



1	Multi-angle aerosol optical depth retrieval method based on
2	improved surface reflectance
3	Lijuan Chen ¹ , Ren Wang ¹ , Ying Fei ² , Peng Fang ² , Yong Zha ² , Haishan Chen ^{1*}
4	¹ Key Laboratory of Meteorological Disaster, Ministry of Education (KLME)/Joint International
5	Research Laboratory of Climate and Environment Change (ILCEC)/Collaborative Innovation
6	Center on Forecast and Evaluation of Meteorological Disasters (CIC-FEMD), Nanjing University
7	of Information Science and Technology, Nanjing 210044, China
8	² Key Laboratory of Virtual Geographic Environment of Ministry of Education, Jiangsu Center for
9	Collaborative Innovation in Geographical Information Resource Development and Application,
10	College of Geographic Science, Nanjing Normal University, Nanjing 210023, China
11	Correspondence: Haishan Chen (haishan@nuist.edu.cn)

12 Abstract

13 Retrieval of terrestrial aerosol optical depth (AOD) has been a challenge for satellite Earth 14 observations, mainly due to the difficulty of estimating surface reflectance caused by 15 land-atmosphere coupling. Current satellite AOD retrieval products have low spatial resolution 16 under complex surface processes. In this study, based on our previous studies of AOD retrieval, 17 we further improved the estimation method of surface reflectance by establishing an error 18 correction model and then obtained a more accurate AOD. A lookup table is constructed using the 19 Second Simulation of Satellite Signal in the Solar Spectrum (6S) to obtain high-precision retrieval 20 of AOD. The retrieval accuracy of the algorithm is verified by AERONET (Aerosol Robotic 21 Network) observations. The results indicate that the retrieved AOD based on the improved method





22	of this study has advantages in fewer missing AOD pixels and finer spatial resolution, as
23	compared to the MODIS AOD product and our previous estimation method. Among the nine
24	MISR angles, the optimal correlation coefficient (R) of retrieved AOD and observed AOD can
25	reach 0.89. Root mean square error (RMSE) and relative mean bias (RMB) can reach a minimum
26	values of 0.20 and 0.32, respectively. This study will help to further improve the accuracy of
27	retrieving multi-angle AOD at large spatial scales and long time series.
28	
29	Keywords: surface reflectance; aerosol optical depth; satellite remote sensing; MISR; MODIS
30	1. Introduction

31 Aerosols are liquid or solid particles suspended in the atmosphere, with particle diameters 32 ranging from approximately 0.001 to 100 µm (Giles et al., 2019). Aerosols have a large impact on 33 the Earth's radiation budget balance and the uncertainties are difficult to estimate (Holben et al., 34 2001; Li et al., 2020; Berhane et al., 2021; Sun et al., 2022), thus direct and indirect effects of 35 aerosols have received widespread attention in the study of climate change mechanisms (Hatzianastassiou et al., 2009; Dao et al., 2014; Daniel et al., 2014; Samset et al., 2018; Li et al., 36 37 2018; Huang et al., 2021). In addition, concentrations of aerosols may pose a serious threat to human health (Lee et al., 2010; Dehghani et al., 2012; Mironova et al., 2015). The parameters of 38 39 the optical properties of aerosols include aerosol optical depth (AOD), scattering phase function, single scattering albedo and absorbing optical depth, etc. As an important parameter, AOD is 40 41 defined as the integral of aerosol extinction coefficient in the vertical direction. AOD describes the 42 attenuation effect of aerosols on light and also reflects an important indicator of the degree of air 43 pollution. Over the past two decades, multi-channel spectrometers carried by multiple





44	geostationary and polar orbit satellites have been used for AOD retrieval. The AOD products
45	obtained from satellite retrieval are widely used in the study of atmospheric environment
46	(Kaufman et al, 1997; Xie et al., 2019; Chen et al., 2021). Although the retrieval accuracy of AOD
47	is constantly improving, there is still a lot of room for improvement in the retrieval results over
48	land.
49	Scholars have conducted studies using multi angle sensors. Flowerdew et al. (1996) utilized
50	Along Track Scanning Radiometer 2 (ATSR-2) dual angle observation data, based on the
51	approximate condition of minimum variation of surface reflectance with wavelength, and using
52	the assumption of independent invariance of ground features and Lambertian bodies, simulated
53	using a bidirectional reflection radiation transfer model, and proposed a dual angle algorithm
54	(ATSR-DV) to invert AOD over land. Kokhanovsky et al. (2009) used the ATSR-DV algorithm to
55	invert the AOD over Germany on October 13, 2005, and compared the retrieval results with
56	MEdium Resolution Imaging Spectrometer (MERIS) and MISR products, indicating that the
57	ATSR-2 algorithm is also suitable for Advanced Along Track Scanning Radiometer (AATSR).
58	Sundstrom et al. (2012) obtained an aerosol model of eastern China based on Aerosol Robotic
59	Network (AERONET) observation data, and used the ATSR-DV algorithm to retrieve the
60	proportion of AOD and coarse to fine particles from AATSR data. Abdou et al. (2005) compared
61	the MISR AOD and the Moderate-resolution Imaging Spectroradiometer (MODIS) AOD products
62	carried by Terra using data from 62 AERONET observation sites. The results showed that over
63	land, the MODIS AOD in the 0.470 um and 0.660 um channels was 35% and 10% higher than
64	MISR. In coastal and desert areas, the MODIS retrieval error was relatively large, while over the
65	ocean, in the 0.470 um and 0.660 um channels, the MISR was 0.1 and 0.05 higher than the





66	MODIS AOD value, respectively, mainly depends on the accuracy of radiometric calibration.
67	Martochik et al. (1997) proposed an algorithm for extracting aerosol optical parameters using
68	MISR multi angle observations. The results showed that in the presence of dense vegetation over
69	land, AOD was extracted using its low reflectivity and multi angle observations. If dense
70	vegetation did not exist, AOD and aerosol models were determined using the reflectance function
71	spectral contrast angle dependence relationship. As a new remote sensing tool, multi angle remote
72	sensing has the ability to provide aerosol characteristics such as optical depth, single scattering
73	albedo, and phase function with sufficient precision, which is more suitable for playing its unique
74	role in aerosol research than traditional single angle optical remote sensing (Dubovik et al., 2019).
75	Multi angle remote sensing retrieval of aerosol optical properties can utilize the angle information
76	contained in satellite signals to better separate the contributions of the surface and atmosphere,
77	making it suitable for some bright surfaces. This provides a new approach for AOD retrieval.
78	The surface reflectance measures the ability of land acquisition to absorb and reflect solar
79	radiation. The surface reflectance is relatively complex on land, and its contribution is received by
80	satellite detectors after atmospheric scattering and absorption. Satellite observations are obtained
81	as a coupling of the two, making it difficult to directly distinguish between surface reflectance and
82	
	atmospheric scattering. Therefore, simultaneous retrieval of atmospheric aerosols and surface
83	atmospheric scattering. Therefore, simultaneous retrieval of atmospheric aerosols and surface reflectance is the goal pursued by quantitative satellite remote sensing (Deuzé et al., 2001). In
83 84	atmospheric scattering. Therefore, simultaneous retrieval of atmospheric aerosols and surface reflectance is the goal pursued by quantitative satellite remote sensing (Deuzé et al., 2001). In optical remote sensing, the blue band has shorter wavelengths, relatively low surface reflectivity,
83 84 85	atmospheric scattering. Therefore, simultaneous retrieval of atmospheric aerosols and surface reflectance is the goal pursued by quantitative satellite remote sensing (Deuzé et al., 2001). In optical remote sensing, the blue band has shorter wavelengths, relatively low surface reflectivity, and there is relatively more reflection and scattering caused by the atmosphere. Therefore, AOD is
83 84 85 86	atmospheric scattering. Therefore, simultaneous retrieval of atmospheric aerosols and surface reflectance is the goal pursued by quantitative satellite remote sensing (Deuzé et al., 2001). In optical remote sensing, the blue band has shorter wavelengths, relatively low surface reflectivity, and there is relatively more reflection and scattering caused by the atmosphere. Therefore, AOD is generally retrieved through the blue band. In the process of AOD retrieval, the overestimation of





88	reflectance will lead to an overestimation of AOD. Separating atmospherically generated
89	reflectance and surface reflectance from apparent reflectance is one of the difficulties of AOD
90	retrieval (apparent reflectance is the reflectance at the top of the atmosphere). In general, signals
91	such as aerosols are weaker compared to surface signals (Dong et al., 2023). Previous studies have
92	shown that when using satellite remote sensing to retrieve AOD, an intercept error of 0.01 in
93	surface reflectance can result in an retrieval error of approximately 0.1 (Zhang et al., 2021).
94	Therefore, accurate estimates of surface reflectance are an important basis for aerosol retrieval.
95	A high-precision AOD product obtained from retrieval is of great significance for monitoring
96	changes in atmospheric pollution and providing decision-making for pollution control. Observing
97	the spatial distribution of AOD is very important for daily monitoring of air pollution. In addition,
98	aerosol particles can affect the energy balance between the land and the atmosphere by absorbing
99	and scattering solar radiation, thus affecting the global climate system. To further improve the
100	retrieval accuracy and resolution of AOD, this study uses data from nine camera angles in the blue
101	band of MISR L1B2T from 2016 to 2018 using an improved retrieval algorithm. Firstly, the study
102	analyzes the retrieval errors of the MISR AOD for nine camera angles before the improved
103	retrieval. Secondly, we established an error correction model to correct the estimated surface
104	reflectance of MISR, thereby improving the surface reflectance at 9 angles. The improved surface
105	reflectance retrieval is used to obtain the MISR AOD with high accuracy. Finally, the improved
106	AOD retrieval method is verified and its estimated results are compared with the previous
107	retrieval.





108 2. MISR, MODIS, and AERONET Data

109 2.1 MISR and MODIS data

110	The MISR sensor is manufactured by National Aeronautics and Space Administration
111	(NASA) Jet Propulsion Laboratory (JPL) in the U.S. The MISR sensor consists of nine cameras,
112	each fixed at a specific angle of view along the orbital direction. The MISR has four bands (Blue:
113	446 nm, Green: 558 nm, Red: 672 nm and Near InfraRed: 866 nm) and nine angle (Table S1). The
114	MISR is capable of imaging the region almost simultaneously by all cameras within 7 min (Diner
115	et al., 1998; Martonchik et al., 2002; Kahn et al., 2007). 36 channels of MISR data are included,
116	all of which can be retrieved for AOD. Typically, medium angles are used for surface observations
117	and large angles are more sensitive to the effects of cloud cover and atmospheric aerosols. The
118	MISR product (MIL2ASAE_3) format is .nc, and the MISR format is HDF-EOS. The data used in
119	this study are MISR Level 1B2 Terrain Projected Data (MI1B2T) data. The projection of the
120	MI1B2T data is the Space Oblique Mercator (SOM), which uses Hierarchical Data Format - Earth
121	Observing System (HDF-EOS) strip record data, and out of 233 paths, each of the 233 orbits
122	consists of 180 mutually independent blocks (Kahn et al., 2005). The study extracted 64 and 65
123	blocks of data covering the Yangtze River Delta region (Fig. S1). MISR data cannot be directly
124	processed using Arcgis and ENVI due to its special storage method. The study uses HDF-EOS To
125	GeoTIFF Conversion Tool (HEGTool) for batch processing of MI1B2T and MI1B2GEOP. Using
126	the HEGTool, the radiation data in the MI1B2T dataset is extracted, and as the radiation data
127	contains 180 blocks, the corresponding blocks, output data types and projections are selected
128	based on the area locations. Solar zenith angle, solar azimuth angle, satellite zenith angle for 9
129	cameras and satellite azimuth angle for 9 cameras data were extracted from the MI1B2GEOP, and





130	the corresponding blocks, output data types and projection information were selected. The TIFF
131	data output by the software is only then available for the next step of processing using ARCGIS
132	and ENVI. For the downloaded MISR data, remote sensing images with cloud pixels less than
133	50% are used for cloud detection and removal of cloud pixels. This study using blue bands to set
134	thresholds to remove cloud pixels. After repeated experiments, if a fixed threshold is used, cloud
135	pixels cannot be removed cleanly from the 9 observation angles of MISR (Fig. S2, Fig. S3).
136	Therefore, this study uses dynamic threshold method to remove MISR cloud pixels from images.
137	The dataset used in the study is shown in table S2.
138	MODIS L1B data are Earth observation data stored in a hierarchical (HDF) format ,
139	providing MOD02QKM, MOD02HKM and MOD021KM data respectively (Hong et al., 2007;
140	Bandaru et al., 2013; Wong et al., 2020). MODIS geolocation data (MOD03/MYD03) contain the
141	MODIS solar/satellite zenith angle for each 1 km EV (Earth View) centre, the latitude, longitude,
142	solar/satellite azimuth and land/sea thresholds. To obtain the MODIS L1B apparent reflectance,
143	MOD03/MYD03 solar zenith angle data are also used. The MODIS L1B data used for the study
144	consisted of radiometric data (MOD02/MYD02) and geolocation data (MOD03/MYD03). The
145	MODIS L1B data were pre-processed with the corresponding MOD03 data for geometric
146	correction. The MODIS Conversion Toolkit (MCTK) was used to achieve radiometric calibration,
147	"bowtie" processing, geometric correction, reprojection and band extraction of MODIS data.
148	MODIS sensors can observe the surface at zenith angles up to approximately 65.5°, and repeated
149	observations of the same surface image over multiple days can be obtained for different angles
150	(-65.5° to 65.5°) of this image. Assuming that the surface of this image does not change
151	significantly during this time, this set of multi-angle observations can be used for the BRDF





152	(Bidirectional Reflectance Distribution Function) model retrieval. MODIS BRDF/Albedo is a
153	standard terrestrial level 3 product, generated from data acquired by the Terra and Aqua satellite
154	platforms and MODIS. This product has a 16-day cycle, with observations on day 9 of the 16-day
155	retrieval cycle being assigned a weight to obtain daily data, which is the global surface albedo
156	daily product data (Hsu et al., 2004). The core dataset of the MODIS BRDF product is
157	MCD43A1.

158 2.2 AERONET data

159 AERONET uses a French produced solar radiometer CE-318 instrument to obtain solar direct spectral radiation measurements at 340nm, 380nm, 440nm, 500nm, 670nm, 870nm, 936nm, 160 161 1020nm, and 1640nm channels every 3 minutes at a field of view angle of 1.5°. The total 162 atmospheric water vapour content can be obtained from the 936 nm channel measurements and the 163 AOD values can be retrieval using the remaining channel measurements with an retrieval error of 164 about 0.01-0.02. Therefore, it can provide aerosol characterization parameters with high accuracy 165 and validate the aerosol parameters from satellite retrievals (Lu et al., 2019). AERONET has more 166 than 600 observing sites globally distributed over land and ocean, using a sun photometer as the 167 basic observing instrument, and most of the sites achieve daily data acquisition and unified 168 transmission to the network for centralized processing. It plays an important role in studying 169 global aerosol radiation effects, aerosol transport, validating radiative transfer models and 170 verifying satellite remote sensing aerosol results. The study therefore examines the accuracy of the 171 satellite remote sensing retrieval of AOD using the AOD measured by AERONET as the true 172 value.

173 AERONET provides observations of AOD, retrieval products and precipitable water





174	distributed over a wide range of aerosol patterns worldwide. There are three quality levels of AOD
175	data. Level 1.0 (unscreened), Level 1.5 (cloud-screened and quality-controlled), and Level 2.0
176	(quality-assured). The study area is mainly the Yangtze River Delta region of China, where
177	AERONET has a large number of sites, but only the Taihu and Xuzhou-CUMT sites provide
178	continuous data, while the rest of the sites have been acquiring data for a relatively short period of
179	time.
180	The amount of available data for AERONET AOD Level 2 is relatively small. To ensure
181	sufficient ground site validation data, this study selected Level 1.5 data with a large and
182	continuous amount of current observation data to verify the satellite remote sensing AOD obtained
183	from retrieval (Dubovik et al., 2000; Li et al., 2009) . In terms of time, the AERONET AOD Level
184	1.5 data selected for this study correspond to the MISR data for the three years from 2016-2018,

186 3. Methodology

187 **3.1** Problems in the previous surface reflectance estimation method

Accurate estimation of the contribution of surface reflectance has been the focus and difficulty in the process of remote sensing retrieval of AOD (Remer et al., 2009; Gupta et al., 2016). Previous study used MODIS BRDF data and MODIS V5.2 Algorithm to determine the MODIS surface reflectance (Chen et al., 2021), and after spectral conversion between the two sensors, MODIS and MISR, the MISR surface reflectance was obtained for nine angles. A look-up table was constructed using the Second Simulation of Satellite Signal in the Solar Spectrum (6S), and then the MISR surface reflectance of the 9 angles was combined to obtain the MISR AOD of





- 195 different angles by retrieval. This study found the variation pattern of the MISR AOD of the 9
- 196 angles. However, the error of the 9-angle retrieved MISR AOD was relatively large compared to
- 197 the AERONET AOD (Table S3).

198 In 6S model, a series of parameters related to the simulated imaging date atmospheric 199 conditions need to be input, including geometric parameters, AOD, water vapor, ozone, elevation, 200 etc. At the same time, these parameters are inputted through a simple and intuitive lookup table. 201 Finally, a linear atmospheric correction formula is generated to obtain the surface reflectance 202 values after 6S atmospheric correction for each pixel in each band one by one. The 6S model 203 inputs the geometric parameter information corresponding to the MISR image at the the Taihu and 204 Xuzhou-CUMT sites, and the input of 550nm AOD parameters is the AERONET AOD of the two 205 sites. Then, the 6S model is used to obtain the MISR atmospheric corrected reflectance. To 206 analyze the overall high AOD values retrieved from 9 angles of MISR, this study compared the 207 MISR atmospheric corrected reflectance at that pixel position with the MISR surface reflectance 208 (Fig. 1) (The calculation method for MISR surface reflectance refers to chen et al. (2021)). It can 209 be seen that the MISR surface reflectance is low compared to the value of MISR atmospheric 210 corrected reflectance at the corresponding locations of the two sites, resulting in a higher retrieve 211 MISR AOD compared to the AERONET AOD. Therefore, it is necessary to establish a correction 212 model to correct the MISR surface reflectance to improve the retrieval accuracy of the MISR 213 AOD.







Figure 1. Comparison of MISR surface reflectance with atmospheric corrected reflectance in the
blue band (At the pixel locations of Taihu and Xuzhou-CUMT sites).

216 **3.2 Improved surface reflectance estimation method**

- 217 In order to develop a correction model to improve the surface reflectance, the design scheme
- 218 of this study is shown below:

a) Atmospheric correction of MODIS L1B using the 6S model to obtain MODIS atmospheric

220 corrected reflectance;

b) The new estimated MISR surface reflectance based on the MODIS atmospheric correction
was calculated by bringing the MODIS atmospheric correction reflectance into Eq. 1 and Eq. 2.
The MISR surface reflectance was combined with the newly estimated MISR surface reflectance
and a regression was fitted (with 60% of the overall sample data randomly selected) to create a
surface reflectance error correction model, as shown in the following formula:

226

$$\rho(\theta_s, \theta_v, \phi)_{MISR_a} = \rho(\theta_s, \theta_v, \phi)_{MODIS_at} \times \frac{BRDF(\theta_s, \theta_v, \phi)_{MISR}}{BRDF(\theta_s, \theta_v, \phi)_{MODIS}}$$
(1)

In Eq. (1), $BRDF(\theta_s, \theta_v, \phi)_{MISR}$, $BRDF(\theta_s, \theta_v, \phi)_{MODIS}$ are BRDFs obtained at MISR and MODIS angles, respectively. θ_s is the solar zenith angle, θ_v is the satellite zenith angle, and ϕ is the relative azimuth angle. $\rho(\theta_s, \theta_v, \phi)_{MISR \ a}$ is the surface reflectance of MODIS at the





230	geometric observation angle of MISR, and $\rho(\theta_s, \theta_v, \phi)_{MODIS_{at}}$ is the MODIS atmospheric								
231	corrected reflectance.								
232	This study selected spectral data containing 28 typical features of different types of								
233	vegetation, soil and water bodies from five standard spectral libraries that come with the ENV								
234	software. Calculate the surface reflectance of different features in the blue bands of MODIS an								
235	MISR using formulas (Chen et al., 2021).								
236	$\rho(\theta_s, \theta_v, \phi)_{MISR} = \rho(\theta_s, \theta_v, \phi)_{MISR_a} \times 0.9834 - 0.0081 $ (2)								
237	The New MODIS surface reflectance ($ ho(heta_s, heta_v,\phi)_{MISR_a}$) at the MISR angle obtained from								
238	Eq. (1) is converted to the MISR surface reflectance by Eq. (2).								
239	c) The MISR surface reflectance estimated by Eq. 2 is transformed by an error correction								
240	model to obtain the final improved MISR surface reflectance. The improved MISR surface								
241	reflectance will be used in the retrieval of the AOD. The MISR correction model was developed								
242	by fitting a linear regression of the previously estimated MISR surface reflectance based on the								
243	MODIS V5.2 algorithm to the MISR surface reflectance estimated based on the MODIS								
244	atmospheric correction (60% of the data were randomly selected) as shown in Eq. 3. The								
245	previously estimated 9-angle MISR surface reflectance was error-corrected by Eq. 3 to obtain the								
246	improved surface reflectance for the 9 angles of the MISR sensor, which was ultimately used to								
247	perform the MISR AOD retrieval for the 9 angles.								
248									

$$\rho(\theta_s, \theta_v, \phi)_{MISR-b}^* = \rho(\theta_s, \theta_v, \phi)_{MISR} \times 0.9209 + 0.0409 \tag{3}$$

249 Where ρ_{MISR-b}^{*} is the improved MISR surface reflectance in Eq. (3).





250 **3.3 Flow of improved multi-angle AOD retrieval**

251	The flow of the improved surface reflectance algorithm for this study is shown in Fig. 2. The
252	MODISL1B data were first atmospherically corrected using 6S, and then the MISR surface
253	reflectance estimated from previous study was combined with the new MISR surface reflectance
254	estimated from Eq. 2 to build a MISR error correction model to obtain the improved MISR surface
255	reflectance (Chen et al., 2021). The study retrieved the MISR AOD for nine camera angles using
256	improved MISR surface reflectance. We use AERONET AOD to validate the improved MISR
257	AOD. Compare the improved AOD with the previously retrieved AOD, and analyze the accuracy
258	and spatial distribution trend of the improved AOD. The AOD retrieval method used in this study
259	is based on chen et al. (2021). The selection of appropriate aerosol type is very critical for the
260	retrieval of aerosol optical depth. It has been shown that continental aerosols can be used for AOD
261	retrieval in the Yangtze River Delta (He et al., 2015). Therefore, we used continental aerosols in
262	our aerosol retrieval for the study area.









265 **4. Results and discussion**

264

266 4.1 Improved MISR surface reflectance variation characteristics

The estimated MISR surface reflectance, the MISR atmospheric corrected reflectance and the improved MISR surface reflectance are presented in the Fig. 3. This is the average of all sample data at the corresponding locations at the two sites in Taihu and Xuzhou-CUMT for the valid dates of 2016-2018. It can be noted that at the two site of Taihu and Xuzhou-CUMT, the 9 camera angle MISR-improved surface reflectance values are overall higher than the MISR surface reflectance and lower than the MISR atmospheric corrected reflectance. The nine camera angle MISR surface reflectance values ranged from 0.02 to 0.04. The improved surface reflectance averages were







274 overall greater than the previously estimated MISR surface reflectance.



To clarify the trend of the improved surface reflectance, the study performed a time-series analysis of the MISR surface reflectance and the improved surface reflectance (Fig. 4). It can be seen that the improved MISR surface reflectances are all higher than the previously estimated MISR surface reflectances. MISR surface reflectance values generally range from 0-0.05, with improved surface reflectance values ranging from approximately 0.05-0.1. The improved surface reflectance values have increased overall.









Figure 4. Surface reflectance time series of MISR sensors in the blue band at 9 observation angles.

285

(The order of time from front to back for 9 angles is shown in Table S2)





286 4.2 Results of the Improved MISR AOD retrieval

287 MISR AOD for 2016-2018 were obtained using an improved surface reflectance retrieval. 288 The study presents retrieval results from nine camera observation angles of the MISR sensor on 12 289 June 2018 (Fig. 5). As can be seen from the spatial distribution of AOD, the retrieval results in the 290 study area do not exceed a value of 1. The overall spatial distribution trend is generally consistent 291 with the results before the improvement (Chen et al., 2021), but differs in the magnitude of the 292 values. Values in the north-eastern and southern regions range from 0.5 to 1, indicating to some 293 extent that the air quality in this region is poor. The five camera observation angles, Ba, Aa, An, 294 Af and Bf, retrieve AOD values in the approximate range of 0.25-0.5. In the central region, the 295 four camera observation angle values of Da, Ca, Cf and Df were mostly in the range of 0-0.25. 296 The values indicate that the air quality in the region is generally good, but there are some areas of 297 light air pollution. The higher AOD in the southern part of Shandong Province and the northern 298 part of Jiangsu Province may be due to increased local aerosol emissions caused by human 299 activities.









Figure 5. Plot of AOD 550nm retrieval results for the improved MISR 9 camera angles on 12
 June 2018.

The study validates our improved MISR AOD spatial distribution results by comparing with MODIS AOD products of the same date (Fig. 6). The MODIS AOD products have a resolution of 304 3km. It can be seen that the trend of spatial distribution of MODIS AOD products is consistent 305 with the improved MISR AOD. However, the MODIS AOD product has more missing data, 306 which can be avoided by the AOD obtained from the retrieval of the improved algorithm, and the 307 AOD retrieval by the improved algorithm has a higher resolution by comparing with the image 308 quality of the MISR AOD product.





Figure 6. MODIS AOD 550nm product spatial distribution on June 12, 2018

324





310 4.3 Verification of the improved MISR AOD

311 There are many AERONET sites in the Yangtze River Delta region, but so far, only the Taihu and Xuzhou-CUMT sites continue to provide data, and other sites have a short time to 312 313 obtain data. Therefore, the Taihu and Xuzhou-CUMT sites with more data are selected for verification. To verify the retrieved MISR AOD, In terms of time, we selected effective AOD 314 315 records in the 550 nm band within a 30 minute interval between the AERONET ground 316 observation site and the Terra satellite. The 9 camera views of MISR require about 7 minutes to 317 observe the same geographical location, with relatively short intervals. Therefore, we will use the 318 calculated AERONET AOD average as the approximate truth value, and compare the average 319 value with the retrieved MISR AOD to verify and reduce errors caused by time difference. In terms of space, we selected pixels observed by MISR sensors from 9 angles and compared them 320 321 with the nearest data observed by AERONET, which can reduce errors caused by spatial 322 differences. The solar photometer does not have a 550 nm wavelength that corresponds to the 323 retrieval results, and the AOD at 550 nm can be calculated by applying Angstrom (Eq. 4).

$$\tau(\lambda) = \beta \lambda^{-\alpha} \tag{4}$$

³²⁵ In the formula, $\tau(\lambda)$ is the AOD at wavelength λ , β is the concentration of the entire ³²⁶ atmospheric aerosol, and α is the wavelength index of Angstrom.

In this study, four parameters will be used to assess the accuracy of the remotely sensed AOD
dataset, namely the correlation coefficient (R), the Root Mean Square Error (RMSE), p-value and
the relative mean bias (RMB). The specific calculation principles for the three parameters R,
RMSE and RMB are shown in Eq. (5)-(7). The validation results of this study's improved AOD
dataset from 2016-2018 at Taihu and Xuzhou-CUMT sites are shown in Fig. 7 and Fig. 8.

In general, the scatter plot is distributed above and below the 1:1 line. R is a parameter used to characterize the correlation between the remote sensing retrieval results and the ground-based retrieval results. At the Taihu site, R reach up to 0.89. At the Xuzhou-CUMT site, R reach up to 0.85. The RMSE is a parameter used to characterise the absolute error of the remote sensing retrieval results, with a minimum root mean square error of 0.21 at the Taihu site and 0.20 at the





Xuzhou-CUMT site. RMB is the parameter used to characterise the relative error of the remote
sensing retrieval results, with a minimum RMB of 0.52 at the Taihu site and 0.32 at the
Xuzhou-CUMT site. In summary, by comparing the results with the validation of the AOD scatter
plot before the improvement, the accuracy of the nine camera observation angles at both sites has
improved after the improvement (Table 1).

342
$$R = \frac{\sum_{i=1}^{N} (A_i - \overline{A})(A_i' - \overline{A'})}{\sqrt{\sum_{i=1}^{N} (A_i - \overline{A})^2 \sum_{i=1}^{N} (A_i' - \overline{A'})^2}}$$
(5)

343
$$RMSE = \sqrt{\sum_{i=1}^{N} (A_i - A_i')^2 / N}$$
(6)

344
$$RMB = \sum_{i=1}^{N} (A_i - A_i') / N$$
(7)

345 where A_i is the retrieve MISR AOD, A'_i is the corresponding AERONET AOD, \overline{A} 346 and $\overline{A'}$ are the mean values of the retrieve MISR AOD and AERONET AOD, respectively. *N* is 347 the number of valid matching results for AERONET AOD and MISR AOD.









348 Figure 7. Comparison between improved MISR AOD and AERONET AOD at Taihu site (N is

349

the number of verification points, red line represents a linear fitting line).









350 Figure 8. Comparison between improved MISR AOD and AERONET AOD at Xuzhou-CUMT

351

site (N is the number of verification points, red line represents a linear fitting line).







352 Figure 9. Comparison of validation of retrieved AOD with MODIS AOD product



354

Table 1. Precision comparison of MISR AOD and AERONET AOD before

and after improvement

Site	Angle	R	RMSE	RMB	Improved R	Improved RMSE	Improved RMB	
	Da	0.77	0.20	1.08	0.80	0.40	0.69	
	Ca	0.70	0.29	1.00	0.72	0.42	0.74	
	Ba	0.77	0.14	0.61	0.80	0.37	0.70	
	Aa	0.81	0.11	0.68	0.82	0.36	0.71	
Taihu	An	0.45	0.29	1.22	0.84	0.21	0.52	
	Af	0.72	0.14	0.87	0.89	0.21	0.68	
	Bf	0.72	0.17	0.60	0.82	0.32	0.67	
	Cf	0.57	0.24	0.65	0.79	0.39	0.73	
	Df	0.77	0.20	0.47	0.79	0.40	0.66	
	Da	0.45	0.36	1.58	0.65	0.43	0.61	
	Ca	0.59	0.34	0.96	0.71	0.32	0.82	
	Ba	0.67	0.27	0.78	0.70	0.29	0.32	
	Aa	0.73	0.24	0.78	0.82	0.22	0.49	
Xuzhou- CUMT	An	0.75	0.20	0.85	0.85	0.20	0.47	
00111	Af	0.72	0.23	0.63	0.74	0.33	0.62	
	Bf	0.62	0.28	0.65	0.75	0.24	0.53	
	Cf	0.68	0.28	0.66	0.72	0.25	0.54	
	Df	0.67	0.30	0.65	0.72	0.26	0.53	

355

356 By comparing the validation results of MODIS AOD products with those of observation sites

357 (Taihu: R=0.59, RMSE=0.19, P<0.05, RMB=0.52; Xuzhou-CUMT: R=0.71, RMSE=0.25, P<0.05,

358 RMB=0.44) (Chen et al., 2021), we find that the improved MISR AOD has a higher correlation





359 with MODIS AOD products in the Taihu and Xuzhou-CUMT sites. The smaller the observation angle of the improved MISR AOD, the closer the error is to that of the MODIS AOD product. The 360 361 observation angle of MISR An is the same as that of MODIS. Therefore, we selected An 362 observation angle and MODIS AOD products at two pixel positions in the Taihu and 363 Xuzhou-CUMT for verification (Fig. 9). The results show that the An AOD retrieval by the 364 improved algorithm correlates well with the MODIS AOD product, and the position errors of the two image elements are close to each other. The RMSE of Xuzhou-CUMT site is slightly higher 365 366 than that of Taihu site, and the RMB of Taihu site is slightly higher than that of Xuzhou-CUMT 367 site.

368 5. Conclusion

369	This study first explored the problem of estimating the surface reflectance in our previous
370	study and then obtained an error correction model for the surface reflectance using a linear fit of
371	the MISR surface reflectance and a new estimate of the MISR surface reflectance. The improved
372	MISR surface reflectance was obtained by means of an error correction model. A new AOD
373	product was retrieved using the improved surface reflectance and a look-up table constructed from
374	6S model. Two AERONET ground observation sites with longer time series were used to validate
375	the AOD obtained by satellites.
376	(1) Overall, the improved AOD and its spatial distribution trends are consistent with our
377	previous results. The AOD estimated by our improved method presented a higher accuracy and a
378	high degree of agreement with the AERONET ground-based observational AOD.
379	(2) More importantly, compared to the MODIS AOD products, the retrieved AOD in this
380	study has fewer missing AOD pixels and finer spatial resolution. The retrievals of An Angle AOD
381	by the improved algorithm are highly correlated with the MODIS AOD products, as shown
382	through validation with the MODIS AOD product.





383	(3) In the future, more aerosol models conforming to the actual situation in the study area
384	can be constructed using the AERONET ground observation data and introduced into the MISR
385	AOD retrieval algorithm to further improve the accuracy of the AOD retrieval results. In this study,
386	the AERONET AOD was used as the true value and as an input to the AOD parameter in the 6S
387	model for atmospheric correction of MISR and MODIS images. A surface reflectivity error
388	correction model was then obtained to retrieve the AOD for the entire region. It should be
389	emphasized that the more AERONET sites used to train the corrected model, the more accurate
390	the AOD results obtained by this method. However, the data of AERONET ground observation
391	sites was limited. In the future, the study area can be expanded to large scale and longer time
392	series.

393 Acknowledgement

This study was supported by the Key Program of the National Natural Science Foundation of
China (42130609), Jiangsu Funding Program for Excellent Postdoctoral Talent (2023ZB482), the
Natural Science Foundation of Jiangsu Province of China (BK20220455), and the National
Science Foundation of China (42201028).

398 Competing interests

399 The contact author has declared that none of the authors has any competing interests.





400 Reference

- 401 Abdou, W. A.; Diner, D. J.; Martonchik J. V.; Bruegge C. J. Comparison of coincident Multiangle
- 402 Imaging Spectroradiometer and Moderate Resolution Imaging Spectroradiometer aerosol
- 403 optical depths over land and ocean scenes containing Aerosol Robotic Network sites. J.
- 404 Geophys Res-atmos. 2005, 110 (D10): 1-12.
- 405 Bandaru, V.; West, T. O.; Ricciuto, D. M.; Izaurralde, R. C. Estimating crop net primary
- 406 production using national inventory data and modis-derived parameters. Isprs. J Photogramm.
 407 2013, 80, 61-71.
- Berhane, S. A.; Bu, L. Aerosol-Cloud Interaction with Summer Precipitation over Major Cities in
 Eritrea[J]. Remote Sens-basel, 2021, 13(4):21.
- 410 Chen, L.; Wang, R.; Han J. Influence of observation angle change on satellite retrieval of aerosol
- 411 optical depth[J]. Tellus B, 2021, 73.
- 412 Chen, L.; Fei, Y.; Wang, R. Retrieval of high temporal resolution aerosol optical depth using the
- 413 GOCI remote sensing data[J]. Remote Sens-basel. 2021, 13:2376.
- 414 Chen, L.; Wang, R.; Wei, G. A surface reflectance correction model to improve the retrieval of
- 415 MISR aerosol optical depth supported by MODIS data[J]. Adv Space Res, 2021,
- 416 67(2):858-867.
- 417 Dao, Y.; Gong, W. Observed holiday aerosol reduction and temperature cooling over East Asia[J].
- 418 J Geophys Res-atmos, 2014.
- 419 Daniel, R.; Steven, S.; Robert, W.; Leo, D. Climate Effects of Aerosol-Cloud Interactions[J].
- 420 Science, 2014, 343, 379-380.
- 421 Dehghani, M.; Saeedi, A. A.; Zamanian, Z. A study of the relationship between indoor and outdoor
 422 particle concentrations in Hafez hospital in Shiraz[J]. Phys.Rev.C, 2012.
- 423 Diner, D. J.; Beckert, J. C.; Reilly, T. H.; Bruegge, C. J.; Conel, J. E. Multi-angle imaging
- 424 SpectroRadiometer (MISR) instrument description and experiment overview. IEEE Trans.
- 425 Geosci. Remote Sens. 1998, 36, 1072–1087.
- 426 Dong, W.; Tao, M.; Xu, X.; Wang, J.; Wang, Y.; Wang, L.; Song, Y.; Fan, M.; Chen, L. Satellite





- 427 Aerosol Retrieval From Multiangle Polarimetric Measurements: Information Content and
- 428 Uncertainty Analysis," IEEE Trans. Geosci. Remote Sens. 2023, 61, 1-13.
- 429 Dubovik, O.; Smirnov, A.; Holben, B. Accuracy assessments of aerosol optical properties retrieved
- 430 from Aerosol Robotic Network (AERONET) Sun and sky radiance measurements[J]. J
- 431 Geophys Res-atmos, 2000, 105(D8): 9791-9806.
- 432 Dubovik, O.; Li.; Mishchenko M. I. Polarimetric remote sensing of atmospheric aerosols:
- 433 Instruments, methodologies, results, and perspectives[J]. Pergamon, 2019, 474-511.
- 434 Deuzé, J. L.; Bréon, F. M.; Devaux, C.; Goloub, P.; Herman, M.; Lafrance, B.; Maignan, F.;
- 435 Marchand, A.; Nadal, F.; Perry, G.; Tanré, D. Remote sensing of aerosols over land surfaces
- 436 from POLDER-ADEOS-1 polarized measurements. J Geophys Res-atmos, 2001, 106(D5):
 437 4913-4926.
- 438 Flowerdew, R. J.; Haigh, J. D. Retrieval of aerosol optical thickness over land using the ATSR-2

439 Dual-look Satellite Radiometer. Geophys Res. Lett. 1996, 23 (4): 351-354.

- Giles, D. M.; Sinyuk, A.; Sorokin, M. G. Advancements in the Aerosol Robotic Network
 (AERONET) Version 3 database-automated near-real-time quality control algorithm with
- 442 improved cloud screening for Sun photometer aerosol optical depth (AOD) measurements[J].
- 443 Copernicus GmbH, 2019, 12(1), 169-209.
- Gupta, P.; Levy, R. C.; Mattoo, S. A surface reflectance scheme for retrieving aerosol optical depth
 over urban surfaces in MODIS dark target retrieval algorithm[J]. Atmos Meas Tech, 2016,
- 446 9(7): 3293-3308.
- Hatzianastassiou, N. The direct effect of aerosols on the radiation budget and climate of the
 Earth-atmosphere system: its variability in space and time. EGU General Assembly
 Conference Abstracts EGU General Assembly Conference Abstracts, 2009.
- 450 He, J.; Zha, Y.; Zhang, J.; Gao, J.; Li, Y.; & Chen, X. (2015). Retrieval of aerosol optical
- 451 thickness from hj-1 ccd data based on modis-derived surface reflectance[J]. International
- 452 Journal of Remote Sensing, 2015, 36(3-4), 882-898.
- 453 Holben, B. N.; Tanré, D.; Smirnov, A. An emerging ground-based aerosol climatology: Aerosol





454	optical depth from AERONET[J]. J Geophys Res, 2001, 106, 12067-12097.
455	Hong, G.; Yang, P.; Gao, B. C.; Baum, B. A.; Hu, Y. X.; King, M. D. High cloud properties from
456	three years of modis terra and aqua collection-4 data over the tropics. J Appl Meteorol Clim.
457	2007, 46(11): 1840-1856.
458	Huang X , Ding A .Aerosol as a critical factor causing forecast biases of air temperature in global
459	numerical weather prediction models[J].Science Bulletin,
460	2021.DOI:10.1016/j.scib.2021.05.009.
461	Hsu, N. C.; Tsay, S. C.; King, M. D. Aerosol properties over bright-reflecting source regions[J].
462	IEEE Trans. Geosci. Remote Sens, 2004, 42(3):557- 569.
463	Kahn, Gaitley, B. J.; Martonchik, J.; Diner, D.; Crean, K. Misr global aerosol optical depth
464	validation based on two years of coincident aeronet observations. J Geophys Res-atmos, 2005,
465	110.
466	Kahn, R. A.; Garay, M. J.; Nelson, D. L.; Yau, K. K.; Bull, M. A. Satellite-derived aerosol optical
467	depth over dark water from MISR and MODIS: Comparisons with aeronet and implications
468	for climatological studies. J. Geophys. Res. 2007, 112.
469	Kaufman, Yoram J. The MODIS 2.1-um channelcorrelation with visible reflectance for use in
470	remote sensing of aerosol.[J]. IEEE T Geosci Remote, 1997, 35, 1286-1298.
471	Kokhanovsky, A. A.; Curier, R. L.; Leeuw, G. D.; Grey, W. M. F. The intercomparison of AATSR
472	dual-view aerosol optical thickness retrievals with results from various algorithms and
473	instruments. Int. J. Remote Sens. 2009, 30 (17): 4525-4537.
474	Lee, S. S.; Donner L. J.; Penner J. E. Thunderstorm and stratocumulus: How does their contrasting
475	morphology affect their interactions with aerosols?[J]. Atmos. Chem. Phys. 2010, 10(2),
476	6819-6837.
477	Li, E.; Zhang, Z.; Tan, Y. A Novel Cloud Detection Algorithm Based on Simplified Radiative
478	Transfer Model for Aerosol Retrievals: Preliminary Result on Himawari-8 Over Eastern
479	China[J]. IEEE T Geosci Remote, 2020, 59(3):1-12.
480	Li, Y.; Xue, Y.; Guang, J. Ground-Level PM2.5 Concentration Estimation from Satellite Data in

481 the Beijing Area Using a Specific Particle Swarm Extinction Mass Conversion Algorithm[J].





- 482 Remote Sens-basel, 2018, 10(12), 1906.
- 483 Li, Z.; Zhao, X.; Kahn, R.; Mishchenko, M.; Remer, L.; Lee, K. H. Uncertainties in satellite
- 484 remote sensing of aerosols and impact on monitoring its long-term trend: a review and
- 485 perspective. Ann. Geophys-germany, 2009, 27, 2755-2770.
- 486 Lu, S.; Xue, Y.; Yang, X.-H.; Leys, J.; Guang, J.; Che, Y.-H.; Fan, C.; Xie, Y.Q.; Li, Y. Joint
- 487 Retrieval of Aerosol Optical Depth and Surface Refliflectance Over Land Using
- 488 Geostationary Satellite Data. IEEE Trans. Geosci. Remote Sens. 2019, 57, 1489–1501.
- 489 Martonchik, J. V. Determination of aerosol optical denth and land surface directional reflectances
- 490 using multiangle imagery. J. Geophys Res-atmos. 1997, 102 (D14): 17015-17022.
- Martonchik, J. V.; Diner, D. J.; Crean, K. A.; Bull, M. A. Regional aerosol retrieval results from
 MISR. IEEE Trans. Geosci. Remote Sens. 2002, 40, 1520–1531.
- 493 Mironova, I. A. Aerosols over continental Portugal (1978-1993): their sources and an impact on
- 494 the regional climate[J]. Atmos. Chem. Phys. 2015, 15(11):6407-6418.
- 495 Remer, L. A; Tanré, D.; Kaufman, Y. J. Algorithm for remote sensing of tropospheric aerosol from
- 496 MODIS: Collection 005[J]. 2009.
- 497 Samset, B. H.; Sand, M.; Smith, C. J. Climate Impacts From a Removal of Anthropogenic Aerosol
- 498 Emissions[J]. Geophys Res Lett, 2018, 45(2).
- 499 Sun, E.; Fu, C.; Yu, W. Variation and Driving Factor of Aerosol Optical Depth over the South
- 500 China Sea from 1980 to 2020[J]. Atmosphere, 2022, 13, 372.
- 501 Sundstrom, A. M.; Kolmonen, P.; Sogacheva, L.; Leeuw, G. D. Aerosol retrieval over China with
- 502 the AATSR dual view algorithm. Remote Sens. Environ. 2012, 116 (1): 189-198.
- 503 Wong, N.; Wong, S.; Lim, A.; Mcneil, D.; Huete, A. R. Impacts of spatial heterogeneity patterns
- on long-term trends of moderate resolution imaging spectroradiometer (modis) land surface
 temperature time series. J Appl Remote Sens. 2020, 14(1): 1.
- 506 Xie, Y.; Xue, Y.; Jie, G. Deriving a Global and Hourly Data Set of Aerosol Optical Depth Over
- Land Using Data From Four Geostationary Satellites: GOES-16, MSG-1, MSG-4, and
 Himawari-8[J].IEEE T Geosci Remote, 2019, 99:1-12.
- 509 Zhang, Y.; Li, Z.; Liu, Z. Retrieval of aerosol fine-mode fraction over China from satellite





- 510 multiangle polarized observations: validation and comparison[J]. Atmos Meas Tech, 2021, 2,
- 511 1655-1672.