# RC1: 'Comment on amt-2024-99', Anonymous Referee #1, 03 Jul 2024

This paper provides a brief overview of the radar-only, radar-lidar and radar-lidar-radiometer cloud and precipitation retrieval products from the JAXA Level 2 production model, with illustrations of key variables based on numerical forecast models and CloudSat/CALIPSO. The results in this paper are well-presented and summarises a long record of work with CloudSat/CALIPSO and other instruments in preparation for EarthCARE. While we appreciate concise papers, ultimately the paper does not provide enough of a detailed description of the algorithm(s), nor a consistent view of the data products in question. I recommend this paper for major revisions to address these issues.

We are sincerely grateful to all reviewers for their careful reading of the manuscript and useful comments that have allowed us to improve its quality. In our revised manuscript, we have made corrections and added a better overview and major details of the algorithms and products based on comments by the reviewers. Below are the reviewers' comments in blue text followed by our replies in black text.

# Major comments

 The names of the products (i.e. C-CLP, AC-CLP, ACM-CLP) have typically been given in the paper titles within this special issue, and this would help the user to navigate through the special issue as well as the EarthCARE L2 data products.

Product names have been added to the title as follows; JAXA Level 2 cloud and precipitation microphysics retrievals based on EarthCARE CPR, ATLID, and MSI: The C-CLP, AC-CLP, ACM-CLP products

2. There's a fundamental ambiguity in the Abstract, which isn't resolved within the paper. We read that, (L16-17) "these products provide a detailed view of [cloud properties] as well as vertical velocity information", but then in the final sentence, (L20) "Level 2 velocity-related products will be described in a future paper." Please make it very clear from the abstract onward that the vertical velocity products are not described in the present paper.

We have deleted "as well as vertical velocity information" (L16–17) and edited the final sentence to read: JAXA Level 2 velocity-related products (i.e., CPR\_VVL, AC\_VVL, and ACM\_VVL) such as hydrometeor fall speed and air vertical velocity will be described in a future paper. (L21-22)

3. Table 1 is key to detailing the production model represented within this paper. The "standard" version of each L2 product are given, and also "research" versions of each product which contain vertical velocity information (i.e. retrieved precipitation mass flux, in-cloud vertical air velocity). This is ambiguous: are the existing products to be updated at a later date to include these additional variables, or are additional products to be released? Asked another way: it's not clear from Table 1 or from the text of this paper whether the JAXA L2 production model as described in Eisinger et al. (2024) is accurate: in Figure 3 of Eisinger et al (2024) the C-CLP, AC-CLP and ACM-CLP products will provide the "standard" quantities relating to cloud mask, phase, shape and cloud microphysics, while additional "research" L2 data products with different names will provide complementary

information: e.g. C-RAS for CPR retrievals of rain and snow properties, C-VVL for vertical velocity; ACM-ICE for ice cloud effective radius as described in Eisinger et al. (2024). In Table 1, the titles of the three last columns indicate that all of these variables fall under "C-CLP", "AC-CLP" and "ACM-CLP", and no other products are named. Please resolve this ambiguity, while citing the Eisinger et al. (2024) paper which was intended as a centralised resource for understanding the full range of products available.

The JAXA L2 products summarized by Eisinger et al. (2024) are correct. Table 1 and the corresponding text have been updated for consistency with Eisinger et al. (2024). The main research products (C-RAS, AC\_RAS, ACM\_RAS, CPR\_VVL, AC\_VVL, and ACM\_VVL products) have been added to Table 1 and the text. The C-CLP, AC-CLP, and ACM-CLP products include both cloud and precipitation microphysics, but are reported only as cloud microphysics, whereas C-RAS, AC\_RAS, and ACM\_RAS include precipitation-only products using Doppler velocity. The C-CLP, AC-CLP, and ACM-CLP products will eventually be updated using Doppler velocity. For ACM\_CLP, this update will be included in ACM\_CDP, which is processed by JAXA Laboratories (LR), but results fulfilling the release criteria could be added to the ACM-CLP products. ACM-ICE is under consideration.

Subsection 2.1.1 is added in the paper as follows;

## "2.1.1 Primary cloud products

Standard cloud property (CLP) products (i.e., CPR\_CLP, AC\_CLP, and ACM\_CLP) include a cloud mask, cloud particle type, cloud particle habit category, cloud microphysics, cloud optical thickness, and cloud water/ice paths (Table 1). The microphysical properties of all hydrometeor types in the standard products are reported in the cloud microphysics product, and precipitation-sized particles are not separated into precipitation products. JAXA L2 research cloud products include velocity-related products such as sedimentation velocity and air vertical velocity (Sato et al., 2009), which are designated CPR\_VVL, AC\_VVL, and ACM\_VVL; precipitation-only products (e.g., rain and snow rates; CPR\_RAS, AC\_RAS, and ACM\_RAS) (Table 1). Details of these research products will be reported in a future paper. All products are reported using the Joint Standard Grid (JSG) with 1-km horizontal and 100-m vertical grid spacing. Note that CPR\_CLP, AC\_CLP, and ACM\_CLP are produced with and without the use of L2 CPR Doppler velocity to show the effect of additional information obtained from Doppler velocity. Similarly, research products will be developed through RAS and VVL, and results fulfilling the release criteria may be added to the standard products (i.e., CPR\_CLP, AC\_CLP, and ACM\_CLP) for release."

- 4. The algorithm description section 2.3 of this paper is very dense and difficult to understand.
  - It begins by citing the ATBD, but a link to this document is not provided—and in any case, I would ask that the authors provide some recapitulation of the main detail provided in the ATBD within this paper, or in the

citations within. As in other papers in this special issue, this should provide at least an overview of the algorithm, the key physical assumptions, etc.

The descriptions of the L2 cloud algorithms in section 2.3 have been improved to provide a better overview and major details of the algorithms. The reference to ATBD has been replaced by Okamoto et al. (2024b), which is a better reference for the interrelations among algorithms and products, corresponding to the subjects covered in this paper. A general summary of the JAXA L2 cloud microphysics algorithms is added in subsection 2.1.3

• It would be helpful to break the description into paragraphs at least, but even more helpfully into sub-sections: cloud mask, phase discrimination, cloud microphysics. Here it would also be useful to provide a description of the intended use of Doppler measurements for vertical velocity products, even if they are to be more fully described in a later paper.

Thank you for your suggestion.

The text in section 2.3 has been divided into subsections related to the preprocessing algorithms for cloud microphysics retrieval (i.e., subsection 2.3.1 Preprocessing for cloud microphysics retrieval; 2.3.1.1Cloud mask; 2.3.1.2 Cloud type; 2.3.1.3 Cloud particle category (CPC)), and cloud microphysics retrieval algorithms (i.e., 2.3.2 Cloud microphysics; 2.3.2.1 Ice cloud microphysics; 2.3.2.2 Liquid cloud microphysics).

The intended use of Doppler velocity has been added to subsections 2.3.1 and 2.3.2 for cloud preprocessing and microphysics retrieval and subsection 2.3.3 for air vertical velocity (i.e., 2.3.3 Intended use of Doppler measurements for air vertical velocity and terminal velocity products).

The following subsections has been added in the paper:

# "2.3 Description of the JAXA Level 2 cloud microphysics product

The following section provides a brief overview and highlights of the standard JAXA L2 cloud microphysics products.

2.3.1 Preprocessing for cloud microphysics retrieval

2.3.1.1 Cloud mask

The ATLID-only cloud mask is processed by the ATLID\_CLA algorithm (Nishizawa et al., 2024), the CPRonly cloud mask is processed by the CPR\_CLP algorithm (Okamoto et al., 2024a), and the MSI-only cloud mask is processed by the MSI\_CLP algorithm (Nakajima et al., 2019). For ATLID, aerosol, cloud, and surface components are discriminated from clear pixel when the Mie backscattering coefficient is significant compared to the noise level (Nishizawa et al., 2024). A cloud mask scheme is then applied; this scheme includes a vertically variable threshold value for the Mie backscattering coefficient (or particle backscattering coefficient when the Rayleigh backscattering coefficient is significant), as well as a spatial continuity test to exclude noisy pixels. The lack of sufficient surface signal is used to identify fully attenuated ATLID pixels below aerosol or cloud layers. Similarly, the CPR cloud mask scheme considers noise level, continuity testing, and surface echo information to determine sufficient radar echo power for cloud and precipitation analysis, as well as full attenuation of the radar signal. The AC\_CLP synergy cloud mask scheme merges the single active sensor cloud mask results from ATLID\_CLA and CPR\_CLP, and then flags cloudy pixels in ATLID, CPR, or both. MSI cloud mask information is not used for the ACM\_CLP cloud mask. The AC\_CLP and ACM\_CLP cloud mask products are currently identical.

## 2.3.1.2 Cloud type

The ATLID cloud type scheme (ATLID\_CLA) uses  $\delta$ ,  $\beta_{att}$ , and temperature to identify the cloud phase and ice particle orientation, which is designated as two-dimensional (2D) ice, three-dimensional (3D) ice or mixed 2D and 3D ice (Okamoto et al., 2024a). The CPR cloud type scheme (CPR\_CLP) uses mainly Z<sub>e</sub> (along with its vertical profile) and temperature to discriminate the hydrometeor phase, ice particle orientation, precipitation type (snow, drizzle, or rain) and melting layer (Okamoto et al., 2024a). The AC\_CLP synergy cloud type scheme combines ATLID\_CLA with CPR\_CLP and reclassifies the cloud type when estimates from the two sensors differ according to the classification rule specified by Kikuchi et al. (2017). The differing particle size sensitivity of CPR and ATLID aid in the identification of mixed-phase clouds and mixed cloud-precipitation types (i.e., cloud water + drizzle or cloud water + rain). The ACM\_CLP and AC\_CLP cloud types are identical. In addition, Doppler velocity will be used to improve differentiation between snow and rain and between cloud and drizzle. Further details of the cloud mask and cloud particle type products were reviewed by Nishizawa et al. (2024a) and Okamoto et al. (2024a).

# 2.3.1.3 Cloud particle category (CPC)

After applying the cloud mask and cloud phase discrimination schemes (Okamoto et al., 2024a), one of the main products of the EarthCARE JAXA L2 cloud product is the cloud particle category product, which enables more detailed comprehensive exploration of the ice particle habit category contained within each JSG grid. Among cloud particle categories, the liquid-phase types are the same as those in the cloud type product (subsection 2.3.1.2). Ice particles are further categorized based on ATLID lidar ratio and depolarization ratio diagrams (Okamoto et al., 2019; 2020; Sato and Okamoto, 2023). This information is anticipated to be instrumental for general remote sensing applications (Van Diedenhoven, 2018; Letu et al., 2016) and the development of ice optical parameterization (Li et al., 2022) and hydrometeor sedimentation velocity parameterization for use in numerical models. The retrieved ice particle habit categories include horizontally oriented 2D plates and their assemblages, 2D columns and their assemblages, bullet rosettes and 3D-oriented aggregate types, droxtal/compact types, Voronoi/irregular/roughened types, and fractal-type snow aggregates (Ishimoto et al., 2008, 2012). ATLID-only CPC is used to train the CPR-based algorithm for ice particle category retrieval from Z<sub>e</sub> and temperature information in regions with CPR-only measurements. The CPR-only CPC product is obtained from CPR\_CLP. CPR\_CLP, and ATLID-only CPC are combined to produce the synergy AC\_CLP CPC product. For ice categories, ATLID-only CPC estimates are used when both CPR\_CLP cnd ATLID-only CPC estimates are used when both CPR\_CLP cnd ATLID-only CPC estimates are available for

the same JSG grid. The Doppler velocity will be further used to improve category identification, particularly for snow types (e.g., graupel or hail).

## 2.3.2 Cloud microphysics

In CPR\_CLP, ACP\_CLP, and ACM\_CLP, forward models corresponding to the derived cloud particle categories are used to analyze the observations from each sensor, and microphysics corresponding to each category are thus obtained. The single scattering properties of ice particles with various shapes and orientations are calculated using physical optics (Borovoi et al., 2012) and modified geometrical optics integral equation methods (Masuda et al., 2012) for ATLID specification (Okamoto et al., 2019), and discrete dipole approximation and finite-difference time domain (FDTD) methods for CPR wavelength (Sato et al., 2011; Ishimoto et al., 2008, 2012); Mie theory is used for the liquid phase and multiple scattering effects are estimated based on Sato et al. (2018, 2019).

The total effective radius for cloud and precipitation information is given as:

$$r_{\rm eff} = \int r_{eq}^3 \frac{dn(r_{eq})}{dr_{eq}} dr_{eq} \left/ \int r_{eq}^2 \frac{dn(r_{eq})}{dr_{eq}} dr_{eq} \right. \tag{1}$$

where  $r_{eq}$  is the melted mass equivalent radius to a sphere,  $dn/dr_{eq}$  is the size distribution function. For both iceand liquid-phase clouds, a maximum of two different particle size distributions (i=1,2) can be considered within one JSG grid to handle the presence of cloud and precipitation modes, i.e.,  $\frac{dn(r_{eq})}{dr_{eq}} = \sum_{i=1}^{2} \frac{dn_i(r_{eq})}{dr_{eq}}$ . The corresponding effective radius is given as:

$$r_{\text{eff,i}} = \int r_{eq}^3 \frac{dn_i(r_{eq})}{dr_{eq}} dr_{eq} / \int r_{eq}^2 \frac{dn_i(r_{eq})}{dr_{eq}} dr_{eq} \quad (i = 1, 2)$$
(2)

For dni/dreq, a modified gamma size distribution,

$$\frac{dn_i(r_{eq})}{dr_{eq}} = \frac{No_i i}{\Gamma(p)r_{m,i}} \left(\frac{r_{eq}}{r_{m,i}}\right)^{p-1} exp\left(-\frac{r_{eq}}{r_{m,i}}\right) \quad (i = 1,2)$$
(3)

in which  $r_m$  is the characteristic radius and the dispersion value is p = 2 (Okamoto, 2002; Sato and Okamoto, 2011), is employed for cloud ice, snow, and rain in cold precipitation. A log-normal size distribution,

$$\frac{dn_i(r_{eq})}{dr_{eq}} = \frac{No_i i}{\sqrt{2\pi}r_{eq}ln\sigma} exp\left\{-\frac{\left[ln(r_{eq}/r_{o,i})\right]^2}{2(ln\sigma)^2}\right\} \qquad (i = 1,2)$$
(4)

in which  $r_o$  is the mode radius and the standard deviation of the distribution is  $\sigma = 1.5$  (Okamoto, 2002), is assumed for warm water, super-cooled liquid, and warm precipitation.

In the following, general approaches for cloud microphysics retrievals are explained based on the AC\_CLP cloud microphysics algorithm, which are common to CPR\_CLP and ACM\_CLP cloud microphysics algorithms.

#### 2.3.2.1 Ice cloud microphysics

For ice clouds, a lidar-only cloud region, lidar-radar overlap cloud region, and radar-only region generally

exist for ice and liquid precipitation. An algorithm to retrieve microphysical properties that considers a mixture of two particle types at maximum (i.e., 2D and 3D ice) has been developed for ice cloud regions observed with CloudSat and CALIPSO synergy (Okamoto et al., 2010) using Z<sub>e</sub>, the attenuated backscattering coefficient β, and the depolarization ratio. A framework to extend the applicability of the microphysics retrieval algorithm from the cloud region to the entire precipitation region in the vertical column was developed to efficiently reflect information from the lidar-radar overlap region to the microphysics retrieved in the CloudSat- or CALIPSOonly region (Sato et al., 2011, 2020). The relationships between microphysical properties ( $r_{eff}$  and IWC) and  $\beta$ or Z<sub>e</sub> in the vertical cloud grids of the lidar-radar overlap region were derived for each profile and used to estimate the microphysical properties in the radar- or lidar-only cloud region (Sato et al., 2011). The EarthCARE JAXA L2 cloud microphysics retrieval algorithms extend these algorithms in the following three aspects: (1) the spatial variability of the microphysics and observables are considered to derive more reliable relationships among cloud microphysics and observables, (2) the microphysics estimates in the ice precipitation region far from the lidar-radar overlap region of a precipitation system are further improved by extending the microphysics estimates from the precipitation region upward rather than downward from the lidar-radar region (Heymsfield et al., 2018), and (3) single-size mode for cloud ice is considered for lidar-only cloud region and lidar-radar overlap cloud region, while two different size modes for cloud ice and ice precipitation (snow) are considered for the CPR-only region existing from the bottom altitude of the lidar-radar overlap region to the top altitude of the melting level. The PIA is used to correct the attenuation of  $Z_e$ . (Iguchi et al., 2000).

Specifically, for (1), the L2 cloud microphysics algorithm uses  $r_{eff}$  and IWC for all horizontal and vertical grids within the radar–lidar overlap region embedded in each cloud system to obtain robust relationships of cloud microphysics with  $Z_e$  and  $\beta$  (e.g.,  $Z_e$ –IWC relationships,  $Z_{e,1}$ =  $a_1IWC_1^{b1}$  are determined for each record, where  $Z_e$  [mm<sup>6</sup> m-<sup>3</sup>] and IWC [g m-<sup>3</sup>]). These relationships are derived for each record using all data within each cloud system (or within a single EarthCARE orbit frame when a sufficient number of points cannot be obtained to derive the statistics) weighted by distance from the target profile record and are used to provide initial estimates of cloud ice microphysics based on  $Z_e$  or  $\beta$  in the CPR-only (ice cloud and ice precipitation) or ATLID-only (ice cloud) regions, respectively.

For (2), the relationship between the microphysics and observables is expected to change from the cloud region to the precipitation region. Because lidar signals are fully attenuated at optically thick precipitation region, new relationships for ice precipitation are derived using CPR data. In this process, CPR data at melting levels or layers around the ice–liquid interfaces of a precipitation system are used. At the top of the melting level, it is assumed that only precipitation mode exists ( $Z_e=Z_{e,2}$ ), and during melting, the mass in each size bin (i.e.,  $r_{eff}$ ) remains constant across several successive layers (Heymsfield et al., 2018). For a given  $r_{eff}$ , dBZ <sub>e</sub> changes due to the different scattering properties for ice and liquid. Therefore,  $r_{eff}$  and IWC (or LWC) are derived and the relationships ( $Z_{e,2}=a_2IWC_2^{b^2}$ ) can be established for ice precipitation (snow) holding the coefficient b<sub>2</sub> at the value derived in (1) (b<sub>2</sub>=b<sub>1</sub>) for each record.

For (3), Z<sub>e,1</sub> and Z<sub>e,2</sub> for the two size modes (cloud ice and snow) in the CPR-only ice precipitation region at

each vertical grid ( $Z_{e,1} + Z_{e,2} = Z_e$ ) are determined as follows. The ratio IWC<sub>2</sub>/(IWC<sub>1</sub>+IWC<sub>2</sub>) =IWC<sub>2</sub>/IWC=*A* increases linearly from 0 at the bottom of the lidar–radar overlap region to 1 at the top of the melting level. *A* is given as,  $A = \int_h^{ht} Z_e \, dh \, / \int_{hm}^{ht} Z_e \, dh$ , with a range of 0 to 1, where the integrated  $Z_e$  from the bottom altitude of the lidar–radar overlap region (h<sub>t</sub>) to a certain altitude h below h<sub>t</sub> within the CPR-only ice precipitation region ( $\int_h^{ht} Z_e \, dh$ ) is normalized using the value integrated to the melting level altitude h<sub>m</sub> ( $\int_{hm}^{ht} Z_e \, dh$ ). As the Z<sub>e</sub>–IWC relationships for both cloud ice and snow are derived, determining the vertical profile of IWC<sub>2</sub>/IWC is equivalent to providing the relationship between Z<sub>e,1</sub> and Z<sub>e,2</sub> for each vertical grid. Therefore IWC<sub>i</sub>, reff.<sub>i</sub> (i=1,2) and other microphysical properties are derived for each JSG grid (Table 1).

In microphysics retrieval for convective/stratiform rain below the melting level, only the precipitation size mode is assumed to exist. The r<sub>eff</sub> and LWC obtained at the rain top altitude of each observation record described in (2) are used to derive the No and x values of the Marshall–Palmer size distribution (dn/dD = No e<sup>-AD</sup> [m<sup>-4</sup>], where D is the particle diameter,  $\Lambda = xR^{0.21}$  [cm<sup>-1</sup>] and R is the rain rate in mm/hr, which is a function of LWC and r<sub>eff</sub>) (Marshall and Palmer, 1948). No and x are assumed to be constant within the vertical profile for rain in a given record and are used to determine the vertical profiles of LWC and r<sub>eff</sub> for the modified gamma size distribution associated with each Z<sub>e</sub> value in the rain region.

Generally, for the same  $Z_e$ , when the mass mixing ratio of the small mode to total IWC is overestimated (underestimated), optical thickness will be overestimated (underestimated); in the 3-sensor ACM\_CLP algorithm, the mass mixing ratio of the two size modes is further constrained by the optical thickness obtained from the MSI. When only a single size mode is present, the r<sub>eff</sub> and IWC of the single mode are adjusted to be consistent with MSI optical thickness retrievals. Doppler velocity is expected to effectively improve particle sizing in regions of ice and liquid precipitation, as well as in the breakup of large snow particles during melting (e.g., Fujiyoshi et al., 2023).

#### 2.3.2.2 Liquid cloud microphysics

A two-size-mode approach similar to the ice cloud microphysics retrieval process is used for water clouds, which considers the coexistence of cloud particles and drizzle. CPR\_CLP derives the liquid microphysics corresponding to each size mode from CPR-only scheme. In AC\_CLP and ACM\_CLP, for JSG grids with ATLID observables, ATLID  $\delta$  and  $\beta_{att}$  (or  $\sigma_{ext}$ ) are used to derive  $r_{eff,1}$  and LWC<sub>1</sub> for cloud water or super-cooled water (Sato et al., 2018, 2019; Sato and Okamoto, 2020). As ATLID is expected to provide a better estimate of the cloud mode than CPR, for the CPR and ATLID overlap region, the ATLID cloud microphysics and  $Z_{e,1}$  estimate are used for microphysics estimation of the drizzle mode.

In water clouds, *in situ* and ground-based radar measurements have shown that cloud particles and drizzlesized particles can coexist above  $-35 \text{ dBZ}_e$  (Baedi et al., 2000). Except at very small (<  $-35 \text{ dBZ}_e$ ) and large values of  $Z_e$ , where only a single mode is likely to occur, the cloud mode can dominate LWC and r<sub>eff</sub>, whereas the precipitation mode can dominate  $Z_e$  (Baedi et al., 2000; Krasnov and Russchenberg, 2005). For this reason, in general, the dependence of total LWC on  $Z_e$  differs significantly from results derived for only cloud particles (LWC<sub>1</sub> and Z<sub>e,1</sub>) or only drizzle-sized particles (LWC<sub>2</sub> and Z<sub>e,2</sub>) (Baedi et al., 2000). PIA is sensitive to total LWC, and in the CPR-only microphysics retrieval scheme, the  $Z_e$ -LWC relationship ( $Z_e = aLWC^b$ , where  $Z_e$  [mm<sup>6</sup> m<sup>-3</sup>] and LWC [g m<sup>-3</sup>]) and LWC for the cloud+drizzle mode for the JSG grids within each record are determined from PIA and  $Z_e$  assuming that b = 5.17 (Baedi et al., 2000). The power  $b_i$  of the  $Z_e$ -LWC relationship for clouds and drizzles are reported to have similar values and assumed to be fixed (i.e.,  $b_1 \sim 1.17$ ; Baedi et al., 2000, Fox and Illingworth, 1997,  $b_2 \sim 1.58$ ; Krasnov and Russchenberg, 2002), while the coefficients  $a_i$  in the  $Z_e$ -LWC relationship could differ between clouds and drizzles by several orders of magnitude reflecting the size distribution difference (Khain et al., 2008). As CPR  $Z_e$  is more sensitive to the drizzle mode (i.e.,  $Z_{e,2}$ ), the  $a_1$  coefficient for cloud mode is assumed to be initially fixed at reported value ( $a_1=0.015$ ; Baedi et al., 2000), and  $a_2$  is derived for each  $Z_e$  and LWC profile, given that LWC<sub>1</sub> + LWC<sub>2</sub> = LWC and  $Z_{e,1} + Z_{e,2} = Z_e$ . Finally,  $Z_{e,i}$ ,  $r_{ef,i}$ , LWC<sub>i</sub> (i=1,2), and other microphysical properties such as the number concentration and particle fall speed are derived for the two size modes.

The liquid cloud microphysics are further constrained by the ATLID observables for the AC\_CLP and ACM\_CLP algorithms, and the MSI for the ACM\_CLP algorithm. Doppler information will be used to improve the microphysics estimates of the precipitation (drizzle) mode.

## 2.3.3 Intended use of Doppler measurements for air vertical velocity and terminal velocity products

The Doppler velocity is intended to be used in at least two approaches; air vertical velocity will be determined by subtracting the  $Z_e$ -weighted particle fall speed corresponding to each cloud particle category obtained without the use of Doppler velocity, and simultaneous retrieval of air vertical velocity and microphysics through an approach similar to that described by Sato et al. (2009), which considers the difference between the vertical structures of  $Z_e$  (reflecting cloud microphysics) and  $V_D$  (which is affected by air vertical velocity and cloud microphysics) to extract the air vertical velocity component. "

#### Minor comments

#### • L91: please provide a DOI for the ATBD

The reference to ATBD has been replaced by Okamoto et al. (2024b) (reply to major comment 4). This reference is cited in section 2.2.

## • L119: what are the two different size distributions?

The description for particle size distributions is provided in subsection 2.3.1 in the revised manuscript. A modified gamma size distribution is assumed for cloud ice and snow, and a log-normal size distribution is assumed for warm water, super-cooled liquid, and warm precipitation. For both ice- and liquid-phase clouds, a maximum of two different particle size distributions can be considered within one JSG grid to handle the presence of multiple cloud modes (i.e., cloud ice, cloud water, or super-cooled water), precipitation modes (drizzle, rain, or snow), and

cloud particles of differing phases. Therefore, two different effective radii with corresponding ice water or liquid water content and other microphysical properties are derived for each active sensor grid within a vertical profile. The retrieval procedure of the two size modes are provided in subsections 2.3.2.1 and 2.3.2.2.