



SI-traceable validation of a balloon-borne spectrometer for water vapor measurements in the upper atmosphere

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Abstract. Despite its crucial role in the Earth's radiative balance, upper air water vapor (H₂O) is still lacking accurate, in-situ,
and continuous monitoring. Especially in the upper troposphere-lower stratosphere (UTLS), these measurements are notori-
ously difficult, and significant discrepancies were reported in the past between different measuring techniques. Here, we pre-
sent a laboratory validation of a recently developed mid-IR quantum-cascade laser absorption spectrometer for balloon-borne
measurements of H₂O in the UTLS (ALBATROSS). The validation was performed using SI-traceable reference gas mixtures
generated based on the permeation method and dynamic dilution. The accuracy and precision of ALBATROSS were evaluated
at a wide range of pressures (30–250 mbar) and H₂O amount fractions (2.5–35 ppm), representative of the atmospheric varia-
bility of H₂O in the UTLS. Best agreement was achieved by implementing a quadratic Speed-Dependent Voigt Profile
(qSDVP) line-shape model in the spectroscopic retrieval algorithm. The molecular parameters required by this parameteriza-
tion were determined empirically using a multi-spectrum fitting approach over different pressure conditions. ALBATROSS
achieves an accuracy better than ±1.5 % with respect to the SI-traceable reference at all investigated pressures and H₂O amount
fractions. The measurement precision was found to be better than 30 ppb (i.e., 0.1 % at 35 ppm H₂O) at 1 s resolution for all
conditions. This performance, unprecedented for a balloon-borne hygrometer, demonstrates the exceptional potential of mid-
IR laser absorption spectroscopy as a new reference method for in-situ measurements of H₂O in the UTLS.

1. Introduction

Water vapor (H₂O) is the strongest greenhouse gas in the Earth's atmosphere and a major driver of the atmospheric dynamics,
microphysics, and interaction with radiation (IPCC, 2021). Its abundance in the atmosphere decreases strongly with altitude,
from typically 1–3 % close to the surface, to around 5 μmol/mol (or parts per million, ppm) in the stratosphere. Despite its
scarcity, upper air H₂O is still of great importance to the radiative balance of the atmosphere. Particularly, in the upper tropo-
sphere-lower stratosphere (UTLS), i.e. at altitudes of about 8–25 km, even small changes in H₂O were shown to have signifi-
cant impact on the rate of global warming (Solomon et al., 2010; Riese et al., 2012). Hence, accurate measurements of H₂O in



the UTLS are crucial for reliable climate predictions. However, due to its low amount fraction and the low temperatures ($T < -60$ °C) of the UTLS, accurate measurements of H₂O in this region are notoriously difficult. Several studies comparing in-situ measurements of UTLS H₂O from both aircraft and balloon platforms found significant discrepancies between different measuring techniques (e.g., Oltmans et al., 2000; Vömel et al., 2007; Rollins et al., 2014; Brunamonti et al., 2019; Singer et al., 2022), with implications for the understanding of ice microphysical processes (e.g., Peter et al., 2006; Krämer et al., 2009). Large relative discrepancies (exceeding ± 100 %) between different hygrometers were also reported from laboratory experiments simulating the UTLS conditions, such as AquaVIT-1 (Fahey et al., 2014), where only a small subset of instruments were able to achieve mean deviations below ± 10 % from a reference value at all conditions.

Currently, cryogenic frost-point hygrometry (CFH/FPH) is considered as a state-of-the-art method for balloon-borne measurements of UTLS H₂O (Fahey et al., 2014) and is routinely used in global long-term monitoring networks, such as the GCOS Reference Upper Air Network (GRUAN) (e.g., Hurst et al., 2011). CFH/FPH instruments are based on the chilled-mirror principle and have an estimated uncertainty of 4–6 % in the UTLS (Hall et al., 2016; Vömel et al. 2016). However, these devices are currently undergoing a fundamental reconception, because the cooling agent fluoroform (HFC-23) used for their operation must be phased out due its high global warming potential (UNEP, 2016). Thus, there is an urgent need for alternative, reliable technologies for the long-term monitoring of UTLS H₂O. Alternative measurement techniques demonstrated for light-weight balloon platforms include Lyman- α fluorescence (Sitnikov et al., 2007; Khaykin et al., 2013) and laser absorption spectroscopy (Durry et al., 2008; Graf et al., 2021).

The aim of this work is to validate the accuracy and precision of a newly developed mid-IR quantum-cascade laser absorption spectrometer (Graf et al. 2021). This direct absorption-based technique has the promise of being a calibration-free method, which makes it exceptionally attractive for demanding field applications. In practice, however, the accuracy can suffer from uncertainties in physical (pressure, temperature) as well as molecule-specific spectroscopic parameters, which will ultimately limit the quality of the retrieved data. While the error contribution of the environmental factors can be estimated based on the measurement uncertainties of these quantities, their large variations encountered in the stratosphere will also impact the intrinsic molecular properties of the absorption line, such as the line strength and pressure broadening parameters. Especially, a detailed knowledge of the latter is a prerequisite for an accurate spectral retrieval.

Here, we address this aspect by a dedicated laboratory investigation conducted at the Swiss Federal Institute of Metrology (METAS). Using a dynamic-gravimetric permeation method (e.g., Haerri et al., 2017; Guillevic et al., 2018), we generated SI-traceable H₂O amount fractions as low as 2.5 ppm in synthetic air with an expanded measurement uncertainty smaller than ± 1.5 %. This allowed us to improve our spectroscopic retrieval algorithm by implementing a more advanced line shape model than the standard Voigt profile, i.e., the quadratic Speed-Dependent Voigt Profile (qSDVP), and then to assess the accuracy and precision of ALBATROSS at a wide range of UTLS-relevant conditions. The molecular parameters required by the qSDVP parameterization were determined experimentally using a multi-spectrum fitting approach (e.g., Cygan and Lisak, 2017).



2. Experimental setup

2.1. ALBATROSS spectrometer

The ALBATROSS spectrometer, described in detail in Graf et al. (2021), leverages on recent advances in optics and laser driving concepts. It incorporates a monolithic, segmented circular multipass cell (SC-MPC), consisting of a rotationally symmetric arrangement of individual mirror segments carved into its inner surface (Graf et al., 2018). This geometry was found to be highly tolerant to thermally induced distortion, robust to mechanical stress, and thus, well suited for open-path applications (Tuzson et al., 2020). The SC-MPC contains 57 mirror segments ($6\times 6\text{ mm}^2$) with a diagonal distance of 108.8 mm, resulting in an effective optical path length (OPL) of 610.7 cm. The last segment of the SC-MPC is designed such that the laser beam is directly focused onto the IR-detector without any additional beam-shaping optics. A distributed feedback quantum cascade laser (DFB-QCL), packaged in a high-heat-load housing with embedded thermoelectric cooling and collimation optics (Alpes Lasers SA, Switzerland), is used as a light source. The laser is tuned across a $\sim 1\text{ cm}^{-1}$ spectral window centered around an isolated H_2O absorption line at 1662.809 cm^{-1} ($\lambda \approx 6.01\text{ }\mu\text{m}$). A thermoelectrically-cooled MCT detector (PVM-2TE-8 1×1 , Vigo Systems, Poland) is used for the detection of the transmitted light. Rapid spectral sweeping of the QCL is achieved by periodic modulation of the laser driving current, following the intermittent continuous wave (ICW) modulation approach (Fischer et al., 2014, Liu et al., 2018). The spectra are acquired at a frequency of 3 kHz and co-averaged to a resolution of 1 Hz. A full description of the laser driving and data acquisition systems can be found in Graf et al. (2021). The spectrometer weights 3.41 kg, including batteries and a Styrofoam thermal insulation enclosure.

For its laboratory operation, ALBATROSS was set up into a closed-path configuration, in which the SC-MPC is closed on both sides by stainless-steel lids. The lids, as well as all tubes in contact with the gas, were treated by a highly inert coating (SilcoNert®2000, SilcoTek, USA) that minimizes the adsorption of molecules on its surface (e.g., Vaitinen et al., 2018; Macé et al., 2022), thereby minimizing memory effects and shortening the response time. The required leak-tightness was achieved by using nitrile O-rings (between the cell body and the lids), and high-vacuum metal fittings (VCR, Swagelok, USA) for the gas handling system. The SC-MPC volume in closed-path configuration is about 140 cm^3 .

2.2. Reference gas generation

The SI-traceable reference gas mixtures used for the validation of ALBATROSS were generated based on the permeation method combined with dynamic dilution (following ISO 6145-7 and ISO 6145-10). This is an established, standardized technique in metrology, particularly for reactive or polar compounds (e.g., Scaringelli et al., 1970; Brewer et al., 2011; Haerri et al., 2017; Guillevic et al., 2018). The setup used here is the same as described in Guillevic et al. (2018). All parts in contact with the gas mixture are passivated with SilcoNert®2000. The mass loss of the permeator is determined by a magnetic suspension balance (MSB, Rubotherm, Germany) under controlled temperature and pressure. This allows for continuous and



unperturbed mass measurements, as the permeator is physically decoupled from the balance. To correct for balance drift and buoyancy, two calibration masses traceable to the SI realization of the kilogram of Switzerland (Fuchs et al., 2012) are automatically placed on the balance plate at regular intervals during the calibration routine. The permeator used here (Fine Metrology, Italy) has a nominal permeation rate of approximately 10 $\mu\text{g}/\text{min}$ at 45 °C for H_2O . The expanded relative measurement uncertainty of the generated H_2O amount fractions, determined upon calibration of the permeator including its drift and the dynamic dilution unit, was found to be better than $\pm 1.5\%$ at all conditions.

2.3. Gas handling system

Figure 1 shows a schematics of the dilution and sampling system used for the measurements. The carrier gas is synthetic air that is passed through a gas purifier (MicroTorr MC400, Saes, Italy), which reduces the H_2O content to about 70 ppb (previously measured by cavity-ring down spectroscopy). The volumetric flow rates in the different branches of the dilution and sampling system are controlled by a series of mass flow controllers (MFCs). The humidity content of the reference gas flow is adjusted by varying the flow rate of the dry synthetic air through MFC 2 between 0.05 slpm (corresponding to ~ 35 ppm H_2O) and 3 slpm (~ 2.5 ppm H_2O) to dilute the H_2O -enriched air stream that passes at a constant flow rate of 0.3 slpm (set by MFC 1) through the permeation chamber of the MSB. The total flow rate through the SC-MPC is kept constant at 0.3 slpm by MFC 3, while a solenoid valve allows the automatic switching between the reference gas mixture and the zero air (i.e., dry synthetic air) supplies. A vacuum pump and an upstream pressure controller (PC-15PSIA, Alicat, USA) are used to set the pressure in the multipass cell at various levels between 30–250 mbar. The sample pressure in the SC-MPC was monitored by a heated capacitance manometer (AA02, MKS Instruments, USA), with an absolute accuracy of 0.12 %. The standard deviation of the measured pressure was within 0.1 % during all experiments. All measurements were performed under controlled temperature conditions, monitored in the vicinity of the multipass cell by a HMP110 sensor (Vaisala, Finland), with an absolute accuracy of 0.2 °C. The average temperature measured during the campaign period was 23.5 ± 0.03 °C.

In addition to the SI-traceable reference gas mixtures, another series of measurements were performed using a secondary reference gas mixture. This allowed us to extend the range of the validation beyond the maximum value of 35 ppm H_2O . The custom-made secondary reference gas was produced by spiking a gas cylinder of synthetic air with a known amount of pure water. For this purpose, a conditioned and evacuated gas cylinder was filled with synthetic air up to 1 bar, then a syringe with distilled water was used to introduce a given amount of H_2O , followed by pressurizing the gas cylinder with synthetic air up to 100 bar, resulting in a humidity content of about 180 ppm H_2O . After a few days of equilibration, the gas was expanded into a SilcoNert®2000-coated stainless-steel cylinder to prevent any potential surface effects during the measurements.

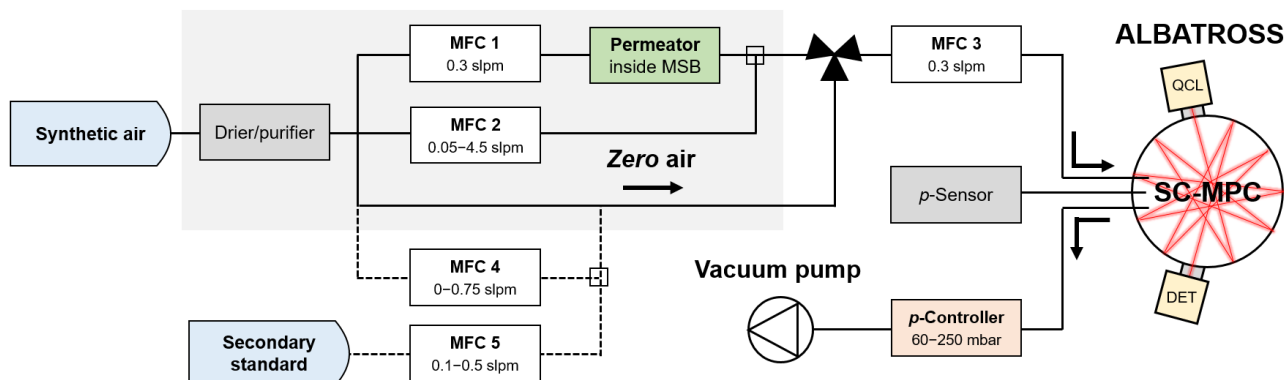


Figure 1. Scheme of the sampling system used for the validation measurements. The grey shaded area indicates the magnetic suspension balance (MSB) and the dilution system used for the generation of the SI-traceable reference gases.

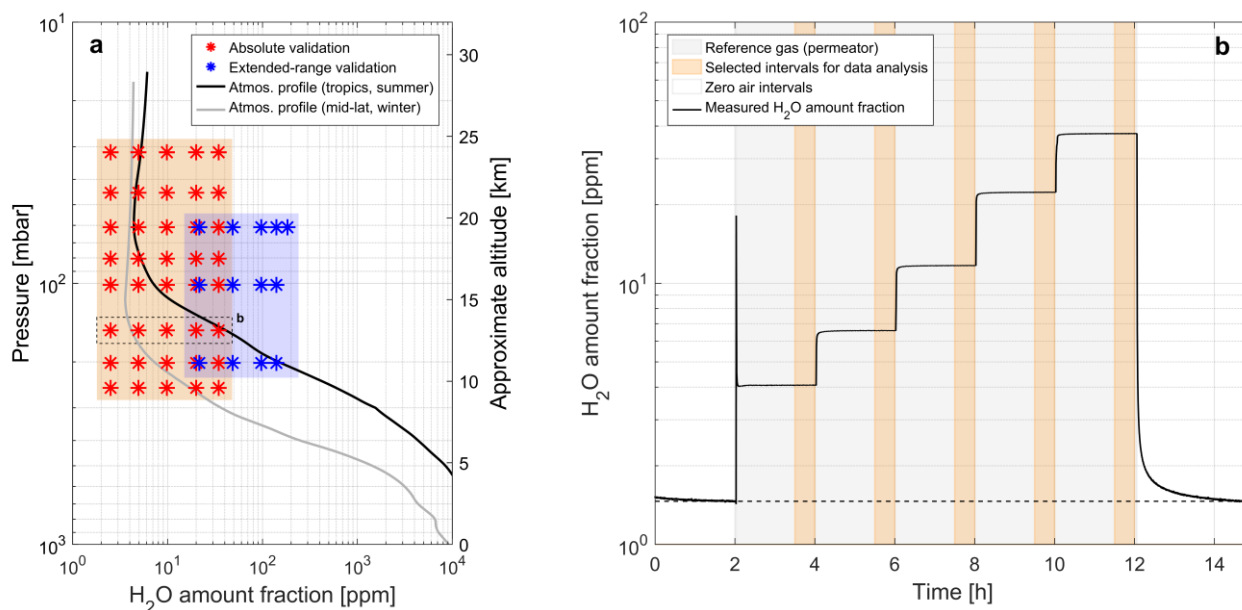
2.4. Measurements overview

5 The H₂O amount fraction levels (or "setpoints") and gas pressure (p) levels were selected to be representative of the atmospheric profile of H₂O in the UTLS. Figure 2a shows the distribution of all H₂O setpoints and p -levels investigated in this work, overlaid with two atmospheric profiles of H₂O amount fraction measured by CFH during recent field campaigns. The profiles correspond to moist, tropical summer conditions (black line) and dry, mid-latitude winter conditions (grey line).

We selected five H₂O setpoints covering the full range offered by the dynamic-gravimetric reference gas generation source
10 (2.5–35 ppm), and eight p -levels between 30–250 mbar, for a total of 40 possible combinations (red crosses in Figure 2a). This allows to fully cover the expected variability range of UTLS H₂O. The generated H₂O amount fractions and selected p -levels, along with their relative uncertainties, are listed in Table 1. The conditions generated using the secondary reference gas mixtures for the extended-range validation (blue crosses in Figure 2a) included five H₂O setpoints (between approximately 22–180 ppm) and three p -levels (60, 100, 200 mbar). These measurements allow to assess the linearity of ALBATROSS up to
15 conditions relevant for the upper troposphere, i.e., roughly 6–10 km altitude.

Figure 2b shows a representative time series of the measurement performed at 150 mbar. Each H₂O setpoint generated by the dynamic gravimetric method is measured for 2 h to allow the equilibration of the H₂O amount fraction in the SC-MPC (Figure 2b). The last 30 min of each interval are then selected for the precision assessment. Before and after each experiment, the SC-MPC is purged with zero air for at least 3 h to obtain the "empty-cell" spectrum, i.e. the transmission signal of the SC-MPC in the absence of the reference gas. This is used to normalize the subsequent raw absorption spectra. Despite the substantial
20 purging by dry zero air, the lowest measured H₂O amount fraction in the SC-MPC was about 1.5 ppm after 3 h of purging,

with continuing trend towards lower values. This is mainly due to the strong surface adsorption/desorption properties of H₂O, causing a memory effect of our sampling line.



5 **Figure 2. Overview of the laboratory validation measurements. Panel (a): experimental settings in terms of pressure (p) and H₂O amount fraction levels investigated for the absolute validation (red stars, SI-traceable reference) and the extended-range validation (blue stars, secondary reference) measurements, overlaid with two atmospheric profiles of UTLS H₂O. The profiles correspond to tropical summer conditions (black: Brunamonti et al., 2018) and mid-latitude winter conditions (grey: Graf et al., 2021) and were smoothed by a ± 1 km moving average. Panel (b): time-series of the H₂O amount fraction measured during the individual experiment at 150 mbar gas pressure, retrieved using the qSDVP line shape model at 1 s resolution. Different color shadings indicate zero air measurements (white), reference gas measurements (grey), and the 30 min intervals selected for the data analysis of each H₂O level**
10 **(orange). The estimated H₂O amount fraction content of the zero air (1.46 ppm) is indicated by a black dashed line.**



H₂O amount fraction, Reference levels

<i>Nominal value [ppm]</i>	<i>Actual value [ppm]</i>
2.5	2.51 ± 0.04
5	4.93 ± 0.07
10	9.84 ± 0.14
20	20.05 ± 0.28
35	34.58 ± 0.51

Pressure levels

<i>Nominal value [mbar]</i>	<i>Actual value [mbar]</i>
30	31.52 ± 0.04
45	45.14 ± 0.05
60	61.13 ± 0.07
80	80.67 ± 0.1
100	101.55 ± 0.12
150	151.75 ± 0.18
200	201.96 ± 0.24
250	251.62 ± 0.3

Table 1. Summary of the H₂O amount fraction levels generated by the dynamic-gravimetric method and their expanded measurement uncertainty (top) and pressure levels and their measurement uncertainty (bottom) used for the absolute validation. The relative expanded measurement uncertainty on the H₂O amount fraction levels varies between 1.4–1.47 % for all conditions.

3. Data analysis

5 3.1. Spectroscopic retrieval

The H₂O amount fractions are retrieved from the measured spectra using the Beer-Lambert law in combination with the ideal gas law. Molecular line parameters are taken from the HITRAN2020 database (Gordon et al., 2022), and the wavenumber-dependency of the absorption coefficient is approximated by a line shape model. Typically, the Voigt profile (VP), a convolution of Gaussian (Doppler-broadening) and Lorentzian (pressure broadening) profiles, is considered as a good compromise to capture the pressure broadening effects while offering a good stability and reliability for airborne hygrometers (e.g., Buchholz et al., 2014; 2017; Graf et al. 2021). However, there is growing experimental evidence that various non-Voigt effects (e.g.,



Dicke narrowing, line narrowing due to speed dependence of pressure-induced broadening, and line asymmetry caused by speed dependence of pressure shift) can impact the observed absorption line profile (e.g., Hodges et al., 2008; Kochanov, 2012; Ngo et al., 2012; Tennyson et al., 2014; Lisak et al., 2015). Thus, one major aim of this work was to identify the line shape model that provides the best accuracy and reliability throughout the entire range of UTLS conditions.

5 Recently, the Hartmann–Tran profile (HTP) has been recommended as a new standard in spectroscopic databases (Tennyson et al., 2014). However, this model requires a large number of line-specific parameters, which are difficult to determine considering the correlations between them and the relatively moderate signal-to-noise ratio (SNR) of our measured spectra. Therefore, we investigated lower-order models that could still reproduce the measured spectra at high fidelity (i.e., to nearly the experimental noise level). Our analysis revealed the quadratic speed-dependent Voigt profile (qSDVP), in which the speed-
10 dependence of the relaxation rates are considered as the sole source of line broadening, to be the most suitable.

The qSDVP model is characterized by the phenomenological rate parameters Γ_2 and Δ_2 to describe the quadratic dependence on the active-molecule speed of the pressure-broadening width and shift, in addition to the collisional width and shift (Γ_0 and Δ_0) averaged over all speeds (Tennyson et al., 2014). While the molecular parameters required by the VP line shape model can be readily obtained from the HITRAN2020 database, the corresponding qSDVP parameters for the H₂O transition used here
15 are not available in the literature. Hence, we had to determine these parameters experimentally. The parameter optimization procedure is discussed in detail in Sections 3.3–3.4.

Once the molecular parameters are defined, the H₂O amount fractions are retrieved by minimizing the squared differences between the measured spectra and the line shape model (i.e., the fitting residuals), using a Levenberg-Marquardt least squares algorithm (Press et al., 2007). The overall measurement uncertainty on the amount fractions retrieved by this method, associated with the uncertainty on the measured environmental parameters that are used in the calculation, is estimated to be less
20 than 0.2 %. This results from combining the absolute uncertainties on the measured p (0.12 %) and T (0.06 %), and their variability throughout the measurement periods (standard deviations 0.05 % and 0.01 %, respectively). The uncertainty on the measured OPL (< 1 mm) contributes by less than 0.01 %.

3.2. Pre-processing

25 Prior to the spectroscopic retrieval, the raw spectra are pre-processed. This includes a normalization step, which can be done either by dividing each raw spectrum by the corresponding empty-cell spectrum (i.e., the transmission through the multipass cell filled with zero air), or by reconstructing the laser intensity baseline using a polynomial function (Graf et al., 2021). Then, the time domain is converted to the frequency domain using the transmission spectrum of a 2-inch long Ge-etalon, and the data are interpolated to an equally-spaced wavenumber grid. For this, a free spectral range (FSR) value of 0.02429 cm⁻¹ was
30 determined by optimizing the retrieved peak position of two neighboring absorption lines of N₂O (measured at 100 mbar) to their line positions given in HITRAN2020 (see Figure S1 in Supplementary material). Finally, the number of data points per



spectrum is reduced by a factor of 4, i.e. from $\sim 2 \times 10^4$ to 5×10^3 , by using the moving average approach, resulting in a spectral-point resolution of $1.69 \times 10^{-4} \text{ cm}^{-1}$.

Figure 3 shows an overview of the measured spectra, normalized using the empty-cell transmission approach at H₂O amount fractions of 2.5 ppm, 10 ppm, 35 ppm, and 140 ppm. These spectra were obtained by co-averaging 50 s of data, corresponding to a total number of 1.5×10^5 individual spectra. This choice is justified by the Allan-Werle deviation analysis, discussed in Section 4.1. The SNR, defined as the ratio of the peak absorption signal to the standard deviation of the measured spectrum, calculated in a frequency interval that excludes the absorption line center, is about 50 at 2.5 ppm H₂O (Figure 3a) and 2000 at 140 ppm H₂O (Figure 3d).

As the empty-cell spectrum can only be used in a closed-path configuration, we also tested the polynomial-baseline normalization approach for the case of open-path configuration of ALBATROSS, which is suitable for the evaluation of flight data. Therefore, we apply both normalization methods, and we show in Supplementary material that they provide the same results in terms of accuracy of the retrieved H₂O amount fractions (Figure S2).

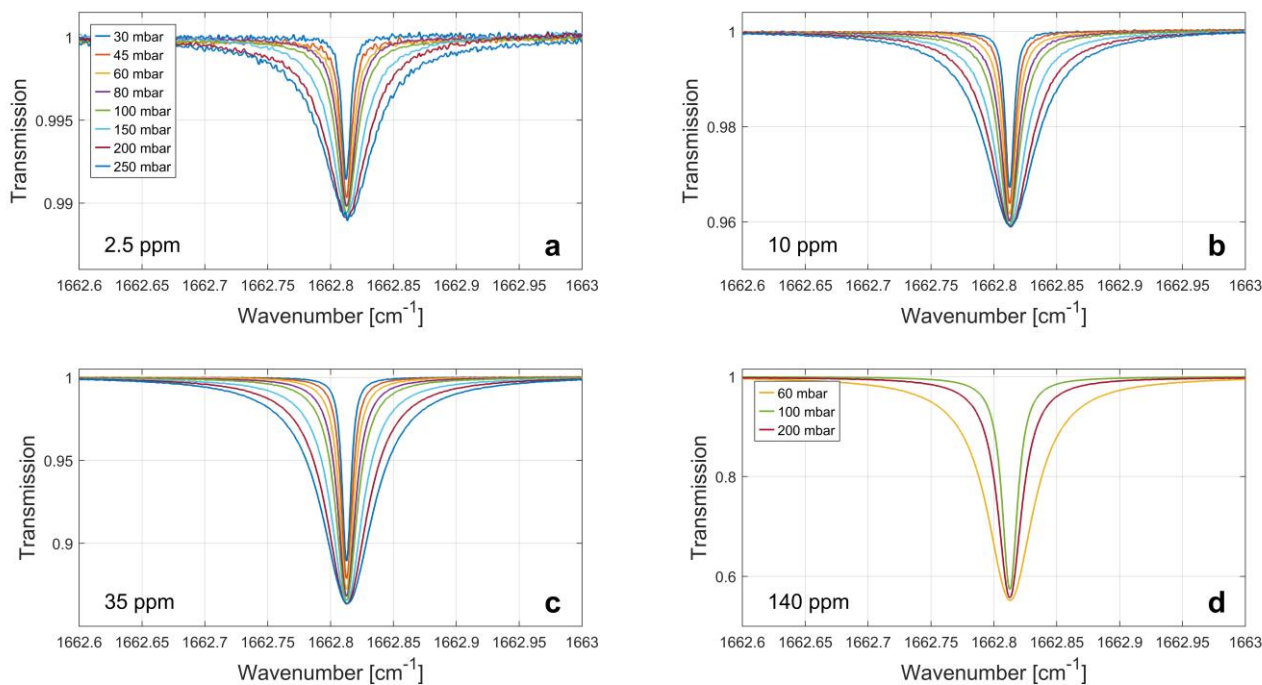


Figure 3. Normalized transmission spectra measured during the laboratory validation experiments at 2.5 ppm (panel a), 10 ppm (b), 35 ppm (c), and 140 ppm H₂O (d), color-coded with gas pressure. The SI-traceable reference gas mixtures (panels a-c) were measured at 8 pressure levels (30–250 mbar), while secondary reference gas mixtures (panel d) at 3 pressure levels (60–200 mbar). All spectra shown here are integrated over 50 s.



3.3. Voigt profile (VP)

As mentioned above, we first evaluated the performance of the VP model using the molecular parameters provided by the HITRAN2020 database, as well as by optimizing these parameters to obtain the best quality of the fit and agreement with the SI-traceable reference. However, neither approach was found to provide a sufficient level of accuracy or a good quality of the fit (QF). Here, the QF is defined, following Cygan et al. (2012), as the ratio of the peak absorption signal to the standard deviation of the fit residuals, calculated over the entire spectrum. This definition accounts not only for the random noise in the experimental spectrum, as in the case of SNR, but also for systematic distortions caused by limitations in the line shape model (Cygan et al., 2012).

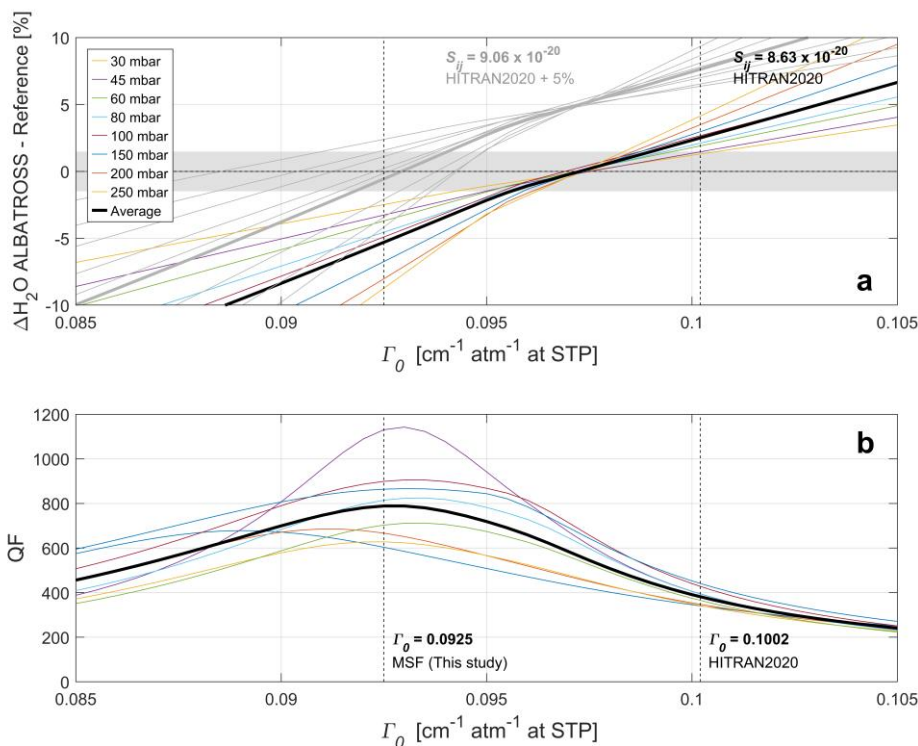
The optimization of the line profile parameters was performed using a multi-spectrum fitting (MSF) approach, in which spectra recorded at different pressures are fitted simultaneously by the least-squares algorithm (e.g., Cygan and Lisak, 2017). This method has the advantage to eliminate the partial correlations between the different line-profile parameters that can occur in least-squares fits to single-pressure spectra. The procedure was implemented in Python using the HITRAN Application Programming Interface (HAPI) routines (Kochanov et al., 2016) as core engine for the spectral fitting.

The performance of the VP was evaluated based on the difference in H₂O amount fraction between the spectroscopic retrievals and the SI-traceable reference values ($\Delta\text{H}_2\text{O}$), and the QF index. Figure 4 illustrates the VP optimization process, in terms of $\Delta\text{H}_2\text{O}$ (panel a) and QF index (b) calculated from fitting the spectra recorded at 35 ppm H₂O and eight different pressures (as shown in Figure 3c), as function of the pressure broadening coefficient (Γ_0). This was obtained by using prescribed values of Γ_0 in the least-squares algorithm. Vertical dashed lines indicate the Γ_0 values from the HITRAN2020 database, namely 0.1002 cm⁻¹ atm⁻¹ at standard temperature (296 K) and pressure (1 atm) (STP conditions), and the value obtained by the MSF approach applied to the spectra measured at 35 ppm H₂O and pressures 30–250 mbar ($\Gamma_0 = 0.0925$ cm⁻¹ atm⁻¹ at STP). All the line shape parameter values used for the retrieval are summarized in Table 2.

As a first step, the retrieval was performed using the Γ_0 and line strength ($S_{ij} = 8.63 \times 10^{-20}$ cm⁻¹/(molecule cm⁻²) at 296 K) parameters specified in HITRAN2020. However, these settings result in an overestimation of the retrieved H₂O amount fractions by up to 5 % compared to the reference value (Figure 4a), and a substantial spread of the deviations obtained at different pressures, indicating an increasing bias correlated with pressure (also shown in Figure 8 below). Therefore, we used the MSF algorithm to optimize the Γ_0 value over the 8 spectra considered here. This leads to the maximum of the QF index (Figure 4b), but in this case, the retrieved H₂O amount fractions underestimate the reference value by up to 9 % (Figure 4a). Finally, we considered the case of varying S_{ij} , as recent studies reported slightly different values of this parameter compared to HITRAN2020 for the H₂O transition considered here (Ptashnik et al., 2016; Conway et al., 2017; Birk et al., 2017; Delahaye et al., 2021). In this case, we observe that increasing the line strength parameter by 5 % compared to HITRAN2020 (consistent with Ptashnik et al., 2016), it is possible to obtain on average a good agreement with the reference values. However, this is not true for each individual pressure level, and a pressure-related bias still affects the results (see spread of grey lines in Figure 4a



at $\Gamma_0 = 0.0925 \text{ cm}^{-1} \text{ atm}^{-1}$ at STP). Hence, we conclude that the VP model is not capable to reproduce the observed spectral characteristics with sufficiently high accuracy.



5 **Figure 4. Voigt profile (VP) optimization. Panel (a):** difference in H_2O amount fraction between the spectroscopic retrievals and the SI-traceable reference values ($\Delta\text{H}_2\text{O}$) at 35 ppm H_2O and different pressures (color-coded), as function of the pressure broadening coefficient (Γ_0). The grey shaded area represents the relative uncertainty of the H_2O reference levels generated by the permeator ($\pm 1.5\%$). Colored lines represent the results obtained using the line strength parameter (S_{ij}) from HITRAN2020, while grey lines using a 5 % higher S_{ij} (given in units of $\text{cm}^{-1}/(\text{molecule} \cdot \text{cm}^{-2})$ at 296 K). Panel (b): quality of the fit (QF) index as function of Γ_0 , obtained from the VP model. Vertical dashed lines indicate the Γ_0 values obtained from the HITRAN2020 database ($0.1002 \text{ cm}^{-1} \text{ atm}^{-1}$ at STP) and by the MSF approach ($0.0925 \text{ cm}^{-1} \text{ atm}^{-1}$ at STP). In both panels, a thick black line shows the average value over all pressure levels.

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3.4. Quadratic speed-dependent Voigt profile (qSDVP)

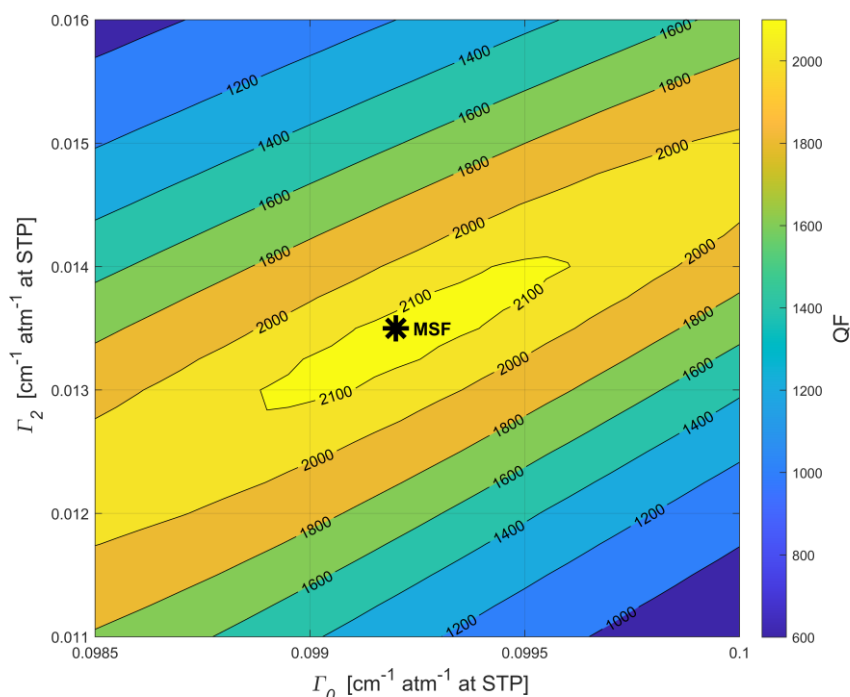
After showing the limitations of the VP model, we proceeded to improve our spectroscopic retrieval algorithm to the qSDVP parameterization. To determine the required parameters, the MSF algorithm was applied to the spectra collected at 140 ppm H₂O and pressure levels 60–200 mbar (shown in Figure 3d). This H₂O level was chosen to achieve the highest possible SNR (i.e., about 2000), which is critical for an accurate estimation of the higher-order line shape parameters. To avoid overfitting, we fixed the Δ_0 and S_{ij} parameters to their values given in HITRAN2020 and set $\Delta_2 = 0$, while the most critical Γ_0 and Γ_2 parameters were determined by the MSF algorithm. As will be shown below, this configuration is highly efficient to fully reproduce the measured spectra at high fidelity, and with high accuracy of the retrieved H₂O amount fractions. The parameters resulting from the MSF calculation are namely $\Gamma_0 = 0.0992 \text{ cm}^{-1} \text{ atm}^{-1}$ and $\Gamma_2 = 0.0135 \text{ cm}^{-1} \text{ atm}^{-1}$ at STP.

To illustrate that this approach provides a robust estimate of the pressure broadening parameters, Figure 5 shows a contour map of the QF index as function of Γ_0 and Γ_2 . This was obtained by consecutively varying Γ_0 and Γ_2 stepwise within a 25×25 grid, centered around their value obtained from the MSF method. The QF index shown here is calculated as the average QF over the three pressure levels measured at 140 ppm H₂O (60, 100, and 200 mbar). The results show that our MSF estimate of Γ_0 and Γ_2 lies in a well-defined maximum of the QF index (i.e., minimum of the standard deviation of the residuals), with values slightly exceeding 2000 (Figure 5). This QF matches the SNR of the measured spectra, which is a strong indication that the line-shape model reproduces the measured spectra to the experimental noise level.

Line shape model parameters

Parameter	Voigt profile (VP)		qSDVP
	HITRAN2020	This study	This study
$S_{ij} [\text{cm}^{-1}/\text{molec cm}^{-2}]$	8.63×10^{-20}	$8.63 \times 10^{-20} (*)$	8.63×10^{-20}
$\Delta_0 [\text{cm}^{-1} \text{ atm}^{-1}]$	3.87×10^{-3}	3.87×10^{-3}	3.87×10^{-3}
$\Gamma_0 [\text{cm}^{-1} \text{ atm}^{-1}]$	0.1002	0.0925	0.0992
$\Delta_2 [\text{cm}^{-1} \text{ atm}^{-1}]$	N.A.	N.A.	0
$\Gamma_2 [\text{cm}^{-1} \text{ atm}^{-1}]$	N.A.	N.A.	0.0135

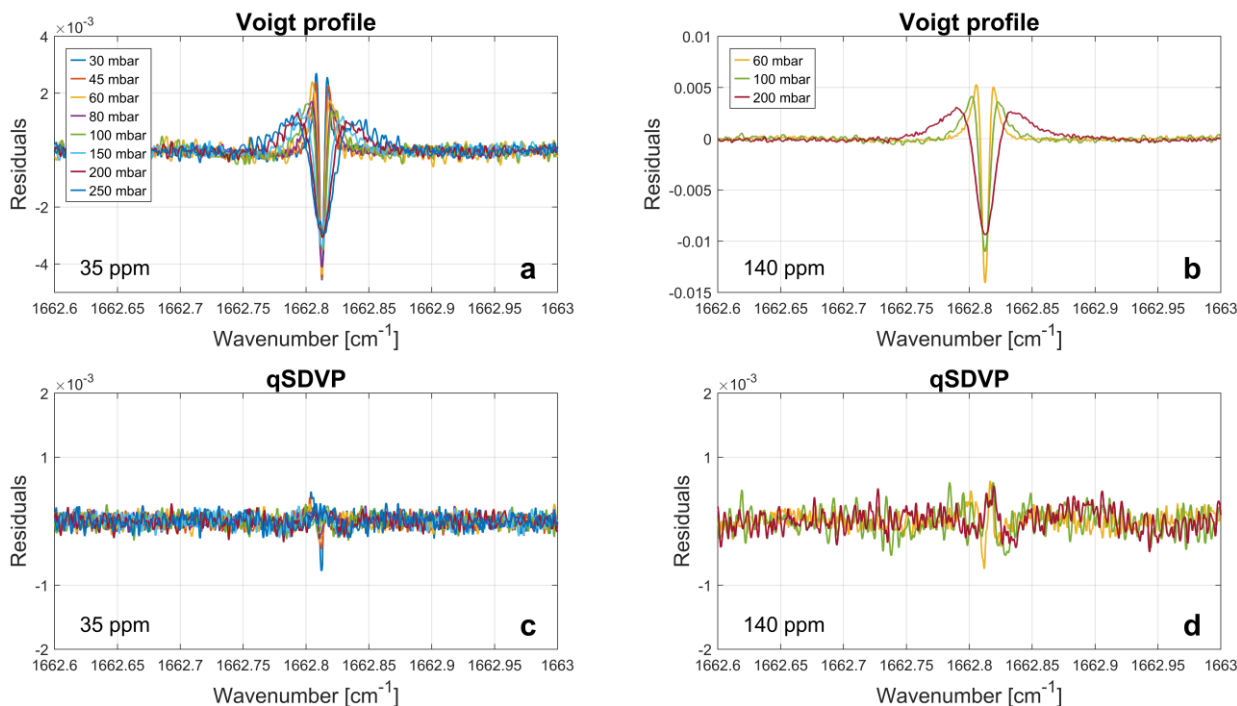
Table 2. Summary of molecular parameters used for the spectroscopic retrieval. For the Voigt profile (VP) model, both the values taken from the HITRAN2020 database and those determined in this study are reported. The quadratic Speed-Dependent Voigt Profile (qSDVP) parameters were determined in this study, as discussed in Section 3. All values are expressed for standard temperature and pressure (STP) conditions (i.e., 296 K, 1 atm). (*) Note that a 5 % higher S_{ij} value was also tested for the VP optimization (Figure 4), and slightly different estimates of S_{ij} for this transition are also reported in the literature (see Section 3.3). The Γ_2 and Δ_2 parameters are not applicable (N.A.) to the VP parameterization.



5 **Figure 5. Determination of the qSDVP fitting parameters. Contour map of quality of the fit (QF) index as function of Γ_0 and Γ_2 , obtained by fitting a qSDVP profile to the spectra recorded at 140 ppm H₂O. The Γ_0 and Γ_2 values obtained from the MSF approach applied to the same spectra ($\Gamma_0 = 0.0992 \text{ cm}^{-1} \text{ atm}^{-1}$, $\Gamma_2 = 0.0135 \text{ cm}^{-1} \text{ atm}^{-1}$ at STP) are indicated by a star. The QF index is calculated as the average QF over the three pressure levels measured at 140 ppm H₂O (60, 100, and 200 mbar). The map was obtained by varying Γ_0 and Γ_2 in a regular 25×25 grid with a resolution of $6 \times 10^{-5} \text{ cm}^{-1} \text{ atm}^{-1}$ in Γ_0 and $2 \times 10^{-4} \text{ cm}^{-1} \text{ atm}^{-1}$ in Γ_2 .**

Finally, Figure 6 shows the fit residuals obtained using the VP (panels a-b) and qSDVP (c-d) line shape models, at 35 ppm H₂O (panels a, c) and 140 ppm H₂O (b, d). The VP retrieval is performed using the line parameters from the HITRAN2020 database, while for the qSDVP we use the Γ_0 and Γ_2 coefficients determined by the MSF method. Using VP yields systematic residuals significantly exceeding the measurement noise and exhibiting the characteristic W-shape, with deviations as large as 1 % of the absorption signal near the line center (Figure 6b). Conversely, using qSDVP, the fit residuals stay below ~0.1 % of the transmission signal for all pressures and H₂O amount fractions (Figure 6c-d). The absence of any structure in the qSDVP residuals beyond the random or optical (i.e., fringe) noise level, justifies the choice of this reduced model against a generalized higher-order (HTP) parameterization.

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15



5 **Figure 6.** Fit residuals obtained using a Voigt profile (VP, panels a-b) and quadratic Speed-Dependent Voigt Profile (qSDVP, panels c-d) line shape model at 35 ppm H₂O (panels a, c) and 140 ppm H₂O (b, d), color-coded with pressure. The VP fit is performed using molecular parameters from the HITRAN2020 database, while the the qSDVP fitting uses the Γ_0 and Γ_2 values determined from the MSF approach.

4. Results

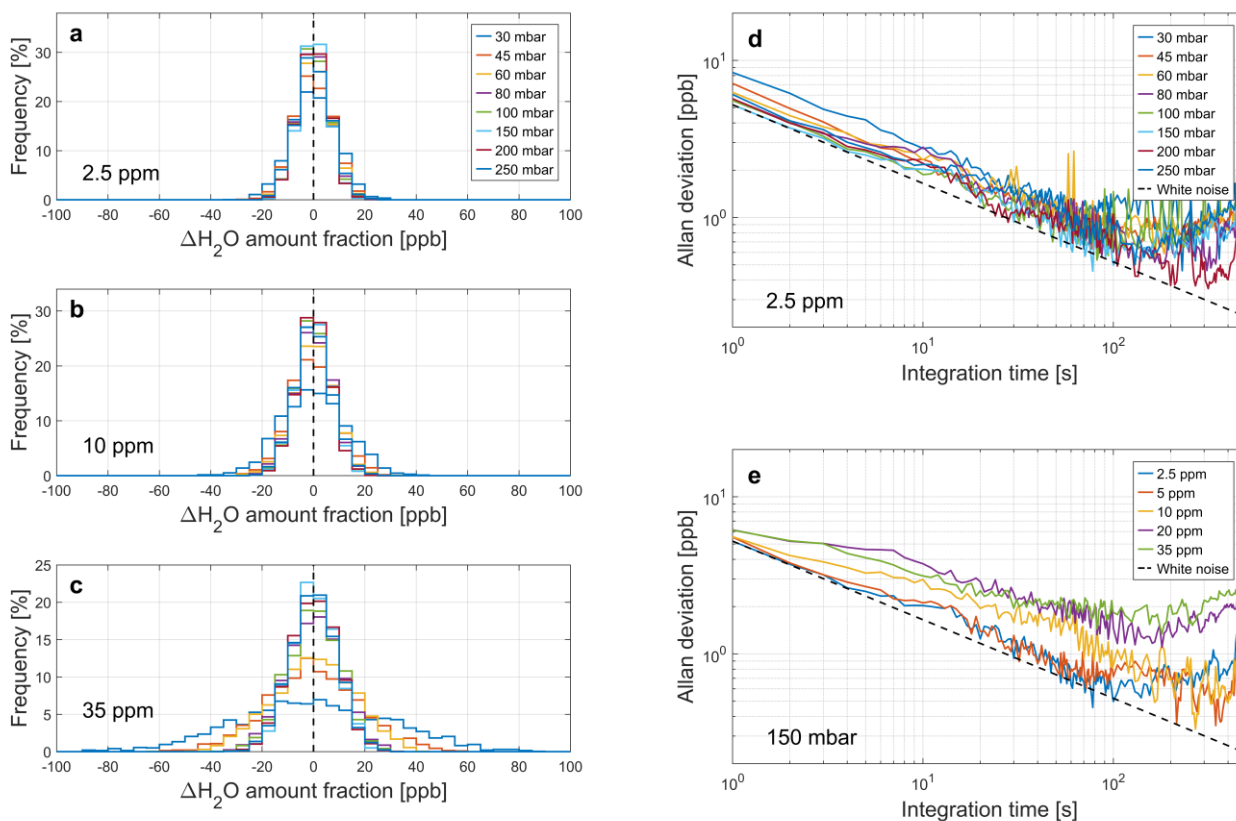
4.1. Precision and long-term stability

10 The precision and long-term stability of ALBATROSS are assessed using 30 min of data recorded at 1 s resolution, taken at the end of each measurement interval (see Figure 2b). In particular, the Allan-Werle deviation technique (Allan, 1966; Werle et al., 1993) was used to determine the measurement precision as a function of integration time, and to distinguish between instrumental drifts and random noise (Werle et al., 1993). The analysis was repeated using both the VP and qSDVP retrievals, revealing no impact on the measurement precision by the choice of line shape model. The results shown in the following were obtained with the qSDVP model and the parameters determined as discussed above.

15 Figure 7 shows frequency of occurrence histograms (left) and Allan-Werle deviation plots (right) for a selection of H₂O amount fractions and pressure levels. The frequency of occurrence distributions, color-coded with pressure, correspond to 2.5, 10, and



35 ppm H₂O (panels a-c). Each distribution is calculated in 40 bins of 5 ppb width, centered around the mean H₂O amount fraction. Allan-Werle deviation as function of integration time are shown for all pressures at 2.5 ppm H₂O (panel d), and for all H₂O setpoints at 150 mbar (e). The theoretical line corresponding to white-noise behavior ($\sim \tau^{-1/2}$ where τ = integration time) is also indicated on panels d-e as a reference.



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Figure 7. Precision and long-term stability assessment. Panels (a-c): frequency of occurrence distributions calculated over 30 min of measurements at 1 s resolution for 2.5 ppm H₂O (a), 10 ppm (b), and 35 ppm (c), color-coded with pressure. Each distribution is calculated using 40 bins of 5 ppb width, centered on the mean H₂O amount fraction measured at the corresponding setpoint and pressure. Panels (d-e): Allan-Werle deviation plots for all pressures at 2.5 ppm H₂O (panel d), and for all H₂O setpoints at pressure level 150 mbar (e). The white-noise behavior ($\sim \tau^{-1/2}$, where τ = integration time) is indicated on panels (d-e) as a reference.

10



The frequency of occurrence distributions are generally very narrow (within ± 30 ppb) at all conditions (Figure 7a-b), with a few exceptions at low pressures and high water amount fractions (Figure 7c). The standard deviation ($\sigma_{I,s}$) of the distributions shown in Figure 7a-c vary between 6–9 ppb (corresponding to 0.23–0.33 % of the reference value) at 2.5 ppm H₂O, 7–13 ppb (0.07–0.13 %) at 10 ppm H₂O, and 9–30 ppb (0.03–0.08 %) at 35 ppm H₂O. On average over all pressures, $\sigma_{I,s}$ lies between 5 0.04 % (14 ppb) at 35 ppm H₂O, and 0.25 % (7 ppb) at 2.5 ppm H₂O.

The Allan-Werle deviation plots indicate a stable operation of the spectrometer up to 100 s, after which drifts start to dominate the measurements. This can be due to mechanical and optical instabilities of the instrument, but also due to flow rate fluctuations of the MFCs. The Allan deviation minimum is always reached at integration times between 50–100 s, corresponding to a precision between 0.5–5 ppb (i.e., 0.02–0.5 % of the reference value). Therefore, an integration time of 50 s was selected for 10 the accuracy assessment and for the determination of the qSDVP parameters.

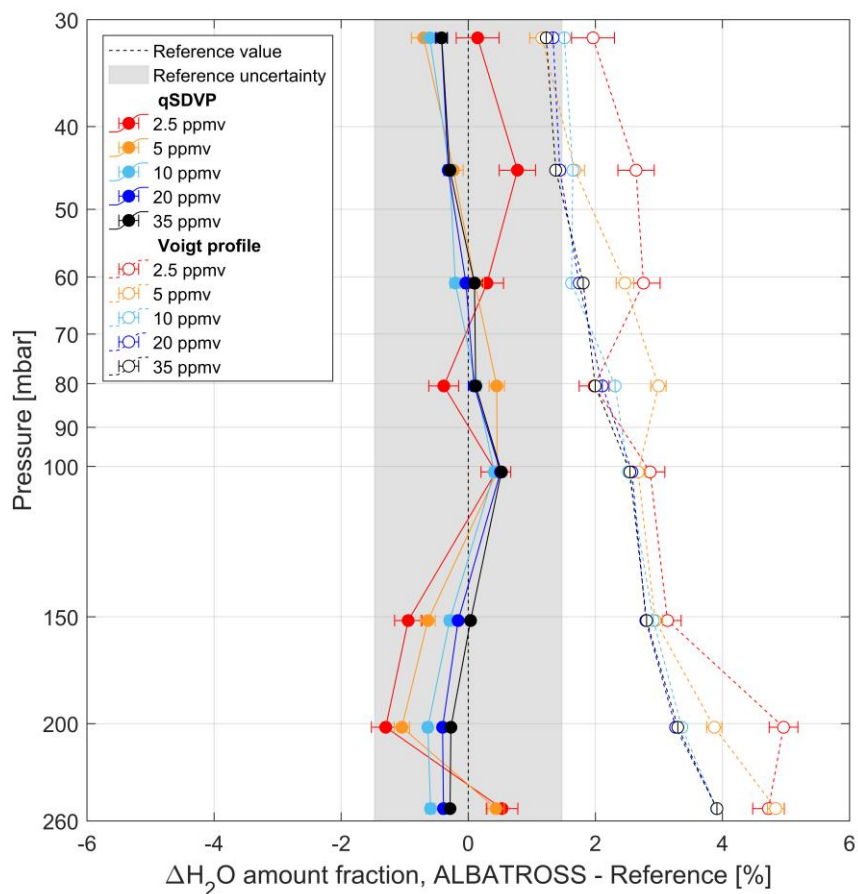
4.2. Accuracy

The accuracy of ALBATROSS is evaluated by comparing the spectroscopically retrieved H₂O amount fractions with the SI-traceable reference H₂O values generated by the dynamic-gravimetric method (as listed in Table 1), at each pressure level. To quantify the improvement in accuracy obtained by the qSDVP parameterization, the results are also compared to those obtained 15 using the VP line shape model and HITRAN2020 molecular parameters.

Figure 8 shows the relative difference in H₂O amount fraction between the ALBATROSS retrievals (integrated over 50 s) and the reference values (ΔH_2O), as function of pressure and color-coded with the H₂O setpoint. The error bars correspond to 1 σ standard deviation. The relative uncertainty on the H₂O reference levels generated by the permeator (± 1.5 %) is indicated by the grey shaded area.

20 The results show that all measurements retrieved using the qSDVP line shape model are found within the ± 1.5 % uncertainty range of the permeator, hence in excellent agreement with the SI-traceable reference values. In contrast, the VP retrievals (with Γ_0 and S_{ij} from HITRAN2020) systematically overestimate the amount fractions by 1.5–5 % compared to the reference, with a bias increasing with pressure.

Using qSDVP, the largest relative deviation from the reference occurs at 2.5 ppm H₂O and 200 mbar (–1.3 %, corresponding 25 to –32 ppb H₂O), while the largest absolute deviation is found at 35 ppm H₂O and 100 mbar (+0.52 %, i.e., +180 ppb H₂O). On average over all pressures, ΔH_2O varies between –0.26 % (i.e., –26 ppb) at 10 ppm H₂O, and –0.06 % (–21 ppb) at 35 ppm H₂O. The pressure-averaged ΔH_2O values and their corresponding standard deviations are summarized in Table 3.



5 **Figure 8. Accuracy assessment. Relative difference in H_2O amount fraction between the spectroscopic retrievals and the reference values (ΔH_2O) as function of pressure, color-coded with H_2O amount fractions. Results obtained using the qSDVP line shape model are shown as filled circles and solid lines, while Voigt profile (VP) results are shown as open circles and dashed lines. The error bars on each data point correspond to the 1σ standard deviation. The expanded relative uncertainty of the H_2O reference levels generated by the dynamic-gravimetric method ($\pm 1.5\%$) is indicated by the grey shaded area.**



Accuracy and precision

Setpoint [ppm]	Absolute deviation [ppb]	Relative deviation [%]
2.5	-2 ± 7	-0.06 ± 0.25
5	-7 ± 7	-0.15 ± 0.14
10	-26 ± 8	-0.26 ± 0.08
20	-28 ± 11	-0.14 ± 0.05
35	-21 ± 14	-0.06 ± 0.04

Table 3. Accuracy and precision of the ALBATROSS measurements performed during the validation. Absolute and relative deviations are expressed as $\Delta\text{H}_2\text{O} \pm 1 \sigma$, where $\Delta\text{H}_2\text{O}$ is the H_2O amount fraction difference between the ALBATROSS retrievals and the SI-traceable reference levels and σ the standard deviation, both averaged over all the considered pressure levels (30–250 mbar).

4.3. Linearity (extended-range validation)

- 5 The aim of the extended-range validation was to characterize the linearity of ALBATROSS beyond the upper limit of 35 ppm H_2O delivered by the dynamic-gravimetric method. For this, the custom-made secondary reference gas, prepared as described in Section 2.3, was diluted to four different H_2O amount fraction levels using two mass flow controllers (MFCs 4-5 in Figure 1). The linearity of the spectrometer is assessed by comparing the retrieved H_2O amount fractions with the calculated amount fractions based on the dilution ratio and the H_2O content of the undiluted reference gas.
- 10 The MFCs (Axetris, Switzerland) were calibrated at METAS using an SI-traceable primary standard with an accuracy of 0.2 % on volumetric flow rate. The dilution ratios are defined as the ratio of the flow rate of the secondary reference gas to the total flow rate (kept constant) of secondary reference and dilution gas. The H_2O amount fraction of the undiluted secondary reference gas, namely 181.47 ± 0.06 ppm, was determined by ALBATROSS at 60 mbar using the qSDVP line shape model and integration time 50 s. The calculated H_2O amount fractions and their corresponding uncertainties obtained for each dilution
- 15 step are listed in Table 4.
- Figure 9 shows a scatter plot of the H_2O amount fractions retrieved by ALBATROSS (using qSDVP and integrated over 50 s) at 60 mbar versus the expected H_2O amount fractions (panel a), and their difference ($\Delta\text{H}_2\text{O}$) at 60, 100, and 200 mbar (panel b). The results of a linear fit (slope, intercept, and determination coefficient R^2) between the measured and the expected H_2O amount fractions at 60 mbar are also displayed in panel a.
- 20 All spectroscopically retrieved H_2O amount fractions are found in very good agreement with the expected values, and with an excellent correlation over the entire investigated range (Figure 9). On average over all pressures, $\Delta\text{H}_2\text{O}$ varies between +0.16 ppm (i.e., +0.76 %) at 22 ppm H_2O , and +1.36 ppm (+0.97 %) at 140 ppm H_2O . The largest relative deviation was found at 48



ppm H₂O and 100 mbar (+1.71 %, i.e., +0.82 ppm). The 1 σ standard deviation varies between 13 ppb (0.06 %) at 22 ppm H₂O, and 30 ppb (0.02 %) at 140 ppm H₂O (see Table 4).

Along with the results obtained with the SI-traceable reference gases in the range of 2.5–35 ppm H₂O, the extended-range validation demonstrates the outstanding performances of ALBATROSS at conditions that fully cover the expected variability of H₂O in the UTLS (see Figure 2a). The accuracy and precision of ALBATROSS in the upper troposphere can be particularly attractive for cirrus clouds microphysical modeling (e.g., Luo et al., 2003; Reinares Martínez et al., 2020) and in-cloud super-saturation studies (e.g., Krämer et al., 2009; Dekoutsidis et al., 2023), requiring accurate H₂O measurements at high vertical/temporal resolution as input parameters.

Linearity (extended-range validation)

<i>Dilution ratio</i>	<i>Calculated H₂O amount fraction [ppm]</i>	<i>Measured H₂O amount fraction [ppm]</i>	<i>Absolute deviation [ppm]</i>	<i>Relative deviation [%]</i>
0.1187 ± 0.0004	21.53 ± 0.07	21.69 ± 0.01	+ 0.16	+ 0.76
0.2637 ± 0.0008	47.86 ± 0.15	48.56 ± 0.02	+ 0.70	+ 1.44
0.5261 ± 0.001	95.47 ± 0.18	96.45 ± 0.02	+ 0.98	+ 1.01
0.7650 ± 0.0007	138.83 ± 0.13	140.19 ± 0.03	+ 1.36	+ 0.97
1	-	181.47 ± 0.06	-	-

Table 4. Summary of the extended-range validation measurements. Dilution ratios and calculated H₂O amount fractions are expressed as expected values ± the uncertainty associated with the volumetric flow rates measurements by the mass-flow controllers. Measured H₂O amount fractions are expressed as mean values ± 1 σ the standard deviation, averaged over the three pressure levels considered here (60, 100 and 200 mbar). Absolute and relative deviations are the difference between measured and calculated H₂O amount fractions. Note that the undiluted H₂O amount fraction of the secondary reference gas (i.e., dilution ratio = 1) was measured by ALBATROSS at 60 mbar and is used to calculate the expected H₂O amount fractions at all levels.

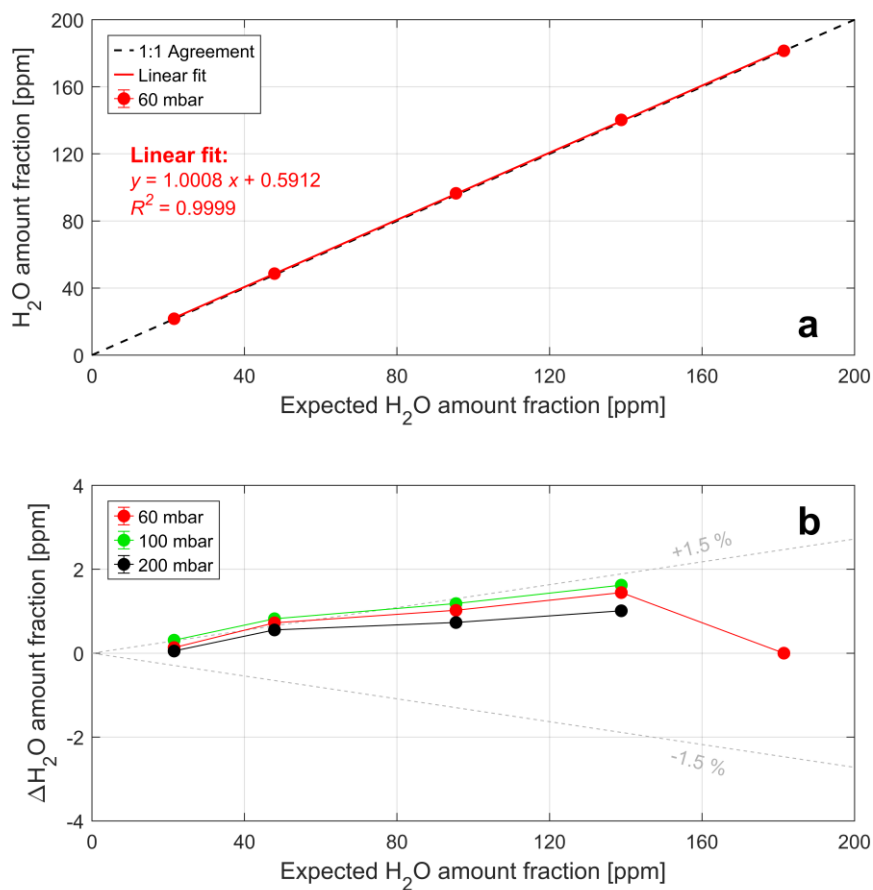


Figure 9. Linearity assessment over the extended-range. Correlation plot of the H₂O amount fractions retrieved by ALBATROSS (y-axis) versus the expected H₂O amount fractions based on dilution ratio (x-axis) at 60 mbar (panel a), and their difference ($\Delta\text{H}_2\text{O}$) as function of the expected amount fractions at 60, 100, and 200 mbar (panel b). The grey dashed lines in panel (b) indicate a relative deviation of $\Delta\text{H}_2\text{O} = \pm 1.5\%$.

5



5. Conclusions

We have presented a detailed laboratory assessment of ALBATROSS, a newly developed quantum-cascade laser absorption spectrometer for balloon-borne measurements of H₂O in the upper troposphere-lower stratosphere (UTLS). The validation was performed using SI-traceable reference gases generated by a dynamic-gravimetric permeation method, capable to deliver H₂O amount fractions as low as 2.5 ppm in synthetic air, with an uncertainty smaller than ±1.5 %. The accuracy and precision of ALBATROSS were evaluated in a wide range of pressure (30–250 mbar) and H₂O amount fractions (2.5–35 ppm), representative of the atmospheric variability of H₂O in the UTLS. In addition, the linearity of ALBATROSS was verified up to 180 ppm H₂O using a custom-made secondary reference gas mixture.

We found that the quadratic-speed-dependent Voigt profile (qSDVP) gives a quality of the fit that is commensurate with our spectral signal-to-noise ratio, and accurately reproduces the measured line shapes without systematic residuals over the entire pressure range. The molecular parameters required by this parameterization (Γ_0 and Γ_2) for the H₂O line investigated here were determined experimentally using a multi-spectrum fitting (MSF) approach over multiple pressure conditions. Furthermore, we demonstrated that the implementation of the qSDVP line shape model, using these empirically determined broadening parameters (given in Table 2) improves the accuracy by up to 5 % compared to the VP model.

The measurements show that ALBATROSS achieves a calibration-free accuracy better than ± 1.5 % with respect to the SI-traceable reference at all investigated pressures and H₂O amount fractions. The measurement precision at 1 s resolution was found to be better than 30 ppb (i.e., 0.1 % at 35 ppm H₂O) at all conditions, and as low as 5 ppb (0.02 % at 35 ppm H₂O) upon integrating the data over 50 s, corresponding to the Allan deviation minimum. These results are particularly remarkable considering the technical challenges of maintaining a stable H₂O amount fraction level in the low-ppm range in a laboratory setting, due to the strong surface adsorption/desorption properties of H₂O.

The performance achieved by ALBATROSS is unprecedented for a light-weight balloon-borne hygrometer, and thus demonstrate the exceptional potential of mid-IR laser absorption spectroscopy as an alternative reference method to cryogenic frost-point hygrometry (CFH) for in-situ measurements of H₂O in the UTLS. This is particularly relevant considering the ongoing phasing out of the cooling agent required by CFH (fluoroform, HFC-23), which urges the need of an alternative solution to maintain the monitoring of UTLS H₂O in global, long-term monitoring networks, such as the GCOS Reference Upper Air Network (GRUAN).

Currently, ALBATROSS is participating in the AquaVIT-4 International Intercomparison of Atmospheric Hygrometers, held at the AIDA cloud simulation chamber (Karlsruhe Institute of Technology, Germany). This will allow to characterize the performance of the spectrometer under a new set of challenging UTLS-relevant conditions, also in terms of temperature, and to validate its accuracy against a set of well-established hygrometers (both laboratory-based and airborne). Finally, new inter-comparison test flights between CFH and ALBATROSS are foreseen to further validate the performances achieved in the laboratory in a real atmospheric environment.



Data availability. The data are available from the authors upon request.

Competing interests. The authors declare that they have no conflict of interest.

Author contributions. SB, MG and TB performed the measurements, under the supervision of BT and CP. SB performed the data analysis with the support from MG, II and BT. BT and LE supervised the project. SB and BT wrote the manuscript with contributions from all authors.

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