| 1      | Role of time-averaging of eddy covariance fluxes on water use efficiency                                 |
|--------|--|
| 2      | dynamics of Maize crop   |
| 3<br>4 | Arun Rao Karimindla, Shweta Kumari, Saipriya SR, Syam Chintala*, and BVN Phanindra<br>Kambhammettu       |
| 5      | Department of Civil Engineering, Indian Institute of Technology Hyderabad, Telangana, India.             |
| 6      | *Corresponding author: E-Mail: ce22resch11012@iith.ac.in; Tel: +91 7997014429                            |
| 7      | Abstract   |
| 8      | Direct measurement of carbon and water fluxes at high frequency makes eddy                               |
| 9      | covariance (EC) as the most preferred technique to characterize water use efficiency (WUE).              |
| 10     | However, reliability of EC fluxes is largely hinged on energy balance ratio (EBR) and inclusion          |
| 11     | of low-frequency fluxes. This study is aimed at investigating the role of averaging period to            |
| 12     | represent EC fluxes and its propagation into WUE dynamics. Carbon and water fluxes were                  |
| 13     | monitored in a drip-irrigated Maize field at 10 Hz frequency and are averaged over 1, 5, 10,             |
| 14     | 15, 30, 45, 60, and 120 minutes considering daytime unstable conditions. Optimal averaging               |
| 15     | period to simulate WUE fluxes for each growth stage is obtained by considering cumulative                |
| 16     | frequency (ogive) curves. A clear departure of EBR from unity was observed during dough and              |
| 17     | maturity stages of the crop due to ignorance of canopy heat storage. Error-Deviation in                  |
| 18     | representing water (carbon) fluxes relative to the conventional 30 min average is within $\pm$ 3 %       |
| 19     | (± 10 %) for 10-120 min averaging and is beyond ± 3 % (± 10 %) for other time-averages.                  |
| 20     | Ogive plots conclude that optimal averaging period to represent carbon, water and <u>WUE</u> water       |
| 21     | fluxes is 15-30 min for 6 <sup>th</sup> leaf and silking stages, and is 45-60 min for dough and maturity |
| 22     | stages. Dynamics of WUE considering optimal averaging periods are in the range of 1.49 $\pm$             |
| 23     | 0.95, $1.37\pm0.74,1.39\pm0.79,and3.06\pm0.69\mu mol\ mmol^{-1}$ for the 6th leaf, silking, dough, and   |
| 24     | maturity stages respectively. Error in representing WUE with conventional 30 min averaging               |
| 25     | is marginal (< 1.5 %) throughout the crop period except for the dough stage (12.12 %). We                |
| 26     | conclude that the conventional 30 min averaging of EC fluxes is not appropriate for the entire           |
| 27     | growth stage. Error in representing WUE with conventional 30 min averaging is marginal (<                |

1.5 %) except during the dough stage (12.12 %). Our findings can help in developing efficient

water management strategies by accurately characterizing WUE fluxes from the EC

30 measurements.

28

29

Keywords: Eddy covariance, Maize crop, Time-average, Energy balance ratio, Ogive
function, Water use efficiency.

## 33 Research Highlights:

- The time-averages that yield the most effective energy balance closure are identified as
   45 and 60 minutes.
- Insufficiently short time-averages such as 1 and 5 minutes, as well as excessively long time-averages such as 120 minutes, resulted in a high relative error in representing
   carbon and water fluxes.
- 39 3. The conventional 30-minute averaging period proved to be insufficient in capturing
  40 low-frequency fluxes, necessitating the use of longer averaging periods.
- 4. <u>4. Different time averaging periods are to be considered to compute the EC fluxes</u>
   <u>considering the crop growth stage.</u> <u>Time averaging of Eddy Covariance fluxes needs to</u>
   <u>be performed in accordance with crop growth stage.</u>
- 44

### 45 **1.0 INTRODUCTION**

Water use efficiency (WUE) is an important eco-hydrologic trait relating two important 46 47 processes of plant metabolism namely carbon fixation (via photosynthesis) and water consumption (via transpiration) (Bramley, 2013). The need for achieving food security with 48 diminishing water resources under changing climate has made WUE as the controlling 49 parameter in planning and design of irrigation strategies (Tang, 2015). Depending on the scale 50 of investigation, WUE can be quantified at: i) leaf, ii) plant, iii) ecosystem, or iv) regional 51 scales (Medrano, 2015). Of these, ecosystem WUE has taken precedence in irrigation and 52 agronomy due to: i) accurate and reliable measurement using micrometeorological techniques, 53 ii) ability to evaluate the role of various water conservation techniques on ecosystem 54 productivity, and iii) understand the relation between carbon and water cycles in response to 55 changes in climate (Tang, 2015; Tong, 2014). 56

Eddy covariance (EC) is a non-destructive, micrometeorological technique for direct measurement of water vapour (H<sub>2</sub>O) and carbon (CO<sub>2</sub>) fluxes between vegetation and atmosphere at high temporal frequency (Aubinet, 1999; Leclerc and Foken, 2014). EC method precisely measures the overall transfer of heat, mass, and momentum between the earth's surface (such as vegetation) and the atmosphere. This is achieved by estimating the covariance

of turbulent fluctuations in vertical wind (referred to as eddies) with respect to the specific flux 62 63 under consideration such as H2O, CO2, temperature. EC represents the scalar fluxes of interest (representative of eco-hydrological processes) from a region upwind of the measurement 64 known as the footprint. At ecosystem scale, WUE is estimated as the ratio of net primary 65 product (NPP: proxy for photosynthesis) to evapotranspiration (ET: proxy for water 66 67 consumption) WUE is estimated as the ratio of net primary product (NPP): proxy for photosynthesis to evapotranspiration (ET): proxy for water consumption (Peddinti, 2020). 68 WUE is a key eco-hydrologic trait that is used to analyse the role of climate change, drought, 69 70 deficit irrigation, and management strategies on ecosystem productivity. Currently, EC is the 71 most accurate and reliable method for estimating carbon and water fluxesexchnages, hence WUE at ecosystem scale (Tong, 2009). A number of studies have demonstrated the efficacy of 72 EC in estimating WUE across a wide range of ecosystems (Tang, 2015; Tong, 2014; Wang, 73 2017). Error sources that affect the accuracy of EC fluxes are grouped into: i) Unrepresentative 74 75 (due to footprint heterogeneity, unsatisfied underlying theory), ii) Measurement uncertainties (due to random errors, interference and contamination, sensor drifts) and iii) Measurement 76 77 biases in fluxes (tilt, frequency losses, air density fluctuations etc). Despite improvements in 78 measurement accuracy, data sampling, and processing techniques, EC method still suffers from the drawback of lack of conservation among the energy terms, resulting in energy balance 79 closure (EBC) problem (Charuchittipan, 2014; Foken, 2011; Reed, 2018). Lack of EBC as 80 observed in EC system is reported across diverse ecosystems ranging from simple bare soils 81 82 (Oncley, 2007), to homogeneous grasslands (Twine, 2000), to heterogeneous croplands (Peddinti and Kambhammettu 2019), to complex forest ecosystem (Charuchittipan, 2014; 83 84 Wilson, 2002). Apart from the errors associated with instrumentation, measurement, and neglected energy sinks, lack of EBC at the EC sites is also attributed to the omission of low 85 frequency secondary circulations in the turbulent flux estimation (Wilson, 2002). This problem 86 can be circumvented by choosing appropriate averaging period during flux estimation, the 87 selection of which is based on: i) 'ensemble block time-averaging method' (Finnigan, 2003; 88 89 Malhi, 2004; Sakai, 2001), and ii) 'ogive method' (Berger, 2001).

A number of studies have highlighted the importance of averaging period in quantifying
the EC fluxes, with an objective to obtain optimal time-averaging period under various canopy
and surface roughness conditions. While smaller averaging periods (15-30 min) are suitable
for managed croplands, flux estimation from forest and tall canopies demand longer averaging
periods (60-120 min) due to the presence of large-sized, slow moving eddies (Finnigan, 2003;

Sakai, 2001; Sun, 2006). Zhang (2013) concluded that time-averaging of EC fluxes has to be 95 96 done in accordance with the observation scale. In an analysis of Chengliu riparian forest in China, they found that lower time-averaging periods (15 min) are suitable for daily variation 97 of EC fluxes, whereas higher time-averaging periods (60 min) are suitable for long-term flux 98 computations. A similar observation was made by Lee (2004) over farmlands. In a wheat field 99 in Yucheng, China, 10 min and 30 min averaging periods were found suitable for diurnal and 100 long-term flux observations respectively. Flux observations over a Maize crop at Daxing 101 experimental station in China conclude that optimal time-averaging period has to be considered 102 103 in accordance with crop growth stage (Feng, 2017). However, they observed a marginal (< 3 104 %) error in representing the fluxes at conventional 30 min averaging relative to the optimal averaging obtained for each growth stage. 105

Maize is the third most important cereal crop in India after rice and wheat, and accounts 106 107 for about 10 % of total food production in the country (Sharma, 2018; Ficci 2014). Inspite of a huge area under cultivation (9.4 MHa), high production (23 million tons), and enormous water 108 109 consumption (18 BCM), both crop productivity (2.5 t ha<sup>-1</sup>) and crop water productivity (CWP) (1.83 kg m<sup>-3</sup>) of Indian Maize are far lower than corresponding world averages (Sharma, 2018). 110 Low CWP (hence, WUE) of Indian Maize can be attributed to: i) a high dependence (85%) on 111 112 erratic, uncertain rainfall, ii) low adoption of hybrid varieties, iii) improper drainage facilities 113 leading to water logging, and iv) unscientific application of irrigation water without analysing soil-water-crop interactions (Shankar, 2012). Thus, an accurate quantification of WUE and its 114 temporal variation during the crop cycle is essential for effective irrigation water management 115 116 of Maize crop (Medrano, 2015).

While the effect of time-averaging on carbon and water fluxes measured at EC sites is 117 reported, the effect on their interaction term, i.e. WUE, which is crucial in irrigation water 118 management is unexplored. Evaluation of time-averaging period on WUE dynamics is 119 necessary to understand the contribution of low and high frequency photosynthetic carbon and 120 121 evaporative water fluxes generated from various field management strategies. Also, most of the EC flux studies are confined to data rich AmeriFLUX, EuroFLUX, and ChinaFLUX sites, 122 with limited focus to Indian fragmented croplands. This motivates the present study, and the 123 objectives of this study are as follows: i) investigate the role of time-averaging of EC fluxes on 124 EBR and WUE dynamics, ii) compute optimal averaging period to simulate carbon and water 125 (hence, WUE) fluxes of Maize crop, and iii) investigate the association of carbon, water, and 126 WUE fluxes between multiple averaging periods. Results of this study can help in designing 127

efficient management strategies using EC datasets in response to changes in WUE during thecrop cycle.

130

## 131 2.0 MATERIALS AND METHODOLOGY

## 132 2.1 Site Description and Instrumentation

Controlled Maize plots situated at Professor Jaya Shankar Telangana State Agricultural 133 University (PJTSAU), Hyderabad, Telangana, India (17°19'17" N, 78°24'35" E, 559 m above 134 sea level) forms the study area. The region is composed of red gravel to sandy loam soils with 135 136 field capacity and wilting point in the ranges of 17.92 - 19.56 % and 8.2 - 9.87% respectively. As per Koppen-Geiger's classification, the region falls under tropical savanna climate zone 137 (Aw) characterized by long dry and short wet seasons (Kottek, 2006). Mean annual 138 139 precipitation of the region is 900 mm (IMD, 2019) with more than 80% occurring during the monsoon months (Jun-Sep). Temperatures are high during summer  $(38.33 \pm 2.12 \text{ }^{4}\text{C})$  and 140 141 low during winter  $(30 \pm 2.20 \stackrel{\text{Q-C}}{\to})$  months. Humidity of the region varies from 35% in summer to 73% in monsoon (CGWB, 2013). Mean seasonal wind speed is in the range of 1.5 to 2.7 m/s 142 143 (Peddinti and Kambhammettu 2019). Hydro-geologically, the study area forms part of the Deccan plateau characterized by multiple layers of solidified flood basalt resulting from 144 volcanic eruptions. Depth to groundwater ranges from 12 m (pre-monsoon) to 6 m (post-145 monsoon) (CGWB, 2013). 146

Meteorological parameters and turbulent fluxes were obtained for one crop season, i.e.
26 May to 06 Sep, 2019 using an open path eddy covariance (EC) flux tower. <u>The flux system</u>
is composed of integrated CO<sub>2</sub>/H<sub>2</sub>O open-path gas analyzer and 3D sonic anemometer
(IRGASON-EB-NC, Campbell Sci. Inc., USA) to measure CO<sub>2</sub> and H<sub>2</sub>O concentrations at 3
m above the canopy. <u>The flux system is composed of a 3D sonic anemometer (CSAT3,</u>

- 152 Campbell Sci. Inc., USA), and an open path fast response infrared gas analyzer (IRGASON-
- 153 EB-IC, Campbell Sci. Inc., USA) to measure CO<sub>2</sub> and H<sub>2</sub>O fluxes at 3 m above the canopy.
- Raw data was collected with a logger (CR1000, Campbell Sci. Inc., USA) at 10 Hz frequency.
- 155 Additionally, slow response meteorological variables including precipitation (TE525-L-PTL,
- 156 Tipping Bucket, Campbell Sci. Inc., USA), soil heat flux (HFP01SC-L-PTL, Campbell Sci.
- 157 Inc., USA), solar radiation (CNR 4, Campbell Sci. Inc., USA), and soil moisture (CS616-L-
- 158 PT-L, Campbell Sci. Inc., USA) were obtained at 10 min intervals.

Formatted: Not Highlight

## 160 2.2 Data Collection and Processing

Table 1 shows the phenological stages of the Maize crop in the study area (Soujanya, 2021). Additionally, leaf-area index (LAI) and mean plant height were measured-monitored during the crop cycle (Table 1). The LAI was measured using the plant canopy analyser, whereas the plant height was measured using a ruler from the base of the plant to its crown. Maize crops of the experimental fields are sown on 25<sup>th</sup> May 2019 and harvested on 6<sup>th</sup> September 2019 with a base period of 104 days.

| 1 | 6 | 7 |
|---|---|---|
|   |   |   |

Table 1: Phenological growth stages and physical properties of the Maize crop

| S.<br>No | Growth stage         | Start date | End date   | Period<br>Length<br>(days) | Leaf Area Index<br>(m <sup>2</sup> m <sup>-2</sup> ) | Plant height<br>(cm) |
|----------|----------------------|------------|------------|----------------------------|--|----------------------|
| 1        | 6 <sup>th</sup> leaf | 26/05/2019 | 12/06/2019 | 18                         | 0.61   | 46.8                 |
| 2        | Silking              | 13/06/2019 | 19/07/2019 | 37                         | 1.56   | 75.2                 |
| 3        | Dough                | 20/07/2019 | 12/08/2019 | 24                         | 3.46   | 133                  |
| 4        | Maturity             | 13/08/2019 | 06/09/2019 | 25                         | 3.03   | 134                  |

168

169 Data from the EC system at 10 Hz frequency was converted to ASCII format using LoggerNet (4.3) software (Campbell Scientific Inc., Logan, Utah, USA), and further 170 aggregated to various averaging periods (1, 5, 10, 15, 30, 45, 60, and 120 minutes). Post data 171 processing was done using EddyPro post-processing software (version 7.0.8, LI-COR, USA). 172 Primary corrections performed on the raw dataset include tilt corrections, turbulent 173 fluctuations, density fluctuations, frequency corrections and quality checks. Tilt corrections 174 175 were made by the double axis rotation method for each averaging period. Either The block 176 average method or and linear detrending method were used considered to correct compute the 177 turbulent fluctuations. Block averaging method was used for detrending the fluxes at 1, 5, 10, 30, 45, and 60 min averaging periods. Longer averaging periods (e.g. 120 min) has resulted in 178 179 inconsistency in the obtained fluxes, which is a weakness of the block averaging (Renhua,

Formatted: Not Highlight

Formatted: Not Highlight

Formatted: Not Highlight

Formatted: Not Highlight

180 2005; Sun et al., 2006), Hence, linear trend removal method was used to compute the fluxes 181 for 120 min averaging period. Density fluctuation corrections were done using Webb-182 Pearman-Leuning (WPL) method. Quality checks were performed following a flagging policy 183 proposed by Mauder and Foken (2006) (0-1-2 system). Flag set to "0" corresponds to the best quality fluxes, "1" corresponds to fluxes acceptable for general analysis, and "2" corresponds 184 to poor quality fluxes that should be removed from the dataset. The resulting fluxes may exhibit 185 spikes, discontinuity, randomness etc. There is a need to perform secondary corrections on the 186 187 data that include flux spike removal (Vickers and Mahrt 1997), friction velocity corrections (to 188 filter night time observations), gap filling and uncertainty analysis (Finkelstein, 2001), 189 skewness & kurtosis removal, spectral corrections, and frequency corrections. To correct flux estimates for low and high frequency losses due to instrument setup, intrinsic sampling limits 190 191 of the devices, and various data processing decisions, spectral corrections are performed. Additionally, slow sensor meteorological data obtained at 1 min interval were processed for 192 193 different time-averaging periods using the EddyPro post-processing software (version 7.0.8, 194 LI-COR, USA).

195

## 196 2.3 Effect of time-averaging on EBR and EC fluxes

197 Violation of law of conservation of energy resulting from the EC observed energy terms is referred as energy balance closure (EBC). Lack of conservation among the measured energy 198 199 terms of the EC tower is referred as energy balance closure (EBC). The available energy ( $R_n$ -200 G) is generally higher than the turbulent fluxes (H+LE), resulting in a positive balance (Eshonkulov, 2019) where R<sub>n</sub>, G, H and LE correspond to net radiation, soil heat flux, sensible 201 heat and latent heat respectively. Apart from instrument and measurement issues, this lack of 202 203 energy closure is thought to be partly from averaging periods and coordinate systems 204 (Finnigan, 2003; Finnigan, 2004; Gerken, 2018). The energy closure fraction, commonly 205 termed as energy balance ratio (EBR) is used to evaluate the quality of EC data by examining energy fluxes at the surface (Chen and Li 2012), given by: 206

 $207 \qquad EBR = \frac{H + LE}{R_n - G}$ 

 $208 \qquad \underline{H} = \rho_a C_p \overline{w'T'}$ 

209 <u>**LE** =  $L_v \overline{w' \rho_v'}$ </u>

(1)

(2).

(3)

Formatted: Font: 12 pt, Italic Formatted: Font: 12 pt, Italic, Font color: Dark Blue Formatted: Font: 12 pt, Font color: Dark Blue Formatted: Normal, Indent: Left: 0 cm

|  | - 6 |
|--|-----|
| is the temperature fluctuation, $L_y$ is the latent heat of vaporization and $\rho_y'$ is the H <sub>2</sub> O gas | 3   |
| concentration fluctuation.   |     |
|  |     |
| EBR helps to determine the averaging period required to calculate H and LE fluxes over a                           | ι   |
| range of landscapes (Chen and Li 2012). A high EBR (EBR $\geq$ 0.7) ensures reliability of EC                      | 2   |
| observations for use with flux estimation (Barr et al., 2006; Kidston et al., 2010).                               | ,   |
|  |     |
| Eddy fluxes are computed as the covariance between instantaneous deviation in vertical                             |     |
| wind speed $(w')$ and scalar component of interest $(s')$ from their respective means, given by                    |     |
|  |     |

where  $\rho_a$  is the air density;  $C_p$  is the specific heat of air,  $w'_i$  is the wind velocity fluctuation,  $T'_i$ 

218  $F \approx \overline{\rho_a w' s'}$ 

219 (<u>24</u>)

210

211

212

213

214

215

216

217

where  $\overline{\rho_a}$  is the mean air density, and the overbar represents the time-average of eddy fluxes, which is of interest in the present study. Depending on the scalar component considered (ex: temperature, water vapour (H<sub>2</sub>O), carbon dioxide (CO<sub>2</sub>) concentration), corresponding eddy fluxes (ex: sensible heat, latent heat, carbon flux) are computed as below.

$$224 \qquad F_{CO_2} \approx \overline{\rho_a w' CO_2'} \tag{35}$$

225 
$$F_{H_2O} \approx \overline{\rho_a w' H_2 O'}$$

Ecosystem WUE is then estimated as the ratio of daytime carbon (net primary product) to water
 fluxes (evapotranspiration), observed <u>during considering</u> daytime unstable atmospheric
 conditions (08:00 am to 04:00 pm) given by:

229 
$$WUE = \frac{NPP}{ET} = \frac{F_{CO_2}}{F_{H_2O}}$$
 (5)

Fluxes originating from real-world sites are composed of both high frequency (turbulence) and low frequency (advection) fluctuations, with a spectral gap in between. Isolating local turbulence component for use with flux studies is achieved by choosing an appropriate averaging period,  $T_1$  (typically 30 minutes) on fast response measurements operating at high frequency  $T_2$  (Manon and Kristian 2020). Optimal averaging period ( $T_1$ ) should be long enough to reduce random error (Berger, 2001) and short enough to avoid non-stationarity associated with advection (Foken & Wichura, 1996). The flux estimates (eq. 2) are further decomposed

Formatted: Font: (Default) Times New Roman, 12 pt Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt. Font color: Dark Blue Formatted: Font: 12 pt, Font color: Dark Blue Formatted: Font: 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt. Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Space Before: 0 pt, After: 8 pt Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue Formatted: Font: 12 pt, Font color: Dark Blue

Formatted: Font: 12 pt, Font color: Dark Blue

Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue

Formatted: Subscript

Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Dark Blue

Formatted: Font: (Default) Times New Roman, 12 pt, Font color: Black

8

(46)

into frequency dependent contributions, known as co-spectra Cows (f) between vertical wind 237 238 velocity (w) and scalar of interest (s) for frequencies 'f' (Manon and Kristian 2020). For an accurate estimation of the flux, it is essential that the EC method is applied under conditions 239 where the flow is stationary, and all eddies carrying flux are sampled. Given that the flow 240 remains stationary, an 'ogive' serves as a check for the essential requirement to sample all 241 scales carrying the flux. Ogive function is well proposed to check if all low frequency fluxes 242 are included in the turbulent flux measured with the EC method (Foken & Wichura, 1996; 243 Foken et al., 2005). It is used to investigate the energy balance losses caused by low frequency 244 245 fluxes. Ogive analysis is performed to investigate the flux contribution from each frequency 246 range and to arrive at most suitable averaging period to capture most of the turbulent fluxes (Desjardans, 1989; Charuchittipan, 2014). Ogive function thus provides the cumulative sum of 247 co-spectral energy starting from the highest frequency, given by: 248

$$Q49 \qquad Og_{ws}(f_0) = \int_{f_0}^{\infty} Co_{ws}(f) df \qquad (67)$$

The point of convergence on the Ogive plot to an asymptote corresponds to optimal averaging 250 251 period  $(T_1)$  for use with averaging of high frequency turbulence fluxes. In other words, the 252 point at which the ogive plot flattens out represents the optimal averaging period. A total of 253 eight averaging periods, i.e., 1, 5, 10, 15, 30, 45, 60, and 120 minutes were considered to 254 investigate the role of time-averaging on EBR, EC and WUE-and EC-fluxes, and further to 255 arrive at the optimum averaging period for use with WUE estimation. The biophysical and physiological characteristics such as plant height, crop water requirement, LAI, etc. changes 256 with respect to the crop growth stage (Chintala et al., 2024) and have a significant effect on the 257 258 EC fluxes. Since these factors vary over growth stages, For this reason, time-averaging of EC fluxes is separated based on crop growth stage. 259

### 260

### 261 2.4 Performance Evaluation

The ability of various averaging periods to close the energy balance and compute the EC fluxes is evaluated using three goodness of fit indicators, namely: a) coefficient of determination  $(R^2)$ , b) root mean squared error (RMSE), and c) relative error (RE). While  $R^2$ and RMSE aim to quantify the error in closing the energy balance, RE is aimed to compute the error in estimating EC fluxes with conventional 30 min averaging period relative to optimal averaging period.

|  |        | Formatted              |
|--|--------|------------------------|
| Root mean square error (RMSE) measures overall accuracy in closing the energy balance for  |        | Formatted              |
| a given averaging period by penalizing large errors heavily, given by:   |        | Formatted              |
| $RMSE = \left[ \frac{\sum_{i=1}^{n} (R_{n} - G)_{i} - (H + LE)_{i}}{n} \right]^{0.5}$  |        | Formatted              |
|  |        | Formatted              |
| $RMSE = \left[\frac{\sum_{i=1}^{n} ((R_n - G)_i - (H + LE)_i)^2}{n}\right]^{0.5} $ (78)  |        | Formatted              |
|  | $\leq$ | Formatted              |
| Where where n is the number of observations.   |        | Formatted              |
| $Coefficient of determination (R^2) \underline{and Pearson \ correlation \ coefficient \ (r)} \underline{is} \underline{are \ thea} \ measures$  |        | Formatted              |
| of the strength of linear association between turbulent fluxes and available energy, given by:   |        | Formatted              |
|  |        | Formatted              |
| $= \left(\sum_{i=1}^{n} \left[ (R_{r} - G)_{i} - \overline{(R_{r} - G)} \right]^{2} \left[ (H + LE)_{i} - \overline{(H + LE)} \right]^{2} \right)^{2}$   | 1      | Formatted              |
| $R^{2} = \begin{cases} \frac{\sum_{i=1}^{n} [(R_{n}-G)_{i} - \overline{(R_{n}-G)}]^{2} [(H+LE)_{i} - \overline{(H+LE)}]^{2}}{\sqrt{\sum [(R_{n}-G)_{i} - \overline{(R_{n}-G)}]^{2} [(H+LE)_{i} - \overline{(H+LE)}]^{2}}} \end{cases} $ (89) |        | Formatted              |
| $\left( \sqrt{\sum \left[ (n_n - 0) \right] - (n_n - 0)} \left[ (n + LE) \right] - (n + LE) \right] \right)$   | /      | Formatted              |
|  |        | Formatted              |
| $r = \left\{ \frac{\sum_{i=1}^{n} [(R_n - G)_i - \overline{(R_n - G)}] [(H + LE)_i - \overline{(H + LE)}]}{\left( \sqrt{\sum_{i=1}^{n} [(R_n - G)_i - \overline{(R_n - G)}]^2 [(H + LE)_i - \overline{(H + LE)}]^2}} \right\}$               | $\ $   | Formatted              |
| $\left(\sqrt{\sum \left[(R_n - G)_i - (R_n - G)\right]^2 \left[(H + LE)_i - (H + LE)\right]^2}\right)$   |        | Formatted              |
| Relative error (RE) provides the disparity in the fluxes estimated with conventional (30 min)  |        | Formatted              |
|  |        | Formatted              |
| relative to the fluxes estimated with optimal averaging period, given by:  |        | Formatted              |
| $RE = \left[\frac{\{F_{opt} - F_{30min}\}}{F_{opt}}\right] \times 100 \tag{910}$   |        | Formatted              |
| $KL = \begin{bmatrix} F_{opt} \end{bmatrix} \times 100 \tag{710}$  |        | Formatted              |
| where $F_{opt}$ and $F_{30}$ are the flux of interest considering optimal and conventional (30 min)  |        | Formatted              |
|  |        | Formatted              |
| averaging periods.   |        | Formatted              |
| <u>Averaging period corresponding to A-high R<sup>2</sup> (close to 1), low RMSE (close to zero), and low</u>  |        | Formatted              |
| RE (close to zero)-is considered to be the optimal choice in representing the EC fluxes.   |        | Formatted              |
|  |        | Formatted              |
|  |        | Formatted              |
| 3.0 RESULTS AND DISCUSSION   |        | Formatted              |
| 50 RESULTS AND DISCUSSION  |        | Formatted              |
| 3.1 Diurnal variations in energy balance components  |        | Formatted              |
|  |        | Formatted              |
| To understand the energy variation in response to rapid changes in meteorological  |        | Formatted              |
| conditions, we analysed the diurnal variations in energy balance components. Figure 1 shows  |        | Formatted              |
|  |        | Formatted<br>Formatted |
|  |        | Formatted              |
|  |        | A                      |

275 
$$R^{2} = \begin{cases} \frac{\sum_{i=1}^{n} [(R_{n}-G)_{i} - \overline{(R_{n}-G)}]^{2} [(H+LE)_{i} - \overline{(H+LE)}]^{2}}{\sqrt{\sum [(R_{n}-G)_{i} - \overline{(R_{n}-G)}]^{2} [(H+LE)_{i} - \overline{(H+LE)}]^{2}}} \end{cases}^{2} \\ 276 \qquad r = \begin{cases} \frac{\sum_{i=1}^{n} [(R_{n}-G)_{i} - \overline{(R_{n}-G)}]^{2} [(H+LE)_{i} - \overline{(H+LE)}]^{2}}{\sqrt{\sum [(R_{n}-G)_{i} - \overline{(R_{n}-G)}]^{2} [(H+LE)_{i} - \overline{(H+LE)}]^{2}}} \end{cases}$$

277 Relative erro 278 relative to th

279 
$$RE = \left[\frac{\{F_{opt} - F_{30min}\}}{F_{opt}}\right] \times 100$$

280 where Fopt a 281 averaging pe

282 Averaging p

283 RE (close to

284

268

269

270

271

272

273

274

#### 3.0 RESULT 285

#### 3.1 Diurnal 286

To und 287 conditions, w 288

| Formatted | ( |
|-----------|---|
| Formatted |   |
| Formatted |   |
| Formatted |   |
| Formatted | ( |
| Formatted |   |
| Formatted | ( |
| Formatted |   |
| Formatted |   |
| Formatted | ( |
| Formatted |   |
| Formatted | ( |
| Formatted | ( |
| Formatted |   |
| Formatted | ( |
| Formatted |   |
| Formatted |   |
| Formatted |   |
| Formatted | ( |
| Formatted | ( |
| Formatted | ( |
| Formatted |   |
| Formatted | ( |
| Formatted |   |
| Formatted |   |
| Formatted | ( |
| Formatted | ( |
| Formatted |   |
| Formatted | ( |
| Formatted |   |
| Formatted | ( |
| Formatted |   |
| Formatted | ( |
| Formatted | ( |
| Formatted | ( |
| Formatted |   |
| Formatted | ( |
| Formatted | ( |
| Formatted | [ |
|           |   |

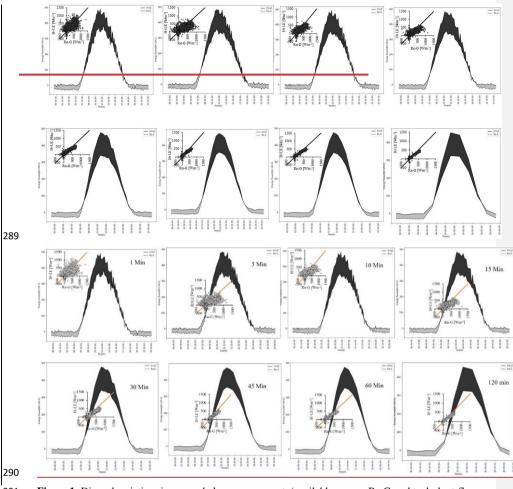


Figure 1: Diurnal variations in energy balance components (available energy: Rn-G and turbulent fluxes:
 H+LE) during the crop cycle with different averaging periods. Inset: Scatter-plots between the two datasets.

294 the diurnal variations in available energy (Rn-G) and turbulent fluxes (H+LE) averaged over the crop cycle for various time-averages. The diurnal variations of (Rn-G) and (H+LE) are bell-295 shaped, with peak occurring at around noon (480.16  $\pm$  14.15 Wm<sup>-2</sup>, 356.23  $\pm$  18.51 Wm<sup>-2</sup>) 296 (Figure 1). The energy balance difference (shaded areas of the figure) is found to be positive 297  $(76.88 \pm 43.14 \text{ Wm}^{-2})$  during daylight hours (08:00 am to 06:00 pm) and is negative (-24 ± 298 11.65 Wm<sup>-2</sup>) for the remaining time. The vertical offset between the two curves, representing 299 the residual of energy balance is highest around the noon (142.39  $\pm$  19.42 Wm<sup>-2</sup>), and is 300 301 consistent between the averaging periods. For an average site-day, the cumulative energy

| 302 | balance difference was found to be constant with a mean of $1811 \pm 91.56$ Wm <sup>-2</sup> at all averaging                    |
|-----|--|
| 303 | periods. The cumulative energy balance difference is crossing the 'zero' line at around 11:30                                    |
| 304 | am. The variation is rough at lower averaging periods due to a high sample size (n= 10859 at                                     |
| 305 | T = 1 min) and is gradually smoothened towards higher averaging periods (n= 811 at $T = 120$                                     |
| 306 | min). The slope of regression lines between (H+LE) and ( $R_n$ -G) considering all averaging                                     |
| 307 | periods are in the range of 0.59 to 0.71 with a mean of 0.65 $\pm$ 0.041. The intercept is ranged                                |
| 308 | from 19.01 to 31.56 Wm <sup>-2</sup> . The best slope ( $\geq 0.70$ ) and intercept ( $\leq 20$ Wm <sup>-2</sup> ) were achieved |
| 309 | with 45 and 60 minutes averaging periods, which is consistent with literature (Gao, 2017;  |
| 310 | Leuning, 2012). This conclude that, longer averaging periods have a good closure over shorter                                    |
| 311 | <u>averaging periods</u> . The strength of linear association between $(R_n-G)$ and $(H+LE)$ around the                          |
| 312 | best fit line, explained by r is high $(0.80 \le r \le 0.9)$ at low averaging periods, i.e., 1, 5, 10                            |
| 313 | minutes, and is very high ( $r > 0.9$ ) for other averaging periods (Table 2). However, the departure                            |
| 314 | of the data from 1:1 line is relatively low both at low short and high long averaging periods.                                   |
| 315 | Our findings show that averaging period has minimal influence in representing the energy   |
| 316 | balance terms. However, data scatter around 1:1 line is high for shorter time-averages due to                                    |
| 317 | large sample size and data randomness. Our findings show that averaging period has minimal                                       |
| 318 | influence in representing the energy balance terms.  |
|     |  |

| 319 | <b>Table 2:</b> Summary of linear regression parameters in closing the energy balance with different |
|-----|--|
| 320 | averaging periods.   |

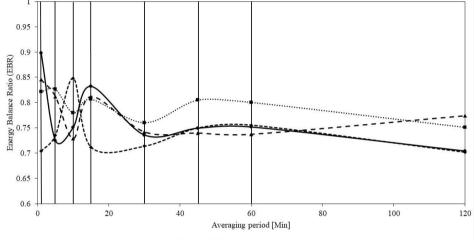
| Averaging Period | Slope | R <sup>2</sup> | Intercept<br>(Wm <sup>-2</sup> ) | r    | N     | RMSE<br>(Wm <sup>-2</sup> ) |
|------------------|-------|----------------|----------------------------------|------|-------|-----------------------------|
| 1min             | 0.63  | 0.66           | 30.31                            | 0.81 | 10859 | 98.38                       |
| 5min             | 0.59  | 0.74           | 31.56                            | 0.86 | 10785 | 76.47                       |
| 10min            | 0.60  | 0.80           | 28.94                            | 0.90 | 10753 | 64.41                       |
| 15min            | 0.63  | 0.84           | 26.56                            | 0.92 | 7150  | 58.18                       |
| 30min            | 0.66  | 0.93           | 20.49                            | 0.96 | 3554  | 38.33                       |
| 45min            | 0.70  | 0.94           | 19.99                            | 0.97 | 2355  | 36.30                       |

Formatted: Not Highlight

| 60min  | 0.71 | 0.94 | 19.01 | 0.97 | 1765 | 35.07 |
|--------|------|------|-------|------|------|-------|
| 120min | 0.67 | 0.93 | 20.77 | 0.96 | 811  | 39.95 |

# 322 **3.2 Effect of averaging period on EBR and EC fluxes**

The variation in energy balance ratio (EBR) with averaging period for individual growth stages of the crop is presented in Figure 2. We observed a clear departure of EBR from



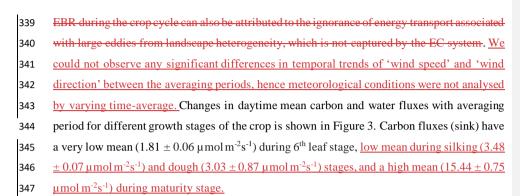


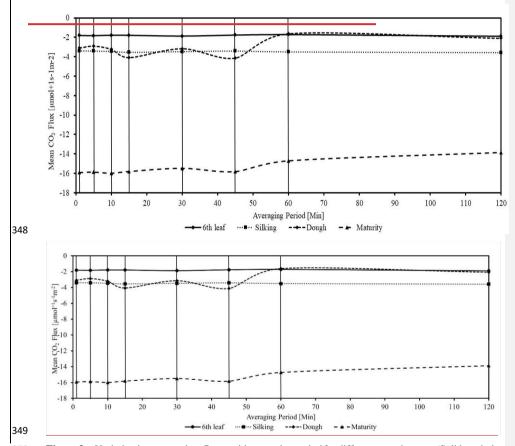
← 6th Leaf … Silking - - Dough - - Maturity

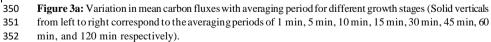
Figure 2: Variation in energy balance ratio (EBR) with averaging period for different growth stages. (Solid
 verticals from left to right correspond to the averaging periods of 1 min, 5 min, 10 min, 15 min, 30 min, 45
 min, 60 min, and 120 min respectively).

unity for all growth stages, particularly with dough and maturity stages due to ignorance of 329 canopy heat storage. EBR is fluctuating between 0.70 and 0.90 at low (1 - 30 min) averaging 330 periods and is fairly constant  $(0.75 \pm 0.03)$  at high ( $\geq 30$  min) averaging periods. Our reported 331 values of EBR during the crop growth are within the typically found range of 0.65 to 1.2 for 332 most of the crops (Feng, 2017; Finnigan, 2003; Wilson, 2002). The mean EBR with 333 334 conventional 30 min averaging period is found to be 0.74, 0.76, 0.71, and 0.74 during 6th leaf, 335 silking, dough, and maturity stages respectively. Low EBR during the crop cycle can also be 336 attributed to the ignorance of energy transport associated with large eddies from landscape 337 heterogeneity. However, EC method assumes the landscape within the footprint of 338 measurement to be flat and homogenous. This violation might have lowered the EBR. Low

Formatted: Not Highlight







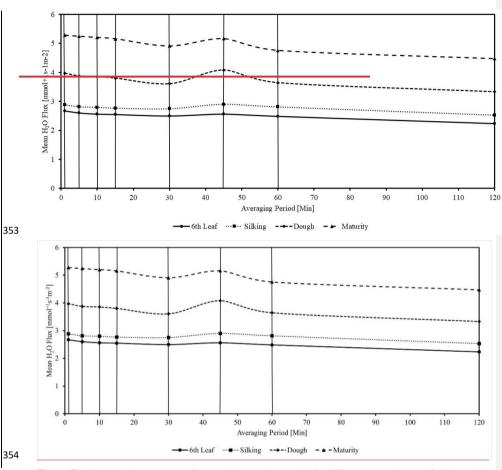


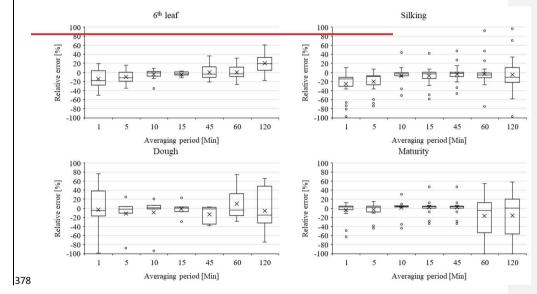
Figure 3b: Variation in mean water fluxes with averaging period for different growth stages (Solid verticals
from left to right correspond to the averaging periods of 1 min, 5 min, 10 min, 15 min, 30 min, 45 min, 60
min, and 120 min respectively).

low mean during silking  $(3.48 \pm 0.07 \,\mu \text{molm}^2\text{s}^4)$  and dough  $(3.03 \pm 0.87 \,\mu \text{molm}^2\text{s}^4)$  stages, 358 359 and a high mean  $(15.44 \pm 0.75 \,\mu \text{mol}^{-2}\text{s}^{-1})$  during maturity stage. Mean carbon fluxes during 6<sup>th</sup> leaf and silking stage are mostly unaffected by averaging period. We observed a gradual 360 361 increase in water vapour fluxes during the crop cycle, from 6<sup>th</sup> leaf ( $2.52 \pm 0.13$  mmol s<sup>-1</sup>m<sup>-2</sup>) to maturity (5.02  $\pm$  0.29 mmol s<sup>-1</sup>m<sup>-2</sup>). As the averaging period is increased, the mean water 362 363 vapour flux is decreased, with an exception at 45 min averaging period. Distribution of 364 errorDeviation in in representing carbon and water fluxes at different averaging periods, 365 relative to the conventional 30 min averaging period i.e. relative error (RE) is presented in 366 Figure 4. The RE is obtained by considering daily averages in the deviations for each growth

Formatted: Not Highlight

Formatted: Not Highlight

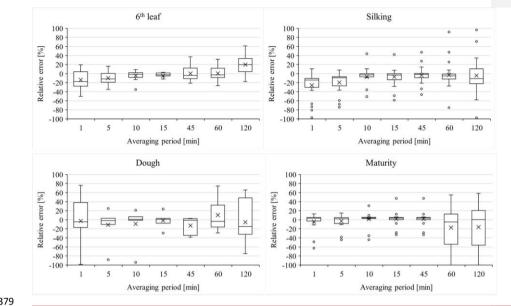
367 stage. During 6th leaf and silking stages, RE in estimating carbon fluxes is high (~ -15 %) with 368 low averaging periods, and is gradually diminishing towards higher averaging periods, with an 369 exception at very high (120 min) average period. For dough and maturity stages, RE is found 370 to be significant with higher averaging periods (60-120 min). RE in estimating water vapour 371 fluxes is found to be insignificant at all averaging periods for the 6<sup>th</sup> leaf and silking stages. 372 However, dough and maturity stages have shown a large variation in RE considering either, 373 too-short (1, 5 min) or too-long (60, 120 min) time averages. A high variability in RE for time 374 scales larger than 45 min indicate the effects of sub mesoscale (non-turbulent) motions. Hence, 375 45 min average period can be considered as optimal in isolating the turbulence components for 376 use with flux representation. RE in estimating water vapour fluxes is found to be insignificant 377 at all averaging periods, irrespective of growth stage.



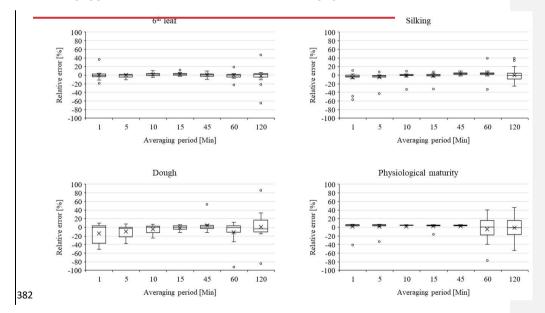
### **Formatted:** Not Highlight

Formatted: Not Highlight
Formatted: Not Highlight

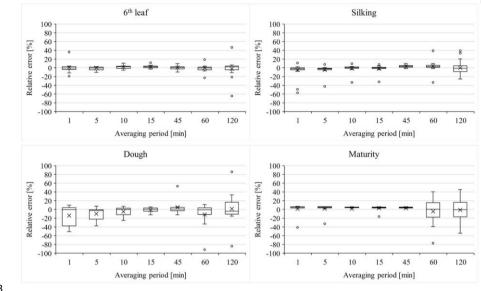




**Figure 4a:** Whisker plots showing the distribution of error in estimating carbon fluxes with various averaging periods relative to the conventional 30 min averaging.





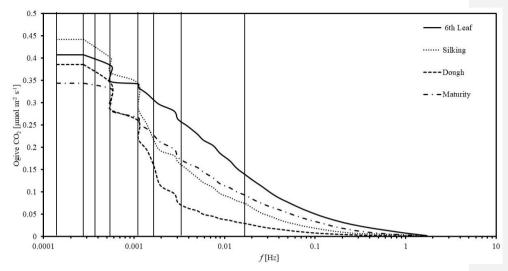


383

Figure 4b: Whisker plots showing the distribution of error in estimating water fluxes with various averaging
 periods relative to the conventional 30 min averaging.

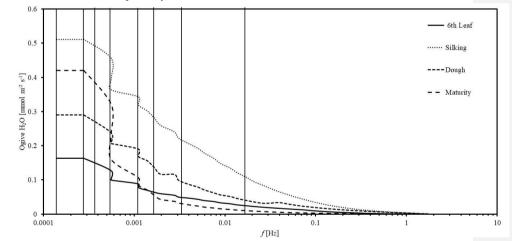
## 387 3.3 Selection of Optimal averaging period

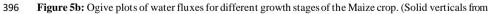
388 Ogive functions representing the cumulative integral of the co-spectral energy starting 389 with highest frequency, i.e., 0.016 Hz (T = 1 min) for carbon, water and water-WUE fluxes are 390 presented



395

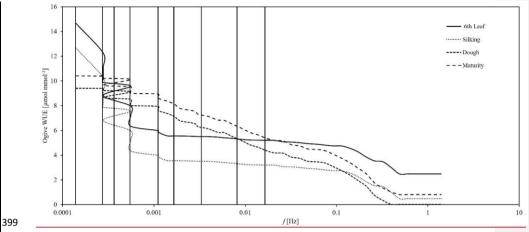
Figure 5a: Ogive plots of carbon fluxes for different growth stages of the Maize crop. (Solid verticals from
 left to right extremes correspond to the averaging periods of 120 min, 60 min, 45 min, 30 min, 15 min, 10
 min, 5min and 1 min respectively).





left to right extremes correspond to the averaging periods of 120 min, 60 min, 45 min, 30 min, 15 min, 10

398 min, 5min and 1 min respectively)





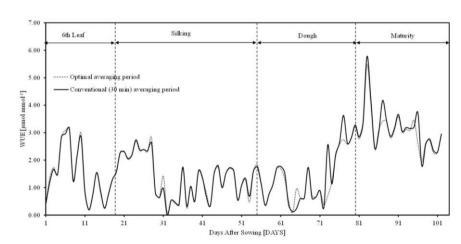
400 Figure 5c: Ogive plots of water use efficiency for different growth stages of the Maize crop. (Solid verticals 401 from left to right extremes correspond to the averaging periods of 120 min, 60 min, 45 min, 30 min, 15 min, 402 10 min, 5min and 1 min respectively)

in Figure 5. Shorter time periods corresponding to daytime unstable atmospheric conditions 403 404 (08:00 am to 04:00 pm) for various growth stages were investigated. Ogive plots of carbon fluxes for 6th leaf and silking stages showed an increasing trend upto 0.011 Hz (15 min) and 405 remained fairly constant before 0.0055 Hz (30 min). This concludes that whole turbulent 406 spectrum can be covered with 15 to 30 min averaging, with negligible flux contribution from 407 408 longer frequencies. Ogive plots of carbon fluxes for dough and maturity stages showed a continuous increasing trend without a defined plateau (horizontal asymptote) in between. This 409 conclude that the conventional 30 min averaging period is inadequate to capture the low 410 411 frequency fluxes, thus demanding for higher averaging periods. We observed a similar behaviour with water fluxes (Figure 5b). The flat part of the Ogive curve representing the 412 optimal averaging period was found to vary across the crop cycle. While 15-30 min time-413 average is suitable for aggregating the EC fluxes during 6<sup>th</sup> leaf and silking stages, 45-60 min 414 415 averaging is more appropriate for dough and maturity stages. Similar to carbon and water 416 fluxes, the ogive plots for WUE were presented in Figure 5c. From this, it is observed that the 417 flat part of ogive is achieved at 15 min time average period for the stages of 6<sup>th</sup> leaf and silking 418 and 45 min time average for the dough and maturity stages which is similar to the carbon and 419 water fluxes. It concludes that the WUE followed a similar behaviour as its individual fluxes i.e. carbon and water fluxes in achieving optimal time averages. The crop biophysical factors 420 like LAI and plant height are minimum during 6th leaf and silking stages contributes low 421 quantity of CO2 and H2O fluxes (refer figure 3a & 3b) whereas they are maximum in the later 422

423 stages of the crop i.e., tasselling dough and maturity by contributing to high quantitiesy of CO2 424 and H<sub>2</sub>O fluxes (refer figure 3a & 3b). Our results are in accordance with the previous studies of Fong et al., 2020 on Cotton, where the responses in NPP and ET were related seasonally to 425 426 plant growth stages. The previous studies on various crops revealed that the GPP NPP and ET fluxes were initially low in the early stages and increases towards maturity stage due to crop 427 phenology and management practices. To capture these low quantity fluxes, low averaging 428 periods i.e., 15 min is sufficient, whereas 45 min time-averaging period can capture high 429 430 quantity fluxes that are prevalent during later growth stages of the crop. As the crop 431 characteristics are dependent on the crop growth stages, a single time-averaging period is not 432 appropriate to capture the dynamics of CO<sub>2</sub> and H<sub>2</sub>O fluxes and as well as their ratio WUE. This clearly demonstrates that, as the plant achieves its higher stage, flux contribution from 433 low-frequency components becomes more valuable. Very low averaging periods (ex: 1 min, 5 434 min) were found unsuitable to capture low-frequency flux components, which is in agreement 435 436 with literature (Feng, 2017).

437

### 438 **3.4 Dynamics of Water use efficiency**

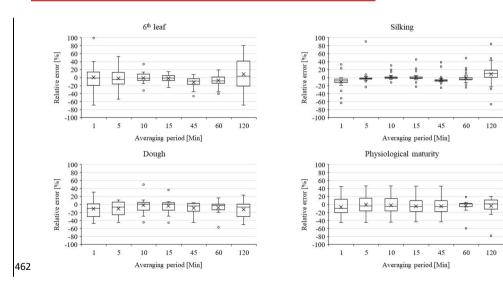


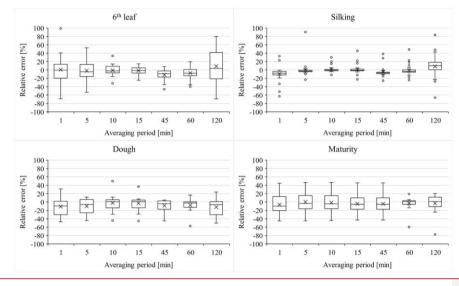
Daily means of water use efficiency (WUE) estimated with conventional 30 min and
 growth specific optimal averaging periods is presented in Figure 6. Mean WUE fluxes for 6<sup>th</sup>

441 442

Figure 6: Seasonal variations in daily mean WUE fluxes obtained with conventional 30 min (blacksolid)
 and optimal averaging periods (reddotted) during the crop cycle.

leaf, silking, dough and maturity stages with conventional 30 min averaging are  $1.48 \pm 0.96$ , 444 445  $1.36 \pm 0.73$ ,  $1.38 \pm 0.95$  and  $3.184 \pm 0.78 \ \mu mol \ mmol^{-1}$  respectively. Corresponding fluxes with stage specific optimal averaging periods are  $1.49 \pm 0.95$ ,  $1.37 \pm 0.74$ ,  $1.39 \pm 0.79$  and 3.06446  $\pm$  0.69 µmol mmol<sup>-1</sup> respectively. Error in estimating mean daily WUE fluxes with 30 min 447 averaging is very low (< 1.45%) during 6<sup>th</sup> leaf and silking stages, low (8.56 to 9.04 %) during 448 maturity stage, and is moderate (11.84 to 12.12 %) during dough stage. This conclude that, 449 choice of optimal averaging period is more crucial for late stage growth periods of the crop. 450 Distribution of error in estimating WUE fluxes with various averaging periods relative to 451 452 conventional 30 min average period (RE) is presented in Figure 7. A close to zero RE with all 453 averaging periods during 6<sup>th</sup> leaf and silking stages conclude that, choice of averaging period has insignificant role in estimating the WUE fluxes, particularly during early growth stages. A 454 slightly high RE (~-5.4%) during dough and maturity stages conclude that, choice of averaging 455 period matters for WUE estimation during late stage periods. Hence, conventional 30 min 456 457 averaging period can be considered for estimating WUE fluxes during 6th leaf and silking stages, whereas optimal averaging period need to be considered for estimating WUE fluxes 458 459 during dough and maturity stages. Correlation charts showing the linear association within 460 carbon, water, and WUE fluxes represented at different averaging periods is presented in Figure 8. For ease with comparison, data for the entire crop cycle was considered. Linear association 461

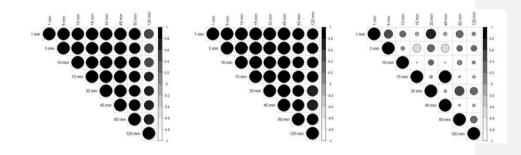




463

Figure 7: Whisker plots showing the distribution of error in estimating WUE fluxes with various averaging
 periods relative to the conventional 30 min averaging.

between any two averaging periods is positive ( $\rho > 0.56$ ) for carbon and water fluxes. Except 466 with 120 min time-averaging, all other averaging periods are strongly correlated ( $\rho > 0.87$ ) 467 with 30 min averaging period. However, a poor linear association in WUE fluxes was observed 468 469 between any two averaging periods, which is attributed to a larger variation in individual WUE 470 fluxes between averaging periods. However, the corresponding individual carbon and water 471 fluxes have recorded low variations between time averages. This conclude that, the need for 472 optimal averaging period is more crucial in estimating WUE fluxes rather than individual carbon and water fluxes. Our findings can improve representation of WUE fluxes using EC 473 474 data, thereby help in developing efficient water management strategies in response to WUE 475 changes.



476

477 Figure 8: Correlation charts showing the linear association of a) Carbon fluxes, b) Water fluxes, and c) 478 WUE fluxes estimated with different averaging periods.

479

508

#### 4.0 CONCLUSIONS 480

This study explores the effect of averaging period of EC fluxes on EBR dynamics and 481 482 WUE in semi-arid Indian conditions. The proposed methodology was applied on drip-irrigated Mmaize field for one crop period (May-Sept 2019). Major findings of this study are: 483

- EBR was found vary marginally at low averaging periods and less significant during • 484 485 higher averaging periods.
- With reference to conventional 30 min averaging period, relative error is within 12% 486 • 487 for 10-45 min averaging periods for carbon fluxes and is within 5% for 15-45 averaging 488 periods for water fluxes.
- From ogive analysis we found the optimal averaging period as 15 30 min for the 6th 489 • leaf, and silking stages, and as 45 - 60 min for the dough and maturity stages. 490
- The mean carbon fluxes are increasing from  $1.81 \pm 0.06 \,\mu mol^{+1}m^{-2}s^{-1}$  (6th leaf stage) 491 492 to  $15.44 \pm 0.75 \ \mu mol^{+1}m^{-2}s^{-1}$  (maturity stage) which indicates that carbon sink is a function of crop growth period. In case of water fluxes, it increased from  $2.52 \pm 0.13$ 493  $\text{mmol}^{+1}\text{m}^{-2}\text{s}^{-1}$  (6th leaf stage) to 5.02  $\pm$  0.29  $\text{mmol}^{+1}\text{m}^{-2}\text{s}^{-1}$  (maturity stage). Variation of 494 carbon and water fluxes are directly influencing WUE dynamics. 495
- 496 • The variation in WUE was increased subsequently with the plant growth and achieved its maximum value of 5.17 µmol mmol<sup>-1</sup> in between dough to maturity stages which 497 concludes that, crop consumes more carbon than water as the crop period progresses. 498
- The correlation between CO2 and H2O fluxes for all averaging periods was found to be 499 • high. However, WUE, which is calculated as the ratio of CO<sub>2</sub> and H<sub>2</sub>O fluxes, is not 500 following the same pattern. While 45 min and 15 min averaged WUE exhibits a good 501 correlation, 30 min averaged WUE is not correlated with other averaging periods. 502 503 Averaging period is found to be an influencing factor in controlling WUE, hence should 504 be considered with caution during the crop growth.

505 This study is limited to understand the role of different time-averaging periods on EC observed. carbon, water fluxes as well as EC derived WUE fluxes contributed by homogeneous Maize 506 507 crop which is having relatively smaller flux footprint in an unstable atmospheric condition. This study is limited to understand the role of different time-averaging periods on EC observed

Not Highlight

Formatted: Space After: 0 pt

Formatted: Font: (Default) Times New Roman, 12 pt,

| 509 | carbon, water fluxes as well as EC derived WUE fluxes. Study findings can help to accurately   |      |
|-----|--|------|
| 510 | characterise WUE of Maize crop considering growth stages for effective implementation of       |      |
| 511 | irrigation strategies,   | Form |
| 512 |  |      |
| 513 | Acknowledgments  |      |
| 514 | The authors acknowledge the anonymous reviewers for their insightful comments. This            |      |
| 515 | research evolved as an extension of a term project in CE6520-Irrigation Water Management       |      |
| 516 | course at IIT Hyderabad.   |      |
| 517 |  |      |
| 518 | Data Availability Statement:   |      |
| 519 | All footprint climatologies, site-level data files, and supplementary material can be accessed |      |
| 520 | via the Zenodo Data Repository ( <u>https://zenodo.org/badge/latestdoi/528291820</u> )         |      |
| 521 | (Shweta07081992, 2022)   |      |
| 522 |  |      |
| 523 | Author Contribution:   |      |
| 524 | Arun Rao Karimindla: Data processing and Analysis, Writing- Original draft. Shweta             |      |
| 525 | Kumari: Conceptualization, Methodology, Project Supervision. Saipriya SR: Data processing      |      |
| 526 | Analysis, and Writing- Original draft. Syam Chintala: Data processing and Analysis, Writing-   |      |
| 527 | Original draft. BVN Phanindra Kambhammettu: Project Administration, Writing-                   |      |
| 528 | Reviewing and Editing.   |      |
| 529 |  |      |
| 530 | Competing interests:   |      |
| 531 | The authors declare that they have no known competing interests or personal relationships that |      |
| 532 | could have appeared to influence the work reported in this paper.                              |      |
| 533 |  |      |
| 534 | 5.0 REFERENCES   |      |
|     |  |      |

Barr, A. G., Morgenstern, K., Black, T. A., McCaughey, J. H., & Nesic, Z. (2006). Surface
energy balance closure by the eddy-covariance method above three boreal forest stands
and implications for the measurement of the CO2 flux. Agricultural and Forest
Meteorology, 140(1–4), 322–337. https://doi.org/10.1016/j.agrformet.2006.08.007

25

ormatted: Font color: Red

| 539<br>540<br>541        | <ul> <li>Bramley, H., Turner, N. C., &amp; Siddique, K. H. M. (2013). Water Use Efficiency. In C. Kole (Ed.), Genomics and Breeding for Climate-Resilient Crops: Vol. 2 Target Traits (pp. 225–268). Springer Berlin Heidelberg. https://doi.org/10.1007/978-3-642-37048-9_6</li> </ul>                     |
|--------------------------|---|
| 542                      | Berger, B. W., Davis, K. J., Yi, C., Bakwin, P. S., & Zhao, C. L. (2001). Long-term carbon  |
| 543                      | dioxide fluxes from a very tall tower in a northern forest: Flux measurement  |
| 544                      | methodology. Journal of Atmospheric and Oceanic Technology, 18(4), 529–542.   |
| 545                      | Central Ground Water Board. (2015). Annual Report 2013.   |
| 546                      | https://cgwb.gov.in/old_website/Annual-Reports/Annual Report-2013-14.pdf  |
| 547<br>548<br>549<br>550 | <ul> <li>Charuchittipan, D., W. Babel, M. Mauder, J. P. Leps, and T. Foken, 2014: Extension of the Averaging Time in Eddy-Covariance Measurements and Its Effect on the Energy Balance Closure. <i>Boundary-Layer Meteorol.</i>, <b>152</b>, 303–327, https://doi.org/10.1007/s10546-014-9922-6.</li> </ul> |
| 551                      | Chen, Y. Y., and M. H. Li, 2012: Determining adequate averaging periods and reference   |
| 552                      | coordinates for eddy covariance measurements of surface heat and water vapor fluxes   |
| 553                      | over mountainous terrain. <i>Terr. Atmos. Ocean. Sci.</i> , 23, 685–701,  |
| 554                      | https://doi.org/10.3319/TAO.2012.05.02.01(Hy).  |
| 555                      | Chintala, S., Karimindla, A. R., & Kambhammettu, B. V. N. P. (2024). Scaling relations  |
| 556                      | between leaf and plant water use efficiencies in rainfed Cotton. Agricultural Water   |
| 557                      | Management, 292, 108680. https://doi.org/https://doi.org/10.1016/j.agwat.2024.108680  |
| 558                      | Desjardins, R. L., MacPherson, J. I., Schuepp, P. H., & Karanja, F. (1989). An evaluation of  |
| 559                      | aircraft flux measurements of CO2, water vapor and sensible heat. Boundary-Layer  |
| 560                      | Meteorology, 47(1), 55–69. https://doi.org/10.1007/BF00122322   |
| 561<br>562<br>563        | Eshonkulov, R., and Coauthors, 2019: Evaluating multi-year, multi-site data on the energy balance closure of eddy-covariance flux measurements at cropland sites in southwestern Germany. <i>Biogeosciences</i> , <b>16</b> , 521–540, https://doi.org/10.5194/bg-16-521-2019.                              |
| 564                      | Feng, J., B. Zhang, Z. Wei, and D. Xu, 2017: Effects of Averaging Period on Energy Fluxes   |
| 565                      | and the Energy-Balance Ratio as Measured with an Eddy-Covariance System.  |
| 566                      | <i>Boundary-Layer Meteorol.</i> , 165, 545–551, https://doi.org/10.1007/s10546-017-0284-8.  |
| 567                      | Ficci, 2014: Maize in India. India Maize Summit'14, 1–32.   |
|                          |   |

568 Finkelstein, P. L., T. Park, N. Carolina, and F. Sims, 2001: z [( w ' O ') Fe. 106, 3503–3509.

| 569<br>570<br>571 | Finkelstein, P. L., & Sims, P. F. (2001). Sampling error in eddy correlation flux<br>measurements. Journal of Geophysical Research, 106, 3503–3509.<br><u>https://api.semanticscholar.org/CorpusID:128980052</u> |
|-------------------|--|
| 572               | Finnigan, J. J., 2004: A re-evaluation of long-term flux measurement techniques part II:   |
| 573               | Coordinate systems. Boundary-Layer Meteorol., 113, 1-41,   |
| 574               | https://doi.org/10.1023/B:BOUN.0000037348.64252.45.  |
| 575               | Finnigan, J. J., R. Clement, Y. Malhi, R. Leuning, and H. A. Cleugh, 2003: Re-evaluation of  |
| 576               | long-term flux measurement techniques. Part I: Averaging and coordinate rotation.  |
| 577               | Boundary-Layer Meteorol., 107, 1–48, https://doi.org/10.1023/A:1021554900225.  |
| 578               | Foken, T., and B. Wichura, 1996: Tools for quality assessment of surface-based flux  |
| 579               | measurements. Agric. For. Meteorol., 78, 83-105, https://doi.org/10.1016/0168-   |
| 580               | 1923(95)02248-1.   |
| 581               | Foken, T., Göockede, M., Mauder, M., Mahrt, L., Amiro, B., & Munger, W. (2005). Post-  |
| 582               | Field Data Quality Control BT - Handbook of Micrometeorology: A Guide for Surface  |
| 583               | Flux Measurement and Analysis (X. Lee, W. Massman, & B. Law, Eds.; pp. 181-208).   |
| 584               | Springer Netherlands. https://doi.org/10.1007/1-4020-2265-4_9  |
| 585               | Foken, T., M. Aubinet, J. J. Finnigan, M. Y. Leclerc, M. Mauder, and K. T. Paw U, 2011:  |
| 586               | Results of a panel discussion about the energy balance closure correction for trace gases.   |
| 587               | Bull. Am. Meteorol. Soc., 92, https://doi.org/10.1175/2011BAMS3130.1.  |
| 588               | Fong, B. N., Reba, M. L., Teague, T. G., Runkle, B. R. K., & Suvočarev, K. (2020). Eddy  |
| 589               | covariance measurements of carbon dioxide and water fluxes in US mid-south cotton  |
| 590               | production. Agriculture, Ecosystems and Environment, 292.  |
| 591               | https://doi.org/10.1016/j.agee.2019.106813   |
| 592               | Gao, Z., H. Liu, G. G. Katul, and T. Foken, 2017: Non-closure of the surface energy balance  |
| 593               | explained by phase difference between vertical velocity and scalars of large atmospheric   |
| 594               | eddies. Environ. Res. Lett., 12, https://doi.org/10.1088/1748-9326/aa625b.   |
| 595               | Gerken, T., and Coauthors, 2018: Investigating the mechanisms responsible for the lack of  |
| 596               | surface energy balance closure in a central Amazonian tropical rainforest. Agric. For.   |
| 597               | Meteorol., 255, 92-103, https://doi.org/10.1016/j.agrformet.2017.03.023.   |
|                   |  |

598 Kidston, J., Brümmer, C., Black, T. A., Morgenstern, K., Nesic, Z., McCaughey, J. H., &

| 599 | Barr, A. G. (2010). Energy balance closure using eddy covariance above two different           |
|-----|--|
| 600 | land surfaces and implications for CO2 flux measurements. Boundary-Layer                       |
| 601 | Meteorology, 136(2), 193-218. https://doi.org/10.1007/s10546-010-9507-y                        |
| 602 | Kole, C., 2013: Genomics and breeding for climate-resilient crops: Vol. 2 target traits. 1-474 |
| 603 | pp.  |
| 604 | Kottek, M., J. Grieser, C. Beck, B. Rudolf, and F. Rubel, 2006: World map of the Köppen-       |
| 605 | Geiger climate classification updated. Meteorol. Zeitschrift, 15, 259–263,                     |
| 606 | https://doi.org/10.1127/0941-2948/2006/0130.   |
| 607 | Leclerc, M. Y., and T. Foken, 2014: Footprints in micrometeorology and ecology. Springer,.     |
| 608 | Lee, X., Q. Yu, X. Sun, J. Liu, Q. Min, Y. Liu, and X. Zhang, 2004: Micrometeorological        |
| 609 | fluxes under the influence of regional and local advection: A revisit. Agric. For.             |
| 610 | Meteorol., 122, 111–124, https://doi.org/10.1016/j.agrformet.2003.02.001.                      |
| 611 | Leuning, R., E. van Gorsel, W. J. Massman, and P. R. Isaac, 2012: Reflections on the surface   |
| 612 | energy imbalance problem. Agric. For. Meteorol., 156, 65-74,                                   |
| 613 | https://doi.org/10.1016/j.agrformet.2011.12.002.   |
| 614 | Malhi, Y., K. McNaughton, and C. Von Randow, 2004: Low Frequency Atmospheric                   |
| 615 | Transport and Surface Flux Measurements. 101-118, https://doi.org/10.1007/1-4020-              |
| 616 | 2265-4_5.  |
| 617 | Manon, M., and M. Kristian, 2020: Estimating local atmosphere-surface fluxes using eddy        |
| 618 | covariance and numerical Ogive optimization. 15, 21387–21432,                                  |
| 619 | https://doi.org/10.5194/acpd-14-21387-2014.  |
| 620 | Mauder, M., and T. Foken, 2006: Impact of post-field data processing on eddy covariance        |
| 621 | flux estimates and energy balance closure. Meteorol. Zeitschrift, 15, 597-609,                 |
| 622 | https://doi.org/10.1127/0941-2948/2006/0167.   |
| 623 | Medrano, H., and Coauthors, 2015: From leaf to whole-plant water use efficiency (WUE) in       |
| 624 | complex canopies: Limitations of leaf WUE as a selection target. Crop J., 3, 220–228,          |
| 625 | https://doi.org/10.1016/j.cj.2015.04.002.  |
| 626 | Oncley, S. P., and Coauthors, 2007: The energy balance experiment EBEX-2000. Part I:           |
| 627 | Overview and energy balance. Boundary-Layer Meteorol., 123, 1-28,                              |
| 628 | https://doi.org/10.1007/s10546-007-9161-1.   |

| 629 | Peddinti, S. R., and B. P. Kambhammettu, 2019: Dynamics of crop coefficients for citrus      |
|-----|--|
| 630 | orchards of central India using water balance and eddy covariance flux partition             |
| 631 | techniques. Agric. Water Manag., 212, 68-77,   |
| 632 | https://doi.org/10.1016/j.agwat.2018.08.027.   |
| 633 | Peddinti, S. R., Kambhammettu, B. V. N. P., Rodda, S. R., Thumaty, K. C., & Suradhaniwar,    |
| 634 | S. (2020). Dynamics of Ecosystem Water Use Efficiency in Citrus Orchards of Central          |
| 635 | India Using Eddy Covariance and Landsat Measurements. Ecosystems, 23(3), 511-528.            |
| 636 | https://doi.org/10.1007/s10021-019-00416-3   |
| 637 | Reed, D. E., J. M. Frank, B. E. Ewers, and A. R. Desai, 2018: Time dependency of eddy        |
| 638 | covariance site energy balance. Agric. For. Meteorol., 249, 467-478,                         |
| 639 | https://doi.org/10.1016/j.agrformet.2017.08.008.   |
| 640 | Sakai, R. K., D. R. Fitzjarrald, and K. E. Moore, 2001: Importance of low-frequency          |
| 641 | contributions to eddy fluxes observed over rough surfaces. J. Appl. Meteorol., 40, 2178-     |
| 642 | 2192, https://doi.org/10.1175/1520-0450(2001)040<2178:IOLFCT>2.0.CO;2.                       |
| 643 | Sharma, B. R., Gulati, Mohan, Gayathri, Manchanda, S., Ray, I., & Amarasinghe, U. (2018).    |
| 644 | Water Productivity Mapping of Major Indian Crops. NABARD and ICRIER, 4(1), 88-               |
| 645 | 100.   |
| 646 | Shankar, V., C. S. P. Ojha, and K. S. H. Prasad, 2012: Irrigation Scheduling for Maize and   |
| 647 | Indian-mustard based on Daily Crop Water Requirement in a Semi- Arid Region. 6,              |
| 648 | 476–485.   |
| 649 | Soujanya, B., B. B. Naik, M. U. Devi, T. L. Neelima, and A. Biswal, 2021: Dry Matter         |
| 650 | Production and Nitrogen Uptake as Influenced by Irrigation and Nitrogen Levels in            |
| 651 | Maize. Int. J. Environ. Clim. Chang., 155-161,   |
| 652 | https://doi.org/10.9734/ijecc/2021/v11i1130528.  |
| 653 | Sun, X. M., Z. L. Zhu, X. F. Wen, G. F. Yuan, and G. R. Yu, 2006: The impact of averaging    |
| 654 | period on eddy fluxes observed at ChinaFLUX sites. Agric. For. Meteorol., 137, 188-          |
| 655 | 193, https://doi.org/10.1016/j.agrformet.2006.02.012.  |
| 656 | Tang, X., Z. Ding, H. Li, X. Li, J. Luo, J. Xie, and D. Chen, 2015: Characterizing ecosystem |
| 657 | water-use efficiency of croplands with eddy covariance measurements and MODIS                |

658 products. *Ecol. Eng.*, **85**, 212–217, https://doi.org/10.1016/j.ecoleng.2015.09.078.

| 659 | Tong, X., J. Zhang, P. Meng, J. Li, and N. Zheng, 2014: Ecosystem water use efficiency in a |
|-----|---|
| 660 | warm-temperate mixed plantation in the North China. J. Hydrol., 512, 221-228,               |
| 661 | https://doi.org/10.1016/j.jhydrol.2014.02.042.  |
| 662 | Tong, X. J., J. Li, Q. Yu, and Z. Qin, 2009: Ecosystem water use efficiency in an irrigated |
| 662 |   |
| 663 | cropland in the North China Plain. J. Hydrol., <b>374</b> , 329–337,                        |
| 664 | https://doi.org/10.1016/j.jhydrol.2009.06.030.  |
| 665 | Twine, T. E., and Coauthors, 2000: Correcting eddy-covariance flux underestimates over a    |
| 666 | grassland. Agric. For. Meteorol., 103, 279-300, https://doi.org/10.1016/S0168-              |
| 667 | 1923(00)00123-4.  |
| 668 | Vickers, D., and L. Mahrt, 1997: graphDBSummary. J. Atmos. Ocean. Technol., 14, 512         |
| 669 | <del>526.</del>   |
| 670 |   |
| 670 | Vickers, D., & Mahrt, L. (1997). Quality control and flux sampling problems for tower and   |
| 671 | aircraft data. Journal of Atmospheric and Oceanic Technology, 14(3), 512–526.               |
| 672 | https://doi.org/10.1175/1520-0426(1997)014<0512:QCAFSP>2.0.CO;2                             |
| 673 | Wang, X., C. Wang, and B. Bond-Lamberty, 2017: Quantifying and reducing the differences     |
| 674 | in forest CO2-fluxes estimated by eddy covariance, biometric and chamber methods: A         |
| 675 | global synthesis. Agric. For. Meteorol., 247, 93-103,                                       |
| 676 | https://doi.org/10.1016/j.agrformet.2017.07.023.  |
| 677 | Wilson, K., E. Falge, M. Aubinet, and D. Baldocchi, 2002: DigitalCommons @ University of    |
| 678 | Nebraska - Lincoln Energy Balance Closure at FLUXNET Sites. Agric. For. Meteorol.,          |
| 679 | 223–243.  |
| 680 | Zhang, P., G. Yuan, and Z. Zhu, 2013: Determination of the averaging period of eddy         |
| 681 | covariance measurement and its influences on the calculation of fluxes in desert riparian   |
|     |   |
| 682 | forest. Arid L. Geogr., <b>36</b> , 401–407.  |