First Atmospheric Aerosol Monitoring Results from Geostationary Environment Monitoring Spectrometer (GEMS) over Asia

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16 Abstract. Aerosol optical properties have been provided from the Geostationary Environment Monitoring Spectrometer 17 (GEMS). It is), the world's first geostationary earth orbit (GEO) satellite instrument designed for atmospheric environmental air 18 quality monitoring. This study describes improvements to the GEMS aerosol retrieval algorithm (AERAOD). These), which 19 include spectral binning, surface reflectance estimation, cloud masking, and post-processing. Furthermore, the study presents, 20 along with validation results. These enhancements are aimed at providing more accurate and reliable aerosol--monitoring 21 results for Asia. The adoption of spectral binning in the lookup table (LUT) approach reduces random errors and enhances the 22 stability of the satellite measurements. In addition, we introduce introduced a new high-resolution database for surface 23 reflectance estimation based on the minimum reflectance method adapted to the GEMS pixel resolution. Monthly background 24 aerosol optical depth (BAOD) values arewere used to consistently estimate the hourly GEMS surface reflectance- consistently. 25 Advanced cloud-removal techniques arehave been implemented to significantly improve the effectiveness of cloud detection 26 and enhance the quality of aerosol retrieval quality. An innovative post-processing correction method based on machine 27 learning is introduced to address artificial diurnal biases in aerosol optical depth (AOD) observations. This In this study 28 investigates specific, we investigated selected aerosol events. It This highlights the capability of GEMS to monitor and provide 29 insights into hourly aerosol optical properties during various atmospheric events. The performance of the GEMS AERAOD 30 products iswas validated against the Aerosol Robotic Network (AERONET) and Cloud-Aerosol Lidar with Orthogonal 31 Polarization (CALIOP) data for the period from November 2021 to October 2022. The GEMS AOD at 443 nm demonstrates 32 a strong correlation with the AERONET AOD at 443 nm (R = 0.792). However, it exhibits biasexhibited biased patterns, 33 including underestimation of high AOD values and overestimation inof low AOD conditions. Different aerosol types (highly 34 absorbing fine, dust, and non-absorbing) exhibited distinct validation results. The GEMS single scattering albedo (SSA) 35 at 443 nm retrievals agreed well with the AERONET dataSSA at 440 nm within reasonable error ranges, with variations 36 observed among the aerosol types. For GEMS AOD at 443 nm exceeding 0.4 (1.0), 42.76% (56.61%) and 67.25% (85.70%) 37 of GEMS SSA data points fall within the ± 0.03 and ± 0.05 error bounds, respectively. Model-enforced post-processing 38 correction improved the GEMS AOD and SSA performances performance, thereby reducing the diurnal variation in the biases. 39 The validation of the GEMS aerosol layer height (ALH) retrievals against the CALIOP data demonstrates a good agreement, 40 with a mean bias of -0.225 km₇ and 55.29% (71.70%) of data within ± 1 km (1.5 km).

42 1 Introduction

43 The regional and global monitoring of aerosol optical properties (AOPs) was conducted using satellite measurements. Low 44 earth orbit (LEO) instruments such as the Advanced Very High-Resolution Radiometer (AVHRR), Moderate Resolution 45 Imaging Spectroradiometer (MODIS), Multiangle Imaging Spectro Radiometer (MISR), Visible Infrared Imaging Radiometer 46 Suite (VIIRS), and Sea-viewing Wide Field-of-view Sensor (SeaWiFS), can provide daily aerosol properties for the global 47 domain (Hsu et al., 2004, 2006, 2017, 2019; Jackson et al., 2013; Jethva et al., 2007; Levy et al., 2013; Lyapustin et al., 2018; Lee 48 et al., 2012; Martonchik et al., 2009; Remer et al., 2005). While Although significant diurnal variations in AOPs have been observed at daily and local scales, (Kassianov et al., 2013; Kuang et al., 2015), emphasizing the importance of geostationary 49 50 satellite measurements for both air quality and climate studies, the temporal resolutionsresolution of LEO satellites (typically 51 tone per day) havehas limitations in investigating the diurnal variation and transboundary transportation of aerosols 52 (Lennartson et al., 2018; Zhang et al., 2018). Geostationary earth orbitorbits (GEO) instruments such as the Advanced Baseline 53 Imager (ABI), Geostationary Ocean Color Imager (GOCI), GOCI-II, Meteorological Imager (MI), and Advanced Himawari 54 Imager (AHI), have contributed to the operational monitoring of the continuous spatio-temporal variations in AOPs at 55 continental spatial scales with temporal resolutions of minutes to hours using the visible and near-infrared channel (Choi et al., 56 2018; Kim et al., 2016; Kim et al., 2014; Kondragunta et al., 2020; Lee et al., 2023; Yoshida et al., 2018).

57 Besides In addition to spatial and temporal resolutions, channel specification is another critical consideration for satellite 58 aerosol retrievals is channel specification. Every above mentioned instrumentretrieval. All instruments except GOCI-II 59 usesused only visible (Vis) and near-infrared channels. However, the near-ultraviolet (UV) spectral region uniquely leverages 60 its the sensitivity to aerosol absorption. Thereby, it Therefore, this study provides valuable insights into aerosol the optical 61 properties of aerosols. A majorsignificant advantage of near-UV measurements is that-the surface reflectance in the near-UV 62 region is darker than that in the visible region. This enables the derivation of AOPs over a bright surface-, typically aerosol 63 source regions. In addition, observations in the UV region are sensitive to aerosols'aerosol radiative absorption and aerosol 64 layer height (ALH) information-because. The contribution of Rayleigh scattering to the total Top of the Atmosphere (TOA) 65 reflectance enhancement is reduced below the aerosol layer owing to aerosol attenuation (Kayetha et al., 2022; Torres et al., 66 2005).

The Ozone Monitoring Instrument (OMI) serves as an example of an LEO sensor that utilizes UV wavelengths for aerosol retrievals. It has measured radiances in the 270–500 nm spectral range and offered global coverage at a spatial resolution of 13×24 km at nadir since 2004 (Levelt et al., 2018). OMI employs two aerosol algorithms. The first one, OMAERO (Curier et al., 2008), developed and maintained by the Royal Netherlands Meteorological Institute (KNMI), is a multiwavelength algorithm that relies on spectral fitting procedures to derive aerosol properties. The other is the OMI near-UV aerosol retrieval algorithm (OMAERUV). It focuses on retrieving key atmospheric aerosol properties, including-the aerosol optical depth (AOD), single scattering albedo (SSA), and absorbing aerosol index (AI) (Torres et al., 2007).

The OMAERUV algorithm has its heritage in the Total Ozone Mapping Spectrometer (TOMS) aerosol retrieval algorithm. It uses reflectance measurements at 354 and 388 nm to determine the AOD and single scattering albedo (SSA) using the twochannel inversion method (Torres et al., 2002; Torres et al., 2007). The globalGlobal statistics reported by Ahn et al. (2014) indicate a correlation coefficient (R) of 0.81. However, OMAERUV providesprovided a lower R (0.63) over Central and East Asia (Zhang et al., 2015). In addition, the Tropospheric Monitoring Instrument (TROPOMI) aerosol algorithm (TropOMAER) was developed as an adaptation of the OMAERUV. A comparison between Aerosol Robotic Network (AERONET) and TropOMAER AOD at 12 locations yielded an R of 0.82 and a root mean square error (RMSE) of 0.19 (Torres et al., 2020).

81 The Geostationary Environment Monitoring Spectrometer (GEMS) is the first UV-Vis hyperspectral satellite instrument in a

82 GEO. It is onboard the Geostationary Korea Multi-Purpose Satellite-2B (GEO-KOMPSAT-2B or GK-2B). GEMS was), 83 launched on February 19, 2020 (Kim et al., 2020). The objective of the GEMS mission is to monitor the hourly air quality in 84 Asia $(5^{\circ}S^{\circ}-45^{\circ}N, 75^{\circ}-145^{\circ}E)$ with a fine spatial resolution $(3.5 \times 7.7 \text{ km}^2 \text{ atin} \text{ South Korea})$. GEMS provides 85 hyperspectral measurements covering 300-500 nm at #0.2 nm spectral sampling and 0.6 nm full width at half maximum 86 (FWHM) spectral resolution-of 0.6 nm. Considering. The GEMS retrieval domain coverage changes with time because of the 87 varying GEMS scan patterns with the solar zenith angle (SZA), the GEMS east-west scan profiles are between morning, noon, 88 and afternoon following the sunlit part of the globe to cover the full field of regard (FOR). The GEMS aerosol retrieval 89 (AERAOD) algorithm is based on the OMAERUV algorithm and the optimal estimation (OE) method by findingdetermining 90 the optimized values of AOD, SSA, and ALH from GEMS measurements at six wavelengths (354, 388, 412, 443, 477, and 91 490 nm). In order to overcome the challenge posed by the limited degree of freedom for signal in the GEMS wavelength range, 92 this This algorithm employs the two-channel inversion method that is used in the OMAERUV algorithm to retrieve the AOD 93 and SSA- to overcome the challenge posed by the limited degrees of freedom for signals in the GEMS wavelength range. 94 Subsequently, these retrievals arewere used as the first guessesestimates for the OE method (Kim et al., 2018). The six 95 wavelengths in the UV-Vis region containcontained information regarding-the aerosol absorption in the UV region and the 96 absorption bands of the oxygen dimer (O₂-O₂) at 477 nm. This Before the GEMS was launched, this method was first tested 97 using the OMI Level 1 data and was-used to derive key aerosol parameters, including AOD, SSA, ALH, UV, and VisAI (Jeong 98 et al., 2016; Kim et al., 2018; Go et al., 2020a, 2020b). Kim et al. (2018) reported that a comparison between AERONET and 99 GEMS AOD at 26 locations in Asia yielded an R of 0.71 and an RMSE of 0.46. The percentage of GEMS SSA within the 00 expected error range of the AERONET inversion data (±0.03) was denoted by 27.54%. Spectral variations of aerosol 01 absorption in the UV-Vis region-were, as investigated by Go et al. (2020a) and it is), were applied to the GEMS aerosol 02 algorithm- to achieve improved AOPs retrieval. This adjustment accounts for the spectral dependence of aerosol absorption, 03 which was previously treated as independent of wavelength. The GEMS AOD demonstrated a strong correlation with the 04 AERONET AOD (R = 0.847 and RMSE = 0.285)), and the percentage of GEMS SSA within the expected error of ± 0.03 05 increased to 41.64% (Go et al., 2020a). To improve the accuracy of the GEMS aerosol retrieval, Go et al., al. (2020b) tested 06 the use of cloud mask information and from MODIS IR channels to remove cirrus and sub-pixel cloud contamination, as well 07 as the total dust confidence index from MODIS IR channels was tested for synergy (Go et al., 2020b), for the classification of 80 aerosol type. The limitations associated with the UV-Vis regions of GEMS were overcome using the IR channels of other 09 satellites, leading to research on the synergistic use of hyperspectral satellite instruments and broadband meteorological 10 imagers.

However, <u>asbecause</u> the testbed for the GEMS algorithm was on the LEO platform, <u>thea</u> time-dependent retrieval bias <u>hadwas</u> not <u>been observed</u> previously <u>observed</u>. The diurnal variations in <u>the</u> satellite-retrieved AOPs may differ from the actual diurnal variations <u>in the AOPs</u>. This discrepancy can be <u>attributed toexplained by</u> the different patterns of bias observed over time among <u>the different geostationaryGEO</u> satellites and retrieval algorithms (Choi et al., 2018; Lennartson et al., 2018; Wei et al., 2019; Zhang et al., 2020). This diurnal bias in AOP measurements can originate from various factors, such as errors in the surface reflectance assumption used in the retrieval algorithm, calibration issues in the Level 1 data, or the presence of short light paths at noon (Ceamanos et al., 2023).

To address this-<u>issue,</u> Zhang et al.. al. (2020) developed an empirical AOD bias-correction algorithm <u>was developed</u>. This algorithm utilizes<u>that utilized</u> the lowest AOD values observed within a-30-<u>day period_days</u> in conjunction with the background AOD to obtain a smoothed bias curve for each pixel of the ABI AOD data (Zhang et al., 2020). This approach helps mitigate the impact of diurnal bias in satellite AOD retrievals to improve the accuracy by removing artifacts from the retrieval. By applying bias-_correction methods, more reliable diurnal variations in AOD can be explained. <u>BeyondIn addition to</u> traditional statistical methods, bias correction methods based on machine learning have <u>started to bealso been</u> proposed. Model-enforced post-processing correction involves the use of a machine–_learning-based model to predict errors in conventional aerosol retrievals (Lipponen et al. 2021, 2022a, 2022b). This method was trained to learn the relationship between the input parameters of the-satellite measurements and the associated retrieval errors. This approach provides a practical and effective method to enhancefor enhancing the accuracy of aerosol retrieval without requiring extensive modifications to existing retrieval algorithms. It leverages machine–learning capabilities to improve the reliability and precision of hourly aerosol measurements obtained from GEO satellite observations.

130 In this paperstudy, we report the first aerosol monitoring resultson AOPs, including the AOD, SSA, and ALH, derived from 31 GEMS operational observations using the GEMS aerosol retrieval algorithm. The remainder of this paper is organized as 32 follows: Section 2 of the paper-describes the GEMS data and the aerosol retrieval algorithm. It also highlights the algorithm 33 updates after the GEMS in orbit test (IOT) period. Section 3 discusses the post-process correction for near-real-time retrieval. 34 Section 4 discusses the GEMS aerosol monitoring results for dust, biomass burning, and absorbing aerosol events over Asia. 35 Section 5 presents an evaluation of the retrieved-GEMS AOD, SSA, and ALH retrievals against AERONET and Cloud-Aerosol 36 Lidar with Orthogonal Polarization (CALIOP) data- and directions for future work. Finally, Section 6 presents thea summary 37 and future work.

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139 2 Data and GEMS aerosol algorithm

140 2.1 Data description

141 2.1.1 GEMS normalized radiance

42 GEMS operation process The National Institute of Environmental Research (NIER) of Korea provides the GEMS Level-1C 43 (L1C) dataset in purpose of improving, which includes various auxiliary variables necessary for retrieval to improve the 44 efficiency of the Level 2 algorithm-process by combining parameters dispersed in different files into one file. In this study, 45 Thethe aerosol retrieval algorithm used radiances only with the quality flags of 0 (Good) or 2 (interpolated radiances), 146 determined by the "bad_pixel_mask" variable. Rather than the GEMS irradiance, we used the KNMI solar reference spectrum 147 to calculate the GEMS-normalized radiance (Dobber et al., 2008). The GEMS irradiance is within the range of -5% to -20% 148 compared with the KNMI solar reference spectrum. It still requires further improvement Further improvements in L1 49 processing. To account for the spectral characteristics of the instrument, the are ongoing. The KNMI solar reference spectrum 50 iswas convolved with thea GEMS spectral response function- (Kang et al., 2020). GEMS-measured irradiances are planned 51 towill be employed when the NIER releases an improved version of the Sun L1C product-is released by the National Institute 152 of Environmental Research (NIER) ...

153 Normalized radiances are defined in the following equation:

154
$$N_{\lambda} = \frac{I_{\lambda}}{ESD \times E_{\lambda}}$$
 (1)

where *I*, *E*, *ESD*, and λ are the GEMS radiance, KNMI solar reference spectrum, earth–sun distance correction factor, and wavelength (354, 388, 412, 443, 477, and 490 nm), respectively. The spectral spectral radiance and irradiance were spectrally binned and averaged within ±2.2 nm fromof each wavelength to enhance the measurement signals. Additionally, earth– sunEarth-Sun distance correction was used to calculate the normalized radiance.

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160 2.1.2 AERONET

AERONET is a global ground-based remote-sensing network that measures aerosol optical, microphysical, and radiative

- 162 properties (Giles et al., 2019; Holben et al., 1998; Sinyuk et al., 2020). The measurement systems useused Cimel sun 63 photometers to measure the solar irradiancesirradiance at eight wavelengths ranging from 340 to 1020 nm and sky radiances 64 at four wavelengths ranging from 440 to 1020 nm. The AERONET data provide global aerosol information, including the 65 spectral AOD and inversion products, such as the SSA, aerosol size distribution, and refractive index. The uncertainties in 166 AODs are wavelength-dependent. It is approximately 0.01 (Vis) to 0.02 (Near-UV) in direct sun measurements (Dubovik et 167 al., 2002). The uncertainties of SSA are ±0.03 when AOD exceeds 0.4 at 440 nm (Dubovik et al., 2002). For the evaluation of 168 GEMS AOD and SSA data from November 2021 to October 2022, we used AERONET V3 Level 1.52.0 data for AOD and 69 AERONET V3 Level 1.52.0 hybrid inversion data for SSA from all sites within the entire GEMS domain, ensuring higher 70 guality compared to Level 1.5. However, we used AERONET V3 Level 1.5, data for AOD, and AERONET V3 Level 1.5 71 hybrid inversion data for SSA for post-process correction to ensure a sufficient volume of data during the modelling and near-
- 172 <u>real-time processing</u>.
- 173

174 2.1.3 CALIOP

The CALIOP instrument is a two-wavelength polarization-sensitive lidar on the <u>Cloud Aerosol Lidarcloud aerosol lidar</u> and Infrared Pathfinder Satellite Observations<u>infrared pathfinder satellite observation</u> (CALIPSO) satellite. It was launched on April 28, 2006 (Winker et al., 2009). CALIOP monitors the global vertical profiles of aerosols and clouds by measuring three signals: the backscatter intensity at 1064 nm and the orthogonally polarized components of the backscattered signal at 532 nm.

Quantitative scattering information from the CALIOP instruments was used as reference data for validating to validate the ALH obtained from passive sensors (Xu et al., 2017; Xu et al., 2019; Nanda et al., 2020; Park et al., 2023). We used the CALIPSO Lidar Level 2 Aerosol Profile V3-41 data to validate the GEMS ALH. CALIOP profiles of the extinction coefficient (β_{ext})() at the 532 nm channel were utilized to calculate the CALIOP ALH using the following equation:

183
$$Z_{aer} = \sum_{i=1}^{n} H(i) \left[\frac{\beta_{ext}(i)}{\sum_{i=1}^{n} \beta_{ext}(i)} \right]$$
(2)

where $\beta_{ext}(i)$ is the CALIOP profile of the 532 nm extinction coefficient at height H(i), and n is the number of layers.

185

186 2.2 GEMS AERAOD retrieval algorithm

187 2.2.1 Aerosol optical properties retrieval algorithm for GEMS

The GEMS AERAOD algorithm produces AOD, SSA, and ALH viadata using the OE method. The preliminaryAn early version of the GEMS AERAOD was developed using OMI L1B normalized radiance (Kim et al., 2018; Go et al., 2020a, 2020b). After the launch, the algorithm was tested using the GEMS observationobservations during the IOT period, and several parts of the algorithm were updated. This section briefly describes the GEMS AERAOD algorithm²/₂ AERAOD L2 data²/₂ and updates² including the Look-Up Tablelookup table (LUT), cloud-masking procedure, surface reflectance estimation, and postprocessing after the IOT period. The general flow of the GEMS AERAOD retrieval algorithm is illustrated in Figure 1.

194 <u>The GEMS algorithm adopts an LUT approach to optimize computation efficiency. The LUT iswas calculated assuming the</u> 195 AOPs of three aerosol types-by using a radiative transfer model (RTM), the Vector Linearized Discrete Ordinate Radiative 196 Transfer code (VLIDORT) (Spurr, 2006). The AOPs of Highlyhighly absorbing fine (HAF), Dust, and Non-absorbing (NA) 197 are integrated from the AERONET inversion data and are-applied for the RTM simulation. -The details of the updated LUT 198 are described in sectionSection 2.1.2. The preliminary algorithm used the OMI climatology Lambertian equivalent reflectance 199 (OMLER v003) datasets as surface reflectance, but (Kleipool et al., 2008). However, for the GEMS AERAOD algorithm, GEMS L2 surface reflectance at 354, 388, 412, 443, 477, and 490 nm are were obtained by using the minimum reflectance method. The details of the surface reflectance estimation are described in sectionSection 2.1.3.

The-GEMS AERAOD provides UV and visible (Vis) AI to indicate-the qualitative radiative absorptivity and particle size information, respectively (Torres et al., 2002). The GEMS UVAI and VisAI were calculated using the following equations:

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$$AI = -100 \left[log \left(\frac{N_{\lambda_1}}{N_{\lambda_2}} \right)_{meas} - log \left(\frac{N_{\lambda_1}(LER_{\lambda_1})}{N_{\lambda_2}(LER_{\lambda_2})} \right)_{calc} \right]$$
(3)

where $N_{A_{\pm}}$ and $N_{A_{\pm}}$ are the normalized radiances at the 354/388 (477/490) nm wavelength pair for UVAI and VisAI, respectively. The subscripts Subscripts meas and *calc* represent the measured and calculated normalized radiances, respectively.

207 Aerosol type among The aerosol types HAF, dust, and NA iswere selected using the UVAI and VisAI. The NA type was 208 detected by aA negative UVAI value, was detected for the NA type. The dust and HAF types were distinguished by using the 209 VisAI. When HAF type was selected when UVAIs are positive and VISAI is negative. The dust type was selected when both 210 AIs were positive, the dust type was selected. Sun glint and cloud masking leave only the pixels appropriate for aerosol 211 retrieval. The glint mask iswas set for glint angles less than 35°. The details of the cloud-masking procedure are described in 212 Section 2.1.4. The *a priori* states of AOD and SSA at 443 nm were obtained by two-channel inversion with neighboring 213 wavelengths (354 and 388 nm) over both land and ocean. The assumption was that, with a priori states of ALH based on the 214 climatology of ALH was based on CALIOP ALH. The a priori states of the AOD and SSA were supplied to solve the 215 Levenberg-Marquardt equation- (Rodgers, 2000). The optimal ALH was retrieveddetermined by fitting the normalized 216 radiance between the measured and calculated values for the OE routine. The detailsDetails of the GEMS aerosol inversion 217 procedure are described byin Kim et al. (2018).

To improve the accuracy of near real-time GEMS AOD retrieval, a model-enforced post-processprocessing correction step was implemented using a random forest (RF) model. By combining GEMS aerosol retrieval with this post-processing correction model, more reliable and accurate near-real-time AOD estimates can be obtained.

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222 2.1.2 LUT calculation

223 In this study, the AOPs were considereddetermined as described by Kim et al. (2018) and Go et al. (2020a). However, the 224 dimensions of the LUT varied (as shown in Table 1) compared with), which is different from Kim et al. (2018). The nodes for 225 the 412 nm SSA node for the NA were added. In addition, the nodes for the AOD in the LUT were extended to include the 226 values at 5.0° and 10.0-because the previous maximum node was 3.6. These modifications enable°, enabling the retrieval of 227 exceptionally severe aerosol events during GEMS observations. The preliminary early version of the GEMS AERAOD retrieval 228 algorithm utilized the normalized radiance at six specific monochromatic wavelengths (354, 388, 412, 443, 477, and 490 nm). 229 However, satellite measurements averaged over a specific wavelength range produce more stable values than measurements 230 obtained at individual monochromatic wavelengths. This increased stability is attributed, owing to the averaging of random 231 errors (i.e., instrument noise).

Consequently, a spectral-binning LUT approach was employed to reduce random errors and improve the<u>measurement</u> stability of the measurements. This allowed for<u>enabled</u> more reliable and consistent observations. Compared with<u>to</u> monochromatic wavelengths, the spectral binning method is computationally intensive. Therefore, the calculations were performed using the Mie theory without considering the non sphericity of the dust. <u>a</u> forward RTM coupled with the Mie theory. The aerosol parameters, including the mean radii and standard deviations of the fine and coarse modes, respectively, of the aerosol bimodal number size distribution, fine mode particle fraction with respect to the total number concentration, and real part of the 238 refractive index, were used to generate the LUT (Kim et al., 2018).

239 The process of spectral binning LUT in the GEMS aerosol algorithm involves three steps: 1) A reference spectrum is generated 240 using an RTM, which provides a spectral interval of 0.1 nm. 2) The calculated spectrum iswas convolved with the GEMS 241 spectral response function and resampled to the target spectral grids with a resolution of 0.2 nm₇ (Kang et al., 2020). 3) The 242 resampled spectrum is averaged at intervals of ± 2.2 nm at six central wavelengths (354, 388, 412, 443, 477, and 490 nm) and 243 saved in the LUT. This range is Intervals of ± 2.2 nm were selected to account for the calculation capacity and reduce the impact 244 of random errors. During the retrieval process, the GEMS L1C normalized radiances, after being averaged at intervals of ± 2.2 245 nm at six central wavelengths, are compared with the calculated spectrum in the LUT. By Through these steps, the spectral-246 binning LUT aims to generate more stable retrieval results for aerosol properties.

247

248 2.1.3 Surface reflectance estimation

249 In this study, several Several improvements were introduced- in this study. These include an updated GEMS surface reflectance 250 estimation method. The preliminary early version of the GEMS AERAOD retrieval algorithm used the OMI surface reflectance 251 climatology–data product (OMLER v003) (Kleipool et al. 2008), with a spatial resolution of $0.5 \times 0.5^{\circ}$. The limitation of the 252 previous surface reflectance data was its, which is too coarse spatial resolution compared with that of GEMS pixels. This 253 resulted pixel size, therefore, resulting in discontinuities in the GEMS AOPs-owing to spatial resolution differences. To address 254 this limitation, the. The updated GEMS surface reflectance hashed a finer spatial resolution $(0.1 \times 0.1^{\circ})$, to address this 255 limitation. This aligns closely aligns with the GEMS pixel resolution of the GEMS. This enhancement enables aenabled more 256 accurate aerosol retrieval at the pixel level. The compiled hourly surface reflectance indirectly reflects the bidirectional 257 reflectance distribution function (BRDF) effect. In addition, a new hourly surface reflectance database was generated using 258 the minimum reflectance method based on the GEMS data- (Herman and Celarier, 1997; Hsu et al., 2004). The algorithm 259 adopts the climatological minimum reflectance method for each pixel over a ± 15 -day window spanning a period of two years. 260 Several tests were performed to evaluate different time windows and methods for constructing an accurate surface reflectance. 261 These tests evaluated the effectiveness of using a ± 15 -day window as well as alternative options such as a previous 30-day 262 window. In addition, different methods, including the minimum reflectance and second minimum reflectance approaches, were 263 evaluated to determine the most suitable one method for generating appropriate surface reflectance values (not included in this 264 study).

265 The background AOD (BAOD) was considered in the retrieval algorithm. The BAOD represents the baseline level of AOD 266 that is consistently present in a region. This was then used to derive the surface reflectance dataset. The Rayleigh scattering, 267 gaseous absorption, and BAOD were corrected during the atmospheric correction process to create a surface reflectance dataset. 268 Recent studies have shown that incorporating BAOD into an algorithm can reduce the uncertainty associated with satellite-269 based AOD remote sensingretrieval (Kim et al., 2014, 2021). Zhang et al. (2016) estimated the BAOD as the lowest fifth 270 percentile of the AERONET AOD over a two-year period years and improved the performance of the VIIRS aerosol algorithm. 271 It has been observed that Asia experiences relatively high BAOD values with seasonal variation. For example, at the Dhaka 272 University site, the monthly BAOD over the past two years varied from a minimum of 0.124 in August to a maximum of 0.685 273 in April. Therefore, considering the seasonal variation and barriations in BAOD for atmospheric correction can help mitigate the 274 uncertainty in satellite-derived AOD retrieval, particularly over Asia. The monthly BAODs were calculated using the following 275 equation for each $0.1 \times 0.1^{\circ}$ box from November 2020 to October 2021:

276 $\tau_{grid,b,m}(lat,lon) = \sum_i W_i \tau_{b,m,i^{-}} / \sum_i W_{i_{-}}$

where $\tau_{grid,b,m}(lat, lon)$ is the interpolated BAOD 443 nm at (lat, lon) in month m. \mathcal{W}_{t} is the inverse distance weighting

(4)

function, which is defined as $e^{-d_t(lat,lon)/d_n}$. $d_t(.(lat, lon))$ is the distance between the AERONET site and the GEMS pixel. and d_{θ} is a constant, respectively. $\tau_{p.m,t}$ is the lowest fifth percentile of AERONET AOD over a two year period years at AERONET site *i* in month *m*.

Figure S1 shows the monthly BAOD obtained based on the AERONET AOD data. Additionally, the fifth percentiles of the AERONET AOD 443 nm values at each AERONET site are plotted as circles for reference. It is evident that regions such as India exhibit a high BAOD of over-approximately 0.15 throughout the year, regardless of the month. However, seasonal variations in BAOD occuroccurred over the IndochineseIndo-Chinese Peninsula, Korea, and China. These areas experience heavy pollution from biomass burning during the dry season and dust events from the deserts. Both these factors contribute to increased atmospheric aerosol concentrations. These enhancements, including the use of hourly GEMS surface reflectance and the incorporation of monthly BAOD, can result in improved improve aerosol retrieval.

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289 2.1.4 Cloud masking procedure

290 The GEMS aerosol algorithm retrieved AOPs only infor cloud-free pixels. Clouds exhibit spatial inhomogeneity and <u>a higher</u> 291 brightness than aerosols. This study aimed to enhance the cloud-masking process in the GEMS aerosol algorithm by addressing 292 the limitations of previous simple cloud-masking techniques. The previous method relied on a (Step_1) fixed threshold for 293 reflectance at 412477 nm and (Step 2) standard deviation test of reflectance at 477 nm within a 3 × 3 pixel area. To improve 294 the performance of cloud masking, an An additional cloud-removal technique has been was introduced in this study- to improve 295 the cloud masking performance. These tests includeincluded the following: (Step 3) a 470/477 nm normalized radiance ratio 296 test. It involves This involved a threshold test for the ratio of the normalized radiance values at 470 nm and 477 nm. This 297 contrasts with the presence of clouds using the absorption bands of O_2 - O_2 -(4) due to the decrease in absorption of O_2 - O_2 at 298 477 nm in the presence of clouds (Kim et al., 2024) (Step 4). The difference between the hourly surface reflectance database 299 and the-calculated scene reflectivity at 412 nm: Significant differences indicate indicates the presence of clouds (Torres et al., 800 2013). (Step 5) Standard deviation test of normalized radiance at 477 nm within a-3 × 3 pixel area: pixels> f(latitude): The B01 threshold for this test can vary based on the latitude, considering the regional differences in cloud characteristics. (5-1Step 6) 802 Standard deviation in at 477 nm within 3×3 pixel > f(latitude) (5.2) after 3.1, standard deviation in 3×3 pixels pixels> 803 f(latitude, number of cloud pixels-detected method (1), (3), (): The threshold for this test can vary based on the latitude and 804 <u>number of pixels detected as clouds from Steps 1 to 4) in 3×3 pixels).</u> A final cloud mask was applied after the aerosol 805 retrieval. This included (6) filtering(Step 7) Filter out the high AOD values usingover the ocean > f(number of cloud pixels): a 806 threshold that is a function of the number of cloud pixels detected by methods (as clouds from steps 1), (3), (4), and (5) to 6 in 807 an 11×11 -pixel over the oceanwindow (Lyapustin et al., 2021). This helps remove residual clouds- over the ocean. By 808 implementing these new methods, the proposed algorithm aims to improve the effectiveness of cloud detection and removal 309 in GEMS pixels.

810 Quantitative analysis was performed to assess the impact of the improvements in each section on the retrieval results of GEMS B11 AOD at 443 nm (Table S1). We analyzed the influence of each update factor on the AOD validation results. The validation 812 periods were January, April, and July 2022. The statistics included R, RMSE, mean bias error (MBE), slope, y-offset, Q value 813 indicating the percentage of AOD retrievals falling within the uncertainty envelope of 0.1 or ±30% of AOD error range, and 814 the Global Climate Observing System (GCOS) requirement is defined as the percentage of AOD retrieval falling within the 815 uncertainty envelope of 0.03 or \pm 10% for AOD error range. The early version of the GEMS AERAOD had an MBE of 0.36, 816 indicating an overestimation of the GEMS AOD. When using KNMI irradiance instead of GEMS irradiance and changing to 817 spectral binning LUT, Set1 resulted in a closer MBE of -0.074 to zero and an increased Q-value of 50.63%, approximately 30% 818 higher than the results of the early version of GEMS AERAOD. Set 2 was the result of the analysis using the GEMS surface

- reflectance instead of the OMI climatology values as the surface reflectance (Section 2.1.3). Set 2 showed a slight decrease in
 the R-value but an improvement in the Q-value by over 7%. Finally, introducing a new cloud removal method (Set3) increased
 R and decreased RMSE, leading to an increase in the Q value compared to Set2.
- B22

323 3 GEMS post-process correction for the near-real-time retrieval

The GEMS AOD <u>exhibitedexhibits</u> a diurnal bias pattern that <u>fluctuatedfluctuates</u> throughout the day. It formed a U-shape, with a minimum at 03:00 UTC (as will be demonstrated in Section 5.1). To improve the accuracy of near real time GEMS AOD retrieval, <u>aA</u> model-enforced post-<u>processprocessing</u> correction step was implemented using <u>athe</u> random forest (RF) model proposed by Lipponen et al. (2021)-:) to improve the accuracy of near real-time GEMS AOD retrieval.

828 This concept was trained to learn the relationship between the hourly GEMS data and AOD errors (GEMS-AERONET AOD) 829 and to predict the AOD errors at the target time. To enable near real time retrieval, the The proposed method consists of two 830 main parts: modelling and prediction-to enable near-real-time retrieval. In the modelling part, the input data for the RF model 831 includes included GEMS data (normalized radiances at six wavelengths, scattering angle, viewing zenith angle (VZA), relative 832 azimuth angle (RAA), SZA UV and VisAI, aerosol type, AOD, and a clear fraction (ClearFrac) (which is the ratio of clear-sky 833 pixels to the total pixels within thea 0.25° radius from the pixel center)). The data also include auxiliary information, such as 834 time, land-sea mask, and elevation. The target data for training were the AOD errors. Each of these was, which were calculated 335 as the difference between the GEMS AOD and AERONET AOD at the corresponding single GEMS pixel where the 836 AERONET site was located. AERONET data are temporally matched within a ±10 min window of the GEMS measurement 837 time. Data from three AERONET sites (Sorong, Jambi, and BMKG_GAW_PALU) with severe subpixel cloud contamination 838 were excluded from the modeling to exclude cloud-contaminated pixels during the modelling process. The predictors and 339 target variables were collected for a time window ranging from N days to one day before the target time. After conducting 840 several tests, N was determined to be 30 days. In the prediction part, the input variables, including the GEMS data and auxiliary 841 information inat the target time, were used for the pretrained pre-trained RF model. Using these inputs, the model 842 predicted predicts the error in the GEMS AOD in near real-time. This predicted error value was then applied to the first GEMS 843 AOD retrieved GEMS AOD fromusing the retrieval algorithms algorithm. This resulted in the production of the post-844 processed postprocessed GEMS AOD.

845 To investigate the performance in areas without AERONET data, we conducted a Leave-One-Site-Out Cross-Validation. This 846 principle involves excluding data from one site and training the model using data from all other sites. The performance of the 847 model was evaluated using data from excluded sites. The station selected for evaluation was excluded from the model-fitting 848 process. For the period ranging from 30 days prior to the current day to 1 h before the target day, modelling was conducted, 849 excluding data from one site. The predictive accuracy of the model was evaluated for one site on the target day. Figure S2 850 shows the statistical maps illustrating the results of the Leave-One-Site-Out Cross-Validation for post-process-corrected GEMS 851 AOD for the one year from November 1, 2021, to October 31, 2022. In Northeast Asia, there was a notably high R, indicating 852 a strong relationship in the AERONET data. However, sites closer to the equator tended to exhibit lower R values, around 0.5. 853 The RMSE followed a similar pattern, with lower values in densely populated Northeast Asia, reflecting a better fit between 854 the predicted and AERONET values in this region. The MBE in Northeast Asia tended to be close to zero, suggesting minimal 855 bias in the predictions. In contrast, the Indian region shows negative MBE values, indicating an underestimation, whereas 856 Southeast Asia shows positive values, signifying an overestimation.

A variable importance analysis for post-processing correction of the GEMS AOD was conducted (Figure S3). The GEMS AOD
 was the most crucial variable, emphasizing its direct influence on the correction process. The VZA and elevation are highly
 important. However, their significance can be attributed not only to their inherent properties but also to their role in conveying

AERONET location-related information. The aerosol type appeared to be less significant in the RF models. This result
 contrasts the notable importance of the GEMS UVAI and VisAI. This discrepancy can be attributed to the inaccurate aerosol-

type classification in the GEMS aerosol algorithm.

- 863 In addition, the diurnal bias pattern in the GEMS SSA-also exhibited fluctuations throughout the day, forming a bell shape 864 with a minimum at 03:45 UTC. This is showndiscussed in Section 5.2. The post-processing method adopted was similar to 865 that used for the AOD. This method was trained to determine the relationship between hourly GEMS data and SSA errors 866 (the difference between GEMS at 443 nm and AERONET SSA at 440 nm) and to predict SSA errors for the target time. The 867 key difference between the RF model predicting the AOD error and that-predicting the SSA error is as follows: the. The 868 second model includes the GEMS SSA as an input variable, and then, 19 input parameters are used to construct as well. A 869 variable importance analysis for the post-processing correction of the GEMS SSA was conducted (Figure S4). The GEMS 870 SSA was the most critical variable in the correction process. The GEMS AOD also emerged as a highly influential variable 871 in the RF model models for GEMS SSA post-process correction. In addition, the aerosol types appeared to have relatively
- 872 <u>low significance within the RF models for SSA correction</u>.

Unlike AOD and SSA, the <u>postprocessingpost-processing</u> of ALH using an RF model is inherently limited. <u>CALIOP is</u> predominantly used as reference data for ALH. Because by the fact that CALIOP is an LEO satellite, and pixels co-located with GEMS ALH data are available only from 03:45 to 07:45 UTC. This <u>rendersrendered</u> it inaccessible as a reference hourly dataset covering 22:45–02:45 UTC. Unlike AEROENT, the use of data from ground-based lidar is severely constrained by the limited number of observation stations and restricted geographical areas in which <u>the</u> lidars are deployed.

378

879 <u>4Aerosol</u>4 Aerosol events

380 4.1 Dust aerosol event (2022.04.08)

Figure 2 presentpresents an example of hourly maps of the GEMS aerosol product, including AOD, SSA, and ALH, UVAI, and VisAI for April 8, 2022. TheseNote that these results are the GEMS AOD, SSA, and SSAALH before post-processing. The selected case is for the dust aerosol event over northwestern China. The GEMS false RGB is shown using R (477 nm), G (412 nm), and B (354 nm) bands similar to those of the OMI false RGB method (Levelt et al., 2006).

885 As shown in Figure 2, different the GEMS retrieval regions domain coverage changed with respect to time are shown asowing 886 to the varying GEMS scan profile variespatterns with the SZA. Overall, the GEMS AOD showshowed a significantly good 387 agreement with the AERONET AOD measurements. It captures higher values in the Beijing-Hebei-Tianjin (BTH) region and 888 lower values overin South Korea and Japan. High GEMS AOD values were evident along the dust plume, attaining tworeaching 889 2.0 at 06:45 UTC. In the case of SSA, the retrieval results demonstrated a relatively lower accuracy (notably in the BTH region) 890 compared with AOD. In general, from 22:45 to 05:45 UTC, the SSA values displayed good concordance with both the 891 AERONET and GEMS SSA. However, from 06:45-to-_07:45 UTC, the SSA numbers did not match overthose for Beijing. 892 Compared with the Beijing region, the results arewere more consistent infor the dust plume. The SSA values remained 393 relatively stable at approximately 0.92-0.96 over time. However, the GEMS SSA tended to have a positive bias compared with 894 the AERONET values. This is showndiscussed in Section 5.2. The GEMS ALHs were ~3-4 km for the dust plume over the B95 Taklamakan Desert and ~1.0 km over the Beijing region. The GEMS ALH exhibited continuous spatial and temporal patterns. 896 The UVAI provides information regarding the radiative absorption of aerosols. It attained a maximum of four for dust plumes, 897 thereby indicating significant aerosol absorption. However, over Beijing, the SSA was ~1. This indicated a marginal absorption 898 owing to the different aerosol emission source. VisAI provides information on the aerosol size. In regions with a dust plume, 899 the VisAI value was higher than that in the background areas. This indicated the presence of coarse aerosol particles.

401 4.2 Biomass burning event (2022.03.19)

402 Figure 3 illustratesshows maps of the GEMS aerosol product at 06:45 UTC on March 19, 2022. It This represents a biomass-403 burning event overin mainland Southeast Asia. These results were obtained for the GEMS AOD and SSA before post-404 processing. During the dry season in this region, highly absorbing fineFine pollution particles are prevalent in this region 405 during the dry season (Yin et al., 2019). The GEMS AOD > 1.6. This indicated a significant aerosol loading and enhancement 406 during the event. The GEMS SSA was - is approximately 0.88. This indicated indicates aerosol absorption during this event. 407 The ALH ranged from 2 to 3 km within the biomass-burning plume. The GEMS ALH was not retrieved along the east-to-west 408 straight line at ~22.5-°N, which are is a bad pixelspixel in the CCD. The GEMS UVAI showed revealed hotspots and fine 409 features associated with this event. Thus, it eaptured captures the aerosol absorption in the ultraviolet spectrum. GEMS VisAI 410 exhibited higher values than the background did not clearly show signals from small particles caused by biomass burning, 411 indicating that signals from the surface were not completely removed. There may be limitations in considering aerosol size 412 information using GEMS VisAI (Go et al., 2020b). This case study demonstrates that the GEMS provides valuable insights 413 into aerosol properties during specific events such as biomass burning, and can capture temporal and spatial variations in AOD, 414 SSA, ALH, UVAI, and VisAI.

415 Figure 3g shows a comparison of the CALIOP extinction coefficients at 532 nm, the CALIOP ALH, and the GEMS ALH over 416 the CALIOP path (the green line on the GEMS false RGB image in Figure 3a). Figure 3g illustrates a clearthe precise 417 relationship between the GEMS AOD and the accuracy of the GEMS ALH. The accurate Accurate retrieval of ALH requires 418 the presence of a sufficient amount of aerosols in the atmosphere. The GEMS ALH closely followsfollowed the latitudinal 419 variation in the CALIOP ALH. As the latitude increased from 18° to 21°, the GEMS ALH followed the CALIOP ALH and 420 showed exhibited an increase in altitude. In the latitude range of $24^{\circ}-28^{\circ}$, the GEMS AOD decreased, and the GEMS ALH 421 exhibited scattered variations owing to weaker signals. In the scatter plot comparing CALIOP ALH and GEMS ALH (Figure 422 3h), 39.88% of the pixels are within the expected error range of 0.5 km, and 68.10% of the pixels are within the expected error 423 range of 1 km. As the GEMS AOD values decreased, the GEMS ALH pixels were more likely to befall outside the expected 424 error range.

425

426 4.3 Absorbing aerosol event (2021.12.04, 2021.12.23)

427 Figure 4 shows an example of the GEMS AOD before and after post-processing for an absorbing aerosol case over the Indo-428 Gangatic Plane (IGP) at 04:45 UTC on December 4, 2021. During the wintertime in this region, atmospheric Atmospheric haze 429 is prevalent in this region during the winter (Ram et al., 2012). Recent studies have shown that primary aerosols and precursors 430 for of secondary aerosols emitted from fossil fuel combustion and biomass burning are released into the atmosphere (Singh et 431 al., 2021). Figure 4a shows thea GEMS false RGB image with the AERONET stations represented by circles. The color 432 indicates the Colors indicate AERONET AOD. Two distinct aerosol plumes arewere observed. The northwest shows showed an 433 AOD of -approximately 0.8, whereas the southeast hashowed a value of -approximately 1.3. Figure 4b shows the GEMS 434 AOD data. The spatial distribution of the GEMS AOD is was similar to that of the AERONET AOD, as shown in Figure 4a. 435 However, the values are were marginally lower than those of the AERONET AOD. Meanwhile However, the post-processed 436 AOD increased after post processing showed an elevated value, particularly in the moderate original AOD range (~ 0.7). 437 Moreover,), bringing the GEMS AOD was-closer to the AERONET AOD (Figure 4c). Specifically, at the Gandhi--College 438 site (25.871-°N, 84.128-°E) and Lahore (31.480-°N, 74.264-°E), postprocessing post-processing resulted in more reasonable 439 values.

- Figure 5 shows the maps of the GEMS SSA and the GEMS SSA after post-processing for an absorbing aerosol case over India, Bangladesh, and mainland Southeast Asia at 03:45 UTC on December 23, 2021. Figure 5a shows the The GEMS false-color RGB image-with AERONET stations, represented by circles-, is shown in Figure 5a. The color indicates the AERONET SSA at 440 nm. The AERONET SSA values are ~0.9 in India and Bangladesh, and ~0.93 in Thailand. Before postprocessingpostprocessing, the GEMS SSAs exhibitexhibited values of ~0.96 in the Indian region and ~1.0 in the-other areas. However, following postprocessingafter post-processing, the GEMS SSA values converged to be and became more similar to the AERONET SSA values. Nonetheless, a marginal tendency fortoward overestimation remained.
- 726

727 5 Validation in GEMS AERAOD product

728 This section evaluates the GEMS AOD and SSA at 443 nm according to the aerosol type and measurement time using the 729 AERONET data in the entire GEMS domain. We used AERONET version 3 level 1.52.0 data to validate both the AOD and 730 SSA to ensure a larger dataset for validation purposes, as it is quality-assured. Figure 6 illustrates a map of the AERONET 731 sites used for GEMS AOD and SSA validation, in conjunction with site-specific data counts. The AERONET AOD data 732 generally showed higher counts forin South Korea, China, and Taiwan. Meanwhile, sitesSites in South and Southeast Asia 733 typically hadhave fewer data points. Similarly, the The number of AERONET SSA data points showed a distribution similar to 734 that of the AOD. However, AERONET sites #38, #39, and #47 in India had over 400 validation points. In addition, we retrieved 735 the GEMS ALH and compared it with the CALIPSO level 2 extinction coefficient profiles at 532 nm, as well as with the 736 CALIOP ALH defined by Equation (2).

737

738 5.1 Aerosol optical depth

In this section, the GEMS AOD at 443 nm iswas validated against AERONET data across the entire GEMS domain from November 1, 2021, to October 31, 2022. The GEMS AOD data were spatially collocated within a 0.25° radius of the AERONET stations and temporally within a 30-min-minute window of the GEMS measurement time. When a specific aerosol type in the GEMS was present in more than 90% of the pixels within the validation radius, an-aerosol-type validation was conducted.

- 744 Figure 7 presents the results for all the pixels and each aerosol type (HAF, dust, and NA). The statistics include R, RMSE, 745 mean bias error (MBE), slope, y offset, Q value indicating the percentage of data points within the maximum (0.1 or 30% 746 AOD) error range, and the Global Climate Observing System (GCOS) requirement (defined as the maximum (0.03 or 10% 747 AOD)). The total GEMS AOD demonstrated a good correlation with the AERONET AOD, with $\frac{1}{2} = 0.781$, RMSE = 748 0.221, and MBE = 0.792, RMSE of 0.227, and MBE of 0.038047 (Figure 7a). The Q value was calculated to be 54.84as 749 52.93%, with 18.3917% of the AOD satisfying the GCOS requirements. However, the slope and y-intercept wereare 0.589572 750 and 0.193202, respectively. This indicated an overestimation for a flow AERONET AOD and an underestimation flow AERONET AOD action flow AERONET AOD 751 AERONET AOD. In the case of a low AERONET AOD, there There is evidence of cloud contamination effects. These result 752 in the case of low AERONET AOD. This results in an overestimation of the retrieved GEMS AOD.
- The validation shows the showed differences by according to aerosol type. The HAF type showed the highest R and Q values compared with the other aerosol types (Figure 7b). Pixels that deviated beyond the error range owing to the-GEMS AOD underestimation were notably observed in two main categories: sites in the Indian region (which still showed bias notwithstanding the consideration of BAOD) and sites located in Beijing with an AERONET AOD of approximately 2.0_{\pm} and a GEMS AOD of approximately 1.0. Among the three aerosol types, the dust type had the fewest samples, accounting for 1 / 1015 of the NA -(Figure 7c). The R-value was 0.786821, and the slope was the highest among the three types. Pixels that deviated beyond the error range owingdue to GEMS AOD underestimation were primarily observed in the Indian region. In

760 contrast, pixels exceeding the error range owing to GEMS AOD overestimation were located in Northeast Asia. Currently, the '61 GEMS uses the same aerosol model (number-size distribution parameters and-real refractive index) over the entire domain for 762 each aerosol type. However, given the varying bias patterns observed in the dust typetypes, it is necessary to consider regional 763 variations in the GEMS aerosol model (and, thus, the LUT) in future studies. The NA-type was selected most frequently among 764 the three aerosol types (Figure 7d). Figure 7d shows that a significant number of pixels arewere influenced by cloud 765 contamination, which is was particularly evident in regions with low NA AOD values. It appears that the The GEMS aerosol 766 cloud masking process requires further improvement, particularly over the ocean. The current cloud-masking process may not 767 effectively distinguish small clouds (i.e., broken clouds) near the equatorial regions. This resultsresulted in an overestimation 768 of the AOD owing to cloud contamination. This phenomenon has frequently been observed frequently at AERONET stations 769 located near the equator. The underestimation of high AOD values inby the GEMS aerosol algorithm can be attributed to the 770 effecteffects of the current aerosol model assumptions used in the algorithm. This emphasizes the importance of 771 understanding the AOPs to better characterize these them in the atmosphere, particularly in the UV region.

772

773 Figure \$2\$5 and Table 2 present the hourly AOD validation results and statistical metrics, including N, R, slope, y-intercept, 774 RMSE, MBE, Q value, and GCOS. It is important to note that the GEMS varies its E-E-W scan profile of the GEMS varied 775 depending on the SZA. Therefore, the sites used for the validation may not have remained consistent over time. For example, 776 the AERONET stations around 22:45 UTC and 23:45 UTC were mostlymainly used for validation in the eastern region of 777 GEMS, whereas those around 06:45 UTC and 07:45 UTC were expected to be located in the western region of GEMS. A 78 systematic error analysis is will be planned in a-future study studies. Nevertheless, the hourly validation results of the GEMS 79 AOD provide significant insights. The hourly slopes of the GEMS AOD exhibited a-diurnal variations, starting at 780 0.730725 at 22:45 UTC; decreasing to 0.534490 and 0.555 by 533 at 1:45 UTC and 2:45 UTC, respectively; and 781 subsequently increasing to 0.647606 and 0.617632 at 06:45 and 7:45 UTC, respectively. However, the R-values remained 782 relatively stable over time. Most time intervals exhibited R values of approximately 0.777 or higher except for 22at 00:45. 783 Figure <u>S2S5</u> and Table 2 show that the diurnal variation in GEMS AOD diddoes not precisely reflect the actual diurnal AOD 784 variation. Thus, it is necessary to correct and produce a consistent dataset over time to investigate the diurnal variations in 785 aerosol properties. A machine-learning model using RF was used to train the hourly dependent error characteristics, remove 786 artifacts in the retrieval processes, and maintain the physical signals.

787 Figure 8a shows the comparison results for GEMS AOD after model-enforced post-processing correction with AERONET 788 data. For near real time post-processing correction, data from the past 30 days were used for training. Therefore, these results 789 were evaluated over 11 months: from December 1, 2021, to October 31, 2022. Figure 8a shows that all the statistical metrics 790 improved. In particular, the slope was closer to one at 0.809857, and the y-intercept was closer to zero at 0.068049. Additionally, 791 the R, RMSE, and MBE were 0.899920, 0.159135, and -0.005001, respectively. The Q value and GCOS requirements alsowere 792 improved to 79.13 by 82.17% and 36.0837.29%, respectively. The bias near low AOD values of approximately zero was 793 reduced significantly reduced. Furthermore, the high AOD values were closer to the 1:1 line. Figure 8b shows the bias of the 794 GEMS AODs before and after post-processprocessing correction with respect to time for all the AOD pixels. After applying 795 the model-enforced post-processprocessing correction to the GEMS AOD data, significant improvements in bias were 796 observed over the diurnal cycle. The original GEMS AOD exhibited an hourly-dependent bias-characteristic-. It formed a U-797 shape, with a minimum value near noon, at 03:45 UTC. However, with the implementation of the model-enforced post-798 processing correction, the diurnal bias was mitigated effectively mitigated. This resulted in a bias value close to zero throughout 799 the day and a-decreased standard deviation. Figure 8c illustrates the diurnal variation in the bias of a-low AOD (AERONET 800 AOD < 0.4). The GEMS AOD (red circles) exhibited exhibits a positive bias of ~0.1. It was mostly mainly corrected to values 801 close to zero after post-processing (blue circles). However, certain a positive bias was observed at approximately

802 22:45 and 23:45 UTC, and at-06:45 and 07:45 UTC. Figure 8d shows the diurnal variation in the bias of high AOD (AERONET 803 AOD > 0.4). The diurnal variation in GEMS AOD (red circles) shows a clear U-shaped pattern with a maximum negative bias 804 of approximately -0.2 at 0.3 UTC. However, after post-processing, the bias was still negative but less than -0.1, which is 805 significantly closer to zero. By incorporating the predicted error, we obtained obtain an improved GEMS AOD that considers 806 the uncertainties and biases inherent in the retrieval process. This approach helps reduce these biases, including a-low AOD 807 overestimation, high AOD underestimation, and artificial diurnal bias in near-real-time AOD retrievals. TheA reduction in 808 artifactual diurnal bias is crucial for ensuring the reliability of hourly GEMS AOD data. This-is because it eliminates time-809 dependent discrepancies and provides a more representative hourly aerosol distribution. Users can now rely on corrected 810 GEMS AOD data for various applications without being influenced by diurnal variations in the original measurements. 811 Variable importance analysis for the post processing correction of the GEMS AOD was conducted (Figure S3). GEMS AOD 812 was the most important variable, emphasizing its direct influence on the correction process. VZA and elevation exhibited high 813 importancethe influence of diurnal variations in the original measurements-However, their significance can be attributed not 814 only to their inherent properties but also to their role in conveying AERONET location-related information. Aerosol type 815 appeared to have less significance in the RF models. This result contrasted with the notable importance of GEMS UVAI and 816 VisAI. This discrepancy can originate from inaccurate aerosol type classification in the GEMS aerosol algorithm.

817

818 5.2 Single-scattering albedo

819 This section presents a comparison of the GEMS SSA at 443 nm with the AERONET SSA at 440 nm inover the entire GEMS 820 domain. The validation period and collocation criteria for the AERONET sites were identical to those for the GEMS AOD. 821 Similar to the AOD, when a particular aerosol type in the GEMS was detected for over 90% of the pixels within a 0.25° 822 radius, we performed aerosol-type validation. Figure 9 and Table 3 displayshow the validation results for all pixels and each 823 aerosol type. The statistics Statistics, including N values and percentages are, were within the expected error ranges (0.03 and 824 0.05). The uncertainty of SSA is ±0.03 when AERONET AOD 440 nm is over 0.4 (Dubovik et al., 2002). The gray dashed 825 lines indicate an uncertainty envelope of ± 0.03 in SSA, whereas the black dashed lines indicate an uncertainty envelope of 826 ±0.05 in SSA. These reference lines help to assess the agreement between the GEMS SSA and AERONET data within a 827 reasonable error range. When aerosols are not abundant in the atmosphere, capturing Capturing SSA signals from satellite 828 observations is challenging- when atmospheric aerosols are not abundant. Therefore, for validation, separate analyses were 829 conducted for the cases where the GEMS AOD was > 0.4 (indicated by the red open circles) and the GEMS AOD was > 1.0830 (as indicated by the blue open circles). Notwithstanding the largesignificant uncertainties associated with the satellite 831 measurements, the GEMS aerosol product showed-a good overall agreement with the AERONET SSA. When GEMS AOD 832 exceeds 0.4, the percentage of GEMS SSA within the expected error range of ± 0.03 is denoted by 42.7634.22%, and that within 833 the expected error range of ± 0.05 is denoted by $\frac{67.25}{51.38\%}$. When the aerosol signal is strong (when GEMS AOD exceeds 834 1.0), the percentage of GEMS SSA within the expected error of ± 0.03 (0.05) increases to $\frac{56.61\%}{(83.7048.85\%)}$ (84.48%). 835 However, the percentage percentages within the expected error range and scatter plots varied depending on the aerosol type. 836 For the HAF type, the SSAs showed exhibited the largest spread. This indicated indicates a lower accuracy. It was This is likely 837 to be a result of an-ineffective aerosol-type selection (red circles). However, when the AOD exceeds exceeded 1.0 (blue circles), 838 these tend, it tended to approach the 1:1 line, (blue circles). Moreover, the percentage falling within the expected error range 839 of ± 0.03 increases significantly. For the dust type, the GEMS SSA exhibited a positive bias of approximately 0.04 compared 840 with the AERONET SSA (red circles). Similarly, when the AOD exceedsexceeded 1.0, these biases decreased, 841 approaching the 1:1 line (blue circles). However, the systematic bias observed in the GEMS SSA for the dust type indicates 842 the need to refine the assumed dust AOPs in the LUT. The NA type in the GEMS was observed to have a significantly low lower 843 variability compared with the AERONET SSA. The GEMS SSAs showed values close to one compared with the

- AERONET data. According to Lee et al. (2010), the NA type is identified when the SSA is above 0.95. However, many NAtype pixels were observed, with AERONET SSA values below 0.95 in the NA type. This indicates potential inaccuracies in the classification of the absorbing and NA GEMS aerosol types. <u>NeverthelessHowever</u>, when the AOD iswas high (blue circles), these classification errors tend tended to decrease. This results resulted in values closer to those of the AERONET SSA.
- 848 Figure <u>\$4\$6</u> and Table 4 present the hourly SSA validation results and statistic metrics, including the N and percentage 849 within the expected error range of ± 0.03 (± 0.05). The GEMS and AERONET SSA exhibited varying distributions over time. 850 The difference between the GEMS and AERONET SSA was most significant at 03:45 UTC and 04:45 UTC, with a positive 851 bias. This difference decreased at 22:45 and 23:45 UTC orand at 05:45 and 06:45 UTC (Figure S4S6). Similar to the GEMS 852 AOD, the GEMS SSA showed exhibited diurnal variations. These values are also reflected in the EE% values shown in Table 853 4. At 22:45 and 23:45 UTC, the percentage within the expected error range of ± 0.03 exceeded $\frac{60\%-64}{100}$. However, it 854 reduced decreased to less than 3019% at 03:45 UTC and 23% at 04:45 UTC before increasing again. Further studies are 855 required to understand the bias and accuracy variations in the SSA and improve the retrieval results. This can also be 856 attributed to the shorter path length in the observation geometry when the influence of the surface reflectance increases, 857 similar to that in AODs.
- 858 Figure 10a presents the comparison results for the GEMS SSA after post-process correction and the AERONET data. The 859 near real time post-process correction utilized data from the preceding 30 days for training. The validation period was from 860 DecemberNovember 1, 2021, to October 31, 2022. Notably, all-the statistical metrics demonstrated improvements. 861 Specifically, the percentage of GEMS SSA falling within the expected error range of ± 0.03 was recorded at 68.3354%, 862 whereas the percentage within the range of ± 0.05 was indicated at 88.8695%. Furthermore, the SSA-values exhibited a closer 863 alignment with the 1:1 line. Figure 10b depicts the difference between the GEMS and AERONET SSA over the 864 measurement time. Notably, the bias pattern observed in the GEMS SSA exhibits exhibited artifactual characteristics, thereby 865 forming a bell-shaped curve. In particular, during the time interval from 01:45 to 05:45 UTC, the mean bias of GEMS SSA 866 consistently surpassed the expected error range of ± 0.03 . However, the implementation of model-enforced post-867 processprocessing correction was demonstrated to be highly effective in mitigating this artificial diurnal bias. This correction 868 methodology resulted in a significant improvement insignificantly improved the GEMS SSA values within the expected 869 error range. Thereby Therefore, it enhanced the overall accuracy of the SSA retrieval. Variable importance analysis for the 870 post-processing correction of the GEMS SSA was conducted (Figure S5). The GEMS SSA was the most important variable 871 in the correction process. The GEMS AOD also emerged as a highly influential variable in the RF models for GEMS SSA 872 post process correction. Also, aerosol types appeared to have relatively lower significance within the RF models for SSA 873 correction.
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875 5.3 Aerosol layer height

876 From November 1, 2021, to October 31, 2022, the GEMS and CALIOP data were co-located for comparison. In this section, 877 the level-2 aerosol extinction coefficients at 532 nm arewere used to calculate the CALIOP ALH. This is shown in Equation 878 expressed as (2-). GEMS ALH pixels within a 0.05° radius surrounding each CALIOP pixel were averaged and compared with 879 the CALIOP ALHs within a time window of 1 h from of the GEMS observation time. The validation Validation was conducted 880 when the GEMS AOD values were largergreater than 0.2. This wasis because the error in ALH retrieval increased when the 881 presence of aerosols in the atmosphere was insufficient. Figure 11a shows a histogram of the differences between the GEMS 882 and CALIOP ALH. The total co-located number of data is 77,318, and the mean difference is -0.225 km. The median of 883 differences is differences is differences follows a Gaussian distribution of the differences follows a Gaussian 884 distribution, although it is skewed marginally skewed in athe positive direction. Figure 11b shows a comparison between the

925 GEMS and CALIOP ALH. These were distributed predominantly at altitudes <u>of</u> less than 2 km. The percentage of data falling 926 within the expected error of ± 1 km was 55.3%, and the percentage falling within the expected error of ± 1.5 km was 71.7%. 927 The variability <u>ofin</u> the GEMS ALH was comparable to that of the CALIOP ALH.

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929 <u>5.4 Limitations of the current GEMS AOPs</u> and future work

930 Figure S7 shows seasonal and regional variations as a function of UTC for each of the following four regions: Korea (33° N-931 39° N and 124° E–132° E), North China (33° N–34° N and 110° E–124° E), South China (21° N–33° N and 110° E–122° E), 932 Indochina Peninsula (8° N–22° N and 92° E–110° E). The Indian region was excluded from the regional analysis because the 933 observable area within the entire region of India varied significantly depending on the GEMS scan profiles. After gridding the 934 GEMS AOPs into a $0.1^{\circ} \times 0.1^{\circ}$ grid box, monthly averages were calculated. After monthly averaging, seasonal averages were 935 calculated for each pixel only when all three months within a season had data available for the given pixel. Regional averages 936 were calculated when more than 50% of the available values were within the domain. For the AOD, U-shaped or flat diurnal 937 variations were observed in all four regions. In the case of SSA, higher values were observed during June, July, and August 938 (JJA) in Korea, North China, and South China, which are considered to be influenced by aerosol hygroscopic growth owing 939 to relatively high atmospheric humidity. However, the Indochina Peninsula showed the highest SSA values in SON (September, 940 October, and November) and the lowest values in DJF (December, January, and February), which is consistent with the 941 relatively low SSA values observed at the Chiang Mai AERONET site from 2011 to 2016 during DJF (Liang et al., 2019). 942 However, there are limitations to the investigation of diurnal variations in ALH. The diurnal variations in the ALH were not 943 consistent with the diurnal variations in the mixing layer height. One reason for the uncertainty in the ALH is that it is retrieved 944 from the OE depending on the uncertainty of the a priori AOD, SSA, and ALH. Before post-processing, GEMS AOD and SSA 945 exhibited diurnal bias patterns compared to the AERONET data (details in Sections 5.1 and 5.2). These uncertainties affect the 946 uncertainty in the diurnal variation of ALH. Because the GEMS ALH cannot be post-corrected using CALIOP data (details in 947 Section 3), we are considering post-process-corrected ALH using ground-based lidar observation networks (i.e., the Korea 948 Aerosol Lidar Observation Network and the Asian dust and aerosol lidar observation network) in future studies. Therefore, 949 one of the limitations of this study is that the GEMS ALH has limitations in the detailed investigation of diurnal variations in 950 ALH.

Several methods can be employed to improve the results of the GEMS aerosol algorithm. First, additional satellite data could
 be integrated for cloud detection. Incorporating data from other satellite sensors with IR channels, such as the AMI, can provide
 complementary information for cloud masking. Secondly, it is necessary to consider the AOPs used in the LUT to improve the
 GEMS aerosol algorithm. It is essential to incorporate additional ground-based observations in the UV region, such as those
 from the Pandora Instrument and SKYNET. Collecting ground-based observations in the UV region and incorporating them
 into the LUT can enhance the algorithm's performance. Finally, regional LUTs with data from diverse regions that consider
 variability in AOPs based on regional characteristics are crucial.

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959 <u>6 Summary</u>

AERAOD algorithm was developed using OMI as <u>the input data before the GEMS</u> launch, modifications were made considering theto consider GEO observation characteristics during the IOT period. A new hourly surface reflectance database was created using the minimum reflectance method with <u>a fine spatial resolution-that aligned with the GEMS pixel resolution</u>. In addition, monthly BAOD maps were incorporated to estimate the hourly GEMS surface reflectance. <u>A newNew</u> cloud--

- removal techniques have significantly improved the effectiveness of cloud detection and enhanced the quality of aerosol
 retrievalsretrieval. To avoid discrepancies between the observed and simulated radianceradiances that may arise due tobecause
 of the monochromatic assumption of the LUT calculation, we applied a spectral binning approach to the LUT calculation.
 Finally, post-processprocessing correction methods based on machine learning were used to remove the non-physical diurnal
 biases in the AOD and SSA retrieval. This reduced the biases over time and provided more reliable hourly GEMS aerosol
 products in near-real-time.
- The GEMS aerosol product wasproducts were investigated for three specific events: dust events over Northeast Asia, biomass
 burning in Southeast Asia, and the absorption of aerosols over<u>absorbing aerosol in India</u>. These events highlight the capability
 of the GEMS to monitor and provide insights into aerosol properties during various atmospheric events while also emphasizing
 the importance of post-processing for data accuracy and agreement with ground measurements.
- 1015 The GEMS aerosol products were validated against the <u>AERONET and CALIOP data for the entire GEMS domain for one</u> 1016 year (from November 2021 to October 2022). The performance of the GEMS aerosol algorithm was validated to verify its 1017 applicability forin studying the distribution of AOPs across Asia. The validation results for each product are summarized below:
- 1018 The GEMS AOD showshowed a good correlation with the AERONET AOD (R = 0.792). However, it exhibits certain 1019 biasspecific biased patterns- were observed. Notably, anthe underestimation of AOD in high AERONET AOD and the 1020 overestimation of AOD in low AERONET AOD occurred owing tobecause of cloud contamination. Different aerosol types 1021 exhibited varying validation results: the HAF type with the highest R and Q values; the dust type with underestimation in India 1022 but overestimation in Northeast Asia; and the NA type with cloud contamination issues, particularly for low AOD. This 1023 indicated indicates the need for an improvement in of the cloud - masking process, particularly over the ocean. Certain deviations 1024 beyond the error range of the GEMS AOD were observed in India and Beijing. The underestimation of the high AOD values 1025 can be attributed to the aerosol model. Diurnal variation variations in the retrieval performance waswere evident, with varying 1026 slopes and other comparison statistics throughout the day. AsBecause, the testbed for the GEMS algorithm was on the LEO 1027 platform, thea time-dependent retrieval bias hadwas not been observed previously observed. Therefore, we adopted a model-1028 enforced post-process correction and findfound that this enhancesenhanced GEMS AOD performance, reducing reduced the 1029 overall biases. This These corrected data ensures ensure the reliability for of various applications.
- The GEMS SSA at 443 nm was validated against-the AERONET SSA at 440 nm over the entire GEMS region. The GEMS SSA's agreement with the AERONET data was evaluated within a reasonable error range of ±0.03 (±0.05). For GEMS AOD exceeding 0.4, 42.76 (67.25)% of GEMS SSA is within ±0.03 (0.05) error. This increases to 56.61 (85.70)%%) for the strong aerosol signals (GEMS AOD above> 1.0). However, the accuracy varies varied among the aerosol types. The HAF type has / aexhibited higher variability and lower accuracy. The dust type hashad a marginal positive bias, particularlymainly when the / AOD iswas high. Similar to the AOD, the post-processprocessing correction for the GEMS SSA data yielded significant / enhancements in the statistical metrics.
- The GEMS and CALIOP data were then compared. The GEMS ALH was compared with the CALIOP ALH when the GEMS
 AOD exceeded 0.2. The results showed a mean difference of -0.225 km, with 55.29% of data being within ±1 km and 71.70%
 being within ±1.5 km. The GEMS ALH exhibited variability similar to that of CALIOP ALH.
- Several methods can be used to further, improve the results of the GEMS acrossed algorithm. First, additional satellite data could
 be integrated for cloud detection. Incorporating data from other satellite sensors with IR channels such as the AMI can provide
 complementary information for cloud masking. Second, it is necessary to consider the AOPs used in the LUT to improve the
- 1043 GEMS aerosol algorithm. It is particularly important to incorporate more ground based observations in the UV region, such
- 1044 as those from the Pandora Instrument and SKYNET. Collecting ground based observations in the UV region and incorporating

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1078 these into LUT can enhance the performance of this algorithm. Finally, regional LUTs with data from diverse regions that 1079 consider the variability in AOPs based on regional characteristics are crucial. Overall, the improvements toin the GEMS aerosol 1080 algorithm contributehave contributed to advancing our understanding of aerosol properties and their effects on the environment. 1081 Thereby Therefore, it provides valuable information for diverse applications, including air quality monitoring, air quality data 1082 assimilation, and health impact assessments in Asia. 1083 1084 Code availability. The GEMS L2 AERAOD algorithm is not available publicly. 1085 1086 Data availability. GEMS L2 AERAOD wasdata were downloaded from the Environmental Satellite Center website 1087 (https://nesc.nier.go.kr/en/html/datasvc/index.do). 1088 1089 Author Contribution. YC, JK, SG, and MK designed the experiments. WL and DL provided support for the-data collection. 1090 SL, MK, HC, OT, and SP contributed to the algorithm development. YC wrote the manuscript with contributions from all the 1091 co-authors. JK reviewed and edited the manuscript. JK provided support and supervision. All-the authors analyzed the 1092 measurement data and prepared the manuscript. 1093 1094 Competing Interests. At least one of the (co-)authors is a member of the editorial board of Atmospheric Measurement 1095 Techniques. 1096

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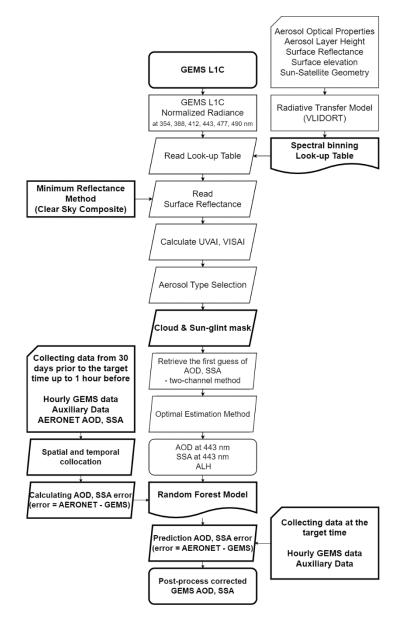
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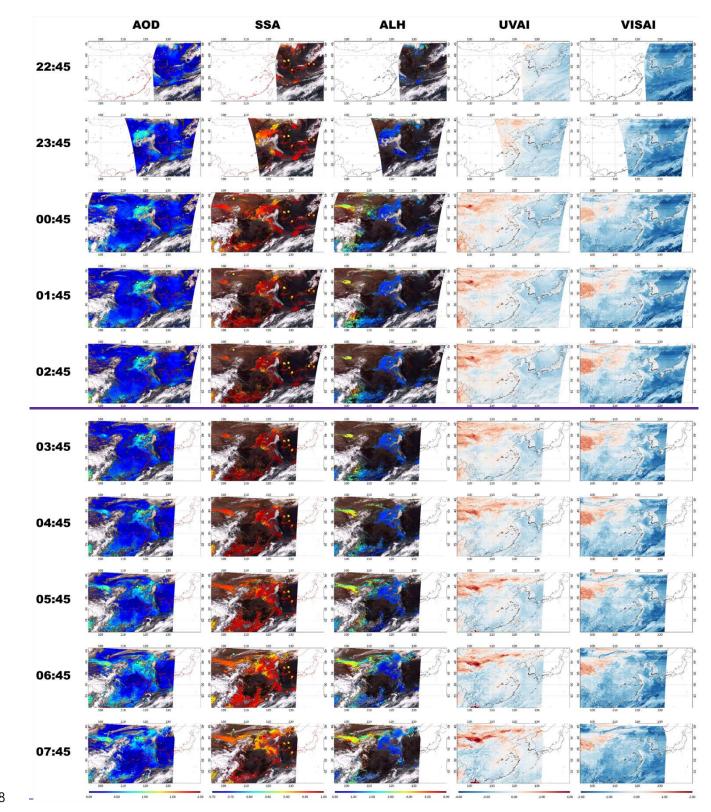
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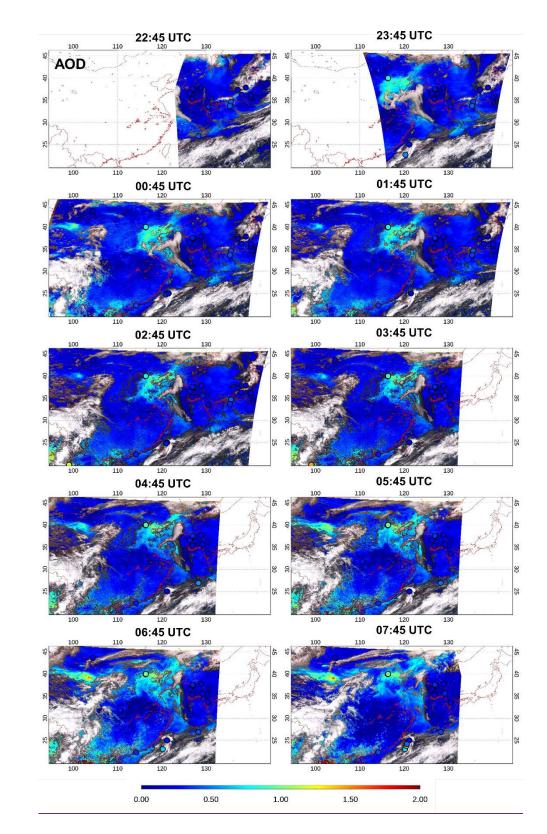
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1386 Figure 1: The flowchart of the GEMS AERAOD retrieval algorithm and the modifications in the study (in bold boxes)

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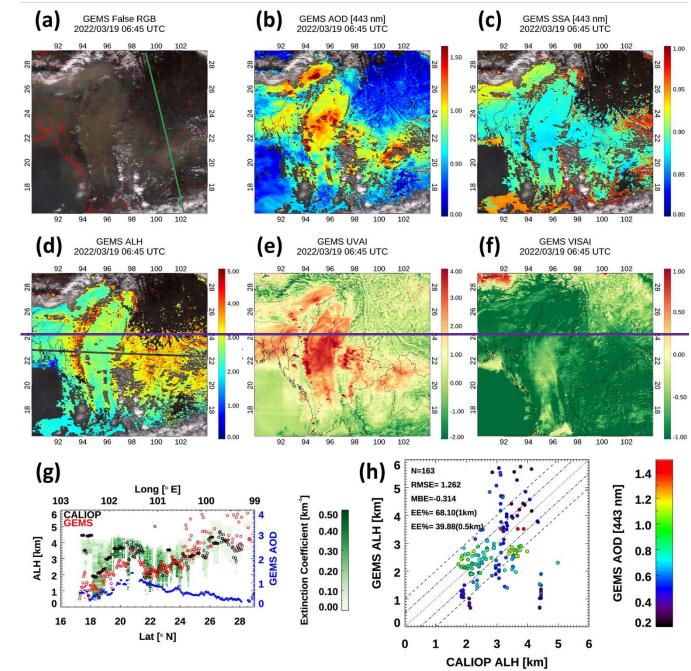


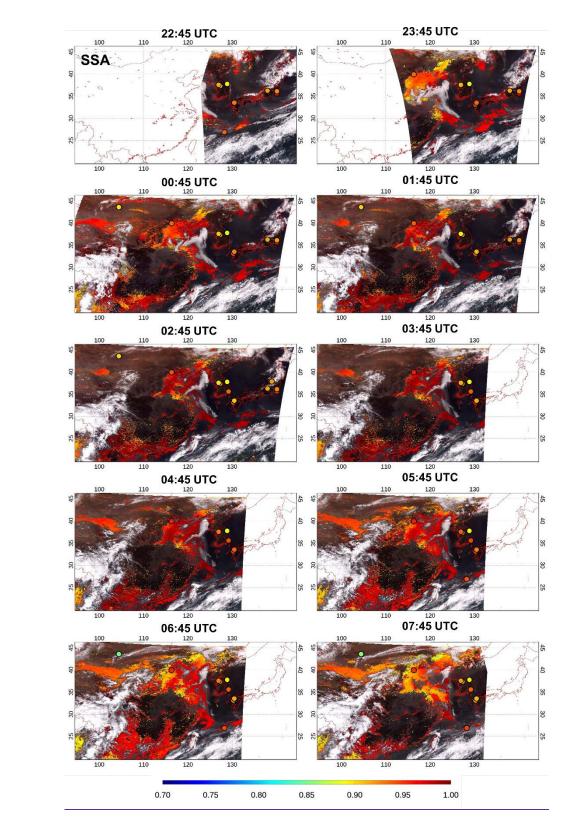


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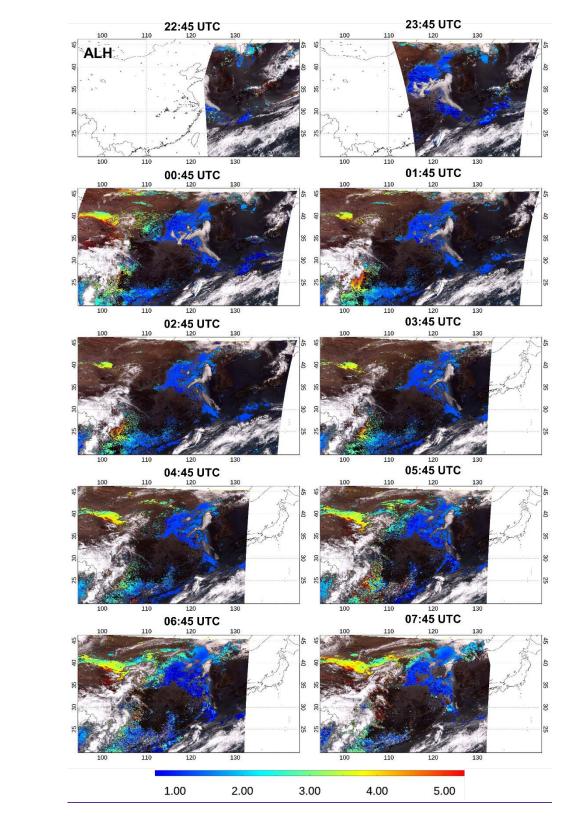
1890Figure 2: Hourly GEMS aerosol products for the dust case on April 8, 2022 over northwestern China. Time-series maps of AOD at1891443 nm, SSA, ALH, UVAI at 443 nm, and VISAIALH (km) from 22:45 UTC to 07:45 UTC. The circle denotes an AERONET1892station, and the filled color indicates the AERONET AOD and SSA at 443 nm in the AOD and SSA columns. GEMS SSA, and ALH

1393 are displayed only when GEMS AOD > 0.2.

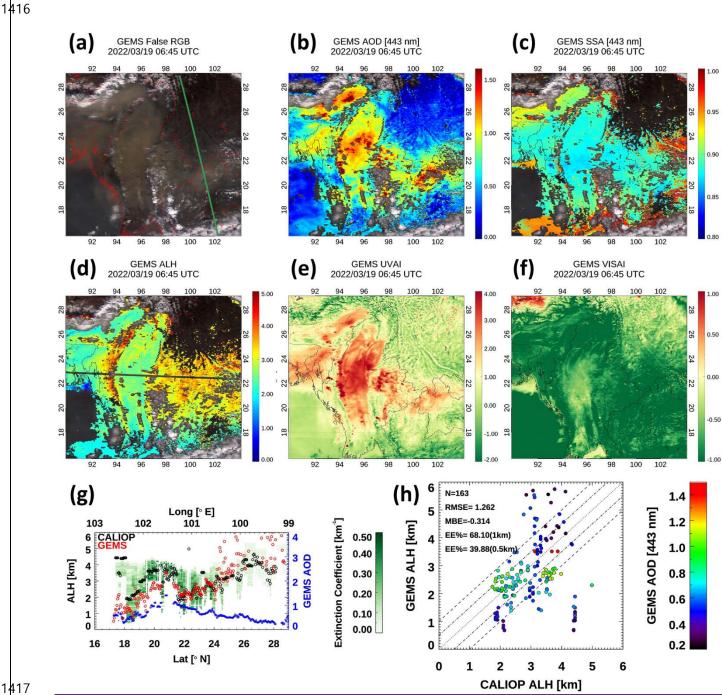




1398Figure 2: Continue



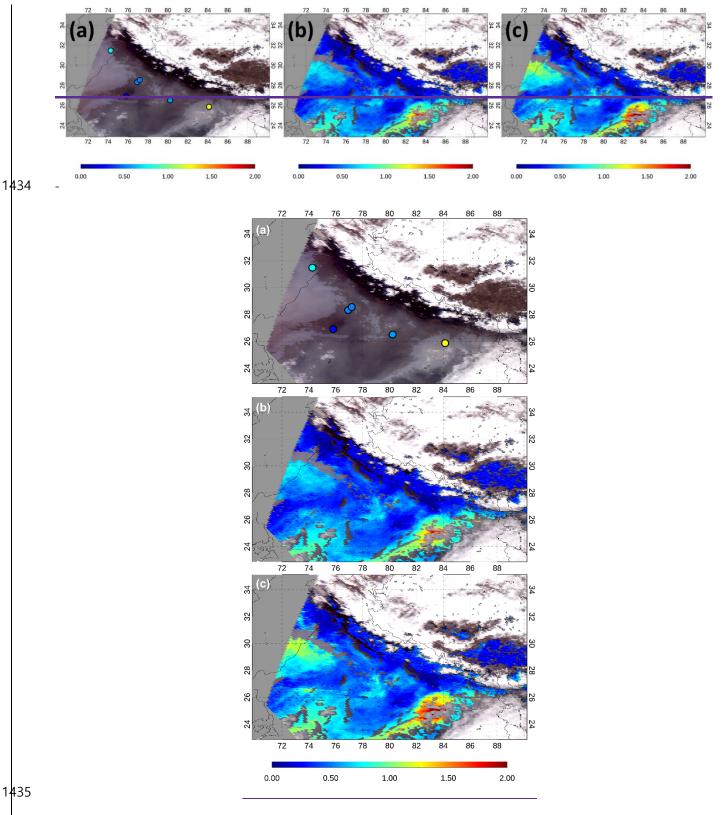
1400 <u>Figure 2: Continue</u>



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1418 Figure 3: The example of GEMS aerosol products for biomass burning over mainland Southeast Asia. The maps of- (a) GEMS 1419 False RGB, (b) AOD, (c) SSA, (d) ALH, (e) UVAI, and (f) VisAI. The green line in GEMS False RGB indicates the overpass path of 1420 CALIOP. The GEMS SSA and ALH are displayed only when the GEMS AOD is over 0.2. (g) GEMS ALH compared with CALIOP 1421 extinction coefficient in the domain. The background color represents the CALIOP extinction coefficient. The black open circles 1422 denote the CALIOP ALH, whereas the red open circles represent the GEMS ALH. The blue squares represent the GEMS AOD. (h) 1423 Comparison of GEMS and CALIOP ALH when GEMS AOD > 0.2. The dashed and dash-dotted lines indicate an uncertainty 1424 envelope of ±1 km and ±0.5 km in ALH, respectively. The dotted lines represent the 1:1 line. The color in the circles represents the 1425 GEMS AOD.

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1436 Figure 4: The example of the GEMS AOD before and after post-processing for an absorbing aerosol case over Indo-Gangatic Plane 1437 at 04:45 UTC on December 4, 2021. (a) GEMS false RGB. The circle denotes an AERONET station, and the filled color indicates the AERONET AOD at 443 nm, (b) GEMS AOD_a and (c) GEMS AOD after post-process correction. 1438

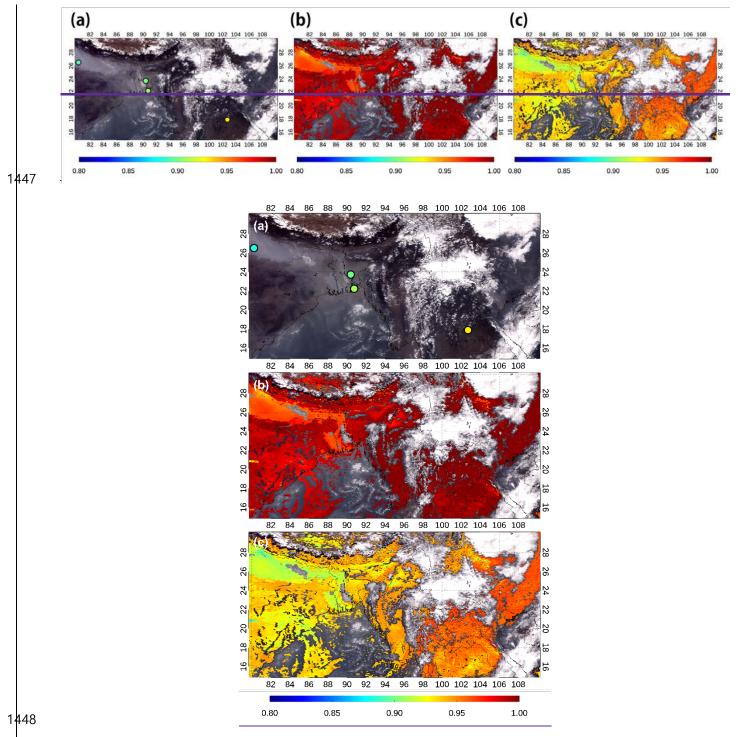
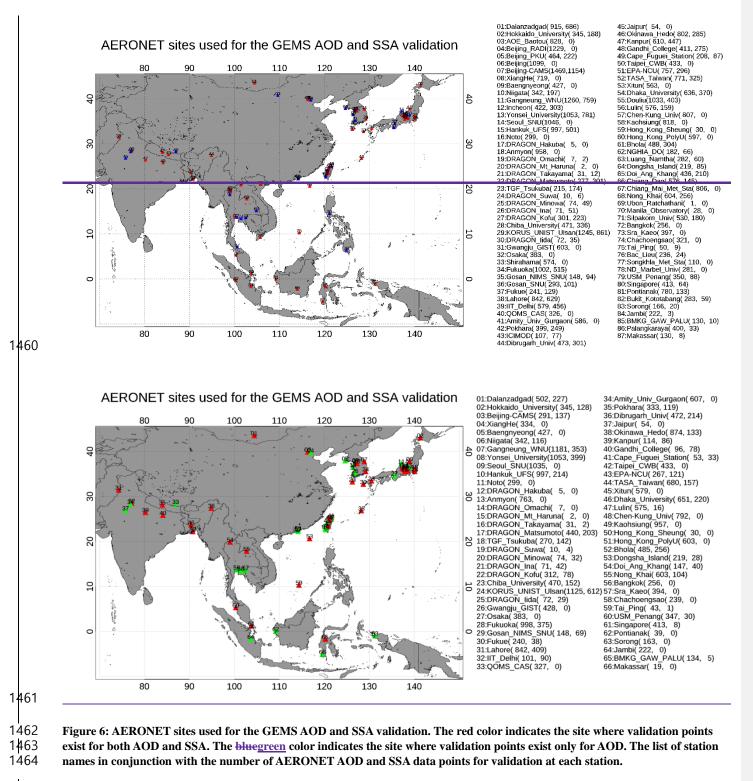
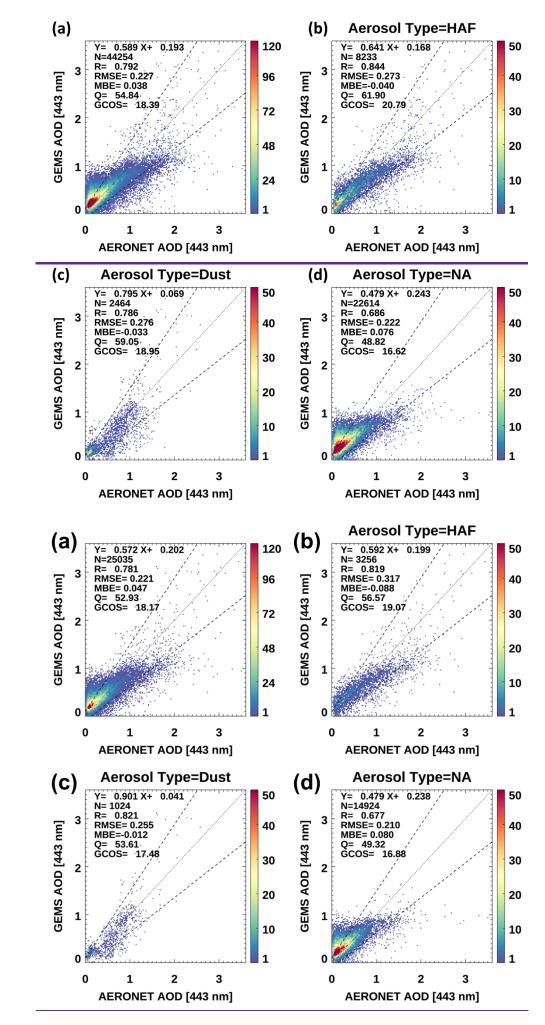


Figure 5: The example of GEMS SSA and the GEMS SSA after post-processing for an absorbing aerosol case over India,
 Bangladesh, and mainland Southeast Asia at 03:45 UTC on December 23, 2021. (a) GEMS false RGB. The circle denotes an
 AERONET station, and the filled color indicates the AERONET SSA at 440 nm, (b) GEMS SSA, and (c) GEMS SSA after post process correction.

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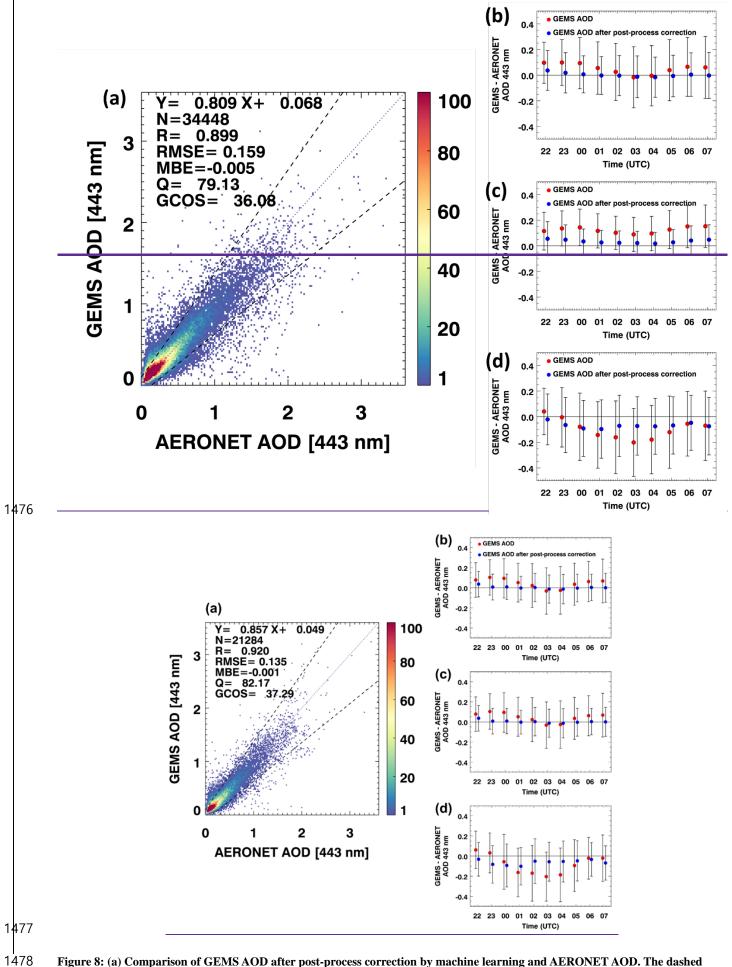


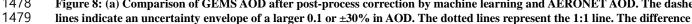
1472	Figure 7: Comparison of GEMS and AERONET AOD for (a) total and individual aerosol types: (b) HAF, (c) dust, and (d) NA. The
1473	dashed lines indicate an uncertainty envelope of maximum (0.1 or 30%) in AOD. The dotted lines represent the 1:1 line. Data from
1474	November 1, 2021, to October 31, 2022, are used for comparison.

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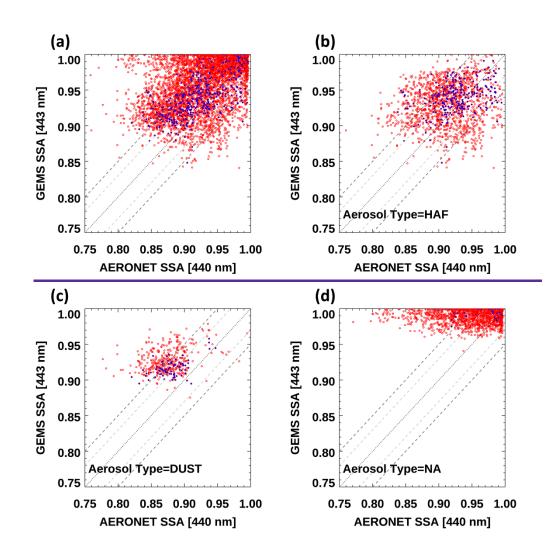
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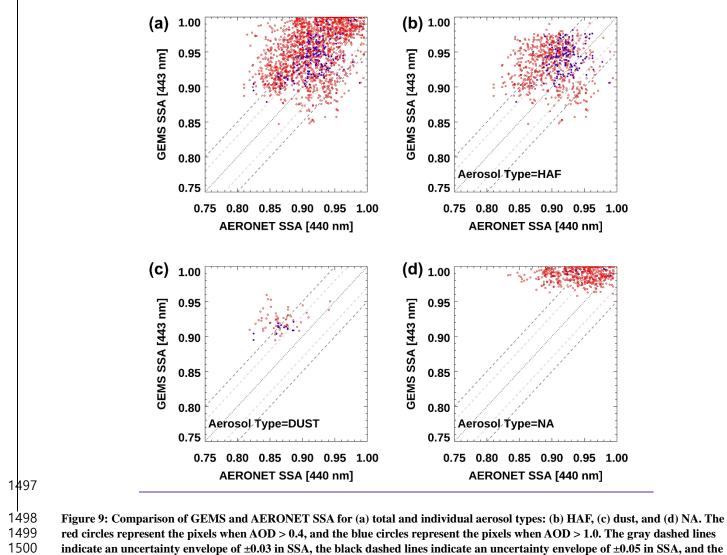
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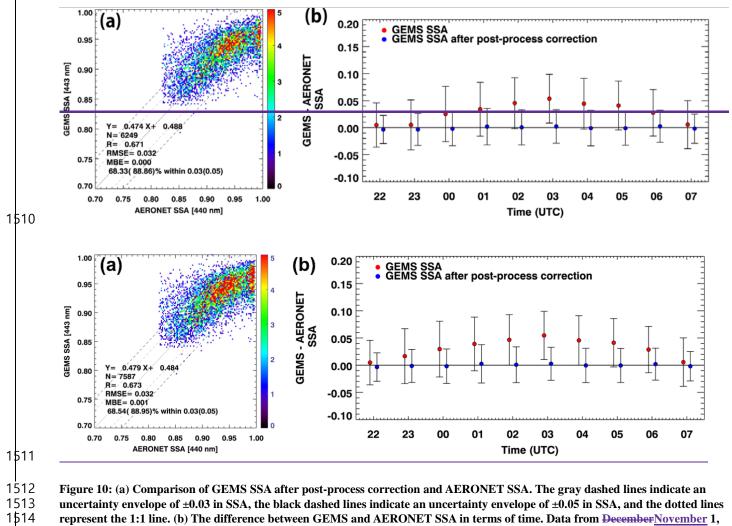
1485	between GEMS AOD and AERONET AOD in terms of time. (b) All pixels, (c) pixels when AERONET AOD < 0.4, and (d) pixels
1486	when AERONET AOD > 0.4. The red circles represent the GEMS AOD, and the blue circles represent the GEMS AOD after post-
1487	process correction. The error bars correspond to the standard deviation. Data from December November 1, 2021, to October 31,
1488	2022 ₁ are used for comparison.
1489	<u>ــــــــــــــــــــــــــــــــــــ</u>



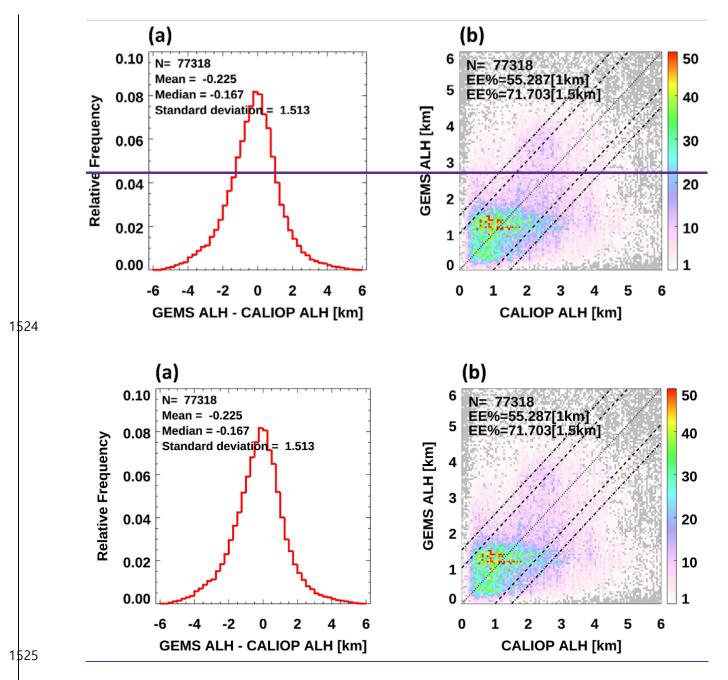


1501 dotted lines represent the 1:1 line. Data from November 1, 2021, to October 31, 2022, are used for comparison.

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uncertainty envelope of ±0.03 in SSA, the black dashed lines indicate an uncertainty envelope of ±0.05 in SSA, and the dotted lines represent the 1:1 line. (b) The difference between GEMS and AERONET SSA in terms of time. Data from DecemberNovember 1, 1515 2021, to October 31, 2022, are used for comparison.



1526Figure 11: (a) Histogram of difference between GEMS and CALIOPALH and (b) comparison of GEMS and CALIOPALH. The1527dashed lines indicate an uncertainty envelope of ±1 km in ALH. The dash-dotted lines indicate an uncertainty envelope of ±1.5 km1528in ALH. The dotted lines represent the 1:1 line. Data from November 1, 2021, to October 31, 2022, are used for comparison.

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Table 1: Dimension of LUT in GEMS Aerosol algorithm.

	Number		1
Variable Name [Unit]	of	Entries	4
	Entries		
Wavelength [nm]	6	354, 388, 412, 443, 477, 490	
SZA [°]	12	0.01, 5, 10, 15, 20, 27, 34, 41, 48, 55, 62, 69	
VZA [°]	12	0.01, 5, 10, 15, 20, 27, 34, 41, 48, 55, 62, 69	
RAA [°]	11	0.01, 15, 30, 45, 60, 80, 100, 120, 140, 160, 180	
Surface reflectance [-]	4	0.0, 0.05, 0.1, 0.2	
AOD at 443 nm [-]	8	0.0, 0.1, 0.4, 0.8, 1.5, 2.0, 2.8, 3.6, 5.0, 10.0	
SSA of 442 mm []	8	1.0, 0.98, 0.96, 0.94, 0.91, 0.88, 0.85, 0.82 for HAF and Dust	
SSA at 443 nm [-]	0	1.0, 0.99, 0.98, 0.97, 0.96, 0.94, 0.92, 0.90 for NA	
ALH above the surface [km]	5	0.5, 1.5, 3.0, 4.5, 6.0	
Elevation [km]	3	0, 3, 6	

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1543 Table 2: Statistic of hourly comparison of GEMS and AERONET AOD in Figure <u>S2S5</u>.

Time	Ν	Slope	y-intercept	R	RMSE	MBE	Q (%)	GCOS (%)
22:45	925<u>801</u>	0. 730<u>725</u>	0. 180 177	0. 715 738	0. 188 <u>181</u>	0. <u>100</u> 094	<u>58.3860.42</u>	23. 24 <u>.97</u>
23:45	<u>19641413</u>	0. <u>684728</u>	0. 190 193	0. 830 752	0. 212 187	0. 076<u>115</u>	59.32 <u>53.93</u>	20.93<u>19.89</u>
00:45	4 593 2879	0. <u>584600</u>	0. 217 221	0. 767<u>698</u>	0. 224 218	0. 088 <u>112</u>	51<u>48</u>.32	16.74<u>15.56</u>
01:45	5632 <u>3345</u>	0. <u>534490</u>	0. 200 211	0. 774<u>715</u>	0. 211<u>209</u>	0. 054<u>063</u>	<u>54.8352.68</u>	17.47<u>16.95</u>
02:45	<u>64003718</u>	0. 555<u>5</u>33	0. 183 193	0. 795 780	0. 221 214	0. 029<u>039</u>	<u>54.5352.66</u>	18.55 <u>17.86</u>
03:45	<u>61393504</u>	0. 569 <u>577</u>	0. <u>165171</u>	0. <u>824830</u>	0. 233 238	- 0. 013<u>011</u>	56.5 4 <u>53.48</u>	17.04<u>16.67</u>
04:45	<u>61573556</u>	0. 593<u>592</u>	0. 169<u>176</u>	0. <u>822</u> 824	0. 230 238	<u>-</u> 0. 000 <u>001</u>	55.19 <u>53.12</u>	18.16<u>17.97</u>
05:45	<u>56423186</u>	0. <u>586518</u>	0. 204 233	0. 773 <u>725</u>	0. 235<u>043</u>	0.041043	<u>52.8750.00</u>	19.25 <u>18.33</u>
06:45	4 261 2117	0. <u>647606</u>	0. 218 241	0. 794<u>766</u>	0. 233<u>239</u>	0. 065<u>069</u>	<u>54.8952.01</u>	19.4 <u>679</u>
07:45	<u>25411299</u>	0. <u>617</u> 632	0. 224 227	0. 771 754	0. 247 245	0. 054 063	<u>56.55</u> 54.89	19.4 <u>886</u>

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1548Table 3: Comparison of GEMS and AERONET SSA for different aerosol types in Figure 9. N represents the number of data, and1549EE% denotes the percentage within the expected error range of ±0.03 (±0.05).

	C	SEMS AOD > 0.4		GEMS AOD > 1.0
Aerosol Type	Ν	EE% ±0.03 (±0.05)	Ν	EE% ±0.03 (±0.05)
All	<u>5227</u> 1841	4 2.76(67.25 <u>34.22(61.38</u>)	<u>454174</u>	56.61(83.70<u>48.85(84.48</u>)
HAF	1559 764	4 1.95(70.24 <u>31.68(62.43</u>)	277 136	61.01(87.73 54.41(89.71)
Dust	<u>43771</u>	20.37(50.57<u>12.68(45.07</u>)	<u>8215</u>	39.02(73.17 <u>13.33(66.67</u>)
NA	1850<u>536</u>	4 <u>5.14(65.62</u> <u>32.46(56.72</u>)	31 7	51.61(70.97<u>4</u>2.86(57.14)

1550

1551 Table 4: Statistic of comparison of GEMS and AERONET SSA in Figure <u>S4S6</u>.

	GEMS AOD > 0.4			GEMS AOD > 1.0		
Time	Ν	EE% ±0.03 (±0.05)	Ν	EE% ±0.03 (±0.05)		
2:45	<u>13749</u>	64.96 (86.13<u>6</u>7.35(89.80)	<u>2313</u>	52.17 (86.96<u>61.54(92.31</u>)		
3:45	288 <u>76</u>	60.76 (83.68<u>6</u>4.47(82.89)	<u>6718</u>	74.63 (92.54<u>77.78(94.44</u>)		
0:45	4 <u>20100</u>	57.6 2 (82.38.00(87.00)	93<u>21</u>	73.12 (88.17<u>90.48(100.00</u>)		
)1:45	4 <u>5</u> 4 <u>138</u>	56.61 (79.07<u>57.25(81.16</u>)	<u>11329</u>	63. 72 (88.50.41(96.55)		
)2:45	<u>655190</u>	39.69 (62.90<u>31.58(56.84</u>)	237<u>72</u>	4 5.99 (73.00<u>31.94(56.94</u>)		
)3:45	<u>859</u> 391	27.82 (53.20<u>18.67(44.76</u>)	339 206	25.07 (57.23<u>15.05(46.60</u>)		
4:45	<u>822406</u>	28.22 (55.60).41(52.46	335 209	27.76 (62.39<u>23.44(58.85</u>)		
)5:45	<u>621223</u>	36<u>30.49(61</u>.88 (63.12)	222 94	38.29 (67.57<u>28.72(65.96</u>)		
)6:45	<u>620175</u>	48.23 (73.23 <u>37.14(69.71</u>)	<u>25583</u>	51.37 (77.65 40.96(75.90)		
)7:45	<u>35193</u>	60.68 (79.49<u>53.76(73.12</u>)	160 46	63.12 (84.38<u>5</u>4.35(76.09)		

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