

1 **Stability assessment of organic sulfur and organosulfate compounds**
2 **in filter samples for quantification by Fourier Transform-Infrared**
3 **Spectroscopy**

4 Marife B. Anunciado¹, Miranda De Boskey², Laura Haines², Katarina Lindskog², Tracy
5 Dombek², Satoshi Takahama³, Ann M. Dillner¹

6 ¹Air Quality Research Center, University of California Davis, Davis, California, United States

7 ²Research Triangle Institute, Research Triangle Park, North Carolina, United States

8 ³Laboratory of Atmospheric Processes and their Impacts (LAPI), ENAC/IIE, Ecole
9 Polytechnique Fédérale de Lausanne (EPFL), Switzerland

10 *Correspondence to:* Ann M. Dillner (amdillner@ucdavis.edu)

11 **Abstract.** Organic sulfur and sulfate compounds, tracers for sources and atmospheric processes,
12 are not currently measured in national monitoring networks such as the Interagency Monitoring of
13 Protected Visual Environments (IMPROVE). The goal of this paper is to begin to assess the
14 stability of organic sulfur and sulfate containing compounds on polytetrafluoroethylene (PTFE)
15 filters and the suitability of Fourier-transform infrared (FT-IR) spectroscopy to measure these
16 compounds. Stability assessment is needed because PTFE samples collected by IMPROVE are
17 typically stored 6-9 months prior to analysis. For this study, two organosulfur compounds,
18 methanesulfonic acid (MSA) and hydroxymethanesulfonate ion (HMS), and two organosulfate
19 compounds, methyl sulfate (MS) and 2-methyltetrol sulfate (2-MTS), are collected individually
20 on PTFE filters. Gravimetric mass measurements are used to assess mass stability over time. FT-
21 IR spectra are evaluated to assess the capability of measuring the compound from PTFE filters by

22 assessing the compound stability or chemical changes over time. Ion chromatography (IC) and
23 Inductively Coupled Plasma Optical Emission Spectroscopy (ICP-OES) are used as an additional
24 tool to assess stability or chemical changes over time. MS has the highest potential to be measured
25 by FT-IR in IMPROVE samples. For MS, a simple organosulfate, the mass changes are within
26 measurement uncertainty and FT-IR spectra indicate no compositional change over a 4-month
27 period, suggesting MS can be measured using FT-IR. IC and ICP-OES support the conclusion
28 that MS is stable on the filter. However, for 2-MTS, the other organosulfate measured in this study,
29 spectral changes after a month on the filter suggests it decomposes into other organosulfates or an
30 inorganic sulfate. MSA in IMPROVE samples can be measured, but only as a lower bound, due
31 to volatility off of the filter as indicated by FT-IR and gravimetry. FT-IR and IC both show that
32 MSA is not chemically changing over the course of the study. Measurements by all methods
33 indicate HMS is unstable on PTFE filter and IC and FT-IR indicate that it likely converts to
34 inorganic sulfate. Future work includes the evaluation of these compounds in as ambient aerosol
35 sample matrix to determine any differences in stability, identifying interferences that could limit
36 quantification and developing calibrations to measure the compounds or functional groups in
37 ambient samples.

38 **1 Introduction**

39 Organic sulfur compounds exist in particulate form in the atmosphere and can be the result of
40 natural processes (e.g. marine sulfur, volcanic emissions) (Aneja and Cooper, 1989; Bates et al.,
41 1992) or activities of anthropogenic sources (e.g. combustion, sulfur-rich wastewaters, smelting)
42 (Grübler, 1998; Smith et al., 2011). Organic sulfur compounds can be categorized as
43 organosulfur compounds such as sulfones (RSO_2) and sulfonic acids (RSO_3^-) having C-S bonds,
44 while organosulfates (ROSO_3^-) have a C-O-S bond in the structure (Song et al., 2019). Two

45 sulfonic acid compounds, methanesulfonic acid (MSA), a tracer for marine aerosol, and
46 hydroxymethanesulfonate ion (HMS), measured in high haze conditions, along with two
47 organosulfates compounds, methyl sulfate (MS) and 2-methyltetrol sulfate (2-MTS) were
48 selected for evaluation.

49 Methanesulfonic acid forms from photochemical oxidation of dimethyl sulfide (DMS) (von
50 Glasow and Crutzen, 2004; Kwong et al., 2018a). DMS is a naturally occurring sulfur species
51 produced by marine algae or phytoplankton and is an important precursor of sulfur dioxide, non-
52 sea salt inorganic sulfate and organosulfur compounds, including MSA (Barnes et al., 1994;
53 Hoffmann et al., 2016). This makes MSA a tracer for marine aerosol (Allen et al., 1997; Becagli
54 et al., 2013; Saltzman et al., 1986). Ion chromatography (IC) has been used to measure MSA in
55 ambient aerosol collected on PTFE filters (Amore et al., 2022) and nucleopore filters (Allen et
56 al., 2002). MSA has also been measured in water soluble fractions of ambient aerosol using
57 proton nuclear magnetic resonance (HNMR) (Decesari et al., 2000). Fourier-transform infrared
58 spectroscopy (FT-IR) has been used to characterize liquid and solid MSA in the laboratory
59 studies (Lee et al., 2019; Zhong and Parker, 2022; Chackalackal and Stafford, 1966) as has
60 Raman spectroscopy (Zhong and Parker, 2022), but to the best of our knowledge, neither FT-IR
61 or Raman spectroscopy have been used to measure MSA in complex mixtures like ambient
62 aerosol samples.

63 Hydroxymethanesulfonate (HMS), formed by sulfite and formaldehyde in aqueous phase, is a
64 strong acid that is stable at low pH (Seinfeld and Pandis, 2016) and is a tracer for aqueous
65 processes (Chen et al., 2022). During severe winter haze in the North China Plain, HMS was
66 measured using real-time single particle mass spectrum instruments and filter-based IC methods
67 during periods of high SO₂ and HCHO concentrations and low oxidant concentrations in

68 particles with high liquid water content (Ma et al., 2020). Very high concentrations of HMS
69 have been measured in Fairbanks, Alaska during pollution events in a cold, dark and humid
70 environment (Campbell et al., 2022). In the Interagency Monitoring of Protected Visual
71 Environments (IMPROVE) Network, there is evidence of an ubiquitous presence of HMS in ion
72 chromatograms of samples collected at 150 sites in the United States (Moch et al., 2020).

73 However, HMS can be challenging to measure (Moch et al., 2018). Single particle mass
74 spectrometry techniques have identified m/z 111 as characteristic for HMS (Chapman et al.,
75 1990; Lee et al., 2003; Song et al., 2019). However, methyl sulfate and other organic sulfur
76 compounds have the same characteristic m/z which makes quantifying HMS using mass
77 spectrometry challenging (Dovrou et al., 2019; Lee et al., 2003). High-resolution aerosol mass
78 spectrometry (HR-AMS) has been used to measure HMS and organosulfates, however the
79 majority of compounds mostly fragment into inorganic sulfate and a non-sulfur containing
80 organic fraction, leading to an underestimation of HMS and overestimation of inorganic sulfate
81 (Dovrou et al., 2019; Song et al., 2019). HMS has been measured in field and laboratory studies
82 by IC (Dovrou et al., 2019; Campbell et al., 2022), however notable challenges have been
83 documented. HMS and sulfate are not fully resolved in all IC methods (Campbell et al., 2022)
84 leading to poor resolution that can introduce error into the results for both HMS and sulfate
85 (Dovrou et al., 2019; Ma et al., 2020). In IC methods where HMS and sulfate are well resolved,
86 HMS and sulfite may be unresolved and co-elute with bisulfite (Moch et al., 2018; Wei et al.,
87 2020). Additionally, HMS may degrade to sulfite and formaldehyde at the high-pH eluent used
88 in IC (Moch et al., 2020). Degradation of both HMS and sulfite may occur in aqueous solutions
89 prior to analysis or in the column during analysis and it's suspected that some of the sulfite
90 oxidizes to sulfate in solution or in the column (Moch et al., 2020). The formation of HMS in

91 the atmosphere occurs at moderate pH and pH differences, between the filter and atmospheric
92 condition (e.g. cloud, fog, pH), can contribute to HMS sample mass loss off the filter leading to
93 an underestimation of HMS (Moch et al., 2020). With proper columns and eluent composition,
94 IC has been shown to separate HMS and sulfate peaks with only a small underestimation of
95 HMS due to sulfate conversion (Dovrou et al., 2019; Campbell et al., 2022). At least one
96 laboratory study (published in Japanese) has characterized HMS by FT-IR (Sato et al., 1984) but
97 FT-IR has not been used to measure HMS in ambient aerosol samples to the best of our
98 knowledge.

99 Organosulfates are the most abundant form of organic sulfur compounds in atmospheric particles
100 (Hettiyadura et al., 2015; Stone et al., 2012; Hawkins et al., 2010; Frossard et al., 2011; Olson et
101 al., 2011). Organosulfates are secondary organic aerosols (SOA) from oxidation of mostly
102 biogenic, but also some anthropogenic volatile organic compounds, in the presence of acidic
103 sulfate (Hettiyadura et al., 2015; Stone et al., 2012; Wang et al., 2021) and have been suggested
104 to be tracers for SOA (Wang et al., 2021; Chen et al., 2021). Organosulfates have been measured
105 in ambient aerosol globally including at four sites in Asia where on average they contribute <1%
106 of PM_{2.5}, 2.3% of organic carbon and 3.8% of total sulfate (Stone et al., 2012). In Arctic haze
107 aerosols in the spring, organosulfates contributed to 13% of organic matter (OM) (Hansen et al.,
108 2014) and contributed to OM at varying levels across the US, with higher levels in summer
109 (Chen et al., 2021). Most studies have used a liquid chromatography method coupled to a mass
110 spectrometer (LC-MS) for measuring organosulfates (Hettiyadura et al., 2015; Wang et al.,
111 2021). FT-IR has been used to measure total organosulfate functional groups (Hawkins et al.,
112 2010) and Raman (Lloyd and Dodgson, 1961; Bondy et al., 2018) and FT-IR (Lloyd et al., 1961)
113 have been utilized to characterize organosulfates in laboratory studies. FT-IR has been used to

114 measure the organosulfate functional group using peak fitting and showed that the organosulfate
115 functional group contributes up to 10% of organosulfate in the Arctic region, when inorganic
116 sulfate concentrations are considered high (Frossard et al., 2011), and 4-8% of OM in the Pacific
117 marine boundary layer, during periods of high organic and sulfate concentrations (Hawkins et al.,
118 2010). While other studies showed little to no organosulfates, likely due to low sulfate
119 concentrations in Mexico City (Liu et al., 2009) and Bakersfield, CA (Liu et al., 2012).

120 Methyl sulfate is the smallest organosulfate (Kwong et al., 2018b) and measured mostly
121 in trace amounts (Hettiyadura et al., 2017, 2015; Wang et al., 2021). However, it is commercially
122 available and therefore useful for laboratory studies. 2-Methyltetrol sulfates are tracers for
123 secondary organic aerosols (SOA) formation in atmospheric particles derived from isoprene
124 (Surratt et al., 2010; Chen et al., 2020) and one of the most abundant organosulfates measured in
125 ambient aerosol. In the eastern US, 2-MTS accounts for the highest percentage summertime
126 particulate organosulfate (11%) (Chen et al., 2021). In Centreville, AL, 2-MTS accounts for
127 more than half of organosulfates during summer of 2013 (Hettiyadura et al., 2017). In Shanghai,
128 China (summer of 2015-2016, 2018-2019), 2-MTS was the most abundant organosulfate (31%)
129 of 29 organosulfates (Wang et al., 2021).

130 The IMPROVE network is a rural particulate matter monitoring network with ~165 sites
131 across the United States (<http://vista.cira.colostate.edu/improve/>). Polytetrafluoroethylene
132 (PTFE), nylon and quartz filters are used to collect PM_{2.5} every one in three days, have a field
133 latency period of up to 7 days and are analyzed by multiple analytical techniques. PTFE filters
134 are stored at room temperature and analyzed between 3 and 12 months after collection (typically
135 6 to 9 months) for PM mass, elements and filter-based light absorption. Recently, FT-IR
136 analysis, a non-destructive method, has been performed on IMPROVE samples to reproduce

137 routinely measured compositional data (Debus et al., 2022) and measure the functional group
138 composition of the organic fraction (Ruthenburg et al., 2014, Kamruzzaman et al., 2018).
139 Organic sulfur compounds and organosulfate functional groups were not measured in these
140 studies. FT-IR analysis has also been conducted on other networks and in chamber and field
141 studies to measure organic functional groups (Boris et al., 2021, 2019; Laurent and Allen, 2004;
142 Ruthenburg et al., 2014; Yazdani et al., 2022; Russell et al., 2011). The extracts of nylon filters
143 have been analyzed by ion chromatography (IC) in the IMPROVE network for more than three
144 decades to routinely measure the inorganic ions, sulfate, nitrate, chloride and nitrite.
145 ([http://vista.cira.colostate.edu/improve/wp-content/uploads/2020/02/1_Anion-Cation-Analysis-](http://vista.cira.colostate.edu/improve/wp-content/uploads/2020/02/1_Anion-Cation-Analysis-by-Ion-Chromatography-SOP-revision-7.pdf)
146 [by-Ion-Chromatography-SOP-revision-7.pdf](http://vista.cira.colostate.edu/improve/wp-content/uploads/2020/02/1_Anion-Cation-Analysis-by-Ion-Chromatography-SOP-revision-7.pdf)). Although not routinely measured in IMPROVE
147 samples, HMS has been identified in IMPROVE samples using IC (Moch et al., 2020).
148 Organosulfates and MSA have been previously identified in extracts of IMPROVE samples
149 (Chen et al., 2021) using hydrophilic interaction liquid chromatography interfaced to
150 electrospray ionization high-resolution quadrupole time-of-flight mass spectrometry (HILIC-
151 ESI-HR-QTOFMS).

152 The goal of this paper is to assess the stability of four organic sulfur and sulfate
153 containing compounds on polytetrafluoroethylene (PTFE) filters and the suitability of Fourier-
154 transform infrared (FT-IR) spectroscopy to measure these compounds in the IMPROVE network.
155 Measuring organic sulfur compounds on a continuous basis across the US would provide a rich
156 data set to evaluate their sources, concentration, seasonality and trends over time. Four organic
157 sulfur compounds, two organosulfur compounds, methanesulfonic acid,
158 hydroxymethanesulfonate, and two organosulfates, methyl sulfate and 2-methyltetrol sulfate, are
159 evaluated. For these compounds to be measurable in IMPROVE by FT-IR, there must be

160 minimal losses or other changes to the compound during the latency period between collection
161 and analysis (3 – 12 months), and there must be minimal interferences in the spectra. To achieve
162 this goal, each compound was dissolved in solution, aerosolized and collected on PTFE filters.
163 Collected samples were weighed and analyzed by FT-IR every few days for two months or more.
164 Characterization of the FT-IR spectra as well as changes (or lack therefore of) in the mass
165 loading and spectra over time indicate the potential for the compounds to be measured by FT-IR
166 in IMPROVE samples. Filter samples, extracted for IC analysis at different time points, indicate
167 stability or chemical changes in the compound on the filter, assists with interpreting gravimetric
168 mass and FT-IR spectra changes.

169 **2 Methods**

170 Two organic sulfur (C-S) compounds, MSA and HMS, and two organosulfates (C-O-S), MS and
171 2-MTS, were selected. The four compounds were selected for this study based on following
172 three criteria. The compound 1.) has been measured in atmospheric particulate matter and is of
173 interest to the atmospheric science community, 2.) is water soluble so it can be put into solution
174 for atomization, and 3.) is available in high purity form to minimize uncertainty in mass
175 measurement. Filter samples of the organic sulfur compounds were prepared for FT-IR,
176 gravimetry, IC and ICP-OES analyses by aerosolizing each compound individually and
177 collecting it on PTFE filters (Pall Corporation, 25 mm diameter). One set of filters was
178 generated for analysis by gravimetry and FT-IR at UC Davis and another set (or sets depending
179 on what was being evaluated) of filters were generated for analysis by IC and ICP-OES at
180 Research Triangle Institute (RTI) following gravimetric analysis at UC Davis. Analyses of
181 these laboratory filter samples were performed to characterize the compound within infrared
182 spectra and to determine the stability of these compounds over time.

183 **2.1 Preparation of laboratory filter samples**

184 Three commercially available standards were used for this study: HMS sodium salt (>97%
185 purity, TCI America), MSA (100% purity, Sigma Aldrich) and MS sodium salt (100 % purity,
186 Sigma Aldrich). 2-methytetrol sulfate ammonium salt was synthesized following a published
187 method (Cui et al., 2018). Each compound was collected on PTFE filters by first preparing an
188 aqueous solution with a concentration of 0.005 M. For HMS, the solution was acidified with
189 hydrochloric acid (HCl) prior to aerosolization to obtain samples with atmospherically relevant
190 pH (pH 2), as pH plays a role in the stability of HMS. 2 μ L of 1 M HCl was added to HMS
191 solution to obtain the final volume of 200 mL for aerosolization. Aerosols were generated using
192 an atomizer (Kamruzzaman et al., 2018; Ruthenburg et al., 2014) and dried with a diffusion
193 dryer (Model 3074B Filtered Air Supply, TSI Inc., St. Paul, MN) which produces a high
194 concentration of poly-disperse submicrometer sized particles allowing for short collection times
195 and adequately representing the expected response from particles of similar size range in the
196 atmosphere. Dry particles were collected on PTFE filters (Pall Corporation, 25 mm diameter)
197 using an IMPROVE sampler with varying collection times (40 to 720 s) at a flow rate of 22.4
198 L/min.

199 **2.1.1 Gravimetric mass determination**

200 Filter mass, before and after particle collection, was measured using an ultra-
201 microbalance (XP2U, Mettler-Toledo, Columbus, OH) with 0.1 μ g sensitivity. Ionizing
202 cartridges (Staticmaster® Model 2U500, Grand Island NY) housed on a flexible stand
203 (Staticmaster® Model BF2-1000, Grand Island NY) and Haug strip (Mettler Toledo 11140160,
204 Columbus, OH) were utilized to help eliminate static for more stable, accurate measurements.
205 Prior to particle collection, the mass of a filter was determined by the average of 5 mass

206 measurements taken on separate days. Only filters that weighed within measurements precision
207 for 25 mm filters ($\pm 6 \mu\text{g}$) for the 5 measurements were used. After particle collection, filters
208 were allowed to achieve equilibrium at room temperature for 24-hrs. Filters were weighed for
209 three consecutive days in the 1st week, twice per week during 2nd – 4th weeks and once in the
210 weeks thereafter. Filters were stored at room temperature (21°C – 27 °C) and relative humidity
211 (30% \pm 10%) to mimic the storage conditions for ambient IMPROVE Teflon filters. The
212 experiment was ended when the weights were stable for a month or more.

213 **2.2 Infrared spectra collection and processing**

214 FT-IR spectra of the filter samples of each compound were collected using a Tensor II FT-IR
215 spectrometer (Bruker Optics, Billerica, MA) with a liquid nitrogen cooled mercury cadmium
216 telluride (MCT) detector over the spectral range of 4 000–400 cm^{-1} . Filters were placed in a
217 house-built sample chamber that is purged of water and CO_2 (PureGas) for 4 min before spectra
218 acquisition (Debus et al., 2019). Transmission mode measurements were made using 512 scans
219 for each filter at 4 cm^{-1} resolution and ratioed to the most recent (less than 1 h) background
220 spectrum to obtain absorbance spectra using OPUS software (Bruker Optics, Billerica,
221 MA)(Debus et al., 2019).

222 To better visualize functional groups in the organosulfur compounds and minimize the impact of
223 the PTFE scattering and absorption on the spectra, several steps were taken. Spectra were
224 baseline corrected from 1500 cm^{-1} to 500 cm^{-1} , using blank correction and smoothing spline
225 fitting (Kuzmiakova et al., 2016) The spectral region from 4000 cm^{-1} to 1500 cm^{-1} were
226 baselined using an automated version of the Kuzmiakova et al., 2016 smoothing spline process in
227 AirSpec (Reggente et al., 2019). Regions with large PTFE absorption (1300-1100 cm^{-1} , 700-600
228 cm^{-1} and 500-420 cm^{-1}) were grayed out in spectra plots and are not considered for peak

229 identification. A baseline corrected spectrum of each compound is shown in supplemental
230 material Figure S1.

231 **2.3 IC and ICP-OES sample analysis**

232 PTFE filters of each organic sulfur and organosulfate sulfur compound were generated and
233 weighed at UC Davis prior to shipping the filters cold overnight to RTI for IC and ICP-OES
234 analysis. On the day the filters arrived at RTI, the filters were extracted in 50 ml of deionized
235 water ($18\text{M}\Omega\text{ cm}^{-1}$, Millipore Milli-Q Darmstadt, Germany), sonicated for 30 minutes in an ice
236 bath and placed on a shaker table in a cold room for 8 hours prior to analysis. The PTFE filter
237 remained in the extraction vial for the duration of these experiments.

238 IC analysis was performed on Dionex Thermo Scientific ICS-3000 and ICS-6000 (Sunnyvale,
239 CA) instruments using suppression and conductivity detection. For MSA, extracts were analyzed
240 using the AS19 analytical and AG19 guard columns (anion hydroxide method) for initial
241 extraction efficiency tests and AS28 analytical and AG28 guard columns (hydroxide method) for
242 the subsequent analyses to evaluate changes over time. HMS and MS extracts were analyzed
243 with AS12A analytical and AG12A guard columns (anion carbonate method), which has been
244 shown to provide sufficient separation of HMS and sulfate, but not separation of sulfite/bisulfite
245 and HMS (Dovrou et al., 2019). MS extracts were analyzed with AS12A analytical and AG12A
246 guard columns (anion carbonate method), the same method as HMS. An IC method for analyzing
247 2-MTS had not been developed and evaluated prior to this work. Eluent concentrations and flow
248 rates were optimized for best separation of all ions of interest.

249 ICP-OES, used to measure total sulfur on the filter, was performed on a Thermo Scientific iCAP
250 7600 duo analyzer (Bremen, Germany). The ICP-OES was run in axial mode using a sprint valve
251 and data were collected at 180.731 nm. The ICP-OES system was calibrated using the sulfate

252 calibration standards and validated using the sulfate calibration verification solutions described
253 below.

254 IC and ICP-OES systems were calibrated with calibration standards prepared via serial dilutions
255 of single source stock standards using a primary source. A secondary source was used to prepare
256 calibration verification solutions to validate the instrument calibration for all compounds except
257 for 2-MTS, for which a second source standard was unavailable.

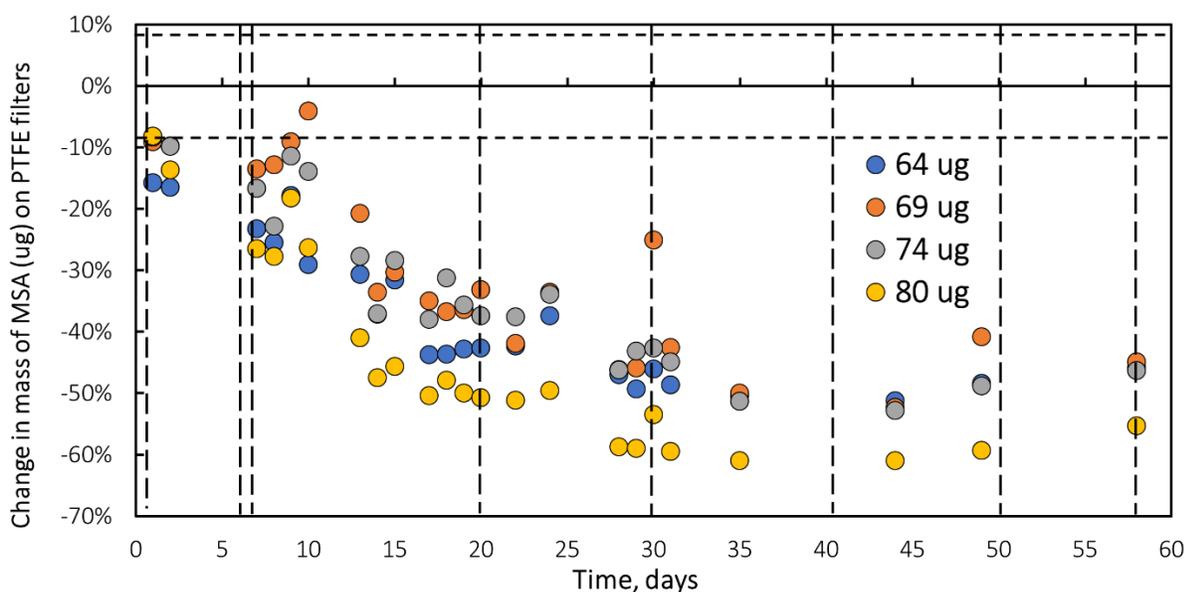
258 Primary and secondary sources of National Institute of Standards and Technology (NIST)-
259 traceable solutions were purchased and used to prepare calibration standards and calibration
260 verification solutions respectively, for sulfate analyses by both IC and ICP-OES. When NIST-
261 traceable solutions were unavailable, salts were used to prepare calibration standards and
262 calibration verification solutions for MSA, HMS, and MS. Vendor information for primary and
263 secondary sources are provided in Table S2 in supplemental material. Certified American
264 Chemical Society (ACS)-grade sodium carbonate (Na_2CO_3) obtained from Fisher Scientific
265 (Fairlawn, NJ) and sodium bicarbonate (NaHCO_3) obtained from EMD Sciences (Gibbstown,
266 NJ) were used to prepare IC eluent when using anion carbonate methods for analyses. Potassium
267 hydroxide eluent generator cartridges purchased from Thermo Scientific were used for eluent
268 preparation for analyses conducted with anion hydroxide methods. NIST-traceable, 1000 $\mu\text{g}/\text{mL}$
269 stock solutions of yttrium (Y) and cesium (Cs) obtained from High Purity Standards (Charleston,
270 SC) were used to for internal standard and ionization suppression, respectively for ICP-OES
271 measurements.

272 3 Results and Discussion

273 3.1 Methanesulfonic acid

274 3.1.1 Gravimetry

275 Mass changes, measured by gravimetry, for four methanesulfonic acid filter samples with masses
276 ranging from 64 μg to 80 μg are shown in Figure 1. Mass decreases steadily during the first
277 month to approximately 50% of the initial mass. During the second month of measurements, the
278 mass remains constant ($50 \pm 6\%$).

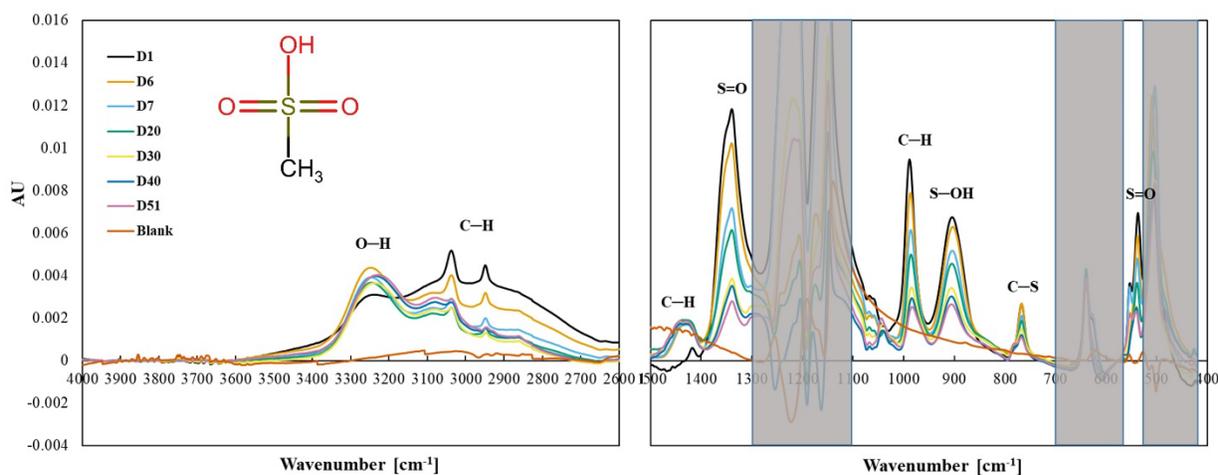


279
280 **Figure 1. Change in mass of methanesulfonic acid (MSA) collected on PTFE filters over a**
281 **2- month period under laboratory conditions (24 °C). Dotted vertical lines indicate FT-IR**
282 **analysis. Horizontal broken line indicates mass balance precision. Colors indicate initial**
283 **masses of samples.**

284 3.2 FT-IR

285 Methanesulfonic acid ($\text{CH}_3\text{SO}_3\text{H}$) is composed of a methyl group attached to a sulfonic acid
286 [$\text{S}(=\text{O})_2\text{—OH}$], via a C-S bond. Methanesulfonic acid aerosols collected on PTFE filters have

287 peaks associated with CH₃, SO₃, S-OH and C-S bonds (Figure 2). Observed peaks (Figure 2)
 288 can be ascribed to portions of the molecule based on previous FT-IR and Raman work (Lee et
 289 al., 2019; Zhong and Parker, 2022; Chackalackal and Stafford, 1966). The peaks at 1342 cm⁻¹
 290 and 536 cm⁻¹ arise from S=O bonds in MSA and are shifted compared to inorganic peaks at 1130
 291 cm⁻¹, 620 cm⁻¹ (Larkin, 2018) or organic sulfate SO₄ peaks at ~1380 cm⁻¹ (Larkin, 2018; Lin-
 292 Vien et al., 1991). The peak at 895 cm⁻¹ is attributable to S—OH (Zhong and Parker, 2022) and
 293 the peak at 766 cm⁻¹ is attributable to C-S (Lee et al., 2019). C-H peaks are observed at 3039 cm⁻¹
 294 ¹, 2951 cm⁻¹, 1414 cm⁻¹ and 987 cm⁻¹ (Chackalackal and Stafford, 1966). The broad peak at 3248
 295 cm⁻¹ is suggested to be water as (Zeng et al., 2014) showed that this peak in MSA infrared
 296 spectra increases with increasing RH. These peaks, particularly strong peaks, were similar to
 297 spectral absorbance of MSA from reference spectra (Spectral Database for Organic
 298 Compounds,SDBS, 2022), and Table S1 in Supplemental Materials compares the observed and
 299 reference peaks.



300
 301 **Figure 2. Changes in the spectra of MSA over a 2- month period, denoted by number of**
 302 **elapsed days. The shaded area indicates the absorbance regions of PTFE filter.**

303 The MSA infrared peaks of SO_3 (1342 cm^{-1}), C—H (987 cm^{-1}), S—OH (895 cm^{-1}), C-S (766 cm^{-1}), and S=O (536 cm^{-1}) decrease rapidly in the first 30 days, consistent with the decline in mass
304 during that time. The spectra suggest that MSA is volatilizing off the filter, even though this is
305 inconsistent with the low vapor pressure of MSA (Knovel - Yaws' Critical Property Data for
306 Chemical Engineers and Chemists - Table 12. Vapor Pressure - Organic Compounds, $\log P = A -$
307 $B/(T + C)$, 2022), 0.00022 mmHg at 20°C). The three spectra obtained on days 30, 41 and 51
308 show only small decreases which is mostly consistent with the lack of mass changes during those
309 days.
310
311 Not all peaks show consistent loss during the first month and little change during the second
312 month. The C-H peaks at 3039 cm^{-1} and 2942 cm^{-1} behave slightly differently, reaching stability
313 (with some minor random variability) earlier (day 20) than most other peaks. The weak peak at
314 1414 cm^{-1} (ascribed to CH_3) increased and broadened to $1445\text{ cm}^{-1} - 1400\text{ cm}^{-1}$ after the initial
315 spectra was collected and remained fairly stable for the duration of the experiment. The peak at
316 3250 cm^{-1} , increased rapidly followed by fluctuations, and then varies somewhat but remained
317 fairly consistent for the duration of the experiment, not unlike the behavior of the 1414 cm^{-1}
318 peak. One possible cause for these spectra changes in the spectra is water vapor condensing on
319 the particles after collection. MSA is hygroscopic and although it effloresces at about $\text{RH} = 50\%$,
320 (Peng and Chan, 2001; Tang, 2020; Zeng et al., 2014), Zeng et al., 2014 shows that there is some
321 water associated with the particles below $50\% \text{ RH}$, which is consistent with the FT-IR spectra
322 (Figure 3), particularly the 3248 cm^{-1} peak. If water is indeed the cause, the change in the 3248
323 cm^{-1} peak can be explained by a very low RH in the particle generation and collection system
324 (lower initial peak) and a higher RH ($\text{RH} = 30 \pm 10\%$) in the lab (increase in peak). The change

325 in the 1414 cm^{-1} peak above 1400 cm^{-1} behaves similarly and can be associated with OH
326 suggesting uptake of water after the initial FT-IR spectra was collected.
327 Another possibility is that ammonia is absorbing onto the MSA. Ammonium absorbs in both the
328 3250 cm^{-1} and 1500-1400 cm^{-1} regions (Boer et al., 2007; Zawadowicz et al., 2015) and when
329 comparing MSA spectra to ammonium sulfate spectra, they show very similar peak absorbance
330 and shapes in these two regions suggesting. This suggests that the cause of these peaks is
331 ammonium (Supplemental Material, Figure S2).
332 A third possible cause for these changes is that MSA may be fragmenting into formaldehyde
333 (CH_2O), that partitions into the gas phase, and sulfite (SO_3) (Kwong et al., 2018a). However,
334 this would show a decrease in CHC-H peaks and a shift in the SOS=O peaks, neither of which
335 are observed in the spectra. The rapid decrease in peak height during the first month and then
336 little decrease or no trend during the second month, suggests that MSA is volatilizing off the
337 filter initially, but then has a slow decline, offset by increases in water or ammonium.

338 **3.1.3 IC and ICP-OES**

339 Twenty PTFE filters with MSA (14 μg – 32 μg ; 60 μg – 144 μg) were shipped to RTI for
340 extraction and analysis by IC and ICP-OES. Each extract was analyzed by both IC and ICP-OES.
341 Recoveries of MSA from PTFE filters was $55 \pm 5\%$ for IC and $51 \pm 5\%$ for ICP-OES.
342 Calibration verification solution recoveries were $96 \pm 5\%$ for IC and $101 \pm 3\%$ for ICP-OES
343 analyses, suggesting that the lower recoveries from the PTFE filter are due to incomplete
344 extraction and/or losses occurring during shipment. To evaluate losses during shipping, six
345 samples were collected, weighed, and analyzed by FT-IR, shipped to RTI and then sent back to
346 UC Davis and analyzed with gravimetric and FT-IR analysis 9 days after initial measurements.
347 The mass loss ($19 \pm 7\%$) during this period was similar to mass loss for filters that remained in

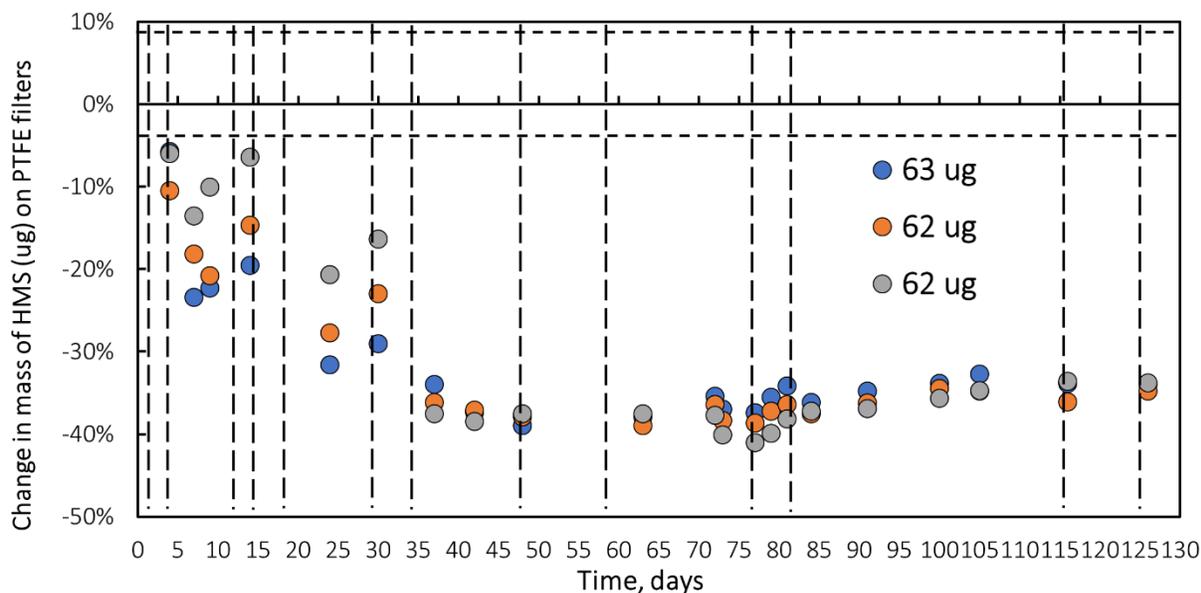
348 the UC Davis laboratory for ~ 8 days ($22 \pm 6\%$). Spectral changes in shipped filters were similar
349 to changes in spectra that occurred during the first week for filters that remained at UC Davis.
350 To evaluate changes in mass and composition on PTFE filters over time, 6 filters ($42 \mu\text{g} - 63 \mu\text{g}$)
351 were shipped to RTI and extracted on days 0, 30, and 61. These filters had consistent recoveries
352 over time of $57 \pm 6\%$ for IC and $55 \pm 6\%$ for ICP-OES. The IC results indicate that MSA did
353 not change chemically during this time period, supporting the FT-IR spectral results. The mass
354 and spectral data indicate that the lower extraction efficiency was not due to loss of MSA off the
355 filter during shipping and suggest that the limitation is extraction efficiency. Despite low
356 extraction efficiencies and difference in behavior over time from FT-IR measurement, the FT-IR,
357 gravimetry and IC results suggest that a lower-bound of MSA can be measured on PTFE filters
358 in IMPROVE by FT-IR.

359 **3.2 Hydroxymethanesulfonate**

360 **3.2.1 Gravimetry**

361 Mass changes in three hydroxymethanesulfonate sodium salt ($\text{HOCH}_2\text{SO}_3\text{Na}$) filter samples with
362 mass loadings of $\sim 62 \mu\text{g}$ per filter are shown (Figure 3). Mass decreases steadily for 1.3 months
363 to a maximum loss of 38% and then remains constant ($38 \pm 3\%$) for the rest of the experiment
364 (4.1 months). Similar results were obtained for filters weighing approximately $30 \mu\text{g}/\text{filter}$ of
365 HMS.

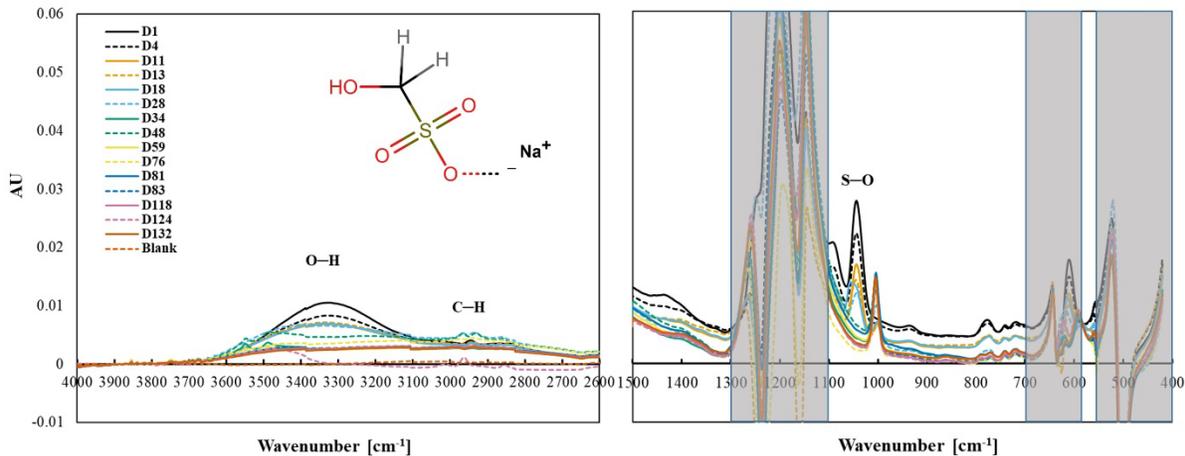
366



367
 368 **Figure 3. Mass behavior of hydroxymethanesulfonate (HMS) over 125 days under**
 369 **laboratory conditions (~24°C). Dotted vertical lines indicate FT-IR analysis. Horizontal**
 370 **broken line indicates mass balance precision.**

371 **3.2.2 FT-IR**

372 HMS ($\text{HOCH}_2\text{SO}_3\text{H}$) is a sulfonic acid compound with C-S bond, where the S bond is part of a
 373 sulfonic acid group $[\text{S}(=\text{O})_2-\text{OH}]$, and the carbon is attached to an -OH functional group,
 374 similar to MSA except with the OH functional group attached to the carbon. The chemical used
 375 in our study has a sodium cation on the sulfonic acid group. HMS aerosol collected on PTFE
 376 filters have three infrared peaks (1094 cm^{-1} , 1041 cm^{-1} and 611 cm^{-1}) between 1500 cm^{-1}
 377 , although the peak at 611 cm^{-1} is obscured by PTFE absorption, and O-H and C-H peaks
 378 between 4000 cm^{-1} and 1500 cm^{-1} (Figure 4).



379
 380 **Figure 4. Changes on the functional group frequency region of hydroxymethanesulfonate**
 381 **(HMS) over 4.1 months.**

382
 383 Observed peaks at 1094 cm^{-1} and 1041 cm^{-1} are similar to the 1080 cm^{-1} and 1040 cm^{-1} peaks
 384 identified as S-O or S=O bonds in FT-IR spectra of Na HMS (Sato et al., 1984) and of HMS
 385 (Larkin, 2018; Shurvell, 2006). A weak band at 611 cm^{-1} is like due to C-S (Lin-Vien et al.,
 386 1991; Sato et al., 1984) or S-O (Sato et al., 1984) but is obscured by PTFE absorbance. Above
 387 1500 cm^{-1} and similar to many aliphatic organic molecules, an O-H (broad peak centered near
 388 3300 cm^{-1}) and C-H (below 3000 cm^{-1}) peaks were observed (Pavia et al., 2008; Shurvell, 2006).
 389 Observed peaks are similar to spectral absorbance, although not all peaks in the reference spectra
 390 (AIST: SDBS, 2022, See Table S1 in Supplemental Material) are observed in the measured
 391 spectra. HMS has a very low vapor pressure (0.00000073 mmHg), (U.S. EPA. Comptox
 392 Chemicals Dashboard, 2022) indicating that HMS should not volatilize off the filter.

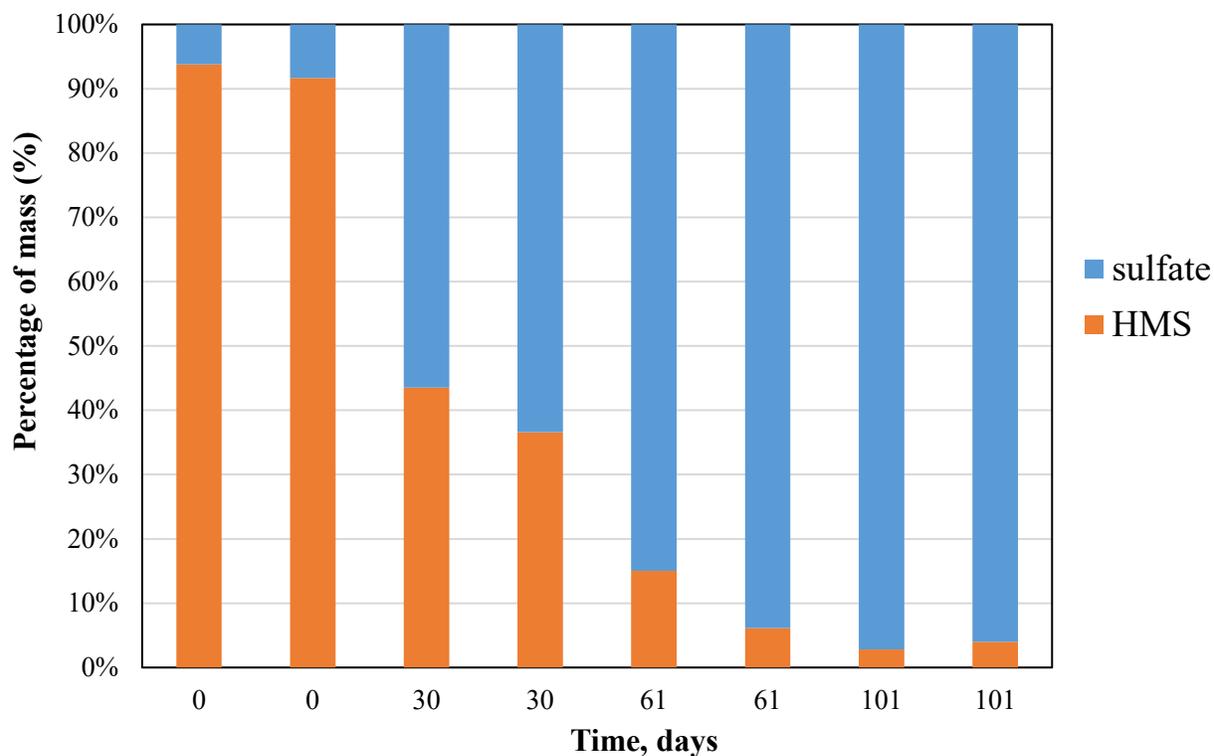
393 All peaks below 1500 cm^{-1} decrease and are no longer visible by day 34, consistent with
 394 the decline but not the extent of decline in mass as the HMS peaks are completely gone and the
 395 mass has only decreased by 40%. This behavior is most clearly observed in the peak at 1041 cm^{-1}
 396 1 , but also observed in 1094 cm^{-1} peak (Figure 4). Similar to the S-O/S=O peaks, C-H peaks

397 decline during the first 34 days and then completely disappear. The O-H peak, centered around
398 3300 cm^{-1} , disappears more slowly, but like the S-O and C-H peaks, is gone by the end of the
399 experiment. Counter balancing the loss of mass, a new peak becomes visible at 1003 cm^{-1} after
400 11 days and increases for the rest of the study. The peak at 1003 cm^{-1} is tentatively identified as
401 bisulfate (Boer et al., 2007; Krost and McClenny, 1994). The small peak centered around 3450
402 cm^{-1} becomes evident as the O-H peak disappears and may indicate the presence of condensed
403 water (Boer et al., 2007).

404 **3.2.3 IC and ICP-OES**

405 Sixteen HMS PTFE filters were analyzed by IC with recoveries of $65 \pm 4\%$ and calibration
406 verification solution recoveries of $94 \pm 5\%$. ICP-OES analysis was not performed on these
407 filters. Eight additional HMS filters ($65\text{ }\mu\text{g} - 100\text{ }\mu\text{g}$) were shipped to RTI and 2 filters were
408 extracted and analyzed by IC and ICP-OES on each of the following days 0, 30, 61, and 101.
409 The sulfur mass losses in IC (Figure S3, Supplemental Material) over the 101 days is $\sim 60\%$ loss
410 from the initial weighed mass or about 39% loss assuming a constant extraction efficiency of
411 65% for all samples and is in agreement with the 38% decrease in mass on the filter measured by
412 gravimetry. Similar results were obtained for ICP-OES. IC analysis (Figure 5) confirms the
413 samples are mostly HMS on day zero but over time the HMS is converted to sulfate (sulfate and
414 bisulfate are indistinguishable in IC), supporting assignment of the 1003 cm^{-1} infrared peak to
415 inorganic bisulfate. A small amount of the HMS may be converting to sulfate in solution or the
416 column, but the measured changes are much larger than what is expected due to that mechanism
417 alone. The small amount of HMS that is measured by IC on day 60 and 101 are near detection
418 limits for IC which corroborate the absence of HMS in the FT-IR spectra after two months. The
419 IC and FT-IR results both show a conversion of HMS into sulfate indicating that HMS is not

420 stable and cannot be quantified reliably on PTFE filters in IMPROVE by either FT-IR or IC. In
421 Moch et al. (2020), HMS did not degrade over time under cold storage conditions on nylon
422 filters from IMPROVE suggesting that storage or perhaps filter type may play an important role
423 in HMS degradation on filters.



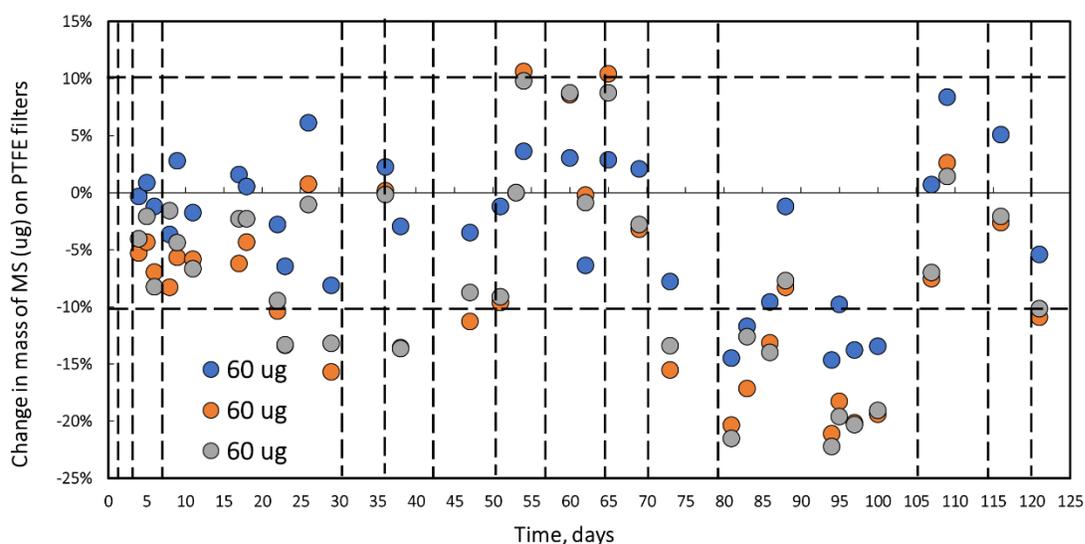
424
425 **Figure 5. Percentage of HMS and sulfate measured by IC. Eight PTFE filters were**
426 **extracted and analyzed in pairs over 101 days.**

427 3.3 Methyl sulfate

428 3.3.1 Gravimetry

429 Mass measurements of methyl sulfate salt over a 4 -month period were within measurement
430 uncertainty for the three filters loaded with approximately 30 μg of methyl sulfate (the change
431 over time was indistinguishable from zero). For the 60 μg filters (Figure 6), the first 2 months

432 and the last two weeks mass change were within measurement uncertainty ($\pm 6 \mu\text{g}$). However,
433 for about a month, between day 70 and 100, much of the data (except day 91) is outside of
434 measurement uncertainty, indicating mass loss of between 10 and 20%. During this period only
435 one spectrum (day 79) was collected, and it does not support mass loss. Day 100 spectra, and all
436 spectra collected through the end of the study, for all three 60 μg loadings, filter samples
437 confirms no changes in the MS compound. Mass and spectral data indicate stability of mostly
438 stable MS on filters under ambient laboratory condition (24°C).



439
440 **Figure 6. Mass behavior of methyl sulfate (MS) standard over 4 months. Dotted vertical**
441 **lines indicate FT-IR analysis. Horizontal broken line indicates mass balance precision for**
442 **PTFE 25- mm diameter filters.**

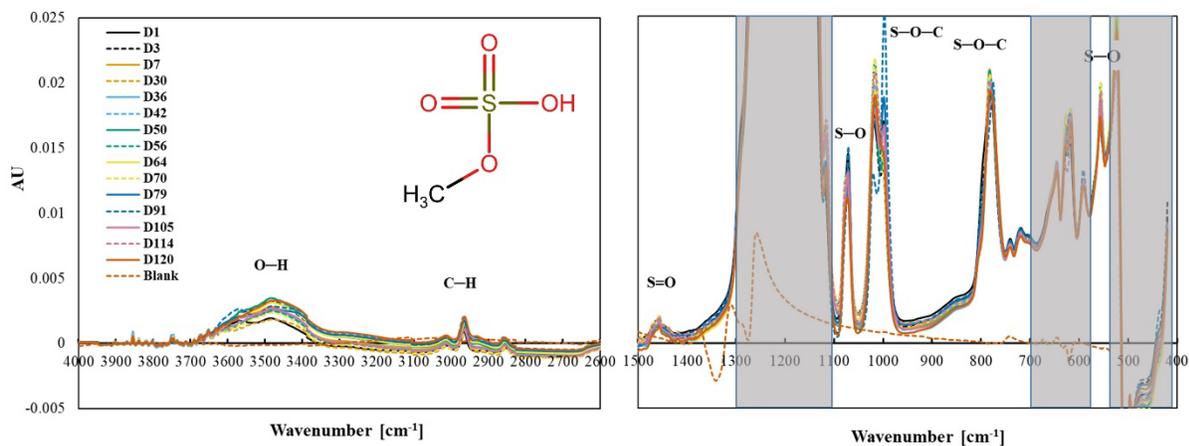
443 3.3.2 FT-IR

444 Methyl sulfate ($\text{CH}_3\text{SO}_4\text{H}$) is composed of a methyl (CH_3) group attached to a sulfate (SO_4)
445 group with C-O-S bond. The chemical used in this study is sodium methyl sulfate ($\text{CH}_3\text{SO}_4\text{Na}$).
446 Methyl sulfate aerosol collected on PTFE filters has two peaks between 4000 cm^{-1} and 2000 cm^{-1}
447 1 and eight peaks between 1500 and 500 cm^{-1} (Figure 7), similar to reference spectra in Table S1

448 (Spectral Database for Organic Compounds, SDBS, 2022). The doublets observed at 1020 cm^{-1} ,
449 1000 cm^{-1} and the nearly overlapping peaks at 795 cm^{-1} and 784 cm^{-1} are identified as S-O-C by
450 both FT-IR (Chihara, 1958; Lloyd et al., 1961; Lloyd and Dodgson, 1961; Segneanu et al., 2012;
451 Shurvell, 2006) and Raman (Okabayashi et al., 1974). The S—O peaks from sulfate are observed
452 around 1073 cm^{-1} and 591 cm^{-1} in the spectra, similar to the previous study where peaks from
453 591 cm^{-1} to 593 cm^{-1} , and at 1063 cm^{-1} (solid) and 1081 cm^{-1} (solution) were ascribed to sulfate
454 in potassium methyl sulfate (Chihara, 1958). There is a weak S=O peak at 1458 cm^{-1} (Segneanu
455 et al., 2012). The O-H group (assuming the Na ion was replaced by H for some of the
456 molecules) at 3500 cm^{-1} is also fairly weak. All collected spectra for one sample are shown, with
457 the exception of day 91 data, which appeared to be anomalous (Supplemental Material, Figure
458 S4 all spectra including day 91).

459 The stability of the spectra over 4 months suggests that MS is stable when collected on a PTFE
460 filter. There were no major or consistent changes in peaks associated with S—O—C, S—O, Na-O
461 and C—H.

462 The pattern of change in the peak height of the 3500 cm^{-1} peak does not correlate with the
463 change in mass. Day 1 and day 30 spectra have smaller O-H peak intensities prior to mass
464 decline, compared to and the O-H peak in the spectrum from day 79, when mass had decreased
465 the mass is low, is higher than both. No consistent difference in the FT-IR spectra was observed
466 on day 79 (low gravimetric mass day) or the following days (days 105 – 120) suggesting that the
467 mass loss seen in the gravimetric data is erroneous. The high melting and boiling points of MS
468 are 96 °C and 298 °C, respectively and the low vapor pressure of 0.0038 mmHg (U.S. EPA.
469 Comptox Chemicals Dashboard, 2022) indicating that MS is not volatilizing off the filter.



470
 471 **Figure 7. Changes in the spectra of MS over a 4- month period. The shaded area indicates**
 472 **the absorbance regions of PTFE filter.**

473

474 3.3.3 IC and ICP-OES

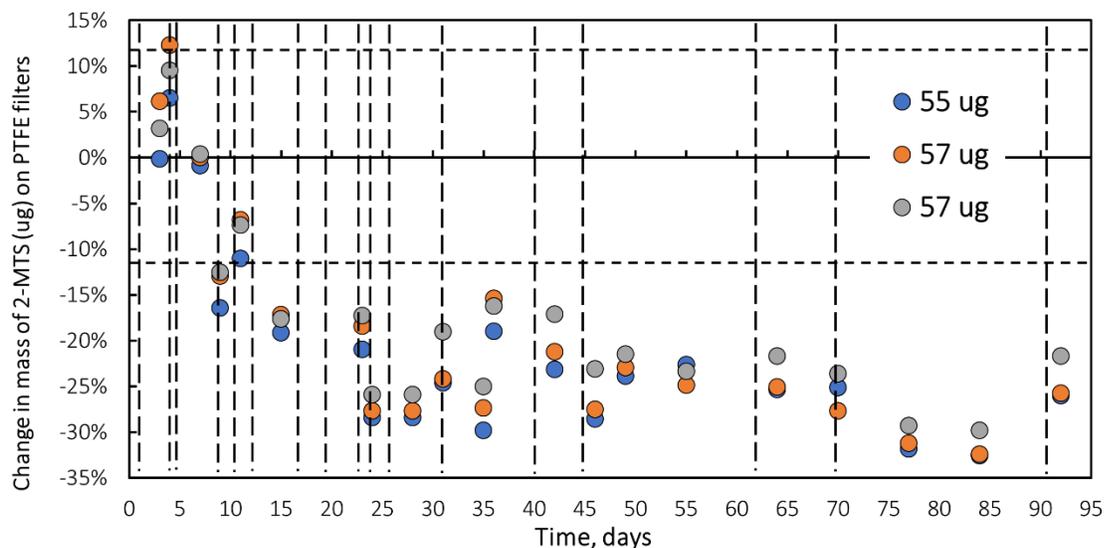
475 Sixteen PTFE samples, eight with mass loadings between 38 μg – 57 μg and eight with higher
 476 mass loadings, between 98 μg – 118 μg , were prepared for IC and ICP-OES analysis. Eight
 477 PTFE filters were extracted on day 0 and recoveries were 60 \pm 1% for IC and 87 \pm 3% for ICP-
 478 OES analyses. Recoveries of calibration verification solutions were 100 \pm 1% for IC and 97 \pm
 479 6% for ICP-OES. Additional filters were extracted on day 39 and 61 showed recoveries
 480 consistent with recoveries measured during the initial extracts, 59 \pm 2% for IC and 86 \pm 4% for
 481 ICP-OES. The overall lower mass recovery is indicative that not all MS is extracted from the
 482 filter and the lower recoveries by IC compared to ICP-OES, suggests that MS is converting to
 483 another sulfur compound in solution. The consistency in recoveries over time indicates stability
 484 of MS on PTFE over this time period, which is in agreement with the mass stability of as shown
 485 by gravimetric over the course of the experiment. Unfortunately, filters were not extracted during
 486 the time period when the gravimetric results show a small deviation from stability.

487 MS is stable on a PTFE filter as indicated by gravimetry, FT-IR and IC suggesting that at least
488 some atmospherically relevant organosulfates can be measured on PTFE filters in IMPROVE by
489 FT-IR.

490 3.4 2-Methyltetrol sulfate

491 3.4.1 Gravimetry

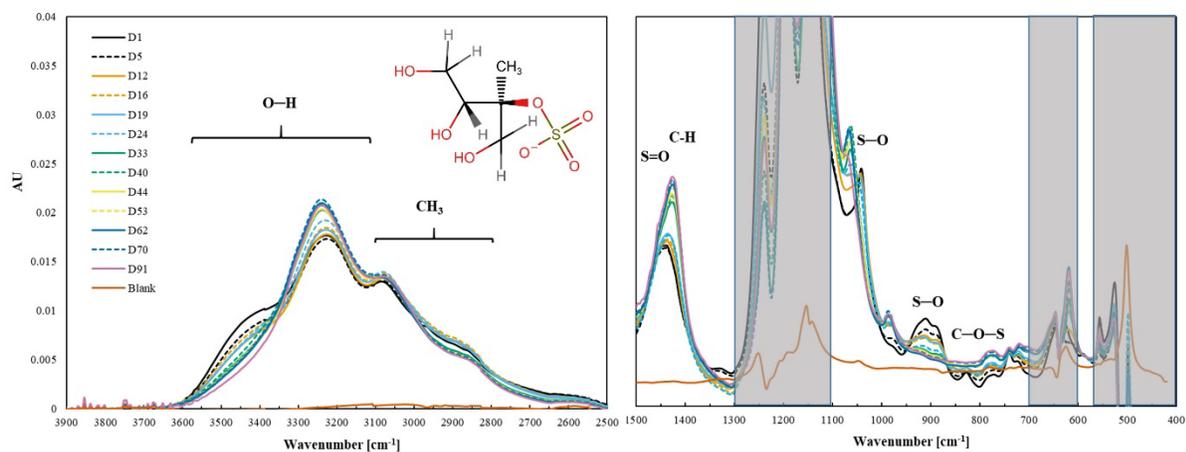
492 Mass changes in three 2-methyltetrol sulfate ($C_5H_{11}SO_7$) filter samples with concentrations from
493 55 μg to 57 μg are shown in Figure 8. Mass decreases during the first 23 days to 73% of the
494 initial mass ($25\% \pm 2\%$). No additional mass loss was observed after 25 days. Mass behavior
495 indicates a loss of 2-MTS on filters under ambient laboratory condition (24°C). Similar results
496 were obtained for filters weighing approximately 40 μg .



497
498 **Figure 8. Change in mass of 2-MTS collected on PTFE filters over a 3-month period under**
499 **laboratory conditions. Dotted vertical lines indicate FT-IR analysis. Horizontal broken red**
500 **line indicates mass balance precision.**

501 3.4.2 FT-IR

502 2-Methyltetrol sulfate ($C_5H_{11}SO_7$) is a branched compound with 3 units of $-OH$, one methyl
503 group and a sulfate group. Like methyl sulfate, 2-methyltetrol sulfate is an organosulfate and has
504 a C-O-S bond. 2-MTS collected on PTFE filters has broad, organic-related peaks between 4000
505 cm^{-1} and 2500 cm^{-1} and four peaks between 1500 and 500 cm^{-1} (Figure 9). Observed peaks can
506 be ascribed to functional groups with the molecule based on previous FT-IR and Raman work
507 (Lloyd et al., 1961; Lloyd and Dodgson, 1961; Bondy et al., 2018; Fankhauser et al., 2022).



508
509 **Figure 9. Changes in the spectra of 2-MTS over a 3-month period. The shaded area**
510 **indicates the absorbance regions of PTFE filter.**

511 The observed peak at 1041 cm^{-1} is ascribed to S-O stretch (Bondy et al., 2018; Fankhauser et al.,
512 2022), as it is similar to the peak of MS at 1050 cm^{-1} (Lloyd et al., 1961; Lloyd and Dodgson,
513 1961). The doublet at 908 cm^{-1} and 898 cm^{-1} correspond to symmetric and asymmetric stretch of
514 S-O of 2-MTS (Fankhauser et al., 2022). The weak peak at 827 cm^{-1} is tentatively assigned to C-
515 O-S stretch based on Raman spectra of 3-MTS (Bondy et al., 2018). The peak at 1446 cm^{-1} is
516 tentatively assigned to asymmetric S=O stretch based on density functional theory of FT-IR
517 spectra of 2-MTS (Fankhauser et al., 2022) and the assignment of S=O to peak at 1448 cm^{-1} in
518 methyl sulfate (Bondy et al., 2018). However, this peak was suggested to be due to C-H using

519 density functional theory of Raman spectra of 2-methyltetrol sulfates (Bondy et al., 2018) and
520 commonly ascribed to CH₂ in organic molecules at ~1465 cm⁻¹ (Pavia et al., 2008). We decided
521 to assign this to S=O because the sulfate related peaks are strong and the C-H peaks are very
522 weak in this molecule. In the higher frequency region, the very subtle peak at 2879 cm⁻¹ is
523 ascribed to C-H stretch and the large broad peaks at 3065 and cm⁻¹, 3210 cm⁻¹ and shoulder at
524 3426 cm⁻¹ may be attributable to –OH stretch (Larkin, 2018; Shurvell, 2006; Bondy et al., 2018;
525 Fankhauser et al., 2022) or inorganic –NH stretch (Boer et al., 2007; Larkin, 2018). Our tentative
526 peaks assignments of ammonium at 3210 cm⁻¹ and 3065 cm⁻¹ and OH at 3426 cm⁻¹ are based on
527 changes in spectra discussed below.

528 Most of 2-MTS infrared peaks decreased and disappeared over time, consistent with the
529 decline in mass. The S-O peaks at 1041 cm⁻¹, 908 cm⁻¹ and 898 cm⁻¹ decreased and were gone
530 by day 40. The weak C-O-S peak at 827 cm⁻¹ disappeared by day 26. The C-H region at 3000-
531 2800 cm⁻¹ decreased, similar to the changes in S-O-C peaks. However, some peaks shifted or
532 new peaks formed. New peaks (or possibly shifts from 1041 cm⁻¹ and 908 cm⁻¹ doublet) at 1066
533 cm⁻¹ and at 987 cm⁻¹ appear on day 40 and then increase slightly over time. The S=O peak at
534 1446 cm⁻¹ behaves differently from other 2-MTS peaks and increases slightly until day 33 when
535 the peak height increased significantly and shifted to 1420 cm⁻¹ where it continues to grow
536 increase over time. Two peaks at 3210 cm⁻¹ and 3065 cm⁻¹ increased whereas the region around
537 3450-3550 cm⁻¹ decreased, indicating that the shoulder at 3450 cm⁻¹ arises from a different bond
538 than the peaks at 3210 cm⁻¹ and 3065 cm⁻¹.

539 The changes in the spectra indicate a change in the chemical composition on the filter. 2-MTS is
540 no longer present by mid-way through the experiment as evidenced by the disappearance of S-O,
541 C-O-S, C-H and O-H peaks. This result is supported by a study of ambient 2-MTS that showed

542 that the atmospheric lifetime of 2-MTS is to be about 16 days (Chen et al., 2020). Intermediate
543 oxidation products of 2-MTS transformation include other organosulfates such as 2-
544 methylglyceric acid organosulfate ($C_4H_7SO_7$, MGOS) and glycolic acid organosulfate ($C_2H_3SO_6$,
545 GAS) (Zhao et al., 2020; Wei et al., 2020) and the final product is likely inorganic $(NH_4)_2SO_4$
546 (Harrill, 2020; Zhao et al., 2020). Raman spectra of MGOS (Bondy et al., 2018) indicate that the
547 spectra at the end of the experiment could be 2-MGOS. The increased peak at 3210 cm^{-1} and
548 3065 cm^{-1} could be associated with $-OH$ stretch and CH_3 asymmetric stretch of 2-MGOS and the
549 1420 cm^{-1} (C-H), 1066 cm^{-1} (S-O) and 987 cm^{-1} (SO_4^{2-}) match the final spectra well. The
550 carbonyl group in 2-MGOS is very weak in the Raman spectra and indistinguishable from a
551 blank filter in the final FT-IR spectra in this project. Alternatively, the final spectra could be
552 ammonium sulfate as indicated by peaks at 3210 cm^{-1} , 3065 cm^{-1} and 1420 cm^{-1} which are
553 indicative of ammonium, however, the 1066 cm^{-1} and 987 cm^{-1} peaks are lower than typically
554 observed for inorganic ammonium sulfate (Boer et al., 2007; Zawadowicz et al., 2015)
555 suggesting this is not inorganic sulfate. There are limited FT-IR or Raman spectra of the many
556 oxidation products of 2-MTS so definitive identification is not possible and it is likely that there
557 is a mixture of oxidation products present on the filter.

558 To further evaluate the possible compounds on the filter at the end of the experiment, mass loss
559 calculations were performed using the molecular weight of 2-MTS and each product (MGOS,
560 GAS and ammonium sulfate). If 2-MTS is converted completely to MGOS, the mass loss would
561 be only 7%. If 2-MTS is converted completely to GAS, the mass loss would be 28% and if 2-
562 MTS is converted completely to ammonium sulfate the mass loss would be 39%. These values
563 span the observed mass loss at 25% which suggesting that the compounds on the filter are of an
564 intermediate type and still in the organosulfate form not inorganic ammonium sulfate.

565 **3.4.3 IC and ICP-OES**

566 Sixteen PTFE samples, eight with mass loading between 11– 28 μg and eight with higher mass
567 loadings, between 37 – 46 μg were extracted immediately upon receipt at RTI. The IC showed
568 low sensitivity to 2-MTS detection and results are not reported so no additional information was
569 available about what compounds the 2-MTS may have change into. Given that ICP-OES
570 measures sulfur and not individual compounds, the results from this method do not provide
571 insight into chemical conversions on the filter but are briefly discussed in the supplemental
572 material (Figure S5).

573 **4 Conclusions**

574 The stability and therefore potential for FT-IR to measure organosulfur and
575 organosulfates collected on PTFE filters varies by compound. MS has the highest potential to be
576 measured on PTFE filters in IMPROVE samples by FT-IR, due to its minimal mass change and
577 no spectral changes. Consistent recoveries by IC and ICP-OES over multiple months of analysis
578 support the conclusion that MS is stable on the filters. Consistent results from analysis at UC
579 Davis and RTI suggest robustness to storage, shipping and handling conditions. MS is one of
580 many organosulfates observed in the atmosphere and not necessarily representative of
581 organosulfates in general as indicated by 2-MTS. Gravimetric mass suggests some (30%) mass
582 loss from 2-MTS samples on the PTFE filter, over a three-month period. FT-IR suggests that 2-
583 MTS is unstable on PTFE filter and changing into different compound(s), likely still an
584 organosulfate. FT-IR and gravimetry show that MSA can be measured from PTFE filters but
585 due to volatility off the filters a lower bound of MSA is measured (i.e., less than the amount of
586 MSA in the atmosphere). IC further confirmed that MSA did not chemically change while on the
587 filter. Infrared peaks in HMS spectra mid-way through the experiment indicate that HMS is not

588 stable on PTFE filters and likely converts to bisulfate. IC indicates that HMS is changing to
589 (bi)sulfate over time. Further investigations of measurements by FT-IR on PTFE of other
590 organosulfates are warranted to evaluate the extent to which the organosulfate functional group
591 can be quantified from IMPROVE PTFE filters. Further work to determine the stability and
592 ability to measure these compounds in aerosol mixtures as found in ambient samples is needed
593 before confidently using FT-IR on IMPROVE samples to measure organic sulfur compounds.

594 **5 Data Availability**

595 Data is available at <https://doi.org/10.25338/B8BH14> (Anunciado et al., 2023).

596 **6 Author Contributions**

597 Conceptualization: AMD, TD; Formal Analysis: MBA, TD, ST, AMD; Funding: AMD, TD;
598 Investigation: MBA, MDeB, LH, KL; Project Administration: AMD; Supervision: AMD, TD;
599 Writing First Draft: MBA, TD; Writing Revisions and Editing: MBA, ST, TD, MDeB, LH, KL
600 AMD; software: ST, MA; Visualization: MA

601

602 **7 Competing Interests**

603 The contact author has declared that none of the authors has any competing interests.

604

605 **8 Acknowledgements**

606 AMD and MA wish to acknowledge funding from Research Triangle Institute (Agreement No.
607 62997) and the Interagency Monitoring of Protected Visual Environments (Grant No.
608 P21AC11294). The authors would like to acknowledge Jason Surratt and his team for providing
609 the 2-MTS compound used in this study.

610 **9 References**

- 611 Spectral Database for Organic Compounds, SDBS: [https://sdb.sdb.aist.go.jp/sdb/cgi-](https://sdb.sdb.aist.go.jp/sdb/cgi-bin/direct_frame_top.cgi)
612 [bin/direct_frame_top.cgi](https://sdb.sdb.aist.go.jp/sdb/cgi-bin/direct_frame_top.cgi), last access: 22 February 2022.
- 613 Allen, A. G., Dick, A. L., and Davison, B. M.: Sources of atmospheric methanesulphonate, non-sea-salt
614 sulphate, nitrate and related species over the temperate South Pacific, *Atmospheric Environment*, 31,
615 191–205, [https://doi.org/10.1016/1352-2310\(96\)00194-X](https://doi.org/10.1016/1352-2310(96)00194-X), 1997.
- 616 Allen, A. G., Davison, B. M., James, J. D., Robertson, L., Harrison, R. M., and Hewitt, C. N.: Influence of
617 Transport over a Mountain Ridge on the Chemical Composition of Marine Aerosols during the ACE-2
618 Hillcloud Experiment, *Journal of Atmospheric Chemistry*, 41, 83–107,
619 <https://doi.org/10.1023/A:1013868729960>, 2002.
- 620 Amore, A., Giardi, F., Becagli, S., Caiazzo, L., Mazzola, M., Severi, M., and Traversi, R.: Source
621 apportionment of sulphate in the High Arctic by a 10 yr-long record from Gruebadet Observatory (Ny-
622 Ålesund, Svalbard Islands), *Atmospheric Environment*, 270, 118890,
623 <https://doi.org/10.1016/j.atmosenv.2021.118890>, 2022.
- 624 Aneja, V. P. and Cooper, W. J.: Biogenic Sulfur Emissions, in: *Biogenic Sulfur in the Environment*, vol. 393,
625 American Chemical Society, 2–13, <https://doi.org/10.1021/bk-1989-0393.ch001>, 1989.
- 626 Barnes, I., Becker, K. H., and Mihalopoulos, N.: An FTIR product study of the photooxidation of dimethyl
627 disulfide, *Journal of Atmospheric Chemistry*, 18, 267–289, <https://doi.org/10.1007/BF00696783>, 1994.
- 628 Bates, T. S., Lamb, B. K., Guenther, A., Dignon, J., and Stoiber, R. E.: Sulfur emissions to the atmosphere
629 from natural sources, *Journal of Atmospheric Chemistry*, 14, 315–337,
630 <https://doi.org/10.1007/BF00115242>, 1992.
- 631 Becagli, S., Lazzara, L., Fani, F., Marchese, C., Traversi, R., Severi, M., di Sarra, A., Sferlazzo, D.,
632 Piacentino, S., Bommarito, C., Dayan, U., and Udisti, R.: Relationship between methanesulfonate (MS-
633 in atmospheric particulate and remotely sensed phytoplankton activity in oligo-mesotrophic central
634 Mediterranean Sea, *Atmospheric Environment*, 79, 681–688,
635 <https://doi.org/10.1016/j.atmosenv.2013.07.032>, 2013.
- 636 Boer, G. J., Sokolik, I. N., and Martin, S. T.: Infrared optical constants of aqueous sulfate–nitrate–
637 ammonium multi-component tropospheric aerosols from attenuated total reflectance measurements—
638 Part I: Results and analysis of spectral absorbing features, *Journal of Quantitative Spectroscopy and*
639 *Radiative Transfer*, 108, 17–38, <https://doi.org/10.1016/j.jqsrt.2007.02.017>, 2007.
- 640 Bondy, A. L., Craig, R. L., Zhang, Z., Gold, A., Surratt, J. D., and Ault, A. P.: Isoprene-Derived
641 Organosulfates: Vibrational Mode Analysis by Raman Spectroscopy, Acidity-Dependent Spectral Modes,
642 and Observation in Individual Atmospheric Particles, *The Journal of Physical Chemistry A*, 122, 303–315,
643 <https://doi.org/10.1021/acs.jpca.7b10587>, 2018.
- 644 Boris, A. J., Takahama, S., Weakley, A. T., Debus, B. M., Fredrickson, C. D., Esparza-Sanchez, M., Burki, C.,
645 Reggente, M., Shaw, S. L., Edgerton, E. S., and Dillner, A. M.: Quantifying organic matter and functional
646 groups in particulate matter filter samples from the southeastern United States – Part 1: Methods,

647 Atmospheric Measurement Techniques, 12, 5391–5415, <https://doi.org/10.5194/amt-12-5391-2019>,
648 2019.

649 Boris, A. J., Takahama, S., Weakley, A. T., Debus, B. M., Shaw, S. L., Edgerton, E. S., Joo, T., Ng, N. L., and
650 Dillner, A. M.: Quantifying organic matter and functional groups in particulate matter filter samples from
651 the southeastern United States – Part 2: Spatiotemporal trends, Atmospheric Measurement Techniques,
652 14, 4355–4374, <https://doi.org/10.5194/amt-14-4355-2021>, 2021.

653 Campbell, J. R., Battaglia, M. Jr., Dingilian, K., Cesler-Maloney, M., St Clair, J. M., Hanisco, T. F., Robinson,
654 E., DeCarlo, P., Simpson, W., Nenes, A., Weber, R. J., and Mao, J.: Source and Chemistry of
655 Hydroxymethanesulfonate (HMS) in Fairbanks, Alaska, Environmental Science & Technology, 56, 7657–
656 7667, <https://doi.org/10.1021/acs.est.2c00410>, 2022.

657 Chackalackal, S. M. and Stafford, F. E.: Infrared Spectra of Methane-, Fluoro-, and Chlorosulfonic Acids,
658 Journal of the American Chemical Society, 88, 4815–4819, <https://doi.org/10.1021/ja00973a010>, 1966.

659 Chapman, E. G., Barinaga, C. J., Udseth, H. R., and Smith, R. D.: Confirmation and quantitation of
660 hydroxymethanesulfonate in precipitation by electrospray ionization-tandem mass spectrometry,
661 Atmospheric Environment. Part A. General Topics, 24, 2951–2957, [https://doi.org/10.1016/0960-
662 1686\(90\)90475-3](https://doi.org/10.1016/0960-1686(90)90475-3), 1990.

663 Chen, C., Zhang, Z., Wei, L., Qiu, Y., Xu, W., Song, S., Sun, J., Li, Z., Chen, Y., Ma, N., Xu, W., Pan, X., Fu, P.,
664 and Sun, Y.: The importance of hydroxymethanesulfonate (HMS) in winter haze episodes in North China
665 Plain, Environmental Research, 211, 113093, <https://doi.org/10.1016/j.envres.2022.113093>, 2022.

666 Chen, Y., Zhang, Y., Lambe, A. T., Xu, R., Lei, Z., Olson, N. E., Zhang, Z., Szalkowski, T., Cui, T., Vizuete, W.,
667 Gold, A., Turpin, B. J., Ault, A. P., Chan, M. N., and Surratt, J. D.: Heterogeneous Hydroxyl Radical
668 Oxidation of Isoprene-Epoxydiol-Derived Methyltetrol Sulfates: Plausible Formation Mechanisms of
669 Previously Unexplained Organosulfates in Ambient Fine Aerosols, Environmental Science & Technology
670 Letters, <https://doi.org/10.1021/acs.estlett.0c00276>, 2020.

671 Chen, Y., Dombek, T., Hand, J., Zhang, Z., Gold, A., Ault, A. P., Levine, K. E., and Surratt, J. D.: Seasonal
672 Contribution of Isoprene-Derived Organosulfates to Total Water-Soluble Fine Particulate Organic Sulfur
673 in the United States, ACS Earth and Space Chemistry, 5, 2419–2432,
674 <https://doi.org/10.1021/acsearthspacechem.1c00102>, 2021.

675 Chihara, G.: Measurement of Infrared Absorption Spectra by Absorption on Japanese Hand-made Paper
676 and Its Application to Paper Chromatography., Chemical & Pharmaceutical Bulletin, 6, 143–147,
677 <https://doi.org/10.1248/cpb.6.143>, 1958.

678 Debus, B., Takahama, S., Weakley, A. T., Seibert, K., and Dillner, A. M.: Long-Term Strategy for Assessing
679 Carbonaceous Particulate Matter Concentrations from Multiple Fourier Transform Infrared (FT-IR)
680 Instruments: Influence of Spectral Dissimilarities on Multivariate Calibration Performance, Applied
681 Spectroscopy, 73, 271–283, <https://doi.org/10.1177/0003702818804574>, 2019.

682 Debus, B., Weakley, A. T., Takahama, S., George, K. M., Amiri-Farahani, A., Schichtel, B., Copeland, S.,
683 Wexler, A. S., and Dillner, A. M.: Quantification of major particulate matter species from a single filter
684 type using infrared spectroscopy – application to a large-scale monitoring network, Atmospheric
685 Measurement Techniques, 15, 2685–2702, <https://doi.org/10.5194/amt-15-2685-2022>, 2022.

686 Decesari, S., Facchini, M. C., Fuzzi, S., and Tagliavini, E.: Characterization of water-soluble organic
687 compounds in atmospheric aerosol: A new approach, *Journal of Geophysical Research: Atmospheres*,
688 105, 1481–1489, <https://doi.org/10.1029/1999JD900950>, 2000.

689 Dovrou, E., Lim, C. Y., Canagaratna, M. R., Kroll, J. H., Worsnop, D. R., and Keutsch, F. N.: Measurement
690 techniques for identifying and quantifying hydroxymethanesulfonate (HMS) in an aqueous matrix and
691 particulate matter using aerosol mass spectrometry and ion chromatography, *Atmospheric
692 Measurement Techniques*, 12, 5303–5315, <https://doi.org/10.5194/amt-12-5303-2019>, 2019.

693 Fankhauser, A. M., Lei, Z., Daley, K. R., Xiao, Y., Zhang, Z., Gold, A., Ault, B. S., Surratt, J. D., and Ault, A.
694 P.: Acidity-Dependent Atmospheric Organosulfate Structures and Spectra: Exploration of Protonation
695 State Effects via Raman and Infrared Spectroscopies Combined with Density Functional Theory, *The
696 Journal of Physical Chemistry A*, 126, 5974–5984, <https://doi.org/10.1021/acs.jpca.2c04548>, 2022.

697 Frossard, A. A., Shaw, P. M., Russell, L. M., Kroll, J. H., Canagaratna, M. R., Worsnop, D. R., Quinn, P. K.,
698 and Bates, T. S.: Springtime Arctic haze contributions of submicron organic particles from European and
699 Asian combustion sources, *Journal of Geophysical Research: Atmospheres*, 116,
700 <https://doi.org/10.1029/2010JD015178>, 2011.

701 von Glasow, R. and Crutzen, P. J.: Model study of multiphase DMS oxidation with a focus on halogens,
702 *Atmospheric Chemistry and Physics*, 4, 589–608, <https://doi.org/10.5194/acp-4-589-2004>, 2004.

703 Grübler, A.: A Review of Global and Regional Sulfur Emission Scenarios, Mitigation and Adaptation
704 Strategies for Global Change, 3, 383–418, <https://doi.org/10.1023/A:1009651624257>, 1998.

705 Hansen, A. M. K., Kristensen, K., Nguyen, Q. T., Zare, A., Cozzi, F., Nøjgaard, J. K., Skov, H., Brandt, J.,
706 Christensen, J. H., Ström, J., Tunved, P., Krejci, R., and Glasius, M.: Organosulfates and organic acids in
707 Arctic aerosols: speciation, annual variation and concentration levels, *Atmospheric Chemistry & Physics*,
708 14, 7807–7823, <https://doi.org/10.5194/acp-14-7807-2014>, 2014.

709 Harrill, A. J.: Aqueous-Phase Processing of 2-Methyltetrol Sulfates by Hydroxyl Radical Oxidation in Fog
710 and Cloud Water Mimics: Implications for Isoprene-Derived Secondary Organic Aerosol, 2020.

711 Hawkins, L. N., Russell, L. M., Covert, D. S., Quinn, P. K., and Bates, T. S.: Carboxylic acids, sulfates, and
712 organosulfates in processed continental organic aerosol over the southeast Pacific Ocean during
713 VOCALS-REx 2008, *Journal of Geophysical Research: Atmospheres*, 115,
714 <https://doi.org/10.1029/2009JD013276>, 2010.

715 Hettiyadura, A. P. S., Stone, E. A., Kundu, S., Baker, Z., Geddes, E., Richards, K., and Humphry, T.:
716 Determination of atmospheric organosulfates using HILIC chromatography with MS detection,
717 *Atmospheric Measurement Techniques*, 8, 2347–2358, <https://doi.org/10.5194/amt-8-2347-2015>, 2015.

718 Hettiyadura, A. P. S., Jayarathne, T., Baumann, K., Goldstein, A. H., de Gouw, J. A., Koss, A., Keutsch, F.
719 N., Skog, K., and Stone, E. A.: Qualitative and quantitative analysis of atmospheric organosulfates in
720 Centreville, Alabama, *Atmospheric Chemistry and Physics*, 17, 1343–1359, <https://doi.org/10.5194/acp-17-1343-2017>, 2017.

722 Hoffmann, E. H., Tilgner, A., Schrödner, R., Bräuer, P., Wolke, R., and Herrmann, H.: An advanced
723 modeling study on the impacts and atmospheric implications of multiphase dimethyl sulfide chemistry,
724 *Proceedings of the National Academy of Sciences*, 113, 11776–11781, 2016.

725 Kamruzzaman, M., Takahama, S., and Dillner, A. M.: Quantification of amine functional groups and their
726 influence on OM/OC in the IMPROVE network, *Atmospheric Environment*, 172, 124–132,
727 <https://doi.org/10.1016/j.atmosenv.2017.10.053>, 2018.

728 Knovel - Yaws' Critical Property Data for Chemical Engineers and Chemists - Table 12. Vapor Pressure -
729 Organic Compounds, $\log P = A - B/(T + C)$:
730 [https://app.knovel.com/kn/resources/kt009ZN2S3/kpYCPDCECD/eptble/itable?b-
731 =&columns=1%2C2%2C3%2C6%2C4%2C5%2C10%2C11%2C13%2C14%2C12%2C7%2C8%2C9](https://app.knovel.com/kn/resources/kt009ZN2S3/kpYCPDCECD/eptble/itable?b=&columns=1%2C2%2C3%2C6%2C4%2C5%2C10%2C11%2C13%2C14%2C12%2C7%2C8%2C9), last
732 access: 7 September 2022.

733 Krost, K. J. and McClenny, W. A.: FT-IR Transmission Spectroscopy for Quantitation of Ammonium
734 Bisulfate in Fine-Particulate Matter Collected on Teflon® Filters, *Applied Spectroscopy*, 48, 702–705,
735 1994.

736 Kuzmiakova, A., Dillner, A. M., and Takahama, S.: An automated baseline correction protocol for infrared
737 spectra of atmospheric aerosols collected on polytetrafluoroethylene (Teflon) filters, *Atmospheric
738 Measurement Techniques*, 9, 2615–2631, <https://doi.org/10.5194/amt-9-2615-2016>, 2016.

739 Kwong, K. C., Chim, M. M., Hoffmann, E. H., Tilgner, A., Herrmann, H., Davies, J. F., Wilson, K. R., and
740 Chan, M. N.: Chemical Transformation of Methanesulfonic Acid and Sodium Methanesulfonate through
741 Heterogeneous OH Oxidation, *ACS Earth and Space Chemistry*, 2, 895–903,
742 <https://doi.org/10.1021/acsearthspacechem.8b00072>, 2018a.

743 Kwong, K. C., Chim, M. M., Davies, J. F., Wilson, K. R., and Chan, M. N.: Importance of sulfate radical
744 anion formation and chemistry in heterogeneous OH oxidation of sodium methyl sulfate, the smallest
745 organosulfate, *Atmospheric Chemistry and Physics (Online)*, 18, [https://doi.org/10.5194/acp-18-2809-
746 2018](https://doi.org/10.5194/acp-18-2809-2018), 2018b.

747 Larkin, P. J.: Chapter 6 - IR and Raman Spectra—Structure Correlations: Characteristic Group Frequencies,
748 in: *Infrared and Raman Spectroscopy (Second Edition)*, edited by: Larkin, P. J., Elsevier, 85–134,
749 <https://doi.org/10.1016/B978-0-12-804162-8.00006-9>, 2018.

750 Laurent, J.-P. and Allen, D. T.: Size Distributions of Organic Functional Groups in Ambient Aerosol
751 Collected in Houston, Texas Special Issue of *Aerosol Science and Technology* on Findings from the Fine
752 Particulate Matter Supersites Program, *Aerosol Science and Technology*, 38, 82–91,
753 <https://doi.org/10.1080/02786820390229561>, 2004.

754 Lee, J.-K., Lee, J.-S., Ahn, Y.-S., and Kang, G.-H.: Restoring the Reactivity of Organic Acid Solution Used for
755 Silver Recovery from Solar Cells by Fractional Distillation, *Sustainability*, 11, 3659,
756 <https://doi.org/10.3390/su11133659>, 2019.

757 Lee, S.-H., Murphy, D. M., Thomson, D. S., and Middlebrook, A. M.: Nitrate and oxidized organic ions in
758 single particle mass spectra during the 1999 Atlanta Supersite Project, *Journal of Geophysical Research:
759 Atmospheres*, 108, SOS 5-1-SOS 5-8, <https://doi.org/10.1029/2001JD001455>, 2003.

760 Lin-Vien, D., Colthup, N. B., Fateley, W. G., and Grasselli, J. G.: CHAPTER 14 - Organic Sulfur Compounds,
761 in: *The Handbook of Infrared and Raman Characteristic Frequencies of Organic Molecules*, edited by:
762 Lin-Vien, D., Colthup, N. B., Fateley, W. G., and Grasselli, J. G., Academic Press, San Diego, 225–250,
763 <https://doi.org/10.1016/B978-0-08-057116-4.50020-1>, 1991.

764 Liu, S., Takahama, S., Russell, L. M., Gilardoni, S., and Baumgardner, D.: Oxygenated organic functional
765 groups and their sources in single and submicron organic particles in MILAGRO 2006 campaign,
766 *Atmospheric Chemistry & Physics*, 9, 6849–6863, <https://doi.org/10.5194/acp-9-6849-2009>, 2009.

767 Liu, S., Ahlm, L., Day, D. A., Russell, L. M., Zhao, Y., Gentner, D. R., Weber, R. J., Goldstein, A. H., Jaoui,
768 M., Offenberg, J. H., Kleindienst, T. E., Rubitschun, C., Surratt, J. D., Sheesley, R. J., and Scheller, S.:
769 Secondary organic aerosol formation from fossil fuel sources contribute majority of summertime organic
770 mass at Bakersfield, *Journal of Geophysical Research: Atmospheres*, 117,
771 <https://doi.org/10.1029/2012JD018170>, 2012.

772 Lloyd, A. G. and Dodgson, K. S.: Infrared studies on sulphate esters. II. Monosaccharide sulphates,
773 *Biochim Biophys Acta*, 46, 116–120, [https://doi.org/10.1016/0006-3002\(61\)90653-9](https://doi.org/10.1016/0006-3002(61)90653-9), 1961.

774 Lloyd, A. G., Dodgson, K. S., Price, R. G., and Rose, F. A.: Infrared studies on sulphate esters. I.
775 Polysaccharide sulphates, *Biochim Biophys Acta*, 46, 108–115, [https://doi.org/10.1016/0006-3002\(61\)90652-7](https://doi.org/10.1016/0006-3002(61)90652-7), 1961.

777 Ma, T., Furutani, H., Duan, F., Kimoto, T., Jiang, J., Zhang, Q., Xu, X., Wang, Y., Gao, J., Geng, G., Li, M.,
778 Song, S., Ma, Y., Che, F., Wang, J., Zhu, L., Huang, T., Toyoda, M., and He, K.: Contribution of
779 hydroxymethanesulfonate (HMS) to severe winter haze in the North China Plain, *Atmospheric Chemistry
780 and Physics*, 20, 5887–5897, <https://doi.org/10.5194/acp-20-5887-2020>, 2020.

781 Moch, J. M., Dovrou, E., Mickley, L. J., Keutsch, F. N., Cheng, Y., Jacob, D. J., Jiang, J., Li, M., Munger, J.
782 W., Qiao, X., and Zhang, Q.: Contribution of Hydroxymethane Sulfonate to Ambient Particulate Matter: A
783 Potential Explanation for High Particulate Sulfur During Severe Winter Haze in Beijing, *Geophysical
784 Research Letters*, 45, 11,969-11,979, <https://doi.org/10.1029/2018GL079309>, 2018.

785 Moch, J. M., Dovrou, E., Mickley, L. J., Keutsch, F. N., Liu, Z., Wang, Y., Dombek, T. L., Kuwata, M.,
786 Budisulistiorini, S. H., Yang, L., Decesari, S., Paglione, M., Alexander, B., Shao, J., Munger, J. W., and
787 Jacob, D. J.: Global Importance of Hydroxymethanesulfonate in Ambient Particulate Matter: Implications
788 for Air Quality, *Journal of Geophysical Research: Atmospheres*, 125, e2020JD032706,
789 <https://doi.org/10.1029/2020JD032706>, 2020.

790 Okabayashi, H., Okuyama, M., Kitagawa, T., and Miyazawa, T.: The Raman Spectra and Molecular
791 Conformations of Surfactants in Aqueous Solution and Crystalline States, *Bulletin of the Chemical
792 Society of Japan*, 47, 1075–1077, <https://doi.org/10.1246/bcsj.47.1075>, 1974.

793 Olson, C. N., Galloway, M. M., Yu, G., Hedman, C. J., Lockett, M. R., Yoon, T., Stone, E. A., Smith, L. M.,
794 and Keutsch, F. N.: Hydroxycarboxylic Acid-Derived Organosulfates: Synthesis, Stability, and
795 Quantification in Ambient Aerosol, *Environmental Science & Technology*, 45, 6468–6474,
796 <https://doi.org/10.1021/es201039p>, 2011.

797 Pavia, D. L., Lampman, G. M., Kriz, G. S., and Vyvyan, J. A.: *Introduction to Spectroscopy*, Cengage
798 Learning, 745 pp., 2008.

799 Peng, C. and Chan, C. K.: The water cycles of water-soluble organic salts of atmospheric importance,
800 Atmospheric Environment, 35, 1183–1192, [https://doi.org/10.1016/S1352-2310\(00\)00426-X](https://doi.org/10.1016/S1352-2310(00)00426-X), 2001.

801 Reggente, M., Höhn, R., and Takahama, S.: An open platform for Aerosol InfraRed Spectroscopy analysis
802 – AIRSpec, Atmospheric Measurement Techniques, 12, 2313–2329, [https://doi.org/10.5194/amt-12-](https://doi.org/10.5194/amt-12-2313-2019)
803 2313-2019, 2019.

804 Russell, L. M., Bahadur, R., and Ziemann, P. J.: Identifying organic aerosol sources by comparing
805 functional group composition in chamber and atmospheric particles, Proceedings of the National
806 Academy of Sciences, 108, 3516–3521, <https://doi.org/10.1073/pnas.1006461108>, 2011.

807 Ruthenburg, T. C., Perlin, P. C., Liu, V., McDade, C. E., and Dillner, A. M.: Determination of organic matter
808 and organic matter to organic carbon ratios by infrared spectroscopy with application to selected sites in
809 the IMPROVE network, Atmospheric Environment, 86, 47–57,
810 <https://doi.org/10.1016/j.atmosenv.2013.12.034>, 2014.

811 Saltzman, E. S., Savoie, D. L., Prospero, J. M., and Zika, R. G.: Methanesulfonic acid and non-sea-salt
812 sulfate in pacific air: Regional and seasonal variations, Journal of Atmospheric Chemistry, 4, 227–240,
813 <https://doi.org/10.1007/BF00052002>, 1986.

814 Sato, S., Higuchi, S., and Tanaka, S.: Structural Examinations of “Sodium Formaldehyde Sulfoxylate” by
815 Infrared and Raman Spectroscopy, Nippon Kagaku Kaishi, 1984, 1151–1157,
816 <https://doi.org/10.1246/nikkashi.1984.1151>, 1984.

817 Segneanu, A. E., Gozescu, I., Dabici, A., Sfirloaga, P., and Szabadai, Z.: Organic Compounds FT-IR
818 Spectroscopy, IntechOpen, <https://doi.org/10.5772/50183>, 2012.

819 Seinfeld, J. H. and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate
820 Change, Wiley, 2016.

821 Shurvell, H. F.: Spectra– Structure Correlations in the Mid- and Far-Infrared, in: Handbook of Vibrational
822 Spectroscopy, American Cancer Society, <https://doi.org/10.1002/0470027320.s4101>, 2006.

823 Smith, S. J., van Aardenne, J., Klimont, Z., Andres, R. J., Volke, A., and Delgado Arias, S.: Anthropogenic
824 sulfur dioxide emissions: 1850–2005, Atmospheric Chemistry and Physics, 11, 1101–1116,
825 <https://doi.org/10.5194/acp-11-1101-2011>, 2011.

826 Song, S., Gao, M., Xu, W., Sun, Y., Worsnop, D. R., Jayne, J. T., Zhang, Y., Zhu, L., Li, M., Zhou, Z., Cheng,
827 C., Lv, Y., Wang, Y., Peng, W., Xu, X., Lin, N., Wang, Y., Wang, S., Munger, J. W., Jacob, D. J., and McElroy,
828 M. B.: Possible heterogeneous chemistry of hydroxymethanesulfonate (HMS) in northern China winter
829 haze, Atmospheric Chemistry and Physics, 19, 1357–1371, <https://doi.org/10.5194/acp-19-1357-2019>,
830 2019.

831 Stone, E. A., Yang, L., Yu, L. E., and Rupakheti, M.: Characterization of organosulfates in atmospheric
832 aerosols at Four Asian locations, Atmospheric Environment, 47, 323–329,
833 <https://doi.org/10.1016/j.atmosenv.2011.10.058>, 2012.

834 Surratt, J. D., Chan, A. W. H., Eddingsaas, N. C., Chan, M., Loza, C. L., Kwan, A. J., Hersey, S. P., Flagan, R.
835 C., Wennberg, P. O., and Seinfeld, J. H.: Reactive intermediates revealed in secondary organic aerosol

836 formation from isoprene, *Proceedings of the National Academy of Sciences*, 107, 6640–6645,
837 <https://doi.org/10.1073/pnas.0911114107>, 2010.

838 Tang, K.: Chemical Diversity and Biochemical Transformation of Biogenic Organic Sulfur in the Ocean,
839 *Frontiers in Marine Science*, 7, 2020.

840 U.S. EPA. Comptox Chemicals Dashboard:
841 <https://comptox.epa.gov/dashboard/chemical/details/DTXSID80805075> (accessed November 14, 2022)
842 Methyl 10,10-diethoxydec-2-ene-4,6,8-triynoate, 2022.

843 Wang, Y., Zhao, Y., Wang, Y., Yu, J.-Z., Shao, J., Liu, P., Zhu, W., Cheng, Z., Li, Z., Yan, N., and Xiao, H.:
844 Organosulfates in atmospheric aerosols in Shanghai, China: seasonal and interannual variability, origin,
845 and formation mechanisms, *Atmospheric Chemistry and Physics*, 21, 2959–2980,
846 <https://doi.org/10.5194/acp-21-2959-2021>, 2021.

847 Wei, L., Fu, P., Chen, X., An, N., Yue, S., Ren, H., Zhao, W., Xie, Q., Sun, Y., Zhu, Q.-F., Wang, Z., and Feng,
848 Y.-Q.: Quantitative Determination of Hydroxymethanesulfonate (HMS) Using Ion Chromatography and
849 UHPLC-LTQ-Orbitrap Mass Spectrometry: A Missing Source of Sulfur during Haze Episodes in Beijing,
850 *Environmental Science & Technology Letters*, 7, 701–707, <https://doi.org/10.1021/acs.estlett.0c00528>,
851 2020.

852 Yazdani, A., Dudani, N., Takahama, S., Bertrand, A., Prévôt, A. S. H., El Haddad, I., and Dillner, A. M.:
853 Fragment ion–functional group relationships in organic aerosols using aerosol mass spectrometry and
854 mid-infrared spectroscopy, *Atmospheric Measurement Techniques*, 15, 2857–2874,
855 <https://doi.org/10.5194/amt-15-2857-2022>, 2022.

856 Zawadowicz, M. A., Proud, S. R., Seppäläinen, S. S., and Cziczo, D. J.: Hygroscopic and phase separation
857 properties of ammonium sulfate/organics/water ternary solutions, *Atmospheric Chemistry and Physics*,
858 15, 8975–8986, <https://doi.org/10.5194/acp-15-8975-2015>, 2015.

859 Zeng, G., Kelley, J., Kish, J. D., and Liu, Y.: Temperature-Dependent Deliquescent and Efflorescent
860 Properties of Methanesulfonate Sodium Studied by ATR-FTIR Spectroscopy, *The Journal of Physical*
861 *Chemistry A*, 118, 583–591, <https://doi.org/10.1021/jp405896y>, 2014.

862 Zhao, X., Shi, X., Ma, X., Zuo, C., Wang, H., Xu, F., Sun, Y., and Zhang, Q.: 2-Methyltetrol sulfate ester-
863 initiated nucleation mechanism enhanced by common nucleation precursors: A theory study, *Science of*
864 *The Total Environment*, 723, 137987, <https://doi.org/10.1016/j.scitotenv.2020.137987>, 2020.

865 Zhong, L. and Parker, S. F.: Structure and vibrational spectroscopy of methanesulfonic acid, *Royal Society*
866 *of Open Science*, 5, 181363, <https://doi.org/10.1098/rsos.181363>, 2022.

867