Atmospheric propane (C_3H_8) column retrievals from ground-based FTIR observations at Xianghe, China

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Abstract. Propane (C_3H_8) is an important trace gas in the atmosphere, as it is a proxy for oil and gas production and has a significant impact on atmospheric chemical reactions related to the hydroxyl radical and tropospheric ozone formation. In this study, solar direct absorption spectra near 2967 cm^{-1} recorded by a ground-based Fourier Transform InfraRed spectrometer (FTIR) are applied to retrieve C_3H_8 total columns between June 2018 and July 2022 at Xianghe in North China. The systematic and random uncertainties of the C_3H_8 column retrieval are estimated to be 18.4% and 18.1%, respectively. The mean and standard deviation of the C_3H_8 columns derived from the FTIR spectra at Xianghe are $1.80\pm0.81(1\sigma) \times 10^{15} \ molecules/cm^2$. Good correlations are found between C_3H_8 and other non-methane hydrocarbons, such as C_2H_6 (R=0.84) and C_2H_2 (R=0.79), as well as between C_3H_8 and CO (R=0.72). However, the correlation between C_3H_8 and CH_4 is relatively weak (R=0.45). Moreover, the FTIR C_3H_8 measurements at Xianghe are also compared against MKIV measurements at several sites around the world. The new FTIR measurements at Xianghe provide us an insight into the C_3H_8 column variations and underlying processes in North China.

1 Introduction

Methane (CH_4) and non-methane hydrocarbons (NMHC), such as ethane (C_2H_6), acetylene (C_2H_2), propane (C_3H_8), propene (C_3H_6), and isoprene (C_5H_8), are important trace gases that play significant roles in atmospheric chemical reactions related to hydroxyl radical (OH) abundance and tropospheric ozone (O_3) formation (Sze, 1977; Donahue and Prinn, 1990; Tan et al., 2012; Lelieveld et al., 2015). Human activities contribute greatly to the emissions of CH_4 and NMHCs, especially in urban areas (Bourtsoukidis et al., 2019; Saunois et al., 2020). Atmospheric C_2H_6 and C_3H_8 emissions are dominated by oil and gas sources, and they are co-emitted with CH_4. Therefore, numerous studies used the ratio of C_2H_6 and/or C_3H_8 to CH_4 to understand the CH_4 trend (Kort et al., 2016; Franco et al., 2016; Rigby et al., 2017).

The major sink of C_2H_6 and C_3H_8 is the reaction with OH, and the lifetime of C_3H_8 and C_2H_6 is about 2-4 weeks in summer and 2 months in winter (Jacob, 1999; Xiao et al., 2008). Compared to CH_4 with a lifetime on an order of 10 years (IPCC, 2013),

the short-lived gases C_2H_6 and C_3H_8 are not well-mixed on the global scale, and are therefore more representative of regional pollution as is carbon monoxide (CO) (Toon et al., 2021).

Atmospheric C_3H_8 concentrations at the surface are observed by National Oceanic and Atmospheric Administration (NOAA) - Global Monitoring Laboratory (GML) flask sampling measurements at 12 sites (https://gml.noaa.gov/hats/gases/C3H8.html). In addition, the HIAPER Pole-to-Pole Observations (HIPPO), Atmospheric Tomography (ATom), and In-service Aircraft for a Global Observing System (IAGOS) aircraft campaigns provide in-situ gas analyzer measurements of C_3H_8 with a wide latitudinal coverage, particularly in the Pacific Ocean, Atlantic Ocean, Europe and North America (Wofsy, 2011; Thompson et al., 2022; Li et al., 2022). Toon et al. (2021) has demonstrated the use of C_3H_8 absorption lines in the mid-infrared region (Harrison et al., 2010), in solar absorption spectra from MkIV interferometers for retrieving the C_3H_8 total columns or vertical profiles at several locations in Sweden, the USA, and Antarctica. Solar absorption infrared spectra are also being collected by ground-based Fourier Transform Spectrometers (FTIR) within the Network for the Detection of Atmospheric Composition Change - InfraRed Working Group (NDACC-IRWG) (De Mazière et al., 2018). Currently, there are more than 20 NDACC-IRWG global sites, with a good global latitudinal coverage from 78°S to 80°N (https://www2.acom.ucar.edu/irwg/sites). However, to our knowledge, no site has reported C_3H_8 retrievals from spectra observed by a Bruker 125HR spectrometer within the NDACC-IRWG.

Xianghe (39.75 °N, 116.96 °E) is located in North China, about 50 km east of the mega-city Beijing (Yang et al., 2020). According to the Emissions Database for Global Atmospheric Research (EDGAR) v6.0 (Crippa et al., 2020) and the Multi-resolution Emission Inventory for China (MEIC) inventory (Wang et al., 2015; Li et al., 2017), there is a large CH_4 emission source in North China coming from fuel exploitation and oil refineries. Therefore, we expect that the C_2H_6 and C_3H_8 concentrations are relatively high in this region. In June 2018, a Bruker IFS 125HR spectrometer, compliant with the NDACC-IRWG protocol, started recording solar absorption spectra in the mid-infrared spectral range. The spectra have been used to retrieve several atmospheric components, e.g., O_3, CH_4, CO, C_2H_2, C_2H_6, HCN and H_2CO (Ji et al., 2020; Zhou et al., 2020, 2021, 2023; Vigouroux et al., 2020; Sha et al., 2021). In this study, we investigate the C_3H_8 retrieval from ground-based FTIR spectra at Xianghe, and discuss the C_3H_8 column variation in North China, based on these new FTIR measurements.

The remainder of this paper is organized as follows. Section 2 describes the Xianghe FTIR site and C_3H_8 retrieval method, Section 3 presents the C_3H_8 variations and correlations with other species. Moreover, the C_3H_8 measurements at Xianghe are compared to ground-based MkIV measurements at other places. Finally, Section 4 draws a conclusion.

2 Method

2.1 Xianghe FTIR spectra measurement

The Xianghe FTIR measurement system started in June 2018, and has been well described in previous studies (Yang et al., 2020; Zhou et al., 2021, 2023). Briefly, the FTIR measurement system contains 3 parts: a solar tracker system, a weather station, and a Bruker IFS 125HR Fourier-transform infrared (FTIR) spectrometer. Short-Wave infrared (SWIR) and Near-infrared (NIR) spectra (4000-11000 cm^{-1}) with a spectral resolution of 0.02 cm^{-1} are recorded with an InGaAs detector,

and these spectra are used to derive greenhouse gases total column abundances as a contribution to the Total Carbon Column Observing Network (TCCON). Mid-infrared (MIR) spectra (1800-4500 cm^{-1}), with a spectral resolution of 0.0035-0.0070 cm^{-1}, are recorded with an InSb detector. To enhance the signal-to-noise ratio (SNR) of the spectra, we add specific optical filters into the light path when recording each MIR spectrum as recommended by NDACC-IRWG (Blumenstock et al., 2021; Zhou et al., 2023). A typical MIR spectrum, with a spectral resolution of 0.0051 cm^{-1}, used for C$_3$H$_8$ retrieval is shown in Figure 1. Note that, we only operate the FTIR measurement during the daytime and under clear-sky conditions, as the sun is the light source. In general, we carry out 4 to 10 MIR spectral measurements of this type per day for about 200 days per year. Each spectrum takes about 10 minutes to record. The spectra taken between June 2018 and July 2022 (about 4 years) are used in this study.
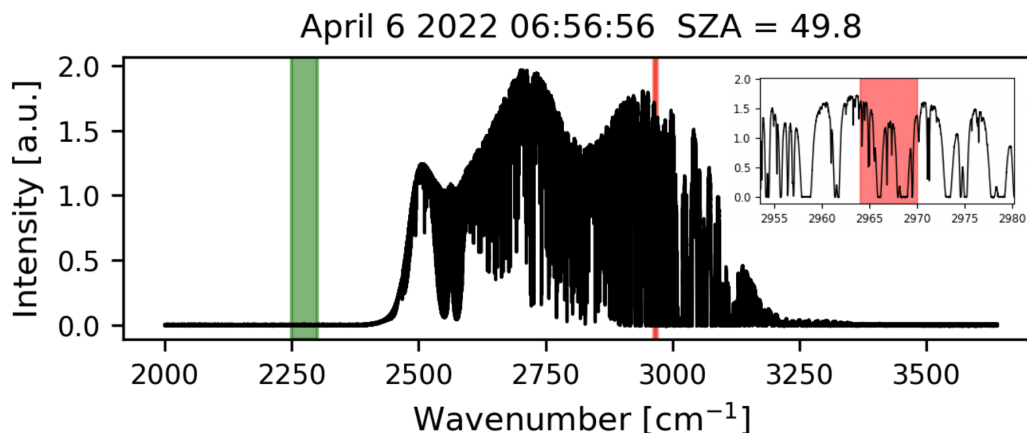
Figure 1. A typical MIR spectrum observed at Xianghe on 6 April 2022 with a solar zenith angle of 49.8°. The red and green windows indicate the micro-windows windows used for the C$_3$H$_8$ retrieval and for calculating the noise (Eq. 5), respectively. The insert in the right-hand corner shows a zoom on the retrieval micro-window.

2.2 Retrieval method

To derive C$_3$H$_8$ mole fractions from the observed spectra, we follow the optimal estimation methodology (Rodgers, 2000). The forward model ($\mathbf{F}$) simulates the absorption spectra ($\boldsymbol{y}$) observed by the FTIR system. It includes modelling of the solar spectra at the top of the atmosphere (TOA), the physics of the radiative transfer from the TOA to the ground-based FTIR, and the FTIR spectrometer line shape function (ILS). Then, the observed spectra ($\boldsymbol{y}$) can be written as

$$\boldsymbol{y} = \boldsymbol{F}(\boldsymbol{x}, \boldsymbol{b}) + \boldsymbol{\epsilon}, \tag{1}$$

where $\boldsymbol{x}$ is the state vector (retrieved parameters), $\boldsymbol{b}$ is the forward model parameters (not retrieved), and $\boldsymbol{\epsilon}$ is the error, including the measurement noise and forward model errors. We wish to find the optimal state ($\boldsymbol{x}$) that minimize the cost function ($\boldsymbol{J}(\boldsymbol{x})$),

given by

$$J(\boldsymbol{x}) = [\boldsymbol{y} - \boldsymbol{F}(\boldsymbol{x})]^T \mathbf{S}_{\boldsymbol{\epsilon}}^{-1} [\boldsymbol{y} - \boldsymbol{F}(\boldsymbol{x})] + [\boldsymbol{x} - \boldsymbol{x_a}]^T \mathbf{S}_R [\boldsymbol{x} - \boldsymbol{x_a}], \quad (2)$$

where $\mathbf{S}_{\boldsymbol{\epsilon}}$ is the measurement error covariance matrix; $\mathbf{S}_R$ is the regularization matrix; $\boldsymbol{x_a}$ is the *a priori* state vector. The Levenberg-Marquardt (LM) method is used to iteratively solve the above equation:

$$\boldsymbol{x}_{i+1} = \boldsymbol{x}_i + [(1+\gamma)\mathbf{S}_R + \mathbf{K}_i^T \mathbf{S}_{\boldsymbol{\epsilon}}^{-1} \mathbf{K}_i]^{-1} \left\{ \mathbf{K}_i^T \mathbf{S}_{\boldsymbol{\epsilon}}^{-1} [\boldsymbol{y} - \boldsymbol{F}(\boldsymbol{x_i})] - \mathbf{S}_R [\boldsymbol{x}_i - \boldsymbol{x}_a] \right\}, \quad (3)$$

where $\mathbf{K}$ is the Jacobian matrix, γ is a parameter to adjust the regularization of a priori information in each iteration step (Rodgers, 2000). Upon convergence, the final state is called $\boldsymbol{x}_r$, which can be related to the true state ($\boldsymbol{x}_t$):

$$\boldsymbol{x}_r = \boldsymbol{x}_a + \mathbf{A}(\boldsymbol{x}_t - \boldsymbol{x}_a) + \varepsilon, \quad (4)$$

where $\mathbf{A}$ is the averaging kernel matrix, representing the sensitivity of the retrieved parameters to the true parameter, and ε is the retrieval uncertainty propagated from Eq.1.

2.3 Retrieval strategy

In this study, we use the SFIT4 v1.0 retrieval algorithm (Pougatchev et al., 1995; Hase et al., 2004) to perform the forward model simulation as well as the LM inversion. The well-established SFIT4 code has been used extensively to retrieve total/partial column of atmospheric species in the NDACC-IRWG community (Zhou et al., 2016; De Mazière et al., 2018; Ortega et al., 2019).

The key C_3H_8 retrieval parameters used in this study are listed in Table 1. The retrieval window is set to 2964.5-2970.0 cm^{-1}, where we have the strongest C_3H_8 absorption line (Harrison et al., 2010). Apart from C_3H_8, several interfering gases (H_2O, CH_4, O_3, C_2H_6 and HDO) also have absorption lines in this window as shown in Figure 2. To reduce the impact of uncertainties about the abundances of these species, CH_4, O_3, C_2H_6, and HDO columns are retrieved along with the target gas mole fractions. For these three species, their profile shapes are fixed and only the scaling factors are retrieved. As H_2O absorption lines are strong (Table 1) and H_2O variability are relatively large in the atmosphere, we perform a profile retrieval for H_2O. The state vector includes CH_4, O_3, C_2H_6, and HDO columns, as well as 47-layers' C_3H_8 and H_2O mole fractions.

The chosen spectroscopic parameters are crucial in the remote sensing technique. In this study, we have tested several line lists, particularly for H_2O (HDO) and CH_4 (see Table 2), including DLR2016 (Loos et al., 2017), HITRAN2020 (Gordon et al., 2022) and ATM2020 (https://mark4sun.jpl.nasa.gov/pseudo.html). The ATM2020 line list is created by Geoff Toon (NASA, JPL) based on HITRAN2020 together with some additional atmospheric and laboratory measurements. It includes pseudo linelists (PLL) for certain species as the ones we use for C_3H_8, based on laboratory cross section measurements by Harrison et al. (2010). For C_2H_6, we use HITRAN2020. We tested more than 1000 spectra recorded in 2019 at Xianghe, and we observed that the lowest root-mean-square error (RMSE) of the fitting residual is obtained when the ATM2020 spectral database is used for CH_4 and H_2O. Table 1 lists the spectral datasets finally used for each species in the C_3H_8 retrieval strategy.

Table 1. The retrieval window, interfering species, spectroscopy, fitting parameters for C_3H_8 at Xianghe.

Parameters	settings
Retrieval window (cm^{-1})	2964.5-2970.0
Profile retrieval species	C_3H_8, H_2O
Column retrieval species	C_2H_6, CH_4, HDO
Retrieved parameters	slope, phase, instrument line shape, wavenumber shift
	solar intensity, solar wavenumber shift
A priori profile	NCEP for H_2O, HDO; WACCM for C_2H_6, C_3H_8, CH_4
Spectroscopy	PLL for C_3H_8; ATM20 for H_2O, HDO, CH_4; HITRAN2020 for C_2H_6
Regularization	Tikhonov $\mathbf{L}_1$ method
DOFS	1.1
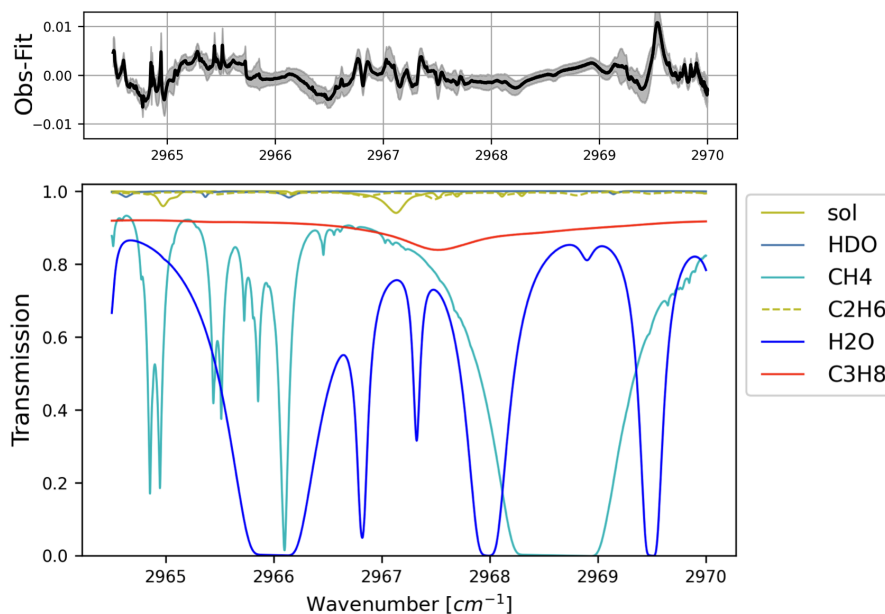
Figure 2. The transmittances of main species and solar lines (bottom), as well as the fitting residual (top) mean (black line) and standard deviation (grey shadow) from all 2783 FTIR C_3H_8 retrievals at Xianghe between June 2018 and July 2022. The mean RMSE is 0.317%.

The *a priori* profiles for C_3H_8, C_2H_6, and CH_4 are derived from the Whole Atmosphere Community Climate Model (WACCM) version 6. We use the averages of the monthly means between 1980 and 2040 (61 years) as the *a priori* profiles. Since the variations of temperature and humidity are quite large in the atmosphere, using fixed *a priori* profiles often results in a bad fitting, especially for the first iteration. To provide a better estimation of temperature and humidity profiles, for

Table 2. The fitting RMSE of the retrieval window for all spectra in 2019 from several different line lists.

H$_2$O (HDO)	CH$_4$	RMSE (mean$\pm1\sigma$)
ATM2020	ATM2020	0.313$\pm$0.081
HITRAN2020	ATM2020	0.327$\pm$0.091
DLR	ATM2020	0.328$\pm$0.091
ATM2020	HITRAN2020	0.417$\pm$0.095
HITRAN2020	HITRAN2020	0.445$\pm$0.097

each measurement, the H$_2$O (HDO) and temperature vertical profiles are derived from the closest 6-hourly NCEP reanalysis data (Saha et al., 2014), and linearly interpolated to the measurement time.

According to Eq.2, the cost function $J(x)$ is composed of the measurement and *a priori* information, each contracted with a weight matrix $\mathbf{S}_\epsilon$ and $\mathbf{S}_R$, respectively. In this study, the diagonal of the $\mathbf{S}_\epsilon$ is calculated as $1/\text{SNR}^2$, and the non-diagonal values are set to 0. The SNR is calculated as

$$\text{SNR} = \frac{\overline{I_r}}{\sigma_{I_n}}, \tag{5}$$

where $\overline{I_r}$ is the max radiation intensity in the C$_3$H$_8$ retrieval window (2964.5-2970.0 cm^{-1}; red window in Figure 1) and σ_{I_n} is the standard deviation (std) of the intensity in the noise window (2250.0-2300.0 cm^{-1}; green window in Figure 1). The Tikhonov $\mathbf{L}_1$ regularization method (Tikhonov, 1963) is applied to generate the $\mathbf{S}_R$, with

$$\mathbf{S}_R = \alpha \mathbf{L}_1^T \mathbf{L}_1, \tag{6}$$

$$L_1 = \begin{bmatrix} -1 & 1 & 0 & \dots & 0 & 0 \\ 0 & -1 & 1 & \dots & 0 & 0 \\ \vdots & \vdots & \vdots & \ddots & \vdots & \vdots \\ 0 & 0 & 0 & \dots & -1 & 1 \end{bmatrix}. \tag{7}$$

To determine the α value in Eq. 6, we apply the degree of freedom for signal (DOF) method proposed by Steck (2002). The trace of the averaging kernel matrix ($\mathbf{A}$) is the DOF, indicating the pieces of independent information of the retrieval (Rodgers, 2000). First, we use the optimal estimation method (OEM) to get an estimated DOF. Using the OEM method, $\mathbf{S}_R = \mathbf{S}_a^{-1}$, where $\mathbf{S}_a$ is a priori covariance matrix, which is derived from a covariance matrix on the WACCM monthly means between 1980 and 2040 (($S_R)_{i,i} = (S_a)_{i,i}^{-1} = \sigma_i^{-2}$; diagonal values), and the non-diagonal values are set as $(S_R)_{i,j} = (S_a)_{i,j}^{-1} = (\sigma_i \sigma_j)/e^{(d_{i,j}/4)}$, where $d_{i,j}$ is the vertical distance between layer i and layer j, in km. The DOF derived from the OEM is about 1.1, indicating that there is only column information for the C$_3$H$_8$ retrieval. Knowing that, we tune the α value in Eq.6 to make the DOF derived from the Tikhonov method close to the DOF that is derived from the OEM; this approach results in setting α to 1000.

2.4 Retrieval uncertainty

The retrieval error (ϵ) of the FTIR C_3H_8 column contains three parts as

$$(\mathbf{A} - \mathbf{I})(\boldsymbol{x}_t - \boldsymbol{x}_a) \quad ... \quad smoothing\ error \tag{8}$$

$$\mathbf{G}_y \mathbf{K}_b (\boldsymbol{b}_t - \boldsymbol{b}_a) \quad ... \quad model\ parameter\ error \tag{9}$$

5 $\quad \mathbf{G}_y \epsilon \quad ... \quad measurement\ error \tag{10}$

where $\mathbf{G}_y$ is the contribution function; $\boldsymbol{b}_t$ and $\boldsymbol{b}$ are the true and used model inputs, respectively. Table 3 lists the systematic and random uncertainty of each component. The vertical distributions of the systematic and random uncertainties are shown in Figure 3. For the smoothing error, we separate the contributions into target species (C_3H_8), interfering species (H_2O, HDO, CH_4, C_2H_6), and retrieved parameters (slope, phase, wavenumber shift, instrument line shape, solar intensity and shift). For the

10 model parameter contributions, we calculate the C_3H_8 uncertainty contribution coming from spectroscopy, solar zenith angle (SZA), temperature profile, curvature parameter, and zero level shift (zshift). Since CH_4 and H_2O have stronger absorptions than C_3H_8, and their absorption lines are not perfectly fitted, the impact from the spectroscopy uncertainty of CH_4 and H_2O are calculated as well.

The systematic and random uncertainties of each parameter are also listed in Table 3. It is assumed that 10% of the a priori

15 profile is used to derive the diagonal values of the systematic covariance matrix $(\mathbf{S}_a)_{i,i}^{sys} = \sigma_i^2$, and the off-diagonal values of $\mathbf{S}_a^{sys}$ are calculated as $(\mathbf{S}_a)_{i,j}^{sys} = \sigma_i \sigma_j$ (von Clarmann, 2014). The covariance matrix derived from the WACCM 61-years' monthly means are set to the random covariance matrix $\mathbf{S}_a^{ran}$. Regarding the model parameter uncertainties in the Table 3, the systematic/random $\mathbf{S}_b$ matrix is created by the mean and standard deviation of the differences between NCEP and ERA5 at Xianghe. The random deviation is about 2 K, and the systematic deviation is about 1.5 K for the whole vertical range.

20 For the target spectroscopic parameters, the relative uncertainties of C_3H_8 is set to 4% according to the pseudo database. For the CH_4 and H_2O spectroscopy parameters, the relative uncertainty of 5% is derived from the HITRAN2020 dataset (Gordon et al., 2022). Note that the spectroscopy uncertainty in Table 3 is the sum of the uncertainties from the line intensity, pressure dependent parameter (linePAir) and temperature dependent parameter (lineTAir). For the uncertainties of background curvature, zero offest (shift), and SZA, we use the default values provided by the SFIT4 algorithm (https://wiki.ucar.edu/

25 display/sfit4/SFIT4+Version+1.0.xx+Release), which are recommended by the NDACC-IRWG community.

Based on our uncertainty estimation, the total systematic and random uncertainty of the C_3H_8 column are both about 18%, and the dominating contribution is the uncertainty on the background curvature parameter in the forward model. To represent the variability of the C_3H_8, we select all days with at least 3 individual measurements on each day, and calculate the daily std. The average of all the daily stds is about 15.3%, and it is close to our estimated random uncertainty.

Table 3. The systematic and random (sys/ran) retrieval uncertainties for the total columns of C_3H_8. The '-' means that the uncertainty is less than 0.1%. 1 σ of the target or interfering species is the std derived from the WACCM model monthly means between 1980 and 2040. The relative std in the bottom row is the average of daily std of C_3H_8 columns on all days with at least 3 measurements, which is to represent the variability of the retrieval.

Error source	Parameter	Parameter uncertainty (sys/ran)	C_3H_8 column uncertainty [%]
Smoothing error	Target species (C_3H_8)	10/1σ %	0.2/0.5
	Interfering species (H_2O, CH_4, HDO, C_2H_6)	10/1σ %	0.7/0.6
	Retrieved parameters		0.6/0.6
Model parameter error	Spectroscopy for C_3H_8	4.0/- %	4.1/-
	Spectroscopy for H_2O and CH_4	5.0/- %	2.5/-
	SZA	0.03/0.03°	0.1/0.1
	Curvature	0.1/0.1 %	17.2/17.2
	Temperature	1.5/2.0 K	2.7/3.9
	Zshift	0.15/0.15 %	2.9/2.9
Measurement error		-/$\frac{1}{SNR}$	-/1.0
Total			18.4/18.1
Std			-/15.3
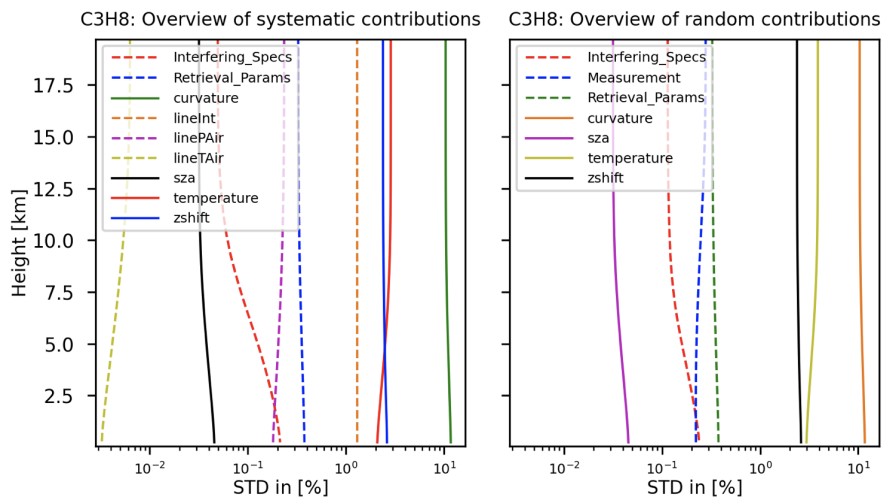
Figure 3. The vertical profiles of the systematic (left) and random error (right) of the FTIR C_3H_8 retrieval from each component.

3 Results and discussions

3.1 FTIR C_3H_8 retrievals at Xianghe

Figure 4 shows the *a priori* profile and retrieved profiles of C_3H_8. The vertical profile of C_3H_8 from the WACCM model shows that the C_3H_8 mole fraction is high near the surface and decreases with increasing altitude. Such a vertical shape is expected as the C_3H_8 emissions are at the surface, and its atmospheric lifetime is too short to achieve a well-mixed troposphere. Although we perform a profile retrieval on C_3H_8, we only have about 1 DOF. In addition, the Tikhonov regularization matrix constrains the vertical shape when the DOF is typically close to 1.0. As a result, the retrieved C_3H_8 profiles have a very similar vertical shape as the *a priori* profile. However, the FTIR measurements show that the a priori column overestimates the C_3H_8 column concentration by about 100%. The column averaging kernel indicates the sensitivity of the retrieved C_3H_8 column to the C_3H_8 partial column in each height. Figure 4 shows that the retrieved C_3H_8 column has good sensitivity to all the layers, and slightly varies with SZA.
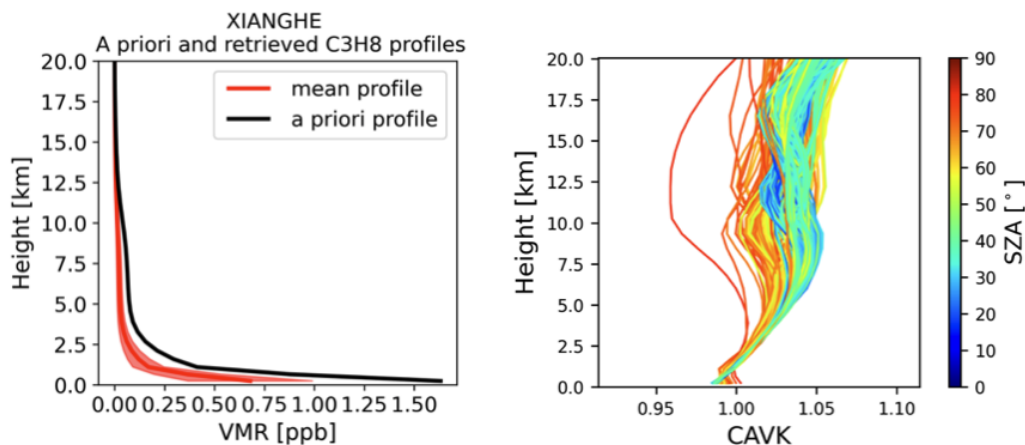
Figure 4. The *a priori* and retrieved C_3H_8 profiles (left), and the column averaging kernel (CAVK) varying with SZA (right).

The time series and seasonal variation of FTIR C_3H_8 column measurements are presented in Figure 5. To better visualize the seasonal variation, the column measurements are fitted by a periodic function $y(t) = A_0 + \sum_{k=1}^{3} (A_{2k-1}\cos(2k\pi t) + A_{2k}\sin(2k\pi t))$, where A_0 is the offset, and A_1 to A_6 are the periodic amplitudes, representing the seasonal variation. The obtained mean and std of C_3H_8 columns at Xianghe are $1.80\pm0.81 \times 10^{15}$ $molec./cm^2$. The C_3H_8 columns show a high mean value in July and a low value in October. The difference between the median values in July (maximum) and October (minimum) is 1.2×10^{15} $molec./cm^2$. Although the median values of C_3H_8 columns in June-August are larger than those in October-March, we notice that extremely high C_3H_8 columns often occur in the latter period.
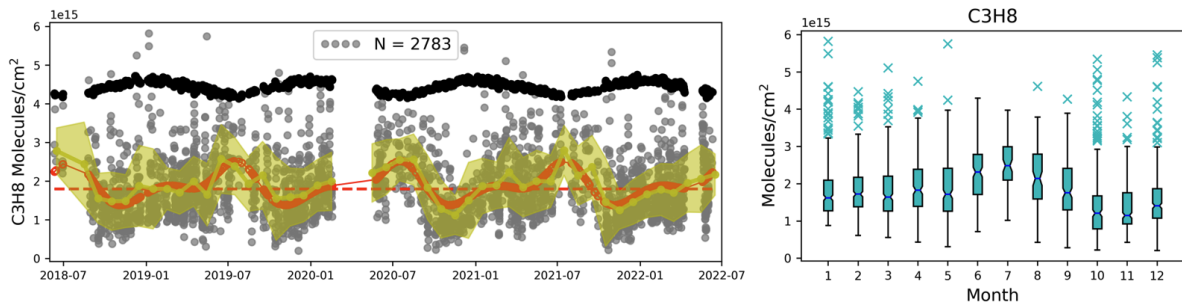
Figure 5. Left panel shows the time series of FTIR a priori C_3H_8 columns (black dots), retrieved columns (grey dots), monthly means (yellow line), monthly stds (yellow shade), periodic function fitting (red solid line) and the fitted offset (red dashed line). Right panel: the monthly box plot of the C_3H_8 columns. The bottom and top bars represent the 10% and 90% percentiles of the datasets and the blue crosses are the extremely high values above 90%.

3.2 Correlations with CO, CH_4, C_2H_2 and C_2H_6 at Xianghe

As mentioned above, the infrared spectra observed by the Xianghe FTIR system have been also used to retrieve CO, CH_4, C_2H_2 and C_2H_6 columns using NDACC-IRWG recommended retrieval recipes (Ji et al., 2020; Zhou et al., 2023), which allows us to investigate the correlation between C_3H_8 and these species. We are particularly interested in the correlation on a regional

5 scale. Therefore, to reduce the impact from the background, we calculate the Δgas (Δgas = gas - monthly median) for all these species. Figure 6 shows the correlation scatter plots between ΔC_3H_8 and ΔCH_4, ΔCO, ΔC_2H_2, and ΔC_2H_6. High correlation coefficients (R) are found between ΔC_3H_8 and ΔC_2H_6 (R=0.84), and between ΔC_3H_8 and ΔC_2H_2 (R=0.79). It indicates that the C_2H_2, C_2H_6 and C_3H_8 (NMHCs) are co-emitted in this region. The slope of ΔC_2H_6 and ΔC_3H_8 is 6.03$\pm$0.03, which suggests a corresponding mixing ratio of C_2H_6 and C_3H_8 mole fractions during the production in North China. CO, as a

10 pollutant tracer, also has a good correlation with C_3H_8 (R=0.72). According to the MEIC inventory, both CO and NMHC are emitted from the energy production, industry, residential and transport sectors.

    The FTIR measurements show that the correlation between ΔC_3H_8 and ΔCH_4 is relatively weak (R=0.45). Note that the variation of the CH_4 column is also affected by the stratospheric partial column (Sepúlveda et al., 2014). The DOF of the FTIR CH_4 retrieval is about 2.5 allowing us to derive the tropospheric and stratospheric CH_4 partial columns separately (Zhou et al.,

15 2018). However, even after eliminating the interference from the stratosphere, the tropospheric CH_4 partial column still has a weak correlation with C_3H_8 (R=0.43). It is probably due to the fact that the CH_4 major emissions in North China are from rice cultivation, waste, and animals instead of the oil and gas production (Ji et al., 2020), and the CH_4 measurements include the emissions from much farther away as compared to the C_3H_8 measurements because of its long lifetime (Callewaert et al., 2023).

20   To further investigate the ratio of ΔC_2H_6 to ΔC_3H_8, the time series of their ratios, together with the monthly correlation coefficients between both time series between June 2018 and June 2022 are illustrated in Figure 7. The ratio of each month is

derived from the linear fitting using all co-located ΔC_2H_6 and ΔC_3H_8 hourly measurements in that month. A relatively low correlation between these two species is found in summer as compared to other three seasons. The mean and std of the ratios are 5.4±2.1 for the whole period. The ratio is lowest in summer and highest in winter, with seasonal means of 6.6, 3.8, 5.4, and 8.3 in spring, summer, autumn, and winter, respectively.
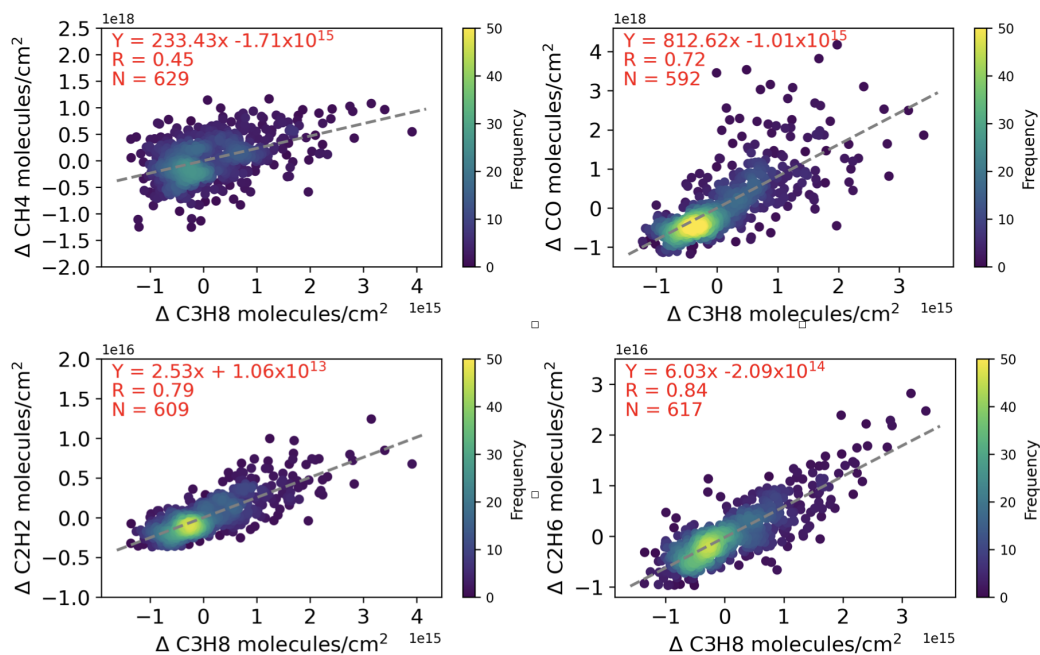
Figure 6. The correlation plots between co-located ΔC_3H_8 and ΔCH_4, ΔCO, ΔC_2H_2 and ΔC_2H_6 hourly means at Xianghe between June 2018 and July 2022. The grey dashed line is the linear fit, N is the number of the FTIR measurements, R is the Pearson correlation coefficient.

3.3   FTIR measurements at Xianghe against MkIV measurements

Here, the C_3H_8 and C_2H_6 columns derived from the FTIR measurements at Xianghe are compared to the ground-based MKIV C_3H_8 retrievals at 6 sites in Sweden and the USA (Figure 8). MKIV data uses the GFIT inverse retrieval code to derive the C_3H_8 columns from the MKIV observed spectra between 2964.5 and 2970 cm^{-1} with a specral resolution of 0.5 cm^{-1}. The mean uncertainties of the MKIV retrieved C_3H_8 and C_2H_6 column are estimated to be around 8×10^{15} $molec./cm^2$ and 7×10^{14} $molec./cm^2$, respectively (Toon et al., 2021). Note that, the C_3H_8 and C_2H_6 retrievals from the MkIV spectrometers at 12 sites have been discussed in Toon et al. (2021), and we only select 6 sites as the measurements are very limited at other 6 sites. The locations and measurement time coverages of sites used in this study are listed in Table 4.

Figure 8 shows that the C_2H_6 column is the largest at Xianghe, apart from several extremely high values at JPL-B and FTS. The seasonal variations of C_2H_6 columns are similar at these sites, especially for JPL-B, MTB and Xianghe, with a high value in northern spring and a low value in northern autumn. Note that, it is hard to derive the seasonal variation of C_2H_6 columns
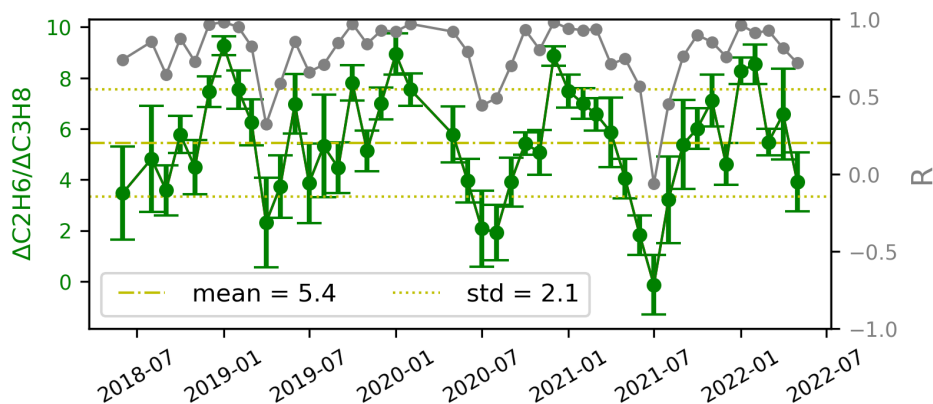
Figure 7. The time series of the ratio of ΔC_2H_6 to ΔC_3H_8 monthly means and stds (green, on the left-hand vertical axis scale), together with their monthly correlation coefficients (grey, on the right-hand vertical axis scale) between June 2018 and June 2022.

at ESN, FAI, TMF and FTS, because measurements were carried out in several months. The mean and std of C_2H_6 columns at JPL-B are $1.96 \pm 0.52 \times 10^{16}$ $molec./cm^2$, which is about 25% less than that at Xianghe ((Xianghe-JPL)/Xianghe $\times 100\%$). Keep in mind that the C_3H_8 columns at MTB and Xianghe have been multiplied by 10 in Figure 8. The C_3H_8 column at Xianghe is quite low as compared to other sites, which is only larger than that at MTB (mountain site), but much less than those at the mid-latitude sites. The mean and std of C_3H_8 columns at JPL-B are $2.14 \pm 1.33 \times 10^{16}$ $molec./cm^2$, which is about 12 times larger than that at Xianghe. The seasonal variations of C_3H_8 columns are similar at JPL-B and Xianghe too, with a high value in northern summer and a low value in northern winter. The good correlations (R>0.6) between C_3H_8 and C_2H_6 columns at JPL-B and FTS have been demonstrated in Toon et al. (2021), which is similar to what we observe at Xianghe. However, the ratio of ΔC_2H_6 to ΔC_3H_8 at JPL-B and FTS are 0.16 ± 0.10 and 0.78 ± 0.10, respectively, which are much less than the ratio observed at Xianghe of 6.03 ± 0.03. It indicates that the emission of C_3H_8 is much larger in the Los Angeles basin, California than that in North China.

4   Conclusions

The Xianghe FTIR 125HR system measures the solar absorption spectra following the NDACC-IRWG guidance. For the first time, the FTIR MIR spectra at Xianghe are used for the C_3H_8 column retrieval, using the well-established SFIT4 code, between June 2018 and July 2022. In this study, the retrieval strategy, retrieval uncertainty, and retrieval information are presented and discussed. Due to the wide and weak absorption of C_3H_8, we only derive the C_3H_8 column instead of its vertical profile. The systematic and random uncertainties of the C_3H_8 retrieved column are estimated to be 18.4% and 18.1%, respectively. In the C_3H_8 retrieval window, CH_4 and H_2O absorption lines are not perfectly fitted, indicating there is still room left to improving the line lists of these two species.
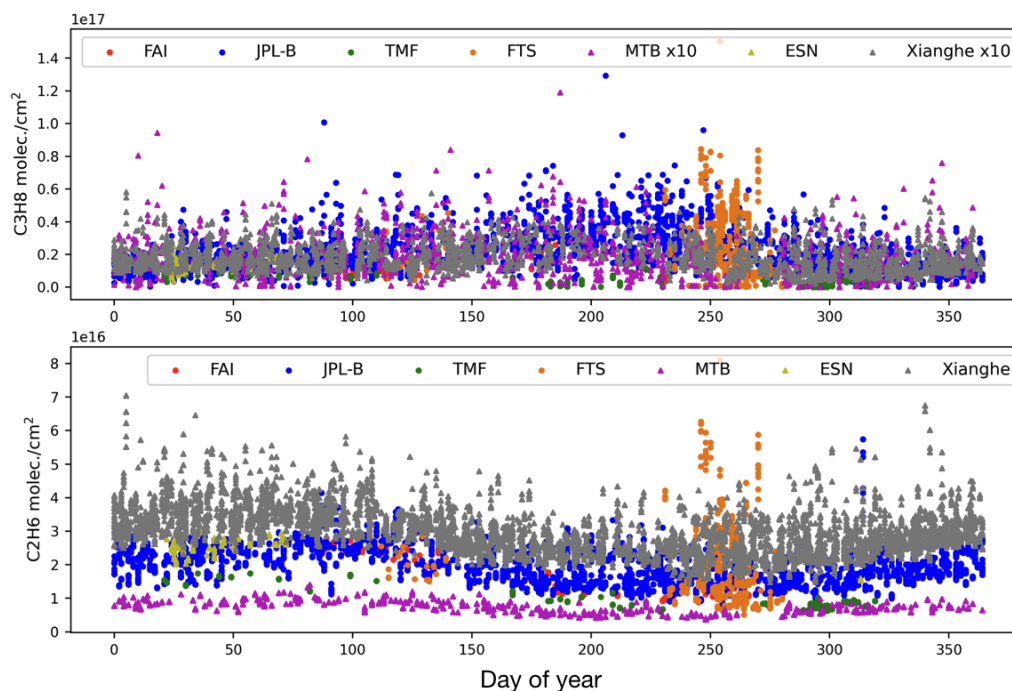
Figure 8. The C_3H_8 (upper panel) and C_2H_6 (lower panel) columns observed by ground-based Bruker IFS 125HR at Xianghe and MkIV spectrometer at 6 sites. Note that the C_3H_8 columns observed at MTB and Xianghe are multiplied by 10 to have a better view.

Table 4. The locations, and data time coverages of the MkIV measurements at 6 sites, together with their mean C_3H_8 and C_2H_6 columns. The bottom row is the Xianghe FTIR measurements in this study.

Site	Country	Latitude	Longitude	Altitude (km)	Time coverage	C_3H_8 (molec./cm^2)	C_2H_6 (molec./cm^2)
Esrange(ESN)	Sweden	67.89°N	21.08°E	0.271	Nov 1999 - Mar 2020	1.4×10^{16}	2.6×10^{16}
Fairbanks(FAI)	USA	64.83°N	147.61°W	0.182	Mar-Sep 1997	1.4×10^{16}	1.8×10^{16}
Mt. Barcroft(MTB)	USA	37.58°N	118.23°W	3.801	Oct 1998 - Aug 2002	1.4×10^{15}	7.3×10^{15}
Ft. Sumner(FTS)	USA	34.48°N	104.22°W	1.260	Oct 1989 - Sep 2021	2.6×10^{16}	1.9×10^{16}
TMF, Wrightwood (TMF)	USA	34.38°N	117.68°W	2.257	Jul-Sep 1988; Nov 1996 Jan-Aug 1998; Oct 2009	2.7×10^{15}	8.3×10^{15}
JPL B183(JPL-B)	USA	34.20°N	118.17°W	0.345	Jun 1985 - Jan 2022	2.1×10^{16}	2.0×10^{16}
Xianghe	China	39.75°N	116.96°E	0.036	Jun 2018 - Jul 2022	1.8×10^{15}	3.0×10^{16}

The mean and std of the C_3H_8 column derived from the FTIR measurements at Xianghe are $1.80\pm0.81 \times 10^{15}\ molec./cm^2$. A month-to-month variation is observed with a high value in July and a low value in October. The difference between the

median values in July (maximum) and October (minimum) is $1.2\times 10^{15}\ molec./cm^2$. As C_3H_8 are co-emitted with CH_4, CO, C_2H_2, and C_2H_6 during oil and gas production, we calculate the correlation between ΔC_3H_8 and these species at Xianghe. Good correlations are found between C_3H_8 and C_2H_6, between C_3H_8 and C_2H_2, as well as between C_3H_8 and CO. However, the correlation between C_3H_8 and CH_4 is relatively weak, which is probably due to CH_4 emission in North China being

5   dominated by rice, cultivation, and waste, instead of oil and gas production and fossil fuels combustion. By comparing the C_3H_8 and C_2H_6 columns at Xianghe with 6 other sites around the world, provided by the ground-based MkIV spectrometers, we find that the C_2H_6 column at Xianghe is the largest. However, the C_3H_8 column at Xianghe is only larger than those observed at the mountain sites and polar sites, and it is much less than the C_3H_8 columns observed at mid-latitude sites in the USA. Currently, the reported uncertainty of MKIV C_3H_8 measurements is relatively large of about $8\times 10^{15}\ molec./cm^2$,

10 which is much larger than the mean FTIR C_3H_8 measurements at Xianghe. Further investigation is needed to understand the large difference between FTIR and MKIV C_3H_8 measurements.

In summary, we successfully retrieve C_3H_8 columns from the FTIR MIR spectra at Xianghe, which provides us with a new dataset to understand the variation of C_3H_8 in North China. The retrieval strategy of C_3H_8 in this study should work at other Bruker 125HR FTIR sites as well, especially for those close to a city or oil and gas field, e.g., Paris, Toronto, and

15 Boulder. Nevertheless, efforts are still needed within the NDACC-IWRG community to generate a global harmonized FTIR C_3H_8 column dataset.

Data availability. The ground-based MkIV C_3H_8 and C_2H_6 retrievals are publicly available via https://mark4sun.jpl.nasa.gov/ground.html (last access date: 27 September 2022). The FTIR C_3H_8 retrievals at Xianghe are available upon request. The WACCM model data are publicly available via https://www.acom.ucar.edu/waccm/download.shtml (last access date: 27 March 2024).

20 *Competing interests.* The authors declare that they have no conflict of interest.

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Author contributions. PW and MZ design the study and wrote the manuscript. MZ, BL, BD, MDM investigated the SFIT4 retrieval strategy.

25 WN, CH and QC operate the FTIR measurements at Xianghe. GT provides the MKIV measurements. All authors have read and commented the manuscript.

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