

1 **Impacts of anemometer changes, site relocations and processing**

2 **methods on wind speed trends in China**

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14 Manuscript for *Atmospheric Measurement Techniques*

15 ~~March~~ ~~Aug~~ ~~September~~ ~~2433~~, 2023

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17 **Abstract**

18 *In-situ* surface wind observation is a critical meteorological data source for various
19 research fields. However, data quality is affected by factors such as surface friction
20 changes, station relocations, and anemometer updates. Previous methods to address
21 discontinuities have been insufficient, and processing methods have not always adhered
22 to World Meteorological Organization (WMO) [World Climate Programme](#) guidelines.
23 We analyzed data discontinuity caused by anemometer changes and station relocations
24 in China's daily *in-situ* near-surface (~10m) wind speed observations and the impact
25 of the processing methods on wind speed trends. By comparing the wind speed
26 discontinuities with the recorded location changes, we identified 90 stations that
27 showed abnormally increasing wind speeds due to relocation. After removing those
28 stations, we followed a standard quality control method recommended by the World
29 Meteorological Organization to improve the data reliability and applied Thiessen
30 Polygons to calculate the area-weighted average wind speed. The result shows that
31 China's recent reversal of wind speed was reduced by 41% after removing the
32 problematic stations, with an increasing trend of $0.017 \text{ m s}^{-1} \text{ year}^{-1}$ ($R^2 = 0.64, P < 0.05$),
33 emphasizing the importance of robust quality control and homogenization protocols in
34 wind trend assessments.

35 **Keywords.** wind speed trends; anemometer changes; station relocations; processing
36 methods; quality control; [data](#) homogenization

37

38 **1. Introduction**

39 *In-situ* surface wind observation is a key meteorological data that has been used in
40 various avenues of research, e.g., wind power evaluation (Tian et al., 2019; Zeng et al.,
41 2019; Liu et al. 2022a), extreme wind hazard monitoring and prevention (Zhou et al.,
42 2002; Tamura, 2009; Liu et al., 2022b), and evapotranspiration analysis (Rayner, 2007;
43 McVicar et al., 2012), to name but a few. The application of robust quality control and
44 homogenization protocols are crucial for generating reliable wind speed time series for
45 further trend and variability analyses (Azorin-Molina et al., 2014; Azorin-Molina et al.,
46 2019).

47 Wind data quality is affected by surrounding surface friction change, station
48 location issues, and anemometer changes in type and height (Masters et al., 2010; Wan
49 et al., 2010; Cao & Yan, 2012; Hong et al., 2014; He et al., 2014; Azorin-Molina et al.,
50 2018; Camuffo et al., 2020). Surrounding surface friction changes are mainly associated
51 with urbanization (Zhang et al., 2022) and vegetation growth (Vautard et al., 2010),
52 which modify wind speed fields around the stations. Because of these issues, stations
53 are relocated to satisfy observing criteria (Trewin, 2010). Station relocation is quite
54 common in rapidly developing countries. For instance, about 60% of stations in China
55 experienced relocation (Sohu, 2004). Some relocation-caused breakpoints have been
56 corrected by parallel observations (i.e. operating observations for an overlapping period
57 at both the old and new observing stations; CMA, 2011; CMA, 2012; WMO, 2020), but
58 not all (Feng et al., 2004; Fu et al., 2011; Patzert et al., 2016; Tian et al., 2019; Yang et
59 al., 2021). Besides relocation caused by rapid urbanization (or vegetation growth),
60 updates to automatic anemographs at the beginning of the 21st century in China also
61 caused discontinuities in wind series (Fu et al., 2011).

62 Scientists have tried different methods to handle discontinuities. Tian et al. (2019)
63 and Yang et al. (2021) deleted stations with recorded changes in latitudes, longitudes or
64 altitudes, but they omitted to check whether those recorded relocations caused an abrupt
65 discontinuity in the time series or if parallel observations have corrected them. This
66 results in some stations being mistakenly deleted and significantly reduced the number

67 of available stations. Other research used statistical methods to detect or correct the
68 time series' abnormal breakpoint (Feng et al., 2004; Wang, 2008). However, without
69 examining the causes behind the discontinuity, this may also mistakenly delete stations
70 with natural abrupt climatic changes (Bathiany et al., 2003). Combining those two
71 methods by matching discontinuity with recorded station relocation is needed. Li et al.
72 (2018) have manually checked the station histories for nine stations in North West
73 China, but an algorithm is required to apply this approach to large datasets.

74 Besides data discontinuities, the processing method also affects the wind series.
75 There are two critical steps in the processing: 1) selecting qualified stations and 2)
76 calculating the average value. As for the first step, World Meteorological Organization
77 (WMO) [World Climate Programme](#) suggests deleting stations with either too much
78 missing data or continuous missing data ([WMO, 2003](#); WMO, 2017). Previous studies
79 only constrained the number of missing values monthly (Zeng et al., 2019), yearly (Tian
80 et al., 2019) or even in the whole period (Yang et al., 2021) but didn't check whether
81 the missing values were continuous. As for the second step, most studies used the
82 station average as the mean wind speed (Li et al., 2017; Zeng et al., 2019; Tian et al.,
83 2019; Yang et al., 2021; Shen et al., 2021; Zha et al., 2021). However, given station
84 distribution and wind speed spatial variation are often inhomogeneous with larger wind
85 but fewer station in Northwest while smaller wind but more stations in Southeast (Feng
86 et al., 2004; Fu et al., 2011; Liu et al., 2019). Therefore, the station average will have
87 spatial biases wind variation in Northwest is underrepresented because of few stations.
88 An improved average method (area weighted average) to rearrange the weight for each
89 station based on the area it represents, e.g. Thiessen Polygon (Fu et al., 2011), is needed.
90 and Thiessen Polygon (Thiessen, 1991) is widely used in which the area is only
91 determined by the station locations while other method like grids is sensitive to the
92 grids chosen. Thiessen Polygon (Fu et al., 2011)

93 Herein, taking stations in China as an example, we analyzed the existing data
94 discontinuities and their potential causes. Furthermore, we propose an improved
95 solution by using an algorithm to compare the statistic breakpoint with the recorded

96 relocation to double-check the discontinuity caused by relocation. Then using WMO's
97 quality control criteria and Thiessen Polygon (Thiessen, 1911), we generated wind
98 speed time series without temporal bias caused by heterogeneous missing values and
99 spatial biases caused by uneven station distribution.

101 **2. Dataset and methodology**

102 **2.1 WMO quality control method**

103 We used the China Surface Climatic Data Daily Data Set (CSD) (Version 3.0) from
104 the China Meteorological Data Service Center (<http://data.cma.cn/en/?r=data/>; last
105 accessed March 2020). The quality control method is recommended by WMO (2017),
106 which required the following criteria before using the daily mean values in a month as
107 monthly mean values: (1) <11 missing daily values in a month; ~~and~~ (2) <5 consecutive
108 missing daily values in a month; ~~(3) Complete monthly values for every month during~~
109 ~~the study period~~ ~~Qualifying stations must have monthly values for every month during~~
110 ~~the study period; Otherwise, the station will be completely excluded from the~~
111 ~~calculation~~. The station excluded by each criterion can be found in Table S1.

113 **2.2 Station location changes in record**

114 CSD provides daily wind speed and location information for 840 stations for 1961-
115 2019. But there are some mistakes in the daily location records. For example, if the
116 station location changed from A to B and back to A within a month, B is potentially a
117 mistaken record. Therefore, we first use mode (the statistic term meaning the value that
118 appears most often, here referring to the location with the highest frequency in a month)
119 to resample the daily location to the monthly location. Second, considering that
120 recorded longitude and latitude has the same spatial resolution of minutes, we defined
121 the threshold of location change as the minimum accuracy of the longitude and latitude
122 record, i.e., one minute. That is 1.85 km for longitude and $1.85 \text{ km} \times \cos\varphi$ for latitude,
123 where φ is the latitude. Third, as for altitude, we allow a 20m measuring error following
124 Tian et al. (2019). A station with more than 20m change in altitude will be considered

125 as relocation. It is noteworthy that CSD labels uncertain altitude records by adding 10
 126 km to the raw data (CMA, 2017), which are considered as no observations in our
 127 analysis. This way, we identified 432 stations as relocations from the 601 qualified
 128 stations after applying the WMO quality control (details in Table S2).

130 2.3 Breakpoint detection and the comparison with recorded relocation

131 We used Pruned Exact Linear Time (PELT) method (Killick, Fearnhead & Eckley,
 132 2012) to detect the jumps in the mean level in the monthly wind speed time series (Fig
 133 4a, Fig 4c). This method is a wrapped function named *findchangepts* in Matlab. PELT
 134 is essentially a traversing method. For a time series with N values ($x_1, x_2 \dots x_N$), the
 135 function uses equations 1 & 2 to calculate the total residual errors (J) for each point (k)
 136 assumed as a breakpoint. The point with the most significant change in the mean (lowest
 137 total residual errors, J) is reported as the breakpoint. The breakpoints here can be caused
 138 by artificial relocations or natural climate changes.

$$139 \quad J(k) = \sum_{i=1}^{k-1} (x_i - \text{mean}([x_1 \dots x_{k-1}]))^2 + \sum_{i=k}^N (x_i - \text{mean}([x_k \dots x_N]))^2 \quad (1)$$

$$140 \quad \text{mean}([x_m \dots x_n]) = \frac{1}{n-m+1} \sum_{r=m}^n x_r \quad (2)$$

141 Then we use relocation records to separate changes brought by artificial relocation
 142 from changes in natural climate. If the breakpoint and one of the relocation dates (some
 143 stations have more than one relocation record) happened in the same two months, we
 144 will consider that the time series is significantly affected by the relocation, and the
 145 station will be deleted. Stations with natural-climate-caused location changes will be
 146 reserved.

147 The change point in the trend of the annual national average wind speed (Fig 2b,
 148 Fig 5b) is detected following the method used by Wang et al. (2011). All the trends
 149 reported are based on the least square fits.

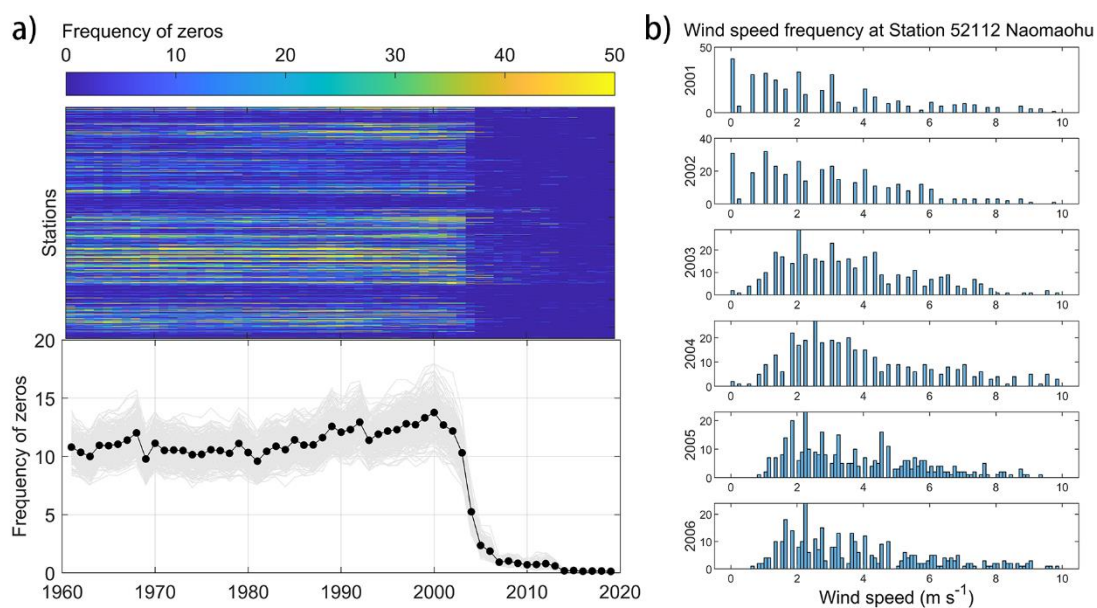
151 3. Results and discussion

152 3.1 Data issue due to anemometer changes

153 We found a clear decline in the frequency of zeros (zero wind speed) in most CSD
154 stations between 2002 and 2007 (Figure 1a), from 10-14 days per year to less than two
155 days per year. This clear drop is not a result of wrongly taking zero values as no
156 observations (NaN) as happened in the Integrated Surface Dataset (ISD, Dunn et al.,
157 2022), as no abrupt increase in NaN frequency was observed (Supplementary Figure
158 S1). Instead, the decline is accompanied by an ~~improvement-decrease~~ in measure
159 ~~uncertainty-record-accuracy~~: i.e., the measurement intervals became narrow (from 0, 0.3,
160 0.5, 0.7, 0.8, 1.0 m s⁻¹, etc. to 0, 0.1, 0.2, 0.3 m s⁻¹, etc.; Figure 1b and Supplementary
161 Figure S2). Taking Station Naomaohu in Xinjiang (station ID: 57432) as an example,
162 from 2002 to 2003, zero values decreased from more than 30 days per year to less than
163 five days per year and wind speed records changed from 0, 0.3, 0.7, 1.0 m s⁻¹, etc. to 0,
164 0.3, 0.5, 0.8, 1.0 m s⁻¹, etc. Since 2004, the measurement record-accuracy was further
165 improved to 0.1, 0.2, 0.3... m s⁻¹ and zeros values almost disappeared (Figure 1b).

166 This change is potentially caused by the transformation in measure frequency,
167 anemometer type and data logging, based on the station history recorded by Xin et al.
168 (2012). As for measurement frequency, in 2003, Station Naomaohu changed from 3
169 observations per day (i.e., 8:00, 14:00 and 20:00, China Standard Time) to four times
170 per day (2:00, 8:00, 14:00 and 20:00, China Standard Time). The increase in the
171 frequency of measurements decreases zeros in daily wind data, as only if all
172 observations report zero wind speeds, will the daily data (i.e., the average of all
173 observations in a day; CMA, 2017) be recorded as zero. Then in 2005, the EL contact
174 anemograph (Yang, 1986; Jin, 2011; Xin et al., 2012, Zhang et al., 2020) requiring
175 manual recording was changed to EC photoelectric encoder self-recording type (Kuang,
176 2016; Jin, 2011; Xin et al., 2012). Both EL and EC type anemographs use cup
177 anemometers to measure wind speed. This anemograph change further decreases the
178 likelihood of recording zero daily wind speed because the updated new anemometers
179 are more sensitive, and even very low wind speeds will be measured with a value
180 instead of recorded as zero (Azorin-Molina et al., 2018). The smooth increasing
181 frequency of zero values from 1960 until 2000 also supports this statement (Figure 1a):

182 the longer the anemometer is used, the less sensitive it will become, and hence a greater
 183 wind speed will be required to record a non-zero value (Azorin-Molina et al., 2018),
 184 overall increasing the zero values. As for the change in data accuracy, there are two
 185 reasons: 1) EL type anemograph only measures the times of electronic contact (200
 186 meters rotation distance per contact) in 10 mins, therefore it has discrete records. For
 187 example, one contact means 0.3 m s^{-1} (200m/600s) and two contacts means 0.7 m s^{-1}
 188 ($400\text{m}/600\text{s}$) (Hu et al., 2009) while EC type has more accurate records using the Grey
 189 Code; 2) the data logging changed from manual reading, calculating and rounding to
 190 instrument automatically calculating and retaining one decimal place. This example
 191 shows us the importance of recording siting criteria, required functional specifications
 192 of wind sensors and maintenance policy. However, those records are missing for most
 193 of the stations which hindered the quality classification and data processing.



194
 195 **Figure 1. Changes in wind speed data caused by anemometer updates. a)** Decrease of
 196 frequency of zeros. Each horizontal bar in the upper figure represents one station and
 197 there are 840 stations in total. The color indicates the frequency of zeros (days per year).
 198 The black dotted line in the lower figure is the average annual frequency of zeros of all
 199 the stations. The 300 grey lines are sample averages, each containing 40% amount of
 200 the total stations. **b)** Frequency (days per year) of daily wind speed measurements
 201 between 2001 and 2006 for Station #52112 Naomaohu ($43^{\circ}45'N$, $94^{\circ}59'E$, 479.0 m

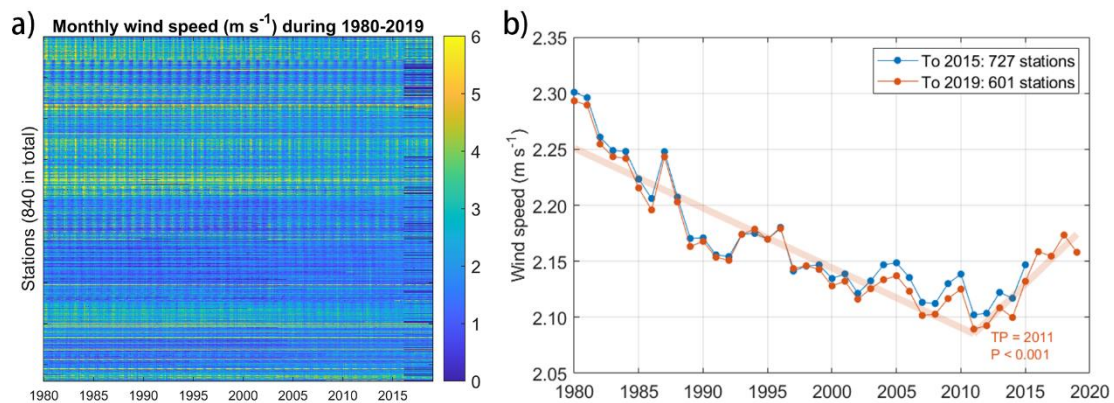
202 a.s.l.)

203

204 3.2 Quality-controlled series

205 Following WMO's criteria, we generated the monthly average wind speed for each
206 station (Figure 2a). We found that since January 2016, there have been 126 stations that
207 no longer have records (distribution see Figure S3). We compared the time series with
208 and without these stations and found the difference is not significant (t-test $P < 0.001$,
209 Figure 2b). To obtain a longer time series including recent years' data, we deleted the
210 126 stations and only used the 601 stations with complete monthly average wind speeds
211 for 1980-2019. The breakpoint was detected in 2011 ($P < 0.001$) with a decreasing trend
212 of $-0.011 \text{ m s}^{-1} \text{ year}^{-1}$ ($R^2 = 0.84$, $P < 0.001$) before the breakpoint and an increasing
213 trend of $+0.022 \text{ m s}^{-1} \text{ year}^{-1}$ ($R^2 = 0.87$, $P < 0.001$) after.

214



215

216 **Figure 2. Monthly average wind speed after being filtered by WMO's criteria. a)**

217 Each horizontal bar represents one station. Months with no data (NaNs) are represented

218 by the deepest blue. **b)** Comparison of the monthly average wind speed for the short-

219 (1980-2015; 727 stations) and long-period (1980-2019; 601 stations)

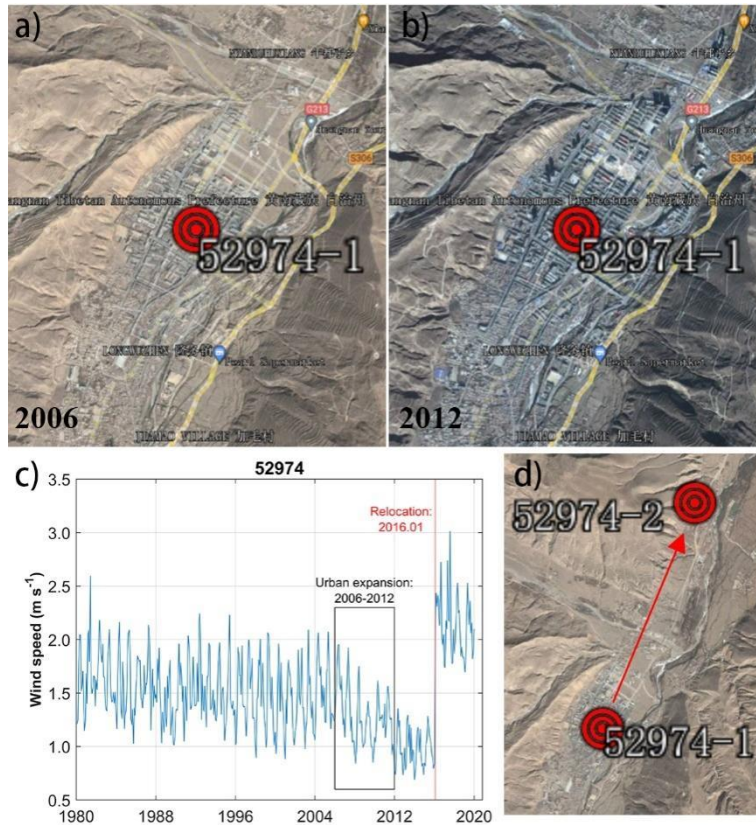
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221 3.3 Station relocations caused by urbanization

222 Another key factor influencing wind speed measurements is the relocation of
223 stations. We found that there is a clear data jump caused by relocations in some of the
224 stations. Taking the station located in Qinghai (station ID 52974) as an example, we
225 detected an abrupt jump in wind speed in January 2016. This date coincides with the

226 relocation of the station from 35°31'N, 102°01'E (ID 52974-1) in December 2015 to
227 35°33'N, 102°02'E (ID 52974-2) in January 2016 (Figures 3c & 3d). The relocation is
228 potentially attributed to the urban growth around the station. As viewed by satellite
229 images from Google Earth Pro, there is a rapid urban expansion from 2006 (Figure 3a)
230 to 2012 (Figure 3b), especially towards the Northeast of the station, during wind speed
231 records also experienced a decrease (Figure 3c). A similar decrease in both daily mean
232 wind speed and maximum wind speed caused by urbanization was also reported in the
233 Yangtze River region (Zhang et al., 2022). To eliminate the effect of buildings on the
234 wind speed measurements, Station 52974 was moved to 4 km away from its previous
235 location (Figure 3d) so that wind speed is properly measured without artificial obstacles
236 in the surroundings. However, this estimation of roughness change based on satellite
237 data is rough. A more proper way as required by the World Climate Data and Monitoring
238 Programme is to record the change in the station logbook (WMO, 2021), which will
239 provide more reliable information about the quality of the data. But most stations don't
240 have such a record. Despite the absence of mete data, we used an established global
241 roughness model through satellite albedo observations to monitor alterations in surface
242 roughness. For the selected station, we employed the roughness estimation technique
243 devised by Chappell & Webb (2016) to analyze changes in roughness across a 5 km x
244 5 km area encompassing the station's location. Our quantitative examination of
245 roughness alterations aligns with the findings derived from satellite imagery analysis,
246 affirming a pronounced increase in roughness between 2000 and 2010 (Supplementary
247 Figure S4). This increase in roughness likely contributed to the observed decline in
248 wind speed and ultimately compelled the relocation of the station.

249



250

251

Figure 3. Example of station relocation caused by rapid urbanization growth. a-b)

252

[Landsat image crop from Google Earth](#) [satellite image](#) near Station 52974 in 2006 and

253

2012, respectively. c) the wind speed change with urbanization and relocation. d)

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[Landsat image](#) [satellite image](#) of the station relocation [crop from Google Earth](#).

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256

Though some stations were influenced by station relocation as shown in Figure 3,

257

a larger fraction (79%) of stations show no change in wind speed after the relocation.

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Further checking the raw record of locations for those stations, we find that one reason

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is that some “relocations” result from wrong location records. For example, Station

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52974 is mistakenly detected with three relocations (Figure 4a). However, only the first

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relocation is real and the latter two are results of location encoding change from 10202

262

(interpreted as 102°02′) to 1022 (interpreted as 10°22′) and back. Another possible

263

reason is that the relocation did happen but the data has been corrected. According to

264

the *Provisional Regulations on Relocation, Construction and Removal of National*

265

Ground Meteorological Observation announced by China’s government in 2012,

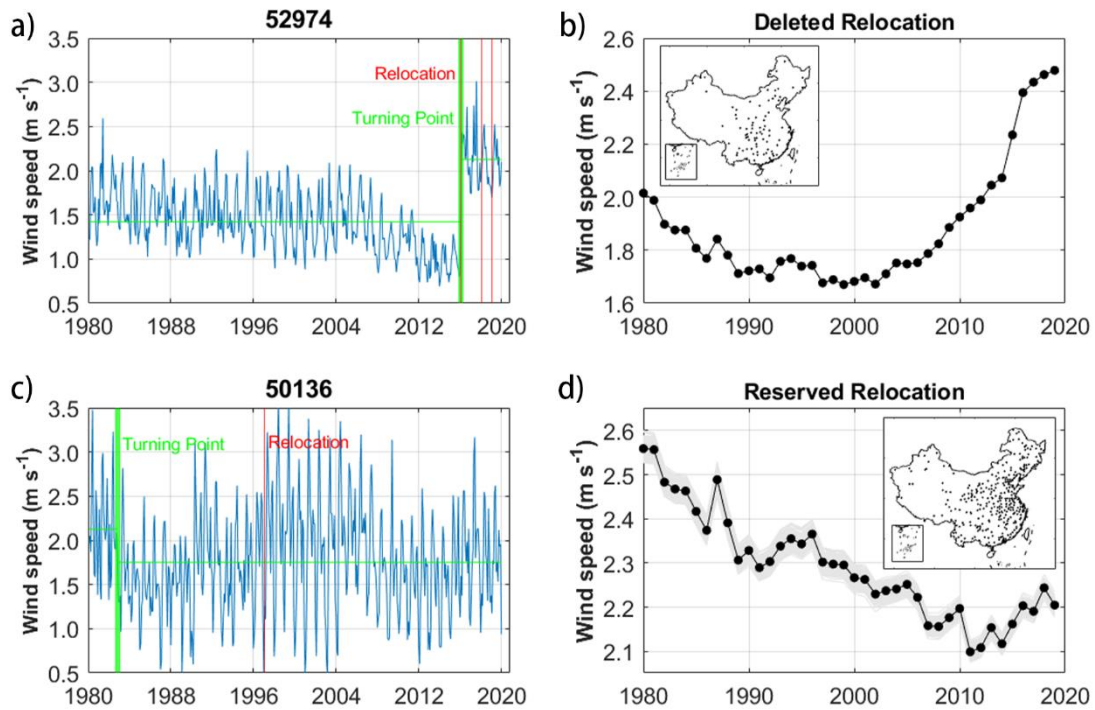
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station relocations should have 1-2 years of parallel observations for data correction

267 (CMA, 2012). This process may fix some of those discontinuities but not all (Feng et
268 al., 2004; Fu et al., 2011; Patzert et al., 2016; Tian et al., 2019; Yang et al., 2021). For
269 example, Station 59287, the only national basic weather station in Guangzhou,
270 experienced two relocations in both 1996 and 2011, which is confirmed by the metadata
271 (CMA, 2011). After correction, the 1996 relocation doesn't show a sharp breaking point
272 but the 2011 one does (Supplementary Figure S54).

273 To examine whether the relocation caused a substantial change in the wind speed
274 record, we identified the most abrupt change in the wind speed time series and checked
275 whether a relocation happened near the change point (see details in *Methods 2.3*). Out
276 of the 432 relocated stations, 90 were deleted because the most significant shift in mean
277 is at the time of the relocation, and hence this is the most likely cause. We then took the
278 average of the “deleted relocation” stations and “reserved relocation” stations
279 separately. The “deleted relocation” group shows an abnormally rapid increase in the
280 recent two decades (Figure 4b). While the “reserved relocation” group is similar to
281 stations without relocation (Supplementary Figure S65). To exclude the impact of
282 different station counts in each category (fewer stations mean higher sensitivity to the
283 individual abnormal station), we performed 300 samples using a random draw of 90
284 stations from the “reserved relocation” group and showed them in grey lines in Figure
285 4d. None of the grey lines shows an abnormal trend as the “deleted relocation” group.
286 This proves that our method is efficient in identifying problematic stations.

287



288

289 **Figure 4. Comparison of deleted relocated stations and reserved ones.** a) The wind
 290 speed data breakpoint and relocations of one example of deleted relocation, Station
 291 52974. b) The station average wind speed of 90 deleted relocated stations. The inset
 292 shows the station distribution across China. c) One example of reserved relocation,
 293 Station 50136. d) The station average wind speed of 342 reserved relocated stations.
 294 The grey lines are the averages of 300 samples, each with 90 randomly drawn reserved
 295 relocated stations. [Maps information are from Department of Natural Resources](#)
 296 [standard map service system of China.](#)

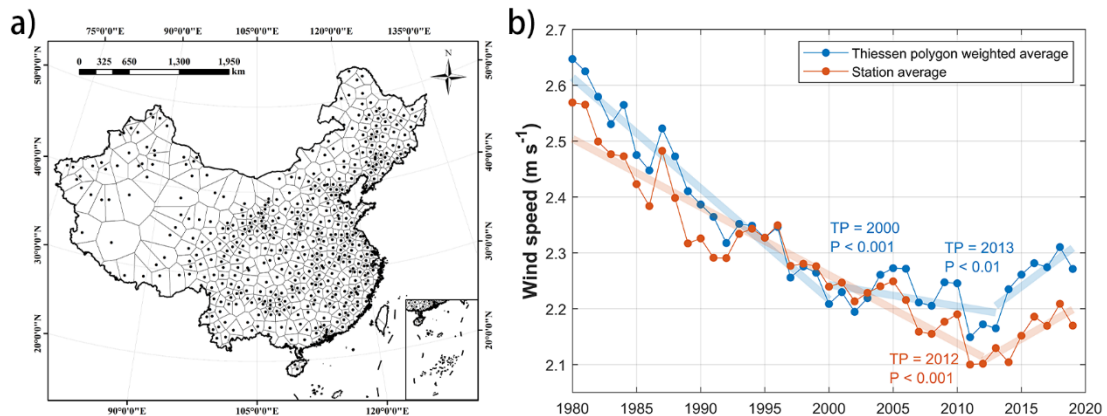
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298 3.4 Average method used to calculate the national average

299 In the station average time series, the breakpoint was detected in 2012 ($P < 0.001$)
 300 with a trend of $-0.012 \text{ m s}^{-1} \text{ year}^{-1}$ ($R^2 = 0.90$, $P < 0.001$) before and $+0.013 \text{ m s}^{-1} \text{ year}^{-1}$
 301 ($R^2 = 0.70$, $P < 0.01$) after (Figure 5b). The increasing trend decreased by 41% after
 302 deleting those relocation-affected stations, compared with the $+0.022 \text{ m s}^{-1} \text{ year}^{-1}$ in
 303 Figure 2b (also reported by Liu et al., 2022a). But the trend is larger than the $+0.011 \text{ m}$
 304 $\text{s}^{-1} \text{ yr}^{-1}$, reported by Yang et al. (2021), with all the recorded location changed stations
 305 deleted without checking whether the station is affected by the relocation.

306 We further used Thiessen Polygon (Thiessen, 1911) to give different weights to
307 each station according to their representing area, i.e., large weight for stations located
308 in sparse stations area (Figure 5a) and compare the result with the station average
309 (Figure 5b). The Thiessen Polygon method, also known as the Voronoi Diagram, is a
310 spatial analysis technique often employed in hydrology and climatology. It involves
311 tessellating a region into polygons based on point data, such that each polygon
312 encompasses only one data point, and every location within a polygon is closer to its
313 associated point than any other. This method is particularly useful for interpolating
314 values across a region when the exact nature of change between points is unknown or
315 when changes are abrupt. By drawing perpendicular bisectors between adjacent data
316 points, the entire area is divided, with each polygon assuming the value of its associated
317 data point. While straightforward and clear in its delineation, the Thiessen Polygon
318 method assumes uniform variation within each polygon~~Thiessen Polygon is essentially~~
319 ~~the finest divided subregion, which splits the region into the smallest representative~~
320 ~~area and ensure there is a station in each subregion; Moreover, because the Thiessen~~
321 ~~Polygon employs perpendicular bisectors to partition the space, every location is~~
322 ~~assigned to its closest station as shown in Figure 5a.~~ The Thiessen polygon weighted
323 average is overall higher than the station average. This can be explained by the
324 increasing weight of stations in North West China with higher wind speeds (Liu et al.,
325 2019). While in the Thiessen polygon weighted average time series, there are two
326 breakpoints in 2000 ($P < 0.001$) and 2013 ($P < 0.01$). The trend changes from quick
327 decrease ($-0.020 \text{ m s}^{-1} \text{ year}^{-1}$, $R^2 = 0.94$, $P < 0.001$) to unstable moderate decrease ($-$
328 $0.004 \text{ m s}^{-1} \text{ year}^{-1}$, $R^2 = 0.17$, $P = 0.14$) and quickly increase ($+0.017 \text{ m s}^{-1} \text{ year}^{-1}$, $R^2 =$
329 0.64 , $P < 0.05$). The increasing trend in the recent decade increased by 31% (from
330 $+0.013 \text{ m s}^{-1} \text{ year}^{-1}$ to $+0.017 \text{ m s}^{-1} \text{ year}^{-1}$) after using the Thiessen polygon approach.
331 This is because the weights of stations in North West and South West are increased
332 when calculating the average and those area has strong increasing wind speed trend
333 (Figure S7). ~~Despite the Thiessen polygon approach already utilizing the nearest station~~
334 ~~observation to represent wind speed in locations lacking direct observations, it remains~~

335 unsatisfactory due to the intricate spatial variability of wind speed attributed to complex
336 terrains. To enhance the accuracy of wind speed interpolation, a more comprehensive
337 model necessitates additional observations within areas characterized by complex
338 terrain.



341
342 **Figure 5. Thiessen polygons and the comparison between Thiessen polygon**
343 **weighted average and station average. a)** The Thiessen polygon map of the 511
344 qualified stations. **b)** The comparison of station average wind speed (orange line) and
345 Thiessen polygon weighted average (blue line) across China for 1980-2019. The linear
346 fitting models are shown in translucent thick lines accordingly. [Maps information are](#)
347 [from Department of Natural Resources standard map service system of China.](#)

348 349 **4. Conclusions**

350 Continuity is crucial for meteorological observation data. However, either the
351 updates in the anemograph, the relocation caused by urbanization or the methods of
352 data logging will affect wind speed data continuity. In this study, we comprehensively
353 examined the discontinuity in wind speed data using a Chinese dataset. We found that
354 updates to the automatic anemometer improved the observation frequency and
355 instrument sensitivity, decreasing the zero-value daily wind speed data and increasing
356 data accuracy. We also propose comparing the discontinuity in time series with recorded
357 station relocation to check whether a relocation caused a breakpoint. We found that 90

358 stations were affected by the relocation and show a quickly increasing wind speed in
359 the recent two decades. After excluding those problematic stations, the wind speed
360 reversal trend is reduced by 41% but still strong ($P < 0.001$, with an increasing trend of
361 $+0.013 \text{ m s}^{-1} \text{ year}^{-1}$). The increasing trend reaches $+0.017 \text{ m s}^{-1} \text{ year}^{-1}$ ($R^2 = 0.64$, $P <$
362 0.05) after using Thiessen Polygon, which gives the stations in North West China a
363 larger weight because their small number but located in a large area.

364 Though lots of methods (Masters et al., 2010; Wan et al., 2010; Cao & Yan, 2012;
365 Hong et al., 2014; He et al., 2014; Azorin-Molina et al., 2018; Camuffo et al., 2020)
366 were proposed to handle those problems, a comprehensive summary of them is
367 ~~lacked~~missing. ~~This study fills this research niche. However, Also,~~ it is hard for external
368 researchers to provide a better solution without a collaboration with National Weather
369 Services and the access to station data records and/or metadata. Therefore, we hope
370 National Weather Services could improve the data quality based on these feedbacks and
371 World Climate Data and Monitoring Programme's guides, and complete the process
372 by introducing an R package with open-source code on GitHub and publishing the
373 metadata. This way, not only the data is easier to get and process, but also researchers
374 can contribute to improve the dataset. One such example is the “rnpn” package to access
375 and process USA National Phenology Network data (<https://github.com/usa-npn/rnpn>).
376 Anyway, all raw data processing has limitations and adds additional uncertainty. As we
377 keep reporting problems in datasets and improving our processing method, we should
378 also pay more attention to increasing the quality and homogeneity of the wind data.
379 This requires raising awareness of the importance of protecting the environment around
380 the observation station and avoiding relocations.

381

382 **Supplemental Information**

383 Document S1. Supplemental Information, Table S1, Figures S1 – ~~S75~~.

384

385 **Acknowledgements**

386 Thanking Robert Dunn (UK Met Office) for discussions and comments on the
387 manuscript. Thank Adrian Chappell for providing the surface roughness data. –The
388 authors wish to acknowledge the reviewers for their detailed and helpful comments to
389 the original manuscript.

390

391 This study was supported by the National Natural Science Foundation of China (grant
392 no. 42071022), the Swedish Formas (2019–00509 and 2017–01408) and VR (2021–
393 02163 and 2019–03954), and the start-up fund provided by Southern University of
394 Science and Technology (no. 29/Y01296122). C. A-M. was supported by VENTS
395 (GVA-AICO/2021/023), the CSIC Interdisciplinary Thematic Platform (PTI) Clima
396 (PTI-CLIMA), the 2021 Leonardo Grant for Researchers and Cultural Creators, BBVA
397 Foundation, and the “Unidad Asociada CSIC-Universidad de Vigo: Grupo de Física de
398 la Atmosfera y del Océano”.

399

400 **Data availability statement**

401 The data that support the findings of this study are available upon request from the
402 authors.

403

404 **Author contributions**

405 **Zhenzhong Zeng:** Conceptualization, Methodology **Yi Liu:** Methodology, Software,
406 Writing – Draft **Lihong Zhou:** Methodology, Data Curation **All other authors:** Writing
407 – Review & Editing

408

409 **Declaration of interest**

410 The authors declare no competing financial interests.

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