	4	6.7 ± 0.6	8.1 ± 8.0
9.45	2	3	15.6 ± 0.6	11.7 ± 12.7
18.89	4	5	22.2 ± 19.7	17.4 ± 4.8

175 observations (Figure S1). Further discussion of this experiment is found in the Supplemental Information (S1.1). An example of a smoothed Gaussian plume model, compared to measured concentrations is also shown in Figure 3 a.

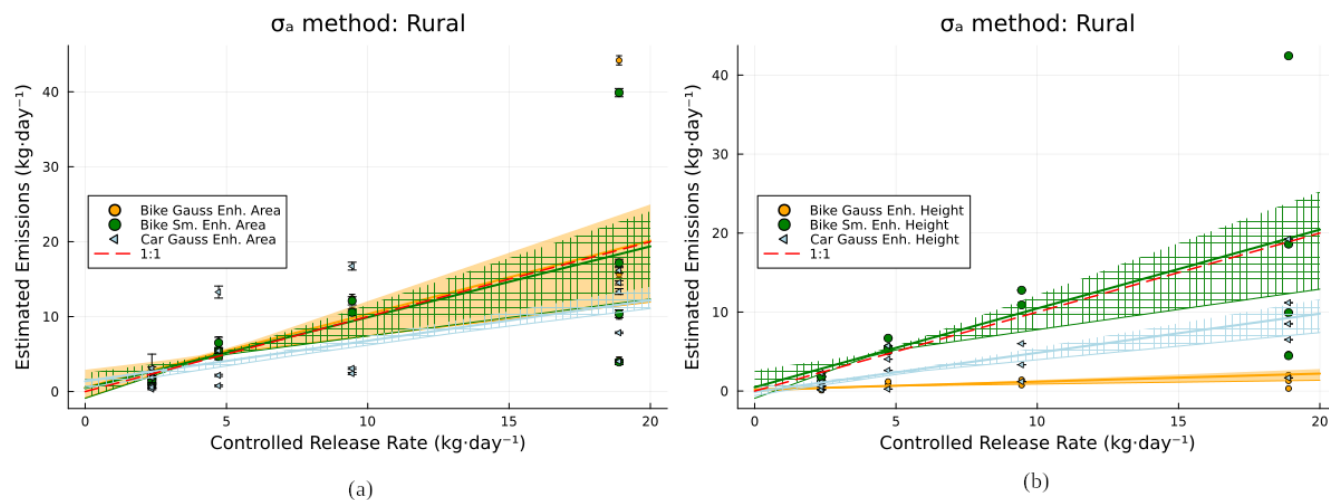
We evaluate four different versions of the Gaussian plume inversion using a controlled release experiment described in Section 2.4. The four versions are: 1. enhancement height, 2. enhancement area, 3. smoothed enhancement height, and 4. smoothed enhancement area. Tests 3 and 4 apply smoothing by convolving the modelled concentrations with the instrument response function in Equation 4.

## 2.4 Controlled Release Experiment

On 20 October 2021, we released CH<sub>4</sub> at a controlled and known emission rate, using a Masterflex proportional flowmeter controller. For the duration of the experiment, the gas outlet was located at ground level on the sidewalk in Toronto's industrial Port Lands neighbourhood, at the coordinates (43.655008°N, -79.325247°W), as shown in Fig S2. There were no buildings between the release location and the downwind transects, which were conducted in a live traffic environment. Wind speed and direction were recorded at the release location with a Vaisala WXT536 weather station at a 2m height directly above the gas outlet for the duration of the experiment. Wind roses from the release period from the three wind sensors, the stationary Vaisala WXT536, and two mobile Airmar WX220 units are discussed in the Supplemental Information (S3.3), but winds were most often from the WSW with a 75th percentile wind speed of 2.0 ms<sup>-1</sup>. Both the bicycle laboratory and vehicle based setup measured multiple transects of each of the four release rates, summarized in Table 1.

We designed this experiment to simulate small urban plume conditions often encountered in our bicycle and vehicle surveys. These plumes are often transected very near the emissions sources. Transect inversions at these distances are highly uncertain, as the horizontal and vertical plume dispersion is generally not well mixed at these distances. Additionally, the uncertainty in emission rate estimates is expected to asymptote as the distance from the source approaches zero (Ars et al., 2017). After we noticed a significant variability in inversion results from the bicycle laboratory's closest transects, we removed transects with a minimum distance to the emissions source of less than 6.25m. Results using closer transects are discussed further in the Supplemental Information (S3.4). Comparing the observations between the ECCC vehicle laboratory and the UofT bicycle lab, the asymmetric smoothing is obvious in the UGGA plumes, and is less evident in the plumes observed with the faster flow-rate





**Figure 4.** The estimated CH<sub>4</sub> emission rates are plotted against the known emission rate from our controlled release experiment performed on 2021-10-20. (a) shows estimates calculated using the enhancement area calculations, and (b) shows the emissions estimates calculated using the enhancement heights. Data from the UGGA bicycle lab estimates calculated using the normal Gaussian plume inversions are shown in orange. The green shows the estimates from the same bike data calculated using the smoothed Gaussian plume inversions. The Gaussian plume estimates from the G2401 vehicle are shown in light blue. The ribbons indicate the total range of bootstrapped linear fits to these data.

Picarro setup, as seen in two plumes in Figure 3. A summary of the estimated emissions rates using the enhancement areas, and enhancement heights for the vehicle as well as the smoothed Gaussian enhancement area and height for the bicycle laboratory are shown in Figure 4.

From our controlled release experiment, we compute a linear regression between the emissions estimates from the bicycle based lab and the known release rate. Using the  $\sigma_a$  method of azimuthal wind fluctuations for determining P-G stability class (Bailey et al., 2000; Mitchell, 1982), the slope and uncertainty of the fit is  $1.18 \pm 0.46$ . For the vehicle laboratory, the slope of the line of best fit is  $0.83 \pm 0.21$ . This demonstrates that for these nearby, close-to-source transects, we can assume a nominal uncertainty on the order of 40% for inversions measured with the bicycle lab.

The largest source of uncertainty in Gaussian plume inversions is typically from the choice of atmospheric stability class (Ars et al., 2017), and therefore similar plots for each stability class are shown in Figure S3. Table 2 gives the slopes of the lines of best fit for the controlled release experiments evaluated with different stability classes calculated using the Gaussian enhancement area for the bicycle lab, the vehicle lab, and the combined datasets.

For the UGGA bicycle lab, results comparing the Gaussian height inversions and the smoothed height inversions are presented in Figure S4, and Figure S5 shows the UGGA smoothed height inversions alongside the car's Gaussian plume height inversions. Generally, the Gaussian height inversions for the UGGA systematically return smaller optimized emission rates. The smoothed enhancement height inversions for the UGGA lab are consistent with the area-based inversions, demonstrating that one must account for the instrument response function of the UGGA only when optimizing the enhancement height of the



**Table 2.** Summary of the slopes for the best fit lines estimated with transects from our controlled release experiments using different stability classes and the Gaussian enhancement area.

Stability Class	Bicycle Data Line of Best Fit	Car Data Line of Best Fit	Combined Data Line of Best Fit
Rural A	$1.66 \pm 0.67$	$1.18 \pm 0.24$	$1.39 \pm 0.29$
Rural B	$1.20 \pm 0.46$	$0.83 \pm 0.21$	$0.98 \pm 0.21$
Rural C	$1.14 \pm 0.42$	$0.75 \pm 0.33$	$0.91 \pm 0.25$
Urban A/B	$1.96 \pm 0.80$	$1.50 \pm 0.27$	$1.71 \pm 0.34$
Urban C	$1.66 \pm 0.67$	$1.18 \pm 0.24$	$1.39 \pm 0.29$
Rural $\sigma_a$	$1.18 \pm 0.46$	$0.83 \pm 0.21$	$0.98 \pm 0.22$
Urban $\sigma_a$	$1.94 \pm 0.80$	$1.19 \pm 0.25$	$1.52 \pm 0.34$

Gaussian plume model. The Gaussian enhancement height inversions for the Picarro also underestimate the emissions with respect to the enhancement area inversions with a best fit slope of  $0.65 \pm 0.16$ . The UGGA is significantly more underestimated, with a best fit slope of  $0.12 \pm 0.05$ .

The smoothed plume enhancement height inversions were affected more severely by outliers, and have a slope and offset further from the 1:1 line than the area inversions.

Based on the results of our controlled release experiment, we argue that when optimizing emissions using the emissions enhancement height, it is necessary to account for the instrument flow-rate, particularly for low flow-rate instruments. We find that simply using the Gaussian plume area, whether or not we take into account the instrumental smoothing, also provides a better estimate of the release rate. Accounting for smoothing in enhancement area inversions does not significantly change the results. These findings agree well with previous studies by Ars et al. (2017), Ars et al. (2020), and Takriti et al. (2021). As a practical application of our insights in inversion comparability, for the remainder of this paper we present emissions derived using the plume area. Further discussion of the different emissions quantification techniques is found in Section 4.

### 3 Quantifying Methane Sources in the Greater Toronto Area

In this section, we quantify emissions from plume transects downwind of large and small methane emissions sources, demonstrating the importance of accounting for the instrument flow-rate for modelling our shorter, near-to-source transects. In Section 3.1, we present the Gaussian plume inversion results for downwind transects of the largest wastewater treatment facility in the Greater Toronto Area, the Ashbridge's Bay Wastewater Treatment Plant (ABWWTP). These plume crossings are, at the closest, 100 m from the modelled sources, and a short plume transect of this facility is nearly 1 km long, which takes between 3-10 minutes to transect by bicycle. Under these conditions, smoothing the Gaussian plume model has minimal impact for modelling our mobile in situ  $\text{CH}_4$  observations. In Section 3.2, we present our results from observed methane plumes from the storage shafts of the Western Beaches Tunnel system, a series of three underground storage basins for retaining Combined



Sewage Overflow (CSO). These transects are quick crossings of closer, smaller plumes, (less than 20m from the source locations, around 150m in length, and 10-20 seconds transecting the plume) and are typical of the small enhancements in our previous urban observations (Ars et al., 2020).

### 240 3.1 Ashbridge's Bay Wastewater Treatment Plant

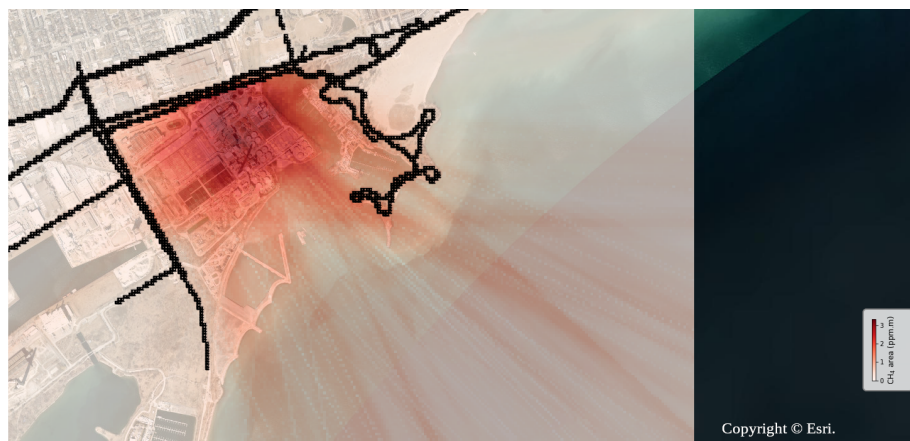
The Ashbridge's Bay Wastewater Treatment Plant is the largest sewage treatment facility in the GTA, processing the waste water of an estimated 1 733 500 people (The City of Toronto, 2022). The FLAME-GTA inventory reports that ABWWTP is the third largest single methane source in the City of Toronto, with an annual emission rate of 787 t CH<sub>4</sub>yr<sup>-1</sup> (Pak et al., 2021). The FLAME-GTA estimate is derived from the volume of biogas produced at ABWWTP, and accounts for biogas collection efficiency, and methane destruction efficiencies (Pak et al., 2021). In Canada's GHG reporting requirements, however, municipal wastewater treatment plants are not required to report methane emissions from the wastewater treatment process. The methane emissions reported to ECCC's GHG reporting program for ABWWTP (0.207 t CH<sub>4</sub> yr<sup>-1</sup>) only include estimated fugitive emissions from the combustion of fossil gas brought into the WWTP (Environment and Climate Change Canada, 2021), and are not an estimate of methane emissions resulting from the biogas produced there.

245 From 2018, we have measured 44 downwind transects of methane plumes from ABWWTP on 21 different days using the UGGA bicycle laboratory, and 14 transects on 5 different days using the ECCC vehicle laboratory. These transects were conducted on public roads and bicycle paths. We revisited this site frequently to evaluate variability in emissions rates on various time scales, from seasonally to sub-hourly. Recently, Gålfalk et al. (2022) observed significant variability in CH<sub>4</sub> emissions on sub-hourly time scales following aerated and rest periods in WWTP infrastructure. To evaluate this possible variability, we conducted multiple transects over the course of a single day.

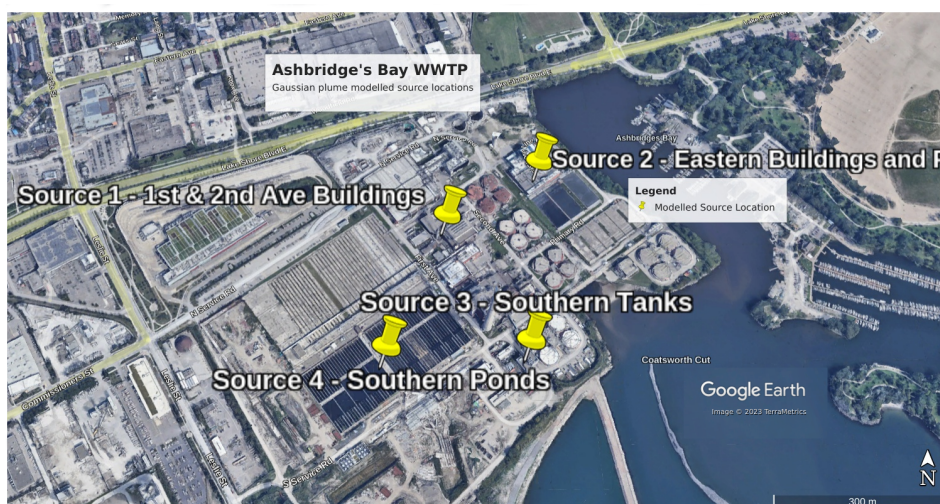
255 To estimate emissions from the ABWWTP, we first located the main emissions sources by summing the concentration-area of upwind concentration enhancements for each transect (Fig. 5a). We do this to locate grid cells which are repeatedly upwind of observed enhancements across a variety of wind directions. Here, we consider only wind direction and not wind speed for these calculations. We average observed in situ methane concentrations in 15m grid cells from each downwind transect. Then, we trace lines upwind of each grid cell, and add the average downwind measured concentration-area to the upwind grid cells. This approach failed to discriminate individual emissions sources, but it demonstrates that the majority of methane emissions we observed likely originate from the buildings located near the centre of the facility. The ponds and tanks also appear to be likely emissions sources.

265 For our Gaussian plume inversions, we use the fewest possible source locations to fit the major observed peaks from the transects with consistent observed winds, such that the wind direction did not change during the transect. We use the smallest number of modelled sources to avoid under-determined plume inversions. Based on our source modelling technique, we choose the buildings between First and Second Avenue as the location of the first point source (Fig. 5 b). The second source we include corresponds to the Eastern buildings and ponds. From fitting plumes advected by Westerly winds, measured along the Eastern bank of Ashbridge's Bay, we observed a persistent plume which required us to add a third source located at the Southern tanks. Similarly, when fitting plumes along Leslie St, the public road to the West of the facility under Easterly winds, we required a

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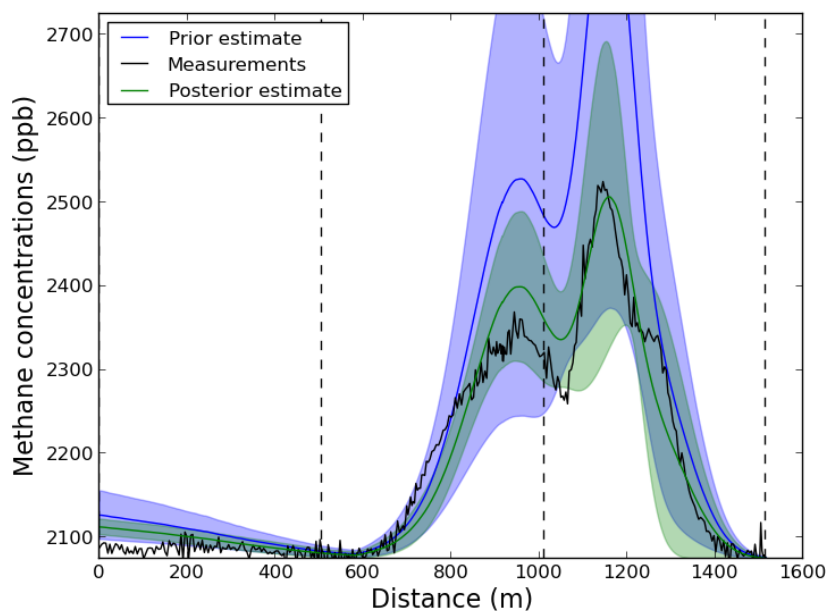
(a) From ESRI (2022).



(b) From Google Earth (2023).

**Figure 5.** (a) Average concentrations in 15m grid cells from all transects are traced upwind. The heatmap is created from the sum of all of the upwind traces, with darker red indicating a higher total upwind concentration area, corresponding to likely emissions source locations. The black boxes represent the grid boxes sampled in the downwind transects. (b) ABWWTP source locations for the Gaussian plume inversion model.

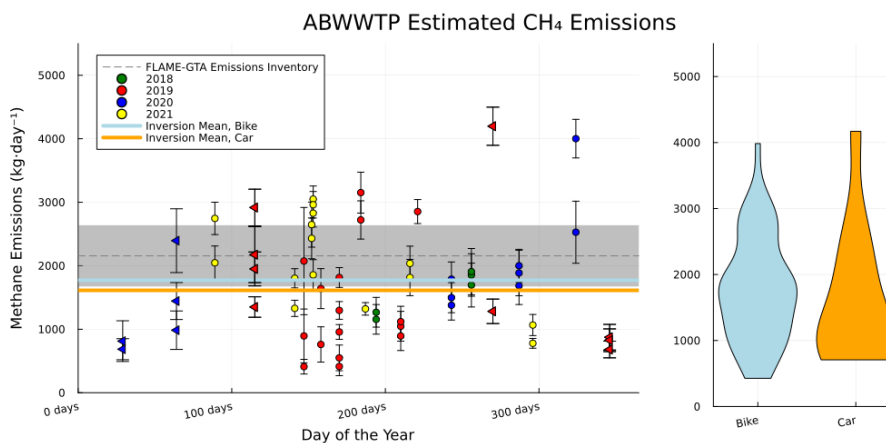
fourth area source located at the Western ponds. All sources are labelled in Figure 5 b. The first three locations are modelled as point sources, since they are potential sites for leaks in biogas collection infrastructure. We modelled the fourth source as a circular area with a 20 m diameter, as this largest pond at the facility is likely to be a diffuse source. Further specifics of the inversion parameters are included in the Supplemental Information (S4.4).



**Figure 6.** An optimized Gaussian plume model for the first transect recorded on 2020-08-31 using the UGGA bicycle laboratory. The transect was along Leslie Street, the major roadway West of ABWWTP, capturing methane plumes advected by Easterly winds.

275 Our prior emission rate comes from scaling the ABWWTP annual emissions reported in the FLAME-GTA inventory (Pak et al., 2021). When modelling the plumes, we found that prescribing 1/3 of the prior emissions to source 1, and the remaining emissions equally, with 2/9 for sources 2, 3, and 4 provided a good first estimate to fit observed data under a variety of wind directions. We used this distribution of emissions as the prior emissions for our inversions.

To estimate emissions based on our observed concentrations, we used a standard Gaussian plume area inversion. These large  
280 plume crossings are significantly less affected by the smoothing of the observed UGGA methane concentrations as plume transect time is several multiples of the mean residence time of the instrument. An example of an optimized Gaussian plume model is shown in Figure 6. A summary of all of our transects is presented in Figure 7. The mean emissions as measured by the UGGA bicycle platform were  $1772 \pm 815 \text{ kg day}^{-1}$  (N=44), and  $1613 \pm 1156 \text{ kg day}^{-1}$  (N=14) for the ECCC Picarro mobile laboratory. For the estimates from each mobile platform, the uncertainty is the standard deviation of all the estimated  
285 emissions. This corresponds to an annual emission rate of  $647 \pm 297 \text{ t yr}^{-1}$  and  $589 \pm 422 \text{ t yr}^{-1}$ , respectively. The greater uncertainty in the results from the vehicle laboratory are likely due to the smaller number of samples, and greater range in the time-of-year that these transects were conducted. A further analysis of temporal variability in the Supplemental Information (S4.2), but we find that a lack of winter observations prevents us from robustly determining if there is any seasonal variability. We note that the range of emissions estimates between the two laboratories are quite similar, with average emissions rates from  
290 each lab agreeing within uncertainty. We used rural stability classes, informed by public historic weather data from ECCC



**Figure 7.** *Left:* Inversion results for all transects of ABWWTP. Bike data are shown with circle markers, and car data are triangles. The x axis denotes the number of days since the start of the calendar year. *Right:* Violin plots showing the relative distribution frequency of different estimated emissions rates from each mobile platform.

(Briggs, 1973). The inversion's sensitivity to the stability class, and other errors are discussed in the Supplemental Information (S4.3). Generally, the average inversion results are highly dependent on stability class, and we calculate a mean flux from the plant which is 70% greater when using urban stability classes. The average annual mean emissions calculated from the both laboratories agree within uncertainty with the FLAME-GTA inventory ( $787 \pm 352 \text{ t yr}^{-1}$ ) (Pak et al., 2021).

### 295 3.2 Combined Sewage Overflow Containment Basins

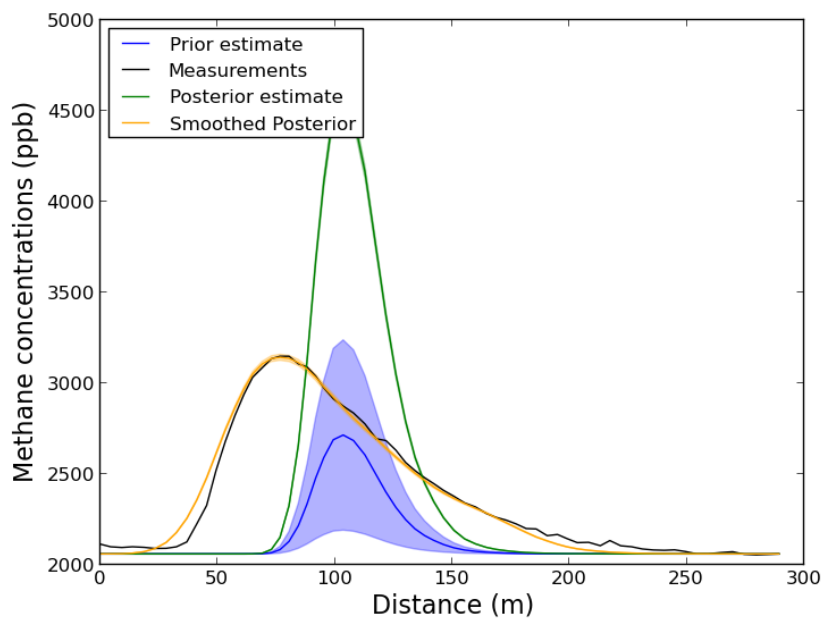
The older downtown part of Toronto has a combined storm-water and wastewater sewage system for collecting and treating urban street rainwater runoff and sewage. Originally constructed in the 1990's as a storage and primary treatment facility to prevent combined sewage overflow (CSO) spillage into Lake Ontario, the Western Beaches Tunnel (WBT) system is an 85 000 m<sup>3</sup> wastewater storage facility designed to contain CSO during high rainfall events. The city recently addressed the operational issues with the Western Beaches Tunnel system, including removing sediment build-up in the system (The City of Toronto, 2020). The locations of the three major storage shafts are shown in Figure 8. On our bicycle surveys along Toronto's waterfront, we observe persistent methane plumes from the CSO ventilation grates near Lake Ontario in three locations: the Glendale Ave storage shaft, the Cowan Ave storage shaft, and the Strachen storage shaft and Battery Park pumping station.

Unlike the plumes observed from the ABWWTP, transects for the CSO basin plumes are shorter, and considerably nearer to the emission source (18-100m). As such, we find that to compute an accurate emission rate, we cannot rely on enhancement height measurements. We use normal Gaussian plume area inversions, which are equivalent to our smoothed plume area inversions. An example of a typical plume transect and fitted model is shown in Figure 9.

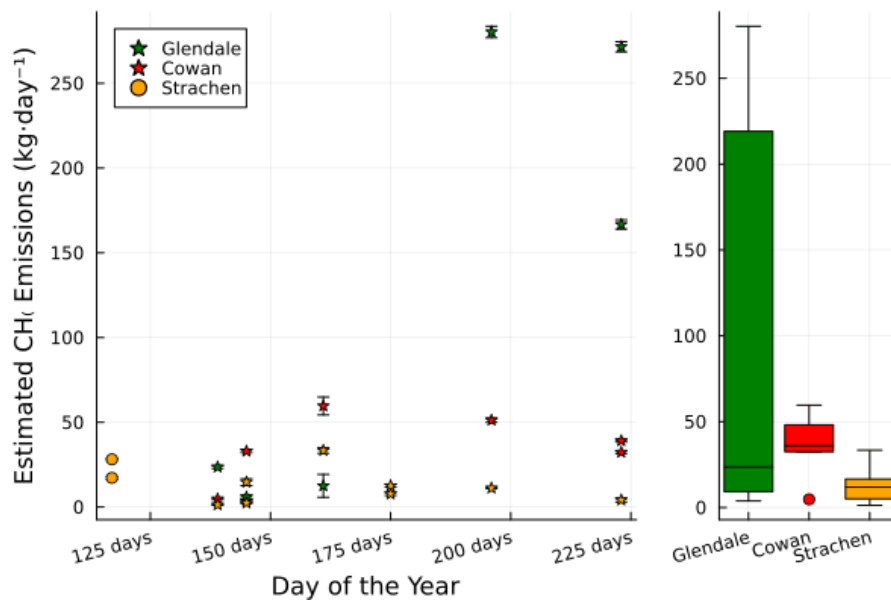
All CSO basin sources were modelled as a single ground-level point source. We chose the centroid of the ventilation grates as the source location at each site. The results of each inversion are shown in Figure 10. The mean daily estimate emission rates



**Figure 8.** The locations of the storage shafts in the Western Beaches Tunnel system. From Google Earth (2023).



**Figure 9.** Measured and modelled  $\text{CH}_4$  plumes from the Glendale storage shaft, from the 2021-08-11, 16:34 UTC transect. In this plot the prior emissions estimate is shown in blue, the normal Gaussian plume enhancement-area posterior is shown in green, and the smoothed Gaussian plume enhancement-area posterior is orange. Measured concentrations are shown in black.



**Figure 10.** *Left:* Inversion results for all transects of the waterfront CSO basins, with stars denoting data from 2021, and circles denote data from 2022. The x axis denotes the number of days since the start of the calendar year. *Right:* box and whisker plots showing the range of estimated emission rates from the basins.

310 are 30.7, 12.3, and 108 kg/day for the Cowen, Strachen, and Glendale storage shafts, respectively. We expect the emissions from the CSOs to be highly heterogeneous and to change seasonally, with significant emissions only during summer (3 months of the year) due to increased precipitation and higher water temperatures. Therefore we would expect annual emissions to be around 14 t CH<sub>4</sub> yr<sup>-1</sup>. These WBT emissions sources are not reported to the GHG reporting program, nor are they presently included in the FLAME-GTA inventory.

#### 315 4 Discussion

Overall in this paper, we present an emissions estimation methodology, tested against a controlled release experiment designed to mimic observations of small sources commonly encountered in the urban environment. The CSO basins are a good example of the small sources we observe regularly during our mobile surveys. As with our controlled release experiment, the inversion results for the CSO basins are consistent between the unaltered Gaussian plume area inversions, and the smoothed plume area inversions. The smoothed enhancement height inversions, where we explicitly account for the instrument response function, also agree well with the estimates from the area inversions. The inversions using the unaltered Gaussian plume enhancement heights are biased low compared with the others. By taking into account the instrument flow-rate with our instrument response function, we are able to reconcile the differences between our peak concentration enhancement inversions, and Gaussian plume area inversions for the LGR UGGA. We did not test our emissions quantification methodology with a larger controlled release

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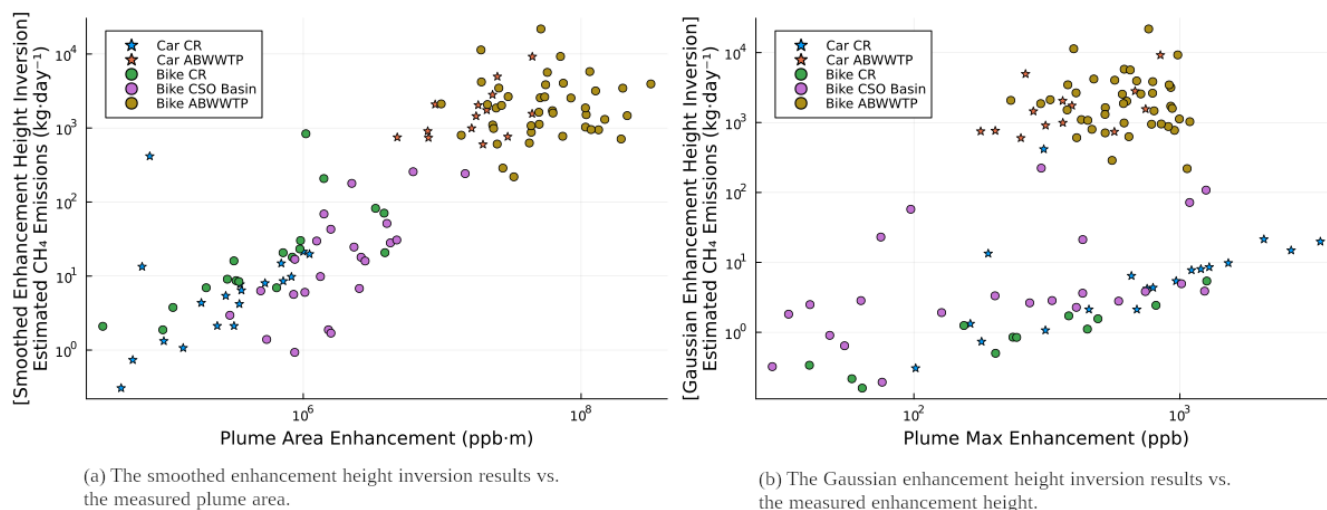
325 experiment, because we were limited by the available resources for conducting such CH<sub>4</sub> releases. However, we use ABWWTP  
as a large source with reasonably well-estimated emissions in the FLAME-GTA inventory (Pak et al., 2021), to also test our  
Gaussian plume inversion methodology on a larger scale.

In recent literature, studies by Von Fischer et al. (2017) and Weller et al. (2019) use a log-log statistical model to estimate  
emissions based on the peak height of the observed concentration. This algorithm does not explicitly account for the distance  
330 to the emissions source, nor does it take into account other factors, such as atmospheric stability or wind speed, because  
those parameters were not measured as part of their initial mobile measurement campaign which consisted of geo-located  
observations with a Picarro G2301 (Von Fischer et al., 2017). The Weller et al. (2019) method has been used in a number of  
studies to estimate fossil gas distribution fugitive emissions rates from mobile survey data (Weller et al., 2020; Ars et al., 2020;  
Defratyka et al., 2021; Maazallahi et al., 2020). The meta-analysis of Vogel et al. (2023a), which analyzed data from mobile  
335 observations in several cities using the exact same approach shows that data from our bicycle UGGA laboratory are the most  
significant outlier, recording a greater percentage of “small” emissions sources than any of the other urban mobile methane  
measurement campaigns. This is likely due to two causes. The first is that our peak heights are reduced by our instrument  
response function. The second is that the survey routes of the bike tend to be constrained to the City of Toronto, along routes  
with suitable cycling infrastructure, potentially missing the larger fossil gas leaks near major roadways (Vogel et al., 2023a).

340 When we synthesize the results from our controlled release experiment and our survey data, we find that plume area is a more  
reliable and consistent proxy for emission rates than peak enhancement concentration (Figures 11 a and b). This corroborates  
previous work suggesting the peak area enhancements are a more consistent metric for estimating and quantifying emissions  
rates, as opposed to a statistical relationship based on peak enhancement heights (Ars et al., 2017, 2020). We consider inversions  
optimized with a Gaussian plume model area to be the best option for quantifying instantaneous emissions rates from mobile  
345 transect data.

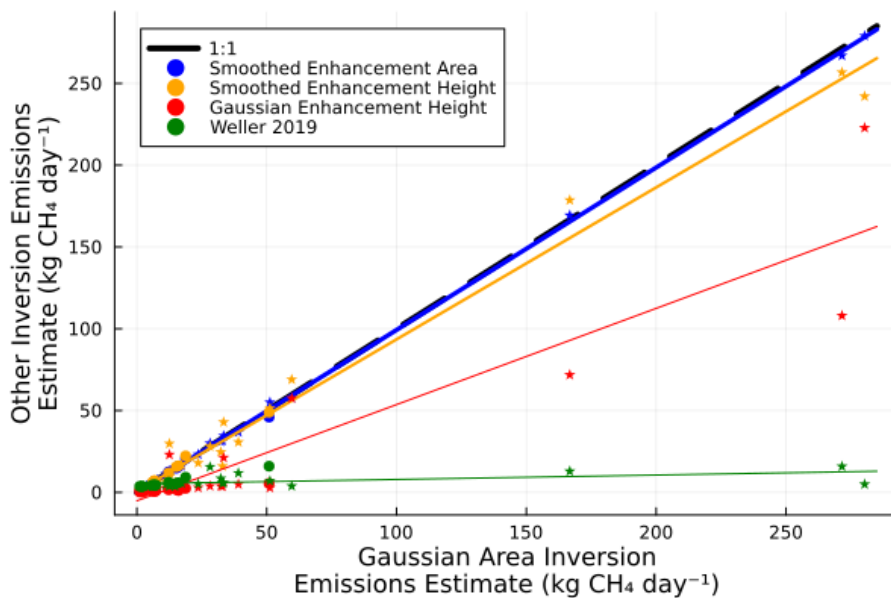
In Figure 12, we compare our estimated emissions rates using four different Gaussian plume inversions, and the statistical  
model for the CSO basin and controlled release experiment transects. There is excellent agreement between the enhancement  
area and the smoothed enhancement area optimizations. This is expected, as the smoothing operation conserves the observed  
total concentration enhancement area from the transect. There is a significant correlation between the area inversions and the  
350 smoothed enhancement height inversion for the small source bicycle transects. This suggests that accounting for the instrument  
response function for the LGR UGGA is necessary to correct the Gaussian plume enhancement height model for short, near-to-  
source transects. The normal Gaussian enhancement height inversion is biased low compared to the other methods, reflecting  
the impact of the asymmetrical smoothing of the lower flow-rate UGGA. The statistical relation from (Weller et al., 2019)  
is the most negatively biased with respect to the other results. Thus, we caution against the use of this statistical algorithm  
355 for quantifying individual sources, especially in cases where the gas concentration observations are affected by the response  
function of the gas analyzer.

We further demonstrate the comparability of the enhancement area as an emissions quantification metric by examining  
the normalized cumulative emissions relative to normalized emitter size for three different emissions metrics: the measured  
enhancement area, the measured enhancement height, and the Weller et al. (2019) statistical algorithm for the UGGA bicycle

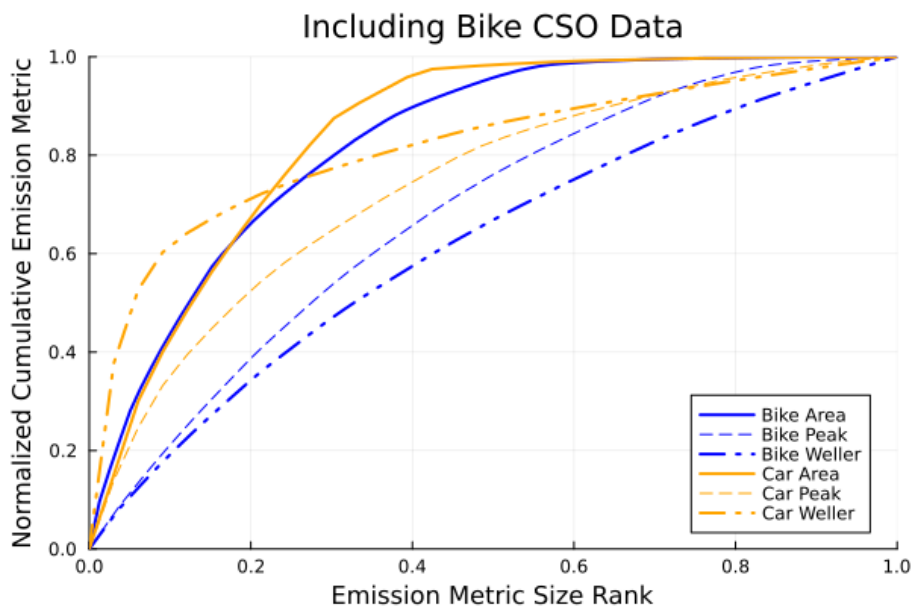


**Figure 11.** (a) Demonstrates that inversion results that are independent of the plume area—but which account for instrument response—are better correlated with the measured plume area than the measured peak enhancement heights, shown in (b). In (b) one inversion from the car ABWWTP dataset is omitted, and two inversions from the bike ABWWTP dataset are omitted for non-physical negative inversion optimizations. Note the logarithmic scales on both axes of the two plots.

360 laboratory and the ECCC Picarro vehicle laboratory for the plumes quantified in this study. We demonstrate in Figure 13 that the most consistent emissions metrics between the two mobile laboratories is the enhancement area. In Vogel et al. (2023a), which uses the Weller et al. (2019) algorithm to compare mobile observations across 12 different cities, the normalized cumulative emissions trend from our bicycle laboratory is an outlier from the rest of the mobile CH<sub>4</sub> measurement laboratories, with the first 17<sup>th</sup> percent of emitters measured were deemed responsible for 50% of observed emissions. For the other laboratories in  
365 that study, the first 8<sup>th</sup> percent or fewer sources were responsible for 50% of the observed cumulative emissions. In addition to the reasons stated in that publication for the differences between our bicycle laboratory and the others, this difference is also likely to be a consequence of the unsuitability of the Weller et al. (2019) methodology for the LGR UGGA instrument, due to the lower enhancements recorded as a result of the instrument response function. In Figure 13, we compute the cumulative emissions as a function of the ranked emission sources (ranked from highest emissions to lowest emissions). The cumulative  
370 emissions are computed for the bicycle and vehicle data when evaluating emissions using the peak area and enhancement height, and also using the Weller et al. (2019) method. Figure 13 demonstrates that the most significant differences between emissions metrics are between estimates using the Weller et al. (2019) technique.



**Figure 12.** Correlation plot comparing the results of various emission estimate calculations on the CSO (stars) and controlled release transects (circles) relative to the uncorrected Gaussian plume area inversion. Lines of best fit for each inversion, relative to the Gaussian Plume area inversion are shown.



**Figure 13.** Normalized cumulative emissions relative to normalized emitter size ranked from high to low for three emissions metrics.



## 5 Conclusions

In this work, we demonstrate the importance and applicability of accounting for the instrument response function when using Gaussian plume inversions to quantify emissions from mobile in situ observations. In the case of the low flow-rate LGR UGGA, convolution by an asymmetric smoothing window greatly improves the agreement between observed and modelled Gaussian plume enhancements, as well as minimizing differences between coincidental observations with other GHG analyzers. For estimating emissions rates, we find that the Gaussian enhancement area inversion is sufficient for both mobile laboratories and is robust to flow-rate issues, as long as the entire plume is sampled. Like previous studies, we demonstrate that the plume enhancement area provides a more consistent observation metric between different experimental setups. Furthermore, we demonstrate that enhancement area inversions and smoothed Gaussian plumes are practically equal to one another. We also demonstrate how approximate sources can be inferred using concentration vector triangulation techniques from downwind measurements.

We utilize the Gaussian enhancement area inversion technique to quantify emissions from the largest WWTP in the City of Toronto. From this analysis, we can demonstrate that the FLAME-GTA inventory estimate for CH<sub>4</sub> emissions from ABWWTP is consistent with our estimates from our mobile plume optimizations. In Canada's GHG reporting program, CH<sub>4</sub> emissions from municipal wastewater treatment plants are not required to be reported. Considering the relative CO<sub>2eq</sub> of the methane emissions we estimate from our mobile observations, we argue that such emissions should be included in the GHG reporting requirements. For facilities that capture and utilize the biogas produced, we suggest that estimates derived from the volume of gas produced, like the estimates in the FLAME-GTA inventory, are a reasonable first guess of the facility's methane emissions for municipal WWTP's. The relevant formula is presented again in the Supplemental Information (S4.2). Since the FLAME-GTA inventory emissions estimates for ABWWTP agree within error of our emissions estimates from both of our mobile laboratories, we do not suggest an update to the FLAME-GTA inventory for this facility, but we do suggest adding the Western Beaches Tunnel system emissions to the FLAME-GTA inventory. If we take our inversion results to represent the facility's average emissions, the system could represent a combined source over 55 t-yr<sup>-1</sup>. However, this is likely an overestimate, as we expect some seasonality in the emissions pattern due to temperature. We recommend additional targeted measurements at the CSO basins throughout the year to determine if the completed construction work has influenced methane emissions from the system.

Importantly, this study also demonstrates how mobile observations from different experimental platforms can be readily compared against one another by accounting for the residence time distribution of analyzers with slower flow-rates and longer mean residence times, such as the LGR UGGA. In future studies, the instrument response function could be more rigorously evaluated for different experimental setups. We suggest that the response function could be determined by a deconvolution of the observed gas concentrations, subjected to a known concentration perturbation, such as a rectangular pulse created with timed electronic valves. This would allow researchers to rigorously evaluate differences between different gas analyzers used for mobile measurement studies. Short of such analysis, we find that using enhancement areas is a suitable technique for comparing observations between different gas analyzers and mobile platforms.



Overall, in this study we demonstrate that estimating small emission rates with Bayesian optimization of Gaussian enhance-  
ment areas for transects close to the source is a useful tool for constraining the magnitude of methane emissions from mobile  
in situ GHG concentration data. Our controlled release experiment demonstrated that repeated measurements are essential for  
410 constraining the source emission rate, and that estimated instantaneous emission rates can vary widely between sequential  
plume transects. We conclude that Gaussian plume inversions are a useful tool to quantify emissions from mobile in situ CH<sub>4</sub>  
measurements, and that inversions should be calculated either using the plume area, or with explicit modelling of the instrument  
response function for inversions using the enhancement height. We demonstrate that estimating emissions using concentration  
enhancement areas is a more robust and comparable method of estimating emissions between various mobile laboratories. We  
415 recognize that additional work is needed to collect and analyze controlled release experiments from other platforms and groups  
globally. This work to develop standard methodologies for use by academic, non-governmental organizations, and private  
sector providers will allow for a more straightforward comparison between emissions estimates from different groups.

*Data availability.* Data from the bicycle lab are openly available from the on the University of Toronto achieve (Wunch et al., 2018; Ars  
et al., 2019; Gillespie et al., 2022a, b, 2023). The ECCC vehicle data are available from a web achieve (Vogel et al., 2023b).

420 *Author contributions.* DW, SA, FV, and LG formulated the research aim for this study. DW and SA assisted with data curation. LG completed  
the formal analysis for this study. DW and FV acquired the funding for this study. LG, SA, JW, LK, TF, MR, MK, AM, and SG completed  
the research investigation. SA and LG devised the research methodology. DW, FV, and MK supervised and guided the experimental work,  
provided resources for the project, and administered the research project. LG and SA and worked on the software and data visualization for  
this publication. LG prepared the manuscript with contributions from all of the co-authors. All authors contributed to the review and editing  
425 of the study.

*Competing interests.* The authors declare that they have no conflict of interest.

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