Atmospheric H₂ observations from the NOAA Cooperative Global-Cooperative Air Sampling Network

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15 Abstract. The NOAA Global Monitoring Laboratory (GML) measures atmospheric hydrogen (H_2) in 16 grab-samples collected weekly as flask pairs at over 50 sites in the Cooperative Global Air Sampling 17 Network. These NOAA H₂ measurements from 2009 to 2021 are publicly available. Measurements 18 representative of background air sampling show higher H_2 in recent years at all latitudes. The marine 19 boundary layer (MBL) global mean H₂ was 552.8 ppb in 2021, 20.2 ±0.2 ppb higher in 2021 compared to 20 2010. A 10 ppb or more increase over the 2010-2021 average annual cycle was detected in 2016 for MBL 21 zonal means in the tropics and in the Southern Hemisphere. Carbon monoxide measurements in the same 22 air samples suggest large biomass burning events in different regions likely contributed to the observed 23 interannual variability at different latitudes. The NOAA H₂ measurements from 2009 to 2021 are now 24 based on the A major focus in recent years involved the adoption of the World Meteorological 25 Organization Global Atmospheric Watch (WMO GAW) H₂ mole fraction X2009 calibration scale, 26 developed and maintained by the Max-Planck Institute for Biogeochemistry (MPI-BGC), Jena, Germany. 27 GML maintains eight H_2 primary calibration standards to propagate the WMOMPH scale. These are 28 gravimetric hydrogen-in-air -mixtures in electropolished stainless steel cylinders (Essex Industries, Set. 29 Louis, MO), which are stable for H₂. These mixtures were calibrated at the MPI-BGC, the WMO Central 30 Calibration Laboratory (CCL) for H₂, in late 2020 and span the range 250-700 ppb. We have used the 31 CCL assignments to propagate the WMOMPI X2009 H₂ calibration scale to NOAA air measurements 32 performed using Gas Chromatography-Helium Pulse Discharge Detector instruments since 2009. To 33 propagate the scale, NOAA uses a hierarchy of secondary and tertiary standards, which consist of 34 high-pressure whole air mixtures in aluminum cylinders, calibrated against the primary and secondary 35 standards respectively. Hydrogen at the ppb-level has a tendency to increase in aluminum cylinders over 36 time. To propagate the scale, NOAA uses a hierarchy of secondary and tertiary standards, which are high-37 pressure tanks with whole air mixtures calibrated against the primary and secondary standards-38 respectively. NOAA secondary and tertiary standards are stored in aluminum eylinders, which have a 39 tendency to grow H₂ over time. We fit the calibration histories of these standards with 0-2nd order 40 polynomial functions of time and use the time-dependent mole fraction assignments on the MPI X2009 41 WMO scale to reprocess all tank air and flask air H_2 measurement records. The robustness of the scale 42 propagation over multiple years is evaluated with the regular analysis of target air cylinders and with 43 long-term same air measurement comparison efforts with WMO GAW partner laboratories. Long-term

44 calibrated, globally distributed and freely accessible measurements of H₂ and other gases and isotopes

45 continue to be essential to track and interpret regional and global changes in the atmosphere composition.

46 The adoption of the WMOMPL X2009 H_2 calibration scale and subsequent reprocessing of NOAA

47 atmospheric data constitute a significant improvement in the NOAA H₂ measurement records.

48

49 1 Introduction

50

51 High quality and sustained observations are essential to track and study changes in atmospheric trace gas
52 distributions. Ambient air measurement programs for trace gases provide objective data to track air
53 pollution levels [Oltmans and Levy, 1994; Thomson et al., 2004; Tørseth et al., 2012; Schultz et al., 2015;
54 Cooper et al., 2020; WMO, 2022], to study how a mix of sources (and sinks) impact the air composition
55 [Ciais et al., 1995; Pétron et al., 2012; Langenfelds et al., 2002; Brito et al., 2015] and to constrain and
56 evaluate fluxes and their trends at scales of interest [von Schneidemesser et al., 2010; Simpson et al.,
57 2012; Propper et al., 2015; Montzka et al., 2018; Friedlingstein, 2022; Heiskanen et al., 2022; Storm et
58 al., 2023] at seales of interest.

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⁶⁰ H₂ is a trace gas in the Earth's atmosphere and its abundance can indirectly impact climate and air quality.
⁶¹ The analysis of H₂ measurements in firn air collected in Antarctica reveal that H₂ levels in the
⁶² high-latitude southern hemisphere grew by some 70% (330 to 550 ppb, 1 ppb = 1 mole of gas per billion
⁶³ (10⁹) moles of air) over the 20th century [Patterson et al., 2021; 2023]. Greenland firn air covers less
⁶⁴ depth and time but results are consistent with a 30% increase in high-latitude northern hemisphere H₂
⁶⁵ from 1950 to the late 1980s [Patterson et al., 2023]. Growing emissions related to fossil fuel burning most
⁶⁶ likely were behind this rise in H₂ [Patterson et al., 2021]. Results also show that H₂ in both polar regions
⁶⁷ leveled off after the 1990s [Patterson et al., 2021, 2023].

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⁶⁹ H₂ has been viewed as a potential low or zero carbon energy carrier for close to five decades [Yap and
⁷⁰ McLellan, 2023]. Since 2020 there has been renewed interest in the hydrogen economy [Yap and
⁷¹ McLellan, 2023] spurred by a rise in announcements of public and private projects to produce low carbon
⁷² H₂, also referred to as "blue" H₂ produced from natural gas with carbon capture, utilization and storage, or
⁷³ "green" H₂ produced using renewable energy [Hydrogen Council and McKinsey & Company, 2023]. In
⁷⁴ 2021, H₂ global demand was over 94 million tonnes or 2.5 % of global final energy consumption [IEA,
⁷⁵ 2022]. This demand was almost entirely driven by refineries and a few industries (ammonia, methanol
⁷⁶ and steel) and H₂ production almost entirely relied on fossil fuels with unabated emissions ("gray H₂",
⁷⁷ [IEA, 2022]). As of December 2023, over 1,400 announced projects globally (worth US\$ 570 billion) are
⁷⁸ anticipated to increase the global H₂ production capacity by 45 million tonnes through 2030 [Hydrogen
⁷⁹ Council and McKinsey & Company, 2023].

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81 Studies of the potential short-term and long-term climate impacts of increased H_2 production and use have 82 called for more research to better understand the current and future H_2 supply chain and end-use

 10 cance for more research to better understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand the carrent and ratio H_2 suppry chain and characteristics are understand to be carrent and ratio H_2 suppry characteristics are understand to be carrent.

84 Bertagni et al., 2022; Warwick et al., 2023]. Global, high quality and sustained atmospheric

85 measurements of H₂ can provide independent information to document its distribution and study its

86 sources and sinks and how they may change.

88 The National Oceanic and Atmospheric Administration (NOAA) Cooperative Global Air Sampling 89 Network comprises over 50 surface and mostly remote sites (<u>https://gml.noaa.gov/ccgg/flask.html</u>). At 90 each site and on a weekly basis, local partners collect air in two 2.5-L glass flasks, and then return the 91 flasks to the NOAA Global Monitoring Laboratory (GML) in Boulder, Colorado, USA, for measurements 92 of major long-lived greenhouse gases, carbon dioxide (CO₂), methane (CH₄), nitrous oxide (N₂O), sulfur 93 hexafluoride (SF₆), as well as carbon monoxide (CO) and hydrogen (H₂) [Conway et al., 1994; Novelli et 94 al., 1999; Dlugokencky et al., 2009]. The network is a contributor to the World Meteorological 95 Organization (WMO) Global Atmospheric Watch (GAW) Programme, which promotes and coordinates 96 international scientific efforts and free access to long-term atmospheric observations [WMO, 2022]. 97

98 CO and H_2 are important trace gases that share sources with CO_2 and CH_4 (fossil fuel burning, biofuel 99 burning and wildfires). Reaction with hydroxyl radicals (OH) is the main sink for CH_4 and CO and an 100 important sink for H_2 . Both H_2 and CO are also produced during the chemical oxidation of CH_4 and 101 nonmethane hydrocarbons. Soil uptake by bacteria accounts for 75% of the total H_2 sink. H_2 and CO have 102 much shorter atmospheric lifetimes than CO_2 and CH_4 : 2-3 months for CO and close to 2 years for H_2 . 103 The H_2 global mean atmospheric lifetime is largely driven by the soil sink strength. The H_2 lifetime 104 related to the oxidation by OH is estimated to be 8-9 years [Price et al., 2007; Warwick et al., 2022].

106 The "Geophysical Monitoring for Climatic Change" was the original program established by NOAA to
107 gather and analyze observations of the background atmosphere composition. GMCC started measuring
108 CO₂ in background air samples in 1968 [Komhyr et al., 1985]. CH₄ was added in 1983 [Steele et al.,
109 1987]. In the late 1980s, GMCC and its successor the Climate Monitoring and Diagnostics Laboratory
110 expanded operations to measure CO in the global network air samples to add a constraint for the study of
111 combustion sources and the global carbon budget. The analytical instrument selected consisted of a gas112 ehromatograph (GC) and a reduction gas analyzer (RGA, from Trace Analytical Inc., California) that
113 eould measure both CO and H₂. ¶

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115 Novelli et al. [19912, 19924] first reported for NOAA on testing the air sampling approach (flask type, 116 stopcock fitting, wet/dry air, untaped versus taped glass flasks to minimize direct sunlight exposure) and 117 an analytical instrument consisting of a gas chromatograph (GC) and a reduction gas analyzer (RGA, 118 from Trace Analytical Inc., California)CC RGA instrumentation that could measure both CO and H₂for-119 CO. Around that time, other laboratories had also adopted the GC-RGA measurement technique for CO 120 and H₂ measurements in discrete air samples or in situ. Khalil and Rasmussen [1989, 1990] reported on 121 H₂ measurements of whole air samples collected weekly in triplicate electropolished stainless steel flasks 122 between October 1985 and April 1989 at the four NOAA atmospheric baseline observatories (Point 123 Barrow, Mauna Loa, Samoa, South Pole), Cape Meares, Oregon, Cape Kumukahi, Hawaii and at the Cape 124 Grim Observatory, Tasmania. These measurements showed that, contrary to CO₂, CH₄, N₂O and CO, 125 background air H₂ levels were higher in the Southern Hemisphere (SH) than in the Northern Hemisphere 126 (NH). 1985-1987 monthly mean observed H₂ ranged between 500-520 ppb at the South Pole and between 127 455 and 520 ppb at Point Barrow. H₂ exhibited a strong seasonal cycle at extratropical latitudes especially 128 in the NH and the seasonal cycles in both hemispheres were offset by 1-2 months only.

130 In 1995, H_2 mole fraction calibration standards were prepared gravimetrically in aluminumScott Marrin 131 Inc. cylinders (Scott Marrin Inc., Riverside, CA) and five of them (spanning 485-603 ppb) were used to 132 define the NOAA H_2 X1996 calibration scale. Working standards used in the NOAA flask analysis 133 laboratory between 1988 and 1996 were reassigned H_2 mole fractions and flask air measurements were 134 reprocessed to be on the X1996 scale. Novelli et al. [1999] described the early NOAA H_2 measurements 135 and reported H_2 time series starting in the late 1980s or early 1990s (depending on the site) for 50 sites in 136 the NOAA Cooperative Global Air Sampling Network.

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138 Simmonds et al. [2000] reported in-situ high-frequency GC-RGA3 measurements of H_2 at the Mace Head 139 baseline atmospheric monitoring station on the Atlantic coast of Ireland for the 1994-1998 period. They 140 found that the background air at Mace Head had lower monthly mean H_2 (470-520 ppb) than background 141 air masses measured at the Cape Grim observatory (510-530 ppb) from July to April. Some of the 40 min 142 H_2 observations showed 10s-200 ppb short-term H_2 enhancements above baseline levels. The authors 143 derived an estimate of European emissions with an inverse model of enhanced H_2 in air masses impacted 144 by upwind sources of pollution. They also observed that nighttime measurements in low wind conditions 145 reflected local depletion of H_2 . The authors derived variable mean deposition velocities and found that the 146 H_2 soil sink was likely a process that occurred year-round in the area.

148 After 1996 and until 2008, the NOAA H_2 measurement program used successive working standards that 149 were assigned based on GC-RGA measurements against the previous standards. With hindsight, the 150 NOAA X1996 calibration scale transfer and the early NOAA H_2 measurements had several limitations 151 which are briefly described below and in more detail in the Supplementary Information section S1. 152

By the late 1990s, same air or collocated air sample measurement comparison between NOAA and the
Commonwealth Scientific and Industrial Research Organisation (CSIRO) for the Cape Grim Observatory
and Alert, Canada, flask air analyses showed an increasing bias for H₂ between the two laboratories
[Masarie et al., 2001; Francey et al., 2003]. Further laboratory tests by several WMO/GAW measurement
laboratories revealed the RGA detector response was non linear and required frequent calibration.
Additionally measurement laboratories found that the H₂ mole fraction for air standards, especially those
stored in high pressure aluminum cylinders, could drift at rates of a few parts per billion (ppb) to tens of
ppbs per year [Novelli et al., 1999; Masarie et al., 2001; Jordan and Steinberg, 2011].

162 To address these compounding issues, in 2008 NOAA GML tested a new analytical instrument: a gas
163 chromatograph with a pulse discharge helium ionization detector (GC-HePDD) [Wentworth et al., 1994].
164 The technique showed very good performance with a stable and linear response over the 0-2000 ppb
165 range and it was adopted for the calibration scale propagation and flask air analysis in 2009 [Novelli et
166 al., 2009]. Around that time GMLNOAA also began testing electropolished stainless steel cylinders
167 (Essex Industries, St. Louis, MO) filled with dry air for stability.

168

169 In 2007-2008, GML prepared six new gravimetric air mixtures in electropolished stainless steel cylinders
170 spanning 250-600 ppb H₂. At that time, the new gravimetric mixtures differed by about +20 ppb
171 compared to two H₂ secondary standards values assigned on the NOAA H₂ X1996 scale. For the next
172 decade, GML kept using the NOAA X1996 calibration scale while also conducting routine measurements

173 of the H_2 secondary standards against the 2007/2008 gravimetric mixtures.

175 The GC-HePDD H_2 measurements on the NOAA H_2 X1996 scale remained biased compared to GAW 176 partner measurements and the NOAA H_2 data from the global network flasks were not released publicly 177 after 2005. SI sections S1-3 and SI Table 1 provide additionalbackground information on issues impacting 178 the 1988-2008 NOAA H_2 measurements on RGAs, and related information from the CSIRO and 179 Max-Planck Institute for Biogeochemistry (MPI-BGC) H_2 measurement programs. The best-more precise 180 and better calibrated NOAA H_2 measurement records date back to 2009/2010 and are the main focus of 181 this paper.

182

183 In Fall 2020, GML initiated an effort to 1) adopt the WMO MPI X2009 H₂ calibration scale [Jordan and 184 Steinberg, 2011] for future measurements and 2) convert GML H₂ measurements made on GC-HePDD 185 instruments (beginning in late 2009) to that scale. This paper describes the MPI X2009 H₂ calibration 186 scale propagation within GML and the revised measurements from the NOAA Cooperative Global Air 187 Sampling Network flask air samples analyzed since late 2009. We show very good agreement for the 188 reprocessed NOAA H₂ data for different WMO /GAW measurement comparison efforts. The revised 189 NOAA GML flask air H₂ dry air mole fraction measurement records for 70 surface sites from 2009-2021 190 are publicly available [Pétron et al., 2023a]. This new dataset complements other WMO /GAW H₂ 191 measurement datasets and provides reliable observational constraints for the study of atmospheric H₂ 192 global distribution and budget since 2009. Future NOAA H₂ dataset updates will be released as we use 193 continued calibration results to reliably track the drift in working-standards and revise their assignments.

195 2 Adoption of the WMO MPI X2009 H₂ calibration scale

196

197 To infer fluxes and trends from atmospheric measurements, scientists need to reliably detect small 198 temporal and spatial gradients in the abundance of trace gases. This requires comparable data across time 199 and across monitoring networks to ensure biases are minimized and do not influence interpretation. The 200 use of a common calibration scale among measurement laboratories ensures data are traceable to a 201 common reference. It is the first step in preventing biases that could stem from using different references. 202

203 In this section, we introduce the NOAA GML H₂ calibration standard hierarchy and describe the adoption 204 of the WMO MPI X2009 H₂ scale. First we introduce the GML H₂ primary standards. Calibration at GML 205 is based on a hierarchy of standards (primary, secondary, tertiary, etc.) primary standards and Then we 206 describe the GML tank air and a dedicated H₂ calibration system used to transfer the scale and the seale-207 transfer from the primary standards to secondary and tertiary standards (2009-April 2019) or from the 208 primary standards to secondaryworking standards (after April 2019). The tertiary standards (until April 209 2019) and secondaryworking standards (after April 2019) are used to calibrate the H₂ instrument response 210 on the flask air analysis systems and value assign discrete air measurements. An important quality 211 assurance procedure within GML is the routine measurement of dedicated quality control cylinders 212 (referred to as "Target" tanks) to track instrument performance. Results are discussed in relation to the 213 uncertainty of the flask air analysis systems and consistency of the MPI X2009 H₂ scale implementation.

214

215 2.1 NOAA GML H₂ primary standards

216

217 In 2007-2008, six mixtures of H_2 in dry air air mixtures spanning a range of H_2 dry air mole fractions 218 were prepared gravimetrically at GML in electropolished stainless steel 34L cylinders ([Novelli et al., 219 2009], and Table 1). The highest H₂ mole fraction tank developed a leak and was lost. The remaining set
220 of five standards covered the range 250 ppb to 600 ppb for H₂. Three standards in electropolished
221 stainless steel cylinders were added in 2019 to extend the upper limit of the calibration range to 700 ppb
222 H₂ and evaluate the stability of the initial set over the intervening years. In 2020, tThese eight standards
223 werehave been designated by GML as NOAA GML'sour highest level H₂ standards. We refer to them as
224 the NOAA H₂ primary standards throughout this document even though they are not used to
225 independently define are not defining the scale.

226

227 The eight primary standards were analyzed by the WMO Central Calibration Laboratory (CCL) for H₂ 228 hosted by the MPI-BGC in Jena, Germany, on their GC-PDD system in November 2020. The results 229 listed in Table 1 are reported on the MPI X2009 H₂ calibration scale [Jordan and Steinberg, 2011]. The 230 CCL uncertainty estimates listed in Table 1 refer to the standard deviation of the 25-32 discreete H₂ 231 measurements made for each standard. Until they are recalibrated by the CCL, we add an 0.5 ppb 1-sigma 232 uncertainty to these assignments. This is the currently reported CCL reproducibility for their GC-HePDD 233 H₂ measurements. It accounts for potential longer term uncertainty in calibration results that would not be 234 evident in the standard deviations of measurements made close in time.

235

236 2.2 MPI X2009 H₂ calibration scale transfer

237

GML has separate, dedicated analytical systems for scale propagation and flask air analyses. Novelli et al.
[2009] describe the GC-HePDD instruments and the operating parameters in detail. GML has used three
GC-HePDD instruments so far. Each is identified by a unique internal instrument identification code: H9
(Agilent 6890 GC, serial number US10326037) for tank calibrations and H8 (Agilent 6890 GC, serial
number US10326011) and H11 (Agilent 7890 GC, serial number US10834030) for flask analyses. The
GC-HePDD instruments' responses are linear (within 0.3%) up to 2000 ppb. They are configured for ppb
level sensitivity and calibrated over the 200-700 ppb range, which is optimal for global and regional
background air analysis.

246

247 The GML H₂ primary standards are used to periodically calibrate the H9 instrument response for the 248 analysis and value assignment of lower level standards. GML used secondary standards from 2009-249 through April 2019 in the calibration hierarchy but has since removed this level to reduce the number of 250 standards which can potentially drift (see discussion of drifting cylinders later in this document). The 251 stability and longevity of the primary standards are critical to ensure the consistency of the GML H₂ 252 measurements over long periods of time as required for trend analysis.

The H₂ secondary and tertiary (or working)-standards) used in GML are whole air mixtures in high pressure aluminum cylinders (Luxfer USA). Most were filled at the GML standard air preparation facility the Niwot Ridge mountain research station using a Rix Industries (Benicia, CA) SA6 oil-free compressor [Kitzis, 2017]. Two additional tertiary standards (CB11551 and CC305198) were purchased from Scott Marrin. All GML tank air mixtures have a unique combination of an alphanumeric cylinder ID and a fill code letter (A-Z) tied to a fill date.

Aluminum tanks are known to be unstable for storing H_2 in air standards [Jordan and Steinberg, 2011]. Therefore regular analyses of working standards on the tank calibrations system are critical for quantifying drift to allow a time dependent value assignment on the MPI X2009 H_2 calibration scale.

265 The calibration history for a secondary, tertiary or working standard only uses retained (valid) calibration 266 event results on H9. GML uses python software developed in-house in house to record calibration data, 267 compute mole fractions, and write instrument output files and to calculate a calibration event result. 268 Another piece of python code is used to analyze tank air calibration histories and evaluate if the stability 269 of H₂ mole fractions in the tank is stable or if it changes over time. For many GML H₂ calibration 270 standards and target air tanks, a linear or quadratic function is the best fit through their calibration history. 271 When this happens, the function coefficients define the tank time-dependent assignment. All mole 272 fraction assignments and associated drift coefficients for standards used to propagate a calibration scale 273 are stored in a database table that can be accessed by the data processing software. The software allows 274 for 0-2nd orderlinear and polynomial drift functions. As new calibration results are available, the drift 275 correction and assignment for a particular tank ID and fill code areis revised as needed and the affected 276 data areis reprocessed.

277

278 2.2.1 Scale transferTank calibration system: 2009-2019 configuration

279

From 2007 through mid-April 2019, the H_2 tank air calibration on the H9 instrument was conducted using a single standard gas (primary or secondary standard) to calibrate the "unknown" (secondary or tertiary) tank Each calibration event consisted of alternating injections of the reference/standard gas and the "unknown" tank air with typically seven or more unknown air injections. The first aliquot in a multi injection measurement sequence on H9 is often slightly biased (due to subtle timing differences with the regulator flush cycle) and is not used. The ratio of the H_2 peak height for each valid "unknown" are injection and the mean peak height of two bracketing reference gas injections (or sometimes only one preceding or following reference gas injection) is multiplied by the reference/standard gas known H_2 mole fraction to calculate the "unknown" air injection mole fraction. Results for a tank air calibration event are defined by the mean and the standard deviation of the calculated H_2 mole fractions for five or more preceding unknown air injections. Typically, the standard deviation for a tank air calibration event on H9 is preceding the mean and the standard deviation of the calculated H_2 mole fractions for five or more preceding unknown air injections. Typically, the standard deviation for a tank air calibration event on H9 is preceding the mean 1 ppb.

292

293 Prior to the 2023 GML H_2 data reprocessing, GML used peak area for the GC-HePDD as described in 294 Novelli et al. [1999]. However, we saw that for some Helium carrier gas tanks (Airgas Ultra High Purity, 295 (99.999% purity), the H_2 chromatogram peak had a tail or a noisy baseline. Since the H_2 peak height was 296 less affected, we use peak height ratios for all GC-HePDD measurements. In 2023, GML switched to 297 Matheson Research Grade Helium carrier gas for the GC-HePDDs (99.9999% purity). 298

299 Two secondary standards with background ambient air level H_2 were in service on H9 to calibrate tertiary-

300 standards: CC119811 (2008/02 to 2013/06) and CA03233 (2013/06 to 2018/11). These two standards-

- 301 were calibrated periodically on H9 against individual members of the primary standard suite. Most-
- 302 ealibration episodes consisted of one to 6 ealibration sequences over 1-3 successive days, each against
- 303 one of the primary standards. For CC119811, 1-point calibration sequence results in 2008, 2009, 2010 and
- 304 2013 against one of the two lowest primary standards (SX-3558 and SX-3543) show a 3 to 5 ppb positive

Ŧ

305 bias which suggests a small non zero intercept in the instrument response during those times. This

306 primary standard dependent bias is not apparent for CA03233 results between 2014 and 2016. Results-

307 against SX-3558 were not used for value assigning either secondary standards and results against

308 SX-3543 were not used for CC119811.

309

The calibration results for the two H₂ secondary standards used between 2009 and Arpil 2019 are plotted in Figure 1 and final assignments are listed in SI Table 2. A small non zero y-intercept for H9 (see next section) likely explains the biased results for CC119811 against the lowest primary standards (SX-3558 and SX-3543). Results against SX-3558 were not used for value assigning either secondary standards and and results against SX-3543 were not used for CC119811.

315

316 CA03233 was stable for H_2 over its time of use and has an assignment of 502.8 ppb H_2 . H_2 in CC119811 317 exhibited a small linear drift and its value assignment is time dependent with a growth rate of 2 ppb/yr. 318 Between 2009 and 2019, these two secondary standards were used on H9 to calibrate seventeen H_2 319 tertiary standards used in the NOAA flaskdiserete air sample analysis laboratory. 320

321 **2.2.2** Scale transferTank air calibration system configuration: 2019-present configuration¶

Beginning in April 2019, GML transitioned H9 to use a multi-point calibration strategy to better define the instrument response. The eight H_2 primary standards are measured relative to a reference air tank (CC49559, filled withat ambient Niwot Ridge dried air) to calibrate the instrument response. A multi-standard response calibration episode for H9 involves the alternating injections from the reference air tank and each primary standard. Each standard is injected 8 times alternating with reference air aliquots. The entire response calibration sequence takes close to 15 hours. GML has performed an H9 instrument response calibration 2 to 3 a few times a year, followed by tank calibrations $\frac{2 \text{ to 3 a few times}}{2 \text{ to 3 a few times}}$

331

The H9 instrument response function is calculated as the best linear fit to the primary standards' mean anormalized chromatogram peak heights and their CCL H_2 mole fraction assignments. H9 calibration accurves are assumed to be valid for several weeks during which time other air cylinders are analyzed relative to the same reference tank.

336

Between April 2019 and December 2022, the H9 instrument response was determined relative to the primary standards-ealibrated nine times. Figure 2a shows the deviations of the H9 linear response functions from the line defined by computing the mean value for the intercept and slope of the 2019-2022 intercept response functions. The instrument response has remained stable within +/-1 ppb over this time period inver the range 200-700 ppb. Figure 2b shows the residuals to the best linear fit for each instrument response calibration episode. We note that the H9 instrument response has been quite stable over the intercept and slope of the linear fit does not go through the origin. The residuals to each the linear fit 200-700 ppb range but that the linear fit does not go through the origin. The residuals to each the linear fit intercept and 5.5 ppb (not shown). Prior to 2019, we assumed a zero intercept for the H9 one intercept for the H9 one intercept around 5 ppb was more likely, it is possible the pre-2019 H9 intercept (with 1 point calibration) were biased by ~1% of the difference between the tank air and it reference/standard H₂ mole fractions. We do not correct for this potential bias at this time.

Since April 2019, A a tank air calibration measurement sequence on H9 hastypically consisteds of 7 tank air injections, each bracketed by reference air injections. The peak heights for the first injections of reference air and tank air on H9 can have a small low bias and are not used. The normalized peak heights for the valid tank air injections are converted to H_2 mole fractions using the most recent H9 instrument response functionealibration episode. The average and standard deviation of the retained injection H_2 mole fractions are stored in a database table.

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357 2.2.3 H₂ standards and calibration approach for the flask air analysis system

358

359 The NOAA Global Cooperative Air Sampling Network dates back to 1967. In recent years, it has-

360 included over 50 surface sites distributed around the world-

361 (https://gml.noaa.gov/cegg/behind_the_seenes/network.html). Partners at each site collect air sample pairs-

362 in two 2.5L glass flasks filled simultaneously once a week and return the samples to NOAA GML in 363 Boulder. H_2 in those flask air samples is measured in addition to long-lived GHGs (CO₂, CH₄, N₂O, SF₆) 364 and CO by the Measurement of Atmospheric Gases that Influence Climate Change (MAGICC) system in 365 the NOAA GML Boulder laboratory. Until mid 2019, GML operated two nearly-identical automated flask

366 air analytical systems: MAGICC-1 (1997-2019) and MAGICC-2 (2003-2014). Since mid-2019, GML has 367 used a new MAGICC-3 system. This new system improved analytical techniques for CO_2 , CH_4 , N_2O , and 368 CO but continues to use the same GC-HePDD instruments from the older systems.

369

370 Two GC-HePDD instruments have been used for hydrogen analysis on the three flask air analysis systems
371 since 2009: H8 (MAGICC-2: 2009-2014 and MAGICC-3: August 2019-September 2020) and H11
372 (MAGICC-1: 2010-July 2019 and MAGICC-3: September 2020-present).

373 On MAGICC-1 and MAGICC-2, the H_2 He-PDD instrument response was calibrated using a single 374 tertiary standard (measured before and after each sample aliquot), similar to the original 1 point 375 calibration approach used on H9.

376 Out of 17Typically, the H₂ tertiary standards used during that time, 3 were used for more than 14 months 377 14 lasted less than a year and 14most displayed H₂ growth over time. Figure 3 shows the calibration 378 histories for H8 and H11 tertiary standards and their start/deployment dates. For each tertiary standard, 379 assigned mole fractions, drift coefficients, and estimated uncertainties are stored in a database (SI Table 380 2sXX). 2 provides a list of the standard cylinder IDs and fill codes and information for their mole-381 fraction assignments: t0 date, the best polynomial function fit coefficients relative to time t0 (ci, i=0,2) 382 and an estimated 1-sigma uncertainty. The uncertainty reported in SI Table 2 is empirically derived and 383 based on the standard calibration history and the standard deviation of the residuals to the best fit (the 384 assignment). The python code that used to calculates a secondary orthe tertiary standard assignment uses 385 a 0.5 ppb 1-sigma H9 reproducibility uncertainty which is added in quadrature to the measurement 386 episode standard deviation to account for longer term uncertainties not evident in the standard deviation 387 of the n-aliquots. We do not formally include an uncertainty for the secondary standard assignment. The 388 H9 reproducibility term is based on the mean of the standard deviation of residuals to the fit for the 389 calibration histories of secondary standards and target tanks over the period 2008-2022 (see section 2.3.1).

390 The 17 number of tertiary standards used successively on the flask analysisolder systems between 2009

and 2019 introduces time dependent issues due to the variable rate of H₂ drift in aluminum tanks and the frequency of the tank calibration historiess on the calibration system. Some of the H11-tertiary standards only have pre-deployment calibration results which do not assess drift during use (CC71649, CA04505, CC105491) and other standards have calibration results during their time in use but do not have post deployment calibrations that may help us evaluate the drift rate for the last couple of weeks or months of use (SI Table 2, notes in column "N"ND46735, ND33801, CB11551, CB11090, CA08107). ThreeA few standards exhibited an increased drift rate towards the end of their life that we did not capture with their infrequent calibrations on H9. This change in drift behavior was observed as increasing biases for measurements of target air tanks and daily test air flasks (see section 3.2). We have applied offline mole fraction corrections to the flask air analysis H₂ results to correct for the end of use drift increase for these three tertiary standards CC71649 (H11), CB11551 (H11) and CC305198 (H8), and their-standards' assignment uncertainty is larger for these time periods (SI Table 2).

403 Since August 2019, the GML has used a newer analytical system (called MAGICC-3 system operates 404 with) in the flask air analysis laboratory with a GC-HePDD (instrument code H8 and later H11) for H₂, 405 new optical analyzers for CO₂, CH₄ (CRDS, Picarro), CO and N₂O (QC-TILDAS, Aerodyne), and a 406 GC-ECD for SF₆. The responses of the instruments on MAGICC-3 are calibrated at the same time using a 407 single set of 11 standards spanning a range of mole fractions for the six trace gases. The MAGICC-3 408 standards were filled at the Niwot Ridge standard air preparation facility on a few different days between 409 December 2017 and May 2018. Their H₂ mole fractions are regularly isare regularly measured on H9 410 against the GML H₂ primary standards.

411 For the MAGICC-3 instrument response calibration, the eleven standards are analyzed sequentially 412 relative to an uncalibrated reference air tank (filled at Niwot Ridge). Air from each standard is injected 6 413 times alternating with the reference air. This entire sequence takes close to 17 hours. The first injection of 414 each standard is often biased low by about 2 ppb for H_2 due to timing issues at the start of each standard 415 sequence and only the remaining 5 injections are used to obtain the average normalized peak height 416 "signal" for each standard.

417 For H_2 , a subset of 8 of the 11 MAGICC-3working standards are used to determine-ealibrate the 418 GC-HePDD response. The time-dependent H_2 value assignment for each standard was derived from 8 or 419 9 calibration events on H9 between June 2018 and December 2022, listed in Table 3 (SI Table 3, SI 420 Figures 1 and 2). We plan on analyzing the MAGICC-3 standards 2 to 3 times a year going forward. The 421 standards' H_2 assignments will be revised as needed. The three cylinders that are not used exhibit complex 422 H_2 growth that is not well captured with periodic calibration episodes and a linear or quadratic fit.

423 The time between MAGICC-3 instrument response calibration sequences was 2 weeks for the first 3 424 months of service of MAGICC-3 and it has been increased to 4-5 weeks as we found the results to be 425 quite stable. A reference air cylinder will last 9 to 12 months on MAGICC-3. When the MAGICC-3 426 reference air cylinder is changed (pressure close 250 psia), a new instrument response calibration episode 427 is done with the new reference air cylinder before flask air samples are analyzed.

428 For the asynchronous calibration to stay valid for up to 5 weeks requires the reference gas composition 429 for the six measured gases to be stable between successive calibration episodesdates. This has been true 430 so far except for one reference air cylinder CA04145 for which a small time dependent H₂ correction was 431 applied between two instrument response calibration dates from 2019-11-06 to 2020-01-16 (see SI Figure432 3 and more details in SI section S4).

433 2.3 Calibration scale transfer quality assurance

434

GML target air tanks are dedicated air mixtures used for measurement quality control over multiple =years
periods. Most are high pressure aluminum cylinders filled at the Niwot Ridge standard preparation
facility. The analysis of target air helps us evaluate the robustness of the calibration scale transfer, and the
consistency of measurements over time and also between different analytical systems. In a perfect
program, we should be able to reproduce a measurement result for a target air tank every time. As noted
earlier, however, the reality is more complicated as H₂ tends to grow with time in aluminum cylinders.
Tracking many aluminum cylinders provides a diverse history of behaviors (stable, or linear vs non-linear
drift), and aids in the understanding of similar cylinders used for calibration.

443

444 2.3.1 Calibration system (H9) Target air tanks

445

446 Some GML target air cylinders are used exclusively to evaluate the stability and performance of the H9 447 measurements. Other target air cylinders are analyzed on H9 and in the flask air analysis laboratory on the 448 H8 and H11 instruments to evaluateunderstand the scale transfer.

449

While H_2 has been increasing in most of our target air tanks, eleven H9 target air tanks have shown either 451 stable H_2 or a linear rate of increase less than 1 ppb/yr. Figure 4 shows the calibration histories for these 452 tanks as well as the residuals from the best fit for each tank. Table 24 has a list of these target tanks and 453 several others binned by linear drift rate. More details for target tanks and their trend best fit coefficients 454 are in SI Table 4-S1. For each bin, the standard deviation of the residuals (differences of the H9 455 calibration results minus the best fit values) is below 0.5 ppb. The standard deviation of all linearly 456 drifting target tanks residuals binned together is 0.4 ppb.

457

458 The regular analysis of target tanks on H9 (right after the instrument response has been calibrated against 459 the primary standards) is used to evaluate the robustness of the calibration scale transfer in GML. Results 460 for tanks with stable or very slowly drifting H₂ indicate that between 2008 and 2021, the scale transfer on 461 H9 has low uncertainty (< 1 ppb).

462 🖷

We have eleven other target tanks for which the best fit to their calibration history is a quadratic function
(SI Figure 4 and SI Table 4S1). The standard deviation of these tanks' residuals binned together is 0.7
ppb. The current set of H9 target air tank results show that residuals for higher mole fraction (>650 ppb)
tanks have a larger standard deviation (0.5-0.8 ppb, SI Figure 4d).

467

468 One tank (CC309852 A, fill date 2009-10-01) with a quadratic drift correction is on the lower end of the

469 GML calibration range with a H₂ mole fraction that grew from 204 ppb in 2009 to 232 ppb in 2021. The

- 470 standard deviation for this target tank residuals is 0.93 ppb. It appears that a quadratic fit does not capture
- 471 the observed growth very well. If we reject the first calibration result in 2009 and only fit the other
- 472 2011-2021 results showing the H₂ mole fraction increasing from 217 ppb to 232 ppb, the best fit is a still a
- 473 quadratic function but with smaller coefficients (c1 = 1.66 ppb and c2 = -0.16 ppb) and the standard

474 deviation of the residuals to this fit is reduced to 0.36 ppb. Similarly, other-Some tanks that were analyzed 475 soon after fill and over several years show a similar-rapid and large initial growth in H_2 (in the first 0.5-2 476 years after fill). In this scenario, the residuals to a best linear or quadratic fit of the full calibration history 477 will be larger and will likely not capture the tank time-dependent H_2 assignment as accurately. For a few 478 of the GML standard and target air tanks, we dropped early calibration results that would bias the best fit 479 derivation and assignment during the time of use of the tank.

480

481 2.3.2 Comparison of measurements of gas mixtures in cylinders with MPI-BGC

482

483 Since 2016, the MPI-BGC GasLab has organized same tank air measurement ("MENI") comparisons 484 between WMO GAW partner laboratories as part of the European ICOS (Integrated Carbon Observation 485 System) Flask and Calibration Laboratory quality control work. In this program, three high pressure 486 eylinders (10L aluminum cylinders (Luxfer UK) are filled with dry airprepared and maintained by the 487 MPI-BGC and sent to measurement laboratories in a round robin loop. Two of the three cylinders had the 488 same air mixture and showed small growth in their H₂ mole fractions over time. The third cylinder 489 contains an "unknown" new mixture for each round robin loop.

490

Between 2016 and 2021, the MENI cylinders came to GML three times and were analyzed two to four 492 times on the H9 instrument during each round robin stop (see SI Figure 5). Some results were rejected 493 due to poor instrument performance or the use of an alternate calibration strategy than the one used to 494 transfer the scale. For the blind and ambient and blind H₂ MENI cylinders the retained NOAA H₂ results 495 agree well with the MPI_BGC measurements (< 1 ppb difference, SI Figure 5 a,b). For the low H₂ 496 cylinder, the 2017/2018 NOAA measurements are biased low by about 2 ppb while the March 10, 2021 497 result is about 2 ppb higher (SI Figure 5c). The MENI program provides an important valuable on-going

498 check for the MPI X2009 H₂ calibration scale transfer in GML.

499

500 2.3.3 Flask analysis systems (H8, H11) target air tanks

501

Figure 5a shows the calibration histories on H9 for target air tanks used in the flask analysis laboratory between 2009 and 2022. H₂ increased in all the target tanks, sometimes rapidly, requiring time dependent value assignments. These time-dependent H₂ assignments are derived from the best linear or quadratic fit to the calibration results on H9. These assignments can be compared to the measurements on the flask analysis systems to evaluate the quality of the scale transfer for the flask analysis system. It should be noted that the non-linear drift of some of these tanks may not be well modeled by a simple quadratic function, leading to higher uncertainty in the value assignments. This is especially true for tanks with limited ealibration histories or gaps in their calibration histories.¶

510

511 Three H₂ target air tanks were in service between 2009 and 2019 and have been used to evaluate the GML 512 calibration scale transfer to the MAGICC-1 and MAGICC-2 H₂ measurements (CC1824-(H), CB08834

513 (B) and CC303036 (A)). These tanks, however, exhibited rapid and large drifts and were not measured on

514 H9 on a regular basis making it more difficult to use them to evaluate potential biases on MAGICC-1 and

515 MAGICC-2 over this time period.

517 The target air tanks ALMX067998 (C) and CB11143 (C) entered service in 2016 and 2019 respectively 518 with more frequent measurements on the calibration system to better define their time dependent value 519 assignments. A new set of six target air tanks were prepared at the Niwot Ridge facility in late 2019 for 520 the MAGICC-3 system. They have been analyzed on MAGICC-3 multiple times a year but only one of 521 them has a H₂ mole fraction that remained below 700 ppb: CB10292 (B).

522

523 With the caveats that the non-linear drift in aluminum cylinders may not be well modeled by a simple 524 quadratic polynomial and that many of the early target tanks were under calibrated, the best polynomial fit 525 to the calibration records for all target air tanks give residuals smaller than 1.2 ppb (Figure 5b). Details for 526 the target tanks, including the best fit coefficients and the standard deviation of residuals to the fits are in 527 SI Table 5S2. The uncertainty on the assignments is larger during extended time periods with no 528 calibration results especially for the 3 earlier target air tanks with a limited calibration history (CC1824 529 (H)) or with calibration histories showing evidence of a change in the drift rate (CB08834 (B) and

530 CC303036 (A), see Figure 5).¶

531

532 In Figure 6, we show the differences between the target tankair analysis results on H8 and H11 and their 533 time-dependent H₂ assignments (based on the best fit to their calibration histories on H9 discussed above). 534 The differences are all within 4 ppb, however there are elearly-times when there are persistent biases 535 between the flask analysis system(s) and the calibration system. Uncertainties on the value assignment of 536 the target air tanks, the value assignments and stability of the standards used to calibrate the flask analysis 537 systems as well as the noise in the H8 and H11 measurements all contribute to the observed differences. 538 Similar offsets on both flask analysis systems (for example CC1824 prior to 2012) may point to the main 539 uncertainty contribution being from the value assignment of the target air tank. Different patterns in the 540 offsets between the two flask analysis systems (for example offsets of different signs for CC303036 (A)-541 and CB08834 (B)-on H8 and H11 in 2011-2013) suggest the offsets are due to value assignments of the 542 flask analysis system standards. Again, this is often due to limited calibration histories not being able to 543 fully map the non-linear drift in the standards. It also indicates there are times with systematic differences 544 (mostly < 2ppb) between the MAGGIC-1/H11 and MAGICC-2/H8 measurements in the flask records. 545

546 The full transition to the new MAGICC-3 system for flask analyses in August 2019 is indicated by the 547 vertical bar in Figure 6. As discussed earlier, one improvement in this new system is that H_2 548 measurements are now calibrated using a multi-point calibration curve from a suite of standards. This 549 makes the measurement results less sensitive to drift or value assignment error in any individual standard 550 since we are fitting multiple standards. These standards are used once a month and thus have much longer 551 lifetimes and longer calibration histories. We also now appreciate the complex H_2 growth patterns that can 552 occur in aluminum cylinders so have undertaken regular calibrations to ensure drift is tracked closely. 553 These changes seem to have reduced the bias observed between the flask analysis system and the 554 calibration system, which gives confidence that future measurements will be higher quality. 555

556 To help us monitor the H_2 calibration scale propagation performance going forward, a new target air tank 557 in an Essex stainless steel cylinder, SX-1009237, was filled in late 2022 to augment the current target 558 tanks. This target air tank should be stable for H_2 and will be used for periodic comparison between 559 measurement systems. Analysis results on H9 and H11 in December 2022 are 526.75 and 527.15 ppb, 560 respectively, consistent with the residuals for other target air tanks at that time.

562 3 NOAA flask air H₂ measurements

563

564 Close to 6000 flask air samples from the Cooperative Global Air Sampling Network are analyzed in 565 GML every year. The network sites are chosen carefully to be representative of large scale air masses and 566 to be able to rely on local support for sampling and shipping logistics. The reprocessing and release of the 567 2009-2021 H₂ global network flask air measurements on the MPI X2009 scale was made possible because 568 of continued efforts to conduct and improve the H₂ measurements, to store all the necessary data, and to 569 develop and update the tools for reliable and traceable reprocessing, comparison, and archiving. 570

571 **3.1 Recapitulation of the GML flask air H₂ analysis system configurations since 2009**¶ 572 ¶

573 The MAGICC-2 H8 and MAGICC-1 H11 instruments started routine flask air H2 analyses on November

574 5, 2009 and February 9, 2010 respectively. The flask air analysis results have been reprocessed using the

575 tertiary standards or working standards time-dependent H₂ assignments on the MPI X2009 scale. As-

576 mentioned earlier, those flask air measurements on H8 and H11 until July 2019 relied on calibration with

577 a single tertiary standard also used as reference to normalize the air sample chromatogram H₂ peak height. ¶

579 After ABeginning in August 2019, the MAGICC-3 response was determined using multiple standards.-

580 multi-standard response the MAGICC-3 system uses a multi-standard instrument response calibration.

581 For H₂, the instrument response curve is derived from eight working standards "known" assignments and

582 their normalized H₂ peak heights, with a reference air that is not a standard. H8 was the H₂ instrument on-

583 MAGICC-3 until September 11, 2020 when it was replaced by H11 which has better precision. The linear

584 fit coefficients for the MAGICC-3 H8 and H11 response curves are stored and used to calculate the H2 dry

585 air mole fraction in unknown air samples.¶

586 ¶

587 3.1₽ Data quality assurance and quality control

588 ¶

589 GML flask air H₂ measurements data quality is evaluated using results from the daily analysis of test air 590 flask pairs and from the agreement between South Pole Observatory (SPO) flasks pairs. \P

591

592 In this section, we first describe the flask sample collection protocol and introduce the data quality control 593 tags used to document sample and measurement data quality issues. GML flask air H₂ measurements data 594 quality is evaluated using results from the daily analysis of test air flask pairs and from the agreement 595 between South Pole Observatory (SPO) flask pairs. Then we assess the GML H₂ measurement short-term 596 noise (repeatability) with statistics from test air flasks and SPOSouth Pole Observatory flask pair

597 differences. Finally, we present a preliminary estimation for the uncertainty of flask air H_2 measurements 598 over 2009-2021, that includes empirical uncertainty estimates for the standards' assignments and the 599 short-term noise of the instruments.

600

601 3.12.1 Flask air sample collection overview and data quality tagging

602

603 Partners in the NOAA Cooperative Global Air Sampling Network collect whole outside air samples in 604 glass flasks in pairs, upwind from any local sources of pollution, people and animals and away from 605 structures or terrain that would affect the wind flow. Two 2.5L glass flasks with two glass stopcocks with 606 Teflon o-rings are connected in series in a portable sampling unit (PSU) made of a rugged case, a battery, 607 a pump, an intake line, and a mechanism to control the pressure of the air samples. Most sampling units 608 include a dryer and are semi-automated, with the exception of those used at relatively dry high latitude 609 locations and a few other locations where a more rugged, manually operated sampling unit is required. At 610 most sites, the operator will carry the equipment outdoors to conduct the sampling. At a few sites, the 611 PSU is indoors and connected to a fixed inlet line drawing air from the outside.

612

613 Before flasks are shipped to sampling sites,ment, the glass flasks are filled with synthetic air in the GML 614 flask logistics laboratory. During the sample collection on site, the flasks are first flushed for several 615 minutes and then filled to a pressure of 4 to 5 psi above ambient pressure in about 1 minute (See video: 616 https://gml.noaa.gov/education/intheair.html).

617

618 Air sample collection and/or measurement issues that are documented or detected and known to affect a 619 sample quality or an analyte measurement result are recorded with data quality control tags in our internal 620 database. For each flask air measurement, internal data quality control tags are translated into a simpler 3 621 column flag indicating if the measurement is retained or rejected for external data users. The GML flask 622 air samples and measurements can also have informational tags and comments, for example if another 623 measurement laboratory analyzed an air sample before it came to GML for analysis (see same air flask 624 measurement comparisons in section 3.3).

625

The global network flasks are filled to target pressure of 17-20 psia, but the final fill pressure can vary by 3-4a few psi, with some of the higher altitude sites having final pressures on the lower range typically. If an air sample pressure is too low for the H₂ GC instrument on the MAGICC system, the H₂ measurement result is tagged as "rejected" for low sample pressure. If H₂ measurements in paired flasks have a 5 ppb or larger difference, the results for the pair are tagged as rejected. If only one member of the pair had an obvious issue (leak, low flask air pressure), only the H₂ measurement for that member is tagged as rejected. Some issues are detected by the MAGICC performance control system and are tagged automatically. Other issues are tagged manually by scientists as part of regular data quality control checks. Scientists also verify the validity of the automatic tags. Members of the team routinely evaluate if follow-up actions are needed to fix a sample collection or measurement issue or reduce the chance of rejecting future sample results for the same issue.

637

638 Some sites can experience brief high-pollution episodes with the H_2 mole fractionsmeasurements in both 639 members of a pair meeting the pair agreement criteria but also being outliers, i.e. outside of the expected 640 long-term variability at the site [Novelli et al., 1999]. Gross H_2 outliers are typically "tagged" manually. A 641 statistical filter is applied before each annual data release [Dlugokencky et al., 1994]. For each site, we 642 applied a smoothing curve fit calculation to determines the time series mean behavior broken down in a 643 long-term trend, a seasonal cycle, and shorter-term (hours to weeks) variations [Thoning et al., 1989; Tans 644 et al., 1989a]. The code is available and a link is provided further down. Measurements that show large 645 residuals from the fit are not representative of the typical background air composition at the site and are 646 tagged as outliers [Novelli et al., 1999]. Gross H_2 outliers are typically "tagged" manually. We also apply-647 a A statistical filter is applied before each annual data release [Dlugokencky et al., 1994]., which The 648 filter works iteratively to find and tag outlier H_2 measurements when their residuals to the smooth curve 649 fit is larger than 3 to 4 times the time series residuals' standard deviation.

650

651 3.12.2 Test air flask analysis results

652

653 Besides the regular analysis of target cylinders, the MAGICC flask analysis system is also routinely-tested 654 daily using flasks filled with "test air" (flasks with site code "TST"). TST flasks are filled in one batch 655 with air from one of We have four rotating high pressure aluminum cylinders for test air (AL47-104, 656 AL47-108, AL47-113, AL47-145), themselves filled at the Niwot Ridge standard preparation site. SI 657 Figure 6 shows their calibration histories on H9 for different fills for these four test air cylinders. H₂ is not 658 stable in the "test air" cylinders and for some tank-fills, H₂ increased rapidly and grew beyond our 659 calibration range upper limit of 700 ppb.

660

661 Every 2 to 3 weeks, the GML flask preparation and logistics laboratory manager fills an even number of
662 TST flasks (14-24) are filled from the same test air cylinder. On typical analysis days, the MAGICC flask
663 air measurement sequence will start with the analysis of air from two TST flasks with the same fill date.
664

665 Global network flask air samples are analyzed at NOAA GML only during the daytime to ensure the 666 system operator is overseeing the full analysis cycle and minimizing the time a flask valve is open for the 667 analysis. This is meant to minimize the risk of losing or contaminating the air samples as many of them 668 are subsequently sent to the University of Colorado Boulder Stable Isotopes Laboratory for CO_2 and CH_4 669 isotope analyses.

670

671 Results from the TST flask pairs with the same fill date and analyzed on successive days give an 672 indication of the short-term repeatability of the measurements. Here, the deviations from the mean H_2 in 673 TST flasks with the same fill date are evaluated. For fill dates with a mean H_2 mole fraction less than 700 674 ppb, we calculate the differences between individual TST flask H_2 and the fill date mean. The standard 675 deviation of the TST flasks H_2 differences from their fill date mean is 1.39 ppb on MAGICC-2/H8 676 (N=872), 0.73 ppb on MAGICC-1/H11 (N=3583), 1.55 ppb on MAGICC-3/H8 (N=504) and 0.68 ppb on 677 MAGICC-3/H11 (N=1085), reflecting the higher measurement noise on H8.

679 Another diagnostic is the comparison of the TST flasks H_2 results and their test air cylinders' 680 time-dependent assignments for the dates the TST flasks were filled based on the best fit of the H9 test air 681 tank calibration results. This analysis is limited to the test air cylinders and fill code(s)-with less than 700 682 ppb H_2 and with tank calibration results on H9 that reasonably capture the increase in H_2 : AL47-108 (F), 683 AL47-113 (D,E,G), AL47-145 (F,G), AL47-104 (I). In SI Figure 7 (a-c), we show the H_2 differences 684 between the TST flask results and their test air cylinder assignments. The differences reflect noise in the 685 flask air measurements and uncertainties (and potentially small biases) in the test air tank-fill assigned H_2 . 686

687 Between 2010 and 2021, ∓the three fills of test air cylinder AL47-113 are in the ambient range and have

688 the most stable H_2 mole fractions. The tank-fill assigned H_2 linear drift rate is 1 ppb/yr in fill D, null in fill

689 E and 0.4 ppb/yr in fill G. Table 35 shows the mean and standard deviation of the differences in H_2

690 between TST flasks and the assigned H_2 in a stable or slowly drifting test air tank-fill. The biases for these

691 subsets of TST air data are less than 1 ppb and the standard deviation is equal to or less than 1.5 ppb and692 is smaller for the most recent MAGICC-3/H11 configuration which has a smaller number of data points.693

694 3.12.3 South Pole Observatory: H₂ differences in flask pairs

695

NOAA GML operates four staffed atmospherie baseline observatories (<u>https://gml.noaa.gov/obop/</u>). The
South Pole Observatory (SPO) in Antarctica and the Mauna Loa (MLO, Hawaii) observatories were built
in connection with the 1957-1958 International Geophysical Year, a global effort bringing together 67nations to study the Earth and in connection with the first launches of artificial satellites in Earth's orbitby the USA and the Soviet Union. The South Pole Observatory in Antarctica was established with supportfrom the US National Science Foundation and NOAA. The other two observatories near Utqiaġvik,
formerly Barrow, (BRW) and Samoa (SMO) were established in 1973 and 1974 respectively. H₂ timeseries for the observatories are shown in section 4. Two flask pairs are typically collected weekly and
close in time at the four NOAA atmospheric baseline observatories using two collection methods. In
method 'S', flasks are filled inside a building by tapping the air continuously pumped for analysis on an
in-situ GHG measurement system. Method, 'P' (or 'G') involves using a portable sampling unit with an
inlet mast and pump set up outside the building, similarly to other global network sites.¶
The South Pole Observatory (site code SPO, sampling location: 89.98°S, 24.80°W, 2815 meters above sea

710 level (masl)) gives scientists access to some of the "cleanest" air on Earth due to its remote location, and 711 thus provides an opportunity to use SPO flask data as a quality assurance tool... On site, GML and partners 712 operate in-situ measurements to monitor the atmosphere composition and properties, and whole air

713 samples have been collected for trace gas analyses in the GML laboratories in Boulder since 1975. ¶
714

715 All four NOAA atmospheric baseline observatories have an upwind clean air sector with no local sources 716 of pollution (https://gml.noaa.gov/obop/). Every week, scientists on location collect discrete air samples

717 when the near surface wind comes from the clean air sector. Two flask pairs are typically collected

718 weekly and close in time at the four NOAA atmospheric baseline observatories using two collection

719 methods. In method 'S', flasks are filled inside a building by tapping the air continuously pumped for

720 analysis on an in-situ GHG measurement system. Method, 'P' (or 'G') involves using a portable sampling

721 unit with an inlet mast and pump set up outside the building, similarly to other global network sites. The

722 scientist doing the air sampling is involved with both sampling techniques. Weekly samples with both

723 methods are typically conducted within minutes of each other. Both flask sampling methods give reliable-

- 724 results for H₂ at the South Pole Observatory.
- 725

726 Staff rotation and flask shipping to and from the South Pole Observatory happen during a limited time 727 window during the Austral summer. While awaiting shipment, SPO flask air samples are stored in crates 728 in a heated storage building. Every year, one large SPO flask shipment arrives in Boulder in 729 December/January and another smaller shipment arrives in February/March. A year's worth of flasks is 730 prepared and shipped to SPO during that same time window. Despite the longer storage for SPO flasks 731 before analysis, we have not detected biases in H₂ measurements of those samples when compared with 732 other high southern latitudes times series. SPO flask air H₂ measurements show close to a 20 ppb seasonal 733 cycle and a ~15 ppb increase in the annual mean levels between 2010 and 2021 (Figure 7). There is very little short-term variability in the surface air over Antarctica for long-lived GHGs, CO and Table H₂. The differences in the H₂ mole fractions in SPO paired samples therefore mostly reflect the short-term noise in the measurements. In SI Table 6 we report statistics for H₂ differences for the two flask sampling methods and the four measurement system configurations between 2009 and 2021 with H8 and H11. As observed for the TST flasks, measurements on H11 are less noisy than on H8, especially on the MAGICC-3 system. The average of the absolute differences for H₂ in SPO flask paired samples is less r41 than 2 ppb ($\sigma \le 1.3$ ppb) and methods S and P H₂ pair averages at SPO agree within 1 ppb on average (σ r42 ≤ 1.7 ppb).

743

744 3.12.4 Flask air H₂ uncertainty estimates

745

746 We have derived preliminary empirical uncertainty estimates for flask air H₂ measurements that fall in the 747 200-700 ppb range. For measurements on MAGICC-1 and MAGICC-2, the total uncertainty estimate 748 comes from the combination of two uncertainties added in quadrature: 1) the uncertainty on the H₂ tertiary 749 standard time-dependent assignment (SI Table 2) and 2) the instrument estimated repeatability (Table 48). 750 If an offline assignment correction is applied to take into account changes in a standard drift rate toward 751 the end of its use, the standard assignment uncertainty is increased. The H8 and H11 instrument 752 repeatability estimates are listed in Table 48. For now, we assume a 0.5 ppb uncertainty on the 753 MAGICC-3 instrument response calibrated with multiple standards. On-going work will allow us to 754 refine this last uncertainty component estimate at a later date. Typical 1-sigma uncertainties for GML 755 flask air H₂ measurements are 1.2 to 1.9 ppb on MAGICC-1, 1.4 to 2.8 ppb on MAGICC-2, 1.6 ppb on 756 MAGICC-3/H8 and 0.8 ppb on MAGICC-3/H11.

757 3.23 Comparison with other GAW laboratories H₂ measurements

758

A small number of laboratories operate well-calibrated long-term measurements of important atmospheric
trace gases. The WMO Global Atmospheric Watch (GAW) coordinates regular technical and scientific
discussions with experts from these laboratories. Another important outcome of the WMO/GAW
collaborations consists of routine comparisons to assess the data compatibility for measurements coming
from different laboratories and programs [Francey et al., 1999; Masarie et al., 2001; Jordan and Steinberg,
2011; Worthy et al., 2023]. The WMO/GAW network compatibility goals for measurements of H₂ in well
mixed background air is 2 ppb (see Table 1 in [WMO/GAW, 2020]). This means that for H₂^{*},
measurement records should not have persistent biases less than 2 ppb to be used in combination with
other qualifying measurements in global budget, trend and large scale gradient analyses.

769 GML participates in several WMO GAW measurement comparison efforts. Same-flask air measurement 770 comparisons consist of one member of a NOAA flask pair collected at a site being analyzed by a partner 771 laboratory before being analyzed by GML. Co-located flask air measurement comparisons involve 2 or 772 more measurement programs having samples collected at the same location and close in time. 773 Historically, these and other "intercomparison" projects have been abbreviated ICPs, which we use in the

774 text below. Here the GML flask air H_2 measurements data compatibility is assessed with results from 775 on-going ICPs.

GML conducts same-flask air measurement comparisons at the Cape Grim Observatory (CGO, 40.68° S, 144.69° W,164 masl) with CSIRO, Australia and at the Ochsenkopf mountain top tower (OXK, 50.03° N, 11.81° E, 1085 masl) with MPI–#BGC, Germany. Sampling at OXK was temporarily suspended between June 2019 and April 2021. The Alert/Dr Neil Trivett Observatory (ALT, 82.45° N -62.51° W, 190 masl)
has facilitated the largest multi-laboratory flask air comparison experiment in the WMO GAW program [Worthy et al., 2023]. NOAA has colocated flask air samples from ALT with CSIRO and the MPI-#BGC.
The CSIRO and MPI-#BGC H₂ measurements are also traceable to the MPI X2009 calibration scale.

785 In Table 57, we summarize the annual mean of the differences for H_2 measurements from different 786 laboratory and flask combinations (same flask, same flask pair or cołlocated flasks) for CGO, OXK and 787 ALT between 2010 and 2021. Columns 2 and 3 show the annual means of the NOAA H_2 differences 788 between the ICP flask and its pair mate at CGO and OXK. All measurements included in the comparisons 789 are retained, meaning they have passed quality control checks.

790

791 Columns 2 and 3 show the annual means of the NOAA H_2 differences between the ICP flask and its pair 792 mate at CGO and OXK. For CGO flask air samples collected before 2019, we find that the NOAA 793 analysis for the NOAA ICP flask first measured at CSIRO often shows higher H_2 than in the non-ICP 794 flask air sample. We suspect several of these ICP flasks had a small but detectable contamination for H_2 . 795 We have applied a rejection tag to NOAA analysis results for CGO ICP flasks with an H_2 mole fraction 2 796 ppb or more above H_2 in the non-ICP pair mate. This affected 165 ICP samples between 2009 and 2018 or 797 37% of all CGO ICP flasks collected between August 2009 and the end of 2021. ¶

799 For OXK, the NOAA analysis result for the ICP flask first measured at MPI-#BGC often shows slightly 800 higher H₂ than in the non-ICP flask (Table 57, 3rd column), and the annual mean bias is less than 1 ppb 801 for all years.

802

803 The last 4 columns in Table 57 show interlaboratory H_2 measurement comparisons for CGO, OXK and 804 ALT flask air samples. The annual mean differences are consistently less than 1.6 ppb for CGO (CSIRO-805 ICP flask and NOAA non ICP flask) and less than 2 ppb for OXK for 9 out of 11 years (MPI/BGC ICP-806 flask and NOAA ICP flask) (Figure 8). For colocated air samples at ALT (NOAA vs CSIRO and NOAA 807 vs MPI/BGC) we compare the mean of flask results for each laboratory and limit the comparison for 808 samples collected within 60 minutes of each other. The ALT annual mean differences vary from year to 809 year, and are less than +/- 2 ppb for 8 years out of 12 for the NOAA vs CSIRO comparison and for 7 810 years out of 10 for the NOAA vs MPI-/BGC comparison. These on-going ICPs are monitored regularly to 811 continually assess the NOAA H₂ data compatibility with data from GAW partners. 812

813 4. NOAA H₂ atmospheric H₂ time series

814 Previous measurement studies have described key features of the H_2 global distribution for different time 815 periods [Khalil and Rasmussen, 1990; Novelli et al., 1999; Langenfelds et al., 2002; Price et al., 2007; 816 Yver et al., 2011]. Some of the spatiotemporal features in the more recent NOAA H_2 measurement records 817 measurements at background sites are described in this section.

819 4.1 H₂ at the NOAA Cooperative Global Air Sampling Network Sites

820 There are 51 sites considered active or recently terminated in the Cooperative Global Air Sampling 821 Network (see map in SI Figure 8 and site information in SI Table 7). The H₂ measurement times series for 822 these sites are shown in SI Figure 9. Note that a few sites that have been discontinued are not shown in 823 this figure. A curve fit is run for each site time series based on Thoning et al. [1989]. First the code 824 optimizes parameters for a function made of a four-term harmonic and a cubic polynomial. The resulting 825 residuals (measurements minus function) are then smoothed with a low-pass filter with a 667-day cutoff 826 and are added to the polynomial part of the function to produce the "trend curve" (shown as the dark blue 827 line in SI Figure 9). The residuals are also smoothed with a low-pass filter with a 80 day cutoff and are 828 added to the function to produce a "smooth curve" at each site.

The data quality control work on our long-term measurement time series includes a data selection step with a statistical filter (see section 3.1.1). Samples with H_2 beyond 3.5 (4 for Ascension Island, ASC) standard deviations of the time series smoothed curve at each site are flagged as not representative of background air conditions and are shown as crosses in SI Figure 9.

833 The annual mean, maximum and minimum H_2 values of the smooth curve for the 51 sites are plotted in 834 Figure 9 (in order of decreasing latitude along the x-axis) for years with retained measurements up to 835 2021. Sampling at the TPI site, on Taiping Island, Taiwan, started in May 2019, which explains the 2 (full 836 sampling year) data points for the site. Sampling at a few network sites was impacted by the COVID-19 837 pandemic resulting in data gaps or delayed return shipping of samples. We recommend data users become 838 familiar with individual sampling site measurement records to best aggregate and interpret signals.

839 The interhemispheric gradient of H₂, with higher levels in the SH, is apparent in the annual means 840 distribution across sites (Figure 9, green circles). The majority of sites in the SH (BKT to SPO on the 841 right side of Figure 9) show smaller seasonal cycle amplitudes (<23 ppb) than NH sites; however, several 842 sites have interannual variations in their H₂ seasonal cycle amplitudes (SI Figure 9). Sites with the lowest 843 H₂ seasonal minima (Figure 9, blue x symbols) likely are the most influenced by soil uptake. A few sites 844 (for ex. TAP (Taiwan), AMY (Republic of Korea), LLN (Taiwan), CPT (South Africa)) show higher 845 smooth curve annual maxima (Figure 9, red crosses), likely reflecting upwind local or regional emissions.

846

847 4.21 H₂ at NOAA Baseline Atmospheric Observatories

848

849 NOAA GML operates four staffed atmospheric baseline observatories (<u>https://gml.noaa.gov/obop/</u>). The
850 South Pole Observatory in Antarctica and the Mauna Loa (MLO, Hawaii) observatories were built in
851 connection with the 1957-1958 International Geophysical Year, a global effort bringing together 67
852 nations to study the Earth and in connection with the first launches of artificial satellites in Earth's orbit
853 by the USA and the former Soviet Union. The South Pole Observatory in Antarctica was established with
854 support from the US National Science Foundation and NOAA. The other two observatories near
855 Utqiaġvik, formerly Barrow, (BRW) and Samoa (SMO) were established in 1973 and 1974 respectively.
856

857 All four NOAA atmospheric baseline observatories have an upwind clean air sector with no local sources 858 of pollution. Every week, scientists on location collect discrete air samples preferentially when the near 859 surface wind comes from the clean air sector. Figure 109 shows the reprocessed H₂ time series for the four 860 NOAA Baseline Atmospheric Observatories between 2009 and 2021. Valid "S" and "P" method flask air 861 H₂ measurements are retained for the South Pole Observatory only. The "S" method flasks show 862 contaminated H₂ at Samoa and show seasonalsome contamination or more variable H₂ at Utqiaġvik 863 (Barrow) until August 2021 when sampling started at a new tower with new sampling lines. The Mauna 864 Loa H₂ in "S" method flasks will be further evaluated and may be retained in future releases.

The Samoa and South Pole H_2 smooth curves flask air measurements show similar maximum levels between 550 and 570 ppb and slightly higher minima at Samoa compared to the South Pole. The seasonal maximuma occurs about -3 months earlier at Samoa than at the South Pole. The interannual variability is seasonal similar at both sites and is dominated by step increases on three occasions: in 2012/2013, 2016 and 2020.

869 The Mauna Loa H₂ time series shows more short-term variability than for Samoa and South Pole.=

870 reflecting the variable latitudes covered by an air mass before it is sampled at the high-altitude-

871 observatory and the strong spatial gradients for H_2 in the NH. The mean seasonal cycle amplitude of theat 872 Mauna Loa H_2 smooth curve is about 40 ppb with maximum levels in April-May and minimum levels in 873 December-January. The observed seasonal maximuma ranges from 550 to 580 ppb and the observed 874 seasonal minimuma ranges from 505 to 520 ppb. The NOAA-measurements indicate that annual mean H_2 875 levels at Mauna Loa after 20168 were higher than in previous years.

876 Of the four observatories, the Barrow H_2 time series shows the lowest levels and the strongest seasonal 877 cycle, about 60 ppb on average. The smooth curve observed seasonal maximum ranges from 520 to 540 878 ppb in April-May and the observed seasonal minimum in September-November ranges from 450 to 490 879 ppb.

Base Despite having larger emissions in the NH, the H_2 interhemispheric gradient shows lower levels in the extratropical NH. This is related to the larger land masses in the NH and the soil sink being the dominant green process for H_2 . Warwick et al. [2022] report model-based estimates for the H_2 lifetime of 8.3 years for the OH sink (from the authors base model configuration) and of 2.5 years for the soil uptake (average of existing literature studies). In their flux inversion, Yver et al. [2011] estimated that the NH high latitudes and the tropics represent 40% and 55% of the global soil sink respectively. The soil sink and OH sink in extratropical northern latitudes both peak in summertime [Price et al., 2007] leading to the posterved stronger H_2 minima.

888 Given the larger variability and stronger seasonality of H_2 in the NH extra tropics, ilt is important to look 889 at data from multiple sites to study and detect interannual and potentially long-term large-scale changes in 890 atmospheric H_2 levels. In the next section, we highlight a few features in the global network H_2 records 891 and present background air zonal mean H_2 time series based on samples collected at marine boundary 892 layer sites.

893 4.2 H₂ at the NOAA Global Cooperative Air Sampling Network Sites

894 There are 51 sites considered active or recently terminated in the Global Cooperative Air Sampling-

895 Network (see map in SI Figure 8). The H₂ measurement times series for these sites are shown in SI Figure-

- 896 9. Note that a few sites that have been discontinued are not shown in this figure. A curve fit is run for-
- 897 each site time series based on Thoning et al. [1989]. First the code optimizes parameters for a function
- 898 made of a four-term harmonic and a cubic polynomial. The resulting residuals (measurements minus-

899 function) are then smoothed with a low-pass filter with a 667-day cutoff and are added to the polynomial

900 part of the function to produce the "trend curve" (shown as the dark blue line in site time series plots in SI

- 901 Figure 9). The residuals are also smoothed with a low-pass filter with a 80 day cutoff and are added to the
- 902 function to produce a "smooth curve" at each site.

903 The data quality control work on our long-term measurement time series includes a data selection step

- 904 [Dlugokeneky et al., 1994]. Samples with H₂ beyond 3.5 (4 for ASC) standard deviations of the time
- 905 series smoothed curve at each site are flagged as not representative of background air conditions and are
- 906 shown as crosses in SI Figure 9.¶
- 907 The annual mean, maximum and minimum H₂ values of the smooth curve for the 51 sites are plotted in
- 908 Figure 10 (in order of decreasing latitude along the x-axis) for years with retained measurements up to-

909 2021. Sampling at the TPI site, on Taiping Island, Taiwan, started in May 2019, which explains the 2 (full

910 sampling year) data points for the site. Sampling at a few network sites was impacted by the covid-19-

- 911 pandemic resulting in data gaps or delayed return shipping of samples. We recommend data users become
- 912 familiar with individual sampling site measurement records to best aggregate and interpret signals.¶
- 913 The interhemispheric gradient of H₂, with higher levels in the SH, is apparent in the annual means-
- 914 distribution across sites in Figure 10 (green circles). The majority of sites in the SH (BKT to SPO on the
- 915 right side of Figure 10) show smaller seasonal cycle amplitudes (<23 ppb) than NH sites; however,
- 916 several sites have interannual variations in their H₂ seasonal cycle amplitudes (SI Figure 9). Sites with the
- 917 lowest H₂ seasonal minima (Figure 10, blue x symbols) likely are the most influenced by soil uptake. A
- 918 few sites (for ex. TAP, AMY, LLN, CPT) show higher smooth curve annual maxima (Figure 10, red
- 919 crosses), likely reflecting upwind local or regional emissions.

920 Two sites, KUM and WIS, had a change of sampling location that resulted in visibly different H₂ mean

- 921 levels and seasonal eyele amplitudes. In mid 2018, the KUM site was moved 30 km NNW along the
- 922 Hawaii island SE coastline when access to a lava field bordering the ocean was lost in the eruption of the
- 923 Kilauea volcano. The KUM sampling location change resulted in higher mean H₂ levels and a smaller-
- 924 seasonal cycle. The WIS site moved 100 km SSE in Israel in early 2015. There are more instances of
- 925 depleted H₂ (in December-March) since the move, potentially reflecting a stronger influence of soil-
- 926 uptake in air masses sampled at the newer location.

927 4.3 H₂ marine boundary layer global and zonal means

⁹²⁸ To extract large scale signals from the global air sampling network, we use the NOAA GML marine ⁹²⁹ boundary layer (MBL) zonal data product [Tans et al., 1989b; Dlugokencky et al., 1994]. Time series ⁹³⁰ from remote MBL sites are smoothed and interpolated to produce a latitude versus time surface of the H₂ ⁹³¹ mean MBL mole fraction (Figure 11). For H₂, the number of sites included in the zonal mean calculations ⁹³² ranges from 29-42 sites until July 2017 when sampling from the Pacific Ocean shipboard (POC) was ⁹³³ stopped, after which 24-27 sites were included in the calculation. Because the Cooperative Global Air 934 Sampling Network is sparse in the tropics and in the SH mid latitudes, the MBL product likely does not
935 equally detect and reflect interannual variability in fluxes in these under-sampled regions, for example
936 biomass burning emissions in Africa and South America.

937 To further isolate changes in background H_2 at different latitudes, we first calculate MBL global (and 938 zonal band) means (shown in SI Figure 10) and then derive anomalies by removing the 2010-2021 939 average year from the global and (or zonal band mean) time series. Figure 12 shows the MBL anomaly 940 for H_2 (black lines) and CO (dashed blue lines) for the global mean and 5 zonal band means (NH and SH 941 Polar (53-90°), NH and SH Temperate (17.5-53°) and Tropics (17.5°S to 17.5°N). The NOAA GML CO 942 measurements are for the same air samples as the H_2 measurements [Pétron et al., 2023b]. Here, we derive 943 the global and zonal means for CO using the 2009-2022 MBL CO measurements and the anomalies are 944 based on the 2010-2021 smooth curve zonal mean results to be consistent with the H_2 data analysis.

945 CO is emitted during incomplete combustion and is a useful marker of biomass burning emissions. CO 946 has a shorter atmospheric lifetime than H_2 which results in shorter-lived CO anomalies from pulse 947 emissions. The data reduction for the anomaly analysis is slightly different from Langenfelds et al. [2002] 948 investigation of CO₂, CH₄, H₂, and CO interannual variability in the CSIRO network 1992-1999 time 949 records. The CSIRO authors employed the same [Thoning et al., 1989] data smoothing technique as we 950 do but used the derivative of the trend curve to analyze correlations in interannual growth rate variations 951 between species. The anomaly approach chosen here allows to retain the timing of abrupt changes in the 952 measurement records.

953 Over 2010-2021, background air H_2 has increased at all latitudes (Figure 12). The global mean MBL H_2 954 shows a non-uniform increase over this time with a noticeable 10 ppb step increase in 2016. The global 955 mean MBL H_2 was 20.2 ±0.2 ppb higher in 2021 compared to 2010 (Figure 12a).

956 The meridional gradient and zonal band mean plots (Figures 11 and Figure 12b-f) highlight the evolution 957 of background air H_2 at different latitudes. By construction, the smooth curve anomalies are not directly 958 proportional to the biomass burning emissions that likely may have caused them. Rather the aAnomalies 959 in the smooth curves are useful to point to time periods when several successive air samples at a site show 960 similar deviations from the average seasonal cycle and multi-year trend.

961 The 2016 H_2 step increase is detected in the Tropics and SH. In the Tropics it coincides with a strong 962 positive CO anomaly that started in November 2015, reached a peak amplitude of 15 ppb mid-January 963 2016 and ended in May 2016. The 2015/2016 H_2 anomaly is first detected at Bukit Kototabang, Indonesia 964 (BKT) and later at Ascension Island (ASC), Cape Grim (CGO) and Crozet Island (CRZ) (SI Figure 11). 965 Some BKT air samples impacted by biomass burning emissions show enhancements of 100s ppb in CO 966 and H_2 . The BKT CO and H_2 data also show enhancements likely related to biomass burning in 2015. The 967 2015 fire season in Indonesia was among the most intense on record as shown by remote sensing products 968 of fire counts, CO and aerosols. Field et al. [2016] found that burning activities to clear peatland for 969 farming likely contributed to larger emissions than expected from dry conditions alone in 2015.

970 There is another step increase in the Polar SH zonal band in early 2020, also coinciding with a pulse971 anomaly in CO (Figure 12f) likely related to large wildfires in Australia in late 2019-early 2020. The972 Cape Grim (CGO) and Crozet Island (CRZ) smoothed curves show a large jump between the late 2019

973 minimum and early 2020 maximum when the CGO CO measurement seasonal minimum is also 10-12 974 ppb higher than in other years (SI Figure 11). van der Welde et al. [2021] estimate that the 2019-2020 975 fires in Australia emitted 80% more CO_2 than "normal" Australian annual fire and fossil fuel emissions 976 combined.

977 In the NH extratropics bands, positive anomalies in H_2 in 2021 coincide with CO pulse anomalies. For the 978 Polar (Temperate) NH zonal band, the CO anomaly lasts from mid-July (June) to December 2021 with a 979 peak in September and an anomaly maximum amplitude of 37 ppb (19 ppb). Record high emissions of 980 CO₂ and CO from boreal forest fires in Eurasia and North America in 2021 have been reported by Zheng 981 et al. [2023].

Previously, Simmonds et al. [2005] and Grant et al. [2010] have reported on the observed variability in 983 the Mace Head continuous H_2 measurement record and linked interannual variability in the baseline 984 annual mean H_2 to larger fire emission events. More recently, Derwent et al. [2023] shared an updated 985 analysis of the February 1994-September 2022 Mace Head in-situ H_2 measurements. The in situ record 986 shows higher monthly mean baseline H_2 levels in recent years and the authors report an increase in 987 monthly mean anomalies after December 2015 (slope of 2.4 +/- 0.5 ppb/yr). They postulate that a 988 "missing" source of increasing intensity after 2010 may be behind the observed sustained increased H_2 , 989 which is markedly different from the 1998-1999 anomalies attributed to biomass burning. Derwent et al. 990 [2023] explore potential candidates for the missing sources. However, in the absence of strong and 991 quantitative direct evidence at this time, additional studies are needed to interpret the observed H_2 992 variability.

993

994 5. Conclusions

995

996 In this paper, we have described how NOAA GML has adopted the MPI X2009 H₂ calibration scale. The 997 work was confined to measurements on GC-HePDD instruments. The GML H₂ primary standards in 998 electropolished stainless steel cylinders have been calibrated once by the MPI-BGC CCL in Fall 2020. 999 We have used the CCL assignments to propagate the scale to secondary and tertiary standards. H_2 1000 increases in most air standards stored in aluminum cylinders. A curve fit wais applied to each standard 1001 calibration history to determine a time-dependent H_2 assignment on MPI X2009. The secondary and 1002 tertiary and working standards H_2 assignments were then used to reprocess results for NOAA flask air H_2 1003 measurements on MPI X2009. These NOAA Cooperative Global Air Sampling Network flask 1004 reprocessed H₂ measurements for 2009-2021 are new-publicly available [Pétron et al., 2023a]. For the 1005 period 2010-2021, same air measurements with GAW partner laboratories have annual mean differences 1006 less than 2 ppb for the Cape Grim comparison with CSIRO and less than 3 ppb for the Ochsenkopf 1007 comparison with MPI BGC. Over 2010-2021, background air H₂ has increased at all latitudes. However, 1008 site time series and marine boundary layer H_2 zonal means show significant interannual variability. We 1009 find that some of strongest H_2 zonal mean anomalies coincide with CO anomalies and therefore were 1010 likely partly driven by large biomass burning events in Indonesia (2015), Australia (2019/2020), and 1011 boreal latitudes (2012 and 2021) [Field et al., 2016; Petetin et al., 2018; Zheng et al., 2023]. A full

1012 analysis of the NOAA Cooperative Global Air Sampling Network H_2 measurement records is beyond the 1013 scope of this paper. This dataset complements WMO/GAW partner laboratories H_2 measurements and it 1014 will be updated and extended routinely moving forward.

1015

1016 Data and Code Availability

1017 The NOAA global network flask air H₂ and CO time series are available at

1018 https://doi.org/10.15138/WP0W-EZ08.

1019

- 1020 We kindly request that users of the NOAA H_2 dataset cite:
- 1021 Pétron, G., Crotwell, A., Crotwell, M., Kitzis, D., Madronich, M.,
- 1022 Mefford, T., Moglia, E., Mund, J., Neff, D., Thoning, K., & Wolter, S.
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1024 Cycle Cooperative Global Air Sampling Network, 2009-2021 [Data set].

- 1025 NOAA GML CCGG Division. Version: 2023-05-25, https://doi.org/10.15138/WP0W-EZ081026
- 1027 The python class used to filter and smooth time series data is available and explained at:

1028 <u>https://gml.noaa.gov/aftp/user/thoning/ccgcrv/ccgfilt.pdf</u> and the method can be referenced as 1029 [Thoning et al., 1989].

1030

1031 Supplement

1032 The supplement for this article is available in a separate file.

1033

1034 Author Contributions

1035 GP and AC designed the scale revision work. GP, AC and JM implemented the scale revision. 1036 GP, AC, MC, MM, DN and JM contributed to the data quality control. GP and JP analyzed 1037 network site time series. AC designed, built and oversaw the H₂ calibration scale transfer and the 1038 flask air analysis system operations, working with Paul Novelli until he retired in 2017. TM and 1039 AC carried out tank calibrations. BH prepared the primary standards. DK was in charge of the 1040 whole air secondary and tertiary standards, reference, target and test air tanks preparation. MM 1041 and EM were responsible for the flask air analysis lab operations, working with Patricia Lang 1042 until her retirement in 2019. EM managed the flask logistics laboratory and flask metadata 1043 entries. DN with support from SW managed the NOAA Cooperative Global Air Sampling 1044 Network. DN managed sampling equipment for sites. JM manages the database and date 1045 releases.⁺, JM, KT and AC developed code and user interfaces for data processing, quality control 1046 and exploration. AJ calibrated the NOAA primary standards. AJ, PK and RL contributed data 1047 from their measurement programs. GP prepared the manuscript with contributions from AC and 1048 AJ and edits from BH, MC, RL, and JP.

1049

1050 Competing Interests

1051 The authors declare that they have no conflict of interest.

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1061 Gary Morris, and Kathryn McKain, Simon O'Doherty and an anonymous referee provided 1062 valuable comments on the manuscript.

1063

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1290 Tables

1291

Serial Number	Fill code	Fill Date	CCL value	CCL uncertainty
SX-3558	А	2008-10-17	248.4	0.1
SX-0614470	А	2019-04-15	352.8	0.1
SX-3543	В	2008-11-03	425.4	0.2
SX-3540	В	2007-08-07	488	0.2
SX-0614471	А	2019-04-19	496.5	0.3
SX-3523	С	2007-07-24	527	0.2
SX-3554	А	2007-08-02	601.2	0.2
SX-0614472	А	2019-04-19	701.9	0.2

1292 Table 1. NOAA GML H_2 primary standards (prepared gravimetrically) and their WMO/MPI X2009 **1293** assignments (dated 2022-02-18). All H_2 dry air mole fractions and their uncertainties are in ppb.

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1296 Table 2: H_2 secondary standards used in the tank calibration laboratory and H_2 tertiary standards used on 1297 the MAGICC-1 and MAGICC-2 systems (2009 to 2019).

Tank Calibration / H9¶										
Tank-ID- (fill)¶	Time of use¶	t0¶	Assignment at t0 (ppb)¶	C1 (ppb/yr)¶	C2¶ (ppb/yr²)¶	N¶	Residuals-standard- deviation (ppb)¶	Fill-date¶		
CC119811- (∧)¶	2/05/2008 to 6/02/2013¶	2010.0689¶	549.4¶	2.0¶	₽¶	47¶	0.50 ¶	01/01/2006¶ SM*¶		
CA03233- (₿)¶	6/02/2013 to 11/01/2018¶	2016.7106 ¶	502.8¶	₽¶	₽¶	19 ¶	0.23¶	08/12/2010¶ NWR¶		
MAGICC-1 / H11¶										
Tank ID − (fill)¶	Time of use¶	tO¶	Assignment at t0 (ppb)¶	C1 (ppb/yr)¶	C2¶ (ppb/yr²)¶	₩¶	Assignment uncertainty (ppb)**¶	Fill date¶		
CA08107- (D)¶	7/22 to- 8/7/2019¶	2019.2959 ¶	562.9¶	15.4 ¶	₽¶	6¶	0.6¶	11/9/2018¶ NWR¶		
€B11090- (B)¶	10/18/2018 to 7/19/2019¶	2019.1482¶	576.3¶	6.9¶	₽¶	4a¶	0.6¶ After 2019-06-21:¶ 1.5¶	9/30/2016¶ NWR¶		
CB11551 (A)¶	2/13 to 10/17/2018 ¶	2018.1878¶	548.8¶	6.7¶	₽¶	3a,b, €¶	0.5¶ After 2018-08-27.¶ 1.5¶	1/1/2015¶ SM*¶		
CC91285- (C)¶	6/19/2017 to- 2/13/2018¶	2017.1711 ¶	538.4 ¶	₽¶	₽¶	8¶	0.5 ¶	8/14/2015¶ NWR¶		
CA08165	10/13/2016 to	2016.9137¶	535.7¶	4 .5¶	0 ¶	3e ¶	0.5¶	12/16/2011¶		

(₿) ¶	06/16/2017¶							NWR¶
CC302566- (₿)¶	3/21/2016 to 10/12/2016¶	2016.3645¶	540.2¶	4.4¶	₩	5¶	0.5¶	8/14/2015¶ NWR¶
CC105491- (₿)¶	8/10/2015 to 3/18/2016¶	2015.1506¶	522.3¶	₽¶	0¶	5d¶	1.0 ¶	1/16/2014¶ NWR¶
ND33801- (B) ¶	8/4/2014 to 8/7/2015¶	2013.8771¶	509.3 ¶	0.9¶	0¶	6e¶	0.5-¶ After 2015-05-14:¶ 1.0-¶	12/27/2012¶ NWR¶
CB09117- (A)¶	2/18 to 8/1/- 2014¶	2013.8912 ¶	635.3¶	28.7¶	0¶	5¶	21	12/17/2012¶ SM¶
ND46735- (A)¶	9/10/2012 to 2/13/2014 ¶	2012.9158 ¶	527.4 ¶	2.5¶	-1.0¶	7e,f ¶	0.5¶ After 2013-12-11. ¶ 1.0 ¶	1/1/2011¶ estimated¶
CA04505- (B)¶	12/9/2011 to 9/7/2012¶	2011.4593 ¶	540.6¶	1.7¶	₽¶	3e,d¶	1.0¶	8/12/2010¶ NWR¶
ND38963- (A)¶	8/12/2010 to- 12/7/2011¶	2011.704 ¶	586.0¶	6.2¶	0¶	4¶	0.5¶	1/1/2009¶ estimated¶
CC71649- (E)¶	1/22 to 8/6 2010 ¶	2009.1184¶	507.1¶	8.4 ¶	₽¶	7b,e¶	1.5¶	9/19/2008¶
			М	AGICC-2 / H8	Ŧ			
Tank ID- (fill)¶	Time of use¶	fi	Assignment at t0 (ppb)¶	ei¶	C2¶	N	Assignment uncertainty (ppb)**¶	Fill date¶
ND38954- (B)¶	3/26/2013 to 3/21/2014¶	2014.2094¶	516.6¶	2.0¶	θ¶	5¶	0.5¶	12/9/2012¶ NWR¶
CA03409- (₿)¶	5/23/2011 to 3/25/2013¶	2011.6278¶	526.6¶	0¶	0¶	5e¶	0.5¶ After 2013-01-21:¶ 1.0¶	1/1/2010¶ estimated¶
ND38415- (A)¶	4/5/2010 to 5/20/2011¶	2010.2502¶	566.1¶	20.9¶	-8.7¶	6¶	0.5¶	1/1/2009¶ estimated¶
CC305198- (A)¶	11/2//2009-to- 4/3/2010¶	2009.7211¶	557.9¶	65.8¶	0¶	3a,b¶	1.5¶ After 2010-01-31.¶ 2.5¶	1/1/2009¶ SM*¶

1298 * Gravimetric blends with CO, H₂, CO₂, CH₄ and N₂O in zero air purchases from Scott Marrin.¶

1299 ** Uncertainty estimates listed for the tertiary standard assignments assume a 0.5 ppb uncertainty for each

1300 ealibration result on H9 and do not formally include the uncertainty on the secondary standard

1301 assignments.

- 1302 a. Assignment does not use existing post-use calibration results that show larger drift¶
- 1303 b. Drift change towards end of use, additional drift correction applied.¶
- 1304 c. Force linear fit in drift calculation code¶
- 1305 d. Only predeployment calibrations¶
- 1306 e. No end-of-use or post-use calibration¶
- 1307 f. Force quadratic fit in drift calculation code

P

1308 Table 3: H_2 working standards used on the MAGICC-3 system. Best polynomial curve fit coefficients to 1309 the August 2019-December 2022 calibration histories.

Tank ID- (fill)¶	₩	Assignment at t0 (ppb)¶	C1 (ppb/yr)¶	€2¶ (ppb/y1²)¶	N¶	Assignment- uncertainty (ppb)¶	Fill date ¶
CA01414 (D¶	2020.0964¶	238.4¶	10.0¶	-1.9¶	9¶	0.5 ppb¶	12/29/2017¶ NWR¶
CA04403- (F)¶	2020.1052¶	474.6¶	10.2 ¶	-1.7¶	9¶	0.5 ppb¶	12/1/2017¶ NWR¶
CB11270- (∧)¶	2020.0012 ¶	515.0¶	2.9¶	-0.5¶	9¶	0.5 ppb¶	12/1/2017¶ NWR¶
CA06388- (₩)¶	2019.9423 ¶	551.2¶	1.1¶	₽¶	9¶	0.5 ppb¶	2/23/2018¶ NWR¶
CA05773- (F)¶	2020.2585¶	565.6¶	1.4¶	0¶	8¶	0.5 ppb¶	5/17/2018¶ NWR¶
CB11034- (₽)¶	2020.0783¶	580.1¶	8.3 ¶	-1.2¶	9¶	0.5 ppb¶	5/17/2018¶ NWR¶
CA05680- (H)¶	2020.0904¶	588.1¶	1.9 ¶	0¶	9¶	0.5 ppb¶	12/1/2017¶ NWR¶
CB11405- (C)¶	2020.1474¶	605.6¶	23.3 ¶	-1.6¶	9¶	0.5 ppb¶	5/17/2018¶ NWR¶

Linear Drift Rate (ppb/yr)	Target Tank IDs	Standard deviation of residuals to best fits (ppb)
0	CA05278, CA06194, CA08247, CC121971, CC311842 ND16439, ND33960	0.46
0-1	ALM-065166, CA05300, CC71607, CC73110	0.42
2-5	CA04551, CA07328, CB10910	0.32
5-10	CC71579	0.36
> 20	CA08145	0.48

1314 Table 24: H9 Target air tanks with zero or linear growth in H_2

System / Instrument	Test air tank id and fill	Differences mean (ppb)	Differences standard deviation (ppb)	Number of samples
MAGICC-2 / H8	AL43-113 D, E	-0.3	1.3	528
MAGICC-1 / H11	AL43-113 D, E, G	+0.3	1.1	1231
MAGICC-3 / H8	AL47-145 G	-0.9	1.5	388
MAGICC-3 / H11	AL43-113 G	+0.4	0.6	144

1318 Table 35. Summary statistics for H_2 differences between test air tank-fill assignment (based on H9 1319 calibration history) and associated TST flask measurements on MAGICC systems

1323 Table 6. Summary statistics for SPO flask pair H₂ differences. Npairs= Number of flask pairs.

¶ System/¶ Instrument	SPO "P" flasks ¶ Absolute differences			SPO "S" flasks ¶ Absolute differences			SPO "S"-"P"¶ Pair mean differences		
	Mean- (ppb)	Std dev (ppb)	Npairs	Mean (ppb)	Std dev¶ (ppb)	Npairs	Mean- (ppb)	Std dev¶ (ppb)	Npairs
MAGICC-2- /H8	1.3	1.0	165	1.1	0.9	87	-0.4	1.5	81
MAGICC-1 / 1111	0.9	0.8	292	0.9	0.8	143	-0.2	1.3	1 44
MAGICC-3- /H8	1.6	1.3	45	1.2	1.2	25	-0.1	1.7	25
MAGICC-3- /1111	0.7	0.6	76	0.8	0.6	35	- 0.5	0.8	43

Uncertainty components	1 sigma uncertainty estimate (ppb)	Source
Tertiary standard time-dependent assignment uncertainty (1 point calibration)	0.5-2.5 Tank specific (see SI Table 2)	Calibration histories, residuals to best fit, TST flasks
MAGICC-3 response curve uncertainty	0.5	Preliminary estimate, will be reassessed.
Measurement repeatability on H8	1.3 (MAGICC-2) 1.5 (MAGICC-3)	TST and SPO flask pair differences (Table
Measurement repeatability on H11	1.1 (MAGICC-1) 0.6 (MAGICC-3)	3 and SI Table 6)

1327 Table 4: Flask air H_2 measurement uncertainty components

 Table 57: Annual mean of H₂ measurement differences (in ppb) for air samples from the Cape Grim Observatory (CGO), Ochsenkopf (OXK) and Alert (ALT). Non background air sample measurement results are included. Collocated (not same air) samples at ALT are matched within a +/- 60 minutes window. [updated 9-25-23]

Year	NC ICP-1 nor	NOAACGOOXKICP-NOAANOAA non ICPNOAA ICPnonICPminusminusCSIRO ICPMBL ICP		ALT NOAA minus CSIRO (not same	ALT NOAA minus MPI (not same air)	
	CGO*	OXK	CSIRO ICP	MPI ICP	air)	air)
2010	-	-0.05	0.72	-0.17	-3.4	-3.5
2011	-	0.15	0.50	-0.02	2.2	-3.9
2012	0.58	0.13	0.40	-0.29	0.66	-2.3
2013	-	0.01	0.23	0.80	1.30	-1.4
2014	-	0.19	1.37	1.61	0.63	-1.1
2015	-	0.85	0.02	0.53	0.52	-1.4
2016	1.32	0.20	1.54	2.91	-0.32	-1.4
2017	1.19	0.56	1.38	2.49	3.2	-
2018	0.91	0.53	1.31	1.69	1.2	-1.3
2019	0.73	-0.07	0.30	1.25	1.0	-0.81
2020	0.18	na	0.19	-	0.01	-0.22
2021	0.33	0.33	0.86	1.71	3.4	-

1334 *Most NOAA ICP flasks from CGO had a small contamination for CO and H₂ prior to 2019. If the 1335 NOAA ICP flask H₂ results are > 2ppb larger than the NOAA non-ICP flask H₂ in the pair, the ICP flask 1336 H₂ has been rejected. Only years with at least 10 valid H₂ pairs are included.

1337 Table 78: Flask air H2 measurement uncertainty components

Uncertainty- components	1 sigma uncertainty- estimate (ppb)	Source
Tertiary standard- time-dependent- assignment uncertainty- (1 point calibration)	0.5-2.5 ¶ Tank specific (see Table 2)	Calibration histories, residuals to best fit, TST flasks
MAGICC-3 response eurve uncertainty	0.5	Preliminary estimate, will be reassessed.
Measurement- repeatability on H8	1.3 (MAGICC-2)¶ 1.5 (MAGICC-3)	TST and SPO flask- pair differences (Tables-
Measurement- repeatability on H11	1.1 (MAGICC-1)¶ 0.6 (MAGICC-3)	5 and 6)

1339 Figures

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1341 Figure 1. Calibration results for GML two H_2 secondary standards (a) CC119811 and b) CA03233) on H9 1342 against one of the primary standards. 2019-2020 multipoint calibration results on H9 are also shown for 1343 CA03233 (pink circles). Only results shown with open circles are used for the assignments.



1347 Figure 2: 2019-2022 H9 standard calibration response curve (RC) results: a) differences from the mean
1348 RC linear fit mean and b) residuals of the response curve fits. Different colors are for different calibration
1349 episodes.



1358 Figure 3. Calibration histories of a) MAGICC-1 / H11 and b) MAGICC-2 / H8 tertiary standards. The1359 colored vertical line indicates when a standard started to be used.1360





1369 Figure 4: Calibration histories and residuals to best fit for H9 target tanks with a stable H_2 mole fraction 1370 or a linear drift less than 1 ppb/yr. Residuals are in ppb.

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1390 Figure 5. Flask air analysis systems (H8 and H11)-target air tanks H9 a) calibration histories and b) 1391 residuals to best linear or quadratic fit.

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1417 time-dependent assignment based on calibration history on H9. The vertical line indicates the transition to 1418 the MAGICC-3 flask analysis system.¶





1425 Figure 7. South Pole Observatory flask air H₂ measurements on H11 and H8. Black symbols are used for 1426 measurements of P flasks and blue symbols are used for measurements of S flasks.

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1429 Figure 7. South Pole Observatory flask air H₂ measurements. Circle and "+" symbols refer to instruments: 1430 H11 or H8. Black is used for measurements of P flasks on the MAGICC-1 or MAGICC-2 system and red-1431 for the MAGICC-3 system. Light green is used for measurements of S flasks on the MAGICC-1 or 1432 MAGICC-2 system and light blue for the MAGICC-3 system. 1433 1434



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1444 Figure 910: Annual maximum (red), mean (green) and minimum (blue) H_2 from the smooth curve fit of 1445 the 2010-2021 measurement time series for each surface site in the global sampling network. Each site is 1446 referred to with a three letter code (see details in SI Table 7at <u>https://gml.noaa.gov/dv/site/</u>). The sampling 1447 sites are shown along the x-axis with decreasing latitudes. An asterisk near the site code indicates if the 1448 site data areis used for the marine boundary layer air zonal and global means H_2 data reduction. 1449

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1460 Figure 11: 2010-2021 marine boundary layer H_2 meridional gradient. Y-axis is the sine of latitude.

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1467 Figure 12: 2010-2021 marine boundary layer global mean and zonal mean H_2 anomaly (black line) and 1468 CO anomaly (dashed blue line) time series.