



1 **Adaptation of RainGaugeQC algorithms for quality control of** 2 **rain gauge data from professional and non-professional** 3 **measurement networks** 4

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9 **Abstract.** Rain gauge measurements are one of the primary techniques used to estimate a precipitation field, but
10 they require careful quality control. This paper describes a modified RainGaugeQC system, which is applied to
11 real-time quality control of rain gauge measurements made every 10-min. This system works operationally at the
12 national meteorological and hydrological service in Poland. The RainGaugeQC algorithms, which have been
13 significantly modified, are described in detail. The modifications were made primarily to control data from non-
14 professional measurement networks, which may be of lower quality than professional data, especially in the case
15 of private stations. Accordingly, the modifications went in the direction of performing more sophisticated data
16 control, applying weather radar data and taking into account various aspects of data quality, such as consistency
17 analysis of data time series, bias detection, etc. The effectiveness of the modified system was verified based on
18 independent measurement data from manual rain gauges, which are considered one of the most accurate
19 measurement instruments, although they mostly provide daily totals. In addition, an analysis of two case studies
20 is presented. This highlights various issues involved in using non-professional data to generate multi-source
21 estimates of the precipitation field.

22 **1 Introduction**

23 **1.1 Precipitation measurements**

24 Precipitation is one of the most important meteorological parameters – due to its great practical importance in
25 water management, flood control and other issues (e.g., Loritz et al., 2021; Sokol et al., 2021). For this reason,
26 conducting measurements and estimating the precipitation field are very important tasks, though also very
27 challenging because of the very high temporal and spatial variability of precipitation and its intermittent nature.
28 The shorter the accumulation time of measurements, the greater the spatial variability of an estimated precipitation
29 field and the greater its uncertainty (Berndt and Haberlandt, 2018; Bárdossy et al., 2021). This is especially true
30 when estimating sub-daily totals and, even more the case for sub-hourly precipitation totals.

31 Until today, the basic measurements of precipitation are in situ measurements carried out by means of rain
32 gauge networks, and this does not change despite the intensive development of remote sensing techniques, such
33 as radar and satellite, from which measurements are highly distorted. Rain gauge measurements are still considered
34 the most accurate, although they are limited to specific, rather sparsely distributed points. Consequently, when
35 estimating the precipitation field, measurement data provided by different techniques are treated as independent
36 estimates of the same physical quantity. Thus, the final estimate of a precipitation field, which is often referred to



37 as quantitative precipitation estimation (QPE), is determined using various methods of combining data from
38 different sources (multi-source estimation), taking into account the strengths and weaknesses of each of these
39 techniques (McKee et al., 2016; Jurczyk et al., 2020b).

40 Since all measurement techniques are subject to significant errors, which have a different temporal and spatial
41 structure, all rainfall measurements need advanced quality control (QC) (Szturc et al., 2022). This applies not only
42 to weather radar measurements (Ośródko et al., 2014; Ośródko and Szturc, 2022; Sokol et al., 2021), but also to
43 rain gauge measurements. The latter are considered accurate at their locations, however field experiments (Wood
44 et al., 2000) and experiences with dual-sensor rain gauges (Ośródko et al., 2022) show that trust in rain gauges is
45 often excessive – errors in their measurements can sometimes be very significant.

46 Quality control of rain gauge data is carried out using various approaches, most commonly by analysing the
47 spatial and temporal distribution of measurements. As such information is insufficient for effective QC, especially
48 in the case of sparse measurement networks, external data from other measurement techniques, most often weather
49 radar and satellite, are often used (Ośródko et al., 2022; Yan et al., 2024). Increasingly, deep learning techniques
50 are also being applied for QC (Sha et al., 2021). It should be noted that QC applied to short rainfall totals, such as
51 the 10-min employed in this work, is considerably more difficult than for longer totals, such as 1-h (Villalobos-
52 Herrera et al., 2022).

53 Due to the particular importance of rain gauge measurements, especially for the adjustment (calibration) of
54 radar and satellite measurements, it is crucial when estimating the precipitation field that rain gauge networks are
55 as dense as possible (Hohmann et al., 2021). This implies a very high financial as well as technical and
56 organisational effort, so that a great deal of work is currently being done to deliver rain gauge data from other
57 networks, not only from the national meteorological and hydrological services (NMHSs). A separate issue is the
58 employment of “opportunistic” measurement techniques, i.e. precipitation data acquired from devices not
59 dedicated to rainfall measurement, e.g. by analysing the attenuation of signals in commercial microwave links
60 used in mobile phone networks, see e.g.: Chwala and Kunstmann (2019), Polz et al. (2020), Graf et al. (2021),
61 Pasierb et al., (2024).

62 **1.2 Non-professional rain gauge networks**

63 Apart from the rain gauge networks of the NMHSs, measurement networks set up and maintained by various
64 institutions – usually state or local authorities taking measurements for their own purposes – can also be a source
65 of rain gauge data. Another possibility is collecting meteorological measurements carried out by individual people
66 with generally low-cost measuring stations, for whom taking measurements, analysing them and comparing with
67 data generated by meteorological services is a hobby activity (Muller et al., 2015; Krennert et al., 2018; Zheng et
68 al., 2018). These are so-called private or citizen weather stations (PWS or CWS).

69 For the purposes of this paper, all measurements carried out by institutions other than NMHSs are considered
70 “non-professional” because they do not guarantee compliance with the standards set by the World Meteorological
71 Organisation (WMO) (WMO-No. 488, 2010) to the same extent as NMHSs measurements. A distinction between
72 professional and non-professional rain gauges has been proposed by, among others, Garcia-Marti et al. (2023). In
73 addition, the aforementioned private stations set up by individual hobbyists need to be distinguished, as direct
74 control of the location, technical conditions or maintenance of such stations is impossible in practice. Such stations
75 should be treated with much less trust, and the high uncertainty of the data is due to a number of reasons, which



76 have been described in detail in the literature (np. Bell et al., 2015; Båserud et al., 2020; Hahn et al., 2022; Urban
77 et al., 2024). Nevertheless, many studies show that such data can be a valuable source of precipitation information
78 (de Vos et al., 2017; 2019; Horita et al., 2018; Nipen et al., 2020; Bárdossy et al., 2021), thanks to the relatively
79 very large number of these stations especially in urban areas, bearing in mind that professional gauges are typically
80 located outside city centres (Overeem et al., 2024).

81 The incorporation of non-professional data is associated with some overall increase in uncertainty in
82 precipitation data. Moreover, dual-sensor rain gauges are rarely used and this reduces the efficiency of the quality
83 control performed. Consequently, QC algorithms for these data should include not only the filtering out of clearly
84 erroneous measurements and a decrease of their quality metric in the form of e.g. a quality index (QI), but for less
85 supervised networks it is also necessary to correct at least the systematic errors associated with the bias of these
86 measurements.

87 1.3 Overview of approaches to QC of rain gauge data

88 The specificity of data from non-professional rain gauges is primarily due to the greater uncertainty of their
89 measurements. This entails the development of more sophisticated, but also more restrictive quality control
90 algorithms. These are generally extensions of the QC methods applied to data from NMHSs, but here they analyse
91 the reliability of individual measurements in more depth. These methods most often rely on verification with
92 professional rain gauges, but also use other measurement data, especially weather radar data.

93 *Spatial distribution of precipitation measurements* – detection of inconsistencies with surroundings. The most
94 common quality control techniques involve checking whether the deviation from the reference measurements,
95 which can also be data from nearby rain gauges, are within preset threshold values. If a measurement exceeds the
96 threshold, then it is treated as an outlier and either its quality index QI (or quality flag) is decreased or the
97 measurement is rejected (de Vos et al., 2017; Båserud et al., 2020). In addition, precipitation data from other
98 sources, primarily weather radar, can be used to quantify the uncertainty of outlying measurement data (Ośródk
99 and Szturc, 2022). Spatial consistency tests are very difficult to perform for a sparse rain gauge network, so the
100 QC in terms of spatial consistency may not be carried out, and in the case of private rain gauges, such data may
101 simply be rejected (Nipen et al., 2020). Alerskans et al. (2022) used a cost function based on a contingency table,
102 which optimises the parameters of the spatial QC algorithm used to detect as many actually erroneous data as
103 possible, while minimising the number of correct data that were found to be erroneous.

104 *Correlation of time series of precipitation measurements with reference data.* The temporal consistency check
105 involves detecting stations from which measurements often have relatively low reliability, but not so much that
106 individual measurements do not pass a spatial consistency check. Analysis of the temporal consistency of rainfall
107 data is most often carried out by analysing the correlation of the time series from the controlled rain gauge with
108 the time series of reference data (Bárdossy et al., 2021; de Vos et al., 2019). Reference data can be either data from
109 professional rain gauges of relatively high quality or from other measurement techniques, primarily weather radar
110 (de Vos et al., 2019). However, the use of radar data is associated with difficulties, most often due to errors in
111 estimation of the precipitation field (Ośródk et al., 2014; Ośródk and Szturc, 2022). Moreover, weather radar
112 measurements are performed at certain heights above the ground surface – from a few hundred metres to as much
113 as a few kilometres – and are then spatially averaged. Analysis of the correlation coefficient of a time series
114 becomes difficult, especially in cases where the rain gauge reports false zero values (no precipitation) due to e.g.



115 a sensor being blocked or some object obstructing the path of the rain (e.g. buildings, vegetation). Another
116 difficulty is caused by non-rainfall periods – time series with predominantly very low rainfall can sometimes
117 disturb the correlation (Hahn et al., 2022).

118 *Detection and removal of bias in precipitation measurements.* The approaches to the issue of quality control
119 of rain gauge data described above do not correct erroneously measured values, but only reduce their *QI* or remove
120 them. However, data correction is an important part of data quality control. First of all, it is about bias correction
121 (unbiasing), which most often results from rainfall underestimation related to rain gauge technology: rain gauge
122 measurements are underestimated due to wind-induced errors, wetting losses, evaporation losses, trace
123 precipitation, etc. The magnitude of the underestimation also depends on the construction of the rain gauge; in
124 particular, tipping-bucket devices are subject to significant bias (Segovia-Cardozo, 2021). This bias can be
125 eliminated, or at least reduced, by, for example, quantitatively analysing all underestimation factors and
126 introducing all important corrections (Zhang et al., 2019). Such adjustments, however, are generally conservative
127 because of the difficulty of considering all relevant factors and the lack of precise data on influencing parameters.
128 Another way is to compare non-professional measurements with a benchmark as reliable as possible, which could
129 be manual rain gauges, preferably lysimetric ones that measure at ground level (Haselow et al., 2019; Schnepper
130 et al., 2023). However, such measurements are not common. Radar observations are more widely available but
131 using them as a benchmark requires the QC and adjustment to professional rain gauge measurements to have been
132 previously carried out. Unbiasing is also calculated on the basis of a larger data set collected during precipitation
133 events typical of the local climate (np. de Vos et al., 2019). The bias factor determined on this basis is treated as a
134 climatological quantity.

135 **1.4 Structure of the paper**

136 This paper presents the RainGaugeQC system (Ośródko et al., 2022) after its adaptation for quality control of rain
137 gauge data from non-professional stations. The paper is structured as follows: after Section 1, Section 2 briefly
138 describes the different kinds of precipitation data on which the RainGaugeQC was developed and verified. Section
139 3 presents the algorithms of the RainGaugeQC system with the emphasis on solutions that are more advanced
140 when compared to the earlier version of the system. Results obtained over several months, as well as analysis of
141 two case studies, were discussed in Section 4. Section 5 summarises the paper with a list of conclusions resulting
142 from the use of the modified RainGaugeQC system.

143 **2 Precipitation data**

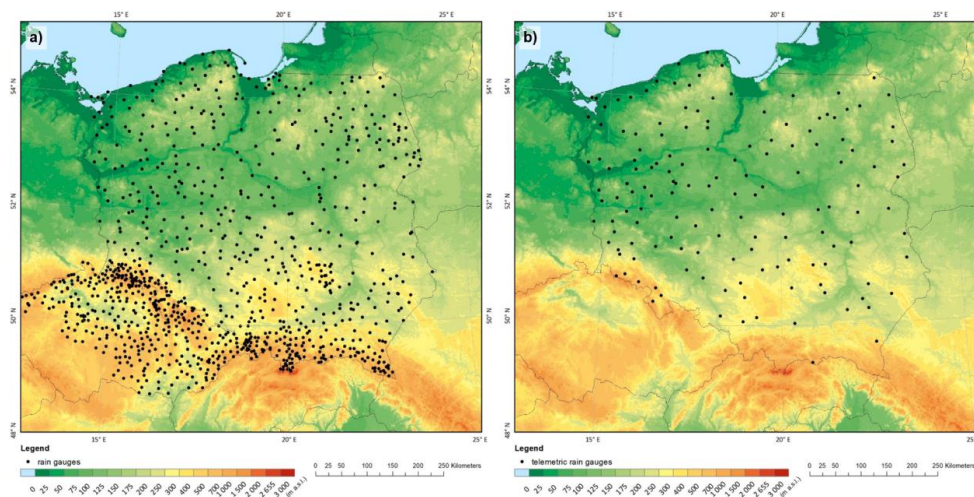
144 **2.1 Available networks of rain gauges**

145 IMGW operationally utilises telemetric rain gauge data from the following measurement networks operated by
146 (Fig. 1):

- 147 • IMGW (Institute of Meteorology and Water Management – National Research Institute) – network of
148 NMHS in Poland (<https://hydro.imgw.pl/#/map>).
- 149 • CHMU (Czech Hydrometeorological Institute) – network of NMHS in the Czech Republic. IMGW uses
150 data from more than 324 stations near the Polish border
151 (https://www.chmi.cz/files/portal/docs/meteo/ok/images/srazkomerne_stanice_en.gif).



152 • General Directorate of the State Forests (DLP) – network of the meteorological monitoring program of
 153 forest areas consisting of 145 stations (<https://www.traxelektronik.pl/pogoda/las/>).
 154 The above data are used to generate operationally (in real-time) a multi-source precipitation field with high
 155 spatial resolution, which is the basis for generating nowcasting precipitation forecasts.
 156 Synthetic information about the above networks is summarised in Table 1.
 157



158
 159 **Figure 1: Telemetric rain gauge networks: a) IMGW and CHMU, b) State Forests.**

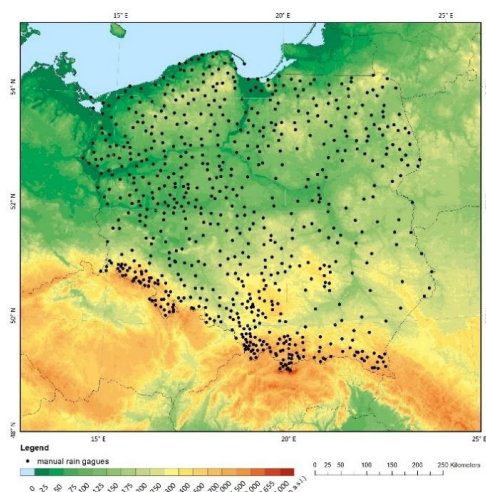
160
 161 **Table 1. Rain gauge networks incorporated into operational processing by RainGaugeQC system and**
 162 **estimation of precipitation field (as of October 2024).**

ID	Network operator	Number of stations	Type of rain gauges	Type of network
1	IMGW	656	Mostly two tipping bucket sensors	Professional
2	CHMU	324 stations located close to Polish territory	Mostly tipping bucket sensors	Professional
3	General Directorate of the State Forests (DLP)	145	Heated	Non-professional

163
 164 For the domain of Poland data from professional rain gauge networks operated by NMHSs in Poland (IMGW)
 165 and the Czech Republic (CHMU) are available. As the territory of the Czech Republic covers a large part of the
 166 analysed domain and, above all, a significant number of rain gauges are located close to mountainous areas on the
 167 border with Poland (Fig. 1), these data are very important for improving the reliability of the estimation of the
 168 precipitation field in southern Poland. The third network, belonging to the State Forestry Authority, is a non-
 169 professional research network so it is uncertain whether all the standards of the WMO recommendations are
 170 followed (WMO-No. 488, 2010).



171 The quality of precipitation data is highly dependent on the type of measuring devices being used. Currently,
172 the IMGW network is still dominated by tipping-bucket type rain gauges, which are considered significantly less
173 accurate than weighing rain gauges (e.g., Colli et al., 2014; Hoffmann et al., 2016).
174



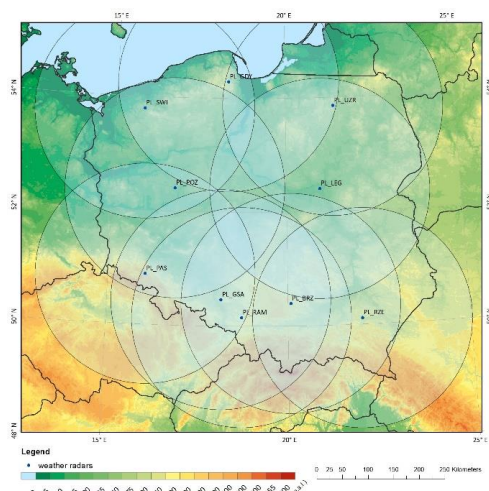
175
176 **Figure 2: IMGW's network of manual rain gauges.**

177
178 A network of Hellmann-type manual rain gauges, providing independent reference data, is used in this study
179 to verify the performance of the developed QC algorithms. As the data from these rain gauges are not available in
180 real time, they cannot be used for rainfall field estimation or operational QC of telemetric data. The IMGW network
181 consists of about 641 manual rain gauges, which provide daily rainfall accumulations (Fig. 2). These data are
182 believed to be much more accurate than measurements from telemetric rain gauges, especially tipping-bucket ones.
183 They are subjected to manual QC before being used.

184 2.2 Weather radar precipitation data

185 Precipitation data from weather radars play a major role in the RainGaugeQC system for quality control of rain
186 gauge data (Ośródka et al., 2022). The data used in this study are provided by the Polish POLRAD radar network
187 operated by IMGW. The network consists of ten Doppler polarimetric radars working in C-band, manufactured by
188 Leonardo Germany (Fig. 3). Three-dimensional raw data and two-dimensional products are generated by the
189 Rainbow 5 system every 5 min with 1-km spatial resolution and a range of 250 km.

190



191

192 **Figure 3: Computational domain of Poland with plotted 215-km ranges of weather radars of the Polish POLRAD radar**
193 **network (as of July 2024).**

194

195 The raw 3D radar data are quality controlled and corrected by the RADVOL-QC system (Ośródka et al., 2014;
196 Ośródka and Szturc, 2022). The product used to estimate the rainfall field is PseudoSRI (Pseudo Surface Rainfall
197 Intensity): cut-off at 1-km altitude above ground and from the lowest elevation out of the SRI range, generated
198 every 5 min and accumulated into 10-min sums taking into account spatio-temporal interpolation between two
199 adjacent measurements. As a result of quality control with the RADVOL-QC system, the corresponding *QI* quality
200 index fields are also assigned to the individual estimated precipitation fields. In addition, some kind of quality
201 control of radar precipitation takes place at the stage when data from individual radars is combined into composite
202 maps. This is done by means of an algorithm that takes into account the time-varying spatial distribution of the
203 quality index (Jurczyk et al., 2020a).

204 Due to the bias present in the weather radar observations, these data are adjusted with rain gauge data, but only
205 from the professional networks, derived from the 1-h moving window. However, if a precipitation accumulation
206 is below a preset threshold, then this period is extended accordingly, up to a maximum of the seasonal
207 accumulation.

208 **2.3 Multi-source precipitation estimates RainGRS**

209 Multi-source precipitation field estimates are generated by the RainGRS system of IMGW. The system combines
210 rain gauge, weather radar, and satellite precipitation data in real time (Szturc et al., 2018; Jurczyk et al., 2020b;
211 2023). The algorithm for combining these data is based on conditional merging according to an algorithm proposed
212 by Sinclair and Pegram (2005), which attempts to enhance the strengths and reduce weaknesses of individual
213 measurement techniques. This approach was modified in RainGRS by taking into account the quantitative
214 information about the spatial distribution of the quality of the individual input data (quantified by *QI*). These
215 estimates are produced every 10 min with a high spatial resolution of 1 km x 1 km.



216 In the study, two versions of multi-source RainGRS precipitation estimates are generated in order to examine
 217 the impact of incorporating non-professional data. In the first version, rain gauge data only from the professional
 218 networks of IMGW and CHMU were taken, while in the second version, data from the non-professional network
 219 of the State Forests were added to this set.

220 3 RainGaugeQC system for QC of rain gauge data

221 3.1 RainGaugeQC system for QC of rain gauge data from a professional network

222 The RainGaugeQC system was originally designed to perform real-time quality control of rain gauge data from
 223 measurement networks maintained by IMGW. This system was described in detail in work by Ośródko et al.
 224 (2022), so in this study, after a very concise presentation of the algorithms, the following sections will describe
 225 only modifications made to adapt it to data from non-professional networks.

226 In the standard version of RainGaugeQC (Ośródko et al., 2022) (see column “Before modification” in Table
 227 2), firstly the simple plausibility tests – the gross error check (GEC) and range check (RC) – were performed on
 228 individual measurements. Then the more complex checks were conducted using a larger amount of rain gauge data
 229 from either a specific time range or a specific area, as well as using external data provided by weather radars.
 230 Firstly, the Radar Conformity Check (RCC) was applied to identify false precipitation on the basis of the radar
 231 measurements. Obstruction or blocking of the sensors was also checked for. Next, the Temporal Consistency
 232 Check (TCC) was performed, but this version was designed only for dual-sensor stations: data from the pairs of
 233 rain gauge sensors were tested for the existence of significant differences between them. The most advanced
 234 algorithm was the Spatial Consistency Check (SCC) which identified outliers by comparing observed values with
 235 data from neighbouring stations.

236 An important outcome of the system was the determination of the quality index (QI) of analysed data, which
 237 is a unitless value with a range [0.0, 1.0], where “0.0” means extremally bad data and “1.0” means perfect data.
 238 This QI metric was determined by the RainGaugeQC for each sensor and then the sensor with higher quality is
 239 taken for further processing.

240

241 **Table 2. A summary of the quality control algorithms used in the RainGaugeQC system before and after**
 242 **modification.**

Abbr.	Algorithm	Before modification	After modification
GEC	Gross Error Check	Gross errors	
RC	Range Check	Exceeding climatological thresholds	
RCC	Radar Conformity Check	Detection of false rainy and non-rainy events	
BSC	Blocked Sensor Check	Detection of blocked sensors	
TCC	Temporal Consistency Check	Comparison of two sensors	Time series comparison with weather radar data
BC	Bias Check	–	Detection and correction of bias with adjusted radar data



SCC	Spatial Consistency Check	Detection of outliers from the local vicinity	Detection of outliers from the local vicinity (updated)
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243

244 3.2 Directions of development in RainGaugeQC

245 The possibility of incorporating non-professional data at IMGW became a motivation for more sophisticated data
 246 quality control. The QC algorithms in the previous version of RainGaugeQC proved unsuitable for non-
 247 professional data, as they are often subject to greater uncertainty than from professional rain gauges, and besides,
 248 these gauges are generally not dual-sensor. On the other hand, the inclusion of new data significantly improved
 249 the performance of the SCC algorithm due to the higher density of the measurement network. Therefore, it was
 250 necessary to redesign the RainGaugeQC system in order to adapt it to rain gauge networks equipped with different
 251 types of sensors, supervised to various degrees, so that the system became more universal. The modified algorithms
 252 tailored to the new challenges associated with incorporating non-professional data are summarised in Table 2 in
 253 the “After modification” column.

254 *TCC.* In the new version of the TCC (time series comparison with adjusted weather radar data) algorithm,
 255 weather radar data is used to compare time series from a specific time interval to check the correlation between
 256 rain gauge measurements and radar observations. The correlation coefficient is used as a metric for the relevant
 257 component of the quality index of the rain gauge data. This allows a reduction in the data quality index of rain
 258 gauges with measurements disturbed for a certain time period due to failure, poor maintenance or bad location.

259 *BC.* The above TCC algorithm is not sensitive to the bias of rain gauge measurements, so the BC (bias check
 260 with adjusted radar data) algorithm is used to detect bias in the data. It also works by analysing long-term data
 261 series, but in this case they are used to compare data accumulations from rain gauges with radar accumulations.
 262 The quantitative estimation of the bias of the rain gauge data allows relevant components of the quality index to
 263 be determined. In the case of private rain gauges, unbiasing is carried out as well as reducing the *QI* value.

264 *SCC.* The SCC (detection of outliers from the local vicinity) algorithm was already introduced in the first
 265 version of the RainGaugeQC system, but significant modifications have been made to the current version. It detects
 266 outliers, i.e. the measurements at a given time-step deviate from the values from rain gauges located in a certain
 267 area. The increase in the number of rain gauges through incorporating non-professional data has made it easier to
 268 determine the degree of outlying for individual data. The quality index reduction for outliers is quantified on the
 269 basis of the spatial variability of the precipitation field derived from the radar data.

270 3.3 New version of TCC algorithm (Time series comparison with weather radar data)

271 The TCC algorithm is designed to eliminate erroneous rain gauge measurements (*G*) by analysing the correlation
 272 on long time series. The reference is radar precipitation (*R*) after adjustment with rain gauge observations only
 273 from professional networks.

274 For the calculation, pairs of rain gauge (*G*) and radar (*R*) data are taken if at least one of the values is greater
 275 than 0.025 mm, and their quality index (*QI*) is at least 0.7 for *G* and 0.8 for *R*. Two time series aggregated from
 276 10-min accumulations: “long” and “short” comprising 10 and 5 days, respectively, are analysed. For long series
 277 the number of non-precipitation pairs c_{dry} is determined provided that both values are less than 0.025 mm. For



278 each series hourly accumulations are determined and then the number of measurement pairs c and correlation
279 coefficient r are calculated.

280 The procedure for assessing data quality is carried out by checking a list of conditions. For a given measurement
281 these conditions are examined sequentially and, depending on the result, further ones are checked or the quality
282 index is reduced accordingly.

283 The check is stopped if the accumulations of both radar and rain gauge precipitation for the long series are
284 below the assumed threshold values:

$$285 \quad (\sum_{10 \text{ days}}(R) < 3.0) \text{ and } (\sum_{10 \text{ days}}(G) < 6.0) \rightarrow \text{TCC stopped} \quad (1)$$

286 If the amount of radar precipitation for the long series is below the assumed threshold and the amount of rain
287 gauge precipitation is above the corresponding threshold, then the check is also stopped, but the quality of the rain
288 gauge data is reduced by a value of 0.05:

$$289 \quad (\sum_{10 \text{ days}}(R) < 3.0) \text{ and } (\sum_{10 \text{ days}}(G) \geq 6.0) \rightarrow \text{TCC stopped, } QI = QI - 0.05 \quad (2)$$

290 This indicates that there are large differences between the two accumulations, but because the rainfall recorded
291 by the radar is too low, the calculation of the correlation coefficient may be not reliable in such cases.

292 The check is passed if the number of measurement pairs is above the preset threshold and correlation coefficient
293 is above 0.3 for short or long series. Then the quality index is reduced on the basis of the relevant correlation
294 coefficient, according to the following formula:

$$295 \quad (c > 6) \text{ and } (r > 0.3) \rightarrow \text{TCC passed, } QI = \begin{cases} QI & r > 0.85 \\ QI - \frac{1-r}{4} & r \leq 0.85 \end{cases} \quad (3)$$

296 If there is an insufficient number of measurements for short series and at the same time the number of non-
297 precipitation data pairs is above a preset threshold, indicating that there is a longer non-precipitation period, then
298 the TCC is stopped and QI is reduced:

$$299 \quad (c_{dry} > 1000) \text{ and } (c_{short} \leq 6) \rightarrow \text{TCC stopped, } QI = QI - 0.05 \quad (4)$$

300 Finally, the number of measurements and correlation coefficient with radar data for short and long periods are
301 examined. If the condition in Formula 5 is met then the check is stopped. If not, the check is failed:

$$302 \quad [(c \leq 6) \text{ or } (r = \text{"no data"})] \rightarrow \text{TCC stopped, } QI = QI - 0.05 \quad (5)$$

303 else \rightarrow TCC failed, $QI = QI - 0.3$.

304 This formula applies to cases when there are too few measurements, or the correlation coefficient could not be
305 calculated or was below the assumed threshold for short or long series.

306 **3.4 New algorithm BC (Detection of bias with adjusted radar data)**

307 The determination of bias in the BC algorithm is carried out by comparing the precipitation accumulations obtained
308 from the time series recorded on a given rain gauge with adjusted radar rainfall as a reference. For the most recent
309 10 days using a 10-min temporal resolution, rain gauge and radar precipitation accumulations, denoted as ΣG and
310 ΣR respectively, are calculated from gauge-radar pairs, for which both measurements have a quality index of at
311 least 0.7 for G and 0.8 for R .



312 Choice of the length of the precipitation accumulation period to determine the bias is not a trivial issue. Long
 313 accumulations better reflect the overall uncertainty of the measurements at a given station, but, on the other hand,
 314 short accumulations better follow the current precipitation characteristics during a particular precipitation event.
 315 Most often, bias is determined on rainfall accumulations from up to a few dozen hours, but sometimes on much
 316 longer accumulations – e.g. Yousefi et al. (2023) used seasonal totals to unbiased radar data with rain gauge data.
 317 The *bias* of the rain gauge measurements is calculated from the ratio of radar to rain gauge precipitation
 318 accumulations:

$$319 \quad bias = \frac{\Sigma R}{\Sigma G} \quad (6)$$

320 The bias determined in this way is used to reduce the quality index *QI* of the controlled rain gauge data. If the
 321 precipitation accumulations ΣG and ΣR are similar, which is checked using the corresponding similarity function,
 322 the quality of the measurement remains unchanged. The similarity function is defined as follows:

$$323 \quad 1.3 \cdot \min(\Sigma G, \Sigma R) + 7.0 > \max(\Sigma G, \Sigma R) \quad (7)$$

324 If the radar and rain gauge precipitation accumulations for a given rain gauge are not similar, then depending
 325 on the bias determined from Formula 6, the value of the quality index *QI* of a given measurement is reduced, but
 326 to a varying extent, according to the formula:

$$327 \quad QI = \begin{cases} QI - 0.05 & bias \in \left[\frac{1}{5}, 5\right] \\ QI - 0.2 & bias \in \left[\frac{1}{10}, \frac{1}{5}\right) \text{ or } bias \in (5, 10] \\ QI - 0.5 & bias \in \left[\frac{1}{20}, \frac{1}{10}\right) \text{ or } bias \in (10, 20] \\ QI - 1.0 & bias \in \left(0, \frac{1}{20}\right) \text{ or } bias \in (20, +\infty) \end{cases} \quad (8)$$

328 In cases when the bias cannot be estimated, the *QI* of a particular measurement is reduced according to the
 329 formula:

$$330 \quad QI = \begin{cases} QI - \min\left(1.0, \frac{|\Sigma G - \Sigma R|}{10.0}\right) & (\Sigma G = 0.0) \text{ or } (\Sigma R = 0.0) \\ QI - 0.2 & (\Sigma G = \text{"no data"}) \text{ and } (\Sigma R = \text{"no data"}) \end{cases} \quad (9)$$

331 In terms of data from private weather stations, they are considered subject to much greater uncertainty due to
 332 the lack of supervision of the technical condition of the rain gauges, poor maintenance, bad location, etc. Such
 333 stations should therefore be treated more rigorously than stations supervised by the institutions responsible for the
 334 measurements. The similarity function (Formula 7) is not applied, as their *QI* values are always reduced by the
 335 formula:

$$336 \quad QI = \begin{cases} QI - 0.1 & bias \in \left[\frac{1}{5}, 5\right] \\ QI - 0.3 & bias \in \left[\frac{1}{10}, \frac{1}{5}\right) \text{ or } bias \in (5, 10] \\ QI - 0.7 & bias \in \left[\frac{1}{20}, \frac{1}{10}\right) \text{ or } bias \in (10, 20] \\ QI - 1.0 & bias \in \left(0, \frac{1}{20}\right) \text{ or } bias \in (20, +\infty) \end{cases} \quad (10)$$

337 when *bias* cannot be estimated, the *QI* value of a given measurement is reduced by the formula:



$$338 \quad QI = \begin{cases} QI - \min\left(1.0, \frac{|\Sigma G - \Sigma R|}{10.0}\right) & (\Sigma G = 0.0) \text{ or } (\Sigma R = 0.0) \\ QI - 0.4 & (\Sigma G = \text{"no data"}) \text{ and } (\Sigma R = \text{"no data"}) \end{cases} \quad (11)$$

339 In addition, unbiasing should be performed for data from private stations, which is not done for other types of
 340 stations, as they only have a reduced QI . Unbiasing is performed on the basis of the bias determined from Formula
 341 6, but limiting its value to factor 4:

$$342 \quad bias_4 = \begin{cases} \frac{1}{4} & bias \leq \frac{1}{4} \\ bias & \frac{1}{4} < bias \leq 4 \\ 4 & bias > 4 \end{cases} \quad (12)$$

343 The above limitation on the value of the $bias_4$ factor is to protect against too large a change in the value of the
 344 corrected precipitation (van Andel, 2021).

345 Finally, the unbiased precipitation accumulation G_{cor} is determined from the formula:

$$346 \quad G_{cor} = bias_4 \cdot G \quad (13)$$

347 As IMGW does not yet have a sufficiently dense network of cooperating private stations (Drożdźoń and
 348 Absalon, 2023), tests have not been carried out to verify the algorithm designed in this study on data from such a
 349 network.

350 3.5 Updated SCC algorithm (Detection of outliers from the local vicinity)

351 The spatial methods for quality control, such as the SCC, are especially effective for dense rain gauge networks
 352 because they utilise observations from nearby stations (Alerskans et al., 2022). Thus, when applied to sparse
 353 networks, it is more likely that a correct value measured by a rain gauge will be classified as erroneous in the case
 354 of intense convective rainfall of a very local nature.

355 Based on the analysis of the performance of the SCC algorithm – as published in a previous paper on the
 356 standard version of RainGaugeQC system (Ośródko et al., 2022) in Appendix C – a modification was made in
 357 relation to the degree of QI reduction depending on the spatial variability of rainfall.

358 The algorithm has not changed in terms of assigning each rain gauge measurement to one of the three classes
 359 of outliers: strong, medium, and weak, and additionally non-outlier. However, the algorithm for reducing the QI
 360 value of each measurement assigned to any of the outlier classes was modified. In the current version of the
 361 algorithm, the magnitude of QI reduction depends on whether a given rain gauge measurement is within an area
 362 of a high spatial variability of precipitation determined from weather radar data of sufficient quality $QI(R)$. In this
 363 case, the outlier is treated less restrictively. The concept of spatial variability function (SVF) was introduced for
 364 this purpose, and is defined as follows:

$$365 \quad SVF = \frac{SVF_{mean}(R_{mean}) + SVF_{var}(R_{var})}{2} \quad (14)$$

366 The SVF consists of two components indicating the degree of spatial variability of the precipitation:

$$367 \quad SVF_{mean}(R_{mean}) = \begin{cases} 1 & R_{mean} \geq 1.0 \text{ mm} \\ \frac{R_{mean} - 0.1 \text{ mm}}{1.0 \text{ mm} - 0.1 \text{ mm}} & 0.1 \text{ mm} < R_{mean} < 1.0 \text{ mm} \\ 0 & R_{mean} \leq 0.1 \text{ mm} \end{cases} \quad (15)$$



$$SVF_{var}(R_{var}) = \begin{cases} 1 & R_{var} \geq 1.0 \text{ mm}^2 \\ \frac{R_{var} - 0.03 \text{ mm}^2}{1.0 \text{ mm}^2 - 0.03 \text{ mm}^2} & 0.03 \text{ mm}^2 < R_{var} < 1.0 \text{ mm}^2, \\ 0 & R_{var} \leq 0.03 \text{ mm}^2 \end{cases}$$

where R_{mean} is the mean radar precipitation (in mm) for wet pixels in the 100 km x 100 km subdomain including 25 km margins (see: Ośródk et al., 2022); R_{var} is the mean variance of radar precipitation (in mm²) in the subdomain calculated analogously to R_{mean} .

On the basis of the value of the SVF function, the reduction in the quality index for individual rain gauge observation is determined, according to its classification into a specific outlier class:

$$QI = \begin{cases} QI - (0.30 \cdot (1 - SVF) + 0.10 \cdot SVF) & \text{strong outlier} \\ QI - (0.20 \cdot (1 - SVF) + 0.05 \cdot SVF) & \text{medium outlier} \\ QI - 0.10 \cdot (1 - SVF) & \text{weak outlier} \end{cases} \quad (16)$$

3.6 Determination of QI

Before all the checks, each rain gauge observation is assigned the perfect QI value (1.0). Depending on the result of a particular QC algorithm, the QI of an examined measurement is decreased by a relevant value. If the final QI value, i.e. after all checks, is below a preset threshold, the observation is considered useless and is replaced with “no data”.

4 Analysis of the RainGaugeQC system performance on non-professional data

The performance of the RainGaugeQC system, designed to control the quality of precipitation data from professional and non-professional rain gauge networks, is shown through a comparison of the statistics calculated for these two rain gauge networks:

- professional network of IMGW, the Polish NMHS, supplemented in the border region by data from CHMU, which is the Czech NHMS,
- non-professional network of the General Directorate of the State Forests.

The most important characteristics of these networks are summarised in Table 1, and the locations of the rain gauges are shown in Fig. 1. Rain gauges from private networks have not been included, as the establishment of their network at IMGW is still at a preliminary stage.

The analysis was carried out for four months – April, July, October 2023 and January 2024 – considered typical of the four seasons. The summer season (July) is dominated by convective precipitation, which is often intense and highly variable in time and space, while the winter season (January) is dominated by stratiform precipitation, often in the form of snow. In the intermediate seasons (April, October) precipitation is less intense – it is generally rain, and is rarely convective.

4.1 Verification metrics

The reliability of the precipitation estimates generated using the RainGaugeQC system was verified by comparison with the reference precipitation accumulations from manual rain gauges that are treated as the closest to the true precipitation at their locations. The following metrics were employed:

- Pearson correlation coefficient:



400
$$CC = \frac{\sum_{i=1}^n (E_i - \bar{E})(O_i - \bar{O})}{\sqrt{\sum_{i=1}^n (O_i - \bar{O})^2 \sum_{i=1}^n (E_i - \bar{E})^2}} \quad (17)$$

- 401 • root mean square error:

402
$$RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^n (E_i - O_i)^2} \quad (18)$$

- 403 • root relative square error:

404
$$RRSE = \frac{\sqrt{\sum_{i=1}^n (E_i - O_i)^2}}{\sqrt{\sum_{i=1}^n (O_i - \bar{O})^2}} \quad (19)$$

- 405 • statistical bias:

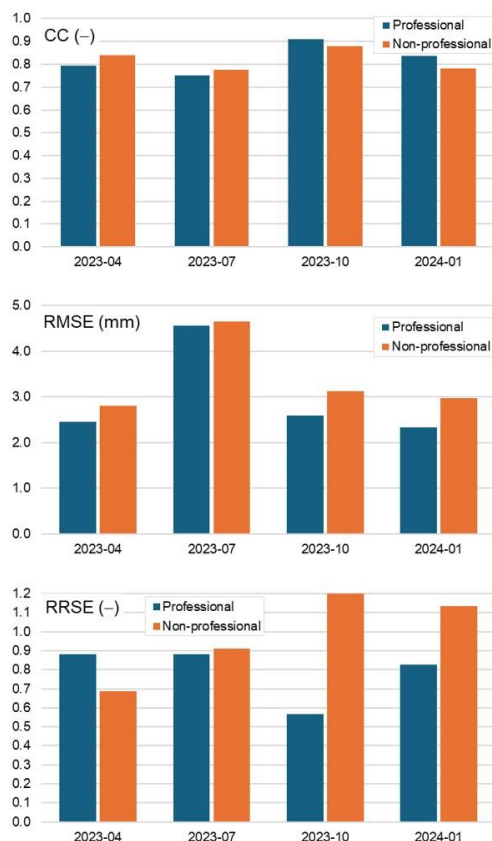
406
$$BIAS = \frac{1}{n} \sum_{i=1}^n (E_i - O_i) \quad (20)$$

407 where E_i is the estimated value, O_i is the reference value, i is the gauge number, n is the number of gauges, whereas
408 \bar{E} and \bar{O} are the mean values of E_i and O_i , respectively.

409 4.2 Non-professional versus professional rain gauge data

410 A comparison of reliability metrics of precipitation estimates obtained from a network of professional and non-
411 professional rain gauges is shown in Fig. 4. Point measurements of rainfall were verified against values at rain
412 gauge locations obtained from the interpolation of manual rain gauges using the inverse distance weighting
413 method. Professional rain gauges situated at manual gauge locations, a relatively common situation in the IMGW
414 network, were not included in the statistics in order not to favour this category of data. Therefore, around 200
415 professional rain gauges were used for verification instead of all 469.

416



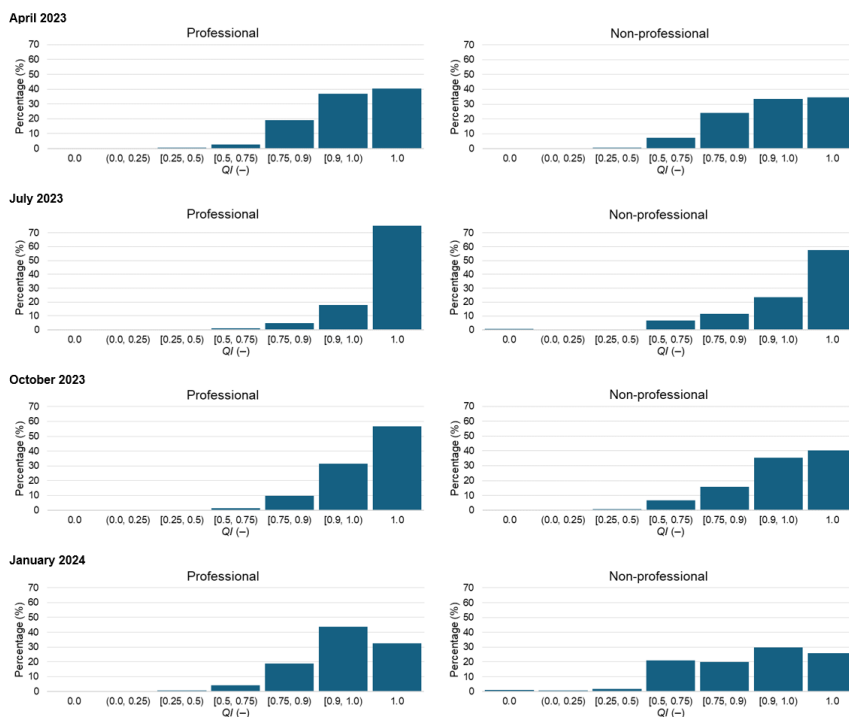
417
418 **Figure 4: Reliability statistics of rainfall estimates calculated for data obtained from the network of professional (navy)**
419 **and non-professional (orange) rain gauges. Spatially interpolated manual rain gauges are used as a reference. Data**
420 **from April, July, October 2023 and January 2024.**

421
422 The reliability of the non-professional data in general is close to that of the professional data, especially as
423 regards the correlation coefficient: on average for both it is about 0.82, and the differences between them are small,
424 at below 0.06. The RMSE metric related to the deviation from the reference data is already clearly worse for the
425 non-professional data, by on average about 0.41 mm. The largest difference was found for January, when it reached
426 0.65 mm. Only in the summer period (July) is the difference between the non-professional and professional data
427 small (0.09 mm), though the error values are highest at that time (4.65 and 4.55 mm, respectively). During this
428 period, convective precipitation is frequent, more intense, and also more dynamic, and as a consequence, the
429 comparison with spatially interpolated reference data can produce large differences. In contrast, a similar but
430 relative RRSE metric gives less conclusive results: in April it is much better for the non-professional data (0.69
431 versus 0.88), while in the other months the non-professional data are worse than the professional, with a significant
432 difference of 0.63 in October.



433 **4.3 Comparison of the QC system performance on professional and non-professional data**

434 In this Section an examination is made of the extent to which the *QI* of rain gauge data for professional and non-
 435 professional stations is reduced by the RainGaugeQC system in different months of the year. The *QI* plays a key
 436 role in the multi-source precipitation field estimation performed by the RainGRS system as the *QI* index is one of
 437 the most important weights during spatial interpolation of rain gauge data and, most importantly, it is a weight
 438 when rain gauge data is combined with the other precipitation estimates – radar and satellite-based. As a result of
 439 this approach, the impact of low-quality data on the final precipitation field estimate can be reduced.
 440



441
 442 **Figure 5: Percentages of data with *QI* values in different ranges (histograms). Data from April, July, October 2023 and**
 443 **January 2024.**

444
 445 Fig. 5 summarises the percentage of rain gauge data in different ranges of *QI* values assigned to individual
 446 measurements as a result of *QI* performed with a modified version of the RainGaugeQC system for four months
 447 representing different seasons, separately for professional and non-professional stations. It can be noted that, in
 448 general, *QI* values are significantly higher for professional data, meaning that QC algorithms indicate higher
 449 uncertainty in non-professional data. While unreduced quality ($QI = 1.0$) characterises 32.5 – 76.1% of all
 450 professional data depending on the season, just 26.0 – 57.6% of non-professional data. On the other hand, lower
 451 values below $QI < 0.75$ at different seasons characterise 1.4 – 4.9% of the professional data and 7.4 – 24.2% of
 452 the non-professional data.



453 There is noticeable seasonal dependence of the number of data with QI in specific value ranges, which is similar
 454 for professional as well as non-professional data. The highest percentage of data with a QI of exactly 1.0, i.e.
 455 perfect data according to the RainGaugeQC system, is observed in July (summer) and equals 76.1% and 57.6%
 456 for professional and non-professional data respectively, while the percentage of data with poor qualities is also
 457 lowest in this month for both types of the data: 1.4% and 7.2%, respectively. Considering the distribution of QI
 458 values in the different ranges, the data from January proved to be the least reliable, when the percentage of data
 459 with low QI values, i.e. in the range between 0.0 and 0.75, is the highest, reaching 4.9% for professional and 24.2%
 460 non-professional data.

461 4.4 Impact of non-professional rain data on the reliability of precipitation estimates

462 The following data sets were applied to test the influence of non-professional rain data on the reliability of
 463 precipitation estimation: (i) professional only and (ii) professional and non-professional together after quality
 464 control with the modified version of RainGaugeQC. From both rain gauge data sets, 10-min multi-source estimates
 465 of precipitation accumulations were generated with the RainGRS system and then aggregated to the daily
 466 accumulations. Table 3 shows the reliability metrics of the daily accumulations calculated for April, July, October
 467 2023 and January 2024, using the manual rain gauge data as a reference. Statistics were determined at the locations
 468 of the manual rain gauges.

469
 470 **Table 3. Reliability metrics of estimates of daily RainGRS precipitation accumulations generated using rain**
 471 **gauge data: professional and professional with attached data from non-professional rain gauges after**
 472 **quality control with the modified version of RainGaugeQC. Measurements from manual rain gauges are**
 473 **used as a reference, data from April, July, October 2023 and January 2024.**

Rain gauge networks	CC (-)	RMSE (mm)	RRSE (-)	BIAS (mm)
<i>April 2023</i>				
Professional (IMGW and CHMU)	0.832	2.74	0.64	1.36
Professional (IMGW and CHMU) and non-professional (State Forests)	0.872	2.40	0.55	1.11
<i>July 2023</i>				
Professional (IMGW and CHMU)	0.835	3.99	0.57	1.03
Professional (IMGW and CHMU) and non-professional (State Forests)	0.847	3.71	0.55	0.93
<i>October 2023</i>				
Professional (IMGW and CHMU)	0.920	2.35	0.43	0.91
Professional (IMGW and CHMU) and non-professional (State Forests)	0.922	2.28	0.41	0.79
<i>January 2024</i>				
Professional (IMGW and CHMU)	0.844	2.55	0.65	1.42
Professional (IMGW and CHMU) and non-professional (State Forests)	0.846	2.52	0.64	1.40

474



475 It can be seen from Table 3 that after the incorporation of non-professional data provided by the General
476 Directorate of the State Forests into RainGRS, all reliability metrics improved. The correlation coefficient, CC,
477 increased for all months analysed only marginally. Greater improvement after the inclusion of non-professional
478 data can be seen in all metrics related to error magnitude: RMSE, RRSE and BIAS, which on average decreased
479 by 0.13 mm, 0.02, and 0.08 mm, respectively.

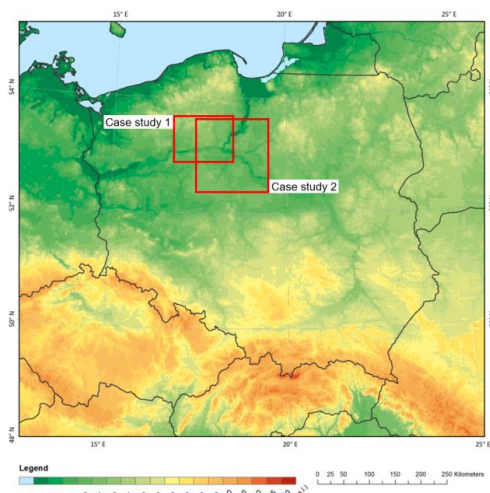
480 Analysing the four metrics used, the most positive impact of incorporating non-professional data was found in
481 April 2023, an intermediate month, when all characteristics improved: CC increased by 0.04, while metrics related
482 to error magnitude decreased: RMSE by 0.34 mm, RRSE by 0.09 and BIAS by 0.35 mm. This observation is
483 consistent with the results shown in Fig. 4, where in April the non-professional data were even more reliable than
484 the professional data in terms of CC and RRSE metrics. The smallest impact of non-professional data was observed
485 in January, when the improvement was negligible.

486 It should be pointed out that the number of non-professional rain gauges available for this study was not large:
487 the ratio between the number of rain gauges in the non-professional and professional networks was about 1:4.
488 Therefore, it can be expected that if there were more of these non-professional rain gauges, then the benefit from
489 them in terms of improvement in the reliability of the precipitation estimates would be even more pronounced.
490 This impact is not only due to the measurement information provided by these rain gauges, but also largely due to
491 the fact that additional rain gauges make quality control of all rain gauges much more effective.

492 **4.5 Impact of non-professional rain gauges on estimated multi-source precipitation field – varying impact** 493 **in different locations**

494 This section presents two case studies illustrating the influence of non-professional precipitation data on the
495 reliability of precipitation estimates generated by the RainGRS system. The location of the study areas is shown
496 on a map of Poland (Fig. 6). Locations in central Poland were chosen because the network of professional rain
497 gauges is sparsest there (see Fig. 1), so the influence of non-professional data on the final estimate of the
498 precipitation field can be expected to be more evident. Two different RainGRS precipitation field estimates were
499 generated using rain gauge data: (i) from professional rain gauges only, (ii) from both professional and non-
500 professional rain gauges. The impact of incorporating non-professional rain gauge data on multi-source field
501 estimates was assessed using manual rain gauge measurements as reference data. The analyses were conducted on
502 daily accumulations because only this kind of data are available from manual rain gauges.

503



504

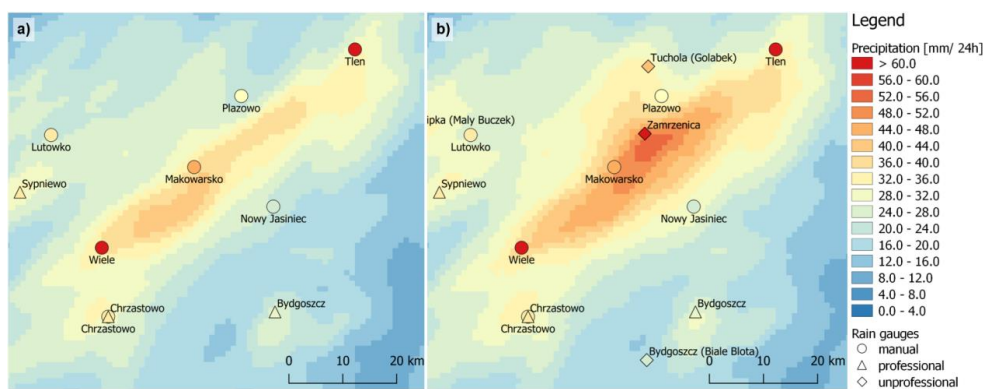
505 **Figure 6. Location of case studies on a map of Poland.**

506

507 **4.5.1 Case study 1: isolated convective precipitation (29-30 July 2023)**

508 On 29 and 30 July 2023 Poland was under the influence of a trough of low pressure and atmospheric front systems
509 moving from west to east. There were some showers and thunderstorms with precipitation locally reaching more
510 than 60 mm per day, which triggered flash flooding in major cities in the north of the country. Fig. 7 presents the
511 daily precipitation accumulations for this day, which shows the effect of including non-professional rain gauge
512 data to the input data to the RainGRS model generating multi-source precipitation field estimates.

513



514

515 **Figure 7: Precipitation maps of multi-source RainGRS estimates from: a) professional, b) professional and non-**
516 **professional data. The symbols are filled with colours that correspond to the precipitation values measured by each rain**
517 **gauge. A fragment of Poland, daily accumulations from 29.07.2023, 06 UTC to 30.07.2023, 06 UTC.**

518

519 In the fields of estimated precipitation accumulations in the vicinity of the thunderstorm cell in Fig. 7, it can
520 be seen that after incorporation of the non-professional data, the accumulations became noticeably higher, as the



521 data from the non-professional rain gauges are generally higher than those from the professional ones – a general
522 increase in values can be seen in Fig. 7b compared to Fig.7a. Using the measurements from the manual rain gauges
523 as reference data, it can be concluded that the obtained increase in the estimated RainGRS precipitation field is
524 closer to the reference precipitation (this is confirmed by the results in Table 3). Regarding the thunderstorm cell
525 moving through the study area, it was compact, small in size (its diameter was about 10 km) and no professional
526 rain gauge was in its path. It was detected by weather radars, so it is visible on the multi-source estimate, but the
527 precipitation values are underestimated compared to the reference precipitation recorded by the manual rain gauges
528 located in the path of this cell.

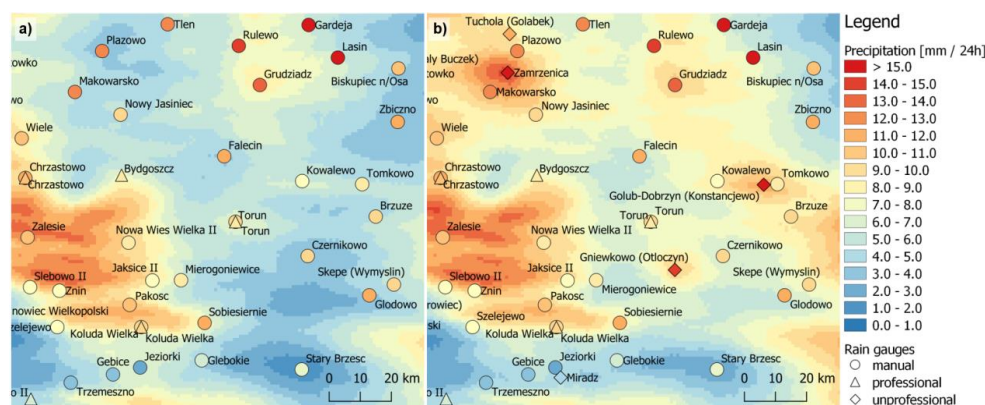
529 When including the non-professional data, a rain gauge in Zamrzenica on the route of this storm cell measured
530 a daily rainfall of 62.3 mm, resulting in a significant increase in the RainGRS precipitation estimate in this area:
531 from 31.6 to 50.6 mm at the Zamrzenica location. However, due to the small number of rain gauges in the area,
532 the high precipitation spread over a much larger region than the close vicinity of the cell. This is evidenced by the
533 lower precipitation measured by the manual rain gauge at Nowy Jasiniec (23.3 mm), while the precipitation
534 estimate increased here from 24.2 to 31.0 mm.

535 Closest to the path of the cell was the Makowarsko manual rain gauge, which measured 46.8 mm. The multi-
536 source estimate after including the non-professional rain gauge increased from 37.8 to 47.1 mm, which is in very
537 good agreement with the reference value. The precipitation estimate at the Płazowo manual rain gauge location
538 also increased: from 22.4 to 33.5 mm, while this rain gauge measured 29.2 mm. The increase in estimates was
539 therefore too high, but nevertheless, after data from non-professional rain gauges were added to the estimate, it
540 was closer to the measurement from the reference rain gauge. The highest value of 68.5 was measured by the Tleń
541 manual rain gauge, but the incorporation of the non-professional data only slightly improved the highly
542 underestimated estimate from 31.5 to 33.7 mm.

543 **4.5.2 Case study 2: winter stratiform precipitation (3-4 January 2024)**

544 At the beginning of January 2024, Poland was in the range of low-pressure systems moving from west to east and
545 associated atmospheric fronts. Rainfall and sleet were observed, with snowfall in the north-east of the country and
546 in the mountains in the south. In the north and centre, there was also freezing rain causing glaze. The example
547 shown in Fig. 8 relates to a lowland area in central Poland, like in the first case study, but here there was stratiform
548 precipitation, which was significantly lower but at a greater extent, as is typical for winter.

549



550
551 **Figure 8: Precipitation maps of multi-source RainGRS estimates from: a) professional, b) professional and non-**
552 **professional data. The symbols are filled with colours that correspond to the precipitation values measured by each**
553 **rain gauge. A fragment of Poland, 24-h accumulations from 3.01.2024, 06 UTC to 4.01.2024, 06 UTC.**
554

555 The RainGRS precipitation field estimation generated values that were underestimated compared to the manual
556 rain gauge measurements: the estimated values were lower by 3.2 mm on average, while their daily accumulation
557 averaged 9.5 mm at the locations of these rain gauges. This is mainly due to the underestimation of weather radar
558 and, to a lesser extent, telemetric measurements.

559 The inclusion of data from non-professional rain gauges, despite their small number, increased the RainGRS
560 estimate at manual rain gauge locations by an average of 1.3 mm. For example, it can be seen that that Zamrzenica
561 non-professional rain gauge had a positive effect on the estimated daily precipitation accumulation (RainGRS) at
562 the manual rain gauge located in Płazowo, where 12.1 mm was measured, and the estimates with and without the
563 incorporation of non-professional data were 9.5 and 3.3 mm, respectively.

564 The impact of the Miradz non-professional rain gauge was slightly different. It measured a value of 3.1 mm
565 and caused the estimates at the location of the two closest manual rain gauges to decrease at Jeziorki from 6.7 to
566 4.7 mm, and at Gębice from 6.0 to 4.9 mm, approaching the values from the manual rain gauges of 1.7 and 3.0
567 mm respectively. On the other hand, the influence of Miradz appeared to negatively affect the estimates at the
568 manual rain gauge locations of Kołuda Wielka and Szelejewo, where values that had been underestimated
569 compared to the reference rainfall were lowered even further.

570 The analysis of the two case studies shows that data from non-professional rain gauges, despite their generally
571 somewhat greater uncertainty, can in most cases play a positive role in the estimation of the precipitation field.

572 5 Conclusions

573 Data from non-professional rain gauge networks, as additional source of precipitation data, increase the density of
574 available rain gauge networks. In consequence they can improve precipitation field estimates at high spatial
575 resolution and can be very helpful to NHMSs for various meteorological and hydrological applications. However,
576 advanced data quality control systems are required to make these data useful for operational applications. At the
577 same time, it should be possible to objectively quantify the uncertainty associated with each individual
578 measurement.



579 The RainGaugeQC system, applied to quality control of rain gauge data, was redesigned in order to adapt it to
580 different rain gauge networks supervised to various degrees. In a modified version of the TCC algorithm, more
581 sophisticated data control was developed applying weather radar data, taking into account various aspects of data
582 quality, such as consistency analysis of data time series. The new BC algorithm was introduced to detect bias of
583 rain gauge measurements comparing rain gauge and radar long-term accumulations. In the SCC algorithm,
584 significant modifications have been made to quantify the quality index reduction for outliers on the basis of the
585 spatial variability of the precipitation field derived from the radar data. The performance of the modified system
586 was verified based on independent measurement data from manual rain gauges, which are considered one of the
587 most accurate measurement instruments. The influence of incorporating non-professional precipitation data on
588 reliability of multi-source precipitation estimates generated by the RainGRS system was also analysed.

589 The main conclusions derived from the analyses carried out in this study can be summarised as follows:

- 590 1. The incorporation of data from non-professional stations into professional rain gauge data, even if they
591 are of poorer quality (Fig. 5), nevertheless improves the reliability of the estimated multi-source
592 precipitation field (Table 3), but on the condition that advanced quality control is carried out.
- 593 2. Despite the quality control performed, the influence of individual rain gauges on the precipitation field
594 estimates may sometimes not be positive, as can be seen from the examples shown in Section 4.5.
595 Furthermore, the same rain gauge may have a different influence, positive or negative, on an estimated
596 precipitation field in various places.
- 597 3. Precipitation field estimates provided by weather radar data play a very important role in the developed
598 RainGaugeQC algorithms. However, it is necessary to perform their advanced quality control beforehand
599 and to adjust them with rain gauge measurements.
- 600 4. An important benefit of including data from non-professional networks is the improvement in
601 performance of individual QC algorithms. This is especially true for the spatial consistency check (SCC),
602 in which the density of a rain gauge network is crucial.
- 603 5. IMGW is in the process of setting up a network of private rain gauges. After its operational launch, it will
604 become possible to test the QC algorithms proposed in this paper on data from these rain gauges.

605 **Acknowledgement**

606 The work described in this paper was carried out primarily within the COSMO consortium (Consortium for Small-
607 scale Modelling) as Priority Task EPOCS “Evaluate Personal Weather Station and Opportunistic Sensor Data
608 Crowdsourcing” during the period 2023-2024.

609

610 *Code availability.* The data processing codes are protected through the economic property rights to the software
611 and are not available for distribution. The codes used for processing follow the methodologies and equations
612 described herein.

613

614 *Data availability.* Out of the data used in this article, the following are publicly available:

615 IMGW rain gauge data in the form of 10-minute accumulations: <https://danepubliczne.imgw.pl/pl/datastore>, tabs:
616 „Dane archiwalne” / „Dane meteorologiczne” / „year” / „Meteo_year-month.zip” / „B00608S_year_month.csv”
617 (B00608S is the code for the 10-min rainfall parameter).



618 Radar data as 1-h files of precipitation accumulation (PAC) maps: <https://danepubliczne.imgw.pl/pl/datastore>,
619 tabs: „Dane archiwalne” / „Mapa zbiorcza sumy opadów za godzinę.” / „year” / „month” /
620 COMPO_PAC.comp.pac_year-month-day.tar.
621 Other data used in this article is available upon request, provided it is not restricted by its producer.
622
623 *Author contributions.* KO, JS, and AJ designed algorithms of the RainGaugeQC system. KO developed the
624 software code and performed the simulations. JS, KO, AJ, and AK prepared the paper. JS made figures. AK carried
625 out statistical calculations.
626
627 *Competing interests.* The contact author has declared that none of the authors has any competing interests.

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