Characterisation of the multi-scheme chemical ionisation inlet-2 and the detection of gaseous iodine species <u>using bromide and nitrate</u>

chemical ionisation methods

Xu-Cheng He^{1,2}, Jiali Shen¹, Siddharth Iyer³, Paxton Juuti⁴, Jiangyi Zhang¹, Mrisha Koirala⁵, Mikko M. Kytökari⁵, Douglas R. Worsnop^{1,6}, Matti Rissanen^{3,5}, Markku Kulmala^{1,7,8,9}, Norbert M. Maier⁵, Jyri Mikkilä⁴, Mikko Sipilä¹, and Juha Kangasluoma^{1,4} ¹Institute for Atmospheric and Earth System Research/Physics, Faculty of Science, University of Helsinki, 00014 Helsinki, Finland ²Finnish Meteorological Institute, 00560 Helsinki, Finland ³Aerosol Physics Laboratory, Faculty of Engineering and Natural Sciences, Tampere University, 33014 Tampere, Finland ⁴Karsa Ltd., 00560 Helsinki, Finland ⁵Department of Chemistry, Faculty of Science, University of Helsinki, 00014 Helsinki, Finland ⁶Aerodyne Research, Inc., Billerica, 01821 MA, USA ⁷Helsinki Institute of Physics, University of Helsinki, 00014 Helsinki, Finland ⁸Joint International Research Laboratory of Atmospheric and Earth System Sciences, School of Atmospheric Sciences, Naniing University, 210023 Naniing, China ⁹Aerosol and Haze Laboratory, Beijing Advanced Innovation Center for Soft Matter Science and Engineering, Beijing University of Chemical Technology, 100029 Beijing, China Correspondence: Jiali Shen (jiali.shen@helsinki.fi) and Xu-Cheng He (xucheng.he@helsinki.fi)

Abstract. The multi-scheme chemical ionisation inlet 1 (MION1) allows fast switching between measuring enables rapid switching between the measurement of atmospheric ions without chemical ionisation and neutral molecules by multiple using various atmospheric pressure chemical ionisation methods. In this study, the upgraded we introduce the upgraded version, the multi-scheme chemical ionisation inlet 2 (MION2) is presented. The new design features improved ion opticsthat increase the

5 incorporates enhanced ion optics, resulting in increased reagent ion concentration, a generally more robust operation and the possibility to run ensuring a robust operation, and enabling the use of multiple chemical ionisation methods with the same ionisation time.

To In order to simplify the regular calibration of MION2, we developed an open-source flow reactor chemistry model (MARFORCE) to quantify called MARFORCE. This model enables quantification of the chemical production of sulfuric acid

- 10 (H₂SO₄), hypoiodous acid (HOI), and hydroperoxyl radical (HO₂). MARFORCE simulates the convection-diffusion-reaction processes inside occurring within typical cylindrical flow reactors with uniform inner diameters. The model also provides includes options to simulate the chemical processes chemical processes in two scenarios: 1) when two flow reactors with different inner diameters are connectedtogether-, and 2) when two flows are merged into one (connected by-using a Y-shape tee), but, although with reduced accuracy. AdditionallyFurthermore, the chemical mechanism files in the model are compatible
- 15 with the widely-used Master Chemical Mechanism , thus allowing (MCM), allowing for future adaptation to simulate other chemical processes in flow reactors.

We further carried out detailed Furthermore, we conducted a comprehensive characterisation of the bromide (Br⁻) and nitrate (NO₃⁻) chemical ionisation methods with different ionisation times. We calibrated performed calibration experiments for H₂SO₄, HOI, and HO₂ by combining gas kinetic experiments with the MARFORCE model. Sulfur The evaluation of sulfur dioxide (SO₂), water (H₂O), and molecular iodine (I₂) were evaluated using involved dilution experiments from a gas cylinder (SO₂), dew point mirror measurements (H₂O), and a derivatization approach in combination combined with high-performance

liquid chromatography quantification (I_2) , respectively. We find-

<u>Our findings indicate</u> that the detection limit is <u>negatively-inversely</u> correlated with the fragmentation enthalpy of the analytereagent ion (Br⁻) cluster, i. e., a. In other words, stronger binding (resulting in a larger fragmentation enthalpy) leads to a lower

- detection limit. Additionally, a moderately longer reaction ionisation time enhances the detection sensitivity thus decreasing, thereby reducing the detection limit. For example, the instance, when using the Br⁻ chemical ionisation method with a 300 ms ionisation time, the estimated detection limit for H₂SO₄ is estimated to be 2.9×10^4 molec. cm⁻³ with a 300 ms ionisation time. A direct comparison suggests that this is even better than. Notably, this detection limit is even superior to that achieved by the widely-used Eisele-type chemical ionisation inlet, as revealed by direct comparisons.
- 30 While the NO₃⁻ chemical ionisation method is generally more robust, we find remains stable in the presence of high humidity, we have observed that the Br⁻ chemical ionisation method (Br⁻-MION2) is significantly affected by air water content. Higher air water content results in lower levels of air water lead to reduced sensitivity for HO₂ and SO₂ within under the examined conditions. On the other hand, a steep sensitivity drop of However, we have found that a sharp decline in sensitivity for H₂SO₄, HOI, and I₂ is only observed occurs only when the dew point is greater than exceeds 0.5-10.5 °C (equivalent to
- 35 20-40 % RH;-, calculated at 25 °C hereafter). Future studies utilising throughout this manuscript). For future studies utilising the atmospheric pressure Br⁻ chemical ionisation method, including Br⁻-MION2, should carefully address the humidity effect on a molecular basisit is crucial to carefully consider the molecular-level effects of humidity. By combining methods such as approaches such as the water-insensitive NO₃⁻-MION2 with Br⁻-MION2, MION2 should be able to provide greater details of air composition than either of these methods can offer more comprehensive insights into the composition of atmospheric air
- 40 than what can be achieved by either method alone.

Combining By employing instrument voltage-scanning, chemical kinetic experiments, and quantum chemical calculations, we find that the detection is not interfered with by iodine oxides under atmospherically relevant conditions. The have conclusively established that the presence of iodine oxides does not interfere with the detection of HIO₃. Our comprehensive analysis reveals that the ions IO_3^- , and ions measured $HIO_3 \cdot NO_3^-$, and $HIO_3 \cdot Br^-$, which are detected using the Br^- and NO_3^- chemical ion-

45 isation methods, are primarily, if not exclusively, produced generated from gaseous HIO₃ molecules – within atmospherically relevant conditions.

1 Introduction

20

Chemical ionisation mass spectrometer spectrometry (CIMS) has been widely used in atmospheric chemistry and aerosol formation studies due to its versatility and high sensitivity in measuring trace level gaseous species (see e.g., Eisele and Tanner 50 (1993); Munson and Field (1966); Hansel et al. (1995); Huey (2007); Kirkby et al. (2011); Ehn et al. (2014); Lee et al. (2014); Berndt et al. (2016); Sipilä et al. (2016); Laskin et al. (2018); He et al. (2021b)). With chemical ionisation methods, an analyte is charged either by 1) receiving charge (proton, electron or ion) from the reagent ion or 2) forming a relatively stable cluster with the reagent ion. Mass spectrometers further measure the charged analyte-containing ions to obtain their molecular information.

55

Various chemicals have been used to produce reagent ions ; the most employed as reagent ions in chemical ionisation methods. The commonly used reagent ions include nitrate (NO₃⁻, Eisele and Tanner (1993)), acetate (C₂H₃O₂⁻, Veres et al. (2008)), iodide (I⁻, Caldwell et al. (1989)), hydronium (H₃O⁺, Lagg et al. (1994)), and sporadically also e.g., bromide (Br⁻, Caldwell et al. (1989)) and ammonium (NH₄⁺, Westmore and Alauddin (1986)). These reagent ions transfer charges to or form clusters with distinct subsets of analytes of interest. The trace gases. However, the detection of an analyte-containing ion is additionally limited influenced by its transmission through the ion optics of the mass spectrometers, since mass spectrometer's ion optics, as collision-induced cluster fragmentation may remove the signature of the original analyte. A strongly bonded analyte-reagent ion cluster has a substantially higher chance to reach can diminish the analyte's signature. Analyte-reagent ion clusters with strong bonds have a higher likelihood of reaching the detector compared to weakly bonded ones (Passananti et al., 2019)

- 65 . Therefore, a smart selection of a clusters (Passananti et al., 2019). Hence, it is crucial to select a chemical ionisation method that can maintain the signature of the analyteis desired preserves the analyte's signature. For example, the NO₃⁻-CIMS has been widely used to detect sulfuric acid (H₂SO₄) (Eisele and Tanner, 1993) and highly oxygenated organic molecules (Ehn et al., 2014). I⁻-CIMS is regularly used to detect semi-volatile organic compounds (Lee et al., 2014), bromine and chlorine-containing species (Liao et al., 2014; Wang et al., 2019) and e.g., dinitrogen pentoxide (N₂O₅) (Thornton et al., 2010). C₂H₃O₂⁻-CIMS
- ⁷⁰ was used to detect small organic acids (Veres et al., 2008) and highly oxygenated organic compounds (Berndt et al., 2016). The bromide chemical ionisation method has recently been used to detect species such as HO₂ (Sanchez et al., 2016) and H₂SO₄ (Wang et al., 2021a). The detection of a series of halogenated species by the Br⁻ chemical ionisation method was first demonstrated by He (2017). Detailed characterisation of the Br⁻ chemical ionisation method utilising the multi-scheme chemical ionisation inlet 1 (MION1) was presented in several of our earlier studies (Wang et al., 2021a; Tham et al., 2021; He
- ret al., 2021b). Multiple species were successfully calibrated using either analytical methods or inter-instrument comparison, including H_2SO_4 , I_2 , Cl_2 and HOI (Tham et al., 2021; Wang et al., 2021a). Among the calibrated species, H_2SO_4 and I_2 were shown to be detected at the collision limit (highest sensitivity). Although H_2SO_4 has been quantified using standard methods (Kürten et al., 2012), the quantification of the measured $I_2 \cdot Br^-$ signal remains challenging. This is primarily contributed by two factors: 1) the current Br^- -MION1/2 have a detection upper limit of a few hundred ppty of I_2 , beyond which the reagent
- 80 ions get depleted and the measurement is non-linear, 2) on the contrary, spectroscopic and other methods could be limited by their high detection limits and may not be able to detect I_2 at appropriate levels. Therefore, the key is to find sensitive methods to quantify gaseous I_2 at tens to hundreds of pptv levels.

Ideally, all of simultaneous measurement of all the mentioned analytes can be measured simultaneously by utilising all

- 85 of the could be achieved by employing their corresponding CIMS methods at the same time concurrently in ambient observations or in-complex laboratory experiments. However, CIMS instruments are expensive costly, and research institutes are regularly limited by available instrumentation. An alternative approach suing often face limitations due to the availability of such instrumentation. As an alternative approach, chemical ionisation inlets expansive of switching reagent ions that have the capability to switch between different reagent ions can be employed. Many switchable systems have been developed previously,
- 90 such as for proton transfer reaction mass spectrometers (Jordan et al., 2009; Breitenlechner et al., 2017; Pan et al., 2017) and other chemical ionisation mass spectrometers (Hearn and Smith, 2004; Smith and Španěl, 2005; Agarwal et al., 2014; Brophy and Farmer, 2015). A common feature of these techniques is using a reduced-pressure ion-molecule reaction chamber, thus unavoidably diluting the gas molecules of interest by orders of magnitude. While the detection limit of an instrument is also affected by other factors, it is commonly observed that chemical ionisation inlets operating at reduced pressures have higher
- 95 limits of detection compared to atmospheric pressure chemical ionisation inlets. For instance, reduced pressure inlets reported detection limits of various organic compounds from sub-pptv (parts per trillion by volume) to hundreds of pptv levels (Lee et al., 2014; Brophy and Farmer, 2015), while the best-performing atmospheric pressure chemical ionisation inlets regularly detect vapours at ppqv (parts per quadrillion by volume) levels for selected acids and highly oxygenated organic vapours using the same time-of-flight mass spectrometer (Jokinen et al., 2012; Ehn et al., 2014; He et al., 2021b).

100

105

110

To lower the limit of detection of reduce the detection limit of switchable reagent ion chemical ionisation systems, we developed the MION1 inlet, which allows for fast switching of reagent ion chemistry operating at atmospheric pressure (Rissanen et al., 2019). This technique has mostly been deployed to detect predominantly been employed for the detection of sulfuric acid and halogenated species using either the NO_3^- or Br^- chemical ionisation methods (Rissanen et al., 2019; Tham et al., 2021; He et al., 2021b; Wang et al., 2021a; Finkenzeller et al., 2022).

However, the remaining problems of However, there are some remaining issues with the MION1inelude 1) the. Firstly, its limit of detection of the MION1 is lower is higher compared to another widely commonly used atmospheric pressure chemical ionisation inlet (noted as referred to as the "Eisele inlet" hereafter) (Eisele and Tanner, 1993; Jokinen et al., 2012; Wang et al., 2021a)and 2) the ion-molecule reaction. Secondly, the ionisation times for different chemical ionisation methods had have to be different , owing due to the designthat, which involves aligning and attaching the chemical ionisation units are aligned and attached at varying distances on a cylindrical tubeat varying distances. These problems could limit its applicability to the detection of. These challenges may restrict its suitability for detecting vapours at extremely low concentrations (e.g., at 10⁵ to

 10^6 molec. cm⁻³ or 5 to 50 4 to 40 ppqv) and the interpretation of the interpreting the species detected by different chemical ionisation methods.

115

In this study, we present an upgraded MION inlet (noted introduce an upgraded version of the MION inlet, referred to as "MION2"hereafter) which addresses these problems. Laboratory experiments were carried out to characterise, which specifically addresses these issues. We conducted laboratory experiments to characterise the performance of this inlet using analytical methods and a newly-developed open-source kinetic model. As halogen anion-based chemical ionisation methods

120 (e.g., based on halogen anions (such as I⁻) are commonly affected by air water content influenced by the water content in the air (Kercher et al., 2009; Mielke et al., 2011; Woodward-Massey et al., 2014; Lee et al., 2014), we additionally performed systematic examinations of also systematically examined the impact of air water content on the detection of Br⁻-MION2.

2 Methods

125 2.1 Characterisation of the MION2 inlet

The ionisation inlet used in this paper utilised in this study is the upgraded multi-scheme chemical ionisation inlet, MION2(, developed by Karsa Ltd.). The inlet is designed to be capable of measuring This inlet is specifically designed to enable the measurement of neutral molecules using chemical ionisation methods and detecting while also facilitating the detection of atmospheric ions by disabling chemical ionisation. It allows fast offers the capability of rapid switching between two or more (up

130 to six) chemical ionisation methods to selectively measure, allowing for selective measurement of gaseous species at ambient pressure. The inlet accommodates three bipolar ion sourcesper reaction time, which Currently, the MION2 inlet supports up to six ion sources.

Due to the geometric limitations of the previous MION1 inlet, the different ionisation sources in MION1 have to employ

135 different ionisation times. The ionisation time is defined by the sample flow rate and the distance from between the ion injection port to and the instrument pinhole (refer to Figure 1). For the longer reaction. The improved geometry of the MION2 inlet overcomes this limitation, allowing for the operation of three bipolar ion sources per ionisation time, all positioned at the same distance from the pinhole. Additionally, for the longer ionisation time, the length of the connecting pipe between the sources can be adjusted, i.e., providing flexibility in modifying the ionisation timeean be modified. Therefore, the new design
140 of MION2 allows it to operate two chemical ionisation methods with the same ionisation time to allow a direct comparison which was not possible with the MION1.

In this study, we deployed the employed the MION2 inlet with two chemical ionisation methods(, namely NO_3^- and Br^-) and two reaction times to understand the inlet characteristics. To clarify the different positions of the ionisation inlet, we refer

- 145 to, along with two different ionisation times (35 and 300 milliseconds, ms, respectively). This configuration was chosen to investigate the characteristics of the inlet. To facilitate clear referencing, we designate the ion source positioned 3 cm away from the mass spectrometer as tower 1 (T1), and the source while the source located 25 cm away from the mass spectrometer is referred to as tower 2 (T2) throughout the paper (see Figure 1).
- 150 Supplementary Figure A1 presents Figure A1 illustrates the conceptual schematic for of one of the ion sources with , depicting the airflow and ion paths. The whole entire source is attached to an electrically grounded a 24 mm inner diame-

ter tube with that is electrically grounded. The sample flow, which is provided by a mass flow controller (MFC) connected to a vacuum pump, is set at a rate of 22.5 standard litres per minute (slpm)sample flow, where the target molecules are charged. The target molecules undergo ionisation by reacting with the reagent ions (NO₃⁻ or Br⁻). In this case, the reaction time of the target target target the target targ

155 targeted

In this configuration, the ionisation time for the target molecules and charged reagent ions for tower 1 was ca. is approximately 35 ms, and for tower 2, for tower 1 and 300 ms -for tower 2. A neutral reagent inflow is a stream introduced, which consists of nitrogen or air enriched with reagent vapour(or-. The reagent vapour is generated by passing nitrogen or air over liquid reagent (nitric acid, HNO₃ in this paper), which, or dibromomethane, CH₂Br₂, in this study). The resulting mixture is then fed into the ion sourceand, where it is ionised by a soft x-ray-X-ray source (Hamamatsu L12535, 4.9 keV). The charged reagent ions are then guided by an electric field guided into the sample flow. The electric field inside by an electric field within the ion sourceis created with. This electric field is generated using concentric stainless steel electrode plates with different size orifices (diameters between orifices of different sizes (ranging from 5-10 mm), added resistances between every two plates, and two in diameter), with resistances placed between each pair of plates. Two high voltages (ca. approximately 2500 V and 250 V) are used in the inlet. The lower of these the two voltages determines if whether the reagent ions pass through the last final orifice in the deflector electrode into the sample flow, effectively turning the ionisation on and off, and enabling the fast,

170 Compared with In contrast to the MION1 (Rissanen et al., 2019), where the source reagent flow was defined by the neutral reagent inflow and the outflow (i.e., exhaust), design (Rissanen et al., 2019), which relied on the reagent inflow and exhaust flow to define the source reagent flow, MION2 has incorporates an additional purge flow to prevent the sample flow from entering the ion source. The purge flow consists of the same nitrogen , or synthetic air , that is used to create or air used to generate the reagent flow. The purge flow splits in two upon Upon entering the ion source: one preventing the sample from entering and

effectively controlling the ionisation process and enabling the rapid switching between ion sources.

175 the other keeping, the purge flow splits into two streams: one stream prevents the sample flow from entering, while the other stream ensures that the neutral reagent from entering does not enter the sample flow. With the

In MION2, the typical flow rates for the reagent, purge, and exhaust are 10, 100, and 50 standard cubic centimetres per minute (sccm), respectively. The reagent concentration in the ion source is estimated to be 2×10^{17} cm⁻³. This design effectively addresses the challenges faced in MION1, one had to choose between possibly where a compromise had to be made between the risk of contaminating the sample pipe with the neutral reagent, or pulling neutral reagent or introducing sample air into the ion sourcepotentially contaminating the ion source, e. g., it sometimes results in salt formation in the ion source or potentially results in uncontrolled ion chemistry within the inlet leading to detection biases, potentially leading to contamination or uncontrolled ion chemistry and resulting in detection biases. In MION2, the water vapour and other contaminants in the sample flow do not have the opportunity to oxidise the surfaces of the electrodes inside the ion source.

Such oxidation would result in reduced ion transmission from the ion source to the sample flow. Operational testing in ambient

measurements indeed showed during ambient measurements has demonstrated that MION2 is significantly more stable than exhibits significantly improved stability compared to MION1. The For example, recent measurements conducted at a coastal site in Finland involved the uninterrupted operation of MION2 for at least two months.

190

195

Additionally, the upgraded ion optics inside the ion sources in of MION2 have been upgraded which increase the reagent ion transmission compared to the MION1 by about an increased the transmission of reagent ions and the observed reagent ion concentration at the mass spectrometer by approximately one order of magnitude -compared to MION1. This improvement was achieved by modifying the last electrode within the ion source to minimize ion residence time and reduce diffusion losses of ions.



Figure 1. Schematic of the MION2 inlet illustrating its gas flows and ion paths. The new design increases the primary ion concentration and allows the operation of multiple chemical ionisation methods with the same ionisation time. L1 and L2 refer to the distances between the ion sources and the pinhole of the mass spectrometer. The exhaust flows are connected to two ports in the middle of the ion source.

2.2 Experimental setup

2.2.1 Calibration of inorganic species

In order to characterise. The experimental setup used for characterising the MION2 inlet , we utilised an experimental setup as shown is illustrated in Figure A2, which comprises three. It consists of three main sections: the flow reactor section, the MION2 chemical ionisation inlet (MION2) section, and an Atmospheric Pressure interface Time-of-Flight mass spectrometer (APi-TOF, Aerodyne Inc., Junninen et al. (2010)).

The flow reactor section consists of includes a calibration source and various several gas feeds. Synthetic air (Woikoski OY, 205 Finland; purity >=≥ 99.999 % with 20.9 % O₂), nitrogen (N₂, Woikoski OY, Finland; purity >=≥ 99.999 %), and sulfur dioxide (SO₂, Air Products, USA; 99.5 % purity) were injected by into the system using mass flow controllers (MFCs) connected to standard gas cylinders or tanksand they . These gases were pre-mixed before reaching the calibration source. Molecular iodine () was produced

I₂ was generated either from a homemade permeation tube or a commercial permeation tube (VICI Metronics)by blowing a . A stream of nitrogen (50 millilitres-per-minute, mlpm) sccm) was passed over the permeation tube at controlled temperatures (ranging from 120 to 140 °C). The temperature of the permeation tubes was controlled by regulated using an electronically controlled heating mantleto enable, allowing for adjustable yet stable iodine concentrations. Water vapour (H₂O) was controlled by an adjustable stream of nitrogen passing adjusting the flow of nitrogen through a water bubbler, providing a controllable source of humidity.

The calibration source was mainly used to calibrate H₂SO₄, HO₂ and HOI. The H₂SO₄ calibration has been was detailed in Kürten et al. (2012) and the HOI calibration has been was presented in Tham et al. (2021) and Wang et al. (2021a). Briefly, a known amount of radical is produced by a mercury lamp (UVP Pen-Ray)in The calibration source in the experimental setup
was constructed using an aluminum box that encloses a 3-quarter-inch quartz tube. The quartz tube was chosen for its high transmission properties for ultra-violet (UV) light emitted from a mercury lamp. Adjacent to the quartz tube, the mercury lamp is housed in an aluminum block that contains a filter-covered hole. The filter used in the aluminum block allows for high transmission of 185 nm light emitted from the lamp. This specific wavelength of light is effective in photolysing H₂O molecules, generating OH radicals.

225

Before conducting the calibration experiment, a mixed flow of nitrogen (N_2), oxygen (O_2), H_2O , and either SO_2 or I_2 was continuously passed through the calibration sourcewhich further reacts with. This flow ensures that the source is thoroughly flushed with the desired gases and vapours, creating a controlled environment for subsequent calibration measurements. The produced OH radicals from the calibration source then undergo reactions with an excess amount of SO_2 or a moderate amount of I_2 to produce either H_2SO_4 or HOI as the final products. As, respectively, Additionally, the HO₂ radical is produced as a

230

To quantify the concentrations of H_2SO_4 , HOI, and HO_2 were predicted using a numerical model adopting, an open-source Python library based on two-dimensional convection-diffusion-reaction equations was developed (Marine Atmospheric paRticle

235 FORmation and ChEmistry, MARFORCE, Shen and He (2023)). This library aims to provide a framework for performing similar calibration tasks. Furthermore, it also allows users to simulate and predict concentrations of other chemical species by adopting different chemical reaction schemes. The MARFORCE library can be used as a tool in future research endeavours involving flow reactor chemistry simulations. As the numerical model needed to do the calibration is not vet available to the public, this study further develops an open-source Python library for such tasks (Shen and He, 2023).

240

The SO_2 calibration is relatively straightforward the calibrator source is not needed straightforward. The SO_2 flow from the SO₂ gas cylinder was diluted with humidified nitrogen and the mixed sample was fed into the inlet. The measured normalised $SO_2 \cdot Br^-$ signal is was further compared with the estimated SO_2 concentration to derive a calibration factor.

245 Absolute concentration was calibrated using To calibrate the absolute concentration of H_2O , a dew point mirror hygrometer (DewMaster Chilled Mirror Hygrometer, EdgeTec) was employed. The dew point mirror drew and measured a hygrometer drew a sample from a branch of the humidified flow before it entered the MION2 inlet tube. By measuring the dew point temperature, the dew point mirror hygrometer provides an accurate and reliable determination of the absolute H_2O concentration in the sample. This calibration method ensures precise measurement of H₂O concentration, which is important for accurate analysis 250 and interpretation of the experimental data.

2.2.2 Calibration of molecular iodine

Although chemical ionisation method has been shown to be extremely sensitive in the molecular iodine () detection (He, 2017; Wang et al., 2) , the quantification of the measured To calibrate the measured signals of $I_2 \cdot Br^-$ signal remains challenging. This is primarily contributed by two factors: 1) the current in Br -- MION1/2 have a detection upper limit of a few hundred pptv of , beyond which 255 the reagent ions get depleted and the measurement is non-linear, 2) on the contrary, spectroscopic and other methods could be limited by their high detection limits and may not be able to detect -MION2, we acquired its stable signals by utilising I_2 at appropriate levels. Therefore, the key is to find sensitive methods to quantify gaseous at tens to hundreds of pptv levelsemitted from a permeation tube, which was regulated at a constant temperature and subjected to a continuous nitrogen stream (50 sccm).

260

The key to this calibration is determining the quantities of I_2 emitted from the permeation tube. We have previously calibrated the I₂ measurement of Br⁻-MION1 using a cavity-enhanced differential optical absorption spectroscopy (CE-DOAS) instrument (Wang et al., 2021a), an UV/Vis spectrophotometer and an inductively coupled plasma mass spectrometer (ICP-MS) (Tham et al., 2021; Wang et al., 2021a). As none of these instruments is was available for this study, we further adapted 270

The collection of the I₂ sample followed exactly the same procedure as described in our previous studies (Tham et al., 2021; Wang et al., 2021a). Briefly, 50 standard cubic centimetres per minute (seem.) sccm nitrogen carrier gas flow was passed through an I₂ permeation tube for 300 minimized, under 120-140 °C. The nitrogen carrier stream containing the released I₂ was bubbled through an Schlenk-type impinger charged with 20 mL of hexane kept at -70 °C by a dry ice/acetone bath. After completion of the sampling process, the absorption flasks were allowed to warm to the ambient temperature and sealed with a Teflon-coated glass stopper. The solution was stored at 4 °C until further processing.

Inspired by Mishra et al. (2000), I₂ was converted into an essentially-a non-volatile and stable derivative, followed by quantification of the latter using gas or liquid chromatography. Mishra et al. (2000) quantified I₂ in aqueous matrices by gas chromatography-mass spectrometry(GC/MS) after I₂ reacting with 2,6-dimethylaniline to form the corresponding 4-iododerivative.

An adaptation of this method was required as the iodine to be determined was diluted in hexane. Specifically, the iodine derivatisation derivatization reaction was conducted directly with the hexane solutions in the presence of an aqueous buffer, to reduce losses associated with a hexane-to-water transfer. To avoid any losses of the volatile I_2 through evaporation, the reaction was conducted in hermetically sealed headspace vials, with efficient phase mass transfer being facilitated by vigorous magnetic stirring.

285 Control of the pH of the buffer was crucial for achieving high derivatisation derivatization yields, with pH at 7.00 providing the most favourable level of conversion after 2 hours. Attempts to perform the derivatisation derivatization reaction under homogeneous conditions in hexane in the presence of a variety of soluble organic bases (e.g., tertiary amines) returned poor yields and led to the formation of several side products, most probably due to iodine oxidation. Experiments using 1.00 mL aliquots of the I₂ sample solutions under investigation produced the derivative at the limit of detection, precluding a reliable quantification of the derivative by the reverse phase high-performance liquid chromatography (RP-HPLC/UV).

To improve the analytical sensitivity, 10 mL aliquots of the iodine sample solution were employed for derivatisation derivatization. To boost the sensitivity further, a high volume (15 μ L) of the concentrated derivatisation derivatization solution was injected into the HPLC system. Unfortunately, the hexane in the injection solution and the high injection volume gave rise to retention time instability and peak distortion. Subsequent optimisation of the chromatographic method provided robust reverse phase chromatographic conditions. Specifically, this was achieved by using relatively weakly eluting isocratic conditions for sample elution, followed by strongly eluting conditions for column cleaning and reconditioning. Using the fully optimised protocol, the derivative could be readily quantified for 0.17 to 11.05 μ g mL⁻¹ initial iodine concentrations, with the LOD (limit of detection) and LOQ (limit of quantification) being 0.012 μ g mL⁻¹ and 0.035 μ g mL⁻¹. Using this method, the hexane solution obtained

by absorption of iodine from the permeation tube was found to contain 0.26 μ g iodine mL⁻¹. Considering a total sample vol-300 ume of 20 mL, the iodine output rate of the permeation tube under the employed conditions was calculated to be 17.3 ng min⁻¹.

It is worth noting that the sensitivity of the current method can be further improved by employing more sensitive separation and/or detection techniques, e.g., liquid chromatography-mass spectrometry (LC/MS) or GC/MS.

305

310

2.2.3 Humidity dependence of analyte detection

An important part integral aspect of the characterisation is examining the impact involves investigating the influence of water on the detection of MION2 when utilising employing the bromide chemical ionisation method. Since is needed As water is essential in the calibration source to produce generate OH radicals which in turn produce, which subsequently yield either H₂SO₄ or HOI, we additionally added a dilution flow which joins after incorporated an additional dilution flow that merges with the calibration source using through a Y piece (see refer to Figure A3). This experimental design enables changing the sample absolute humidity configuration allows for the alteration of the absolute humidity of the sample, independent of the OH production rate in the calibration source. During the experiments with varying humidity, the total flows of the dilution part and the flow reactor section were kept constant, while the relative humidity of the dilution flow was varied by mixing different

315 combinations of dry and humidified flows. In this way, we were able to keep the systematic errors, resulting from the mixing By employing this approach, we maintained a consistent level of systematic errors arising from the blending of the dilution and sample flows, constant. By comparing the relative signal intensities of analyte-containing ions, we could examine the influence of water on the detection of different analytes (e.g., H₂SO₄, HOI and HO₂).

320 2.2.4 Quartz flow reactor setup

In order to study the sensitivity of Br^- -MION2 to other oxidised iodine species, e.g., IO, OIO, HIO₃, I₂O₃, I₂O₄ and HIO₂, a quartz flow reactor with an outer diameter of 2.54 inner diameter of 2.4 cm and a length of 94 cm was used. The residence time inside the quartz tube was 8.5 s. A green LED, with a wavelength of 528 nm, was hung on top and in parallel to the quartz flow reactor to initiate iodine photo-chemistryphotochemistry. In order to keep the temperature and light uniformity in the quartz flow reactor, the flow reactor was wrapped together with the green LED light by aluminium foil. The schematics of the setup

325 are shown in supplementary Figure A4.

2.3 MARFORCE model description

As described above, calibration of H₂SO₄, HO₂, and HOI requires a numerical model to simulate the chemical dynamics in the calibration source and inlet tube. This process can be simplified into a two-dimensional convection-diffusion-reaction prob-

330

lem. The concept of such a model was illustrated elsewhere (Kürten et al., 2012), specifically for the calibration of H_2SO_4 . Our earlier studies also presented a numerical model for HOI calibration with similar principles but a simplified iodine chemistry scheme was instead implemented (Tham et al., 2021; Wang et al., 2021a). HoweverNevertheless, neither of these studies made the calibration scripts openly available their calibration scripts publicly accessible, and the scripts are not flexible in

- 335 adopting-lack adaptability for different chemistry schemes. Therefore, we Consequently, we have developed an open-source two-dimensional flow reactor model (MARFORCE, Marine Atmospheric paRticle FORmation and ChEmistry) for these tasks. MARFORCE was named MARFORCE to address these limitations. MARFORCE is built in Python and is hosted in GitHub (Shen and He, 2023). Thus it can be freely accessed. The model has two major hosted on GitHub (Shen and He, 2023), allowing free access to interested users. The model comprises two main components: 1) the fluid dynamics simulation module
- 340 and 2) the gas-phase photo-chemistry module.photochemistry module.

2.3.1 Convection-diffusion-reaction equation

The MARFORCE model utilises a two-dimensional convection-diffusion-reaction equation to simulate the fluid dynamics, photo-chemistry and chemical reactions in a cylindrical flow reactor. The convection-diffusion-reaction equation has been

345 derived and discussed extensively in the literature (Gormley and Kennedy, 1948; Kürten et al., 2012) and is only briefly discussed here:

$$\frac{\partial c_i}{\partial t} = D_i \left(\frac{1}{r}\frac{\partial c_i}{\partial r} + \frac{\partial^2 c_i}{\partial r^2} + \frac{\partial^2 c_i}{\partial z^2}\right) - \frac{2Q}{\pi R^2} \left(1 - \frac{r^2}{R^2}\right)\frac{\partial c_i}{\partial z} + P \tag{1}$$

where *i* corresponds to a specific chemical (e.g., H₂SO₄), the *c_i* is the concentration, *D_i* is the diffusion coefficient, *r* is the distance in the radial direction, *R* is the radius of the flow reactor, *Q* is the total flow in the flow reactor, *z* is the distance in
the axial direction and *P* shows the production (positive values) or loss (negative values) rate due to chemical reactions. As the flow in tangential direction is symmetrical, the ¹/<sub>r²</sup> ^{∂²c_i} term has been ignored.
</sub>

The diffusion coefficient in the model can be defined in two-three ways: 1) manually defined using experimental values or 2) calculated by kinetic theory or 23) calculated based on elemental composition using Fuller's method (Fuller et al., 1966).

355

The convection and diffusion processes were validated against a theoretical prediction by Alonso et al. (2016). A fixed amount of H_2SO_4 was set at the first cross-section of the MARFORCE simulation and H_2SO_4 was further carried to the outlet of a cylinder only by convection and diffusion processes. Comparing the H_2SO_4 profiles at the outlet yields on average a 0.4 % difference between the MARFORCE model and the theoretical prediction by Alonso et al. (2016) (supplementary Figure A5). This suggests that the convection and diffusion processes in the MARFORCE model are well simulated.

360 This suggests that the convection and diffusion processes in the MARFORCE model are well simulated.

2.3.2 Gas-phase photo-chemistry photochemistry

The photolysis and chemical reactions in the H_2SO_4 , HO_2 and HOI calibrations can be simulated by a set of differential equations which describe the production and loss of various species. To make the MARFORCE model more versatile, the model was

365

designed to accommodate the input file format from the Master Chemical Mechanism (MCM) (Jenkin et al., 1997; Saunders et al., 2003), a near-explicit chemistry mechanism for numerous organic precursors. The scripts used to compile and interpret MCM mechanisms were adapted from O'Meara et al. (2021). The input file extracted from MCM is reshaped and the reaction equations, reaction rate coefficients, reactants, products, their indices, and stoichiometric numbers are generated accordingly. The temperature and pressure dependence of reaction rate coefficients are taken into consideration. Finally, differential equa-

- 370 tions for each species based on its production and loss processes are produced and solved. Additionally, the MARFORCE model leaves an option to set abundant species as constants, so their concentrations are assumed uniform and homogeneous in the flow reactor. These species include, for example, (O_2) , N_2 , SO_2 , I_2 , and H_2O in the H_2SO_4 and HOI calibration experiments. With its flexibility, the MARFORCE model can be readily adapted to simulate organic oxidation or any other experiments using a laminar flow reactor.
- 375

There are two default chemistry schemes provided in the MARFORCE model and they are used for the H₂SO₄ calibration and the HOI calibration, respectively. The reaction rate coefficients utilised in these two schemes are tabulated in supplementary Table A1. The most important procedure of these calibration experiments is to ealeulate obtain the OH concentration. The OH concentration is determined by the photon intensity produced by the calibration source (It-product, amount of photons per eross-section) and the absolute water content in the air passing through the calibration source. It-product refers to the product of UV light intensity at 185 nm and effective illumination time. In this study, we derived the It-product from the N₂O experiment, which was conducted under the same conditions as the H₂SO₄ calibration experiments. The details of the It product It-product determination can be found in Kürten et al. (2012). BrieffyIn brief, the chemical actinometry method utilising was employed, which involves the conversion of N₂O to NO_x (primarily NO)is carried out, to determine the light intensity. As the Since NO is less reactive compared with exhibits lower reactivity compared to OH and it can easily be measured with can be conveniently measured using commercial NO_x monitors, the "It-productcan therefore be derived ..." can be derived accordingly. Considering that the calibration experiment duration is relatively short (a few hours) compared to the potential lifetime of the mercury lamp, it is reasonable to assume that the attenuation of the It-product over time is negligible.

390 The OH initial concentration is further defined as

$$[OH] = \underline{It}\underline{It} \times \sigma_{H_2O} \times \Phi_{H_2O} \times [H_2O]$$
⁽²⁾

where σ_{H_2O} is the absorption cross-section of water vapour, 7.22×10^{-22} cm² (Creasey et al., 2000), and Φ_{H_2O} is the quantum yield (unity in this case).

2.3.3 Flow mixture 395

On top of the regular simulation of a In addition to its ability to simulate a standard cylindrical flow reactor with a uniform size, the MARFORCE model also has limited skills in possesses limited capabilities in two specific conditions: 1) simulating connected two two interconnected flow reactors with different sizes and varying sizes: the model is capable of simulating scenarios where two flow reactors of different sizes are connected. 2) simulating reactions when a dilution flow is merged with the sample flow through a Y-shape tee. Y-shaped tee. These additional features enhance the versatility of the MARFORCE

model, allowing for a more comprehensive analysis of complex flow reactor systems.

The first design aims to cope with the different sizes of the chemical ionisation inlet and the calibration source itself. For example, the MION2 inlet utilises a KF25 connector with an inner diameter of 24 mm while the calibrator calibration source 405 utilises a 3/4" tube with an inner diameter of 15.6 mm. Our model considers an instantaneous transition between the tubes of different sizes, i.e., the chemical distributions at the last grid-cross-section of the first cylinder are copied into the first grid cross-section of the second cylinder while the axial flow speed is adjusted to the cross-section of the second cylinder. As this simplification ignores the convective transport of species to the walls at the transition region, it likely gives the concentration upper limit at the pinhole of the mass spectrometer. Since the inner diameter difference between the calibration source and the MION2 inlet is relatively small in this study, we expect that the resulting uncertainty is well within the overall systematic

410

400

uncertainty of $-50/+100 \frac{\%\%}{2}$.

The second design considers that an additional dilution flow is utilised to reduce the sample water content when entering the Br⁻-MION2 inlet. Similarly, we assume an instantaneous transition at the spot where the dilution flow is added. In this

415 case, both the chemical distribution and axial flow speed are changed since a new branch of flow is added. The simulation is carried out with a two-process procedure: before and after the dilution. First, we carry out a standard simulation before adding the dilution flow. Once the flow is fully developed and the chemical distribution reaches a steady state in the simulation, the last cross-section at the grid right before adding the dilution flow is stored and recalculated into the first cross-section of the next simulation. The second simulation is further carried out after considering the dilution flow, together with the changes in 420 chemical distribution and axial flow speed.

It should be noted that the fluid dynamics processes are overly simplified in the second design and therefore, this option should be used with caution. In this study, this option is necessary only because investigating the detection humidity effect of e.g., H₂SO₄, HO₂ and HOI requires adding a dilution flow after the calibration source. In order to estimate the magnitude of 425 error caused by the simplification of fluid dynamics, we carried out experiments comparing calibration results obtained with the first design (straight tube) and the second design (Y piece) and the results are shown in supplementary Figure A6. We find that the second design additionally introduces a 12 % higher calibration factor in the H_2SO_4 calibration and a 27 % higher calibration factor in the HOI calibration, compared with the calibrations using the first design. Therefore, the application of the second design for the purpose of this study is reasonable and does not introduce an excess amount of uncertainties. This mainly concerns the H_2SO_4 , HO_2 and HOI calibration experiments.

2.4 Quantum chemical calculations

Cluster fragmentation enthalpies were calculated using quantum chemical methods. The initial conformational sampling was performed using the Spartan'18 program. The cluster geometry was then optimized optimised using density function theory (DFT) methods at the ωB97X-D/aug-cc-pVTZ-PP level of theory (Chai and Head-Gordon, 2008; Kendall et al., 1992). Iodine and bromine pseudopotential definitions were taken from the Environmental Molecular Sciences Laboratory (EMSL) basis set library (Feller, 1996; Peterson et al., 2003). Calculations were carried out using the Gaussian 16 program (Frisch et al., 2016). An additional coupled-cluster single-point energy correction at the DLPNO-CCSD(T)/def2-QZVPP (Riplinger and Neese, 2013; Riplinger et al., 2013; Weigend and Ahlrichs, 2005) level of theory was carried out on the lowest energy conformers to refine the DFT calculated enthalpies. The coupled-cluster calculation was performed using the ORCA program version 4.2.1 (Neese, 2012).

The master equation solver for multi-energy well reactions (MESMER) program was used to investigate the ionisation chemistry of $I_2O_3 \cdot HNO_3NO_3^-$. For the $I_2O_3 \cdot HNO_3NO_3^-$ complex, Lennard-Jones potentials of $\sigma = 6.5$ Å and $\epsilon = 300$ K were used, which are identical to those used previously for similar iodine systems (Gálvez et al., 2013). The MesmerILT method was used with a pre-exponential value of 1.26×10^{-9} cm³ molec⁻¹ s⁻¹, which is equal to the $I_2O_3 + HNO_3NO_3^-$ collision rate calculated using the average dipole orientation (ADO) method. The method is described in detail by He et al. (2021a). Varying the collision rate by a factor of 3 has no effect on the MESMER results, indicating that the reported final fragmentation rate coefficients of $I_2O_3 \cdot HNO_3NO_3^-$ are not sensitive to the accuracy of the computed collision rate.

450

3 Results and Discussion

3.1 Calibration of H₂SO₄, HOI and HO₂ using MARFORCE

Gaseous H₂SO₄ concentration is regularly measured around the globe using the nitrate chemical ionisation method. In this study, a direct H₂SO₄ calibration has been carried out for the MION2 inlet at tower 1 using either Br⁻ (Br⁻-MION2-T1, Figure A3) or NO₃⁻ (NO₃⁻-MION2-T1) chemical ionisation methods, and additionally at tower 2 with Br⁻ (Br⁻-MION2-T2, Figure A7) chemical ionisation method. The MARFORCE model is utilised to simulate the evolution of various species at the cross-section of the inlet tube as shown in Figure 2. The predicted H₂SO₄ concentrations are further compared with the measured normalised ratios to derive ealibrator calibration factors (Table 1).

Table 1. Calibration coefficients and detection limits for selected species measured by the MION2 inlet and Eisele-type inlet. It should be noted that these the reported numbers are specific to the experimental conditions and instrument tuning in our experiments. Different instrument tuning can also result in different calibration coefficients and detection limits. Undesired impurities may result in elevated detection limits despite the calibration coefficients being the same for the analytes. Therefore, these numbers should not be applied to another study without carrying out the same calibration experiments described in this study.

Species	Calibration coefficients (MION2)			Detection limit		
			MION2 (Br ⁻)		Eisele inlet	
	Tower 1 (Ionisation Time = 35 ms)		Tower 2	Tower 1	Tower 2 (160 ms)	
			(300 ms)			
	NO_3^-	Br ⁻	Br ⁻	APi2(RH < 0.1%) APi1(RH (RH = 3.7%)	APi1(RH < 0.1%)	APi1
H ₂ SO ₄	$1.3 imes 10^{10}$	8.1×10^9	9.8×10^8	$a \frac{8.5 \times 10^4 e}{10^3} 1 \times 10^5$	$\overset{\underline{if}}{\sim} 2.9 \times 10^4$	li
		(RH = 0.2-23.3%)	(RH = 0.3-11.6%)			-
HOI	n/a	1.8×10^{10}	5.1×10^9	$\frac{a}{2.2 \times 10^5} \frac{fb}{5.9 \times 10^5} 5.9 \times 10^5$	$\overset{\underline{ig}}{\sim} 1.6 \times 10^5$	
		(RH = 3-17%)	(RH = 3-17%)			
HIO ₃	n/a	n/a	n/a	$a \frac{8.0 \times 10^4 e}{1.3 \times 10^5}$	n/a	<u>l</u>
HO ₂	n/a	2.8×10^{11}	1.2×10^{11}	$\frac{\frac{b_{5.2} \times 10^5}{2}}{(\text{RH} = 2.7\%)} \frac{\frac{gc}{2} 3.3 \times 10^6}{(\text{RH} = 0.3\%)} \frac{\frac{kh}{2} 5.7 \times 10^5}{(\text{RH} = 0.3\%)}$	$\frac{\underline{k}h}{\sim} 5.7 \times 10^5$	
		(RH = 2.5%)	(RH = 2.7%)		(RH = 0.3%)	
SO ₂	n/a	2.6×10^{16}	2.1×10^{16}	$^{\circ}0.4 \times 10^{7} \stackrel{hd}{\sim} 1.8 \times 10^{9}$	n/a	
		(RH = 10%)	(RH = 9.9%)	(RH = 0.5%)		
I_2	n/a	8.2×10^9	n/a	$\frac{d_{6.7} \times 10^5}{2} \frac{d_e}{2} 3.3 \times 10^5$	n/a	
		(RH = 26-37%)				
ΙΟ	n/a	n/a	n/a	$a \frac{3 \times 10^{4e}}{1.6 \times 10^5}$	$\underline{\overset{i}{f}}_{\sim}^{f} 2.5 \times 10^{4}$	
OIO	n/a	n/a	n/a	$a3.4 \times 10^{5} 2.0 \times 10^{5}$	$\stackrel{\underline{^e2}\times 10^{5i}f}{\sim} 3.1\times 10^4$	
I_2O_2	n/a	n/a	n/a	$a \frac{7.9 \times 10^{4e}}{1.9 \times 10^5}$	$\underline{\overset{i}{}}_{\sim}^{f} 3.5 \times 10^{4}$	
I_2O_3	n/a	n/a	n/a	$a4.9 \times 10^{4e} 1.9 \times 10^{5}$	$\underline{\overset{i}{}}_{\sim}^{f}4.2 \times 10^{4}$	
I_2O_4	n/a	n/a	n/a	$a \frac{5.1 \times 10^{4e}}{1.9 \times 10^5}$	$\underline{\overset{i}{}}_{\sim}^{f} 3.0 \times 10^{4}$	
I_2O_5	n/a	n/a	n/a	$a\frac{2.5 \times 10^5 2 \times 10^5}{2}$	$\frac{e_{2.0} \times 10^{5i} f}{\sim} 3.7 \times 10^4$	

Unit: molec. cm⁻³ cps cps⁻¹; "n/a" refer to "not available"; Experiments were conducted at room temperature. H_2SO_4 calibration factor is applied to estimate the detection limits for the species were conducted at room temperature. H_2SO_4 calibration factor is applied to estimate the detection limits for the species were conducted at the kinetic limit, their LODs are mere estimations. The detection limits are estimated with 1-min data and one-hour data collection time. The RH report at 25 °C. The calibration factors and LODs have a systematic error of a factor of two (-50 %/+100 %).

Calibration factors used for the LOD calculation are: a. 8.1×10^9 , b. 1.8×10^{10} , c. 2.8×10^{11} , d. 1.03×10^{14} , e. 8.2×10^9 , f. 9.8×10^8 , g. 5.1×10^9 , h. 4.1×10^{10} , i. 3.5×10^9 .



Figure 2. MARFORCE simulation results of the a H_2SO_4 calibration experiment. The x-axis shows the distance from the UVP pen-ray lamp to the instrument pinhole entrance of the chemical ionisation chamber and the y-axis shows the distance in the radial direction. Conditions for the simulation: R = 0.78 cm, L = 26 cm, sample flow = 10.6 slpm, $[SO_2] = 5.78 \times 10^{13}$ cm⁻³, $[O_2] = 2.42 \times 10^{16}$ cm⁻³ and $[H_2O] = 2.8 \times 10^{16}$ cm⁻³.

460 The derived calibration factor of for Br⁻-MION2-T1 (8.1×10^9) is roughly approximately eight times higher than that of Br⁻-MION2-T2 (9.8×10^8). This observation is consistent with the fact that the ionisation time from tower 2 to the pinhole (ea. around 300 ms) is approximately roughly 8.6 times longer than that of tower 1 (35 ms). Longer ionisation time results in a larger fraction A longer ionisation time leads to a greater conversion of Br⁻ and H₂O·Br⁻ being converted to into H₂SO₄·Br⁻ or HSO₄⁻, thus resulting in a lower calibration factor. As for the

465

In the case of NO_3^- -MION2-T1, it achieves exhibits a similar sensitivity as the to H_2SO_4 detection as Br^- -MION2-T1for detection. This . This similarity is likely due to the fact that the ionisation time was kept constant (consistent ionisation time (using tower 1was used) and both) for both methods, since both the NO_3^- and Br^- chemical ionisation methods measure H_2SO_4 at the collision limit, as mentioned in previous studies (Kürten et al., 2012; Wang et al., 2021a).

470



Figure 3. The modelled or calculated vapour concentrations vs. the normalised signals for a) H_2SO_4 , b) HOI, c) SO_2 , d) H_2O . The dasheddotted, solid, and dashed lines are the linear fits of the results from different inlet modes: 1) tower 1 with the NO_3^- chemical ionisation method, 2) tower 1 with the Br⁻ chemical ionisation method, and tower 2 with the Br⁻ chemical ionisation method. The slopes of the fitted lines represent the calibration coefficients, shown in the legend. The colour bar shows the relative humidity in the calibration experiments.

Sanchez et al. (2016) has reported that the bromide chemical ionisation mass spectrometer (Br⁻-CIMS) is capable of detecting HO₂ radicals at ambient relevant concentrations. In this study, we calibrated HO₂ together with H₂SO₄, as HO₂ is a by-product in the chemical production of H₂SO₄ (see supplementary Table A1). As the binding of HO₂ with Br⁻ is significantly weaker than that of H₂SO₄ with Br⁻, the collision induced fragmentation of HO₂ · Br⁻ in the ion-optics of the mass spectrometer is larger (Passananti et al., 2019). Additionally, as the humidity effect of HO₂ will be shown to be strong in a section belowsection 3.3, the calibration coefficient of HO₂ has to be derived with respect to a specific humidity level. The derived HO₂ calibration factors at 2.5 - 2.7 % RH (25 °C) are 2.8×10^{11} and 1.2×10^{11} , respectively, for Br⁻-MION2-T1 and

475

- 480 The HOI calibration was also carried out using the H₂SO₄ calibration source, except that the SO₂ source was replaced with an I₂ source. The calibration procedure was described in greater detail in our earlier studies (Tham et al., 2021; Wang et al., 2021a) . While the calibration experiment remains the same, the MARFORCE model in this study utilises a complete set of iodine chemistry thus producing a more accurate calibration factor for HOI (supplementary Table A1). We estimate that the HOI concentration using the MARFORCE model is only 0.1 0.7 % lower compared to our earlier model which only considers the
- 485 essential steps of the HOI formation. As can be seen in Figure 3 and Table 1, the calibration factor for HOI is roughly two times that of H_2SO_4 . This suggests that HOI is detected at close to the collision limit. It is worth noting that we find instrument setting affects HOI detection significantly since HOI is not strongly bonded to Br^- . The preferred fragmentation pathway is $HOI \cdot Br^- \longrightarrow HOI + Br^-$ (Table 2), and thus a fraction of $HOI \cdot Br^-$ lose the signature of dissociates into HOI and Br^- after passing the ion optics of our the mass spectrometer. A more fragmentation-oriented setting can result in a higher fraction of
- 490 HOI·Br⁻ getting lost in the ion optics, thus resulting in a higher calibration factor, i.e., lower sensitivity. As an example, in our earlier studies (Tham et al., 2021; Wang et al., 2021a), we used a relatively fragmenting setting was used in our earlier studies compared to the one used in this study in an attempt to reduce $(H_2O)_n \cdot Br^-$ clusters and other water-associated clusterswhich complicated the mass spectra (Tham et al., 2021; Wang et al., 2021a). Such a setting resulted in an eight times higher. This experimental setup led to a calibration factor for HOI than that for and it cannot be explained by iodine chemistry schemes used
- 495 in the calibration scripts nor by any differences in the experimental conditions, that was eight times higher than the calibration factor for H_2SO_4 .

3.2 Calibration of H₂O and SO₂

- $H_2O \cdot Br^-$ is a regular peak and one of the primary ions measured by the Br^- -CIMS-MION2. Br^- -CIMS-MION2 is, therefore, able to measure absolute water content if the $H_2O \cdot Br^-$ signal is calibrated against a dew point mirror instrument. Such a calibration has at least two purposes: 1) the calibrated $H_2O \cdot Br^-$: Br^- can be used as an indicator of the fragmentation level of the Br^- -CIMS-MION2 and 2) -CIMS can be more sensitive to compared with compared to regular relative humidity sensors and dew point mirrors, Br^- -MION2 exhibits higher sensitivity towards H_2O . In this study, the calibration of H_2O was calibrated both for the performed for both Br^- -MION-T1 and Br^- -MION-T2as shown, as illustrated in Figure 3. The Interestingly,
- 505 the calibration factors for both of the towers do not differ from each other likely due to the towers did not show significant differences. This can be attributed to the presence of an excess amount of H_2O which establishes an , which establishes a rapid equilibrium with Br^- and $H_2O \cdot Br^-$ rapidly regardless irrespective of the ionisation time.

As a reasonable binding enthalpy of $SO_2 \cdot Br^-$ was predicted using quantum chemical calculations (Table 2), we continued to 510 check whether the Br⁻-MION2 allows us to detect SO₂. A variable amount of SO₂ was mixed with a fixed amount of dilution flow at a constant relative humidity (RH, 10 %) which was measured by the Br⁻-MION2. Clear SO₂ \cdot Br⁻ was measured and it

Table 2. Fragmentation enthalpies (the opposite of binding enthalpies) of analytes with the Br⁻. The cluster geometry was optimised at the ω B97X-D/aug-cc-pVTZ-PP level of theory at 298.15 K (Chai and Head-Gordon, 2008; Kendall et al., 1992). The enthalpies were calculated at the DLPNO-CCSD(T)/def2-QZVPP at 298.15 K

Cluster fragmentation pathway	Fragmentation enthalpies (kcal mol^{-1})
$I_2 \cdot Br^- \longrightarrow I_2 + Br^-$	33.3
$I_2 \cdot H_2 OBr^- \longrightarrow I_2 \cdot Br^- + H_2 O$	8.0
$IO \cdot Br^- \longrightarrow IO + Br^-$	24.5
$IO \cdot H_2OBr^- \longrightarrow IO + H_2O \cdot Br^-$	21.3
$IO \cdot H_2OBr^- \longrightarrow IO \cdot Br^- + H_2O$	9.9
$OIO \cdot Br^- \longrightarrow OIO + Br^-$	23.2
$OIO \cdot H_2OBr^- \longrightarrow OIO + H_2O \cdot Br^-$	22.1
$OIO \cdot H_2OBr^- \longrightarrow OIO \cdot Br^- + H_2O$	11.9
$I_2O_3 \cdot Br^- \longrightarrow IO_3^- + IBr$	24.6
$I_2O_4 \cdot Br^- \longrightarrow I_2O_4 + Br^-$	42.6
$I_2O_4 \cdot H_2OBr^- \longrightarrow I_2O_4 + H_2O \cdot Br^-$	48.8
$I_2O_4 \cdot H_2OBr^- \longrightarrow I_2O_4 \cdot Br^- + H_2O$	10.5
$HIO_3 \cdot Br^- \longrightarrow IO_3^- + HBr$	^a 29.9
$HIO_3 \cdot Br^- \longrightarrow HIO_3 + Br^-$	^a 35.7
$HIO_3 \cdot H_2OBr^- \longrightarrow HIO_3 + H_2O \cdot Br^-$	33.1
$HIO_3 \cdot H_2OBr^- \longrightarrow HIO_3 \cdot Br^- + H_2O$	11.2
$HIO_3 \cdot H_2OBr^- \longrightarrow IO_3 \cdot H_2O^- + HBr$	26.7
$\mathrm{HIO}_2 \cdot \mathrm{Br}^- \longrightarrow \mathrm{HIO}_2 + \mathrm{Br}^-$	^b 29.2
$HIO_2 \cdot H_2OBr^- \longrightarrow HIO_2 + H_2O \cdot Br^-$	15.5
$HIO_2 \cdot H_2OBr^- \longrightarrow HIO_2 \cdot Br^- + H_2O$	1.3
$HIO_2 \cdot H_2OBr^- \longrightarrow IO_2 \cdot H_2O^- + HBr$	27.4
$HOI \cdot Br^{-} \longrightarrow HOI + Br^{-}$	^b 26.9
$HOI \cdot H_2OBr^- \longrightarrow HOI + H_2O \cdot Br^-$	22.9
$HOI \cdot H_2OBr^- \longrightarrow HOI \cdot Br^- + H_2O$	9.6
$HOI \cdot H_2OBr^- \longrightarrow IO \cdot H_2O^- + HBr$	48.4
${\rm H_2O} \cdot {\rm Br^-} \longrightarrow {\rm H_2O} + {\rm Br^-}$	13.2
$\mathrm{HO}_2 \cdot \mathrm{Br}^- \longrightarrow \mathrm{HO}_2 + \mathrm{Br}^-$	23.1
$\mathrm{H_2SO_4}\!\cdot\!\mathrm{Br^-}\longrightarrow\mathrm{HSO_4^-}+\mathrm{HBr}$	^b 27.9
$H_2SO_4 \cdot H_2OBr^- \longrightarrow H_2SO_4 + H_2O \cdot Br^-$	36.1
$H_2SO_4 \cdot H_2OBr^- \longrightarrow H_2SO_4 \cdot Br^- + H_2O$	8.2
$\mathrm{H}_2\mathrm{SO}_4 \!\cdot\! \mathrm{H}_2\mathrm{OBr}^- \longrightarrow \mathrm{HSO}_4 \!\cdot\! \mathrm{H}_2\mathrm{O}^- + \mathrm{HBr}$	22.0
$SO_2 \cdot Br^- \longrightarrow SO_2 + Br^-$	19.4

^{*a*} the fragmentation enthalpy is updated from Wang et al. (2021a) as a lower energy HIO₃ \cdot Br⁻ cluster geometry, which has an additional Br-I interaction, has been located in this study (see Figure A8). ^{*b*} Value adopted from Wang et al. (2021a).

increased linearly with the SO₂ concentration in the sample flow (Figure 3). However, the calibration factor of SO₂ is roughly six orders of magnitude higher than that of H₂SO₄ at 10 % relative humidity (RH)RH. This is consistent with the weaker binding of SO₂ · Br⁻ compared with H₂SO₄ · Br⁻. Additionally, SO₂ calibration is extremely sensitive to RH changes as can be seen in Figure 4. The best detection limit achieved in this study is In this study, the best achieved detection limit was 9.4×10^7 cm⁻³ at an RH below 0.1 %RH and theoretically. Theoretically, it is feasible to further increase the sensitivity by reducing possible to enhance the sensitivity even further by reducing the absolute water content.

515



Figure 4. The effect of humidity on the detection efficiency of a) HOI, b) H_2SO_4 , c) HO_2 , d) SO_2 , and e) I_2 . The measured signals in each set of experiments are normalised by the signal at the lowest RH. Therefore, the normalised signals represent how the increasing RH is affecting the detection limit compared with the initial point. The red and green purple circles and orange triangles show the detection humidity effects of tower 1 and tower 2, respectively. The yellow-blue squares refer to the experiments conducted with a dry flow added before the MION2 inlet. The RH is converted from absolute H_2O concentrations at 25 °C. Error bars represent one standard deviation.

3.3 Detection humidity effect

- 520 The measurement sensitivity of halide ion-anion based chemical ionisation method was regularly reported to be affected by atmospheric water content (Kercher et al., 2009; Mielke et al., 2011; Woodward-Massey et al., 2014; Lee et al., 2014). The humidity effect of atmospheric pressure Br⁻-MION2 could be even stronger owing amplified due to the higher water content in the present in air samples. Although our earlier study Wang et al. (2021a) has demonstrated that the detection of I₂ by Br⁻-MION1 was not affected within a limited humidity variation (40 - 80 % RH at -10 °C), characterisation under a wider range
- 525 of humidity conditions is needed. As the detection humidity effect in this study exclusively refers to the effect of absolute humidity instead of relative humidity, absolute humidity parameters such as dew/frost point or H_2O molecule concentration are commonly presented together with the relative humidity (at 25 °C, otherwise notified).

In this study, we examine the detection humidity effect of H₂SO₄, HOI, HO₂, SO₂ and I₂ with RH from below 1 % to 60 %

at 25 °C. The detection humidity effect of Unlike H₂SO₄, HO₂ and HOI, which require generation from a calibration source, both SO₂ and I₂ is relatively easier to characterise as the calibration source is not involvedhave their own standardised sources. This simplifies their control during the characterisation of the detection humidity effect. Therefore, a straight flow reactor is used to premix the analyte containing air sample to the Br⁻-MION2 (Figure A2). It is worth noting that we do not account for the wall loss of SO₂ and I₂ in the analysis. The wall loss of SO₂ is negligible at the time scale of the calibration processes (few seconds). Despite I₂ being a sticky gas which both condenses to and evaporate from the walls of the flow reactor, an equilibrium can be reached if given sufficient time. When equilibrium is reached, which could take achieved given a sufficient amount of time. In our experiments, it took up to 24 hours in our experiments reach equilibrium. Once equilibrium is established, the condensation and evaporation of I₂ cancel-balance each other outand thus, making the estimation of I₂ concentration is straightforward.

540

On the other hand, the characterisation of the detection humidity effect of H_2SO_4 , HOI and HO₂ is more challenging as the production of these species is nearly proportional to the amount of H_2O passing the calibration source. Therefore, an experimental apparatus was built which enabled humidifying the air sample after the <u>ealibratorcalibration source</u>, thus without disturbing HO_x production processes in the calibration source (Figure A3).

545

550

The results of the humidity characterisation are shown in Figure 4. Despite Although only five species were characterised and observed for their distinct humidity sensitivity is observed for the mentioned five species, a general conclusion can be drawn for that applies to essentially all of the species- an excess: an excessive amount of water content results in lower leads to a decrease in detection sensitivity. The species with stronger binding with Br⁻ exhibits higher tolerance to humidity changes less sensitivity to changes in humidity (e.g, H_2SO_4 , and and I_2), while the weakly bonded ones (HOI, SO₂ and HO₂) are strongly

affected. The humidity tolerance of the measured species can be ordered as $I_2 > HOI > HO_2 > SO_2$ which is the same order as

Interestingly, the detection humidity effect of H₂SO₄ is observed to be non-linear, i.e., the detection sensitivity of H₂SO₄ first increases with higher RH but eventually has a sharp drop at around 40 % RH(25 °C). The enhancement of H₂SO₄ detection at below ca. 33 % RH could be contributed by two mechanisms. First, the diffusivity of H₂SO₄ is lower at higher RH (Hanson and Eisele, 2000). A higher RH, therefore, reduces the wall deposition of H₂SO₄ in the inlet tube, thus effectively increasing the detected H₂SO₄. This is a universal systematic error which affects any factor that influences all H₂SO₄ detection technique techniques with appreciable sampling line residence time. The second possibility is that at low RH regime, H₂O does enhance H₂SO₄ detection by offering more modes through which the excess energy of the cluster can dissipate in the formation of H₂SO₄ ·Br⁻, thus resulting in a relatively more stable cluster (Iyer et al., 2017). Regardless of the sources of the detection humidity effect at the low water content regime, the maximum systematic error is measured to be 37 % by comparing the experiment carried out at 2 % RH (frost point of -25 °C) and the experiment carried out at 33 % (dew point of 7.6 °C) in Figure 4b. As the humidity change in ambient conditions is commonly smaller than during experiments, we expect Based on our findings.

565 we anticipate that the detection humidity effect of H_2SO_4 is would be moderate when the dew point is below ca. approximately 7.6 °C. However, it is important to exercise caution when conducting measurements under higher absolute humidity conditions.

Additionally, a longer reaction time in the ion-molecule reaction chamber-ionisation time by utilising the Br^- -MION2-T2 results in a stronger detection humidity effect as shown in Figure 4. Although such an This phenomenon is the most significant

570 for HOI, i.e., the detection of HOI is more humidity dependent using $Br^--MION2-T2$ than $Br^--MION2-T1$. Although this effect is difficult to quantify, it practically suggests that the Br^- chemical ionisation method should employ a shorter ionisation time (i.e., using the tower 1) when operating MION2 with multiple chemical ionisation methods.

In summary, we find that the detection of $Br^--MION2$ is strongly affected by air water content. The atmospheric pressure 575 Br^- chemical ionisation method is suitable for laboratory experiments where water content is controlled and atmospheric observations in the cryosphere where air water content is low. Nevertheless, the humidity effect should be considered individually for different analytes and the binding enthalpy between the analyte and Br^- is likely a good indicator. As the NO₃⁻-MION2 (or the NO₃⁻ chemical ionisation , in general, in general) is known to have minimal detection humidity sensitivity, it is commonly operated together with the Br^- -MION2. Cross-check Performing a cross-check of mutually measured species (e.g., such as

580 H_2SO_4 , HIO₃, and oxidised organic species) will give essential information on, will provide crucial insights into whether and when the detection capability of Br⁻-MION2 detection is compromised by air water content the water content in the air. In this regardcontext, the new design of Br⁻-MION2allowing as many as , which enables three chemical ionisation methods to have the same ionisation time, is essential.

585 **3.4** Attempts to reduce the detection humidity effect

Several ways Various approaches were explored to reduce mitigate the detection humidity effect. The first and usual way of reducing the detection humidity effect is deploying One commonly used method is to employ a low-pressure chemical ionisation systemwhich was regularly used for, e.g., iodide chemical ionisation systems (Lee et al., 2014) or bromide chemical ionisation system, which has been successfully implemented in iodide chemical ionization systems (Lee et al., 2014)

- 590 and bromide chemical ionization systems (Wang et al., 2021a). However, the reduction of air sample RH is at the cost of reducing measurement sensitivity reducing the relative humidity (RH) of the air sample comes at the expense of reducing the measurement sensitivity for species detected at the collision limit, such as H_2SO_4 , as HIO₃ and I₂, as the air sample is unavoidably diluted unavoidably undergoes dilution in this process. We estimated previously that the Br⁻-FIGAERO inlet had more than 10 times higher detection limit compared to the Br⁻-MION1 inlet (Wang et al., 2021a). For example, the
- 595 Br^- -FIGAERO had an HIO₃ detection limit of 5.1×10^6 cm⁻³ which struggles to detect atmospheric level levels of HIO₃ (commonly below 10^7 cm⁻³) (He et al., 2021b). The lower level of detection limit provided by the Br⁻-MION2 inlet is therefore essential in the detection of iodine species. Another important factor is the reaction of halogen radicals with analytes. Besides halogen anions, halogen radicals can also be produced by chemical ionisation processes. While iodine radical (I·) mostly reacts with halogen species and a very limited minimal number of organic species, bromide radical (Br·) reacts with a
- 600 wider range of organic species as it has a larger reactivity. Regular Conventional low-pressure systems which mix analyte that involve mixing analytes with reagent gases(e.g., FIGAERO inlet) may introduce an additional complexity when interpreting measured, such as the FIGAERO inlet, can introduce additional complexities when interpreting mass spectra. Therefore, we had to seek alternatives to help As a result, alternative approaches were pursued to effectively reduce the detection humidity effect.

605

620

The first method is the dilution method. Instead of measuring the air sample directly, a dry dilution flow was mixed with the air sample at the entrance of the Br⁻-MION2 inlet (see supplementary Figure A4). We tested this method for the SO₂ detection with an air sample flow of 1.8 slpm and a dilution flow of 20.7 slpm (Figure 4). The x-axis for this set of experiments represents humidity in the air sample instead of the humidity after the dilution to compare with the experiments without adding the dilution flow. We observe a significantly reduced detection humidity effect compared to the case without dilution. It is noteworthy that as the air sample is was diluted by a factor of 21.512.5, the detection limit of the instrument is likely enhanced by the same factor. However, since the detection humidity effect for SO₂ is significantly higher than other species (e.g., H₂SO₄, HOI and I₂), the dilution is still effective for SO₂ measurement. For example, no SO₂ · Br⁻ signal would not be measured at 40 % RH (25 °C) if the air sample is not diluted but a noticeable signal would be measured if the air sample is diluted. A similar conclusion is likely applicable to other species but with a different optimal humidity cut-off.

The second method is additionally introducing a core-sampling device that uses the air sample as the core flow and a dry synthetic air flow as the sheath flow (supplementary Figure A9). This takes the advantage of the fact that H_2O diffuses into the sheath flow faster than other analytes with larger molecular weight, thus effectively reducing the RH in the core flow from which the instrument pinhole collects the most sample. On the other hand, Nevertheless, it is important to note that the core-

sampling methodalso inevitably reduces, while helping to mitigate the detection humidity effect, also leads to a reduction in the $SO_2 \cdot Br^-$ signalas. This is because the SO_2 also gets diluted which partially counters the itself gets diluted, partially counteracting the benefits of the reduced detection humidity effect.

- Various sample-to-sheath flow combinations were tested as presented in Figure 5. To compare the detection coefficient to standard conditions, the The measured SO₂·Br⁻ signal from all sets of experiments is normalized was normalised by the experiment with the sample-to-sheath ratio of 21:1 at 0.21 % RH (25 °C). The results reveal indicate that reducing the sample-to-sheath ratio effectively eases-alleviates the SO₂ detection humidity effect. The results show It is observed that different mixing ratios only moderately affect have only a moderate impact on the measured SO₂·Br⁻ when the H₂O is smaller than concentration is below 10¹⁶ cm⁻³ (1 % RH; the-), indicating a low detection humidity effect remains low). This is likely due
- to the fact that the instrument pinhole primarily measures the air sample in the core flow, as it only sucks 0.8 slpmin such conditions. However, the core-sampling device clearly enhances the SO₂ detection efficiency when the H₂O concentration is larger than 10^{16} cm⁻³. The sample-to-sheath ratio of 1:21 enables effective detection of SO₂ at around 4.5×10^{17} cm⁻³ (60 % RH) of H₂O while the sample-to-sheath ratio of 21:1 is not able to detect SO₂ after around 4.3×10^{16} cm⁻³ of H₂O (6 % RH).
- 635 Overall, the sample-to-sheath ratio of 21:1:21 is at least two orders of magnitude more effective in detecting SO₂ when H₂O is greater than 10¹⁶ 2×10¹⁶ cm⁻³. Therefore, the core-sampling method is an effective method for reducing the detection humidity effect of species which are weakly bonded with Br⁻. However, since the detection limit is nevertheless changed by Despite the reduced detection humidity effect, it is important to note that the sample water content still impacts the detection limit of SO₂. Therefore, dedicated experiments have to be carried out to derive need to be conducted to accurately determine
 640 the concentration of the analyte (e.g., SO₂).



Figure 5. Reducing the detection humidity effect with the core sampling method (supplementary Figure A9). This design takes the advantage of the faster diffusion of H_2O than SO_2 from the sample flow to the sheath flow and effectively reduces the RH in the sample flow. Various sample-to-sheath ratios were tested at different H_2O concentrations to find the optimal setting. All the data are normalised to the lowest RH data point in the sample-to-sheath = 21:1 experiment. Due to experimental constraints, the sample-to-sheath ratios = a) 3.5:18.5, b) 2:20 and c) 1:21 experiments started only from the second, third and fourth lowest RH points, respectively. All other experiments collected data in all humidity conditions. The error bars represent the standard deviation of the normalised SO₂ signals.

3.5 Limit of detection

650

The limit of detection (LOD) is an essential parameter for a chemical ionisation inlet system. For the convenience of intercomparison, we define the LOD in this study as:

$$645 \quad \text{LOD} = \mu + 3 \times \sigma \tag{3}$$

where μ is the mean value of one-hour mass spectrometric data with a one-minute time resolution and σ is the standard variation of the same data. Both μ and σ include the experimentally derived calibration coefficient. The species without direct calibration utilise the calibration coefficient of H₂SO₄, thus the LODs for these species generally represent the lower limit. The LODs are measured by passing LOD is determined by introducing pure nitrogen or synthetic air to into the chemical ionisation inletin which case, where none of the species listed in Table 1 is expected. It should be noted are expected to be present. It is important to emphasise that this LOD definition is suitable for disentangling specifically suitable for distinguishing trace gas concentrations from background levels in long-term observations. Values above the LOD can commonly be distinguished from the time series. If one does a careful analysis of the measured mass spectra, a lower value may be recognised.

The reported LODs can be affected by many factors. Some of these factors are 1) the purity of the reagent source (e.g., nitrie acid solutionand dibromomethane solutionHNO₃ or CH_2Br_2 solution), 2) the purity of the sample air used at the LOD determination experiment, 3) the signal-to-noise (electronic background noise) ratio of the instrument, 4) the softness of fragmentation level (controlled by the tuning of the instrument) of the mass spectrometer, and 5) the humidity of the sample air used at the LOD determination experiment (for Br⁻ chemical ionisation)method) and 6) different ways of estimating LODs.

660

Due to the complex nature of the LOD determination, the MION2 inlet was coupled with two independent mass spectrometers (APi1 and APi2, respectively, see Table 1) to test its robustness. The LOD determination experiments were carried out with APi1 and APi2 in two independent laboratory environments with independent reagent sources and sample air. These two instruments were also individually tuned, thus having different signal-to-noise ratios and fragmentation levels. The results of the LOD determination experiments are tabulated in Table 1. Both of the -MION2-T1-APi1 and -MION2-T1-APi2 achieved LODsat the level of 10⁵ cm⁻³ for species that are detected at the collision limit (e. g., , and). In general, the -MION2-T1-APi2 has a slightly lower LOD compared with -MION2-T1-APi1. This could result from the fact that the APi1 has not been serviced for more than four years by the point of the experiments and the multi-channel plate could have degraded.-

- It is worth noting that it may appear that the and LODs of MION2-T1 (1×10^5 cm⁻³ and 3.3×10^5 cm⁻³) are similar to that reported for MION1-T1 (2×10^5 cm⁻³ and 3.8×10^5 cm⁻³) in our earlier study (Wang et al., 2021a). This is because the mass spectrometer used in Wang et al. (2021a) (noted as APi3) had a higher signal-to-noise ratio compared with the APi1. Therefore, getting similar LODs from the MION2-T1-APi1 and MION1-T1-APi3 already suggests that the MION2 inlet has improved its performance. In order to avoid this systematic error, Therefore, comparing the LODs derived in this study with earlier studies may not be meaningful. Hence, we additionally compared the H₂SO₄ LOD of the MION2 inlet with that of
- 675 the the widely-used Eisele-type inlet, both attached to the APil-same mass spectrometer (Table 1). The direct comparison suggests that the MION2-T1-Br⁻-MION2-T1 LOD is roughly 30 % higher than the LOD of the widely-used-Eisele inlet, thus a comparable performance. When we increased the ionisation time from 35 ms (MION2-T1Br⁻-MION2-T1) to 300 ms (MION2-T2Br⁻-MION2-T2), the LOD of MION2-Br⁻-MION2 for H₂SO₄ is further reduced by a factor of three, thus MION2-T2-Br⁻-MION2-T2 performs better than the Eisele inlet. This suggests that the MION2 inlet achieves comparable
- 680 (MION2-T1can achieve comparable (Br⁻-MION2-T1) or even better (MION2-T2) LODs Br⁻-MION2-T2) LOD than the Eisele inlet. Additional tuning of the ionisation time may further increase the advantage for chemical ionisation methods that are less affected by air water content (e.g., -CIMS).

Additionally, the Eisele-type inlet was regularly shown to have $\frac{\text{LODs-a LOD}}{\text{LODs-LOD}}$ as low as 10^4 cm^{-3} (Jokinen et al., 2012), a well-performing mass spectrometer will likely may further reduce the $\frac{\text{LODs-LOD}}{\text{LODs-LOD}}$ of MION2. Nevertheless, the achieved

685 LODs are low enough attained levels of LOD are sufficiently low for atmospheric measurementsas the discussed molecules commonly need to be. The molecules in question typically require concentrations above 10⁶ cm⁻³ to have a significant impact exert a significant influence on atmospheric chemistry and aerosol formation.

3.6 Voltage scanning and cluster formation enthalpy

- 690 Collision induced cluster fragmentation is an unavoidable issue which affects the detection of analytes that are weakly bonded with the reagent ion. Since if a charged cluster is loosely bonded, collisions between charged clusters and air molecules in the atmospheric pressure interface may break a large portion of the charged clusters apart prior to reaching the detector (Passananti et al., 2019). Therefore, charged cluster binding strength is an important factor determining whether an analyte-charger analyte-reagent ion cluster can be measured by the mass spectrometer (Iyer et al., 2016; Lopez-Hilfiker et al., 2016; Wang et al., 2021a). Lopez-Hilfiker et al. (2016) has shown that the level of collision induced cluster fragmentation is associated with the voltage differences between the first and second quadrupoles in the atmospheric pressure interface of the mass spectrometer.
- The voltage difference was shown to be indicative of the fragmentation level of the CIMS and it positively correlates with the cluster formation enthalpy (Iyer et al., 2016).
- In this study, we carried out voltage scan experiments with the same procedures as described in Lopez-Hilfiker et al. (2016). Briefly, we kept the voltage differences inside two individual quadruples constant while changing the voltage difference between these two quadruples to modulate energies in the collision processes and the results are shown in Figure 6. Generally, a higher voltage difference indicates a higher fragmentation level which in turn results in a lower remaining fraction of charged clusters. Charged clusters that are less sensitive to voltage changes, especially in the low voltage difference regime (e.g., ΔV
 40 V), are more stable.

A series of iodine oxides and oxoacids is evaluated together with other inorganic species such as H₂O, HO₂, SO₂ and H₂SO₄ (Figure 6). Based on the results, we categorise the analytes into three categories: 1) analytes which are strongly bonded with Br⁻, 2) analytes which are moderately bonded with Br⁻ and 3) analytes which are weakly bonded with Br⁻. The species H₂SO₄, HIO₃, HIO₂, and I₂O₄ elearly fall can be classified into the first category as since the initial change of in voltage difference does not affect the normalized signalsignificantly, i. e., they have a significant impact on the normalised signal. This indicates that these species are detected at the collision limit. It is also apparent that H₂O, HO₂ and SO₂ belong to the third category since a small increase in the voltage difference leads to substantially reduced normalised ratios. Finally, IO, OIO, I₂O₃ and HOI are moderately bonded with Br⁻. These moderately bonded charged clusters can reach a close to collision limit

715 detection if the instrument is softly tuned (the voltage difference is small) but their detection sensitivity can change dramatically if the instrument fragmentation level is high. Lopez-Hilfiker et al. (2016) defined a parameter ΔV_{50} (dV₅₀, i.e., the dV value at half the maximum of signal remaining) to describe the analyte and reagent ion binding strength. In this study, the dV₅₀ is defined by the following equation:

$$NSR = \frac{SR}{1 + e^{-k \times (dV - dV_{50})}} + SR_{max, pred}$$

(4)



Figure 6. Normalised signal remaining vs. the scanning voltage (ΔV). The normalised signal remaining of each species is normalised by the maximum and minimum values of its values with different ΔV (or dV). The ΔV describes the voltage difference between the skimmer and the second quadruple and can be considered an indicator of the softness of the instrument tuning (Lopez-Hilfiker et al., 2016). A higher ΔV commonly indicates a more fragmenting setting.

720 where NSR is the normalised signal remaining, SR is the signal remaining, dV_{50} is the desired fitted value as represented in Figure 7 and SR_{max,pred} is the fitted value that represents the maximum SR when a compound does not undergo fragmentation while passing through the ion optics.

Additionally, formation free enthalpies of various charged clusters are calculated using quantum chemical calculations (see
Methods) and are compared with the voltages at which 50 % of the charged clusters dissociate (ΔV₅₀) dV₅₀ as shown in Figure
7. These-The two sets of parametersare, consisting of theoretical predictions and measurements of the binding strengthand they show a consistent picture, as shown previously and in this study (Lopez-Hilfiker et al., 2016; Iyer et al., 2016). As a , provide a consistent understanding as demonstrated in previous studies (Lopez-Hilfiker et al., 2016; Iyer et al., 2016). In summary, strongly bonded charged clusters have larger fragmentation free exhibit larger fragmentation-free enthalpies, larger
730 ΔV₅₀ valuesand a lower calibration factor (e. g., dV₅₀ values, and lower calibration factors. Examples of such species include H₂SO₄ · Br⁻, and I₂ · Br⁻). The On the other hand, weakly bonded charged clusters have the opposite properties(e.g., exhibit

opposite properties, including species like $HO_2 \cdot Br^-$, $H_2O \cdot Br^-$ and $SO_2 \cdot Br^-$.



Figure 7. The voltage at which 50 % of analyte-bromide adducts have dissociated ($\Delta 50 \text{dV}_{50}$) vs. the fragmentation free enthalpies of the adducts (Table 2).

3.7 Validation of the measurement of iodine-containing species

- Oxidised iodine vapours have been shown to influence atmospheric oxidation capacity (Sherwen et al., 2016; Wang et al., 2021b) (Saiz-Lopez et al., 2014; Sherwen et al., 2016; Wang et al., 2021b) and particle formation processes (Hoffmann et al., 2001; O'Dowd et al., 2002). Recent publications have suggested proposed iodine oxoacids as the critical driver for iodine particle formation processes (Sipilä et al., 2016; Baccarini et al., 2020; He et al., 2021b, a; Zhang et al., 2022; Liu et al., 2023). However, active debate remains concerning the presence of gaseous HIO₃ and whether HIO₃ plays an important role in atmospheric aerosol nucleation. For example, a recent laboratory study shed doubts on the existence of gaseous HIO₃ as the authors only managed to measure HIO₃ in the particle phase with a photoionisation mass spectrometer but not in the gas phase. They concluded that the particle phase HIO₃ was formed from higher iodine oxides , instead of from gaseous HIO₃ (Gómez Martín et al., 2020). Additionally, they hypothesised Furthermore, they proposed a hypothesis that the IO₃⁻ signal, which was previously interpreted as part of the previously attributed to gaseous HIO₃ measured by the measurements using
- NO₃⁻-CIMS (Sipilä et al., 2016), could also originate from I_2O_y (where y = 2-4) species. Their evidence is primarily the exothermicity of the I_2O_{2-3} reactions with NO₃⁻ which forms IO₃⁻ as part of the products. However, it should be noted that exothermic reactions do not guarantee that the reactions occur at significant rates. For example, reactions such as

$$I_2O_3 + NO_3^- \longrightarrow IO_3^- + \text{products}$$
 (5)

involves breaking several-multiple strong I-O and N-O bondsthat-, which are likely associated with high kinetic barriersand

750 one could expect. As a result, it can be anticipated that this reaction does not occur as fast as the rapidly as the reaction HIO₃

 $+ NO_3^- \rightarrow IO_3^- + HNO_3$ reaction, in which case only one, where only a single proton transfer reaction occurstakes place.

It is worthwhile to note that both our earlier studies (He et al., 2021b; Finkenzeller et al., 2022) and (Gómez Martín et al., 2020, 2022) concluded that is the primary form of . Fortunately, gaseous is well measured by both the and chemical ionisation methods.

- 755 Finkenzeller et al. (2022) calculated the cluster formation enthalpy of as -45.6 kcal mol⁻¹, which indicates that the cluster is extremely stable. Gómez Martín et al. (2020) found that the + → products + reaction is endothermic thus less likely to occur. The same also applies to the chemical ionisation method. As already discussed in the last section, voltage scan experiments indicate that the cluster is in fact the most stable cluster among the examined clusters (see Figure 7). Therefore, is detected at the collision limit with the chemical ionisation method and it does not fragment into species such as .
- In a more recent study, Gómez Martín et al. (2022) alternatively used the nitrate chemical ionisation method and detected gaseous HIO₃, thus confirming consistent with our earlier studies (Sipilä et al., 2016; He et al., 2021b, a) of (Sipilä et al., 2016; He et al., 20 regarding the existence of gaseous HIO₃. However, the authors additionally also suggested that the measured HIO₃ \cdot NO₃⁻ ion, which was previously interpreted as HIO₃eould also be formed from , could potentially be formed through reactions such as below: the following:

765
$$I_2O_3 \cdot HNO_3NO_3^- \longrightarrow IONO_2 + HIO_3 \cdot NO_3^-$$
 (6)

due to the reaction being exothermic. Besides the same reasons noted above, this hypothesis is challenged by the fact that the reaction

$$I_2O_3 \cdot HNO_3NO_3^- \longrightarrow I_2O_3 \cdot NO_3^- + HNO_3$$

$$\tag{7}$$

is a favoured pathway compared to the reaction 6 as shown in Figure 8. We further estimate that the MESMER derived overall rate coefficients at 298 K, 1 atm for reactions 6 and 7 and they are 2.3×10^{-12} cm³ molec⁻¹ s⁻¹, and 1.26×10^{-9} cm³ molec⁻¹ s⁻¹, respectively. Therefore, the yield of the reaction 7 is close to unity and cannot affect the HIO₃ detection.

It is essential to highlight that our previous studies (He et al., 2021b; Finkenzeller et al., 2022) as well as the studies by Gómez Martín et al. (2020, 2022) have consistently concluded that I_2O_4 is the predominant form of I_2O_4 . Fortunately, the

775 gaseous I_2O_4 species can be effectively measured using both the NO_3^- and Br^- chemical ionisation methods. Finkenzeller et al. (2022) calculated the cluster formation enthalpy of $I_2O_4 \cdot NO_3^-$ as -45.6 kcal mol⁻¹, which indicates that the $I_2O_4 \cdot NO_3^-$ cluster is extremely stable. Gómez Martín et al. (2020) found that the $I_2O_4 + NO_3^- \longrightarrow$ products + IO_3^- reaction is endothermic thus less likely to occur. The same principle applies to the Br^- chemical ionisation method as well. As mentioned earlier in the previous section, voltage scan experiments have shown that the $I_2O_4 \cdot Br^-$ cluster is the most stable among the clusters investigated (refer

to Figure 7). Consequently, I_2O_4 is detected at the collision limit using the Br⁻ chemical ionisation method, and it does not fragment into species such as IO_3^- . Given that the measured concentration of I_2O_4 is more than one order of magnitude lower

than that of HIO_3 according to previous studies (Wang et al., 2021a; He et al., 2021b; Finkenzeller et al., 2022), it is unlikely that I_2O_y species have a significant impact on the detection of HIO_3 .



Figure 8. Fragmentation pathways of $I_2O_3 \cdot HNO_3NO_3^-$. The enthalpies are calculated at the DLPNO-CCSD(T)/def2-QZVPP//wB97X-D/aug-cc-pVTZ-PP level of theory.

Most importantly, higher iodine oxides and iodine oxoacids are formed through complex and distinct chain reactions. Laboratory experiments with elevated iodine concentrations could inevitably disturb the ratio of iodine oxides to iodine oxoacids. The concentration of iodine monoxide (IO) is commonly considered a good indicator of the intensity of atmospheric iodine activities and was shown to influence the ratio of and-iodine oxides and iodine oxoacids. (Finkenzeller et al., 2022). We took advantage of capitalised on this phenomenon and carried out conducted chemical perturbation experiments by vary-ing the concentration of ozone (O₃while keeping concentration and), while keeping the concentration of iodine (I₂) and the light intensity constant in a laminar flow reactor. The experiments were repeated replicated for both the Br⁻-MION2-T1 and NO₃⁻-MION2-T1 shown, as illustrated in Figure 9. The measured IO₃⁻ signal is compared with HIO₃ · NO₃⁻ from the NO₃⁻-MION2-T1 and to HIO₃ · Br⁻, I₂O₃ · Br⁻ and I₂O₄ · Br⁻ from the Br⁻-MION2-T1 to find out the origin of IO₃⁻. Interestingly,

the gaseous signals. It is worth noting that the gaseous signals of HIO₃ (HIO₃ · NO₃⁻ and HIO₃ · Br⁻) are perfectly linear with respectively linear relationship with the signals of IO₃⁻. However, the signals of I₂O₃ · Br⁻ and

 $I_2O_4 \cdot Br^-$ show-demonstrate a non-linear dependence on IO_3^- . This suggests that the primary source of IO_3^- is gaseous HIO₃, since if I_2O_y does contribute to IO_3^- , a non-linear correlation between HIO₃ $\cdot NO_3^-$ and HIO₃ $\cdot Br^-$ with IO_3^- would be ex-



Figure 9. The normalised IO_3^- signal vs. the normalised signals of a) $HIO_3 \cdot NO_3^-$, b) $HIO_3 \cdot Br^-$, c) $I_2O_3 \cdot Br^-$, and d) $I_2O_4 \cdot Br^-$. The iodine injection and light intensity were kept constant but the O_3 concentration was varied to modulate the ratio of iodine oxides to oxoacids. Error bars show one standard deviation. Notice the different y-axis scales.

- 800 In summary Therefore, we conclude that the I_2O_y is unlikely to significantly contribute to the IO_3^- signal at a significant level in atmospheric relevant under atmospheric conditions. Experiments carried out with ambient level precursors consistently show conducted with ambient-level precursors consistently demonstrate that gaseous I_2O_4 is significantly considerably less abundant compared with to HIO₃ (He et al., 2021b, a; Finkenzeller et al., 2022). Model simulation simulations of iodine chemistry at the Maïdo observatory has further shown that the sum further indicate that the combined concentration of I_2O_3 and I_2O_4 is
- 805 only at around 1 % of HIO_3 thus, making it unlikely to affect HIO_3 measurements and iodine particle formation processes the formation of iodine particles in boundary layer conditions (Finkenzeller et al., 2022).

4 Summary

In this study, we present an upgraded version of the multi-scheme chemical ionisation inletversion 2 (, known as MION2). It is capable of simultaneously operating in atmospheric ion measurement mode and <u>employing</u> multiple chemical ionisation

810 methods. While the Although the fundamental concept of this inlet is identical to the remains the same as MION1 (Rissanen et al., 2019), this new version improves its performance in the new version significantly improves performance by effectively focusing reagent ions(thus having lower LODs), enhances its, resulting in lower limits of detection (LODs). Moreover, it enhances operational stability and additionally allows to operate enables the concurrent use of multiple chemical ionisation methods with the same ionisation time.

815

We further developed a Python open-source flow reactor kinetic model (MARFORCE, see Shen and He (2023)) to simulate convection-diffusion-reaction equations in cylindrical flow reactors in order to calibrate gaseous species such as H_2SO_4 , HOI, and HO₂. The model is also compatible with the widely-used Master Chemical Mechanism, thus allowing future implementation of other chemical mechanisms.

820

825

The MION2 inlet was further characterised for the detection of various inorganic species using the MION2 inlet with the Br⁻ and NO₃⁻ chemical ionisation methods was further characterised at two different ionisation times. H₂SO₄, HOI, and HO₂ were calibrated by utilising the photo-chemical employing the MARFORCE model, which quantifies their photochemical production in a flow reactorand quantification by the MARFORCE model. We further estimate that the LODs are around. Based on our estimations, the limits of detection (LODs) are approximately 10⁵ molec. cm⁻³ (1-min-1-minute data averaging) for e.g., species like H₂SO₄ and HIO₃ when the ionisation time is at set to 35 ms. When By using a longer ionisation time (300 ms), the LOD for H₂SO₄ is further reduced to 2.9×10^4 molec. cm⁻³ (ea. approximately 1 ppqv). A direct comparison shows demonstrates that the MION2 inlet has exhibits comparable or even better LODs compared to the widely-used Eisele inlet (Jokinen et al., 2012). Therefore, the Hence, this upgraded version of the inlet provides extremely high sensitivity toward measuring offers acceptional carcitricity for the measurement of trace areas relevant to atmospheric particle formation.

830 measuring offers exceptional sensitivity for the measurement of trace gases relevant to atmospheric particle formation.

Additionally, we characterised the detection Furthermore, we conducted an assessment of the detection capabilities of SO_2 and I_2 as they are important since they serve as crucial precursors for H_2SO_4 and HIO_3 respectively. We found that the Br⁻-MION2 inlet is capable of detecting SO_2 by diluting a gas cylinder of a known amount containing a known quantity of SO_2 . Possides our previous methods to calibrate In addition to our previously established methods for calibration gaseous L (Worg

Besides our previous methods to calibrate In addition to our previously established methods for calibrating gaseous I_2 (Wang et al., 2021a; Tham et al., 2021), we successfully adapted employed a derivatization approach in combination conjunction with high-performance liquid chromatography method which quantified to quantify the iodine permeation rateof merely, which was found to be as low as 17.3 ng min⁻¹. The I_2 calibration of using the Br⁻-MION2 further shows inlet further confirms that I_2 is detected at the collision limit, similar to H_2SO_4 and consistent with our earlier estimation, and aligns with our previous As the Br⁻-MION2 measures H₂O in the form of H₂O·Br⁻, we quantified the H₂O detection with a dew point mirror instrument by running them side by side. As a large portion of Br⁻ is converted to H₂O·Br⁻ in the ion-molecule reaction chamber, we predicted the fragmentation pathways of analyte-H₂O·Br⁻ clusters using quantum chemical calculations. We show that H₂O evaporates from the analyte-H₂O·Br⁻ clusters when passing the ion optics of our mass spectrometer due to the weak attachment of H₂O to the charged clusters. However, the chemical signature of the analyte is commonly preserved as the analyte-Br⁻ cluster or deprotonated analyte anion. Additionally, the-

The detection using the Br⁻ chemical ionisation method at atmospheric pressure is affected by excessive air water content.
850 For analytes which susceptible to the presence of excessive moisture in the air. Analytes that are detected at the collision limit(e.g., such as H₂SO₄, HIO₃ and I₂), we find a sharp decrease, exhibit a significant decline in measurement sensitivity after when the dew point is above exceeds 0.5 - 10.5 °C (20 - 40 % RH). The Furthermore, the detection of weakly bonded analytes (e.g., analytes, such as HO₂ and SO₂) show intensified water influence even with a dew point, is more profoundly affected by water content, even when the dew point is below 0 °C. For example, LOD of instance, the limit of detection (LOD)
855 for HO₂ is roughly approximately one order of magnitude higher than that of H₂SO₄ at 2.7 % RHand the LOD of , while the LOD for SO₂ is roughly approximately three orders of magnitude higher than that of H₂SO₄ at below 0.1 % RH.

In order to reduce the detection humidity effect, a mitigate the impact of humidity on detection, two methods, namely the dilution method and a the core-sampling method, were tested in this study. We found that these methods do reduce the detection

- 860 humidity effect. Both of these methodsenable both methods effectively reduce the influence of humidity on detection. By employing these methods, it becomes possible to detect ambient level-levels of SO₂ (below 1 part per billion in volume) with by volume) even at RH levels of up to 50 %RH which is otherwise not possible. It, which would otherwise be challenging. However, it should be noted that the utilisation use of these methods unavoidably dilutes the air sample thus inevitably results in sample dilution, thereby affecting the detection of species which are less severely that are less affected by air water content (e.g., such as H₂SO₄, HOI, and I₂). Therefore, these methods should be deployed only employed selectively, when there
- is a <u>elear aimspecific objective</u>, such as detecting extremely low levels of SO₂ or when the sample's dew point is higher than 10 °C (40 % RH). This <u>suggests implies</u> that atmospheric pressure Br⁻ chemical ionisation is suitable for laboratory experiments with controlled relative humidity and RH and for ambient measurements in relatively cold environments. When interpreting data from obtained through the atmospheric pressure Br⁻ chemical ionisation method, the impact of water should
- 870 be carefully treated using it is crucial to carefully account for the influence of water by employing analytical characterisation or fragmentation enthalpy prediction. As predicting fragmentation enthalpy. Despite these considerations, the MION2 allows to operate water insensitive inlet, which allows for the concurrent operation of the water-insensitive NO₃⁻ chemical ionisation method and water sensitive but more capable the water-sensitive yet more versatile Br⁻ chemical ionisation method to operate water insensitive but more capable the water-sensitive yet more versatile Br⁻ chemical ionisation method to entry it will nevertheless reveal greater details of the atmosphere compared to provides a more comprehensive understanding of

875 atmospheric conditions compared to using either of these methods alone in isolation.

Finally, we validated the measurement of gaseous HIO₃ using both the NO₃⁻ and Br⁻ chemical ionisation methods are validated. The signal of HIO₃ commonly typically consists of IO₃⁻ and either HIO₃ · NO₃⁻ or HIO₃ · Br⁻, depending on the chemical ionisation method utilised. We have experimentally and theoretically validated that all of the employed. Through experimental and theoretical validation, we confirmed that all three ions primarily originate from genuine gaseous HIO₃ and that iodine oxides do not contribute to these ions at the formation of these ions under atmospherically relevant conditions.

Code availability. The MARFORCE model is shared through GitHub repository (https://github.com/momo-catcat/MARFORCE-flowtube). Other data analysis codes can be requested from the corresponding authors.

Data availability. Data is available upon request from the corresponding authors.

880



Figure A1. Schematic of the single source an ionisation scheme source of the MION2 inlet.



Figure A2. Schematic of a typical calibration experiment connecting the MION2 inlet (I.D. 24 mm) with the calibration source (I.D. 15.6 mm).



Figure A3. Schematic of the setup for examining the detection humidity effect of H₂SO₄, HOI and HO₂.



Figure A4. Schematic of the experimental setup for iodine chemistry experiments to produce higher concentrations of iodine oxides and oxoacids.



Figure A5. Comparison of the H_2SO_4 profiles at the outlet of a flow reactor. Theoretical values are predicted using Alonso et al. (2016) and the model results indicate the MARFORCE simulation. In both the theoretical prediction and the MARFORCE model, the tube length is assumed to be two meters, the inlet flow is set to 10 slpm and the diffusivity of H_2SO_4 is set to 0.088 cm²s⁻¹.

geometry optimised at the ω B97X-D/aug-ec-pVTZ-PP level of theory at 298.15 K. Color coding: Iodine = purple, oxygen = red, hydrogen = white, bromine = brown.



Figure A6. Comparing calibration experiments of a) H_2SO_4 and HOI with a straight tube (Figure A2) or additionally with a dilution flow (Figure A3). The difference in the calibration coefficients between the two experimental setups is the result of the less accurate representation of fluid dynamics when the dilution flow is added (Figure A3).



Figure A7. Schematic of the setup for H_2SO_4 , HOI and HO_2 calibration experiment with the tower 2. The difference between this setup and the one shown in Figure A3 is that the position of the MION2 tower is changed from tower 1 to tower 2.



Figure A8. The configuration of $HIO_3 \cdot Br^-$ geometry optimised at the core-sampling device (Karsa Ltd ω B97X-D/aug-cc-pVTZ-PP level of theory at 298.15 K.) which is used for adjusting the sheath and sample flowsColor coding: Iodine = purple, oxygen = red, hydrogen = white, browning = brown.



Figure A9. The configuration of the core-sampling device (Karsa Ltd.) which is used for adjusting the sheath and sample flows. This core-sampling piece features three ports for the dilution flows which pass through a mesh and further mixed with the sample flow.

Table A1. Chemical reactions and the reaction rate coefficients used for H₂SO₄ and HOI calibration experiments

Chemical reactions	Reaction rate coefficients	
H ₂ SO ₄ calibration:		
1. OH + SO ₂ = HSO ₃	$^{a}1.32 \times 10^{-12} \times (Temp/300)^{-0.7}$	
2. $OH + HO_2 = H_2O + O_2$	$^{b}4.8 \times 10^{-11} \times \exp(250/Temp)$	
3. $HO_2 + HO_2 = H_2O_2$	$^{b}(2.2 \times 10^{-13} \times \exp(600/Temp) +$	
	$1.9\times 10^{-33}\times M\times \exp(980/Temp))\times KMT06$	
4. OH + OH = H_2O_2	$^{c}2 \times 6.9 \times 10^{-31} \times (Temp/300)^{-0.8} \times p/(1.38 \times 10^{-23})/Temp/10^{6}$	
4. OH + OH = H_2O	${}^{b}6.2 \times 10^{-14} \times (Temp/298)^{2.6} \times \exp(945/Temp)$	
$6. \operatorname{HSO}_3 + \operatorname{O}_2 = \operatorname{HO}_2 + \operatorname{SO}_3$	$^{b}1.3 \times 10^{-12} \times \exp(-330/Temp)$	
7. $SO_3 + 2H_2O = H_2SO_4$	$^{b}3.9 \times 10^{-41} \times \exp(6830.6/Temp)$	

HOI calibration:

1. IO + IO = I + I	$^{d}0.11 \times 5.4 \times 10^{-11} \times \exp(180/Temp)$
2. $IO + IO = OIO + I$	$^{d}0.38 \times 5.4 \times 10^{-11} \times \exp(180/Temp)$
3. IO + IO = I_2O_2	$^{d}0.45 \times 5.4 \times 10^{-11} \times \exp(180/Temp)$
4. $I_2 + OH = HOI + I$	$e^{2.1} \times 10^{-10}$
5. IO + OIO = I_2O_3	${}^{f}w1a \times \exp(w2a \times Temp)$
6. OIO + OIO = I_2O_4	${}^{f}w1b \times \exp(w2b \times Temp)$
7. IO + OH = $HO_2 + I$	$\frac{g_{1.0}-10}{2}$ $\frac{g_{1.0}\times10^{-10}}{2}$
8. HI + OH = $H_2O + I$	${}^{b}1.6 \times 10^{-11} \times \exp(440/Temp)$
9. HOI + OH = H_2O + IO	$\frac{h_{2.0} \times -13}{2.0 \times 10^{-13}}$
10. I + $HO_2 = HI + O_2$	$^{i}1.47 \times 10^{-11} \times \exp(-1090/Temp)$
11. IO + $HO_2 = HOI + O_2$	${}^{b}1.4 \times 10^{-11} \times \exp(540/Temp)$
12. OH + OH = H_2O_2	$^{c}2\times 6.9\times 10^{-31}\times (Temp/300)^{-0.8}\times p/(1.38\times 10^{-23})/Temp/10^{6}$
13. OH + OH = H_2O	${}^{b}6.2 \times 10^{-14} \times (Temp/298)^{2.6} \times \exp(945/Temp)$
14. $OH + HO_2 = H_2O + O_2$	${}^{b}4.8 \times 10^{-11} \times \exp(250/Temp)$
15. $HO_2 + HO_2 = H_2O_2$	$^{b}2.2 \times 10^{-13} \times KMT06 \times \exp(600/Temp) +$
	$1.9\times 10^{-33}\times M\times KMT06\times \exp(980/Temp)$

^{*a*} Wine et al. (1984); ^{*b*} Atkinson et al. (2004); ^{*c*} Zellner et al. (1988); ^{*d*} Bloss et al. (2001); ^{*e*} Gilles et al. (1999); ^{*f*} Saiz-Lopez et al. (2014); ^{*g*} Bösch (2003); ^{*h*} Chameides and Davis (1980); ^{*i*} Jenkin et al. (1990).

KMT06 = 1 + $(1.4 \times 10^{-21} \times \exp(2200/Temp) \times [H_2O])$, [H₂O] is the absolute water concentration. *M* is the total number of molecules in the atmosphere. *p* is the pressure.

 $w1a = 4.7 \times 10^{-10} - 1.4 \times 10^{-5} \times \exp(-0.75 \times p/1.62265) + 5.51868 \times 10^{-10} \times \exp(-0.75 \times p/199.328);$

 $w2a = -0.00331 - 0.00514 \times \exp(-0.75 \times p/325.68711) - 0.00444 \times \exp(-0.75 \times p/40.81609);$

 $w1b = 1.166 \times 10^{-9} - 7.796 \times 10^{-10} \times \exp(-0.75 \times p/22.093) + 1.038 \times 10^{-9} \times \exp(-0.75 \times p/568.154);$

 $w2b = -0.00813 - 0.00382 \times \exp(-0.75 \times p/45.57591) - 0.00643 \times \exp(-0.75 \times p/417.95061).$

Author contributions. X.-C.H. and J.S. designed and carried out the experiments. J.S. and X.-C.H. wrote the MARFORCE model. J.Z. wrote the documentation of the MARFORCE model. S.I. carried out quantum chemical calculations. N.M.M., M.Koi. and M.M.K. analysed molecular iodine samples. J.K., P.J., M.S. and J.M. provided technical support. X.-C.H. wrote the manuscript with contributions from J.S., N.M.M., P.J. and S.I. Finally, J.K., J.S., M.R., S.I., D.R.W. and M.Kul. commented on and edited the manuscript.

890 Competing interests. Paxton Juuti and Jyri Mikkilä work for Karsa, Ltd. Finland. Juha Kangasluoma works partially for Karsa, Ltd. Finland

Acknowledgements. We thank the ACCC Flagship funded by the Academy of Finland grant number 337549, the Academy professorship funded by the Academy of Finland (grant no. 302958), Academy of Finland projects no. 331207, 346370, 325656, 316114, 314798, 325647, 341349 and 349659. This project has received funding from the European Research Council under the European Union's Horizon 2020 research and innovation programme under Grant Contract No. 742206 and 101002728. The Arena for the gap analysis of the existing Arctic Science Co-Operations (AASCO) funded by Prince Albert Foundation Contract No 2859. M.Kul. thanks the Jane and Aatos Erkko Foundation for providing funding. M.Kul. and X.-C.H thank the Jenny and Antti Wihuri Foundation for funding this research. We also thank Miska Olin, Gustaf Lönn and Heikki Junninen for their helpful discussions and contributions to the MARFORCE model. Simon Patrick O'Meara and Gordon McFiggans are acknowledged for their contributions to the MCM interpreter in the MARFORCE model.

895

References

910

915

- 900 Agarwal, B., González-Méndez, R., Lanza, M., Sulzer, P., Märk, T. D., Thomas, N., and Mayhew, C. A.: Sensitivity and Selectivity of Switchable Reagent Ion Soft Chemical Ionization Mass Spectrometry for the Detection of Picric Acid, The Journal of Physical Chemistry A, 118, 8229–8236, https://doi.org/10.1021/jp5010192, 2014.
 - Alonso, M., Carsí, M., and Huang, C.-H.: Using the fully developed concentration profile to determine particle penetration in a laminar flow tube, Journal of Aerosol Science, 97, 34–37, https://doi.org/10.1016/j.jaerosci.2016.04.002, 2016.
- 905 Atkinson, R., Baulch, D. L., Cox, R. A., Crowley, J. N., Hampson, R. F., Hynes, R. G., Jenkin, M. E., Rossi, M. J., and Troe, J.: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume I - gas phase reactions of O_x, HO_x, NO_x and SO_x species, Atmospheric Chemistry and Physics, 4, 1461–1738, https://doi.org/10.5194/acp-4-1461-2004, 2004.
 - Baccarini, A., Karlsson, L., Dommen, J., Duplessis, P., Vüllers, J., Brooks, I. M., Saiz-Lopez, A., Salter, M., Tjernström, M., Baltensperger, U., Zieger, P., and Schmale, J.: Frequent new particle formation over the high Arctic pack ice by enhanced iodine emissions, Nature Communications, 11, 4924, https://doi.org/10.1038/s41467-020-18551-0, 2020.
 - Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurtén, T., Otkjær, R. V., Kjaergaard, H. G., Stratmann, F., Herrmann, H., Sipilä, M., Kulmala, M., and Ehn, M.: Hydroxyl radical-induced formation of highly oxidized organic compounds, Nature Communications, 7, 13 677, https://doi.org/10.1038/ncomms13677, 2016.

Bloss, W. J., Rowley, D. M., Cox, R. A., and Jones, R. L.: Kinetics and Products of the IO Self-Reaction, The Journal of Physical Chemistry A, 105, 7840–7854, https://doi.org/10.1021/jp0044936, 2001.

- Breitenlechner, M., Fischer, L., Hainer, M., Heinritzi, M., Curtius, J., and Hansel, A.: PTR3: An Instrument for Studying the Lifecycle of Reactive Organic Carbon in the Atmosphere, Analytical Chemistry, 89, 5824–5831, https://doi.org/10.1021/acs.analchem.6b05110, 2017.
- Brophy, P. and Farmer, D. K.: A switchable reagent ion high resolution time-of-flight chemical ionization mass spectrometer for real-time measurement of gas phase oxidized species: characterization from the 2013 southern oxidant and aerosol study, Atmospheric Measurement
- 920 Techniques, 8, 2945–2959, https://doi.org/10.5194/amt-8-2945-2015, 2015.
 - Bösch, H.: Upper limits of stratospheric IO and OIO inferred from center-to-limb-darkening-corrected balloon-borne solar occultation visible spectra: Implications for total gaseous iodine and stratospheric ozone, Journal of Geophysical Research, 108, 4455, https://doi.org/10.1029/2002JD003078, 2003.
- Caldwell, G. W., Masucci, J. A., and Ikonomou, M. G.: Negative ion chemical ionization mass spectrometry—binding of molecules to
 bromide and iodide anions, Organic Mass Spectrometry, 24, 8–14, https://doi.org/10.1002/oms.1210240103, 1989.
 - Chai, J.-D. and Head-Gordon, M.: Long-range corrected hybrid density functionals with damped atom–atom dispersion corrections, Physical Chemistry Chemical Physics, 10, 6615, https://doi.org/10.1039/b810189b, 2008.
 - Chameides, W. L. and Davis, D. D.: Iodine: Its possible role in tropospheric photochemistry, Journal of Geophysical Research: Oceans, 85, 7383–7398, https://doi.org/10.1029/JC085iC12p07383, 1980.
- 930 Creasey, D. J., Heard, D. E., and Lee, J. D.: Absorption cross-section measurements of water vapour and oxygen at 185 nm. Implications for the calibration of field instruments to measure OH, HO₂ and RO₂ radicals, Geophysical Research Letters, 27, 1651–1654, https://doi.org/10.1029/1999GL011014, 2000.
 - Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen,
- 935 L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petäjä, T., Wahner, A., Kerminen, V.-M., Kulmala,

M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature, 506, 476–479, https://doi.org/10.1038/nature13032, 2014.

Eisele, F. L. and Tanner, D. J.: Measurement of the gas phase concentration of H₂SO₄ and methane sulfonic acid and estimates of H₂SO₄ production and loss in the atmosphere, Journal of Geophysical Research: Atmospheres, 98, 9001–9010, https://doi.org/10.1029/93JD00031, 1993.

940

965

970

- Feller, D.: The role of databases in support of computational chemistry calculations, Journal of Computational Chemistry, 17, 1571–1586, https://doi.org/10.1002/(SICI)1096-987X(199610)17:13<1571::AID-JCC9>3.0.CO;2-P, 1996.
- Finkenzeller, H., Iyer, S., He, X.-C., Simon, M., Koenig, T. K., Lee, C. F., Valiev, R., Hofbauer, V., Amorim, A., Baalbaki, R., Baccarini, A., Beck, L., Bell, D. M., Caudillo, L., Chen, D., Chiu, R., Chu, B., Dada, L., Duplissy, J., Heinritzi, M., Kemppainen, D., Kim, C., Krechmer,
- J., Kürten, A., Kvashnin, A., Lamkaddam, H., Lee, C. P., Lehtipalo, K., Li, Z., Makhmutov, V., Manninen, H. E., Marie, G., Marten, R., Mauldin, R. L., Mentler, B., Müller, T., Petäjä, T., Philippov, M., Ranjithkumar, A., Rörup, B., Shen, J., Stolzenburg, D., Tauber, C., Tham, Y. J., Tomé, A., Vazquez-Pufleau, M., Wagner, A. C., Wang, D. S., Wang, M., Wang, Y., Weber, S. K., Nie, W., Wu, Y., Xiao, M., Ye, Q., Zauner-Wieczorek, M., Hansel, A., Baltensperger, U., Brioude, J., Curtius, J., Donahue, N. M., Haddad, I. E., Flagan, R. C., Kulmala, M., Kirkby, J., Sipilä, M., Worsnop, D. R., Kurten, T., Rissanen, M., and Volkamer, R.: The gas-phase formation mechanism of iodic acid as an atmospheric aerosol source. Nature Chemistry. https://doi.org/10.1038/s41557-022-01067-z, 2022.
- Frisch, M. J., Trucks, G. W., Schlegel, H. B., Scuseria, G. E., Robb, M. A., Cheeseman, J. R., Scalmani, G., Barone, V., Petersson, G. A., Nakatsuji, H., Li, X., Caricato, M., Marenich, A. V., Bloino, J., Janesko, B. G., Gomperts, R., Mennucci, B., Hratchian, H. P., Ortiz, J. V., Izmaylov, A. F., Sonnenberg, J. L., Williams, Ding, F., Lipparini, F., Egidi, F., Goings, J., Peng, B., Petrone, A., Henderson, T., Ranasinghe, D., Zakrzewski, V. G., Gao, J., Rega, N., Zheng, G., Liang, W., Hada, M., Ehara, M., Toyota, K., Fukuda, R., Hasegawa,
- J., Ishida, M., Nakajima, T., Honda, Y., Kitao, O., Nakai, H., Vreven, T., Throssell, K., Montgomery Jr., J. A., Peralta, J. E., Ogliaro, F., Bearpark, M. J., Heyd, J. J., Brothers, E. N., Kudin, K. N., Staroverov, V. N., Keith, T. A., Kobayashi, R., Normand, J., Raghavachari, K., Rendell, A. P., Burant, J. C., Iyengar, S. S., Tomasi, J., Cossi, M., Millam, J. M., Klene, M., Adamo, C., Cammi, R., Ochterski, J. W., Martin, R. L., Morokuma, K., Farkas, O., Foresman, J. B., and Fox, D. J.: Gaussian 16 Rev. C.01, 2016.

Fuller, E. N., Schettler, P. D., and Giddings, J. C.: NEW METHOD FOR PREDICTION OF BINARY GAS-PHASE DIFFUSION COEFFI CIENTS, Industrial & Engineering Chemistry, 58, 18–27, https://doi.org/10.1021/ie50677a007, 1966.

Gilles, M. K., Burkholder, J. B., and Ravishankara, A. R.: Rate coefficients for the reaction of OH with Cl₂, Br₂, and I₂ from 235 to 354 K, International Journal of Chemical Kinetics, 31, 417–424, https://doi.org/10.1002/(SICI)1097-4601(1999)31:6<417::AID-KIN3>3.0.CO;2-A, 1999.

Gálvez, O., Gómez Martín, J. C., Gómez, P. C., Saiz-Lopez, A., and Pacios, L. F.: A theoretical study on the formation of iodine oxide aggregates and monohydrates, Physical Chemistry Chemical Physics, 15, 15 572, https://doi.org/10.1039/c3cp51219c, 2013.

Gómez Martín, J. C., Lewis, T. R., Blitz, M. A., Plane, J. M. C., Kumar, M., Francisco, J. S., and Saiz-Lopez, A.: A gas-to-particle conversion mechanism helps to explain atmospheric particle formation through clustering of iodine oxides, Nature Communications, 11, 4521, https://doi.org/10.1038/s41467-020-18252-8, 2020.

Gómez Martín, J. C., Lewis, T. R., James, A. D., Saiz-Lopez, A., and Plane, J. M. C.: Insights into the Chemistry of Iodine New Particle Formation: The Role of Iodine Oxides and the Source of Iodic Acid, Journal of the American Chemical Society, p. jacs.1c12957, https://doi.org/10.1021/jacs.1c12957, 2022.

Gormley, P. and Kennedy, M.: Diffusion from a stream flowing through a cylindrical tube, in: Proceedings of the Royal Irish Academy. Section A: Mathematical and Physical Sciences, vol. 52, pp. 163–169, JSTOR, 1948.

Hansel, A., Jordan, A., Holzinger, R., Prazeller, P., Vogel, W., and Lindinger, W.: Proton transfer reaction mass spectrometry: on-line trace gas

- 975 analysis at the ppb level, International Journal of Mass Spectrometry and Ion Processes, 149-150, 609–619, https://doi.org/10.1016/0168-1176(95)04294-U, 1995.
 - Hanson, D. R. and Eisele, F.: Diffusion of H₂SO ₄ in Humidified Nitrogen: Hydrated H₂SO₄, The Journal of Physical Chemistry A, 104, 1715–1719, https://doi.org/10.1021/jp993622j, 2000.
 - He, X.-C.: From the measurement of halogenated species to iodine particle formation, Ph.D. thesis, University of Helsinki, Helsinki, https://www.action.com/actional-acti
- 980 //helda.helsinki.fi/handle/10138/229173, 2017.
- He, X.-C., Iyer, S., Sipilä, M., Ylisirniö, A., Peltola, M., Kontkanen, J., Baalbaki, R., Simon, M., Kürten, A., Tham, Y. J., Pesonen, J., Ahonen, L. R., Amanatidis, S., Amorim, A., Baccarini, A., Beck, L., Bianchi, F., Brilke, S., Chen, D., Chiu, R., Curtius, J., Dada, L., Dias, A., Dommen, J., Donahue, N. M., Duplissy, J., El Haddad, I., Finkenzeller, H., Fischer, L., Heinritzi, M., Hofbauer, V., Kangasluoma, J., Kim, C., Koenig, T. K., Kubečka, J., Kvashnin, A., Lamkaddam, H., Lee, C. P., Leiminger, M., Li, Z., Makhmutov, V., Xiao, M., Marten, R., Nie, W., Onnela, A., Partoll, E., Petäjä, T., Salo, V.-T., Schuchmann, S., Steiner, G., Stolzenburg, D., Stozhkov, Y., Tauber, C., Tomé, A., Väisänen, O., Vazquez-Pufleau, M., Volkamer, R., Wagner, A. C., Wang, M., Wang, Y., Wimmer, D., Winkler, P. M., Worsnop, D. R., Wu, Y., Yan, C., Ye, Q., Lehtinen, K., Nieminen, T., Manninen, H. E., Rissanen, M., Schobesberger, S., Lehtipalo, K., Baltensperger, U., Hansel, A., Kerminen, V.-M., Flagan, R. C., Kirkby, J., Kurtén, T., and Kulmala, M.: Determination of the collision rate coefficient between charged iodic acid clusters and iodic acid using the appearance time method, Aerosol Science and Technology, 55, 231–242,
- 990 https://doi.org/10.1080/02786826.2020.1839013, 2021a.
 - He, X.-C., Tham, Y. J., Dada, L., Wang, M., Finkenzeller, H., Stolzenburg, D., Iyer, S., Simon, M., Kürten, A., Shen, J., Rörup, B., Rissanen, M., Schobesberger, S., Baalbaki, R., Wang, D. S., Koenig, T. K., Jokinen, T., Sarnela, N., Beck, L. J., Almeida, J., Amanatidis, S., Amorim, A., Ataei, F., Baccarini, A., Bertozzi, B., Bianchi, F., Brilke, S., Caudillo, L., Chen, D., Chiu, R., Chu, B., Dias, A., Ding, A., Dommen, J., Duplissy, J., El Haddad, I., Gonzalez Carracedo, L., Granzin, M., Hansel, A., Heinritzi, M., Hofbauer, V., Junninen, H., Kangasluoma,
- J., Kemppainen, D., Kim, C., Kong, W., Krechmer, J. E., Kvashin, A., Laitinen, T., Lamkaddam, H., Lee, C. P., Lehtipalo, K., Leiminger, M., Li, Z., Makhmutov, V., Manninen, H. E., Marie, G., Marten, R., Mathot, S., Mauldin, R. L., Mentler, B., Möhler, O., Müller, T., Nie, W., Onnela, A., Petäjä, T., Pfeifer, J., Philippov, M., Ranjithkumar, A., Saiz-Lopez, A., Salma, I., Scholz, W., Schuchmann, S., Schulze, B., Steiner, G., Stozhkov, Y., Tauber, C., Tomé, A., Thakur, R. C., Väisänen, O., Vazquez-Pufleau, M., Wagner, A. C., Wang, Y., Weber, S. K., Winkler, P. M., Wu, Y., Xiao, M., Yan, C., Ye, Q., Ylisirniö, A., Zauner-Wieczorek, M., Zha, Q., Zhou, P., Flagan, R. C., Curtius,
- J., Baltensperger, U., Kulmala, M., Kerminen, V.-M., Kurtén, T., Donahue, N. M., Volkamer, R., Kirkby, J., Worsnop, D. R., and Sipilä,
 M.: Role of iodine oxoacids in atmospheric aerosol nucleation, Science, 371, 589–595, https://doi.org/10.1126/science.abe0298, 2021b.
 - Hearn, J. D. and Smith, G. D.: A Chemical Ionization Mass Spectrometry Method for the Online Analysis of Organic Aerosols, Analytical Chemistry, 76, 2820–2826, https://doi.org/10.1021/ac049948s, 2004.
- Hoffmann, T., O'Dowd, C. D., and Seinfeld, J. H.: Iodine oxide homogeneous nucleation: An explanation for coastal new particle production,
 Geophysical Research Letters, 28, 1949–1952, https://doi.org/10.1029/2000GL012399, 2001.
 - Huey, L. G.: Measurement of trace atmospheric species by chemical ionization mass spectrometry: Speciation of reactive nitrogen and future directions, Mass Spectrometry Reviews, 26, 166–184, https://doi.org/10.1002/mas.20118, 2007.
 - Iyer, S., Lopez-Hilfiker, F., Lee, B. H., Thornton, J. A., and Kurtén, T.: Modeling the Detection of Organic and Inorganic Compounds Using Iodide-Based Chemical Ionization, The Journal of Physical Chemistry A, 120, 576–587, https://doi.org/10.1021/acs.jpca.5b09837, 2016.

- 1010 Iyer, S., He, X., Hyttinen, N., Kurtén, T., and Rissanen, M. P.: Computational and Experimental Investigation of the Detection of HO₂ Radical and the Products of Its Reaction with Cyclohexene Ozonolysis Derived RO₂ Radicals by an Iodide-Based Chemical Ionization Mass Spectrometer, The Journal of Physical Chemistry A, 121, 6778–6789, https://doi.org/10.1021/acs.jpca.7b01588, 2017.
 - Jenkin, M. E., Cox, R. A., Mellouki, A., Le Bras, G., and Poulet, G.: Kinetics of the reaction of iodine atoms with hydroperoxy radicals, The Journal of Physical Chemistry, 94, 2927–2934, https://doi.org/10.1021/j100370a036, 1990.
- 1015 Jenkin, M. E., Saunders, S. M., and Pilling, M. J.: The tropospheric degradation of volatile organic compounds: a protocol for mechanism development, Atmospheric Environment, 31, 81–104, https://doi.org/10.1016/S1352-2310(96)00105-7, 1997.
 - Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin, R. L., Kulmala, M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-APi-TOF, Atmospheric Chemistry and Physics, 12, 4117–4125, https://doi.org/10.5194/acp-12-4117-2012, 2012.
- 1020 Jordan, A., Haidacher, S., Hanel, G., Hartungen, E., Herbig, J., Märk, L., Schottkowsky, R., Seehauser, H., Sulzer, P., and Märk, T.: An online ultra-high sensitivity Proton-transfer-reaction mass-spectrometer combined with switchable reagent ion capability (PTR+SRI-MS), International Journal of Mass Spectrometry, 286, 32–38, https://doi.org/10.1016/j.ijms.2009.06.006, 2009.
- Junninen, H., Ehn, M., Petäjä, T., Luosujärvi, L., Kotiaho, T., Kostiainen, R., Rohner, U., Gonin, M., Fuhrer, K., Kulmala, M., and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion composition, Atmospheric Measurement Techniques, 3, 1039–1053, https://doi.org/10.5194/amt-3-1039-2010, 2010.
 - Kendall, R. A., Dunning, T. H., and Harrison, R. J.: Electron affinities of the first-row atoms revisited. Systematic basis sets and wave functions, The Journal of Chemical Physics, 96, 6796–6806, https://doi.org/10.1063/1.462569, 1992.
 - Kercher, J. P., Riedel, T. P., and Thornton, J. A.: Chlorine activation by N₂O₅: simultaneous, in situ detection of ClNO₂ and N₂O₅ by chemical ionization mass spectrometry, Atmos. Meas. Tech., 2, 193–204, https://doi.org/10.5194/amt-2-193-2009, 2009.
- 1030 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Vanhanen, J., Viisa-
- 1035 nen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, Nature, 476, 429–433, https://doi.org/10.1038/nature10343, 2011.
 - Kürten, A., Rondo, L., Ehrhart, S., and Curtius, J.: Calibration of a Chemical Ionization Mass Spectrometer for the Measurement of Gaseous Sulfuric Acid, The Journal of Physical Chemistry A, 116, 6375–6386, https://doi.org/10.1021/jp212123n, 2012.
- 1040 Lagg, A., Taucher, J., Hansel, A., and Lindinger, W.: Applications of proton transfer reactions to gas analysis, International Journal of Mass Spectrometry and Ion Processes, 134, 55–66, https://doi.org/10.1016/0168-1176(94)03965-8, 1994.
 - Laskin, J., Laskin, A., and Nizkorodov, S. A.: Mass Spectrometry Analysis in Atmospheric Chemistry, Analytical Chemistry, 90, 166–189, https://doi.org/10.1021/acs.analchem.7b04249, 2018.
 - Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurtén, T., Worsnop, D. R., and Thornton, J. A.: An Iodide-Adduct High-Resolution Time-of-
- 1045 Flight Chemical-Ionization Mass Spectrometer: Application to Atmospheric Inorganic and Organic Compounds, Environmental Science & Technology, 48, 6309–6317, https://doi.org/10.1021/es500362a, 2014.

- Liao, J., Huey, L. G., Liu, Z., Tanner, D. J., Cantrell, C. A., Orlando, J. J., Flocke, F. M., Shepson, P. B., Weinheimer, A. J., Hall, S. R., Ullmann, K., Beine, H. J., Wang, Y., Ingall, E. D., Stephens, C. R., Hornbrook, R. S., Apel, E. C., Riemer, D., Fried, A., Mauldin, R. L., Smith, J. N., Staebler, R. M., Neuman, J. A., and Nowak, J. B.: High levels of molecular chlorine in the Arctic atmosphere, Nature
- 1050 Geoscience, 7, 91–94, https://doi.org/10.1038/ngeo2046, 2014.
 - Liu, L., Li, S., Zu, H., and Zhang, X.: Unexpectedly significant stabilizing mechanism of iodous acid on iodic acid nucleation under different atmospheric conditions, Science of The Total Environment, 859, 159 832, https://doi.org/10.1016/j.scitotenv.2022.159832, 2023.
 - Lopez-Hilfiker, F. D., Iyer, S., Mohr, C., Lee, B. H., D'Ambro, E. L., Kurtén, T., and Thornton, J. A.: Constraining the sensitivity of iodide adduct chemical ionization mass spectrometry to multifunctional organic molecules using the collision limit and thermodynamic stability of iodide ion adducts, Atmospheric Measurement Techniques, 9, 1505–1512, https://doi.org/10.5194/amt-9-1505-2016, 2016.
- of iodide ion adducts, Atmospheric Measurement Techniques, 9, 1505–1512, https://doi.org/10.5194/amt-9-1505-2016, 2016.
 Mielke, L. H., Furgeson, A., and Osthoff, H. D.: Observation of ClNO₂ in a Mid-Continental Urban Environment, Environmental Science & Technology, 45, 8889–8896, https://doi.org/10.1021/es201955u, 2011.
 - Mishra, S., Singh, V., Jain, A., and Verma, K. K.: Determination of iodide by derivatization to 4-iodo-N,N-dimethylaniline and gas chromatography-mass spectrometry, The Analyst, 125, 459–464, https://doi.org/10.1039/a908363d, 2000.
- 1060 Munson, M. S. B. and Field, F. H.: Chemical Ionization Mass Spectrometry. I. General Introduction, Journal of the American Chemical Society, 88, 2621–2630, https://doi.org/10.1021/ja00964a001, 1966.
 - Neese, F.: The ORCA program system: The ORCA program system, Wiley Interdisciplinary Reviews: Computational Molecular Science, 2, 73–78, https://doi.org/10.1002/wcms.81, 2012.
- O'Dowd, C. D., Jimenez, J. L., Bahreini, R., Flagan, R. C., Seinfeld, J. H., Hämeri, K., Pirjola, L., Kulmala, M., Jennings, S. G., and Hoffmann, T.: Marine aerosol formation from biogenic iodine emissions, Nature, 417, 632–636, https://doi.org/10.1038/nature00775,
 - 2002.
 - O'Meara, S. P., Xu, S., Topping, D., Alfarra, M. R., Capes, G., Lowe, D., Shao, Y., and McFiggans, G.: PyCHAM (v2.1.1): a Python box model for simulating aerosol chambers, Geoscientific Model Development, 14, 675–702, https://doi.org/10.5194/gmd-14-675-2021, 2021.
 Pan, Y., Zhang, Q., Zhou, W., Zou, X., Wang, H., Huang, C., Shen, C., and Chu, Y.: Detection of Ketones by a Novel Technology: Dipolar
- 1070 Proton Transfer Reaction Mass Spectrometry (DP-PTR-MS), Journal of the American Society for Mass Spectrometry, 28, 873–879, https://doi.org/10.1007/s13361-017-1638-7, 2017.
 - Passananti, M., Zapadinsky, E., Zanca, T., Kangasluoma, J., Myllys, N., Rissanen, M. P., Kurtén, T., Ehn, M., Attoui, M., and Vehkamäki, H.: How well can we predict cluster fragmentation inside a mass spectrometer?, Chemical Communications, 55, 5946–5949, https://doi.org/10.1039/C9CC02896J, 2019.
- 1075 Peterson, K. A., Figgen, D., Goll, E., Stoll, H., and Dolg, M.: Systematically convergent basis sets with relativistic pseudopotentials. II. Small-core pseudopotentials and correlation consistent basis sets for the post- *d* group 16–18 elements, The Journal of Chemical Physics, 119, 11 113–11 123, https://doi.org/10.1063/1.1622924, 2003.
 - Riplinger, C. and Neese, F.: An efficient and near linear scaling pair natural orbital based local coupled cluster method, The Journal of Chemical Physics, 138, 034 106, https://doi.org/10.1063/1.4773581, 2013.
- 1080 Riplinger, C., Sandhoefer, B., Hansen, A., and Neese, F.: Natural triple excitations in local coupled cluster calculations with pair natural orbitals, The Journal of Chemical Physics, 139, 134 101, https://doi.org/10.1063/1.4821834, 2013.
 - Rissanen, M. P., Mikkilä, J., Iyer, S., and Hakala, J.: Multi-scheme chemical ionization inlet (MION) for fast switching of reagent ion chemistry in atmospheric pressure chemical ionization mass spectrometry (CIMS) applications, Atmospheric Measurement Techniques, 12, 6635–6646, https://doi.org/10.5194/amt-12-6635-2019, 2019.

- 1085 Saiz-Lopez, A., Fernandez, R. P., Ordóñez, C., Kinnison, D. E., Gómez Martín, J. C., Lamarque, J.-F., and Tilmes, S.: Iodine chemistry in the troposphere and its effect on ozone, Atmospheric Chemistry and Physics, 14, 13 119–13 143, https://doi.org/10.5194/acp-14-13119-2014, 2014.
- Sanchez, J., Tanner, D. J., Chen, D., Huey, L. G., and Ng, N. L.: A new technique for the direct detection of HO₂ radicals using bro-mide chemical ionization mass spectrometry (Br-CIMS): initial characterization, Atmospheric Measurement Techniques, 9, 3851–3861, https://doi.org/10.5194/amt-9-3851-2016, 2016.
 - Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic volatile organic compounds, Atmospheric Chemistry and Physics, 3, 161–180, https://doi.org/10.5194/acp-3-161-2003, 2003.

Shen, J. and He, X.-C.: MARFORCE-Flowtube model, https://github.com/momo-catcat/MARFORCE-flowtube, 2023.

- 1095 Sherwen, T., Evans, M. J., Carpenter, L. J., Andrews, S. J., Lidster, R. T., Dix, B., Koenig, T. K., Sinreich, R., Ortega, I., Volkamer, R., Saiz-Lopez, A., Prados-Roman, C., Mahajan, A. S., and Ordóñez, C.: Iodine's impact on tropospheric oxidants: a global model study in GEOS-Chem, Atmospheric Chemistry and Physics, 16, 1161–1186, https://doi.org/10.5194/acp-16-1161-2016, 2016.
 - Sipilä, M., Sarnela, N., Jokinen, T., Henschel, H., Junninen, H., Kontkanen, J., Richters, S., Kangasluoma, J., Franchin, A., Peräkylä, O., Rissanen, M. P., Ehn, M., Vehkamäki, H., Kurten, T., Berndt, T., Petäjä, T., Worsnop, D., Ceburnis, D., Kerminen, V.-M., Kulmala,
- 1100 M., and O'Dowd, C.: Molecular-scale evidence of aerosol particle formation via sequential addition of HIO₃, Nature, 537, 532–534, https://doi.org/10.1038/nature19314, 2016.
 - Smith, D. and Španěl, P.: Selected ion flow tube mass spectrometry (SIFT-MS) for on-line trace gas analysis, Mass Spectrometry Reviews, 24, 661–700, https://doi.org/10.1002/mas.20033, 2005.
 - Tham, Y. J., He, X.-C., Li, Q., Cuevas, C. A., Shen, J., Kalliokoski, J., Yan, C., Iyer, S., Lehmusjärvi, T., Jang, S., Thakur, R. C., Beck,
- 1105 L., Kemppainen, D., Olin, M., Sarnela, N., Mikkilä, J., Hakala, J., Marbouti, M., Yao, L., Li, H., Huang, W., Wang, Y., Wimmer, D., Zha, Q., Virkanen, J., Spain, T. G., O'Doherty, S., Jokinen, T., Bianchi, F., Petäjä, T., Worsnop, D. R., Mauldin, R. L., Ovad-nevaite, J., Ceburnis, D., Maier, N. M., Kulmala, M., O'Dowd, C., Dal Maso, M., Saiz-Lopez, A., and Sipilä, M.: Direct field evidence of autocatalytic iodine release from atmospheric aerosol, Proceedings of the National Academy of Sciences, 118, e2009951118, https://doi.org/10.1073/pnas.2009951118, 2021.
- 1110 Thornton, J. A., Kercher, J. P., Riedel, T. P., Wagner, N. L., Cozic, J., Holloway, J. S., Dubé, W. P., Wolfe, G. M., Quinn, P. K., Middlebrook, A. M., Alexander, B., and Brown, S. S.: A large atomic chlorine source inferred from mid-continental reactive nitrogen chemistry, Nature, 464, 271–274, https://doi.org/10.1038/nature08905, 2010.
 - Veres, P., Roberts, J. M., Warneke, C., Welsh-Bon, D., Zahniser, M., Herndon, S., Fall, R., and de Gouw, J.: Development of negative-ion proton-transfer chemical-ionization mass spectrometry (NI-PT-CIMS) for the measurement of gas-phase organic acids in the atmosphere,
- International Journal of Mass Spectrometry, 274, 48–55, https://doi.org/10.1016/j.ijms.2008.04.032, 2008.
 Wang, M., He, X.-C., Finkenzeller, H., Iyer, S., Chen, D., Shen, J., Simon, M., Hofbauer, V., Kirkby, J., Curtius, J., Maier, N., Kurtén, T., Worsnop, D. R., Kulmala, M., Rissanen, M., Volkamer, R., Tham, Y. J., Donahue, N. M., and Sipilä, M.: Measurement of iodine species and sulfuric acid using bromide chemical ionization mass spectrometers, Atmospheric Measurement Techniques, 14, 4187–4202, https://doi.org/10.5194/amt-14-4187-2021, 2021a.
- 1120 Wang, S., McNamara, S. M., Moore, C. W., Obrist, D., Steffen, A., Shepson, P. B., Staebler, R. M., Raso, A. R. W., and Pratt, K. A.: Direct detection of atmospheric atomic bromine leading to mercury and ozone depletion, Proceedings of the National Academy of Sciences, 116, 14479–14484, https://doi.org/10.1073/pnas.1900613116, 2019.

- Wang, X., Jacob, D. J., Downs, W., Zhai, S., Zhu, L., Shah, V., Holmes, C. D., Sherwen, T., Alexander, B., Evans, M. J., Eastham, S. D., Neuman, J. A., Veres, P. R., Koenig, T. K., Volkamer, R., Huey, L. G., Bannan, T. J., Percival, C. J., Lee, B. H., and Thornton, J. A.:
- 1125 Global tropospheric halogen (Cl, Br, I) chemistry and its impact on oxidants, Atmospheric Chemistry and Physics, 21, 13973–13996, https://doi.org/10.5194/acp-21-13973-2021, 2021b.
 - Weigend, F. and Ahlrichs, R.: Balanced basis sets of split valence, triple zeta valence and quadruple zeta valence quality for H to Rn: Design and assessment of accuracy, Physical Chemistry Chemical Physics, 7, 3297, https://doi.org/10.1039/b508541a, 2005.
 - Westmore, J. B. and Alauddin, M. M.: Ammonia chemical ionization mass spectrometry, Mass Spectrometry Reviews, 5, 381-465,
- 1130 https://doi.org/10.1002/mas.1280050403, 1986.
 - Wine, P. H., Thompson, R. J., Ravishankara, A. R., Semmes, D. H., Gump, C. A., Torabi, A., and Nicovich, J. M.: Kinetics of the reaction OH + SO₂ + M -> HOSO₂ + M. Temperature and pressure dependence in the fall-off region, The Journal of Physical Chemistry, 88, 2095–2104, https://doi.org/10.1021/j150654a031, 1984.
 - Woodward-Massey, R., Taha, Y. M., Moussa, S. G., and Osthoff, H. D.: Comparison of negative-ion proton-transfer with iodide ion
- 1135 chemical ionization mass spectrometry for quantification of isocyanic acid in ambient air, Atmospheric Environment, 98, 693–703, https://doi.org/10.1016/j.atmosenv.2014.09.014, 2014.
 - Zellner, R., Ewig, F., Paschke, R., and Wagner, G.: Pressure and temperature dependence of the gas-phase recombination of hydroxyl radicals, The Journal of Physical Chemistry, 92, 4184–4190, https://doi.org/10.1021/j100325a038, 1988.
 - Zhang, R., Xie, H.-B., Ma, F., Chen, J., Iyer, S., Simon, M., Heinritzi, M., Shen, J., Tham, Y. J., Kurtén, T., Worsnop, D. R., Kirkby, J.,
- 1140 Curtius, J., Sipilä, M., Kulmala, M., and He, X.-C.: Critical Role of Iodous Acid in Neutral Iodine Oxoacid Nucleation, Environmental Science & Technology, 56, 14166–14177, https://doi.org/10.1021/acs.est.2c04328, 2022.